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## Comparison of BARRA and ERA5 in Replicating Mean and Extreme Precipitation over Australia

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34

35 **Abstract**

36 Reanalysis datasets are critical in climate research and weather analysis, offering consistent  
37 historical weather and climate data crucial for understanding atmospheric phenomena, and  
38 validating climate models. However, biases exist in reanalysis datasets that would affect their  
39 applications under circumstances. This study evaluates BARRA, which is a high-resolution  
40 reanalysis for the Australian region, and ERA5 in simulating mean precipitation and six  
41 selected precipitation extremes for their climatology, temporal correlation, coefficient of  
42 variation and trend. Both models reproduce spatial patterns of mean precipitation well with  
43 minor biases. ERA5 shows stronger temporal correlations, superior inter-annual precipitation  
44 accuracy, and lower biases in coefficient of variation compared to BARRA, especially in  
45 Northern Australia. However, both models exhibit substantial biases in trend, underestimating  
46 increasing trends in Northern Australia. ERA5 underestimates dry days and heavy rainfall,  
47 while BARRA tends to overestimate these extremes. Temporal correlations for extreme  
48 precipitation indices are weaker compared to mean annual precipitation. Notable differences  
49 exist in variability biases, with BARRA showing larger biases, especially for heavy  
50 precipitation in inland regions and Northern Australia. While both datasets replicate the main  
51 trends, biases persist. Overall, the evaluation results support application of both datasets for  
52 climatology analyses, but caution is advised for variability and trend analyses, particularly for  
53 specific extremes.

54

55 **Key words:** BARRA, ERA5, extreme indices, temporal correlation, coefficient of variation,  
56 trend



## 57 **1. Introduction**

58 Reanalysis dataset is created by combining historical observational data from various  
59 sources, such as weather stations, satellites, buoys, and more, with modern data assimilation  
60 techniques and numerical models (Kalnay, et al. 1996; Saha, et al. 2010; Dee et al. 2011;  
61 Kobayashi et al. 2015, Poli et al. 2016; Hersbach 2020). The fundamental aim of reanalysis is  
62 to construct a uniform and coherent historical archive of various atmospheric and  
63 environmental parameters, such as temperature, humidity and wind patterns, on either a  
64 regional or a global scale.

65 These datasets are invaluable for climate studies, weather analysis and model validation  
66 as they provide a uniform representation of historical climate conditions. For instance,  
67 Quagraine et al. (2020) used five global reanalysis datasets to investigate the variability of West  
68 African summer monsoon precipitation, showing all datasets could represent the average  
69 rainfall patterns and seasonal cycle. Dai et al. (2023) utilized the fifth-generation European  
70 Centre for Medium-Range Weather Forecasts (ECMWF) reanalysis (ERA5) data to estimate  
71 rainfall erosivity on the Chinese Loess Plateau, finding rainfall erosivity derived from ERA5  
72 was highly consistent with those derived from the meteorological stations. Cheung et al. (2023)  
73 employed ERA5 to evaluate storm conditions in regional climate simulations, demonstrating  
74 regional climate models can capture climatology of measurements of storm severity over land  
75 including their spatial patterns and seasonality. Numerous studies have used reanalysis datasets  
76 as inputs for regional climate models (RCMs) to evaluate the models' capability in replicating  
77 observed climatic patterns (Solman et al., 2013; Ji et al., 2016; Fita et al., 2016, Di Virgilio et  
78 al., 2019; Capecchi et al., 2023; Di Virgilio et al., 2024; Ji et al., 2024).

79 While reanalysis datasets provide valuable insights into historical weather and climate  
80 conditions, they have limitations and uncertainties, given that they are modelled outputs rather  
81 than direct observations. Many studies have evaluated reanalysis data across various variables



82 and regions. For instance, Betts et al. (2019) assessed ERA5 biases in near-surface variables  
83 over Canada, highlighting its improved performance over ERA-Interim, though precipitation  
84 biases remained significant. Similarly, Hu and Yuan (2021) and Jiang et al. (2021) found that  
85 ERA5 precipitation accurately captured rainfall pattern over the Eastern Tibetan Plateau and  
86 mainland China, but under-estimated intensity. Izadi et al. (2021) found ERA5 performed  
87 better at monthly and seasonal timescales in Iran, underestimating coastal summer precipitation  
88 and overestimating it in mountains. Jiao et al. (2021) and Qin et al. (2021) found ERA5  
89 overestimated summer precipitation and frequency in China but underestimated intensity  
90 during the warm season. Lei et al. (2022) and Shen et al. (2022) noted ERA5's limitations in  
91 simulating extreme precipitation events in China, especially for high-end extremes.

92 Comparisons between reanalysis datasets have also been conducted. Wang et al. (2019)  
93 found that both ERA5 and ERA-Interim exhibited warm biases over Arctic Sea ice, with larger  
94 biases in cold season than warm season. Lei et al. (2020) showed ERA5 improved cloud cover  
95 simulation over eastern China but not over the Tibetan Plateau, when compared to ERA-  
96 Interim. Gleixner et al. (2020) found ERA5 reduced biases in temperature and precipitation  
97 over East Africa compared to ERA-Interim but still struggled with long-term trends. Song and  
98 Wei (2021) found both ERA5 and MERRA-2 captured night precipitation peaks over North  
99 China, but only ERA5 accurately reflected the afternoon peak. Li et al. (2022) concluded that  
100 ERA5 performed better than ERA-Interim, JRA55, and MERRA-2 in capturing precipitation  
101 over the Poyang Lake Basin. A summary of the above literature review can be found in Table  
102 S1.

103 In Australia, reanalyses like NCEP (Kalnay et al., 1996), JRA-55 (Kobayashi et al.,  
104 2015), ERA-Interim (Dee et al., 2011), and ERA5 (Hersbach et al., 2020) are commonly used,  
105 alongside the Australian Bureau of Meteorology's high-resolution (12 km) BARRA reanalysis.



106 BARRA covers Australia, New Zealand, and Southeast Asia (Su et al., 2019), while BARRA-  
107 C offers even higher-resolution (1.5 km) analysis for four capital cities (Su et al., 2021).

108 May et al. (2021) found BARRA reliable, though it showed seasonal and diurnal biases.  
109 Other studies, like Pirooz et al. (2021), compared BARRA with global reanalyses, concluding  
110 BARRA performed better for precipitation and temperature in New Zealand but lagged behind  
111 ERA5 for high gust winds. Du et al. (2023) used BARRA for estimating daily precipitation in  
112 ungauged Australian catchments, while Hobeichi et al. (2023) employed BARRA to train  
113 statistical models for downscaling. Acharya et al. (2019, 2020) found BARRA's precipitation  
114 performance varied by region, with poorer results in tropical areas. Nishant et al. (2022)  
115 suggested higher resolution in BARRA-C didn't always improve precipitation simulations,  
116 while Choudhury et al. (2023) noted ERA5 performed better for mean temperatures than  
117 extremes in Australia. These previous studies on BARRA and BARRA-C have also been  
118 summarized in Table S1.

119 However, there is a gap in the existing studies concerning the intercomparison of  
120 various reanalyses, such as BARRA and ERA5, specifically in relation to precipitation  
121 extremes over Australia. In this study, we aim to bridge this gap by evaluating and comparing  
122 the performance of BARRA and ERA5 in capturing precipitation extremes. While the  
123 traditional evaluation methods focusing on climatology (long-term mean), here we also include  
124 temporal correlation, coefficient of variation and trend in evaluation to quantify their overall  
125 performance, which have not been examined before in previous studies. By assessing climate  
126 means and extremes and quantifying their biases, this study provides a valuable reference for  
127 selecting appropriate datasets for specific applications and cautions against treating reanalysis  
128 data as observations. The paper is organized as follows: Section 2 introduces the reanalysis  
129 datasets and observational data used for evaluation. Section 3 outlines the climate extreme



130 indices and evaluation methodology. Results are presented in Section 4, followed by further  
131 discussion in Section 5. Finally, Section 6 offers a summary and conclusions.

132

## 133 **2. Data**

### 134 **2.1 ERA5**

135 ERA5 is a global atmospheric reanalysis dataset developed by ECMWF (Hersbach, et  
136 al. 2020). ERA5 provides hourly estimates of many atmospheric, land, and oceanic climate  
137 variables. The data is on a ~30 km horizontal grid and resolves the atmosphere using 137 levels  
138 from the surface up to a height of 0.01hPa (~80 km).

139 ERA5 is constructed upon the foundation of the Integrated Forecasting System (IFS)  
140 Cy41r2. This allows ERA5 to benefit from a decade's worth of development in areas such as  
141 model physics, core dynamics, and data assimilation techniques. ERA5 is a significant  
142 advancement over its predecessors (e.g., ERA-Interim) due to its higher spatial and temporal  
143 resolution, improved assimilation techniques, and more sophisticated modelling components.  
144 It provides a detailed and accurate representation of various atmospheric variables, such as  
145 temperature, humidity, wind speed, pressure, and more. The dataset covers the entire globe and  
146 spans from 1940 to the present, making it valuable for various applications in climate research,  
147 meteorology, environmental science, and more.

### 148 **2.2 BARRA**

149 BARRA is a high-resolution regional atmospheric reanalysis dataset developed by the  
150 Australian Bureau of Meteorology, which is available from January 1990 to February 2019 (Su,  
151 et al. 2019). BARRA was constructed based on the Australian Community Climate Earth-  
152 System Simulator (ACCESS) model with assimilation of a wide range of observational data to  
153 create a coherent and consistent representation of past weather and climate conditions. BARRA  
154 covers the Australian continent, New Zealand, part of Asia and some Pacific Islands with a



155 horizontal resolution of 12 km and 70 vertical levels from the surface up to a height of 80 km.  
156 BARRA specifically focuses on providing detailed information about weather patterns and  
157 atmospheric variables over the Australian region, which provides about 100 parameters at  
158 hourly intervals.

### 159 **2.3 AGCD**

160 The observational data in the study are from the Australian Gridded Climate Dataset  
161 (AGCD, Evans et al. 2020). The daily gridded maximum and minimum temperatures, and  
162 precipitation data has a spatial resolution of  $0.05^\circ$  ( $\sim 5\text{km}$ ) and is interpolated from observations  
163 at stations across the Australian continent. Most of those stations are in the more heavily  
164 populated coastal regions with far fewer stations inland and over high elevation areas. For  
165 example, there are very few station observations near the Gibson desert region in Western  
166 Australia, making the gridded observations unreliable over that region. Thus, in the following  
167 figures that region has been masked and not considered for evaluation. Since observations and  
168 reanalyses are not at the same spatial resolutions, we aggregate the observations to the native  
169 grid of ERA5 and BARRA respectively for comparison, including the performance of  
170 statistical significance tests. For comparison purpose, we also interpolate reanalysis to AGCD  
171 grids using the conservative area weighted re-gridding scheme from the Climate Data  
172 Operators (Schulzweida et al., 2006), which will be shown in the Supplementary Information.  
173 The states and sub-regions in the Australian region we discuss in the following can be found  
174 in Figure S1.

175

## 176 **3. Methodology**

### 177 **3.1 ET-SCI**

178 While extreme climate and weather events are generally multifaceted phenomena, in  
179 this study we evaluate climate extremes based on daily precipitation and temperature as defined



180 by Expert Team on Sector-specific Climate Indices (ET-SCI; Alexander & Herold, 2015;  
181 Herold and Alexander, 2016). We use the ClimPACT version 2 software to calculate the ET-  
182 SCI indices (<https://climpact-sci.org/>), focussing on daily precipitation.

183         Although ClimPACT generates 14 precipitation-related core indices, we select six  
184 (Table 1) based on the following considerations: 1) To capture key aspects of climate extremes,  
185 we include absolute indices such as the maximum 1 day precipitation (Rx1day) and total  
186 precipitation (PRCPTOT), threshold-based indices (e.g., number of heavy rain days, R10mm),  
187 percentile indices (e.g., total annual precipitation from very heavy rain days, R99p), and  
188 duration indices such as the consecutive wet (CWD) and dry days (CDD). 2) to capture  
189 extremes which have an impact on society and infrastructure, such as Rx1day, CDD, and CWD,  
190 which significantly affect agriculture, water resources and the economy (Tabari, 2020; Pei et  
191 al., 2021).

192         With the above consideration, six precipitation-related indices were calculated on  
193 native reanalysis grids and observation grids. Since the AGCD observations have the highest  
194 resolution, here we mainly show the evaluation on the native grids of the reanalyses (i.e., the  
195 12-km grid of BARRA and 30-km grid of ERA5). The extreme indices calculated from  
196 reanalysis data have also been regridded to the 5-km resolution, which are included in the  
197 supplementary information to demonstrate that our conclusions are insensitive to the choice of  
198 evaluation resolution.

### 199 **3.2 Evaluation matrices**

200         We evaluate BARRA and ERA5 for their performance in capturing climatology,  
201 coefficient of variation (CV), temporal correlation, and trends of six selected precipitation  
202 extreme indices. The CV is a valuable statistical tool representing the ratio of the standard  
203 deviation to the mean, allowing for the comparison of variation between different data series,  
204 even when their means differ significantly. Temporal correlations of climate extremes measure



205 the similarities between simulations and observations in terms of their inter-annual variabilities,  
206 with larger temporal correlations indicating better performance.

207 We use bias and domain-averaged absolute bias to quantify spatial differences between  
208 reanalyses and observations. Temporal correlation, coefficient of variation, and trend are used  
209 to quantify temporal similarities between reanalyses and observations. The non-parametric  
210 Mann-Kendall test is used to assess the statistical significance of differences and trends. Biases  
211 are assessed at an annual timescale for all extremes.

212

## 213 **4. Results**

### 214 **4.1 Mean climate**

215 This section evaluates and compares the annual mean of daily precipitation between  
216 BARRA and ERA5 against AGCD over Australia.

#### 217 **4.1.1 Bias and temporal correlation**

218 We first evaluate precipitation simulated by BARRA and ERA5 against observations  
219 (AGCD). The mean annual precipitation from the three datasets and biases in BARRA and  
220 ERA5 compared to AGCD are shown in Figure 1 (and Figure S2 on the observation grid).  
221 Results show that both BARRA and ERA5 simulate the spatial patterns of mean annual  
222 precipitation very well with high rainfall in northern Australian, eastern Australia coast and  
223 western Tasmania and low rainfall inland, albeit with clear biases. Compared to AGCD, both  
224 BARRA and ERA5 underestimate precipitation up to 20% for Eastern Australian coast,  
225 southwest western Australia, and western Tasmania, but overestimate annual precipitation up  
226 to 30% inland (Figure S3). Some clear differences in biases between BARRA and ERA5 can  
227 be observed in central western Australia and northern Queensland where BARRA overestimate  
228 precipitation but ERA5 underestimate it. Domain averaged absolute bias in annual precipitation  
229 is about 0.17mm/day (~12.7%) for BARRA and 0.15 mm/day (~10.5%) for ERA5 (Table 2).



230 The skill of simulated precipitation from BARRA and ERA5 are further demonstrated  
231 in the temporal correlations between BARRA/ERA5 and AGCD shown in Figure 2 (and Figure  
232 S4 on the observation grid). Temporal correlation of annual precipitation is larger in southeast  
233 Australia and northern Tasmania for both BARRA and ERA5, which is above 0.85. This  
234 indicates inter-annual variability of precipitation is well captured by BARRA and ERA5. In  
235 contrast, temporal correlation is weaker for western inland and northern Australia. ERA5  
236 generally has larger temporal correlation when compared with BARRA, especially for northern  
237 Australia, where temporal correlation for BARRA is below 0.5. On average, temporal  
238 correlation for ERA5 is 0.85, which is large than 0.77 for BARRA (Table 2).

#### 239 **4.1.2 CV (coefficient of variation) and trend**

240 CV of annual precipitation for AGCD and biases between BARRA/ERA5 and AGCD  
241 are presented in Figure 3 (and Figure S5 on the observation grid). By its definition, CV helps  
242 capture the standard deviation in the dataset relative to its mean. In the observation, CV is  
243 generally smaller for coastal regions except for northwest West Australia and Tasmania than  
244 inland Australia, where annual rainfall is much smaller than coastal regions. Alternatively,  
245 regions with higher annual precipitation generally have smaller CV. Both BARRA and ERA5  
246 reasonably capture the main feature of CV in observation. However, clear biases can be  
247 observed, especially in BARRA that has more than 50% large positive biases in Northern  
248 Australia, up to 20% positive biases for inland, and relatively smaller biases for southeastern  
249 Australia, southwest West Australia and Tasmania. In contrast, ERA5 does not have a clear  
250 bias pattern and biases are relatively smaller when compared to BARRA.

251 To further investigate the variability evident in observations and BARRA/ERA5  
252 simulations, we assess the trends in annual precipitation (Figure 4 and Figure S6 on the  
253 observation grid). AGCD shows strong increasing trends over Northern Australia and  
254 Northeast Australia coastal regions but decreasing trends over Northern Queensland,



255 southwestern West Australia and southern Great Dividing Range including Victoria, although  
256 not all trends are significant. Most of inland regions have relatively small trend in annual  
257 precipitation. Both BARRA and ERA5 reproduce the major trend pattern reasonably well,  
258 however, clear biases can be observed over Northern Australia where both BARRA and ERA5  
259 underestimate biases more than 100%. BARRA overestimated decreasing trend over Northern  
260 Queensland but ERA5 underestimate it (even increasing trend instead).

261 In summary, evaluation of annual mean precipitation indicates both BARRA and ERA5  
262 possess small biases (~20%) in the spatial precipitation patterns. ERA5 shows stronger  
263 temporal correlations than BARRA, particularly in northern Australia. Overall, ERA5  
264 demonstrates higher accuracy in capturing inter-annual precipitation variability. Both BARRA  
265 and ERA5 captured spatial distribution of coefficient of variation reasonably well but with  
266 large biases (~ 50%). BARRA shows much larger biases than ERA5 especially for Northern  
267 Australia. Both BARRA and ERA5 roughly reproduce the pattern of trend but with very large  
268 biases (~100%), especially for Northern Australia where both substantially underestimate the  
269 increasing trend.

270

## 271 **4.2 Climate extremes**

272 This section evaluates the six select precipitation extreme indices (Table 1) from  
273 BARRA and ERA5 over Australia by comparing them against AGCD. Evaluations are  
274 performed primarily using spatial bias maps and temporal correlations. We also assess the  
275 interannual variability and trends in the simulated BARRA and ERA5 indices and compare  
276 these with AGCD to further investigate any discrepancies.

### 277 **4.2.1 Bias and temporal correlation**

278 Annual mean biases in the six precipitation extremes are shown in Figure 5 (and Figure  
279 S8 on the observation grid). For duration-related extremes, there is a clear north-to-south



280 gradient in AGCD (Figure S7) with longer duration of CDD and CWD in northern Australia  
281 than southern Australia (CWD also has a clear west-to-east gradient in Tasmania), which is  
282 well simulated in BARRA and ERA5 (Figure S7). While the spatial distributions are well  
283 captured, clear biases are evident in them (Figure 5). BARRA generally underestimates CDD  
284 especially for central inland and northwest West Australia where biases are up to 40%. ERA5  
285 also under-estimates CDD for central inland, but in contrast its over-estimates CDD for most  
286 of northwestern Australia, overall ERA5 has smaller absolute bias in CDD (6.9 days) than  
287 BARRA (14.5 days) (Table 2). Both BARRA and ERA5 have similar bias pattern for CWD,  
288 which generally overestimate CWD over most of regions except for southern Australian coast,  
289 southwest West Australia and western Tasmania. The positive biases over Northern Australia  
290 may reach 30%. Overall BARRA has slightly larger biases in CWD (2.3 days) than ERA5 (1.7  
291 days) (Table 2).

292 Both BARRA and ERA5 also generally match the spatial distribution of heavy rainfall  
293 days and R90p (Figure S7) in AGCD with large values in Northern Australia, eastern seaboard  
294 and Australian Great Dividing Range, and western Tasmania. However, clear biases can be  
295 observed in BARRA and ERA5 for both R10mm and R90p (Figure 5). BARRA and ERA5  
296 have large negative biases in R10mm over Northern Australia, eastern seaboard, southwest  
297 Western Australia and western Tasmania, but biases in central inland and northwest West  
298 Australia are generally small. Overall, domain averaged absolute bias for ERA5 (1.7 days) is  
299 about half of that for BARRA (3.3 days). Both BARRA and ERA5 also have relatively large  
300 negative biases in R90p for most of northern Australia, eastern coasts, southwest West  
301 Australia and western Tasmania but small positive biases inland, especially for BARRA.  
302 Overall averaged absolute bias is 0.78 mm/day for BARRA and 0.44 mm/day for ERA5 (Table  
303 2).



304 BARRA and ERA5 also reasonably captured the spatial patterns of R99p and Rx1day,  
305 however, quite large biases are in BARRA and ERA5 (Figure 5). BARRA generally  
306 overestimate R99p and Rx1day over northern Australia coasts and along the Great Dividing  
307 Range, in contrast, ERA5 generally underestimate R99p and Rx1day over northern and eastern  
308 coasts, southwest Western Australia and western Tasmania. The domain averaged bias in R99p  
309 is at similar magnitude for BARRA (4.09 mm/day) and ERA5 (3.67 mm/day), however biases  
310 in Rx1day is much larger for BARRA (20.3 mm/day) than ERA5 (7.9 mm/day) (Table 2).

311 Figure 6 (and Figure S9 on the observation grid) presents the temporal correlations  
312 between BARRA/ERA5 and AGCD for the six precipitation extreme indices. Unlike the strong  
313 temporal correlation between BARRA/ERA5 and AGCD for mean annual precipitation (Figure  
314 2), the temporal correlations for these extreme indices are worse except for R90p (Figure 6).  
315 For extremes like R10mm and R90p, the correlation ranges from reasonably good (above 0.6)  
316 to pretty good (above 0.8) between BARRA/ERA5 and AGCD for most of the domain.  
317 Temporal correlation for CWD and R99p are not as good as R10mm and R99p, but they are  
318 comparatively stronger correlations (0.5-0.6) than CWD and Rx1day (~0.5 and less) over most  
319 of the domain. Compared to BARRA, ERA5 has slightly stronger temporal correlations for  
320 those extremes (Table 2).

321

#### 322 4.2.2 CV (coefficient of variation) and trend

323 The observed and simulated CV of precipitation extremes and biases in their CV for  
324 BARRA and ERA5 are shown in Figure S10 and Figure 7 (and Figure S11 on the observation  
325 grid), respectively. Generally, both BARRA and ERA5 have similar CV bias patterns and  
326 magnitude for CDD, CWD and R10mm. In contrast, BARRA is quite different from ERA5 for  
327 other three extremes. BARRA substantially under-estimated CV of R90p over most on inland  
328 regions but ERA5 has much smaller negative biases, even small positive biases, although both



329 have small biases in CV of R90p along most coastal regions and Tasmania. BARRA  
330 systematically overestimate CVs of R99p and Rx1day over northern Australia but ERA5 has  
331 relatively small biases for them. Overall, BARRA has more than twice as much as CV biases  
332 in ERA5 for R90p, R99p and Rx1day (Table 2).

333 Trends of each of the precipitation extreme indices for the three datasets and biases in  
334 trend for BARRA and ERA5 are shown in Figure S12 and Figure 8 (and Figure S13 on the  
335 observation grid), respectively. Generally, both BARRA and ERA5 simulate the main pattern  
336 of trends for those extremes but with large biases. BARRA and ERA5 simulated CDD trend  
337 well for southern Australia but BARRA generally under-estimated trend in CDD over inland  
338 Australia and overestimate trend in northwest Australia. ERA5 only has large positive trend  
339 biases in northern central Australia. The overall domain averaged biases are similar between  
340 BARRA (0.584) and ERA5 (0.566). Both BARRA and ERA5 have small biases in CWD in  
341 central and southern Australia but similar biases pattern in Northern Australia. They also have  
342 similar overall biases in CWD (0.064 for BARRA and 0.060 for ERA5). Both BARRA and  
343 ERA5 under-estimated increasing trend in R10mm in northern Australia, but BARRA  
344 overestimate trend in most of southeast Australia. In contrast, ERA5 under-estimate trend over  
345 there. Overall, ERA5 has slightly larger biases (0.094) than BARRA (0.085). Like R10mm,  
346 both BARRA and ERA5 also underestimate trend of R90p in most of northern Australia but  
347 have small biases in central and southern Australia. They have almost the same overall biases  
348 in R90p. BARRA/ERA5 has similar biases patterns for R99p and rx1day but biases for rx1days  
349 are much larger. Both BARRA and ERA5 have large biases in R99p and Rx1day but biases in  
350 BARRA are generally larger than ERA5.

351 In summary, both BARRA and ERA5 reproduce spatial patterns of extremes well but  
352 display biases. ERA5 underestimates CDD and certain heavy rainfall events, while BARRA  
353 tends to overestimate these extremes. Both reanalyses show discrepancies in various



354 precipitation indices across different regions, with BARRA generally displaying larger biases  
355 compared to ERA5. Temporal correlations between BARRA/ERA5 and observations for  
356 extreme precipitation indices are weaker than those for mean annual precipitation, except for a  
357 few indices where ERA5 demonstrates slightly stronger correlations compared to BARRA.  
358 Both BARRA and ERA5 align in CV patterns and biases for certain extremes but differ notably  
359 in others. BARRA significantly underestimates very heavy precipitation variability over inland  
360 regions, while ERA5 presents smaller biases or even positive biases in these areas. Additionally,  
361 BARRA tends to overestimate extreme precipitation variability in Northern Australia  
362 compared to ERA5. Overall, BARRA shows more than double the biases in variability  
363 compared to ERA5 for specific extreme precipitation indices. Both reanalyses generally  
364 simulate the main trend patterns but exhibit considerable biases. BARRA underestimates or  
365 overestimates trends in certain regions and indices, while ERA5 demonstrates different biases,  
366 including smaller biases overall compared to BARRA across these precipitation extremes.

367

## 368 **5. Discussion**

369 In this study, we assessed the performance of BARRA and ERA5 in simulating mean  
370 precipitation and six selected precipitation extremes. While most previous evaluations have  
371 focused on the climatology of precipitation and its extremes, only a few studies have included  
372 the coefficient of variation (CV) (Teng et al., 2024). Our evaluation encompassed annual  
373 climatology, along with temporal correlation, CV, and trend analysis, providing a  
374 comprehensive assessment of the performance of these two reanalysis datasets.

375 The results indicate that both BARRA and ERA5 demonstrate reasonable skill in  
376 simulating mean precipitation and certain precipitation extremes. However, they encounter  
377 challenges in accurately reproducing temporal correlation, CV, and trends for certain extreme  
378 events, highlighting significant uncertainties in their representation of extremes.



379           While acknowledging the capabilities of these reanalysis datasets, our study also  
380 identifies specific limitations and suggests potential directions for future research. A crucial  
381 consideration in model evaluation is the accuracy of observational data, which substantially  
382 influences evaluation outcomes. In this study, we used the AGCD dataset as the observational  
383 benchmark, which is based on interpolating data from in-situ stations (Evans et al., 2020).  
384 However, the AGCD dataset presents several limitations: 1) Spatial coverage: Sparse station  
385 coverage in northwest and central Australia, and limited observations in high-elevation areas,  
386 result in a concentration of stations in southeastern Australia, southwestern Western Australia,  
387 and eastern Tasmania. The arid interior is notably underrepresented. 2) Data completeness and  
388 homogeneity: Incomplete and inhomogeneous observations due to missing data, changes in  
389 observational techniques, or station relocations can affect the consistency of the dataset. 3)  
390 Interpolation uncertainties: The interpolation method used in AGCD (splining), instead of the  
391 ordinary kriging method used in its predecessor (AWAP), introduces uncertainties, particularly  
392 in areas with sparse data coverage for extreme events like heavy rainfall.

393           These observational uncertainties may contribute to biases in the evaluation results. In  
394 particular, the limited number of monitoring sites over the Great Dividing Range and inland  
395 areas introduces significant uncertainties in estimated observed precipitation for these regions.  
396 Independent studies, such as Chubb et al. (2016), found that daily precipitation is  
397 underestimated by at least 15% in some areas, which could suggest similar underestimation in  
398 BARRA and ERA5 for these regions. Similarly, the sparse gauge network in northwestern  
399 inland areas might miss localized extreme precipitation events.

400           Our analysis focused on six ET-SCI-defined precipitation extreme indices, widely used  
401 in various evaluation studies (Nishant et al., 2020; Ji et al., 2024). However, recognizing the  
402 need for region-specific indices, we suggest future studies extend the analysis to incorporate  
403 additional extreme indices tailored to specific regions and applications.



404 Our findings emphasize that while both BARRA and ERA5 are competent in simulating  
405 the climatology of mean climate, temporal correlation, and CV, challenges remain in accurately  
406 capturing trends, particularly for certain extremes. Notably, ERA5 shows better overall  
407 performance compared to BARRA. Although higher resolution often correlates with better  
408 performance, recent studies have shown that increasing resolution alone does not always  
409 guarantee improvements (Nishant et al., 2022). Considering the critical role of driving data,  
410 model physics, and data assimilation, it may be valuable to update BARRA using the latest  
411 ERA5 data along with improved model physics and data assimilation techniques to enhance its  
412 performance.

413 In this study, we evaluated ERA5 and BARRA on both their native resolutions and a  
414 common resolution (5 km) to match AGCD. The results showed that the evaluations were  
415 consistent across native and common resolutions, suggesting that the performance assessments  
416 were not highly sensitive to changes in resolution.

417

## 418 **6. Summary and Conclusion**

419 Reanalysis datasets play a crucial role in climate research, weather analysis, and various  
420 scientific investigations. Their ability to provide a consistent and comprehensive representation  
421 of historical weather and climate conditions makes them invaluable. These datasets are  
422 particularly essential for studying long-term climate trends, understanding atmospheric  
423 phenomena, and validating climate models.

424 In this study, we evaluate BARRA and ERA5 for their capabilities to simulate mean  
425 precipitation and six selected precipitation extremes for their climatology, temporal correlation,  
426 coefficient of variation (CV) and trend to quantify their overall performance. We evaluated  
427 BARRA and ERA5 at their native resolutions, as well as at a common resolution (i.e., the



428 observation resolution). Both analyses yielded consistent results, indicating that the evaluation  
429 is not sensitive to the remapping process.

430 The assessment of annual mean precipitation reveals that both BARRA and ERA5  
431 adeptly reproduce the spatial precipitation patterns, exhibiting minor biases of around 20%.  
432 Particularly, ERA5 showcases stronger temporal correlations compared to BARRA, especially  
433 evident in northern Australia. ERA5, overall, demonstrates superior accuracy in capturing  
434 inter-annual precipitation variability. However, both models depict the spatial distribution of  
435 the coefficient of variation reasonably well but with larger biases, roughly around 50%.  
436 Particularly, BARRA displays significantly higher biases, especially in Northern Australia.

437 Regarding the replication of trend patterns, both models exhibit substantial biases,  
438 reaching approximately 100%. This is especially notable in Northern Australia, where they  
439 both notably underestimate the increasing trend. Furthermore, while both BARRA and ERA5  
440 possess about the right spatial patterns of extremes, biases are evident. ERA5 tends to  
441 underestimate consecutive dry days (CDD) and certain heavy rainfall events, while BARRA  
442 tends to overestimate these extremes. Discrepancies in various precipitation indices across  
443 regions are apparent, with BARRA generally displaying larger biases compared to ERA5.

444 When examining temporal correlations for extreme precipitation indices compared to  
445 mean annual precipitation, both BARRA and ERA5 show weaker correlations, except for a  
446 few indices where ERA5 slightly outperforms BARRA. While both models align in coefficient  
447 of variation patterns and biases for certain extremes, notable differences arise in others.  
448 BARRA notably underestimates very heavy precipitation variability over inland regions,  
449 whereas ERA5 presents smaller biases or even positive biases in these areas. Moreover,  
450 BARRA tends to overestimate extreme precipitation variability in Northern Australia  
451 compared to ERA5. Specifically, BARRA showcases more than double the biases in variability  
452 compared to ERA5 for specific extreme precipitation indices.



453 In terms of trend patterns, both models generally replicate the observed trends but  
454 exhibit considerable biases. BARRA shows both underestimations and overestimations in  
455 certain regions and indices, while ERA5 displays different biases, including overall smaller  
456 biases compared to BARRA across these precipitation extremes.

457 In summary, our findings suggest that both ERA5 and BARRA are reliable for  
458 climatological analyses, including mean precipitation and precipitation extremes, and can be  
459 confidently used by end-users for such purposes. However, as discussed in the introduction,  
460 caution is advised when using these datasets for variability and trend analyses, particularly for  
461 specific extreme events like Rx1day. The performance of these reanalyses is regionally  
462 dependent, and this should be considered when using them as observational references for  
463 evaluating other model simulations. Additionally, the biases in the variability and trends of  
464 climate extremes present in both datasets must be carefully accounted for when comparing  
465 them with other data sources.

466

467

#### 468 **Data Availability**

469 Details about AGCD are available at the Australian Bureau of Meteorology website  
470 (<http://www.bom.gov.au/metadata/catalogue/19115/ANZCW0503900567>, (accessed on)).  
471 The dataset is available on the NCI (National Computational Infrastructure) server in project  
472 zv2. Detail on how to access the data can be found at [http://climate-](http://climate-cms.wikis.unsw.edu.au/AGCD)  
473 [cms.wikis.unsw.edu.au/AGCD](http://climate-cms.wikis.unsw.edu.au/AGCD), (accessed on). ERA5 data is available on the NCI in Project  
474 rt52. BARRA data is available on the NCI in project cj37.

#### 475 **Author Contributions**

476 KKWC and FJ conceptualized and implemented the research. KKWC, FJ and NN performed  
477 the data analysis and prepared the figures. FJ prepared the draft manuscript. All authors  
478 contributed to the discussion of results, editing and finalization of the manuscript.

#### 479 **Competing Interests**

480 The authors declare that they have no conflict of interest.

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486



487 **References**

- 488 Acharya, S. C., Nathan, R., Wang, Q. J., et al.: An evaluation of daily precipitation from a  
489 regional atmospheric reanalysis over Australia. *Hydrol. Earth Syst. Sci.*, 23, 3387–3403,  
490 <https://doi.org/10.5194/hess-23-3387-2019> , 2019.
- 491 Acharya, S. C., Nathan, R., Wang, Q. J., et al.: Ability of an Australian reanalysis dataset to  
492 characterise sub-daily precipitation. *Hydrol. Earth Syst. Sci.*, 24, 2951–2962,  
493 <https://doi.org/10.5194/hess-24-2951-2020> , 2020.
- 494 Betts, A. K., Chan, D. Z., Desjardins, R. L.: Near-Surface Biases in ERA5 Over the Canadian  
495 Prairies. *Front. Environ. Sci.*, 7, <https://doi.org/10.3389/fenvs.2019.00129> , 2019.
- 496 Capecchi, V., Pasi, F., Gozzini, B., Brandini, C.: A convection-permitting and limited-area  
497 model hindcast driven by ERA5 data: precipitation performances in Italy. *Clim. Dyn.*,  
498 61, 1411–1437, <https://doi.org/10.1007/s00382-022-06633-2> , 2023.
- 499 Cheung, K. K. W., Ji, F., Nishant, N., Herold, N., Cook, K.: Evaluation of Convective  
500 Environments in the NARCLiM Regional Climate Modeling System for Australia.  
501 *Atmosphere*, 14, 690, <https://doi.org/10.3390/atmos14040690> , 2023.
- 502 Choudhury, D., Ji, F., Nishant, N., Di Virgilio, G.: Evaluation of ERA5 simulated  
503 temperature and its extremes for Australia. *Atmosphere*, 14(6),  
504 913. <https://doi.org/10.3390/atmos14060913>, 2023.
- 505 Chubb, T. H., Manton, M. J., Siems, S. T., Peace, A. D.: Evaluation of the AWAP daily  
506 precipitation spatial analysis with an independent gauge network in the Snowy  
507 Mountains. *J. Southern Hemisphere Earth Systems Science*, 66, 55-67, 2016.
- 508 Dai, W., Zeng, Y., Jing, T. G., et al.: Estimation of rainfall erosivity on the Chinese Loess  
509 Plateau: A new combination of the ERA5 dataset and machine learning. *J Hydrol*  
510 (Amst), 624, <https://doi.org/10.1016/j.jhydrol.2023.129892>, 2023.
- 511 Dee, D. P., Uppala, S. M., Simmons, A. J., Berrisford, P., Poli, P., Kobayashi, S., Andrae, U.,  
512 Balmaseda, M. A., Balsamo, G., Bauer, P., et al.: The ERA-Interim reanalysis:  
513 Configuration and performance of the data assimilation system. *Q. J. R. Meteorol. Soc.*,  
514 137, 553–597, <https://doi.org/10.1002/qj.828>, 2011.
- 515 Di Virgilio G., Evans J.P., Di Luca A., Olson R., Argueso D., Kala J., Andrys J., Hoffmann  
516 P., Katzfey J.J., Rockel B.: Evaluating reanalysis driven CORDEX regional climate  
517 models over Australia: model performance and errors. *Climate Dynamics*, 53, 2985-  
518 3005, 2019.
- 519 Di Virgilio, G., Ji, F., Tam, E., Evans, J., Kala, J., Andrys, J., Thomas, C., Choudhury, D.,  
520 Rocha, C., Li, Y., Riley, M.: Evaluation of CORDEX ERA5-forced ‘NARCLiM2.0’  
521 regional climate models over Australia using the Weather Research and Forecasting  
522 (WRF) model version 4.1.2, *Geosci. Model Dev. Discuss.*, <https://doi.org/10.5194/gmd-2024-41> , 2024.
- 524 Du, Y. L., Wang, Q. J., Su, C. H., et al.: Estimating daily precipitation climatology by  
525 postprocessing high-resolution reanalysis data. *Intl. J. Climatol.*, 43, 4151–4165,  
526 <https://doi.org/10.1002/joc.8079>, 2023.



- 527 Evans, A., Jones, D., Smalley, R., Lellyett, S.: An Enhanced gridded rainfall analysis scheme  
528 for Australia. Bureau Research Report No. 41, Bureau of Meteorology, Australia, 2020.  
529 Available at <http://www.bom.gov.au/research/publications/researchreports/BRR-041.pdf>
- 530 Fita, L., Evans, J. P., Argüeso, D., et al.: Evaluation of the regional climate response in  
531 Australia to large-scale climate models in the historical NARCLiM simulations. *Climate*  
532 *Dynamics*, 1–15, <https://doi.org/10.1007/s00382-016-3484-x>, 2016.
- 533 Gleixner, S., Demissie, T., Diro, G. T.: Did ERA5 Improve Temperature and Precipitation  
534 Reanalysis over East Africa? *Atmosphere (Basel)*, 11,  
535 <https://doi.org/10.3390/atmos11090996>, 2020.
- 536 Herold, N., Alexander, L.: *Climpact 2*. <https://github.com/ARCCSS-extremes/climpact2>,  
537 2016.
- 538 Hersbach, H., Bell, B., Berrisford, P., Hirahara, S., Horányi, A., Muñoz-Sabater, J., Nicolas,  
539 J., Peubey, C., Radu, R., Schepers, D., et al.: The ERA5 global reanalysis. *Q. J. R.*  
540 *Meteorol. Soc.*, 146, 1999–2049, <https://doi.org/10.1002/qj.3803>, 2020.
- 541 Hobeichi, S., Nishant, N., Shao, Y. W., et al.: Using Machine Learning to Cut the Cost of  
542 Dynamical Downscaling. *Earths Future*, 11, <https://doi.org/10.1029/2022EF003291>,  
543 2023.
- 544 Hu, X. L., Yuan, W. H.: Evaluation of ERA5 precipitation over the eastern periphery of the  
545 Tibetan plateau from the perspective of regional rainfall events. *Intl. J. Climatol.* 41,  
546 2625–2637, <https://doi.org/10.1002/joc.6980>, 2021.
- 547 Izadi, N., Karakani, E. G., Saadatabadi, A. R., et al.: Evaluation of ERA5 Precipitation  
548 Accuracy Based on Various Time Scales over Iran during 2000-2018. *Water (Basel)*, 13,  
549 <https://doi.org/10.3390/w13182538>, 2021.
- 550 Ji, F., Evans, J. P., Teng, J., Scorgie, Y., Argüeso, D., Di Luca, A.: Evaluation of long-term  
551 precipitation and temperature Weather Research and Forecasting simulations for  
552 southeast Australia. *Climate Research* 67, 99-115, 2016.
- 553 Ji, F., Di Virgilio, G., Nishant, N., Tam, E., Evans, J. P., Kala, J., Andrys, J., Thomas, C.,  
554 Riley, M.: Evaluation of precipitation extremes in ERA5 driven regional climate  
555 simulations. *Weather and Climate Extremes*,  
556 <https://doi.org/10.1016/j.wace.2024.100676>, 2024.
- 557 Jiang, Q., Li, W. Y., Fan, Z. D., et al.: Evaluation of the ERA5 reanalysis precipitation  
558 dataset over Chinese Mainland. *J Hydrol (Amst)*, 595,  
559 <https://doi.org/10.1016/j.jhydrol.2020.125660>, 2021.
- 560 Jiao, D. L., Xu, N. N., Yang, F., Xu, K.: Evaluation of spatial-temporal variation performance  
561 of ERA5 precipitation data in China. *Sci. Rep.* 11. [https://doi.org/10.1038/s41598-021-](https://doi.org/10.1038/s41598-021-97432-y)  
562 [97432-y](https://doi.org/10.1038/s41598-021-97432-y), 2021.
- 563 Kalnay, E., Kanamitsu, M., Kistler, R., Collins, W., Deaven, D., Gandin, L., Iredell, M.,  
564 Saha, S., White, G., Woollen, J., et al.: The NCEP/NCAR 40-year reanalysis project.  
565 *Bull. Am. Meteorol. Soc.*, 77, 437–472. [https://doi.org/10.1175/1520-](https://doi.org/10.1175/1520-0477(1996)077<0437:TNYRP>2.0.CO;2)  
566 [0477\(1996\)077<0437:TNYRP>2.0.CO;2](https://doi.org/10.1175/1520-0477(1996)077<0437:TNYRP>2.0.CO;2), 1996.



- 567 Kobayashi, S., Ota, Y., Harada, Y., Ebita, A., Moriya, M., Onoda, H., Onogi, K., Kamahori,  
568 H., Kobayashi, C., Endo, H., et al.: The JRA-55 Reanalysis: General Specifications and  
569 Basic Characteristics. *J. Meteorol. Soc. Jpn. Ser. II*, 93, 5–48,  
570 <https://doi.org/10.2151/jmsj.2015-001>, 2015.
- 571 Lei, X. Y., Xu, W. L., Chen, S. T., et al.: How Well Does the ERA5 Reanalysis Capture the  
572 Extreme Climate Events Over China? Part I: Extreme Precipitation. *Front. Environ. Sci.*,  
573 10, <https://doi.org/10.3389/fenvs.2022.921658>, 2022.
- 574 Lei, Y. H., Letu, H. S., Shang, H. Z., Shi, J.C.: Cloud cover over the Tibetan Plateau and  
575 eastern China: a comparison of ERA5 and ERA-Interim with satellite observations.  
576 *Clim. Dyn.*, 54, 2941–2957. <https://doi.org/10.1007/s00382-020-05149-x>, 2020.
- 577 Li, X. X., Qin, X. C., Yang, J., Zhang, Y. Z.: Evaluation of ERA5, ERA-Interim, JRA55 and  
578 MERRA2 reanalysis precipitation datasets over the Poyang Lake Basin in China. *Intl. J.*  
579 *Climatol.* 42, 10435–10450. <https://doi.org/10.1002/joc.7915>, 2022.
- 580 May, P. T., Trewin, B., Su, C. H., Ostendorf, B.: Verification of moist surface variables over  
581 northern Australia in a high-resolution reanalysis (BARRA). *J. Southern Hemisphere*  
582 *Earth Systems Science* 71, 194–202. <https://doi.org/10.1071/ES21007>, 2021.
- 583 Nishant N., Evans, J. P., Di Virgilio, G., Downes, S. M., Ji, F., Cheung, K. K. W., Tam, E.,  
584 Miller, J., Beyer, K., Riley, M. L.: Introducing NARClIM1.5: evaluating the  
585 performance of regional climate projections for southeast Australia for 1950-2100.  
586 *Earth's Future*, <https://doi.org/10.1029/2020EF001833>, 2021.
- 587 Nishant, N., Sherwood, S., Prasad, A., Ji, F., Singh, A.: Impact of higher spatial resolution on  
588 precipitation properties over Australia, *Geophysical Research Letters*,  
589 DOI:[10.1029/2022GL100717](https://doi.org/10.1029/2022GL100717), 2022.
- 590 Pei, F., Zhou, Y., Xia, Y.: Assessing the Impacts of Extreme Precipitation Change on  
591 Vegetation Activity. *Agriculture*, 11(6), 487, 2021.
- 592 Pirooz, A. A. S., Moore, S., Carey-Smith, T., et al.: Evaluation of global and regional  
593 reanalyses performance over New Zealand. *Weather and Climate*, 41, 52–70.  
594 <https://doi.org/10.2307/27127989>, 2021.
- 595 Poli, P., Hersbach, H., Dee, D.P., Berrisford, P., Simmons, A.J., Vitart, F., Laloyaux, P., Tan,  
596 D.G.H., Peubey, C., Thépaut, J.-N., et al.: ERA-20C: An Atmospheric Reanalysis of the  
597 Twentieth Century. *J. Climate*, 29, 4083–4097, <https://doi.org/10.1175/jcli-d-15-0556.1>,  
598 2016.
- 599 Qin, S., Wang, K. C., Wu, G. C., Ma, Z. S.: Variability of hourly precipitation during the  
600 warm season over eastern China using gauge observations and ERA5. *Atmos. Res.*, 264,  
601 <https://doi.org/10.1016/j.atmosres.2021.105872>, 2021.
- 602 Quagraine, K. A., Nkrumah, F., Klein, C., et al.: West African Summer Monsoon  
603 Precipitation Variability as Represented by Reanalysis Datasets. *Climate*, 8,  
604 <https://doi.org/10.3390/cli8100111>, 2020.



- 605 Saha, S., Moorthi, S., Pan, H.-L., Wu, X., Wang, J., Nadiga, S., Tripp, P., Kistler, R.,  
606 Woollen, J., Behringer, D., et al.: The NCEP Climate Forecast System Reanalysis. *Bull.*  
607 *Am. Meteorol. Soc.*, 91, 1015–1058. <https://doi.org/10.1175/2010BAMS3001.1>, 2010.
- 608 Schulzweida, U., Kornblueh, L., Quast, R.: CDO user’s guide. *Climate data operators,*  
609 *Version, 1(6)*, 205-209, 2006..
- 610 Shen, L. C., Wen, J. H., Zhang, Y. Q., et al.: Performance Evaluation of ERA5 Extreme  
611 Precipitation in the Yangtze River Delta, China. *Atmosphere (Basel)* 13,  
612 <https://doi.org/10.3390/atmos13091416>, 2022.
- 613 Solman, S.A., Sanchez, E., Samuelsson, P., et al.: Evaluation of an ensemble of regional  
614 climate model simulations over South America driven by the ERA-Interim reanalysis:  
615 model performance and uncertainties. *Climate Dynamics*, 41, 1139–1157,  
616 <https://doi.org/10.1007/s00382-013-1667-2>, 2013.
- 617 Song, Y. Y., Wei, J. F.: Diurnal cycle of summer precipitation over the North China Plain  
618 and associated land-atmosphere interactions: Evaluation of ERA5 and MERRA-2. *Intl.*  
619 *J. Climatol.* 41, 6031–6046, <https://doi.org/10.1002/joc.7166> , 2021.
- 620 Su, C. H., Eizenberg, N., Jakob, D., et al.: BARRA v1.0: kilometre-scale downscaling of an  
621 Australian regional atmospheric reanalysis over four midlatitude domains. *Geosci.*  
622 *Model Dev.*, 14, 4357–4378, <https://doi.org/10.5194/gmd-14-4357-2021> , 2021.
- 623 Su, C. H., Eizenberg, N., Steinle, P., et al.: BARRA v1.0: the Bureau of Meteorology  
624 Atmospheric high-resolution Regional Reanalysis for Australia. *Geosci. Model Dev.*, 12,  
625 2049–2068, <https://doi.org/10.5194/gmd-12-2049-2019> , 2019.
- 626 Tabari, H.: Climate change impact on flood and extreme precipitation increases with water  
627 availability. *Scientific Reports*, 10(1), 1-10, 2020.
- 628 Teng, J., Bennett, J. C., Charles, S., Chiew, F., Ji, F., Potter, N., Fu, G. B., Thatcher, M.,  
629 Remenyi, T.: Trend and variance in regional climate models – validation and  
630 hydrological implications, *Journal of Hydrology*, 642(11), 131817,  
631 <https://doi.org/10.1016/j.jhydrol.2024.131817> , 2024.
- 632 Wang C. X., Graham R. M., Wang, K. G., et al.: Comparison of ERA5 and ERA-Interim  
633 near-surface air temperature, snowfall and precipitation over Arctic sea ice: effects on  
634 sea ice thermodynamics and evolution. *Cryosphere* 13, 1661–1679.  
635 <https://doi.org/10.5194/tc-13-1661-2019> , 2019.
- 636
- 637



Table 1 List of ET-SCI indices evaluated in this study.

Index	Definition	Units	Timescale	Sectors
<b>PRCPTOT</b>	Total wet-day precipitation (Sum of daily precipitation $\geq 1.0$ mm)	mm	Annual/Monthly	Agriculture and food security, water, water resources and food security, forestry/GHGs
<b>CDD</b>	Consecutive dry days (Maximum number of consecutive dry days (when precipitation $< 1.0$ mm))	days	Annual	Health, agriculture and food security, water resources and food security, disaster risk reduction, forestry/GHGs
<b>CWD</b>	Consecutive wet days (Maximum annual number of consecutive wet days (when precipitation $\geq 1.0$ mm))	days	Annual	Coasts, agriculture, transport operations
<b>R10mm</b>	Days when precipitation is at least 10mm	days	Annual/Monthly	Coasts
<b>R90p</b>	Total annual precipitation from very heavy precipitation days (Annual sum of daily precipitation $> 90$ th percentile)	mm	Annual	Coasts, transport operations
<b>R99p</b>	Total annual precipitation from very heavy precipitation days (Annual sum of daily precipitation $> 99$ th percentile)	mm	Annual	Coasts, transport operations
<b>Rx1Day</b>	Amount of precipitation from very wet days (Maximum 1-day precipitation)	mm	Annual/Monthly	Agriculture and food security, water, coasts, disaster risk reduction, forestry/GHGs

Table 2 Domain-averaged absolute biases and temporal correlation between BARRA/ERA5 and AGCD for annual precipitation and precipitation extremes

Indices	Absolute biases in annual mean		Temporal correlation		Absolute biases in CV		Absolute biases in trend	
	BARRA	ERA5	BARRA	ERA5	BARRA	ERA5	BARRA	ERA5
<b>Annual pr</b>	0.169	0.149	0.771	0.854	0.063	0.037	0.008	0.007
<b>CDD</b>	14.543	6.913	0.578	0.650	0.050	0.045	0.584	0.566
<b>CWD</b>	2.346	1.714	0.446	0.527	0.061	0.059	0.064	0.060
<b>R10mm</b>	3.265	1.700	0.688	0.761	0.081	0.053	0.085	0.094
<b>R90p</b>	0.777	0.439	0.761	0.827	0.211	0.082	0.023	0.023
<b>R99p</b>	4.093	3.668	0.562	0.625	0.121	0.060	0.206	0.162
<b>Rx1day</b>	20.333	7.916	0.380	0.486	0.219	0.107	0.848	0.542

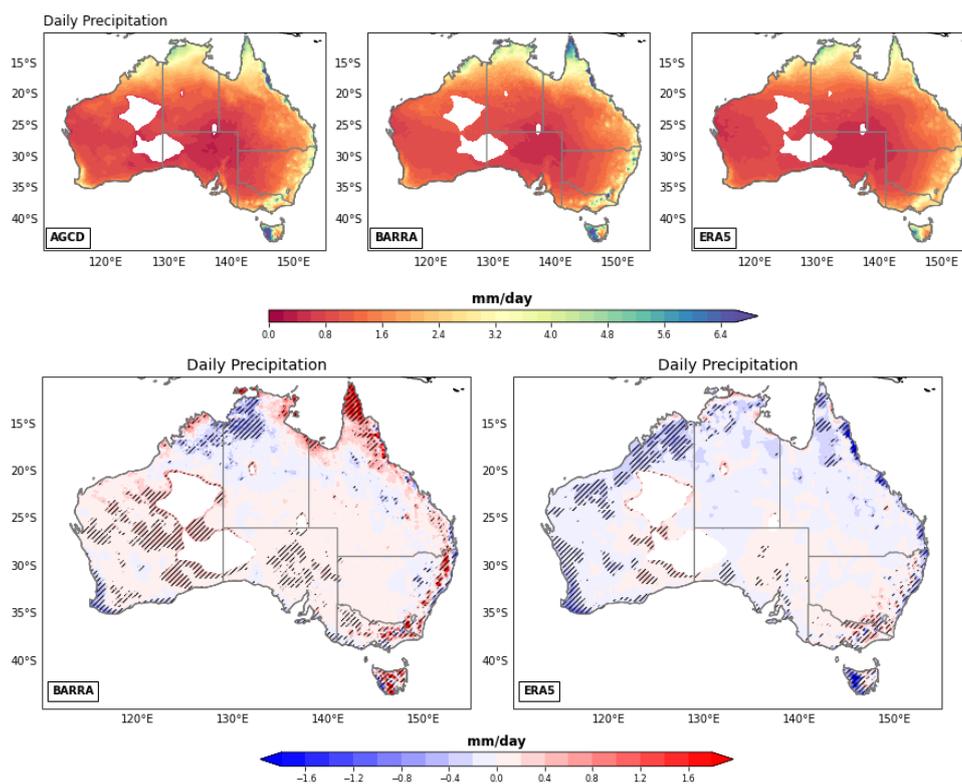


Figure 1 Annual mean precipitation of AGCD, BARRA and ERA5 (upper panels) and annual mean biases between BARRA/ERA5 and AGCD (lower panels). The regions with low density of station observations in AGCD has been masked and not considered in all subsequent evaluation. Unit: mm/day. Stippling indicates areas with biases that are statistically significant at 95% confidence level.

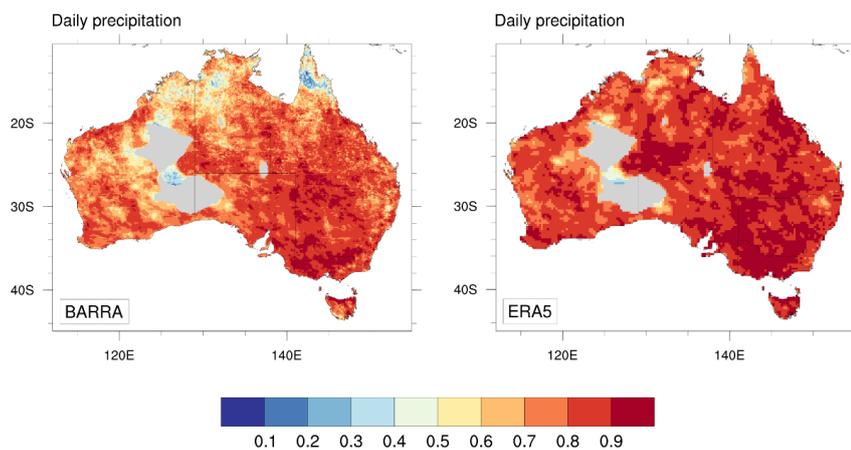


Figure 2 Temporal correlation coefficient of annual precipitation between BARRA/ERA5 and AGCD.

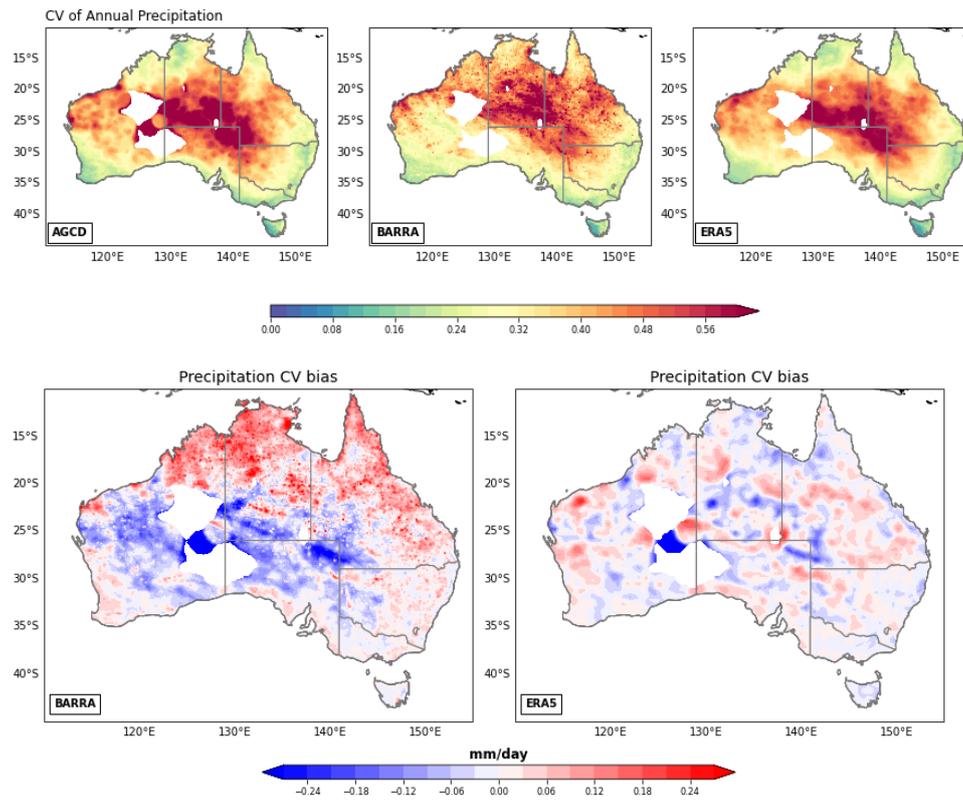


Figure 3 CV of annual precipitation for AGCD, BARRA and ERA5 (upper panels) and biases in CV between BARRA/ERA5 and AGCD (lower panels).

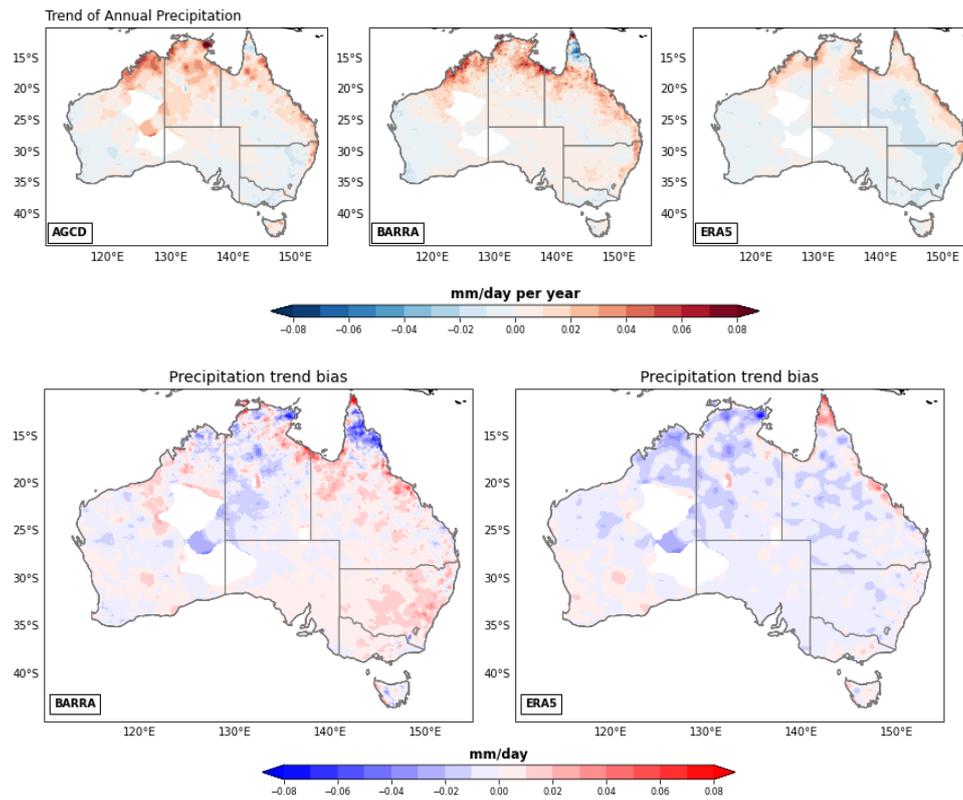


Figure 4 Trend of annual precipitation for AGCD, BARRA and ERA5 (upper panels) and biases in trend between BARRA/EAR5 and AGCD (lower panels).

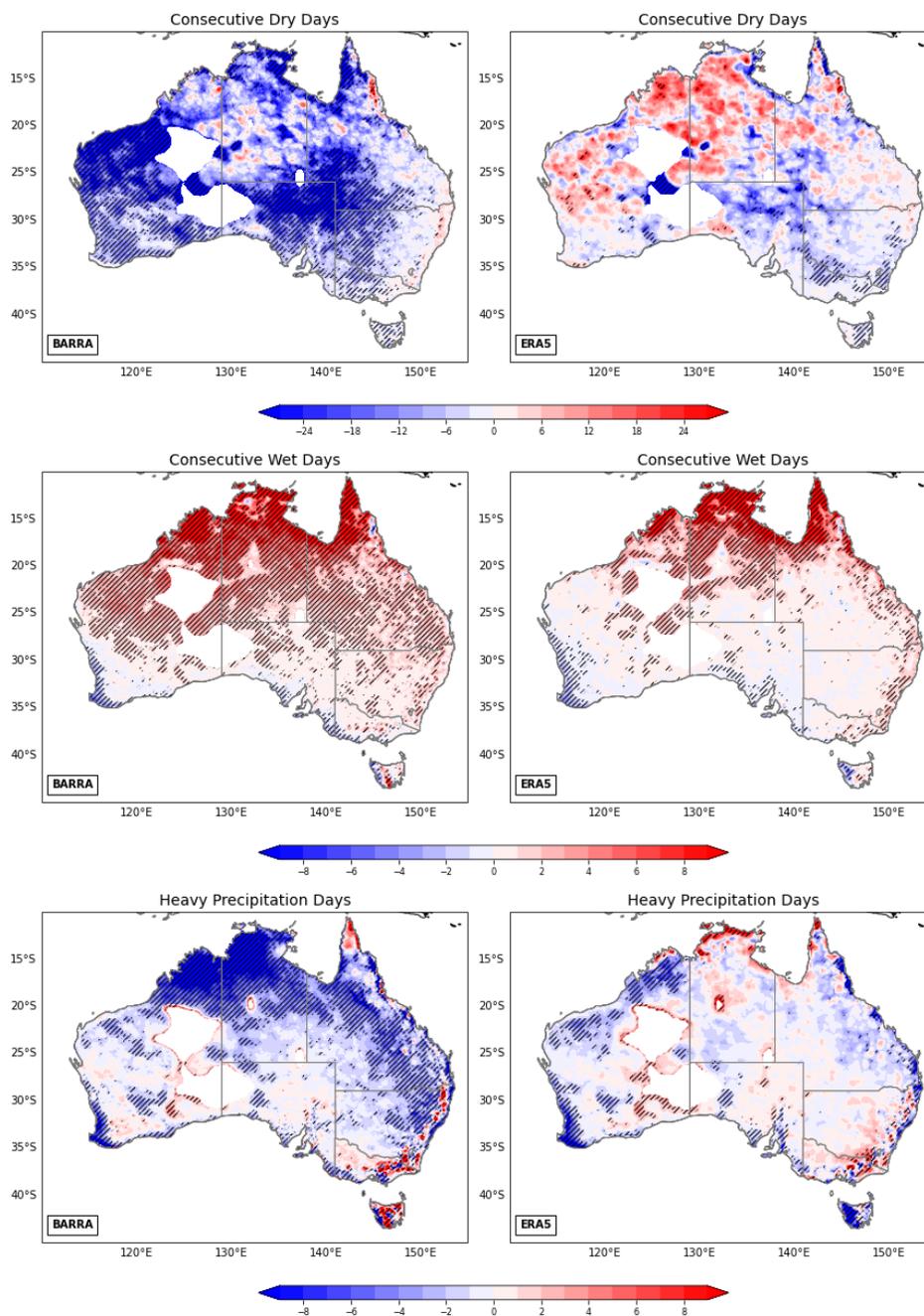
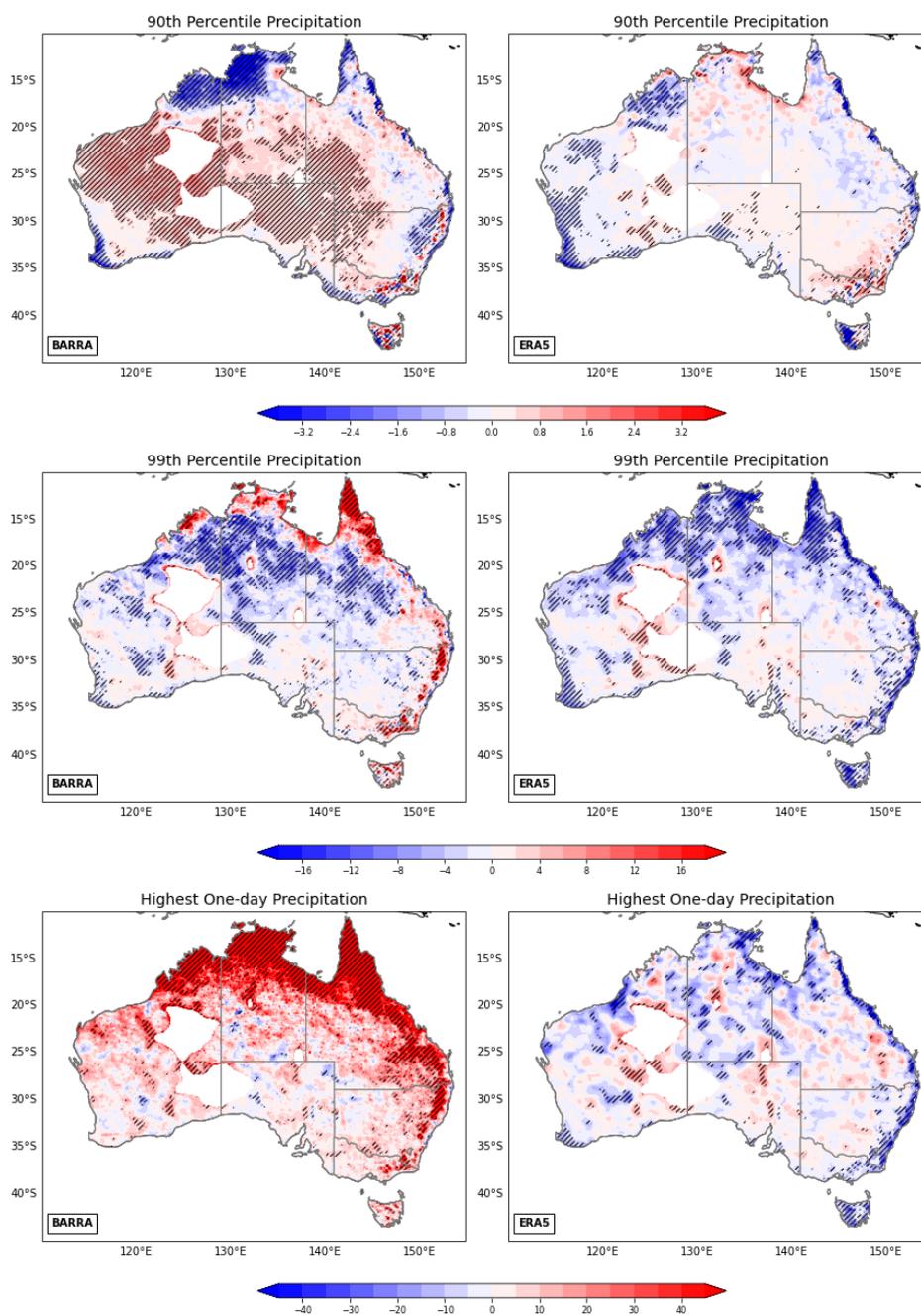


Figure 5 Biases in CDD, CWD, R10mm, R90p, R99p and Rx1Day in BARRA (left column) and ERA5 (right column). Stippling indicates areas with biases that are statistically significant at 95% confidence level.



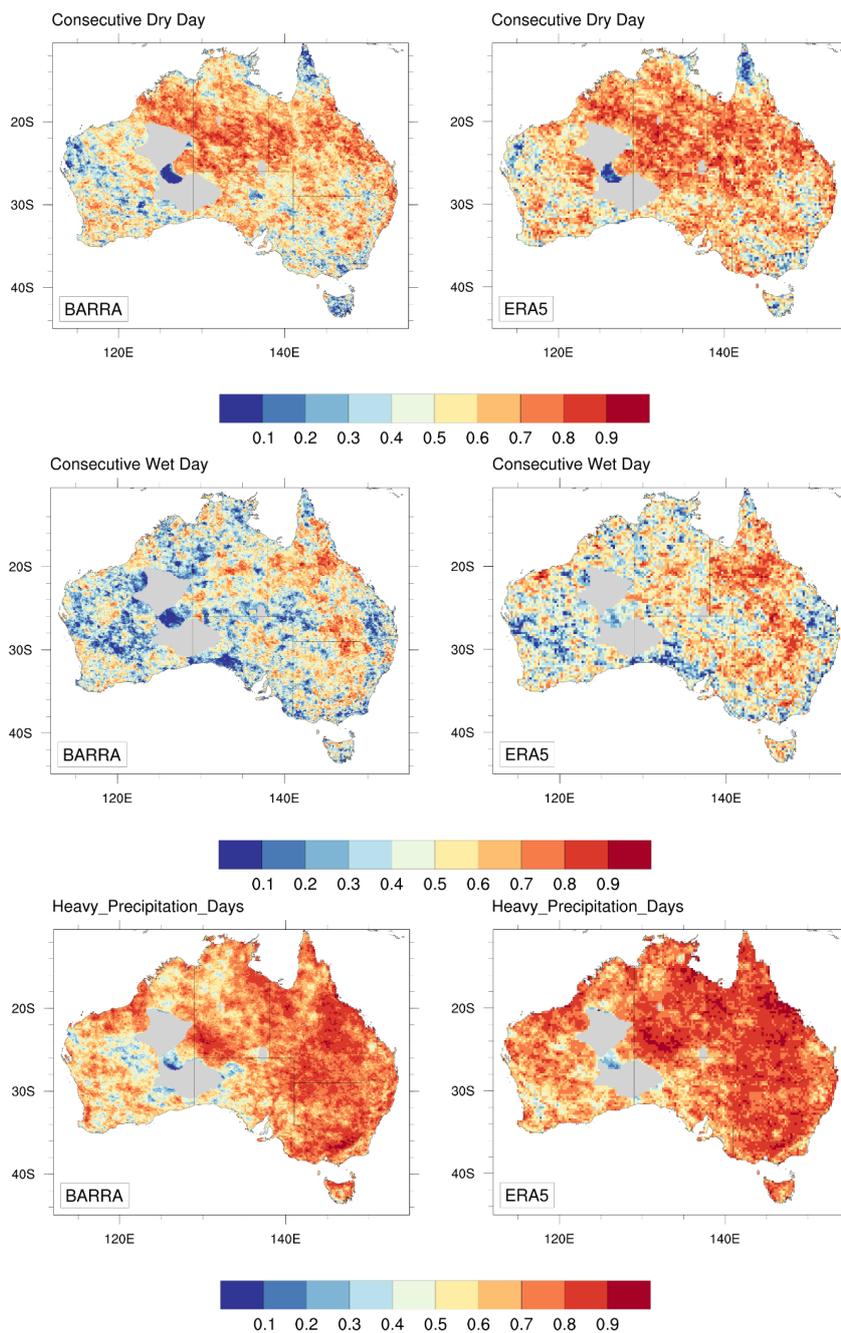


Figure 6 Temporal correlation of CDD, CWD, R10mm, R90p, R99p and Rx1Day between BARRA and AGCD (left column) and between ERA5 and AGCD (right column).

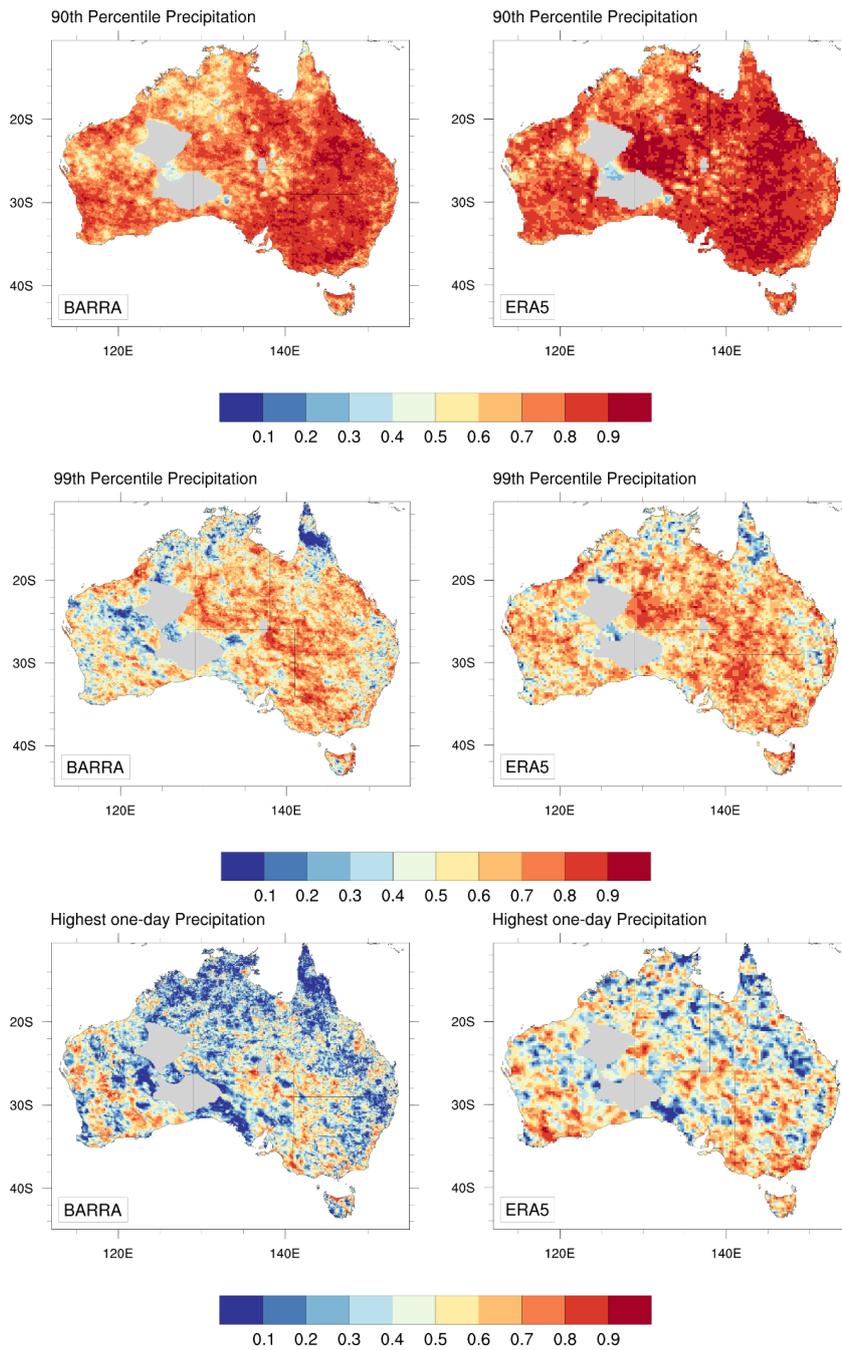


Figure 6 (continued).

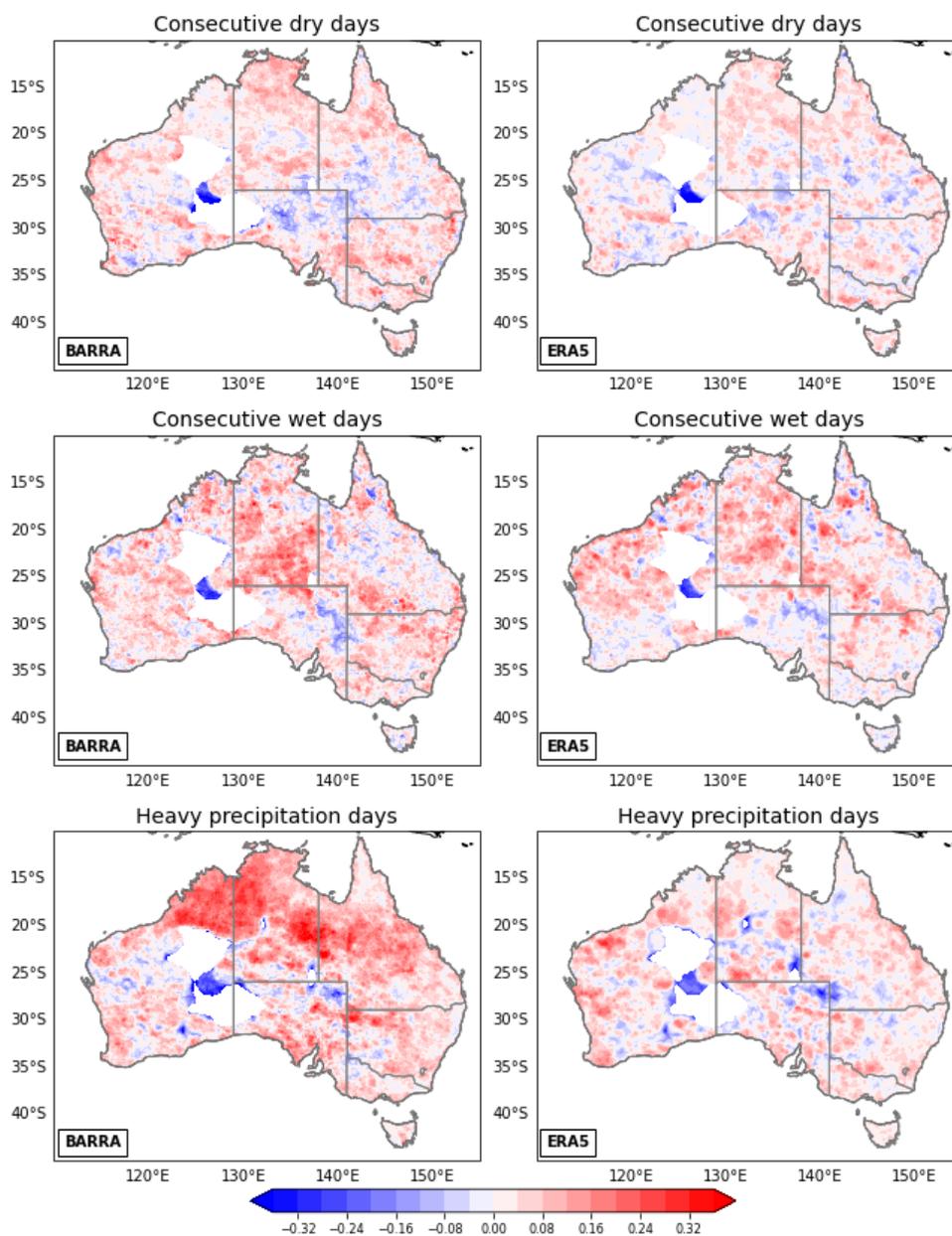


Figure 7 Biases in CV of CDD, CWD, R10mm, R90p, R99p and Rx1Day for BARRA (left column) and ERA5 (right column) relative to AGCD.

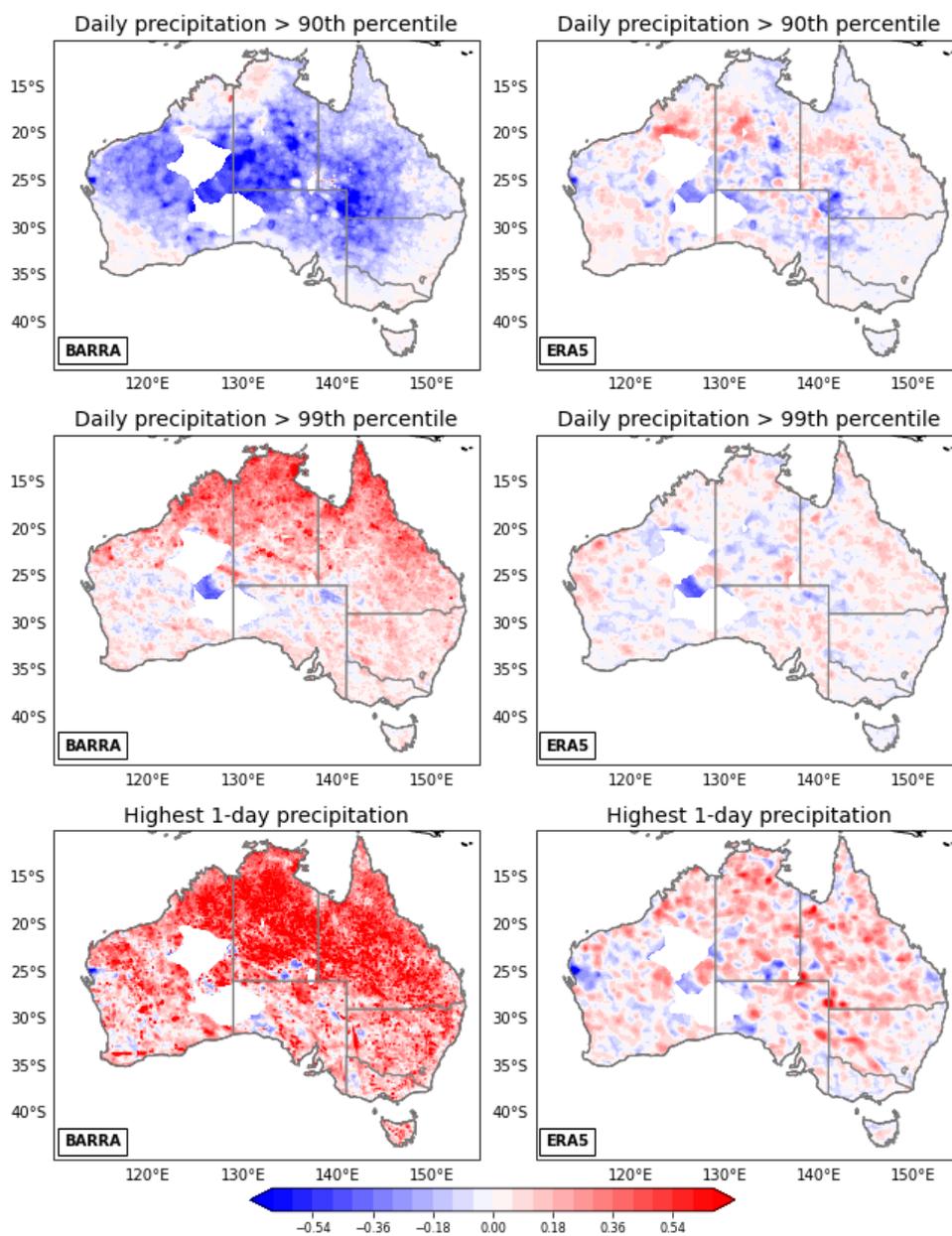


Figure 7 (continued).

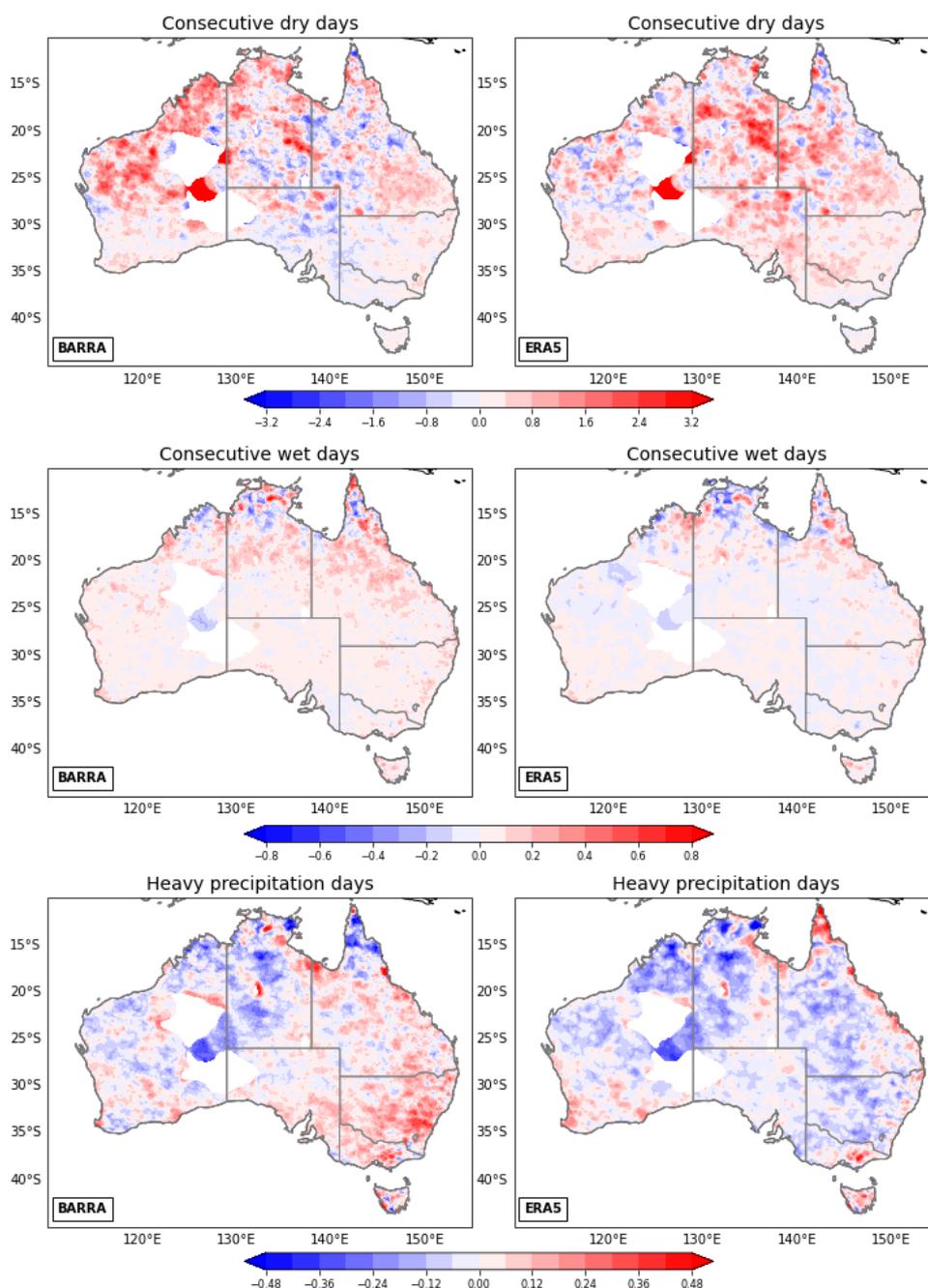


Figure 8 Biases in trends of CDD, CWD, R10mm, R90p, R99p and Rx1Day for BARRA (left column) and ERA5 (right column) relative to AGCD.

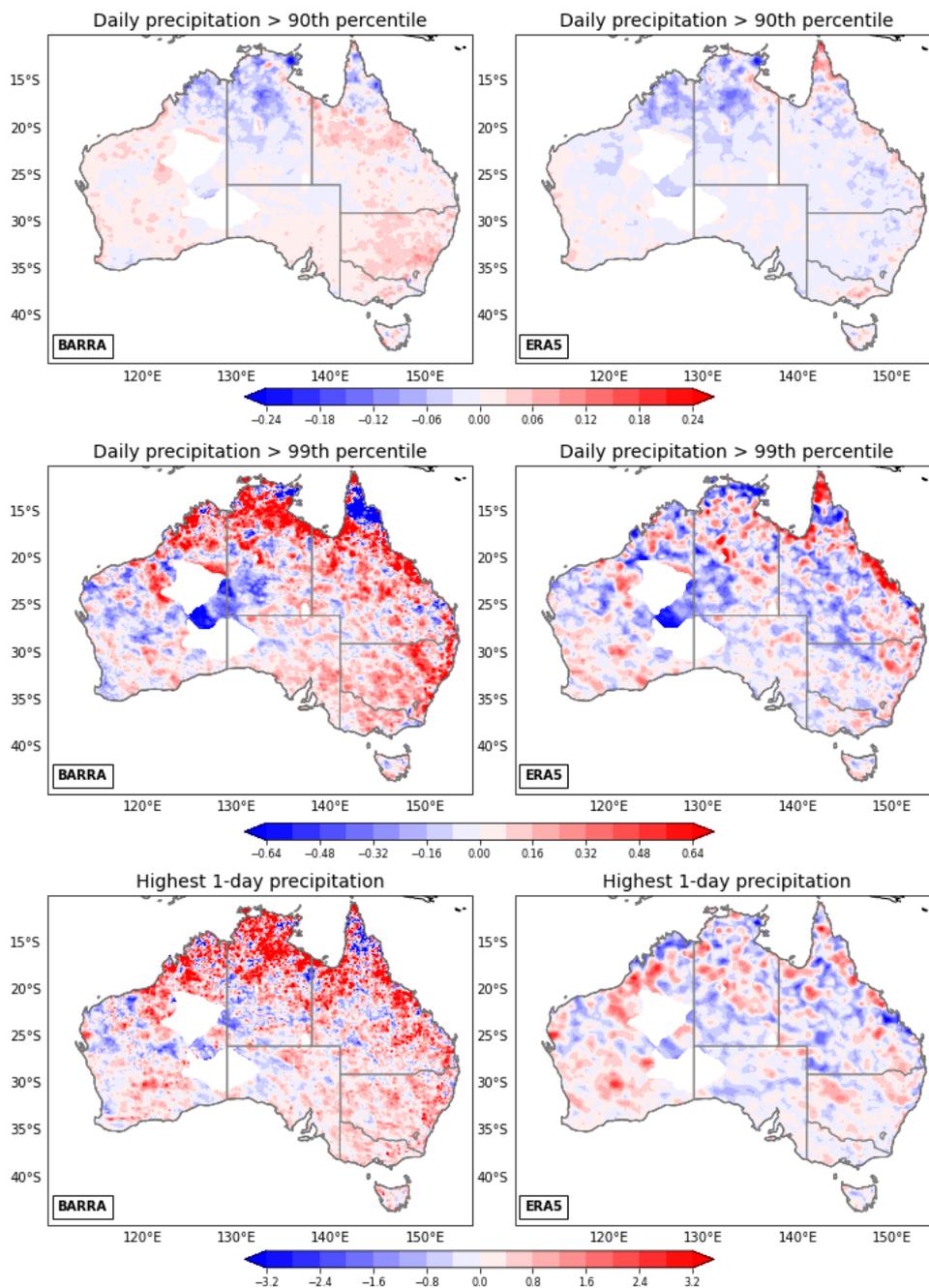


Figure 8 (continued).