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2       **Comparison of BARRA and ERA5 in Replicating Mean and Extreme**

3       **Precipitation over Australia**

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## 37 **Abstract**

38 Reanalysis datasets are critical in climate research and weather analysis, offering consistent  
39 historical weather and climate data crucial for understanding atmospheric phenomena, and  
40 validating climate models. However, biases exist in reanalysis datasets that would affect their  
41 applications under circumstances. This study evaluates BARRA, which is a high-resolution  
42 reanalysis for the Australian region, and ERA5 in simulating mean precipitation and six  
43 selected precipitation extremes for their climatology, temporal correlation, coefficient of  
44 variation and trend. Both datasets reproduce spatial patterns of mean precipitation well with  
45 minor biases. ERA5 shows stronger temporal correlations, superior inter-annual precipitation  
46 accuracy, and lower biases in coefficient of variation compared to BARRA, especially in  
47 Northern Australia. However, both models exhibit substantial biases in trend, underestimating  
48 increasing trends in Northern Australia. ERA5 underestimates dry days and heavy rainfall,  
49 while BARRA tends to overestimate these extremes. Temporal correlations for extreme  
50 precipitation indices are weaker compared to mean annual precipitation. Notable differences  
51 exist in variability biases, with BARRA showing larger biases, especially for heavy  
52 precipitation in inland regions and Northern Australia. While both datasets replicate the main  
53 trends, biases persist. Overall, the evaluation results support application of both datasets for  
54 climatology analyses, but caution is advised for variability and trend analyses, particularly for  
55 specific extremes.

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57 **Key words:** BARRA, ERA5, extreme indices, temporal correlation, coefficient of variation,  
58 trend

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60   **1. Introduction**

61           Reanalysis dataset is created by combining historical observational data from various  
62   sources, such as weather stations, satellites, buoys, and more, with modern data assimilation  
63   techniques and numerical models (Kalnay, et al. 1996; Saha, et al. 2010; Dee et al. 2011;  
64   Kobayashi et al. 2015, Poli et al. 2016; Hersbach 2020). The fundamental aim of reanalysis is  
65   to construct a uniform and coherent historical archive of various atmospheric and  
66   environmental parameters, such as temperature, humidity and wind patterns, on either a  
67   regional or a global scale.

68           These datasets are invaluable for climate studies, weather analysis and model validation  
69   as they provide a uniform representation of historical climate conditions. For instance,  
70   Quagraine et al. (2020) used five global reanalysis datasets (European Centre for Medium-  
71   Range Weather Forecasts Reanalysis ERA-Interim, Dee et al. 2011; ERA5, Herbach et al. 2020;  
72   JRA-55, Kobayashi et al. 2015); MERRA2, (Gelaro et al. 2017); and NCEP-R2, Kanamitsu et  
73   al. 2002) to investigate the variability of West African summer monsoon precipitation, showing  
74   all datasets could represent the average rainfall patterns and seasonal cycle. Dai et al. (2023)  
75   utilized ERA5 data to estimate rainfall erosivity on the Chinese Loess Plateau, finding rainfall  
76   erosivity derived from ERA5 was highly consistent with those derived from the meteorological  
77   stations. Cheung et al. (2023) employed ERA5 to evaluate storm conditions in regional climate  
78   simulations, demonstrating regional climate models can capture climatology of measurements  
79   of storm severity over land including their spatial patterns and seasonality. Numerous studies  
80   have used reanalysis datasets as inputs for regional climate models (RCMs) to evaluate the  
81   models' capability in replicating observed climatic patterns (Solman et al., 2013; Ji et al., 2016;  
82   Fita et al., 2016, Di Virgilio et al., 2019; Capecchi et al., 2023; Di Virgilio et al., 2024; Ji et al.,  
83   2024).

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88 While reanalysis datasets provide valuable insights into historical weather and climate  
89 conditions, they have limitations and uncertainties, given that they are modelled outputs rather  
90 than direct observations. Many studies have evaluated reanalysis data across various variables  
91 and regions. For instance, Betts et al. (2019) assessed ERA5 biases in near-surface variables  
92 over Canada, highlighting its improved performance over ERA-Interim (Dee et al. 2011),  
93 though precipitation biases remained significant. Similarly, Hu and Yuan (2021) and Jiang et  
94 al. (2021) found that ERA5 precipitation accurately captured rainfall pattern over the Eastern  
95 Tibetan Plateau and mainland China, but under-estimated intensity. Izadi et al. (2021) found  
96 ERA5 performed better at monthly and seasonal timescales in Iran, underestimating coastal  
97 summer precipitation and overestimating it in mountains. Jiao et al. (2021) and Qin et al. (2021)  
98 found ERA5 overestimated summer precipitation and frequency in China but underestimated  
99 intensity during the warm season. Lei et al. (2022) and Shen et al. (2022) noted ERA5's  
100 limitations in simulating extreme precipitation events in China, especially for high-end  
101 extremes.

102 Comparisons between reanalysis datasets have also been conducted. Wang et al. (2019)  
103 found that both ERA5 and ERA-Interim exhibited warm biases over Arctic Sea ice, with larger  
104 biases in cold season than warm season. Lei et al. (2020) showed ERA5 improved cloud cover  
105 simulation over eastern China but not over the Tibetan Plateau, when compared to ERA-  
106 Interim. Gleixner et al. (2020) found ERA5 reduced biases in temperature and precipitation  
107 over East Africa compared to ERA-Interim but still struggled with long-term trends. Song and  
108 Wei (2021) found both ERA5 and MERRA-2 (Gelaro et al. 2017) captured night precipitation  
109 peaks over North China, but only ERA5 accurately reflected the afternoon peak. Li et al. (2022)  
110 concluded that ERA5 performed better than ERA-Interim, JRA55 (Kobayashi et al. 2015), and  
111 MERRA-2 in capturing precipitation over the Poyang Lake Basin. A summary of the above  
112 literature review can be found in Table S1.

113 In Australia, reanalyses like NCEP (Kalnay et al., 1996), JRA-55 (Kobayashi et al.,  
114 2015), ERA-Interim (Dee et al., 2011), and ERA5 (Hersbach et al., 2020) are commonly used,  
115 alongside the Australian Bureau of Meteorology's high-resolution (12 km) BARRA reanalysis.  
116 BARRA covers Australia, New Zealand, and Southeast Asia (Su et al., 2019), while BARRA-  
117 C offers even higher-resolution (1.5 km) analysis for four capital cities (Su et al., 2021).

118 May et al. (2021) found BARRA reliable, though it showed seasonal and diurnal biases.  
119 Other studies, like Pirooz et al. (2021), compared BARRA with global reanalyses, concluding  
120 BARRA performed better for precipitation and temperature in New Zealand but lagged behind  
121 ERA5 for high gust winds. Du et al. (2023) used BARRA for estimating daily precipitation in  
122 ungauged Australian catchments, while Hobeichi et al. (2023) employed BARRA to train  
123 statistical models for downscaling. Acharya et al. (2019, 2020) found BARRA's precipitation  
124 performance varied by region, with poorer results in tropical areas. Nishant et al. (2022)  
125 suggested higher resolution in BARRA-C didn't always improve precipitation simulations,  
126 while Choudhury et al. (2023) noted ERA5 performed better for mean temperatures than  
127 extremes in Australia. These previous studies on BARRA and BARRA-C have also been  
128 summarized in Table S1.

129 However, there is a gap in the existing studies concerning the intercomparison of  
130 various reanalyses, such as BARRA and ERA5, specifically in relation to precipitation  
131 extremes over Australia. In this study, we aim to bridge this gap by evaluating and comparing  
132 the performance of BARRA and ERA5 in capturing precipitation extremes. While the  
133 traditional evaluation methods focusing on climatology (long-term mean), here we also include  
134 temporal correlation, coefficient of variation and trend in evaluation to quantify their overall  
135 performance, which have not been examined before in previous studies. By assessing climate  
136 means and extremes and quantifying their biases, this study provides a valuable reference for  
137 selecting appropriate datasets for specific applications and cautions against treating reanalysis

138 data as observations. The paper is organized as follows: Section 2 introduces the reanalysis  
139 datasets and observational data used for evaluation. Section 3 outlines the climate extreme  
140 indices and evaluation methodology. Results are presented in Section 4, followed by further  
141 discussion in Section 5. Finally, Section 6 offers a summary and conclusions.

142

## 143 **2. Data**

### 144 **2.1 ERA5**

145 ERA5 is a global atmospheric reanalysis dataset developed by ECMWF (Hersbach, et  
146 al. 2020). ERA5 provides hourly estimates of many atmospheric, land, and oceanic climate  
147 variables. The data is on a ~30 km horizontal grid and resolves the atmosphere using 137 levels  
148 from the surface up to a height of 0.01hPa (~80 km).

149 ERA5 is constructed upon the foundation of the Integrated Forecasting System (IFS)  
150 Cy41r2. This allows ERA5 to benefit from a decade's worth of development in areas such as  
151 model physics, core dynamics, and data assimilation techniques. ERA5 is a significant  
152 advancement over its predecessors (e.g., ERA-Interim) due to its higher spatial and temporal  
153 resolution, improved assimilation techniques, and more sophisticated modelling components.  
154 It provides a detailed and accurate representation of various atmospheric variables, such as  
155 temperature, humidity, wind speed, pressure, and more. The dataset covers the entire globe and  
156 spans from 1940 to the present, making it valuable for various applications in climate research,  
157 meteorology, environmental science, and more.

### 158 **2.2 BARRA**

159 BARRA is a high-resolution regional atmospheric reanalysis dataset developed by the  
160 Australian Bureau of Meteorology, which is available from January 1990 to February 2019 (Su,  
161 et al. 2019). BARRA was constructed based on the Australian Community Climate Earth-  
162 System Simulator (ACCESS) model with assimilation of a wide range of observational data to

163 create a coherent and consistent representation of past weather and climate conditions. BARRA  
164 covers the Australian continent, New Zealand, part of Asia and some Pacific Islands with a  
165 horizontal resolution of 12 km and 70 vertical levels from the surface up to a height of 80 km.  
166 BARRA specifically focuses on providing detailed information about weather patterns and  
167 atmospheric variables over the Australian region, which provides about 100 parameters at  
168 hourly intervals.

169 The ACCESS model, which was applied to generate BARRA, originated from the  
170 UKMO's Unified Model (UM), which can be configured in global mode or regional mode. For  
171 regional simulations, the global version of ACCESS becomes ACCESS-R. ACCESS-R was  
172 initialized by ERA-Interim reanalysis data, which also provides boundary conditions during  
173 simulation. A series of observations have been assimilated into BARRA, including land and  
174 ship (buoy) synoptic observations, upper-air observations from radiosondes and wind profilers,  
175 satellite derived radiances and winds (Su et al. 2019). However, no precipitation observations  
176 were directly assimilated.

### 177 **2.3 AGCD**

178 The observational data in the study are from the Australian Gridded Climate Dataset  
179 (AGCD, Evans et al. 2020). The daily gridded maximum and minimum temperatures, and  
180 precipitation data has a spatial resolution of 0.05° (~5km) and is interpolated from observations  
181 at stations across the Australian continent. Most of those stations are in the more heavily  
182 populated coastal regions with far fewer stations inland and over high elevation areas. For  
183 example, there are very few station observations near the Gibson desert region in Western  
184 Australia, making the gridded observations unreliable over that region. Thus, in the following  
185 figures that region has been masked and not considered for evaluation. Since observations and  
186 reanalyses are not at the same spatial resolutions, we aggregate the observations to the native  
187 grid of ERA5 and BARRA respectively for comparison, including the performance of

188 statistical significance tests. For comparison purpose, we also interpolate reanalysis to AGCD  
 189 grids using the conservative area weighted re-gridding scheme from the Climate Data  
 190 Operators (Schulzweida et al., 2006), which will be shown in the Supplementary Information.  
 191 The states and sub-regions in the Australian region we discuss in the following can be found  
 192 in Figure S1.

193

### 194 3. Methodology

#### 195 3.1 ET-SCI

196 While extreme climate and weather events are generally multifaceted phenomena, in  
 197 this study we evaluate climate extremes based on daily precipitation and temperature as defined  
 198 by Expert Team on Sector-specific Climate Indices (ET-SCI; Alexander & Herold, 2015;  
 199 Herold and Alexander, 2016). We use the ClimPACT version 2 software to calculate the ET-  
 200 SCI indices (<https://climpact-sci.org/>), focussing on daily precipitation.

201 Although ClimPACT generates 14 precipitation-related core indices, we select seven  
 202 (Table 1) based on the following considerations: 1) To capture key aspects of climate extremes,  
 203 and 2) to capture extremes which have impacts on society and infrastructure such as agriculture,  
 204 water resources and economy (Tabari, 2020; Pei et al., 2021). Accordingly, we include absolute  
 205 indices such as the maximum 1-day precipitation (Rx1day) and total precipitation (PRCPTOT),  
 206 a threshold-based index (e.g., number of heavy rain days, R10mm), percentile indices (e.g.,  
 207 total annual precipitation from very heavy rain days, R99p), and duration indices such as the  
 208 consecutive wet (CWD) and dry days (CDD).

209 With the above consideration, the seven aforementioned precipitation-related indices  
 210 were calculated on native reanalysis grids and observation grids. While the availability of  
 211 AGCD and ERA5 starts much earlier, the analysis period is 1990–2018, which is the duration  
 212 of BARRA. Since the AGCD observations have the highest resolution, here we mainly show

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221 the evaluation on the native grids of the reanalyses (i.e., the 12-km grid of BARRA and 30-km  
 222 grid of ERA5). The extreme indices calculated from reanalysis data have also been regridded  
 223 to the 5-km resolution using bilinear interpolation, which are included in the supplementary  
 224 information to demonstrate that our conclusions are insensitive to the choice of evaluation  
 225 resolution.

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### 227 3.2 Evaluation metrics

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228 We evaluate BARRA and ERA5 for their performance in capturing climatology (29  
 229 years in our case), coefficient of variation (CV), temporal correlation, and trends of seven  
 230 selected precipitation extreme indices. The CV is a valuable statistical tool representing the  
 231 ratio of the (yearly) standard deviation to the mean, allowing for the comparison of variation  
 232 between different data series, even when their means differ significantly. Temporal correlations,  
 233 which are computed at an annual time step, of climate extremes measure the similarities  
 234 between simulations and observations in terms of their inter-annual variabilities, with larger  
 235 temporal correlations indicating better performance. For trend analyses, we applied simple  
 236 linear trend line fitting to the yearly time series of climate indices. All the above metrics are  
 237 computed at each grid point in the datasets' native grids as well as the AGCD grid after re-  
 238 gridding. Differences between BARRA/ERA5 and AGCD then form the bias maps. After  
 239 averaging over all grid points, the domain averages will then be discussed in the following.

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240 We use bias and domain-averaged absolute bias to quantify spatial differences between  
 241 reanalyses and observations. Temporal correlation, coefficient of variation, and trend are used  
 242 to quantify temporal similarities between reanalyses and observations. The non-parametric  
 243 Mann-Kendall test is used to assess the statistical significance of differences and trends. Biases  
 244 are assessed at an annual timescale for all extremes.

245

## 4. Results

### 4.1 Mean climate

This section evaluates and compares the annual mean of daily precipitation between BARRA and ERA5 against AGCD over Australia.

#### 4.1.1 Bias and temporal correlation

We first evaluate precipitation simulated by BARRA and ERA5 against observations (AGCD). The mean annual precipitation from the three datasets and biases in BARRA and ERA5 compared to AGCD are shown in Figure 1 (and Figure S2 on the observation grid). Results show that both BARRA and ERA5 simulate the spatial patterns of mean annual precipitation very well with high rainfall in northern Australian, eastern Australia coast and western Tasmania and low rainfall inland, albeit with clear biases. Compared to AGCD, both BARRA and ERA5 underestimate precipitation up to 20% for Eastern Australian coast, southwest western Australia, and western Tasmania, but overestimate annual precipitation up to 30% inland (Figure S3). Some clear differences in biases between BARRA and ERA5 can be observed in central western Australia and northern Queensland where BARRA overestimate precipitation but ERA5 underestimate it. Domain averaged absolute bias in annual precipitation is about 0.17mm/day (~12.7% relative bias with respect to domain average) for BARRA and 0.15 mm/day (~10.5% relative bias) for ERA5 (Table 2).

The skill of simulated precipitation from BARRA and ERA5 are further demonstrated in the temporal correlations between BARRA/ERA5 and AGCD shown in Figure 2 (and Figure S4 on the observation grid). Temporal correlation of annual precipitation is larger in southeast Australia and northern Tasmania for both BARRA and ERA5, which is above 0.85. This indicates inter-annual variability of precipitation is well captured by BARRA and ERA5. In contrast, temporal correlation is weaker for western inland and northern Australia. ERA5 generally has larger temporal correlation when compared with BARRA, especially for northern

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277 Australia, where temporal correlation for BARRA is below 0.5. On average, temporal  
278 correlation for ERA5 is 0.85, which is larger than 0.77 for BARRA (Table 2).

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#### 279 4.1.2 CV (coefficient of variation) and trend

280 CV of annual precipitation for AGCD and biases between BARRA/ERA5 and AGCD  
281 are presented in Figure 3 (and Figure S5 on the observation grid). By its definition, CV helps  
282 capture the standard deviation in the dataset relative to its mean. In the observation, CV is  
283 generally smaller for coastal regions including Tasmania except for northwest West Australia  
284 and Tasmania than inland Australia, where annual rainfall is much smaller than coastal regions.

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285 Alternatively, regions with higher annual precipitation generally have smaller CV. Both  
286 BARRA and ERA5 reasonably capture the main feature of CV in observation. However, clear  
287 biases can be observed, especially in BARRA that has more than 50% large positive biases in  
288 Northern Australia, up to 20% positive biases for inland, and relatively smaller biases for  
289 southeastern Australia, southwest West Australia and Tasmania. In contrast, ERA5 does not  
290 have a clear bias pattern, and biases are relatively smaller when compared to BARRA.

291 To further investigate the variability evident in observations and BARRA/ERA5  
292 simulations, we assess the trends in annual precipitation (Figure 4 and Figure S6 on the  
293 observation grid). AGCD shows strong increasing trends over Northern Australia and  
294 Northeast Australia coastal regions but decreasing trends over Northern Queensland,  
295 southwestern West Australia and southern Great Dividing Range including Victoria, although  
296 not all trends are significant. Most of inland regions have relatively small trend in annual  
297 precipitation. Both BARRA and ERA5 reproduce the major trend pattern reasonably well,  
298 however, clear biases can be observed over Northern Australia where both BARRA and ERA5  
299 underestimate trend more than 100% (i.e., trend of 0.08 mm/day per year with bias of similar  
300 magnitude). BARRA overestimated decreasing trend over Northern Queensland but ERA5  
301 underestimate it (even increasing trend instead).

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305 In summary, evaluation of annual mean precipitation indicates both BARRA and ERA5  
306 possess small biases (~20%) in the spatial precipitation patterns. ERA5 shows stronger  
307 temporal correlations than BARRA, particularly in northern Australia. Overall, ERA5  
308 demonstrates higher accuracy in capturing inter-annual precipitation variability. Both BARRA  
309 and ERA5 captured spatial distribution of coefficient of variation reasonably well but with  
310 large biases (~ 50%). BARRA shows much larger biases than ERA5 especially for Northern  
311 Australia. Both BARRA and ERA5 roughly reproduce the pattern of trend but with very large  
312 biases (~100%), especially for Northern Australia where both substantially underestimate the  
313 increasing trend.

314

## 315 4.2 Climate extremes

316 This section evaluates the seven select precipitation extreme indices (Table 1) from  
317 BARRA and ERA5 over Australia by comparing them against AGCD. Evaluations are  
318 performed primarily using spatial bias maps and temporal correlations. We also assess the  
319 interannual variability and trends in the simulated BARRA and ERA5 indices and compare  
320 these with AGCD to further investigate any discrepancies.

### 321 4.2.1 Bias and temporal correlation

322 Annual mean biases in six precipitation extremes are shown in Figure 5 (and Figure S8  
323 on the observation grid). For duration-related extremes, (CDD and CWD), there is a clear north-  
324 to-south gradient in AGCD (Figure S7) with longer duration of CDD and CWD in northern  
325 Australia than southern Australia (CWD also has a clear west-to-east gradient in Tasmania),  
326 which is well simulated in BARRA and ERA5 (Figure S7). While the spatial distributions are  
327 well captured, clear biases are evident in them (Figure 5). BARRA generally underestimates  
328 CDD especially for central inland and northwest West Australia where biases are up to 40%.  
329 ERA5 also under-estimates CDD for central inland, but in contrast its over-estimates CDD for

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333 most of northwestern Australia, overall ERA5 has smaller absolute bias in CDD (6.9 days) than  
334 BARRA (14.5 days) (Table 2). Both BARRA and ERA5 have similar bias pattern for CWD,  
335 which generally overestimate CWD over most of regions except for southern Australian coast,  
336 southwest West Australia and western Tasmania. The positive biases over Northern Australia  
337 may reach 30%. Overall BARRA has slightly larger biases in CWD (2.3 days) than ERA5 (1.7  
338 days) (Table 2).

339 For threshold-based extremes (PRCPTOT, R10mm, R90p, R99p, Rx1day), both  
340 BARRA and ERA5 also generally match the spatial distribution of heavy precipitation days  
341 and R90p (Figure S7) in AGCD with large values in Northern Australia, eastern seaboard and  
342 Australian Great Dividing Range, and western Tasmania. However, clear biases can be  
343 observed in BARRA and ERA5 for both R10mm and R90p (Figure 5). BARRA and ERA5  
344 have large negative biases in R10mm over Northern Australia, eastern seaboard, southwest  
345 Western Australia and western Tasmania, but biases in central inland and northwest West  
346 Australia are generally small. Overall, domain averaged absolute bias for ERA5 (1.7 days) is  
347 about half of that for BARRA (3.3 days). Both BARRA and ERA5 also have relatively large  
348 negative biases in R90p for most of northern Australia, eastern coasts, southwest West  
349 Australia and western Tasmania but small positive biases inland, especially for BARRA.  
350 Overall averaged absolute bias is 0.78 mm/day for BARRA and 0.44 mm/day for ERA5 (Table  
351 2).

352 BARRA and ERA5 also reasonably captured the spatial patterns of R99p and Rx1day,  
353 however, quite large biases are in BARRA and ERA5 (Figure 5). BARRA generally  
354 overestimate R99p and Rx1day over northern Australia coasts and along the Great Dividing  
355 Range. In contrast, ERA5 generally underestimate R99p and Rx1day over northern and eastern  
356 coasts, southwest Western Australia and western Tasmania. The domain averaged bias in R99p

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360 is at similar magnitude for BARRA (4.09 mm/day) and ERA5 (3.67 mm/day), however biases  
361 in Rx1day is much larger for BARRA (20.3 mm/day) than ERA5 (7.9 mm/day) (Table 2).

362 Figure 6 (and Figure S9 on the observation grid) presents the temporal correlations  
363 between BARRA/ERA5 and AGCD for the six precipitation extreme indices. Unlike the strong  
364 temporal correlation between BARRA/ERA5 and AGCD for mean annual precipitation (Figure  
365 2), the temporal correlations for these extreme indices are worse except for R90p (Figure 6).  
366 For extremes like R10mm and R90p, the correlation ranges from reasonably good (above 0.6)  
367 to pretty good (above 0.8) between BARRA/ERA5 and AGCD for most of the domain.  
368 Temporal correlation for CDD, CWD and R99p are not as good as R10mm and R99p. CDD  
369 has more regions with stronger correlations (0.5-0.6) or above than CWD and Rx1day. for the  
370 latter correlation is about ~0.5 or less over most of the domain. Compared to BARRA, ERA5  
371 has slightly stronger temporal correlations for those extremes (Table 2).

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#### 373 4.2.2 CV (coefficient of variation) and trend

374 The observed and simulated CV of precipitation extremes and biases in their CV for  
375 BARRA and ERA5 are shown in Figure S10 and Figure 7 (and Figure S11 on the observation  
376 grid), respectively. Generally, both BARRA and ERA5 have similar CV bias patterns and  
377 magnitude for CDD, CWD and R10mm. In contrast, BARRA is quite different from ERA5 for  
378 other three extremes. BARRA substantially under-estimated CV of R90p over most on inland  
379 regions but ERA5 has much smaller negative biases, even small positive biases, although both  
380 have small biases in CV of R90p along most coastal regions and Tasmania. BARRA  
381 systematically overestimate CVs of R99p and Rx1day over northern Australia but ERA5 has  
382 relatively small biases for them. Overall, BARRA has more than twice as much as CV biases  
383 in ERA5 for R90p, R99p and Rx1day (Table 2).

388 Trends of each of the precipitation extreme indices for the three datasets and biases in  
 389 trend for BARRA and ERA5 are shown in Figure S12 and Figure 8 (and Figure S13 on the  
 390 observation grid), respectively. Generally, both BARRA and ERA5 simulate the main pattern  
 391 of trends for those extremes but with large biases. BARRA and ERA5 simulated CDD trend  
 392 well for southern Australia but BARRA generally under-estimated trend in CDD over inland  
 393 Australia and overestimate trend in northwest Australia. ERA5 only has large positive trend  
 394 biases in northern central Australia. The overall domain averaged biases are similar between  
 395 BARRA (0.584) and ERA5 (0.566). Both BARRA and ERA5 have small biases in CWD in  
 396 central and southern Australia but similar biases pattern in Northern Australia. They also have  
 397 similar overall biases in CWD (0.064 for BARRA and 0.060 for ERA5). Both BARRA and  
 398 ERA5 under-estimated increasing trend in R10mm in northern Australia, but BARRA  
 399 overestimate trend in most of southeast Australia. In contrast, ERA5 under-estimate trend over  
 400 there. Overall, ERA5 has slightly larger biases (0.094) than BARRA (0.085). Like R10mm,  
 401 both BARRA and ERA5 also underestimate trend of R90p in most of northern Australia but  
 402 have small biases in central and southern Australia. They have almost the same overall biases  
 403 in R90p. BARRA/ERA5 has similar biases patterns for R99p and Rx1day but biases for  
 404 rx1days are much larger. Both BARRA and ERA5 have large biases in R99p and Rx1day but  
 405 biases in BARRA are generally larger than ERA5.

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406 In summary, both BARRA and ERA5 reproduce spatial patterns of extremes well but  
 407 display biases. ERA5 underestimates CDD and certain extreme precipitation indices (e.g.,  
 408 Rx1day), while BARRA tends to overestimate these extremes. Both reanalyses show  
 409 discrepancies in various precipitation indices across different regions, with BARRA generally  
 410 displaying larger biases compared to ERA5. Temporal correlations between BARRA/ERA5  
 411 and observations for extreme precipitation indices are weaker than those for mean annual  
 412 precipitation, except for a few indices where ERA5 demonstrates slightly stronger correlations

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415 compared to BARRA. Both BARRA and ERA5 align in CV patterns and biases for certain  
416 extremes (CV, R10mm, R90p) but differ notably in others (PRCPTOT, trend, CDD, R99p,  
417 Rx1day). BARRA significantly underestimates very heavy precipitation variability over inland  
418 regions, while ERA5 presents smaller biases or even positive biases in these areas. Additionally,  
419 BARRA tends to overestimate extreme precipitation variability in Northern Australia  
420 compared to ERA5. Overall, BARRA shows more than double the biases in variability  
421 compared to ERA5 for specific extreme precipitation indices. Both reanalyses generally  
422 simulate the main trend patterns but exhibit considerable biases. BARRA underestimates or  
423 overestimates trends in certain regions and indices, while ERA5 demonstrates different biases,  
424 including smaller biases overall compared to BARRA across these precipitation extremes.

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## 426 5. Discussion

427 In this study, we assessed the performance of BARRA and ERA5 in simulating mean  
428 precipitation and six selected precipitation extremes. While most previous evaluations have  
429 focused on the climatology of precipitation and its extremes, only a few studies have included  
430 the coefficient of variation (CV) (Teng et al., 2024). Our evaluation encompassed annual  
431 climatology, along with temporal correlation, CV, and trend analysis, providing a  
432 comprehensive assessment of the performance of these two reanalysis datasets.

433 The results indicate that both BARRA and ERA5 demonstrate reasonable skill in  
434 simulating mean precipitation and certain precipitation extremes (e.g., CWD and R90p).  
435 However, they encounter challenges in accurately reproducing temporal correlation, CV, and  
436 trends for certain extreme events, highlighting significant uncertainties in their representation  
437 of extremes.

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438 While acknowledging the capabilities of these reanalysis datasets, our study also  
439 identifies specific limitations and suggests potential directions for future research. A crucial

442 consideration in model evaluation is the accuracy of observational data, which substantially  
443 influences evaluation outcomes. In this study, we used the AGCD dataset as the observational  
444 benchmark, which is based on interpolating data from in-situ stations (Evans et al., 2020).  
445 However, the AGCD dataset presents several limitations: 1) Spatial coverage: Sparse station  
446 coverage in northwest and central Australia, and limited observations in high-elevation areas,  
447 result in a concentration of stations in southeastern Australia, southwestern Western Australia,  
448 and eastern Tasmania. The arid interior is notably underrepresented. 2) Data completeness and  
449 homogeneity: Incomplete and inhomogeneous observations due to missing data, changes in  
450 observational techniques, or station relocations can affect the consistency of the dataset. 3)  
451 Interpolation uncertainties: The interpolation method used in AGCD (splining), instead of the  
452 ordinary kriging method used in its predecessor (AWAP), introduces uncertainties, particularly  
453 in areas with sparse data coverage for extreme events like heavy rainfall.

454 These observational uncertainties may contribute to biases in the evaluation results. In  
455 particular, the limited number of monitoring sites over the Great Dividing Range and inland  
456 areas introduces significant uncertainties in estimated observed precipitation for these regions.  
457 Independent studies, such as Chubb et al. (2016), found that daily precipitation is  
458 underestimated by at least 15% in some areas, which could suggest similar underestimation in  
459 BARRA and ERA5 for these regions. Similarly, the sparse gauge network in northwestern  
460 inland areas might miss localized extreme precipitation events.

461 Our analysis focused on seven ET-SCI-defined precipitation extreme indices, (including  
462 mean precipitation), widely used in various evaluation studies (Nishant et al., 2020; Ji et al.,  
463 2024). However, recognizing the need for region-specific indices, we suggest future studies  
464 extend the analysis to incorporate additional extreme indices tailored to specific regions and  
465 applications.

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Our findings emphasize that while both BARRA and ERA5 are competent in simulating the climatology of mean climate, temporal correlation, and CV, challenges remain in accurately capturing trends, particularly for certain extremes. Notably, ERA5 shows better overall performance compared to BARRA. Although higher resolution often correlates with better performance, recent studies have shown that increasing resolution alone does not always guarantee improvements (Nishant et al., 2022). Considering the critical role of driving data, model physics, and data assimilation, it may be valuable to update BARRA using the latest ERA5 data along with improved model physics and data assimilation techniques to enhance its performance.

In this study, we evaluated ERA5 and BARRA on both their native resolutions and a common resolution (5 km) to match AGCD. The results showed that the evaluations were consistent across native and common resolutions, suggesting that the performance assessments were not highly sensitive to changes in resolution.

## 6. Summary and Conclusion

Reanalysis datasets play a crucial role in climate research, weather analysis, and various scientific investigations. Their ability to provide a consistent and comprehensive representation of historical weather and climate conditions makes them invaluable. These datasets are particularly essential for studying long-term climate trends, understanding atmospheric phenomena, and validating climate models.

In this study, we evaluate BARRA and ERA5 for their capabilities to simulate mean precipitation and six selected precipitation extremes for their climatology, temporal correlation, coefficient of variation (CV) and trend to quantify their overall performance. We evaluated BARRA and ERA5 at their native resolutions, as well as at a common resolution (i.e., the

492 observation resolution). Both analyses yielded consistent results, indicating that the evaluation  
493 is not sensitive to the remapping process.

494 The assessment of annual mean precipitation reveals that both BARRA and ERA5  
495 adeptly reproduce the spatial precipitation patterns, exhibiting minor biases of around 20%.  
496 Particularly, ERA5 showcases stronger temporal correlations compared to BARRA, especially  
497 evident in northern Australia. ERA5, overall, demonstrates superior accuracy in capturing  
498 inter-annual precipitation variability. However, both models depict the spatial distribution of  
499 the coefficient of variation reasonably well but with larger biases, roughly around 50%.  
500 Particularly, BARRA displays significantly higher biases, especially in Northern Australia.

501 Regarding the replication of trend patterns, both models exhibit substantial biases,  
502 reaching approximately 100%. This is especially notable in Northern Australia, where they  
503 both notably underestimate the increasing trend. Furthermore, while both BARRA and ERA5  
504 possess about the right spatial patterns of extremes, biases are evident. ERA5 tends to  
505 underestimate consecutive dry days (CDD) and certain heavy rainfall events, while BARRA  
506 tends to overestimate these extremes. Discrepancies in various precipitation indices across  
507 regions are apparent, with BARRA generally displaying larger biases compared to ERA5.

508 When examining temporal correlations for extreme precipitation indices compared to  
509 mean annual precipitation, both BARRA and ERA5 show weaker correlations, except for a  
510 few indices ([CDD](#), [R10mm](#), [R90p](#)) where ERA5 slightly outperforms BARRA. While both  
511 models align in coefficient of variation patterns and biases for certain extremes, notable  
512 differences arise in others. BARRA notably underestimates very heavy precipitation variability  
513 over inland regions, whereas ERA5 presents smaller biases or even positive biases in these  
514 areas. Moreover, BARRA tends to overestimate extreme precipitation variability in Northern  
515 Australia compared to ERA5. Specifically, BARRA showcases more than double the biases in  
516 variability compared to ERA5 for specific extreme precipitation indices.

517 In terms of trend patterns, both models generally replicate the observed trends but  
518 exhibit considerable biases. BARRA shows both underestimations and overestimations in  
519 certain regions and indices, while ERA5 displays different biases, including overall smaller  
520 biases compared to BARRA across these precipitation extremes.

521 In summary, our findings suggest that both ERA5 and BARRA are reliable for  
522 climatological analyses, including mean precipitation and precipitation extremes, and can be  
523 confidently used by end-users for such purposes. However, as discussed in the introduction,  
524 caution is advised when using these datasets for variability and trend analyses, particularly for  
525 specific extreme events like Rx1day. The performance of these reanalyses is regionally  
526 dependent, and this should be considered when using them as observational references for  
527 evaluating other model simulations. Additionally, the biases in the variability and trends of  
528 climate extremes present in both datasets must be carefully accounted for when comparing  
529 them with other data sources.

530

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#### 532 **Data Availability**

533 Details about AGCD are available at the Australian Bureau of Meteorology website  
534 (<http://www.bom.gov.au/metadata/catalogue/19115/ANZCW0503900567>, (accessed on)).  
535 The dataset is available on the NCI (National Computational Infrastructure) server in project  
536 zv2. Detail on how to access the data can be found at [http://climate-](http://climate-cms.wikis.unsw.edu.au/AGCD)  
537 [cms.wikis.unsw.edu.au/AGCD](http://climate-cms.wikis.unsw.edu.au/AGCD), (accessed on). ERA5 data is available on the NCI in Project  
538 rt52. BARRA data is available on the NCI in project cj37.

#### 539 **Author Contributions**

540 KKWC and FJ conceptualized and implemented the research. KKWC, FJ and NN performed  
541 the data analysis and prepared the figures. **KKWC and FJ** prepared the draft manuscript. All  
542 authors contributed to the discussion of results, editing and finalization of the manuscript.

#### 543 **Competing Interests**

544 The authors declare that they have no conflict of interest.

545

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552 comments and detailed suggestions for us to improve the manuscript.

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715 Table 1 List of ET-SCI indices evaluated in this study.

Index	Definition	Units	Timescale	Sectors
<b>PRCPTOT</b>	Total wet-day precipitation (Sum of daily precipitation $\geq 1.0$ mm)	mm	Annual/Monthly	Agriculture and food security, water, water resources and food security, forestry/GHGs
<b>CDD</b>	Consecutive dry days (Maximum number of consecutive dry days (when precipitation $< 1.0$ mm))	days	Annual	Health, agriculture and food security, water resources and food security, disaster risk reduction, forestry/GHGs
<b>CWD</b>	Consecutive wet days (Maximum annual number of consecutive wet days (when precipitation $\geq 1.0$ mm))	days	Annual	Coasts, agriculture, transport operations
<b>R10mm</b>	Days when precipitation is at least 10mm	days	Annual/Monthly	Coasts
<b>R90p</b>	Total annual precipitation from very heavy precipitation days (Annual sum of daily precipitation $> 90$ th percentile)	mm	Annual	Coasts, transport operations
<b>R99p</b>	Total annual precipitation from very heavy precipitation days (Annual sum of daily precipitation $> 99$ th percentile)	mm	Annual	Coasts, transport operations
<b>Rx1Day</b>	Amount of precipitation from very wet days (Maximum 1-day precipitation)	mm	Annual/Monthly	Agriculture and food security, water, coasts, disaster risk reduction, forestry/GHGs

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719 Table 2 Domain-averaged absolute biases and temporal correlation between BARRA/ERA5  
720 and AGCD for annual precipitation and precipitation extremes

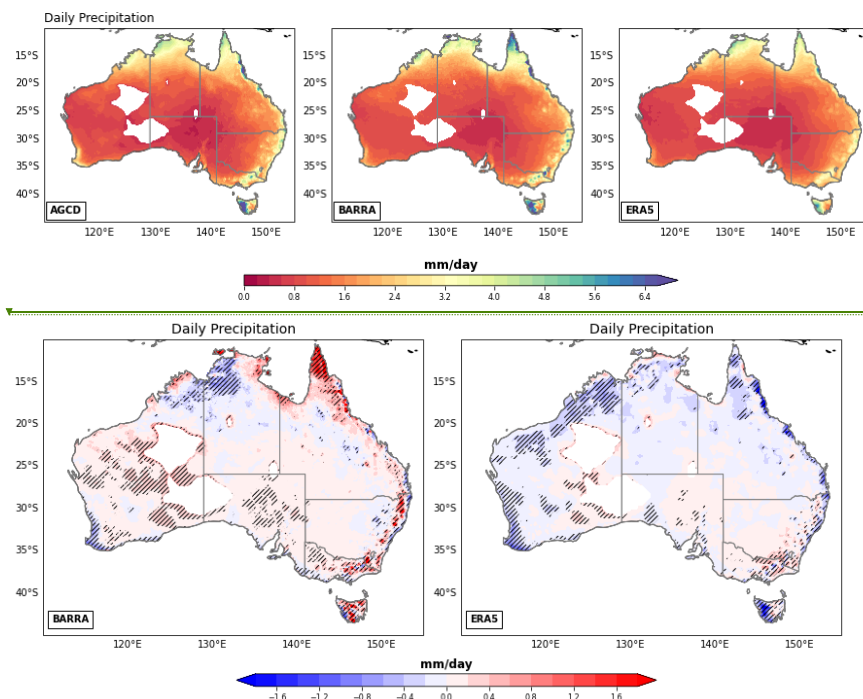
Indices	Absolute biases in annual mean		Temporal correlation		Absolute biases in CV		Absolute biases in trend	
	BARRA	ERA5	BARRA	ERA5	BARRA	ERA5	BARRA	ERA5
<b>Annual pr</b>	0.169	0.149	0.771	0.854	0.063	0.037	0.008	0.007
<b>CDD</b>	14.543	6.913	0.578	0.650	0.050	0.045	0.584	0.566
<b>CWD</b>	2.346	1.714	0.446	0.527	0.061	0.059	0.064	0.060
<b>R10mm</b>	3.265	1.700	0.688	0.761	0.081	0.053	0.085	0.094
<b>R90p</b>	0.777	0.439	0.761	0.827	0.211	0.082	0.023	0.023
<b>R99p</b>	4.093	3.668	0.562	0.625	0.121	0.060	0.206	0.162
<b>Rx1day</b>	20.333	7.916	0.380	0.486	0.219	0.107	0.848	0.542

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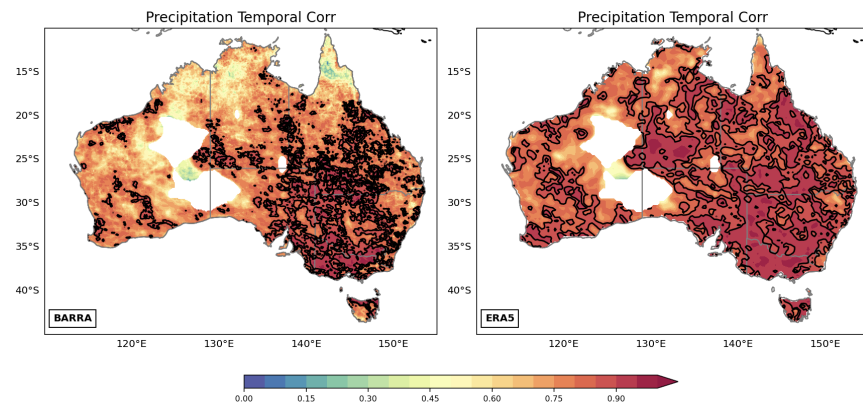
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727 Figure 1 Annual mean precipitation of AGCD, BARRA and ERA5 (upper panels) and annual  
 728 mean biases between BARRA/ERA5 and AGCD (lower panels). The regions with  
 729 low density of station observations in AGCD has been masked and not considered in  
 730 all subsequent evaluation. Unit: mm/day. Stippling indicates areas with biases that  
 731 are statistically significant at 95% confidence level.

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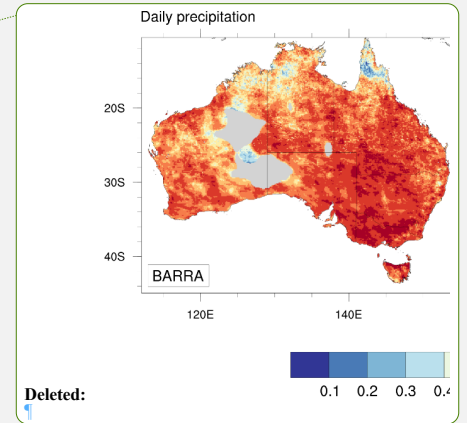
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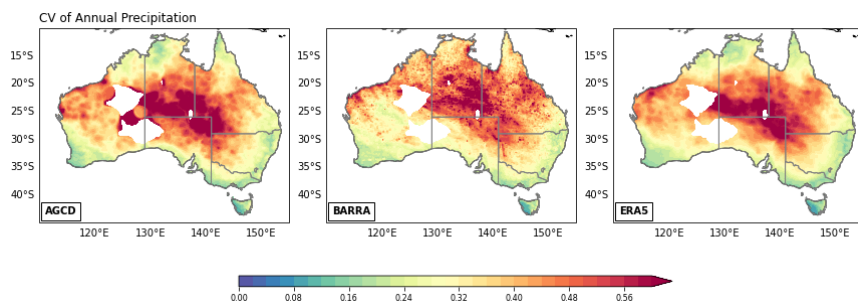
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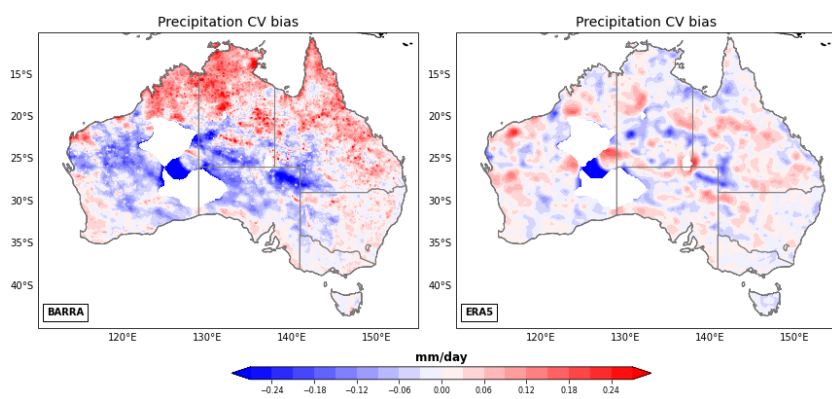
Figure 2 Temporal correlation coefficient of annual precipitation between BARRA/ERA5 and AGCD. A black contour at value 0.85 has been added for reference.



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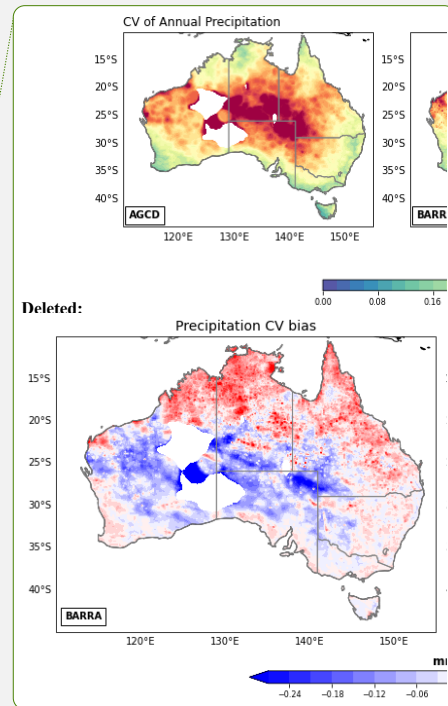


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747 Figure 3 CV of annual precipitation for AGCD, BARRA and ERA5 (upper panels) and biases  
748 in CV between BARRA/ERA5 and AGCD (lower panels).

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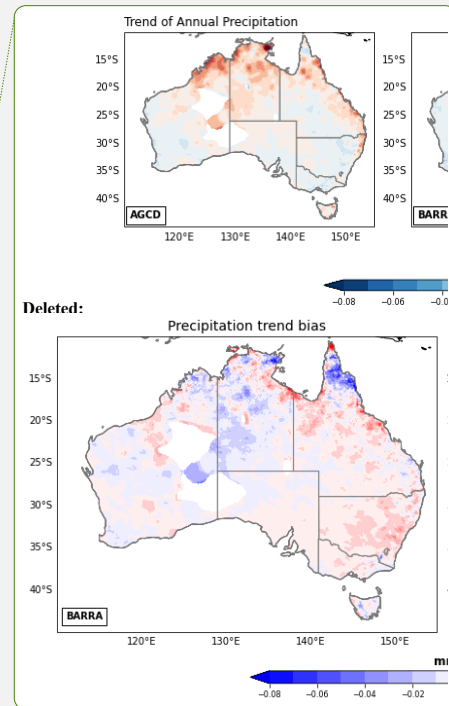
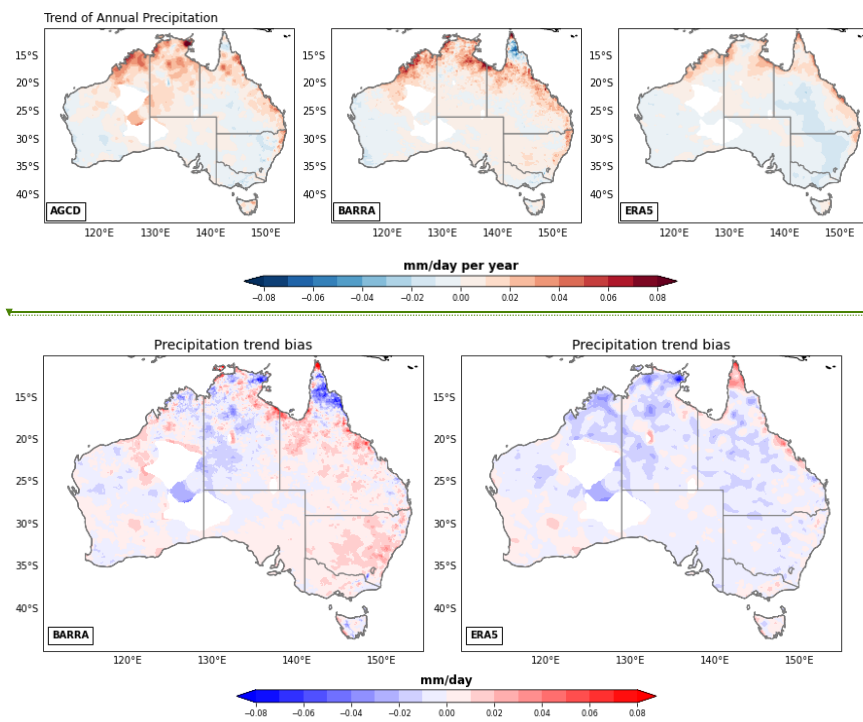
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756 Figure 4 Trend of annual precipitation for AGCD, BARRA and ERA5 (upper panels) and  
757 biases in trend between BARRA/ERA5 and AGCD (lower panels).

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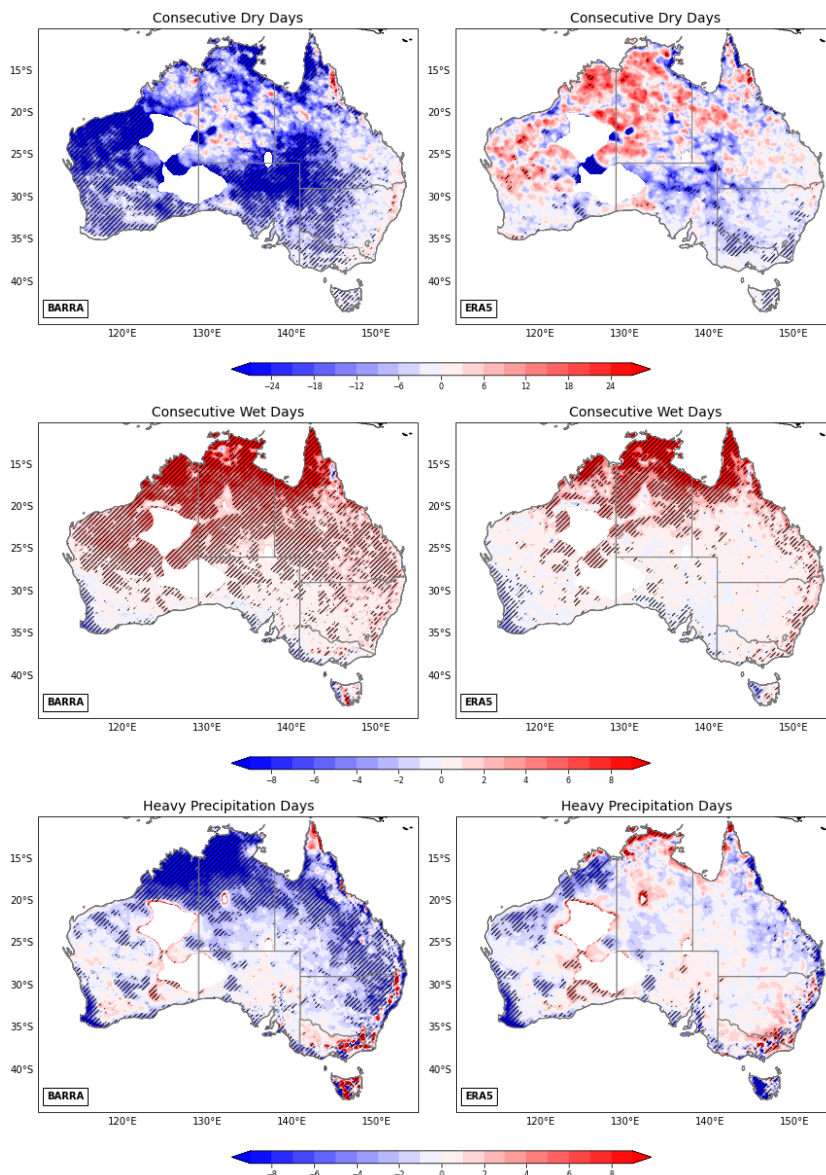
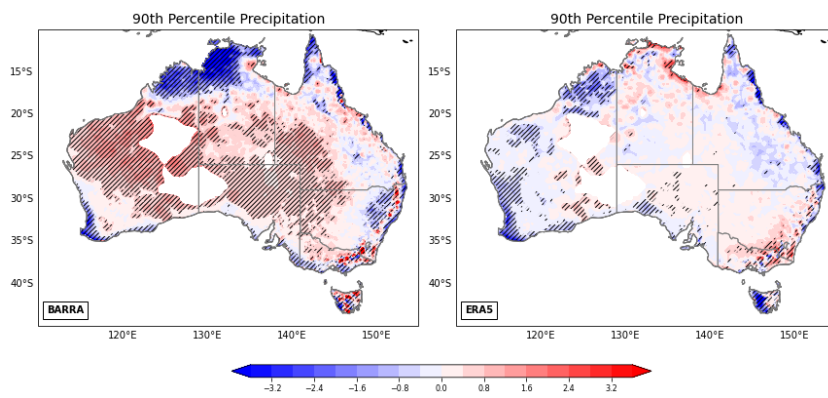


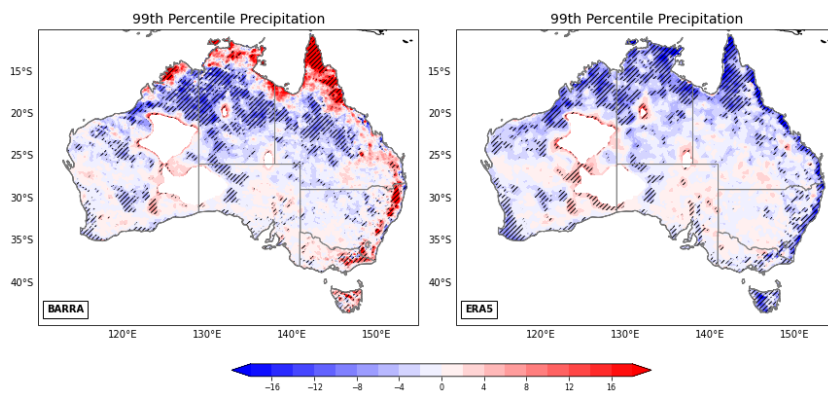
Figure 5 Biases in CDD, CWD, R10mm, R90p, R99p and Rx1Day in BARRA (left column) and ERA5 (right column). Stippling indicates areas with biases that are statistically significant at 95% confidence level. A black contour at 40% has been added to the panels for Rx1day (last row) for reference.

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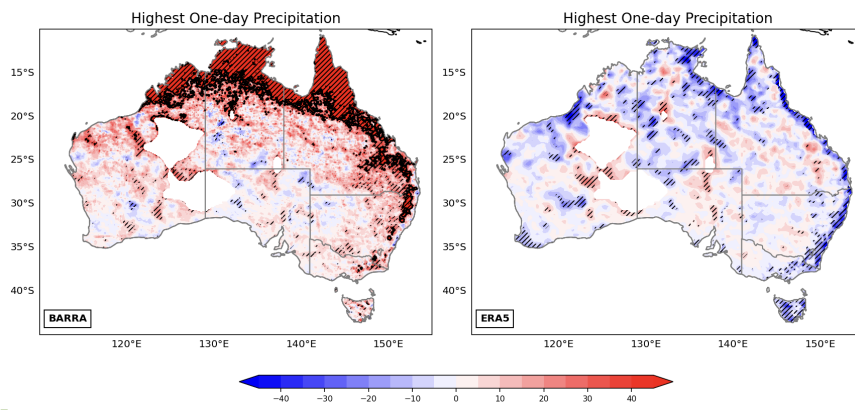
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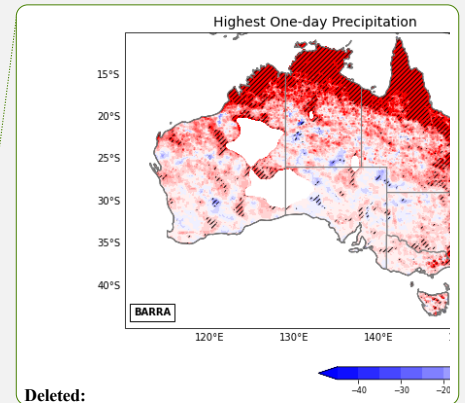
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Figure 5 (continued).





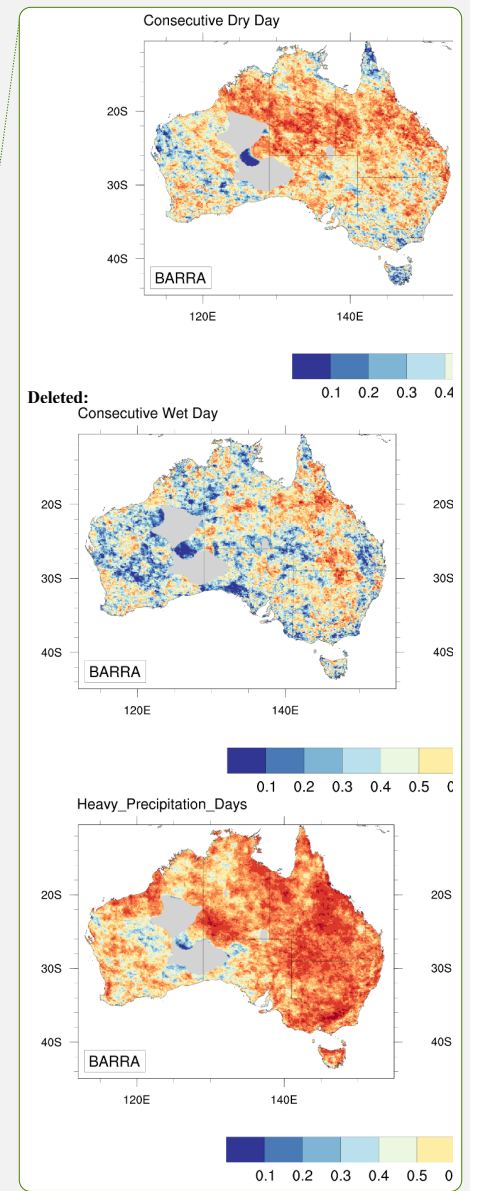
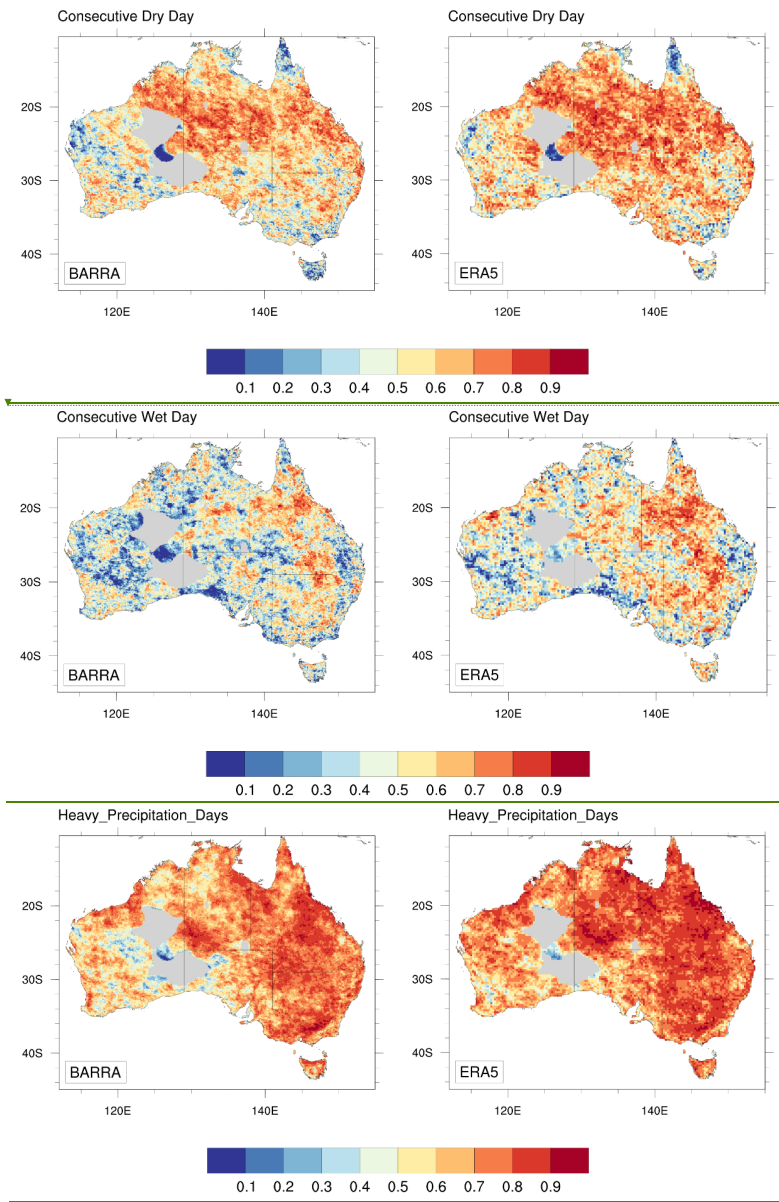


Figure 6 Temporal correlation of CDD, CWD, R10mm, R90p, R99p and Rx1Day between BARRA and AGCD (left column) and between ERA5 and AGCD (right column).

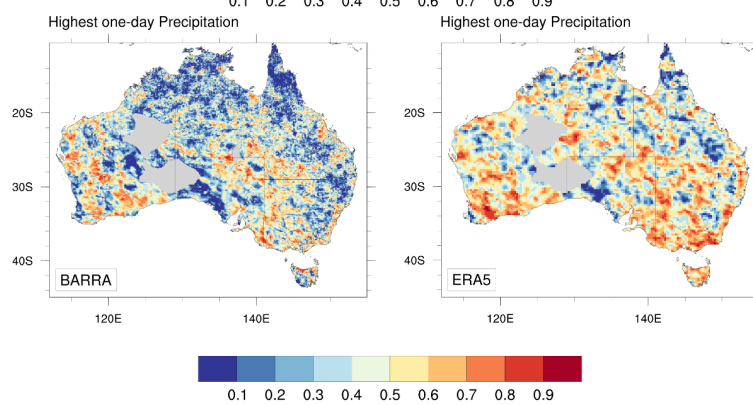
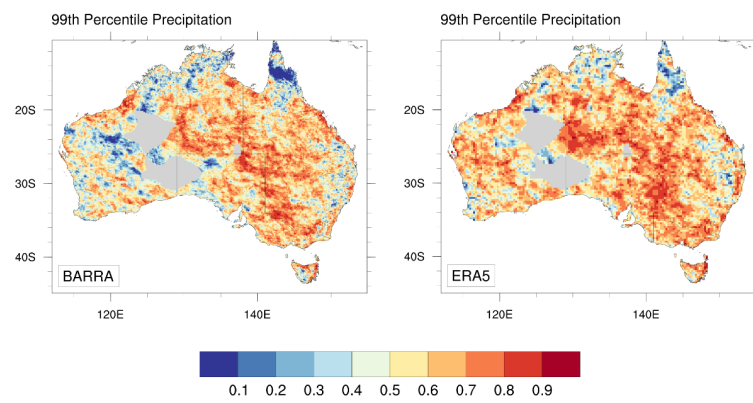
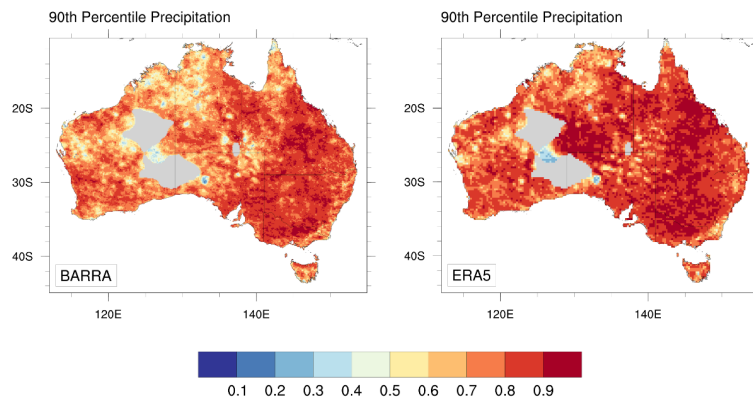
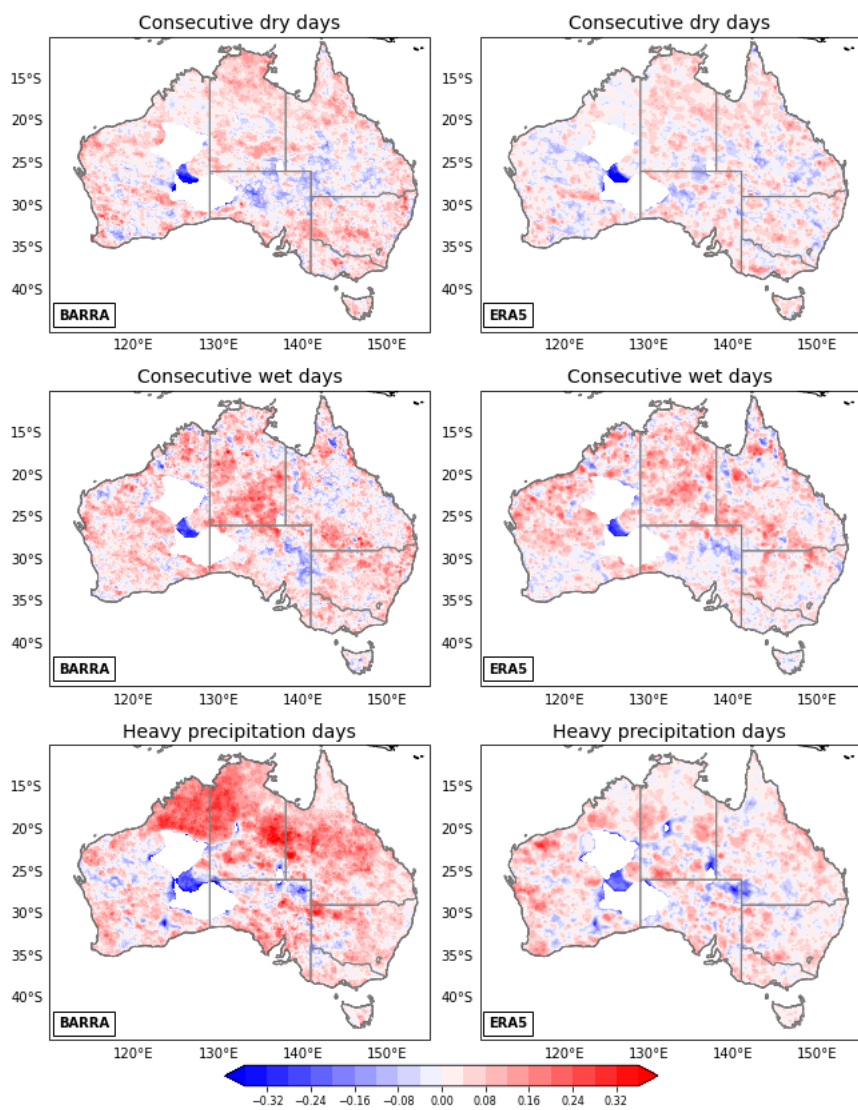


Figure 6 (continued).



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795 Figure 7 Biases in CV of CDD, CWD, R10mm, R90p, R99p and Rx1Day for BARRA (left  
 796 column) and ERA5 (right column) relative to AGCD.

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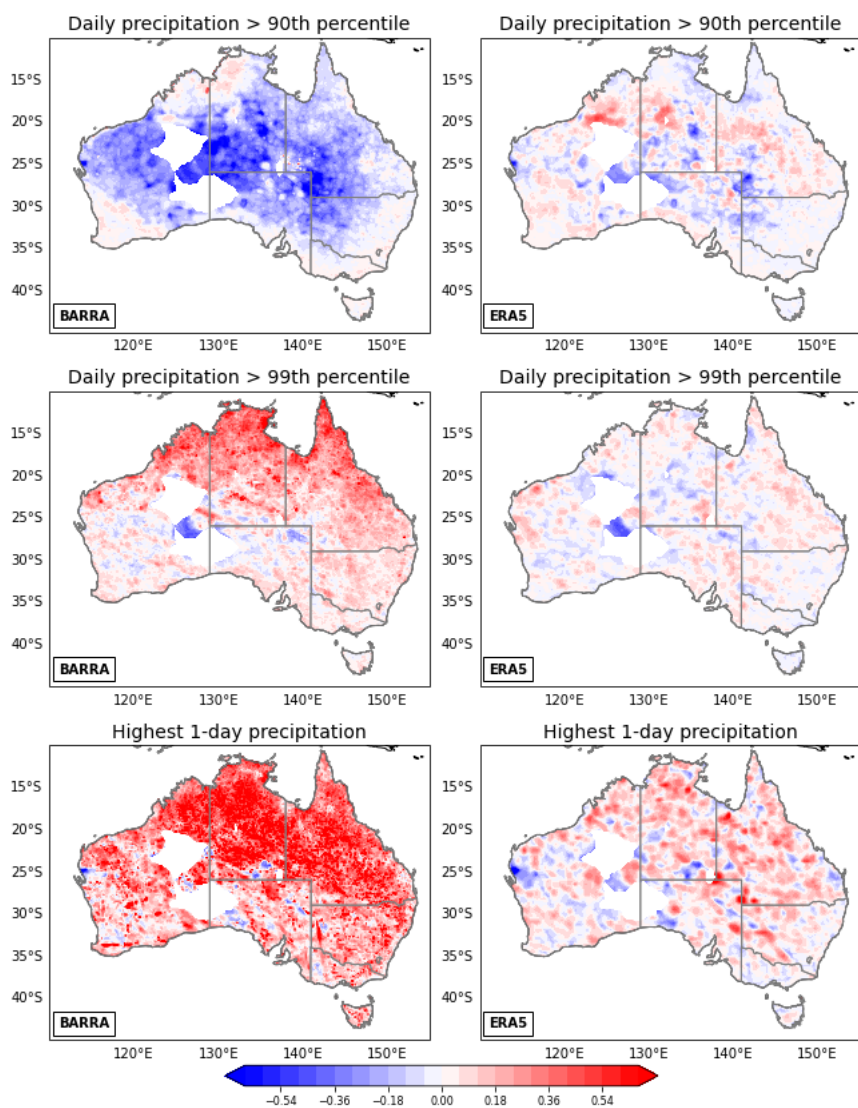
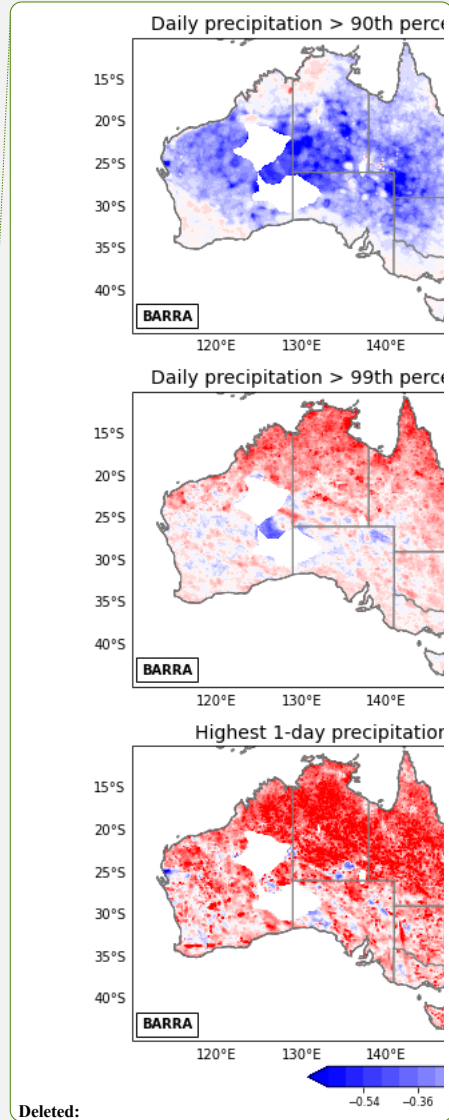


Figure 7 (continued).



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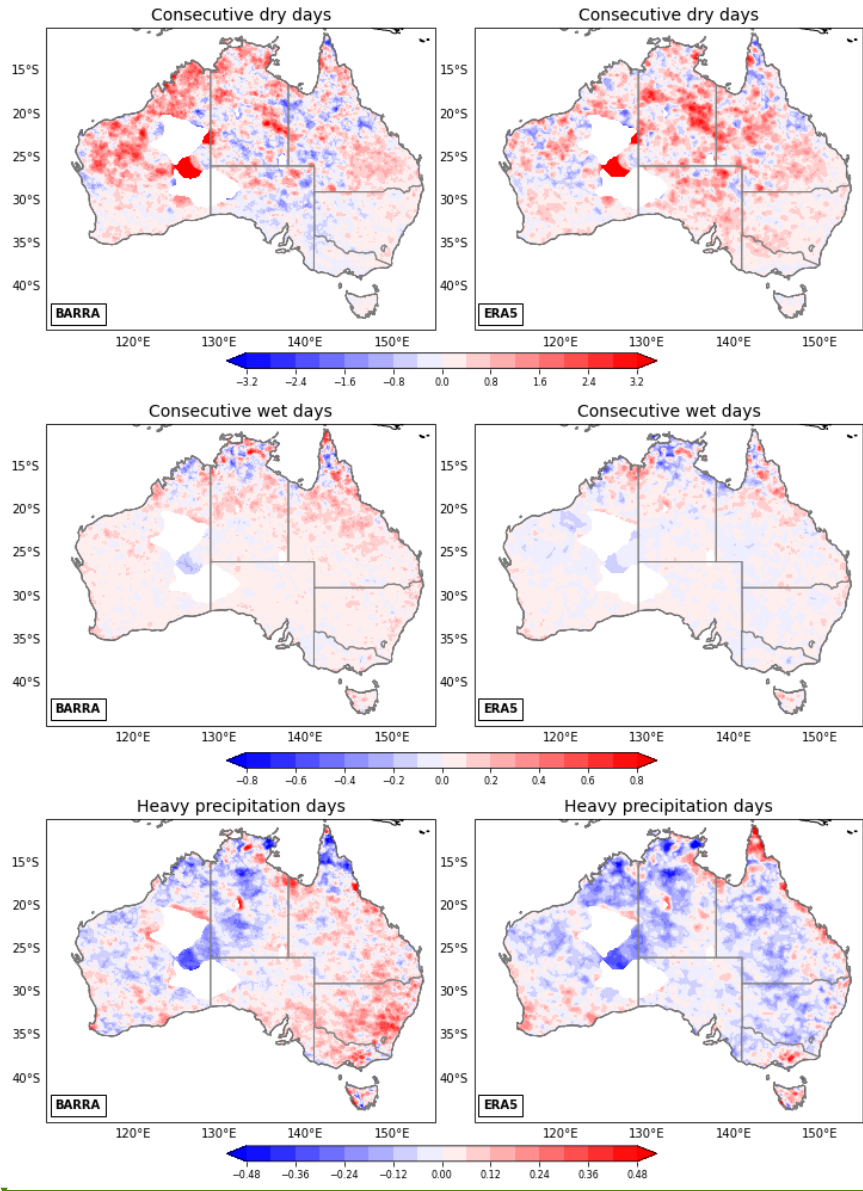
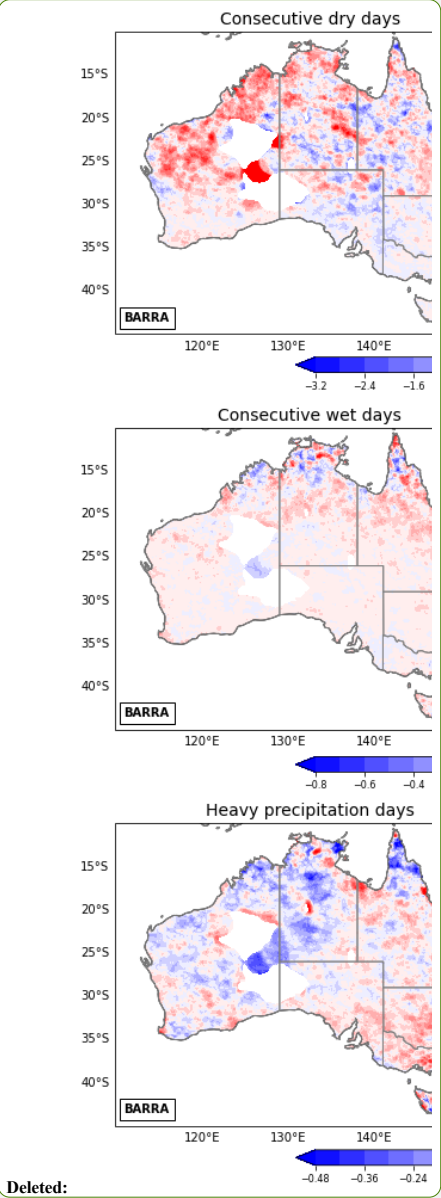
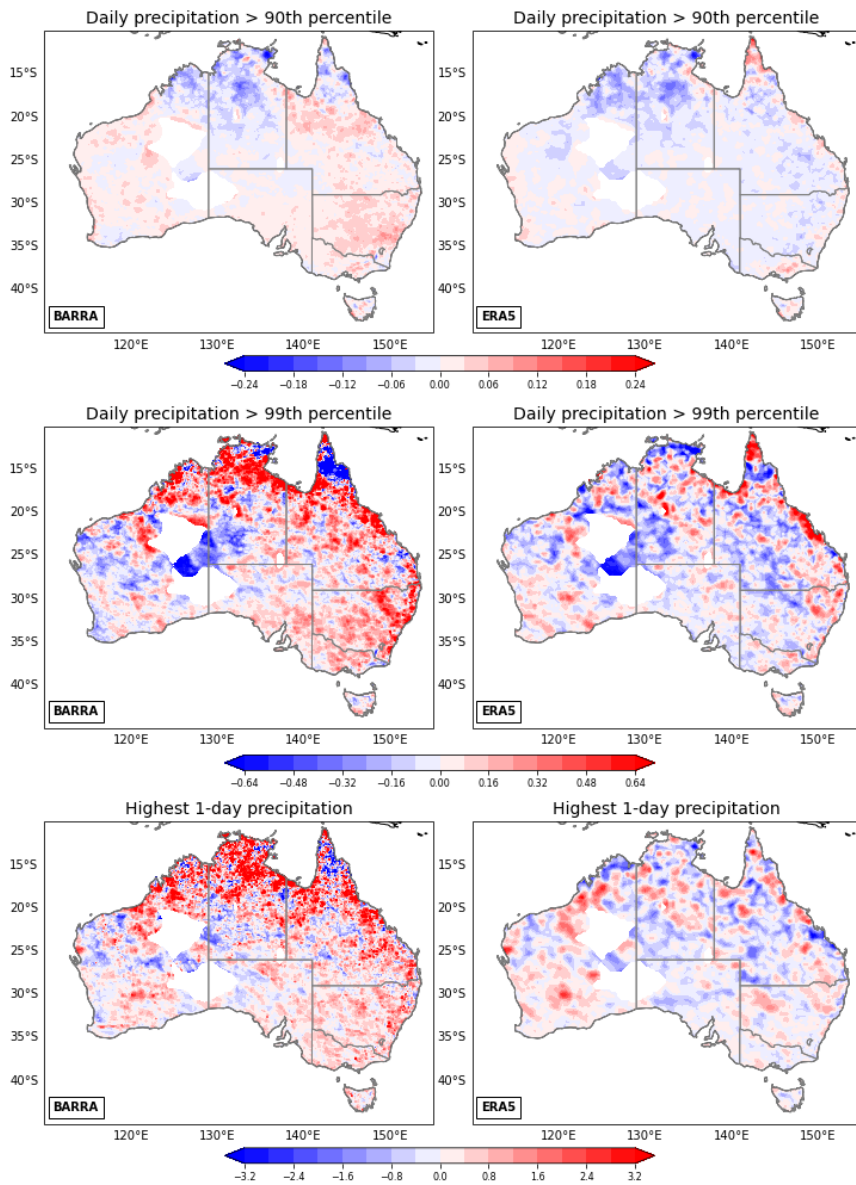


Figure 8 Biases in trends of CDD, CWD, R10mm, R90p, R99p and Rx1Day for BARRA (left column) and ERA5 (right column) relative to AGCD.



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Figure 8 (continued).

