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2 **Comparison of BARRA and ERA5 in Replicating Mean and Extreme**

3 **Precipitation over Australia**

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37 **Abstract**

38 Reanalysis datasets are critical in climate research and weather analysis, offering consistent
39 historical weather and climate data crucial for understanding atmospheric phenomena, and
40 validating climate models. However, biases exist in reanalysis datasets that would affect their
41 applications under circumstances. This study evaluates BARRA, which is a high-resolution
42 reanalysis for the Australian region, and ERA5 in simulating mean precipitation and six
43 selected precipitation extremes for their climatology, temporal correlation, coefficient of
44 variation and trend. Both [datasets](#) reproduce [daily timescale probability density distributions](#)
45 [and](#) spatial patterns of mean precipitation well with minor biases. ERA5 shows stronger
46 temporal correlations, superior inter-annual precipitation accuracy, and lower biases in
47 coefficient of variation compared to BARRA, especially in Northern Australia. However, both
48 models exhibit substantial biases in trend, underestimating increasing trends in Northern
49 Australia. ERA5 underestimates dry days and heavy rainfall, while BARRA tends to
50 overestimate these extremes. Temporal correlations for extreme precipitation indices are
51 weaker compared to mean annual precipitation. Notable differences exist in variability biases,
52 with BARRA showing larger biases, especially for heavy precipitation in inland regions and
53 Northern Australia. While both datasets replicate the main trends, biases persist. Overall, the
54 evaluation results support application of both datasets for climatology analyses, but caution is
55 advised for variability and trend analyses, particularly for specific extremes.

56

57 **Key words:** BARRA, ERA5, extreme indices, temporal correlation, coefficient of variation,
58 trend

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60 **1. Introduction**

61 Reanalysis dataset is created by combining historical observational data from various
62 sources, such as weather stations, satellites, buoys, and more, with modern data assimilation
63 techniques and numerical models (Kalnay, et al. 1996; Saha, et al. 2010; Dee et al. 2011;
64 Kobayashi et al. 2015, Poli et al. 2016; Hersbach 2020). The fundamental aim of reanalysis is
65 to construct a uniform and coherent historical archive of various atmospheric and
66 environmental parameters, such as temperature, humidity and wind patterns, on either a
67 regional or a global scale.

68 These datasets are invaluable for climate studies, weather analysis and model validation
69 as they provide a uniform representation of historical climate conditions. For instance,
70 Quagraine et al. (2020) used five global reanalysis datasets (European Centre for Medium-
71 Range Weather Forecasts Reanalysis ERA-Interim, Dee et al. 2011; ERA5, Herbach et al. 2020;
72 JRA-55, Kobayashi et al. 2015); MERRA2, (Gelaro et al. 2017); and NCEP-R2, Kanamitsu et
73 al. 2002) to investigate the variability of West African summer monsoon precipitation, showing
74 all datasets could represent the average rainfall patterns and seasonal cycle. Dai et al. (2023)
75 utilized ERA5 data to estimate rainfall erosivity on the Chinese Loess Plateau, finding rainfall
76 erosivity derived from ERA5 was highly consistent with those derived from the meteorological
77 stations. Cheung et al. (2023) employed ERA5 to evaluate storm conditions in regional climate
78 simulations, demonstrating regional climate models can capture climatology of measurements
79 of storm severity over land including their spatial patterns and seasonality. Numerous studies
80 have used reanalysis datasets as inputs for regional climate models (RCMs) to evaluate the
81 models' capability in replicating observed climatic patterns (Solman et al., 2013; Ji et al., 2016;
82 Fita et al., 2016, Di Virgilio et al., 2019; Capecchi et al., 2023; Di Virgilio et al., 2024; Ji et al.,
83 2024).

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88 While reanalysis datasets provide valuable insights into historical weather and climate
89 conditions, they have limitations and uncertainties, given that they are modelled outputs rather
90 than direct observations. Many studies have evaluated reanalysis data across various variables
91 and regions. For instance, Betts et al. (2019) assessed ERA5 biases in near-surface variables
92 over Canada, highlighting its improved performance over ERA-Interim (Dee et al. 2011),
93 though precipitation biases remained significant. Similarly, Hu and Yuan (2021) and Jiang et
94 al. (2021) found that ERA5 precipitation accurately captured rainfall pattern over the Eastern
95 Tibetan Plateau and mainland China, but under-estimated intensity. Izadi et al. (2021) found
96 ERA5 performed better at monthly and seasonal timescales in Iran, underestimating coastal
97 summer precipitation and overestimating it in mountains. Jiao et al. (2021) and Qin et al. (2021)
98 found ERA5 overestimated summer precipitation and frequency in China but underestimated
99 intensity during the warm season. Lei et al. (2022) and Shen et al. (2022) noted ERA5's
100 limitations in simulating extreme precipitation events in China, especially for high-end
101 extremes.

102 Comparisons between reanalysis datasets have also been conducted. Wang et al. (2019)
103 found that both ERA5 and ERA-Interim exhibited warm biases over Arctic Sea ice, with larger
104 biases in cold season than warm season. Lei et al. (2020) showed ERA5 improved cloud cover
105 simulation over eastern China but not over the Tibetan Plateau, when compared to ERA-
106 Interim. Gleixner et al. (2020) found ERA5 reduced biases in temperature and precipitation
107 over East Africa compared to ERA-Interim but still struggled with long-term trends. Song and
108 Wei (2021) found both ERA5 and MERRA-2 (Gelaro et al. 2017) captured night precipitation
109 peaks over North China, but only ERA5 accurately reflected the afternoon peak. Li et al. (2022)
110 concluded that ERA5 performed better than ERA-Interim, JRA55 (Kobayashi et al. 2015), and
111 MERRA-2 in capturing precipitation over the Poyang Lake Basin. A summary of the above
112 literature review can be found in Table S1.

113 In Australia, reanalyses like NCEP (Kalnay et al., 1996), JRA-55 (Kobayashi et al.,
114 2015), ERA-Interim (Dee et al., 2011), and ERA5 (Hersbach et al., 2020) are commonly used,
115 alongside the Australian Bureau of Meteorology's high-resolution (12 km) BARRA reanalysis.
116 BARRA covers Australia, New Zealand, and Southeast Asia (Su et al., 2019), while BARRA-
117 C offers even higher-resolution (1.5 km) analysis for four capital cities (Su et al., 2021).

118 May et al. (2021) found BARRA reliable, though it showed seasonal and diurnal biases.
119 Other studies, like Pirooz et al. (2021), compared BARRA with global reanalyses, concluding
120 BARRA performed better for precipitation and temperature in New Zealand but lagged behind
121 ERA5 for high gust winds. Du et al. (2023) used BARRA for estimating daily precipitation in
122 ungauged Australian catchments, while Hobeichi et al. (2023) employed BARRA to train
123 statistical models for downscaling. Acharya et al. (2019, 2020) found BARRA's precipitation
124 performance varied by region, with poorer results in tropical areas. Nishant et al. (2022)
125 suggested higher resolution in BARRA-C didn't always improve precipitation simulations,
126 while Choudhury et al. (2023) noted ERA5 performed better for mean temperatures than
127 extremes in Australia. These previous studies on BARRA and BARRA-C have also been
128 summarized in Table S1.

129 However, there is a gap in the existing studies concerning the intercomparison of
130 various reanalyses, such as BARRA and ERA5, specifically in relation to precipitation
131 extremes over Australia. In this study, we aim to bridge this gap by evaluating and comparing
132 the performance of BARRA and ERA5 in capturing precipitation extremes. While the
133 traditional evaluation methods focusing on climatology (long-term mean), here we also include
134 temporal correlation, coefficient of variation and trend in evaluation to quantify their overall
135 performance, which have not been examined before in previous studies. By assessing climate
136 means and extremes and quantifying their biases, this study provides a valuable reference for
137 selecting appropriate datasets for specific applications and cautions against treating reanalysis

138 data as observations. The paper is organized as follows: Section 2 introduces the reanalysis
139 datasets and observational data used for evaluation. Section 3 outlines the climate extreme
140 indices and evaluation methodology. Results are presented in Section 4, followed by further
141 discussion in Section 5. Finally, Section 6 offers a summary and conclusions.

142

143 **2. Data**

144 **2.1 ERA5**

145 ERA5 is a global atmospheric reanalysis dataset developed by ECMWF (Hersbach, et
146 al. 2020). ERA5 provides hourly estimates of many atmospheric, land, and oceanic climate
147 variables. The data is on a ~30 km horizontal grid and resolves the atmosphere using 137 levels
148 from the surface up to a height of 0.01hPa (~80 km).

149 ERA5 is constructed upon the foundation of the Integrated Forecasting System (IFS)
150 Cy41r2. This allows ERA5 to benefit from a decade's worth of development in areas such as
151 model physics, core dynamics, and data assimilation techniques. ERA5 is a significant
152 advancement over its predecessors (e.g., ERA-Interim) due to its higher spatial and temporal
153 resolution, improved assimilation techniques, and more sophisticated modelling components.
154 It provides a detailed and accurate representation of various atmospheric variables, such as
155 temperature, humidity, wind speed, pressure, and more. The dataset covers the entire globe and
156 spans from 1940 to the present, making it valuable for various applications in climate research,
157 meteorology, environmental science, and more.

158 **2.2 BARRA**

159 BARRA is a high-resolution regional atmospheric reanalysis dataset developed by the
160 Australian Bureau of Meteorology, which is available from January 1990 to February 2019 (Su,
161 et al. 2019). BARRA was constructed based on the Australian Community Climate Earth-
162 System Simulator (ACCESS) model with assimilation of a wide range of observational data to

163 create a coherent and consistent representation of past weather and climate conditions. BARRA
164 covers the Australian continent, New Zealand, part of Asia and some Pacific Islands with a
165 horizontal resolution of 12 km and 70 vertical levels from the surface up to a height of 80 km.
166 BARRA specifically focuses on providing detailed information about weather patterns and
167 atmospheric variables over the Australian region, which provides about 100 parameters at
168 hourly intervals.

169 [The ACCESS model, which was applied to generate BARRA, originated from the](#)
170 [UKMO's Unified Model \(UM\), which can be configured in global mode or regional mode. For](#)
171 [regional simulations, the global version of ACCESS becomes ACCESS-R. ACCESS-R was](#)
172 [initialized by ERA-Interim reanalysis data, which also provides boundary conditions during](#)
173 [simulation. A series of observations have been assimilated into BARRA, including land and](#)
174 [ship \(buoy\) synoptic observations, upper-air observations from radiosondes and wind profilers,](#)
175 [satellite derived radiances and winds \(Su et al. 2019\). However, no precipitation observations](#)
176 [were directly assimilated.](#)

177 **2.3 AGCD**

178 The observational data in the study are from the Australian Gridded Climate Dataset
179 (AGCD, Evans et al. 2020). The daily gridded maximum and minimum temperatures, and
180 precipitation data has a spatial resolution of 0.05° (~5km) and is interpolated from observations
181 at stations across the Australian continent. Most of those stations are in the more heavily
182 populated coastal regions with far fewer stations inland and over high elevation areas. For
183 example, there are very few station observations near the Gibson desert region in Western
184 Australia, making the gridded observations unreliable over that region. Thus, in the following
185 figures that region has been masked and not considered for evaluation. Since observations and
186 reanalyses are not at the same spatial resolutions, we aggregate the observations to the native
187 grid of ERA5 and BARRA respectively for comparison, including the performance of

188 statistical significance tests. For comparison purpose, we also interpolate reanalysis to AGCD
 189 grids using the conservative area weighted re-gridding scheme from the Climate Data
 190 Operators (Schulzweida et al., 2006), which will be shown in the Supplementary Information.
 191 The states and sub-regions in the Australian region we discuss in the following can be found
 192 in Figure S1.

193

194 3. Methodology

195 3.1 ET-SCI

196 While extreme climate and weather events are generally multifaceted phenomena, in
 197 this study we evaluate climate extremes based on daily precipitation and temperature as defined
 198 by Expert Team on Sector-specific Climate Indices (ET-SCI; Alexander & Herold, 2015;
 199 Herold and Alexander, 2016). We use the ClimPACT version 2 software to calculate the ET-
 200 SCI indices (<https://climpact-sci.org/>), focussing on daily precipitation.

201 Although ClimPACT generates 14 precipitation-related core indices, we select seven
 202 (Table 1) based on the following considerations: 1) To capture key aspects of climate extremes,
 203 and 2) to capture extremes which have impacts on society and infrastructure such as agriculture,
 204 water resources and economy (Tabari, 2020; Pei et al., 2021). Accordingly, we include absolute
 205 indices such as the maximum 1-day precipitation (Rx1day) and total precipitation (PRCPTOT),
 206 a threshold-based index (e.g., number of heavy rain days, R10mm), percentile indices (e.g.,
 207 total annual precipitation from very heavy rain days, R99p), and duration indices such as the
 208 consecutive wet (CWD) and dry days (CDD).

209 With the above consideration, the seven aforementioned precipitation-related indices
 210 were calculated on native reanalysis grids and observation grids. While the availability of
 211 AGCD and ERA5 starts much earlier, the analysis period is 1990–2018, which is the duration
 212 of BARRA. Since the AGCD observations have the highest resolution, here we mainly show

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221 the evaluation on the native grids of the reanalyses (i.e., the 12-km grid of BARRA and 30-km
 222 grid of ERA5). The extreme indices calculated from reanalysis data have also been regridded
 223 to the 5-km resolution using bilinear interpolation, which are included in the supplementary
 224 information to demonstrate that our conclusions are insensitive to the choice of evaluation
 225 resolution.

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227 3.2 Evaluation metrics

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228 We evaluate BARRA and ERA5 for their performance in capturing daily precipitation
 229 probability density functions (PDFs), climatology, (29 years in our case), coefficient of
 230 variation (CV), temporal correlation, and trends of seven selected precipitation extreme indices.

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231 Each PDF is evaluated using the skill score defined by Perkins et al. (2007), which quantifies
 232 the common area between two reanalyses and observations. For each grid, the maximum
 233 rainfall range between reanalysis and observation is divided into 200 bins to compute
 234 normalized histograms over the same range. The common area is determined by calculating
 235 the minimum of the pairwise frequencies, and the standardized overlap area is obtained by
 236 summing these minimum frequencies and multiplying by the bin width. The CV is a valuable

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237 statistical tool representing the ratio of the (yearly) standard deviation to the mean, allowing
 238 for the comparison of variation between different data series, even when their means differ
 239 significantly. Temporal correlations, which are computed at an annual time step, of climate

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240 extremes measure the similarities between simulations and observations in terms of their inter-
 241 annual variabilities, with larger temporal correlations indicating better performance. For trend
 242 analyses, we applied simple linear trend line fitting to the yearly time series of climate indices.
 243 All the above metrics are computed at each grid point in the datasets' native grids as well as
 244 the AGCD grid after re-gridding. Differences between BARRA/ERA5 and AGCD then form

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the bias maps. After averaging over all grid points, the domain averages will then be discussed in the following.

We use bias and domain-averaged absolute bias to quantify spatial differences between reanalyses and observations. Temporal correlation, coefficient of variation, and trend are used to quantify temporal similarities between reanalyses and observations. The non-parametric Mann-Kendall test is used to assess the statistical significance of differences and trends. Biases are assessed at an annual timescale for all extremes.

4. Results

4.1 Mean climate

This section evaluates and compares the annual mean of daily precipitation between BARRA and ERA5 against AGCD over Australia.

4.1.1 Daily precipitation PDF

We first compare the Perkins Score for daily precipitation between BARRA, ERA5, and AGCD (Figure 1). Both reanalyses generally capture the daily precipitation PDF well, with Perkins Scores exceeding 0.9 across most of Australia, indicating a strong agreement with observations. However, there are regional variations in performance. Scores are relatively lower in northern Australia and Tasmania, suggesting greater discrepancies in these areas. ERA5 exhibits noticeably higher scores inland compared to coastal regions, reflecting its improved representation of precipitation in interior regions. In contrast, BARRA does not show a clear spatial pattern, with more variability across different locations. Overall, ERA5 outperforms BARRA, particularly in inland areas, likely due to its global data assimilation approach. However, BARRA shows better agreement in the southeastern coastal regions.

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273 [indicating its advantage in capturing local precipitation patterns influenced by complex terrain](#)
274 [and coastal effects.](#)

275

276 4.1.2 Bias and temporal correlation

277 We evaluate precipitation simulated by BARRA and ERA5 against observations
278 (AGCD). The mean annual precipitation from the three datasets and biases in BARRA and
279 ERA5 compared to AGCD are shown in Figure 2 (and Figure S2 on the observation grid).

280 Results show that both BARRA and ERA5 simulate the spatial patterns of mean annual
281 precipitation very well with high rainfall in northern Australian, eastern Australia coast and
282 western Tasmania and low rainfall inland, albeit with clear biases. Compared to AGCD, both
283 BARRA and ERA5 underestimate precipitation up to 20% for Eastern Australian coast,
284 southwest western Australia, and western Tasmania, but overestimate annual precipitation up
285 to 30% inland (Figure S3). Some clear differences in biases between BARRA and ERA5 can
286 be observed in central western Australia and northern Queensland where BARRA overestimate
287 precipitation but ERA5 underestimate it. Domain averaged absolute bias in annual precipitation
288 is about 0.17mm/day (~12.7% [relative bias with respect to domain average](#)) for BARRA and
289 0.15 mm/day (~10.5% [relative bias](#)) for ERA5 (Table 2).

290 The skill of simulated precipitation from BARRA and ERA5 are further demonstrated
291 in the temporal correlations between BARRA/ERA5 and AGCD shown in Figure 3 (and Figure
292 S4 on the observation grid). Temporal correlation of annual precipitation is larger in southeast
293 Australia and northern Tasmania for both BARRA and ERA5, which is above 0.85. This
294 indicates inter-annual variability of precipitation is well captured by BARRA and ERA5. In
295 contrast, temporal correlation is weaker for western inland and northern Australia. ERA5
296 generally has larger temporal correlation when compared with BARRA, especially for northern

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302 Australia, where temporal correlation for BARRA is below 0.5. On average, temporal
303 correlation for ERA5 is 0.85, which is [larger](#) than 0.77 for BARRA (Table 2).

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304 4.1.3 CV (coefficient of variation) and trend

305 CV of annual precipitation for AGCD and biases between BARRA/ERA5 and AGCD
306 are presented in Figure [4](#) (and Figure S5 on the observation grid). By its definition, CV helps
307 capture the standard deviation in the dataset relative to its mean. In the observation, CV is
308 generally smaller for coastal regions [including Tasmania](#) except for northwest West Australia
309 ~~and Tasmania~~ than inland Australia, where annual rainfall is much smaller than coastal regions.
310 Alternatively, regions with higher annual precipitation generally have smaller CV. Both
311 BARRA and ERA5 reasonably capture the main feature of CV in observation. However, clear
312 biases can be observed, especially in BARRA that has more than 50% large positive biases in
313 Northern Australia, up to 20% positive biases for inland, and relatively smaller biases for
314 southeastern Australia, southwest West Australia and Tasmania. In contrast, ERA5 does not
315 have a clear bias pattern, and biases are relatively smaller when compared to BARRA.

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316 To further investigate the variability evident in observations and BARRA/ERA5
317 simulations, we assess the trends in annual precipitation (Figure [5](#) and Figure S6 on the
318 observation grid). AGCD shows strong increasing trends over Northern Australia and
319 Northeast Australia coastal regions but decreasing trends over Northern Queensland,
320 southwestern West Australia and southern Great Dividing Range including Victoria, although
321 not all trends are significant. Most of inland regions have relatively small trend in annual
322 precipitation. Both BARRA and ERA5 reproduce the major trend pattern ~~reasonably well~~,
323 however, clear biases can be observed over Northern Australia where both BARRA and ERA5
324 underestimate ~~trend more than 100% (i.e., trend of 0.08 mm/day per year with bias of similar~~
325 ~~magnitude)~~. BARRA overestimated decreasing trend over Northern Queensland but ERA5
326 underestimate it (even increasing trend instead).

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332 In summary, evaluation of annual mean precipitation indicates both BARRA and ERA5
 333 possess small biases (~20%) in the spatial precipitation patterns. ERA5 shows stronger
 334 temporal correlations than BARRA, particularly in northern Australia. Overall, ERA5
 335 demonstrates higher accuracy in capturing inter-annual precipitation variability. Both BARRA
 336 and ERA5 captured spatial distribution of coefficient of variation reasonably well but with
 337 large biases (~50%). BARRA shows much larger biases than ERA5 especially for Northern
 338 Australia. Both BARRA and ERA5 roughly reproduce the pattern of trend but with very large
 339 biases (~100%), especially for Northern Australia where both substantially underestimate the
 340 increasing trend.

341

342 4.2 Climate extremes

343 This section evaluates the [seven](#) select precipitation extreme indices (Table 1) from
 344 BARRA and ERA5 over Australia by comparing them against AGCD. Evaluations are
 345 performed primarily using spatial bias maps and temporal correlations. We also assess the
 346 interannual variability and trends in the simulated BARRA and ERA5 indices and compare
 347 these with AGCD to further investigate any discrepancies.

348 4.2.1 Bias and temporal correlation

349 Annual mean biases in [six](#) precipitation extremes are shown in Figure [6](#) (and Figure S8
 350 on the observation grid). For duration-related extremes, [\(CDD and CWD\)](#), there is a clear north-
 351 to-south gradient in AGCD (Figure S7) with longer duration of CDD and CWD in northern
 352 Australia than southern Australia (CWD also has a clear west-to-east gradient in Tasmania),
 353 which is well simulated in BARRA and ERA5 (Figure S7). While the spatial distributions are
 354 well captured, clear biases are evident in them (Figure [6](#)). BARRA generally underestimates
 355 CDD especially for central inland and northwest West Australia where biases are up to 40%.
 356 ERA5 also under-estimates CDD for central inland, but in contrast its over-estimates CDD for

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362 most of northwestern Australia, overall ERA5 has smaller absolute bias in CDD (6.9 days) than
 363 BARRA (14.5 days) (Table 2). Both BARRA and ERA5 have similar bias pattern for CWD,
 364 which generally overestimate CWD over most of regions except for southern Australian coast,
 365 southwest West Australia and western Tasmania. The positive biases over Northern Australia
 366 may reach 30%. Overall BARRA has slightly larger biases in CWD (2.3 days) than ERA5 (1.7
 367 days) (Table 2).

368 For threshold-based extremes (PRCPTOT, R10mm, R90p, R99p, Rx1day), both
 369 BARRA and ERA5 also generally match the spatial distribution of heavy precipitation days
 370 and R90p (Figure S7) in AGCD with large values in Northern Australia, eastern seaboard and
 371 Australian Great Dividing Range, and western Tasmania. However, clear biases can be
 372 observed in BARRA and ERA5 for both R10mm and R90p (Figure 6). BARRA and ERA5
 373 have large negative biases in R10mm over Northern Australia, eastern seaboard, southwest
 374 Western Australia and western Tasmania, but biases in central inland and northwest West
 375 Australia are generally small. Overall, domain averaged absolute bias for ERA5 (1.7 days) is
 376 about half of that for BARRA (3.3 days). Both BARRA and ERA5 also have relatively large
 377 negative biases in R90p for most of northern Australia, eastern coasts, southwest West
 378 Australia and western Tasmania but small positive biases inland, especially for BARRA.
 379 Overall averaged absolute bias is 0.78 mm/day for BARRA and 0.44 mm/day for ERA5 (Table
 380 2).

381 BARRA and ERA5 also reasonably captured the spatial patterns of R99p and Rx1day,
 382 however, quite large biases are in BARRA and ERA5 (Figure 6). BARRA generally
 383 overestimate R99p and Rx1day over northern Australia coasts and along the Great Dividing
 384 Range. In contrast, ERA5 generally underestimate R99p and Rx1day over northern and eastern
 385 coasts, southwest Western Australia and western Tasmania. The domain averaged bias in R99p

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391 is at similar magnitude for BARRA (4.09 mm/day) and ERA5 (3.67 mm/day), however biases
392 in Rx1day is much larger for BARRA (20.3 mm/day) than ERA5 (7.9 mm/day) (Table 2).

393 Figure 7 (and Figure S9 on the observation grid) presents the temporal correlations
394 between BARRA/ERA5 and AGCD for the six precipitation extreme indices. Unlike the strong
395 temporal correlation between BARRA/ERA5 and AGCD for mean annual precipitation (Figure
396 3), the temporal correlations for these extreme indices are worse except for R90p (Figure 7).
397 For extremes like R10mm and R90p, the correlation ranges from reasonably good (above 0.6)
398 to pretty good (above 0.8) between BARRA/ERA5 and AGCD for most of the domain.
399 Temporal correlation for CDD, CWD and R99p are not as good as R10mm and R99p. CDD
400 has more regions with stronger correlations (0.5-0.6) or above than CWD and Rx1day. for the
401 latter correlation is about ~0.5 or less over most of the domain. Compared to BARRA, ERA5
402 has slightly stronger temporal correlations for those extremes (Table 2).

403

404 4.2.2 CV (coefficient of variation) and trend

405 The observed and simulated CV of precipitation extremes and biases in their CV for
406 BARRA and ERA5 are shown in Figure S10 and Figure 8 (and Figure S11 on the observation
407 grid), respectively. Generally, both BARRA and ERA5 have similar CV bias patterns and
408 magnitude for CDD, CWD and R10mm. In contrast, BARRA is quite different from ERA5 for
409 other three extremes. BARRA substantially under-estimated CV of R90p over most on inland
410 regions but ERA5 has much smaller negative biases, even small positive biases, although both
411 have small biases in CV of R90p along most coastal regions and Tasmania. BARRA
412 systematically overestimate CVs of R99p and Rx1day over northern Australia but ERA5 has
413 relatively small biases for them. Overall, BARRA has more than twice as much as CV biases
414 in ERA5 for R90p, R99p and Rx1day (Table 2).

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423 Trends of each of the precipitation extreme indices for the three datasets and biases in
 424 trend for BARRA and ERA5 are shown in Figure S12 and Figure 9 (and Figure S13 on the
 425 observation grid), respectively. Generally, both BARRA and ERA5 simulate the main pattern
 426 of trends for those extremes but with large biases. BARRA and ERA5 simulated CDD trend
 427 well for southern Australia but BARRA generally under-estimated trend in CDD over inland
 428 Australia and overestimate trend in northwest Australia. ERA5 only has large positive trend
 429 biases in northern central Australia. The overall domain averaged biases are similar between
 430 BARRA (0.584) and ERA5 (0.566). Both BARRA and ERA5 have small biases in CWD in
 431 central and southern Australia but similar biases pattern in Northern Australia. They also have
 432 similar overall biases in CWD (0.064 for BARRA and 0.060 for ERA5). Both BARRA and
 433 ERA5 under-estimated increasing trend in R10mm in northern Australia, but BARRA
 434 overestimate trend in most of southeast Australia. In contrast, ERA5 under-estimate trend over
 435 there. Overall, ERA5 has slightly larger biases (0.094) than BARRA (0.085). Like R10mm,
 436 both BARRA and ERA5 also underestimate trend of R90p in most of northern Australia but
 437 have small biases in central and southern Australia. They have almost the same overall biases
 438 in R90p. BARRA/ERA5 has similar biases patterns for R99p and [Rx1day](#) but biases for
 439 rx1days are much larger. Both BARRA and ERA5 have large biases in R99p and Rx1day but
 440 biases in BARRA are generally larger than ERA5.

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441 In summary, both BARRA and ERA5 reproduce spatial patterns of extremes well but
 442 display biases. ERA5 underestimates CDD and certain [extreme precipitation indices \(e.g.,](#)
 443 [Rx1day\)](#), while BARRA tends to overestimate these extremes. Both reanalyses show
 444 discrepancies in various precipitation indices across different regions, with BARRA generally
 445 displaying larger biases compared to ERA5. Temporal correlations between BARRA/ERA5
 446 and observations for extreme precipitation indices are weaker than those for mean annual
 447 precipitation, except for a few indices where ERA5 demonstrates slightly stronger correlations

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451 compared to BARRA. Both BARRA and ERA5 align in CV patterns and biases for certain
452 extremes (CV, R10mm, R90p) but differ notably in others (PRCPTOT, trend, CDD, R99p,
453 Rx1day). BARRA significantly underestimates very heavy precipitation variability over inland
454 regions, while ERA5 presents smaller biases or even positive biases in these areas. Additionally,
455 BARRA tends to overestimate extreme precipitation variability in Northern Australia
456 compared to ERA5. Overall, BARRA shows more than double the biases in variability
457 compared to ERA5 for specific extreme precipitation indices. Both reanalyses generally
458 simulate the main trend patterns but exhibit considerable biases. BARRA underestimates or
459 overestimates trends in certain regions and indices, while ERA5 demonstrates different biases,
460 including smaller biases overall compared to BARRA across these precipitation extremes.

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462 5. Discussion

463 In this study, we assessed the performance of BARRA and ERA5 in simulating mean
464 precipitation and six selected precipitation extremes. While most previous evaluations have
465 focused on the climatology of precipitation and its extremes, only a few studies have included
466 the coefficient of variation (CV) (Teng et al., 2024). Our evaluation encompassed annual
467 climatology, along with temporal correlation, CV, and trend analysis, providing a
468 comprehensive assessment of the performance of these two reanalysis datasets.

469 The results indicate that both BARRA and ERA5 demonstrate reasonable skill in
470 simulating mean precipitation and certain precipitation extremes (e.g., CWD and R90p).
471 However, they encounter challenges in accurately reproducing temporal correlation, CV, and
472 trends for certain extreme events, highlighting significant uncertainties in their representation
473 of extremes.

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474 While acknowledging the capabilities of these reanalysis datasets, our study also
475 identifies specific limitations and suggests potential directions for future research. A crucial

478 consideration in model evaluation is the accuracy of observational data, which substantially
479 influences evaluation outcomes. In this study, we used the AGCD dataset as the observational
480 benchmark, which is based on interpolating data from in-situ stations (Evans et al., 2020).
481 However, the AGCD dataset presents several limitations: 1) Spatial coverage: Sparse station
482 coverage in northwest and central Australia, and limited observations in high-elevation areas,
483 result in a concentration of stations in southeastern Australia, southwestern Western Australia,
484 and eastern Tasmania. The arid interior is notably underrepresented. 2) Data completeness and
485 homogeneity: Incomplete and inhomogeneous observations due to missing data, changes in
486 observational techniques, or station relocations can affect the consistency of the dataset. 3)
487 Interpolation uncertainties: The interpolation method used in AGCD (splining), instead of the
488 ordinary kriging method used in its predecessor (AWAP), introduces uncertainties, particularly
489 in areas with sparse data coverage for extreme events like heavy rainfall.

490 These observational uncertainties may contribute to biases in the evaluation results. In
491 particular, the limited number of monitoring sites over the Great Dividing Range and inland
492 areas introduces significant uncertainties in estimated observed precipitation for these regions.
493 Independent studies, such as Chubb et al. (2016), found that daily precipitation is
494 underestimated by at least 15% in some areas, which could suggest similar underestimation in
495 BARRA and ERA5 for these regions. Similarly, the sparse gauge network in northwestern
496 inland areas might miss localized extreme precipitation events.

497 Our analysis focused on [seven](#) ET-SCI-defined precipitation extreme indices ([including](#)
498 [mean precipitation](#)), widely used in various evaluation studies (Nishant et al., 2020; Ji et al.,
499 2024). However, recognizing the need for region-specific indices, we suggest future studies
500 extend the analysis to incorporate additional extreme indices tailored to specific regions and
501 applications.

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504 Our findings emphasize that while both BARRA and ERA5 are competent in simulating
505 the climatology of mean climate, temporal correlation, and CV, challenges remain in accurately
506 capturing trends, particularly for certain extremes. Notably, ERA5 shows better overall
507 performance compared to BARRA. Although higher resolution often correlates with better
508 performance, recent studies have shown that increasing resolution alone does not always
509 guarantee improvements (Nishant et al., 2022). Considering the critical role of driving data,
510 model physics, and data assimilation, it may be valuable to update BARRA using the latest
511 ERA5 data along with improved model physics and data assimilation techniques to enhance its
512 performance.

513 In this study, we evaluated ERA5 and BARRA on both their native resolutions and a
514 common resolution (5 km) to match AGCD. The results showed that the evaluations were
515 consistent across native and common resolutions, suggesting that the performance assessments
516 were not highly sensitive to changes in resolution.

517

518 6. Summary and Conclusion

519 Reanalysis datasets play a crucial role in climate research, weather analysis, and various
520 scientific investigations. Their ability to provide a consistent and comprehensive representation
521 of historical weather and climate conditions makes them invaluable. These datasets are
522 particularly essential for studying long-term climate trends, understanding atmospheric
523 phenomena, and validating climate models.

524 In this study, we evaluate BARRA and ERA5 for their capabilities to simulate [daily](#)
525 [precipitation, followed by](#) mean precipitation and six selected precipitation extremes for their
526 climatology, temporal correlation, coefficient of variation (CV) and trend [on monthly timescale](#)
527 to quantify their overall performance. We evaluated BARRA and ERA5 at their native
528 resolutions, as well as at a common resolution (i.e., the observation resolution). Both analyses

529 yielded consistent results, indicating that the evaluation is not sensitive to the remapping
530 process.

531 [Perkins skill score analysis of daily precipitation indicates that both BARRA and ERA5](#)
532 [resemble PDF of AGCD well, with ERA5 slightly outperforms BARRA in inland regions.](#) The
533 assessment of annual mean precipitation reveals that both BARRA and ERA5 adeptly
534 reproduce the spatial precipitation patterns, exhibiting minor biases of around 20%. Particularly,
535 ERA5 showcases stronger temporal correlations compared to BARRA, especially evident in
536 northern Australia. ERA5, overall, demonstrates superior accuracy in capturing inter-annual
537 precipitation variability. However, both models depict the spatial distribution of the coefficient
538 of variation reasonably well but with larger biases, roughly around 50%. Particularly, BARRA
539 displays significantly higher biases, especially in Northern Australia.

540 Regarding the replication of trend patterns, both models exhibit substantial biases,
541 reaching approximately 100%. This is especially notable in Northern Australia, where they
542 both notably underestimate the increasing trend. Furthermore, while both BARRA and ERA5
543 possess about the right spatial patterns of extremes, biases are evident. ERA5 tends to
544 underestimate consecutive dry days (CDD) and certain heavy rainfall events, while BARRA
545 tends to overestimate these extremes. Discrepancies in various precipitation indices across
546 regions are apparent, with BARRA generally displaying larger biases compared to ERA5.

547 When examining temporal correlations for extreme precipitation indices compared to
548 mean annual precipitation, both BARRA and ERA5 show weaker correlations, except for a
549 few indices ([CDD](#), [R10mm](#), [R90p](#)) where ERA5 slightly outperforms BARRA. While both
550 models align in coefficient of variation patterns and biases for certain extremes, notable
551 differences arise in others. BARRA notably underestimates very heavy precipitation variability
552 over inland regions, whereas ERA5 presents smaller biases or even positive biases in these
553 areas. Moreover, BARRA tends to overestimate extreme precipitation variability in Northern

554 Australia compared to ERA5. Specifically, BARRA showcases more than double the biases in
555 variability compared to ERA5 for specific extreme precipitation indices.

556 In terms of trend patterns, both models generally replicate the observed trends but
557 exhibit considerable biases. BARRA shows both underestimations and overestimations in
558 certain regions and indices, while ERA5 displays different biases, including overall smaller
559 biases compared to BARRA across these precipitation extremes.

560 In summary, our findings suggest that both ERA5 and BARRA are reliable for
561 climatological analyses, including mean precipitation and precipitation extremes, and can be
562 confidently used by end-users for such purposes. However, as discussed in the introduction,
563 caution is advised when using these datasets for variability and trend analyses, particularly for
564 specific extreme events like Rx1day. The performance of these reanalyses is regionally
565 dependent, and this should be considered when using them as observational references for
566 evaluating other model simulations. Additionally, the biases in the variability and trends of
567 climate extremes present in both datasets must be carefully accounted for when comparing
568 them with other data sources.

569

570 **Data Availability**

571 Details about AGCD are available at the Australian Bureau of Meteorology website
572 (<http://www.bom.gov.au/metadata/catalogue/19115/ANZCW0503900567>, (accessed on)).
573 The dataset is available on the NCI (National Computational Infrastructure) server in project
574 zv2. Detail on how to access the data can be found at [http://climate-](http://climate-cms.wikis.unsw.edu.au/AGCD)
575 [cms.wikis.unsw.edu.au/AGCD](http://climate-cms.wikis.unsw.edu.au/AGCD), (accessed on). ERA5 data is available on the NCI in Project
576 rt52. BARRA data is available on the NCI in project cj37.

577 **Author Contributions**

578 KKWC and FJ conceptualized and implemented the research. KKWC, FJ and NN performed
579 the data analysis and prepared the figures. [KKWC](#) and FJ prepared the draft manuscript. All
580 authors contributed to the discussion of results, editing and finalization of the manuscript.

581 **Competing Interests**

582 The authors declare that they have no conflict of interest.

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591 [insightful comments and detailed suggestions for us to improve the manuscript.](#)

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759 Table 1 List of ET-SCI indices evaluated in this study.

| Index | Definition | Units | Timescale | Sectors |
|----------------|--|-------|----------------|--|
| PRCPTOT | Total wet-day precipitation (Sum of daily precipitation ≥ 1.0 mm) | mm | Annual/Monthly | Agriculture and food security, water, water resources and food security, forestry/GHGs |
| CDD | Consecutive dry days (Maximum number of consecutive dry days (when precipitation < 1.0 mm)) | days | Annual | Health, agriculture and food security, water resources and food security, disaster risk reduction, forestry/GHGs |
| CWD | Consecutive wet days (Maximum annual number of consecutive wet days (when precipitation ≥ 1.0 mm)) | days | Annual | Coasts, agriculture, transport operations |
| R10mm | Days when precipitation is at least 10mm | days | Annual/Monthly | Coasts |
| R90p | Total annual precipitation from very heavy precipitation days (Annual sum of daily precipitation > 90 th percentile) | mm | Annual | Coasts, transport operations |
| R99p | Total annual precipitation from very heavy precipitation days (Annual sum of daily precipitation > 99 th percentile) | mm | Annual | Coasts, transport operations |
| Rx1Day | Amount of precipitation from very wet days (Maximum 1-day precipitation) | mm | Annual/Monthly | Agriculture and food security, water, coasts, disaster risk reduction, forestry/GHGs |

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763 Table 2 Domain-averaged absolute biases and temporal correlation between BARRA/ERA5
764 and AGCD for annual precipitation and precipitation extremes

| Indices | Absolute biases in annual mean | | Temporal correlation | | Absolute biases in CV | | Absolute biases in trend | |
|------------------|--------------------------------|-------|----------------------|-------|-----------------------|-------|--------------------------|-------|
| | BARRA | ERA5 | BARRA | ERA5 | BARRA | ERA5 | BARRA | ERA5 |
| Annual pr | 0.169 | 0.149 | 0.771 | 0.854 | 0.063 | 0.037 | 0.008 | 0.007 |
| CDD | 14.543 | 6.913 | 0.578 | 0.650 | 0.050 | 0.045 | 0.584 | 0.566 |
| CWD | 2.346 | 1.714 | 0.446 | 0.527 | 0.061 | 0.059 | 0.064 | 0.060 |
| R10mm | 3.265 | 1.700 | 0.688 | 0.761 | 0.081 | 0.053 | 0.085 | 0.094 |
| R90p | 0.777 | 0.439 | 0.761 | 0.827 | 0.211 | 0.082 | 0.023 | 0.023 |
| R99p | 4.093 | 3.668 | 0.562 | 0.625 | 0.121 | 0.060 | 0.206 | 0.162 |
| Rx1day | 20.333 | 7.916 | 0.380 | 0.486 | 0.219 | 0.107 | 0.848 | 0.542 |

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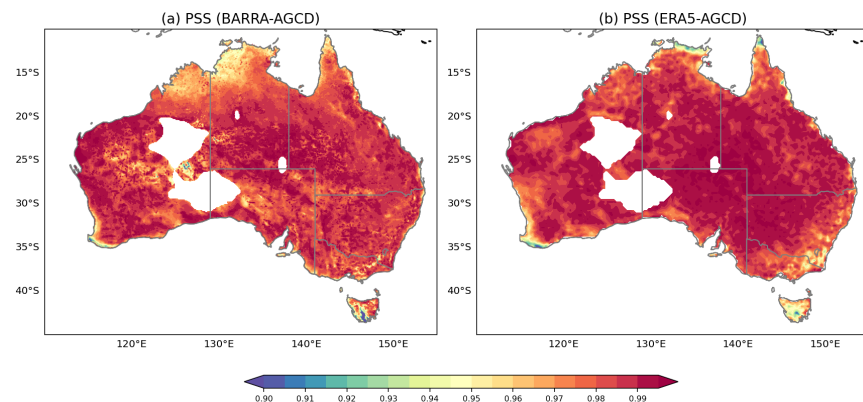


Figure 1 Perkins Skill Score (PSS) of daily precipitation PDF between (a) BARRA and AGCD, (b) ERA5 and AGCD. The regions with low density of station observations in AGCD has been masked and not considered in all subsequent evaluation.

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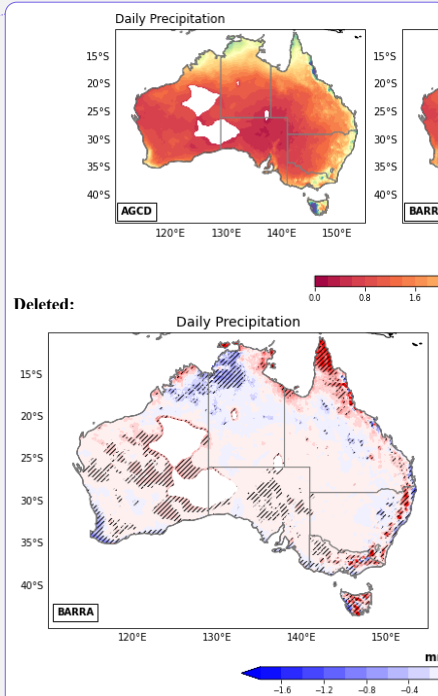


Figure 1

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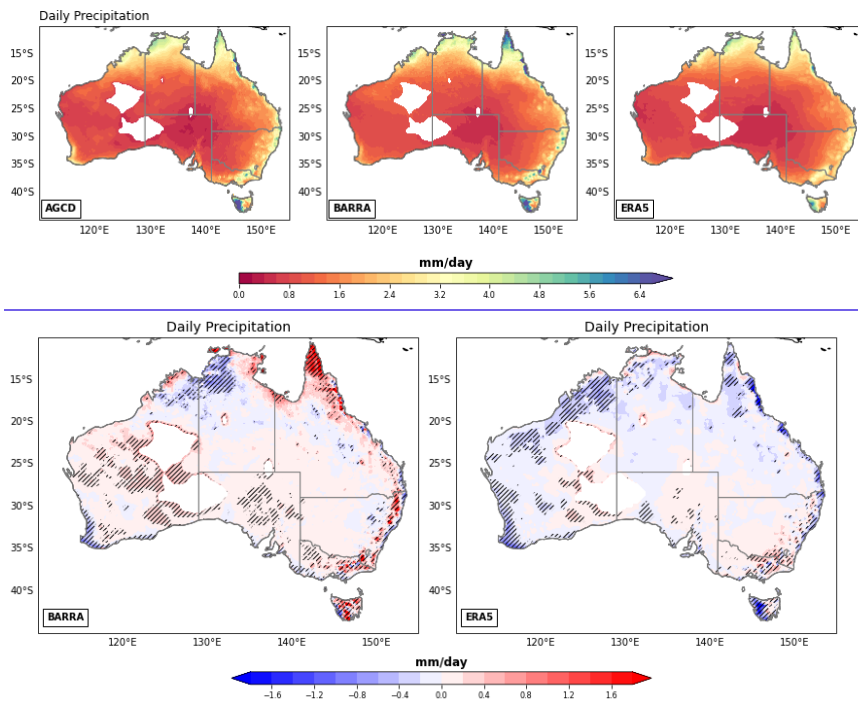
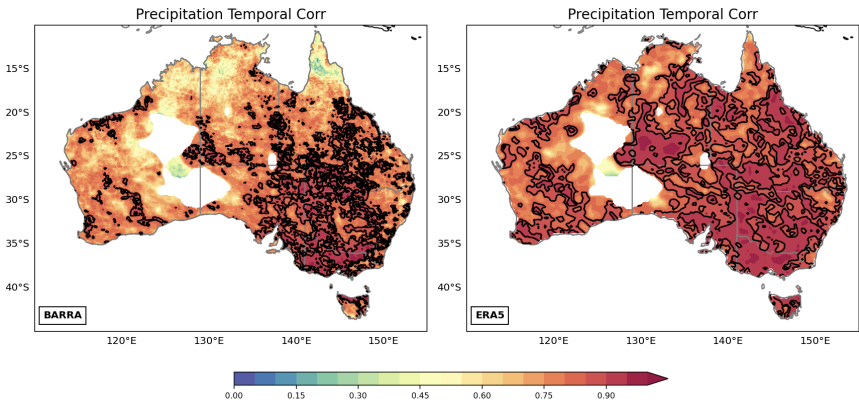


Figure 2 Annual mean precipitation of AGCD, BARRA and ERA5 (upper panels) and annual mean biases between BARRA/ERA5 and AGCD (lower panels). The regions with low density of station observations in AGCD has been masked and not considered in all subsequent evaluation. Unit: mm/day. Stippling indicates areas with biases that are statistically significant at 95% confidence level.

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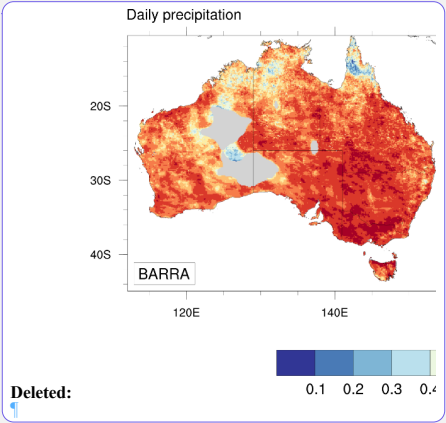


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793 Figure 3 Temporal correlation coefficient of annual precipitation between BARRA/ERA5
794 and AGCD. A black contour at value 0.85 has been added for reference.

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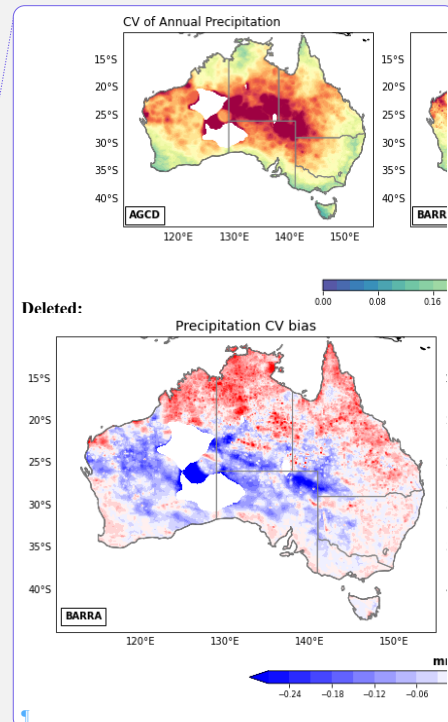
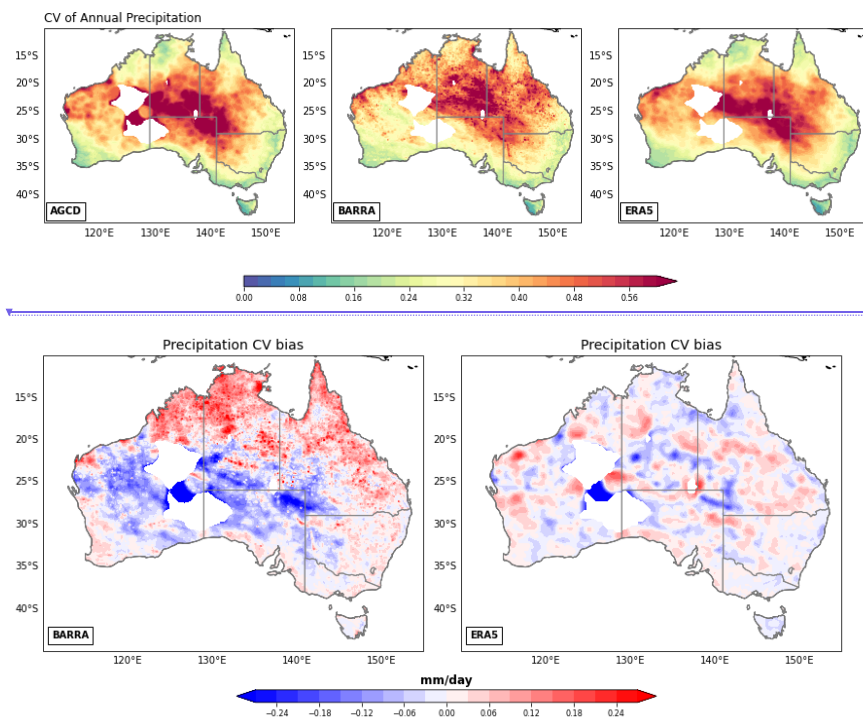
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803 Figure 4 CV of annual precipitation for AGCD, BARRA and ERA5 (upper panels) and biases
804 in CV between BARRA/ERA5 and AGCD (lower panels).

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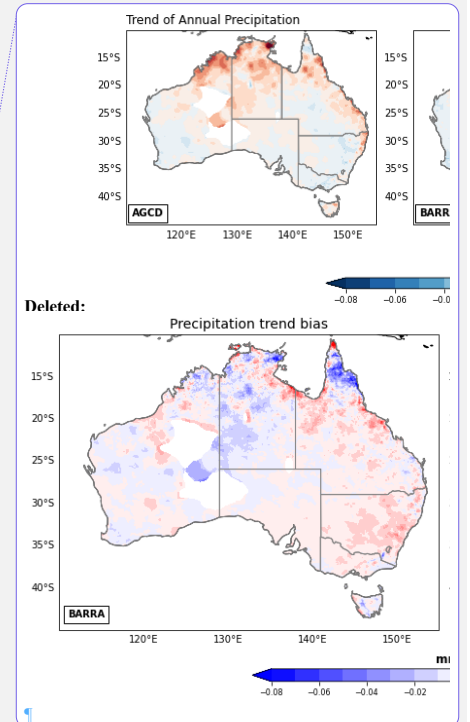
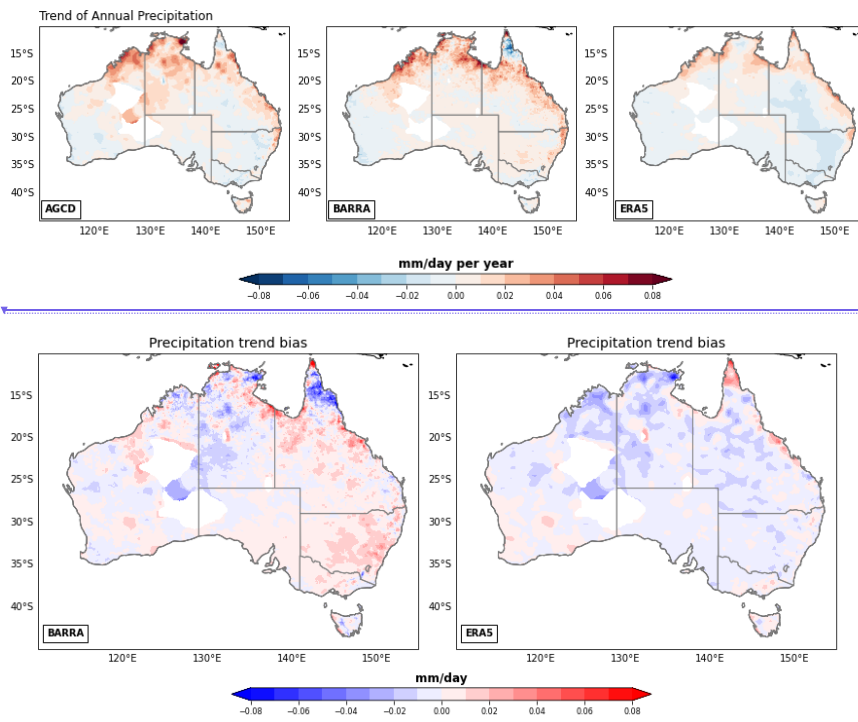
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814 Figure 5 Trend of annual precipitation for AGCD, BARRA and ERA5 (upper panels) and
815 biases in trend between BARRA/ERA5 and AGCD (lower panels).

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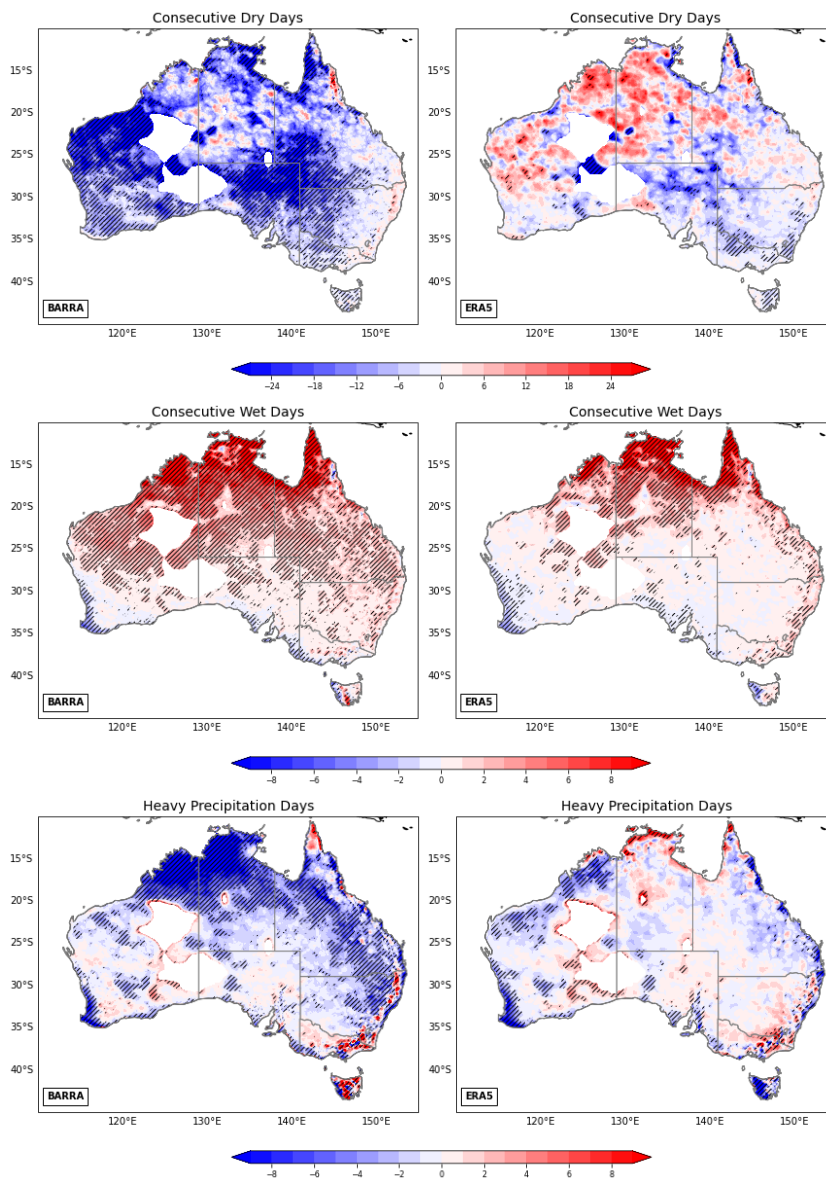
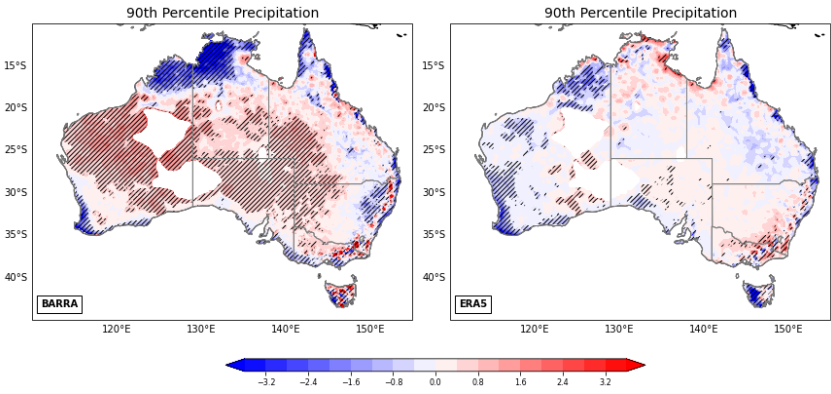


Figure 6 Biases in CDD, CWD, R10mm, R90p, R99p and Rx1Day in BARRA (left column) and ERA5 (right column). Stippling indicates areas with biases that are statistically significant at 95% confidence level. A black contour at 40% has been added to the panels for Rx1day (last row) for reference.

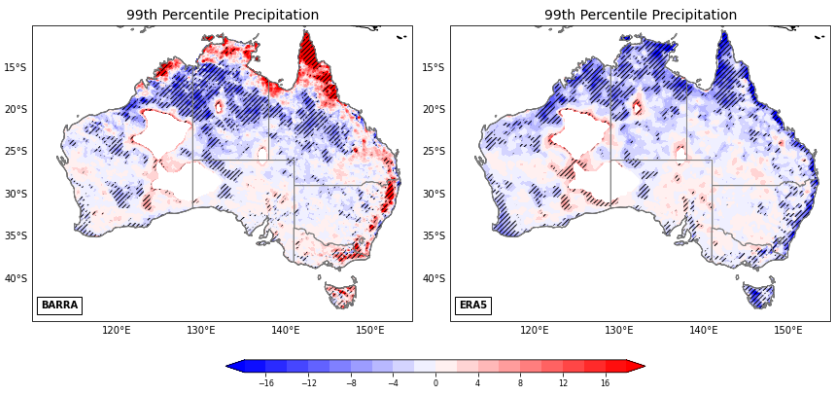
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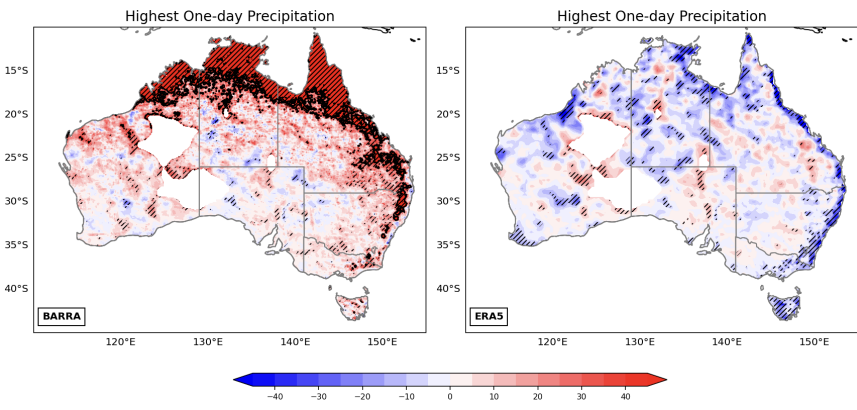
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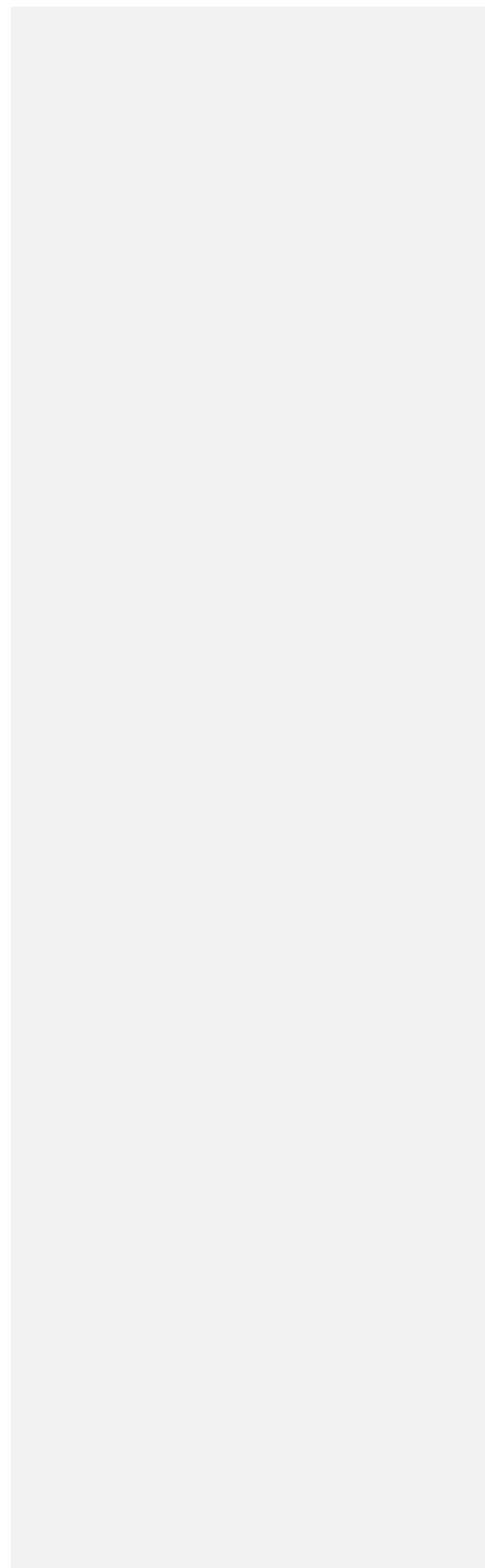


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Figure 6 (continued).

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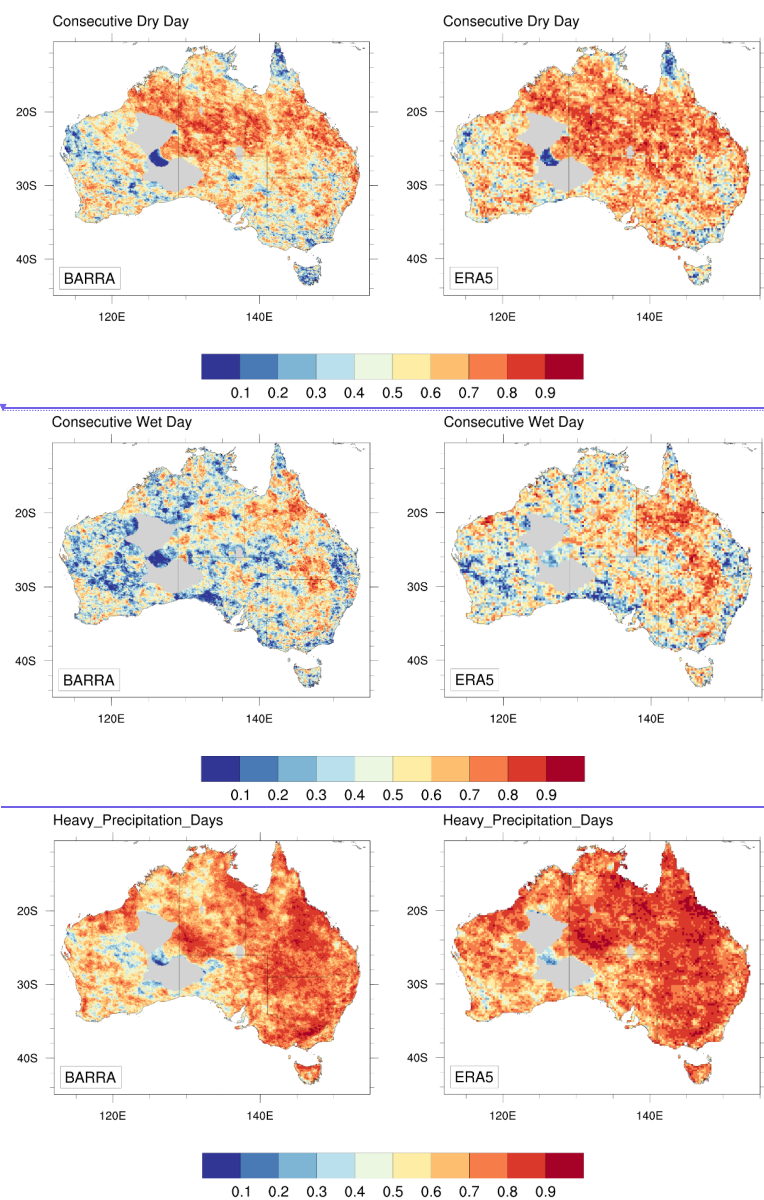
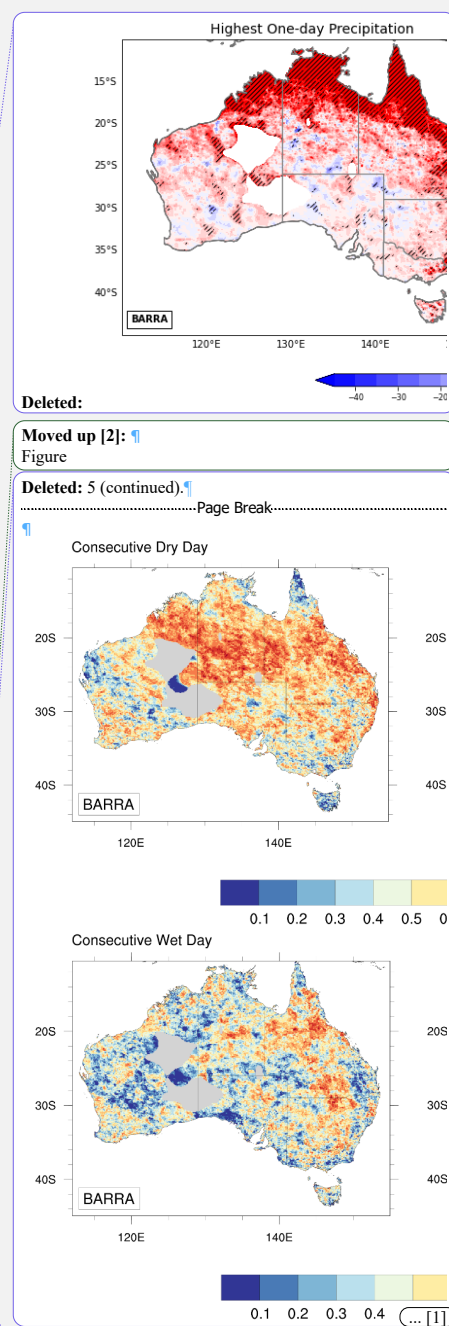


Figure 7 Temporal correlation of CDD, CWD, R10mm, R90p, R99p and Rx1Day between BARRA and AGCD (left column) and between ERA5 and AGCD (right column).



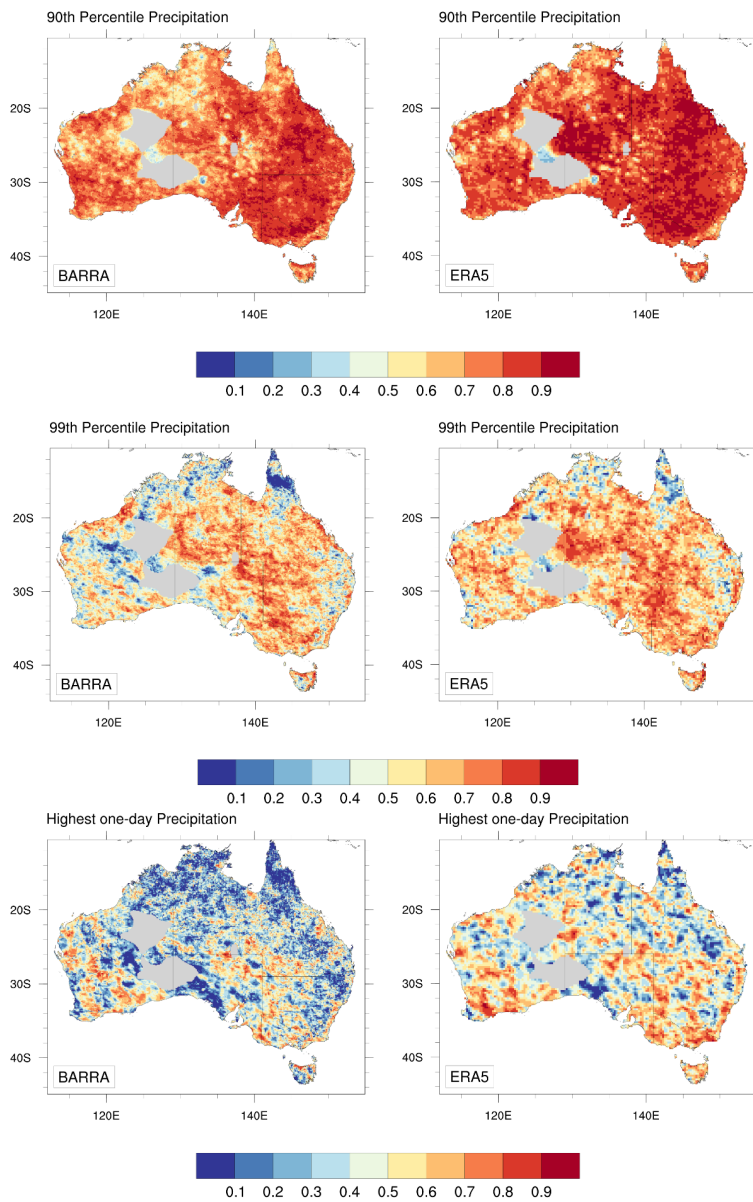
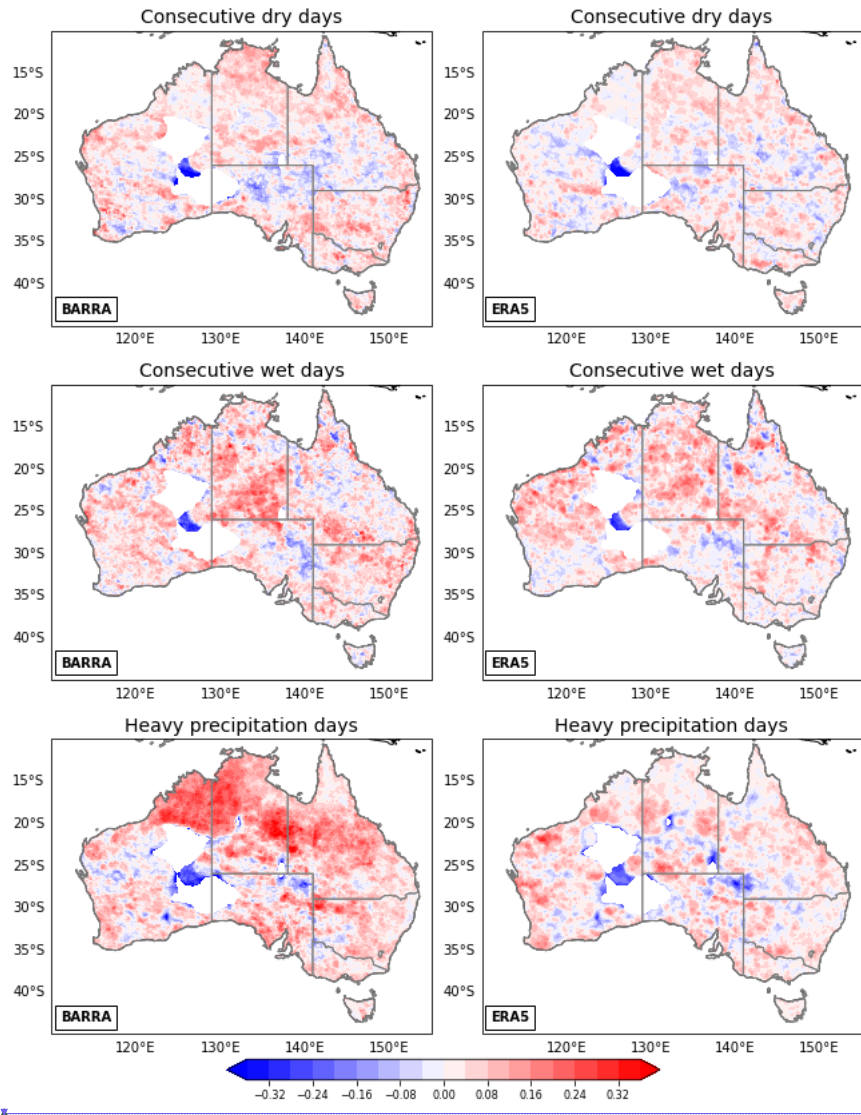


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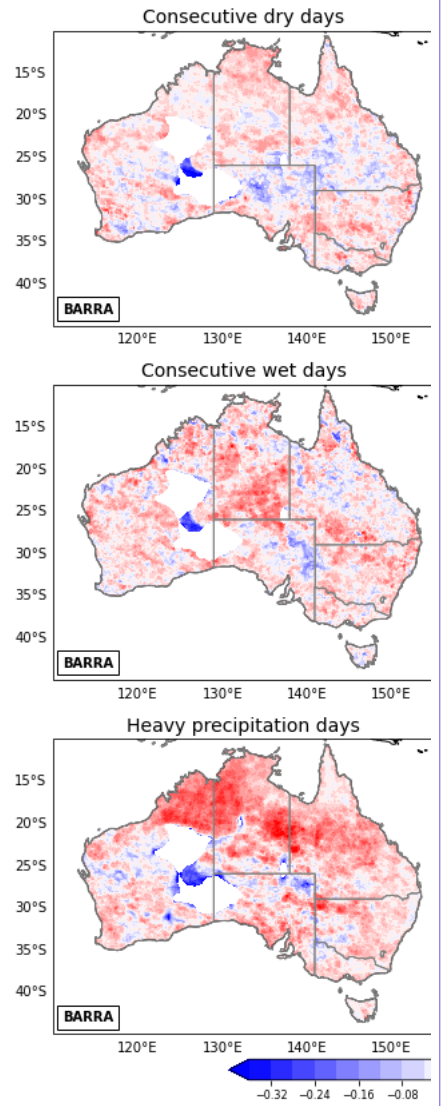
868 Figure 8 Biases in CV of CDD, CWD, R10mm, R90p, R99p and Rx1Day for BARRA (left
869 column) and ERA5 (right column) relative to AGCD.

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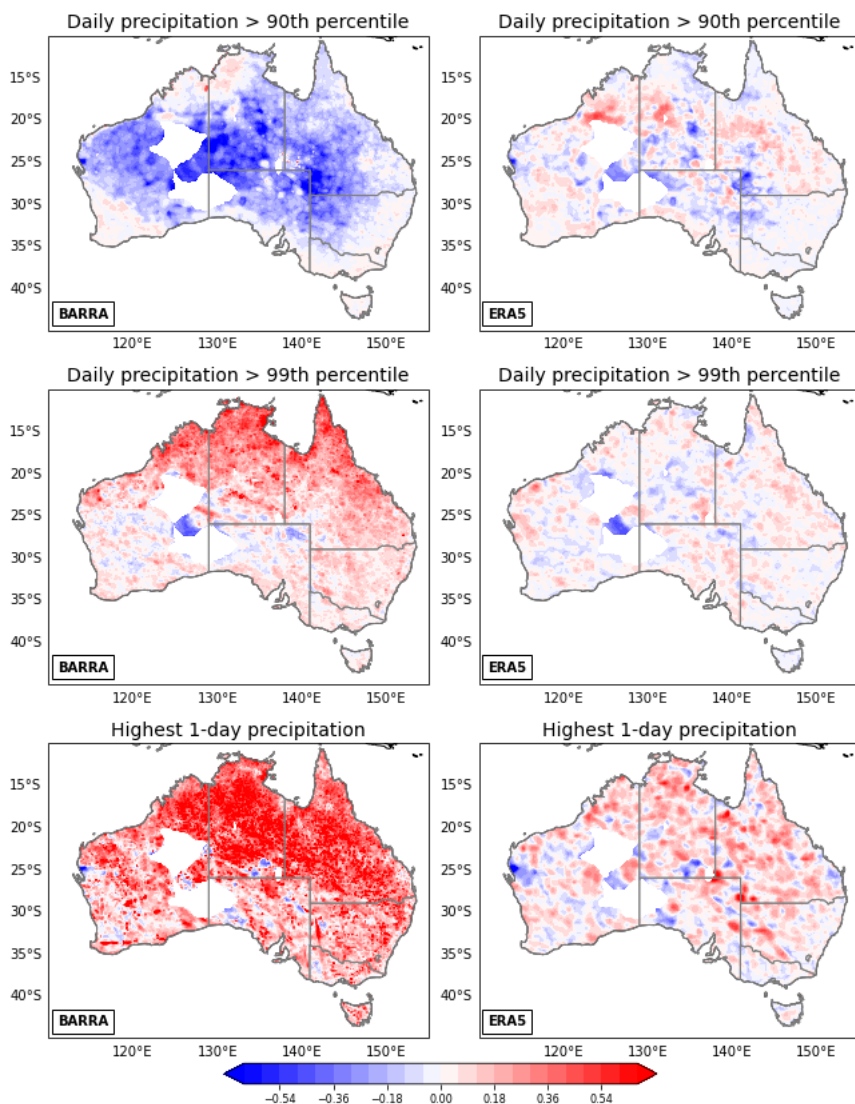
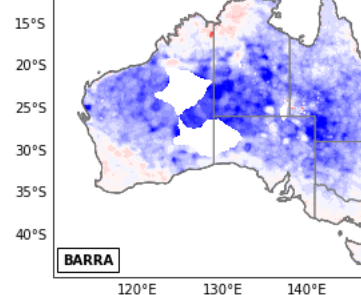


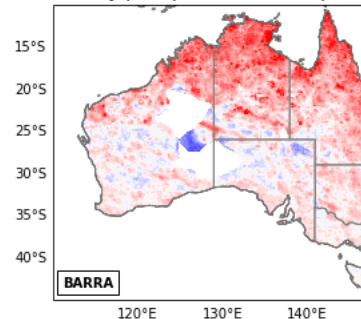
Figure 8 (continued)

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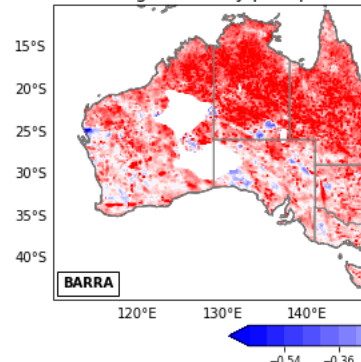
Daily precipitation > 90th perc



Daily precipitation > 99th perc

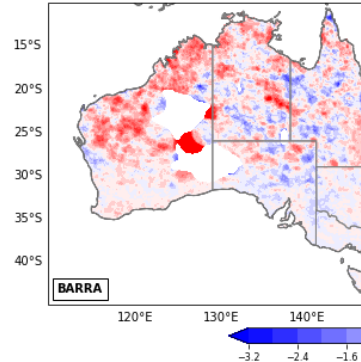


Highest 1-day precipitation

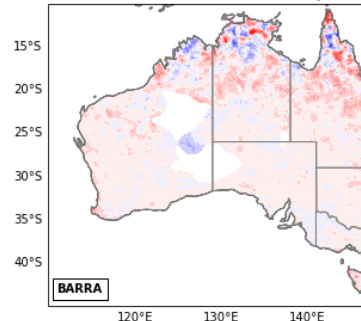


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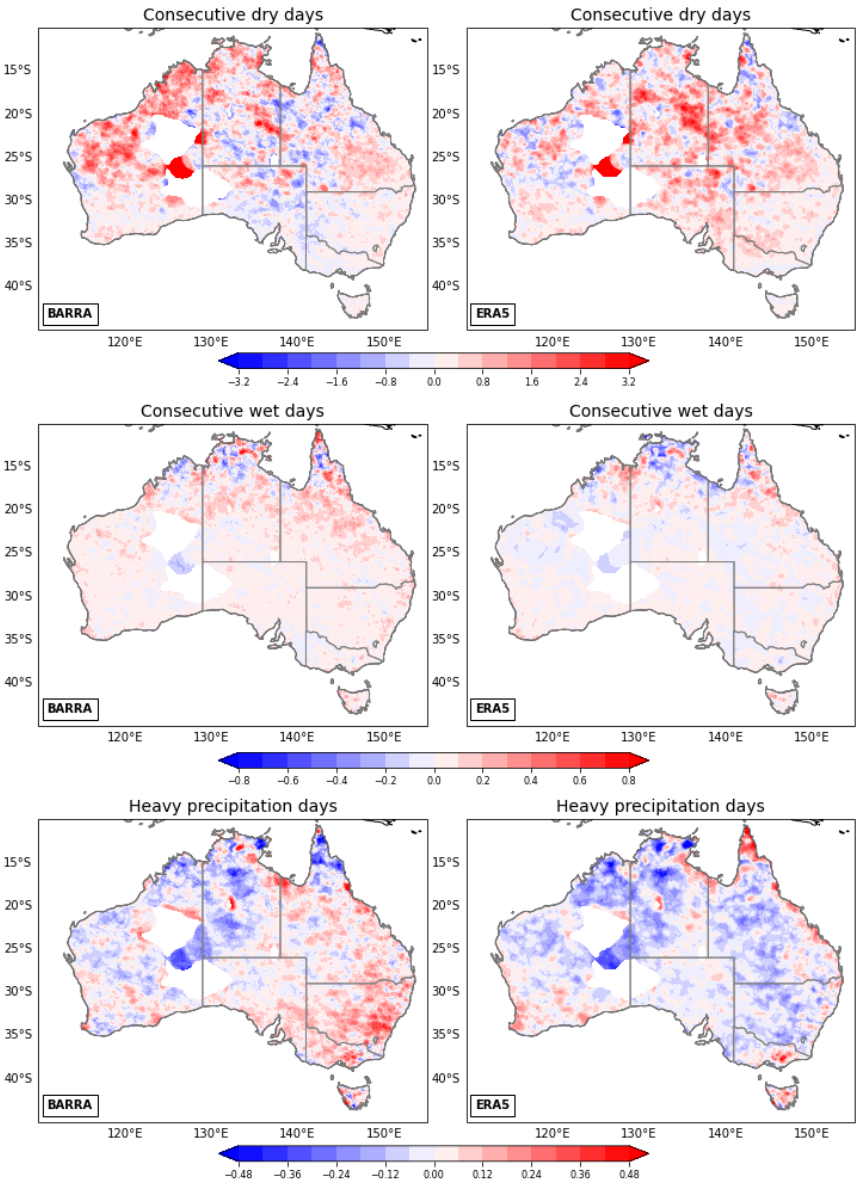
Consecutive dry days



Consecutive wet days



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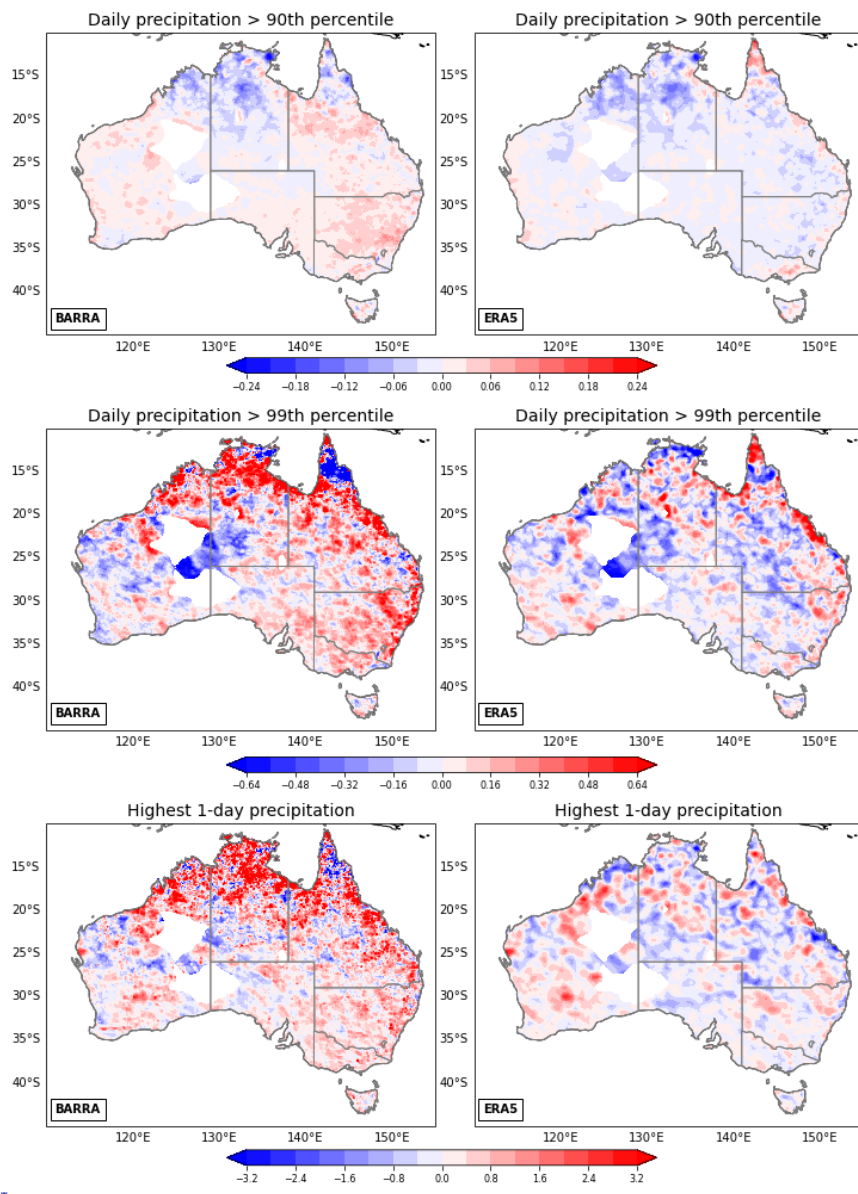
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Figure 9 Biases in trends of CDD, CWD, R10mm, R90p, R99p and Rx1Day for BARRA (left column) and ERA5 (right column) relative to AGCD.

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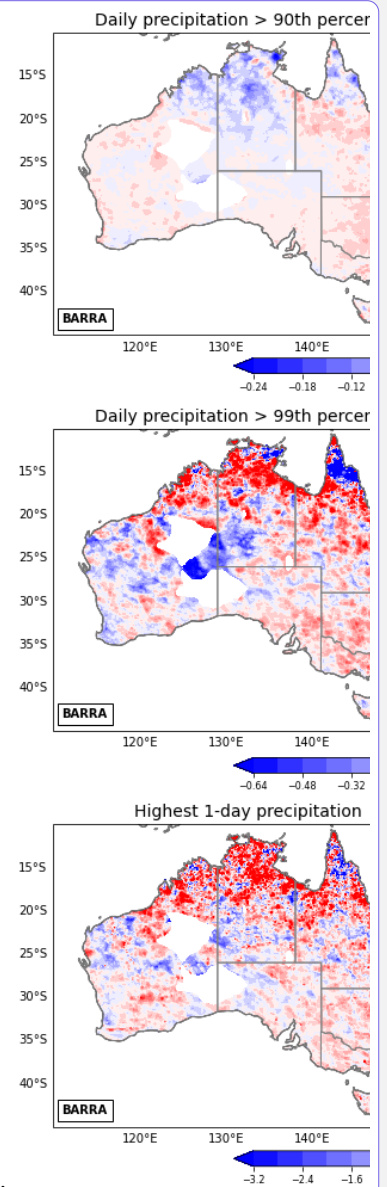
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Figure 9 (continued).

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