General Comments: Voigt et al. present a manuscript detailing the monitoring and modeling of a lake in semiarid southern Spain. Using a multi-systems approach, they model lake levels as a function of parameters that are commonly measured in modern systems as well as characterize potential proxy variables for the reconstruction of paleohydrological conditions. I find their manuscript carefully written and well-presented. I have a series of minor questions and inquiries as line comments below. In addition, I have one main desire which would be the inclusion of more detail on how their salts-based and isotope-based modelling might influence paleohydrologic studies. This linkage is teased in the front of the manuscript but perhaps could be expanded on somewhat in the discussion.

We thank the reviewer for his detailed comments, which will considerably help to improve the manuscript. The discussion of implications of our results for paleoclimate studies has also been suggested by reviewer 1. These will be addressed in the revised manuscript as outlined in the author response to reviewer 1. Below, we respond to each of the specific comments (blue) and indicate how the comment will be addressed in the revised manuscript (yellow).

Figure 1: Is the extrapolation of lake SA and volume consistent with previous lake levels? Or simply an extension based on the established topography? In panel (a), what do the numerical values inset on each lake contour indicate? I assume total lake volume at that water level?

The small numbers in panel (a) indicate the water level in (m) along the contour line. This bathymetric map has been published previously by Castro et al. (2003). We digitalized this map, estimated the surface area for each lake contour line and linearly interpolated between the contour lines. The lake volume was then estimated based on the surface area of slices at discrete depth intervals of 0.01 m:

Lines 155-161: E_SA is estimated from potential evapotranspiration and lake surface area? Or is there additional treatment of evaporation-driven water losses from the lake? Aren't there considerations needed based on wind fields, surface roughness, air-water temperature differences, etc. when estimating something like E_SA? This is not my expertise, so perhaps a lake of this size and particular setting make these less of an issue, but my general understanding is that E_SA is typically non-trivial.

We assume that evapotranspiration from the lake surface (E_SA) is equal to potential evapotranspiration multiplied by the simulated lake surface area as specified in line 161. Indeed, the estimation of lake evaporation is not a simple matter as there are a number of factors that can affect the evaporation rates, notably the climate and physiography of the water body and its surroundings. A wide variety of methods for estimating open water evaporation has been reported in the literature and used in practice, including pan evaporation, mass balance, energy budget models, among others. The fact that the simulated water level agrees well with the observed water level indicates that the potential evapotranspiration provides a reliable estimate of lake evaporation in our case. We will add this information in the result section (Line 273 ff.), as follows:

"Lake water level variations estimated from hydrological mass balance (Eq. 1), accounting for precipitation on the lake surface, evaporation from the lake surface and basin discharge, agree well with observations over the period of lake water isotope monitoring (model-data deviations are lower than 5 cm) (Fig. 2). This supports that major water sources are integrated, and the lake is disconnected from groundwater aquifers. It further indicates that potential evapotranspiration provides a reliable estimate of lake evaporation for the studied lake site."

Lines 173-184: This section also does not mention in any depth the relevance of the E_SA term. Obviously, from proper evaporation, we do not have any meaningful losses of salinity. However, depending on the seasonal strength of the wind field then we might expect some losses due to lake spray aerosols. Again, perhaps this is a non-issue due to the size of this lake and the environmental conditions, but some text on the E_SA term here could help more fully explain your methodology even if the processes I describe are completely negligible.

This is an interesting thought. Loss of lake water by aerosols would not only increase lake evaporation, but also lead to removal of salts from the lake water. However, considering the small size of the lake and in view of the good agreement of the observed and simulated water level (hydrological balance) we think that this process is of minor importance in our studied case.

Line 212-217: Can you expand some on how these sensitivity experiments were handled? As Monte Carlo-type simulations? Were distributions for variables normal, uniform, etc.? Were all variables allowed to freely vary within their uncertainty bounds or was each sensitivity experiment the variation of a single parameter?

We evaluated the sensitivity of the model results to the input parameters by varying one of them within the limits specified in section 3.5 and keeping other model input parameters constant. No Monte Carlo simulation was applied.

Line 221-223: What is the error window of the bathymetric model? Perhaps the uncertainty of the lake bathymetry is small enough that any volume error from bathymetry (as opposed to lake level) is negligible, but given all the other considerations, perhaps this is worth including.

The lake bathymetric model has a resolution of 0.5 m. This is quite coarse, which means that the linearly interpolated bathymetric model can significantly deviate from the real situation. However, results of the sensitivity experiments showed that misestimation of the initial lake water level (and thus lake volume) has relatively little influence on the simulated lake water isotope composition (line 349-353). We think it's worth to keep these analysis results to justify the use of the linear bathymetric model.

Line 235-239: Another possible consideration is the difference between air temperature and lake temperature. Although this is a small lake and probably reasonably tracks monthly air temperature, diurnal variations in air temperature may result in larger-thanexpected evaporation occurring during nighttime when air temperatures drop but the lake remains warm.

Our model simulations are performed on a daily timestep. As the lake investigated in our study is small and shallow, we assume that the lake water temperature is equal to that of the atmosphere on a daily scale. Diurnal variations are not accounted for in our simulation, but can contribute significant uncertainty to our simulation results. In Section 5.1, we discuss the importance of diurnal variations of the relative humidity along with the evaporation rate. Temperature differences between lake water and the atmosphere further increase this uncertainty. We modified the paragraph as follows to include this information:

"On the other hand, the effective relative humidity during evaporation may be lower than its daily average value as the evaporation rate is usually higher during daytime when relative humidity is lower (Gibson, 2002). Also, temperature differences between lake water and the atmosphere may influence evaporation rates and effective relative humidity. Constraining the effective relative humidity that drives the isotope composition of the evaporation flux with precision is thus challenging."

Line 247-255: How well-constrained is *n*, really? A margin of 0.1 (as a SD of a normal distribution?) may be adequate, but I wonder if a larger range should have been considered.

The "margins" given for each input parameter in Section 5.1 do not refer to uncertainty margins but to the value each parameter has changed individually, keeping other parameters constant, to quantify its impact on the simulation. For the turbulence coefficient n a change of 0.1 appears to be reasonable. The

reader can think of a larger change (e.g. 0.2) that would result in a larger change (double) of the simulated results.

Line 270 / Figure 2: I wonder how many discrete rain events you have based on weather station data? Would it be useful (and legible) to include indicators for each rain event observed during the period of observation? You could perhaps limit indicators to rain events above some threshold value. This might be helpful in seeing the step-wise linkage between precipitation and lake level. I see the stepwise increases in cumulative precipitation at the top of Fig. 2. I wonder if a bar chart of each rain event (or binned 1 or 2-week cumulative values) would serve the visual explanation of this data better than a running accumulation timeseries? Just thoughts!

Thanks for this thoughtful comment. Indeed, it was quite difficult to illustrate the impact of precipitation on the lake water level along with the much larger quantities of basin discharge and evaporation. In the revised version of the manuscript, we will add panel (e) of Figure 5 as an additional panel in Figure 2:

Line 385-392 / Figure 5: This is great work and shows the strength of the sensitivity of isotope data to the important tunable parameters. I wonder exactly how the authors arrived at this particular solution and whether or not an automated solution approach (perhaps via Markov-Chain Monte Carlo) could identify whether or not multiple possible solutions exist. Further, I wonder if the variability in each parameter is well constrained by the bounds selected here. Certainly, we can imagine *n* to vary more than between 0.4 and 0.6, but also if we consider things like diurnal variability it may be possible to imagine that a 5% +/- bound on relative humidity is too conservative. Obviously, you must strike a middle point between unbounded variables and exact values – I think the authors have done a good job here generally but should, perhaps, include some additional reasoning on their stated bounds.

The input parameters for the simulation result illustrated by the dashed yellow line in Figure 6 were determined empirically. To achieve this, we divided the dataset into rainy and dry seasons and manually adjusted the input parameters within their uncertainty margins to improve model-data agreement. While this simulation result is not necessarily the best fit, it serves as an example of how deviations between modeled and observed data can be reduced. Quantifying the exact input values required for the best fit lies beyond the scope of this manuscript.

In response to the reviewer's comment, we performed a Monte Carlo simulation to validate our empirical approach. This simulation used uniform distributions for the input parameters (relative humidity, the isotope equilibrium coefficient for atmospheric water vapor, and the turbulence coefficient) and allowed them to vary within the uncertainty margins defined in the sensitivity experiments. To account for seasonal variations, the dataset was divided into rainy (January–May 2021 and October 2021–January 2022) and dry (June–September 2021) seasons. A total of 500 model runs were conducted, and the deviation between simulated and observed δ¹⁸O, δ²H, and ¹⁷O-excess of lake water was **calculated for each run. The top 10 % of simulations (those with the lowest model-data deviations) were analyzed.**

The Monte Carlo simulation results corroborate our empirical findings, showing similar trends in relative humidity and the turbulence coefficient. Furthermore, the analysis suggests that model-data deviations can be reduced by accounting for variability in the isotope equilibrium coefficient between atmospheric water vapor and precipitation. These findings are consistent with the results of the sensitivity experiments (see Fig. A6).

As the Monte Carlo simulation does not introduce new insights beyond those provided by the empirical approach and sensitivity experiments, we believe its inclusion in the manuscript is not necessary. However, we will modify the results section as outlined below, to clarify that this simulation is only one alternative to reduce model-data discrepancies.

Figure R1: Similar to Fig. 5 of the main text, but here the dashed yellow lines show the top 10% of the simulated isotope mass balance obtained from the Monte Carlo simulation. The Monte Carlo simulation suggests better model-data agreement using Veq = 0.9 ± 0.1, n = 0.51 ± 0.06, observed RH – 3.1 ± 1.3% from January to May 2021 and from October 2021 to January 2022 and using Veq = 1.1 ± 0.1, n = 0.47 ± 0.05 and observed RH + 1.9 ± 2.5%, from June to September 2021. Analytical errors are smaller than symbol size.

"No model-data agreement is found for a constant set of relative humidity, the turbulence coefficient and the isotope composition of atmospheric water vapor. Instead, seasonal variations in all three parameters need to be taken into account. Considering the results of the sensitivity experiments, better model-data agreement in the rainy season can achieved when using a relative humidity that is lower than the observed value and a slightly higher turbulence coefficient (Fig. 5a-b, yellow curve). In contrast, during the dry season, using a higher relative humidity value and lower turbulence coefficient is necessary to reduce the offset between the modelled and observed isotope composition of lake water (Fig. 5a-b, yellow curve). The simulated 17O-excess of lake water coincides only with observations when using a higher 17O-excess of atmospheric water vapor (33 per meg) (Fig. 5c, yellow curve)."

We will further modify the figure caption for Figure 5 as follows to clarify that this is only one possible solution:

"The dashed yellow line shows an alternative simulation that reduces the deviation between observed and simulated data. For this simulation, the dataset was divided in a rainy (January to May 2021 and October 2021 to January 2022) and a dry season (June to September 2021). Values of Veq = 1, n = 0.6, and observed RH – 5% were used for the rainy season and Veq = 1, n = 0.4 and observed RH +5 for the dry season. Analytical errors are smaller than symbol size."