

Enhancing the performance of 1D-2D flood models using satellite laser altimetry and multi-mission surface water extent maps from Earth Observation (EO) data

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Abstract. Digital elevation models (DEMs) are essential datasets, particularly for flood inundation mapping in one-dimensional (1D) to two-dimensional (2D) flood models. Given the significant uncertainties associated with DEMs that can affect flood modelling accuracy, minimizing these inaccuracies is essential. This study aims to improve the performance of 1D-2D flood models using satellite Earth Observation (EO) data, focusing on the lower Chao Phraya (CPY) basin.

Two workflows are proposed: DEM analysis and flood map analysis. The DEM analysis evaluates 10 DEM products, including three local DEMs provided by Thai agencies (LDD, JICA, and a merged LDD-JICA DEM) and seven global DEMs derived from EO data (ASTER GDEM V3, SRTM V3, MERIT, GLO30, FABDEM V1-2, TanDEM-X, and TanDEM-EDEM). The evaluation process uses ICESat-2 ATL08 data processing, vertical datum reference processing, and evaluation of DEMs using ICESat-2 ATL08 benchmark processing. The DEMs are assessed using satellite laser altimetry data from the Ice, Cloud, and Land Elevation Satellite-2 (ICESat-2) as the benchmark. The evaluation employs standardized metrics, including point-wise, grid-wise, and track-wise comparisons, to identify the most suitable DEM for integration into the flood model. Results indicate that the merged LDD-JICA DEM and FABDEM V1-2 DEM exhibit the highest accuracy among local and global products, respectively, with root mean square errors (RMSE) of 1.93 m and 1.95 m, and percent biases (PBIAS) of -15.38% and 4.59%. The flood map analysis workflow involves comparing flood extent maps derived from multi-mission satellite datasets, and simulated flood maps generated from 1D-2D flood models using the best available DEMs. This workflow utilizes surface water extent (SWE) maps from the WorldWater project, obtained from the Sentinel-1 and Sentinel-2 imaging satellites, and flood maps from the Geo-Informatics and Space Technology Development Agency (GISTDA) in Thailand to validate flood maps produced by the 1D-2D flood model based on the merged LDD-JICA DEM and FABDEM V1-2 DEM. The results reveal that flood maps based on the FABDEM V1-2 DEM slightly outperform those based on the merged LDD-JICA DEM, with an improvement of approximately 13.55–25.56% in the critical success index (CSI). This study highlights the potential of leveraging satellite EO data to enhance the accuracy and reliability of 1D-2D flood models, thereby improving flood inundation predictions for effective flood management.

1 Introduction

Nowadays, flooding is one of the most common hazards globally, impacting health, economies, and livelihoods worldwide.

35 Flood models play a crucial role in forecasting floods and assessing flood risks, thereby assisting decision-makers in effective water management, particularly through one-dimensional (1D) - two-dimensional (2D) flood models. These models simulate various aspects of flooding, including flow, water levels, flood inundation extents, flood depths, flood maps, and flood duration (DHI Water and Environment, 2019). The Digital Elevation Model (DEM) serves as a primary input parameter for 1D-2D flood models, enabling accurate simulation of flood overflow from rivers, floodplains, and inundated areas, particularly in flat and low-lying regions. The DEM significantly influences the simulation of flood inundation in both 1D-2D and 2D flood models ((Saksena & Merwade, 2015); (J. Shen & Tan, 2020); (Wu et al., 2007); (Morrison et al., 2022)), urban areas (McClean et al., 2020), coastal areas (Darnell et al., 2008), and flood warning systems (Lamichhane & Sharma, 2018). Ultimately, the reliability of flood inundation predictions relies on the accuracy and resolution provided by the DEM, directly impacting the representation of flow geometry characteristics within flood models.

45 Currently, the advancements in survey technologies, such as Unmanned Aerial Vehicles (UAVs) (Perera & Nalani, 2022), Light Detection and Ranging (LiDAR) (Raj et al., 2020), and Mobile Mapping Systems (MMS) (Schwarz & El-Sheimy, 2007), have significantly enhanced the accuracy, quality, and resolution of DEMs. These technologies enable the production of high-resolution terrain data; however, they remain costly, time-consuming, and less feasible for monitoring dynamic land-use changes or covering large river basins. For example, following the severe flooding in 2011, Thailand's Royal Irrigation Department collaborated with the Japan International Cooperation Agency (JICA) to survey a 27,000-square-kilometer area. This effort produced a high-resolution 2x2 meter DEM and required approximately seven months to complete (Japan International Cooperation Agency (JICA), 2012), underscoring the significant resources needed for such large-scale surveys. EO technologies offer a promising alternative by providing global DEMs with comparable resolution and quality. EO-based DEMs, such as ASTER GDEM3 (Abrams et al., 2020), SRTMv3 (Farr et al., 2007), MERIT (Yamazaki et al., 2017), GLO30

55 (AIRBUS, 2020), FABDEMv1-2 (Hawker et al., 2023), TanDEM-X (Krieger et al., 2007), and TanDEM-EDem, are freely available for download and utilize advanced techniques of EO and machine learning, to generate elevation estimates. These satellite-derived DEMs cover remote or inaccessible areas, offering a cost-effective and efficient solution for generating high-resolution terrain data. Moreover, global DEMs derived from EO are increasingly being utilized as inputs for 1D-2D flood models, providing a practical and scalable option for flood risk assessment and forecasting in regions with limited resources.

60 However, validating the DEM before integrating it into the 1D-2D flood model is essential. The Ice, Cloud, and Land Elevation Satellite-2 (ICESat-2) is a satellite equipped with a laser altimeter, capable of measuring ice sheet and glacier elevation change, sea ice freeboard, land elevation, and water elevation (Neumann et al., 2019), providing opportunities for validating DEMs even in remote and hard-to-reach areas worldwide, such as Finland (X. Wang & Liang, 2023), Spain (Zhu et al., 2022), East Antarctica (Hao et al., 2022), Alaska in the USA (C. Wang et al., 2019), and the Qinghai-Tibet Plateau in China (Weifeng et al., 2024). Additionally, ICESat-2 has been used to assess the suitability of global DEMs for hydrodynamic modeling in data-

scarce regions (Nandam & Patel, 2024) and to enhance the accuracy of 2D hydraulic models in the upstream Yellow River (Frias et al., 2024). Moreover, while an efficient DEM enhances the efficiency of 1D-2D flood simulation, it is important to systematically validate flood maps. Currently, satellite earth observation (EO) data can be utilized for monitoring and providing surface water extent (SWE) with synthetic-aperture radar (SAR) sensors, such as RADARSAT (Raney et al., 1991), ENVISAT
70 ASAR (Lv et al., 2005), COSMO-SkyMed (Pulvirenti et al., 2014), and TerraSAR-X (Martinis et al., 2013), which is the only way to validate flood inundation maps from flood models over regional scales. The WorldWater project developed a robust and scalable EO solution for inland SWE monitoring, which can be utilized by a large community of stakeholders involved in local water management (Tottrup et al., 2022). The project used free and open optical and SAR satellite imagery from the Sentinel-1 and Sentinel-2 missions to generate monthly SWE maps over four years, which are accessible from
75 <https://worldwater.earth/>. The product offers new opportunities for validating modelled flood maps with higher SWE resolution.

While satellite EO provides SWE maps that delineate water bodies and inundated areas, they cannot be directly compared to flood maps from 1D-2D flood models. The output of 1D-2D flood models are riverine flood maps. Additional flood classification processing is necessary to ensure comparability between SWE maps and the output of a flood model. However,
80 flood type classification using SWE maps poses challenges and difficulties. Many studies focus on classifying flood types based on meteorological condition rather than using SWE maps, such as (Nied et al., 2014), (Turkington et al., 2016), decision tree using meteorological data ((Stein et al., 2019), and (Yan et al., 2023)). Riverine flood classification specifically involves identifying floods caused by river overflow from SWE maps. Here, we used expanding segmentation labels (ESL) (Van Der Walt et al., 2014), connected component labeling (CCL) ((Rosenfeld & Pfaltz, 1966) and (AbuBaker et al., 2007)), masking
85 off riverine and permanent water, and morphological image processing (MIP) (Soille, 2003) techniques applied to the SWE maps to separate riverine flood areas from other inundated areas.

This study presents two new workflows supporting flood modelling and forecasting in the lower Chao Phraya (CPY) River basin in Thailand and elsewhere: 1) Comprehensive DEM Evaluation: A detailed assessment of 10 DEM products, including three local and seven global DEMs, was conducted using ICESat-2 as a benchmark for the Thailand domain. DEM performance
90 in the lower CPY basin was evaluated using statistical methods, including bias (mean error, ME), mean absolute error (MAE), mean square error (MSE), and root mean square error (RMSE), with comparisons made at point and grid level, as well as track-wise comparisons. The highest-performing DEM from this evaluation was subsequently integrated into a 1D-2D flood model to simulate flood inundation. 2) Systematic Comparison of Flood Maps: Simulated 2D inundation patterns were compared with flood maps derived from satellite EO-based surface water extent (SWE) using a riverine flood classification process. The model's performance was assessed using three statistical metrics: probability of detection (POD), false alarm ratio
95 (FAR), and critical success index (CSI). These methods will improve the performance of the operational hydrologic-hydraulic forecasting system for the Chao Phraya River, managed by the Hydro-Informatics Institute (HII) in Thailand.

2 Study Area

The study area is located in the central part of Thailand, as shown in Figure 1(a). The delta area of the lower CPY River basin in Thailand forms the study area depicted in Figure 1(c). The size of the study area is approximately 16,643 km², including about 70% irrigation area and 20% urban area. The topography of the study area is characterized by a flat terrain, predominantly consisting of a low-lying alluvial floodplain. The northern part of the study area is a mountainous region with four main rivers: the Ping, Wang, Yom, and Nan rivers. These rivers converge to form the CPY river, which then flows into the study area. The eastern and western parts of the study area are connected to the Bang Pakong River and the Mae Klong basin, respectively.

The southern part of the study area borders the Gulf of Thailand.

The study area is located in a tropical climate and is influenced by northeast and southwest monsoons. The northeast monsoon brings cool and dry air from November to February, while the southwest monsoon brings humid air from May to October. The precipitation is approximately 1,100 mm during the rainy season and 170 mm during the dry season. The flooding in the study area is caused by the main rivers and their tributaries. The tributaries of the CPY river include Tha-Chin, Noi, and Lopburi.

Flooding problems are more severe along the main course of the CPY river compared to others. Nevertheless, flooding mechanisms are complicated, arising from the combined effects of extreme precipitation, river overflows, insufficient river conveyance, land-use change, and sea-level rise. This results in frequent flooding, as shown Figure 3(c).

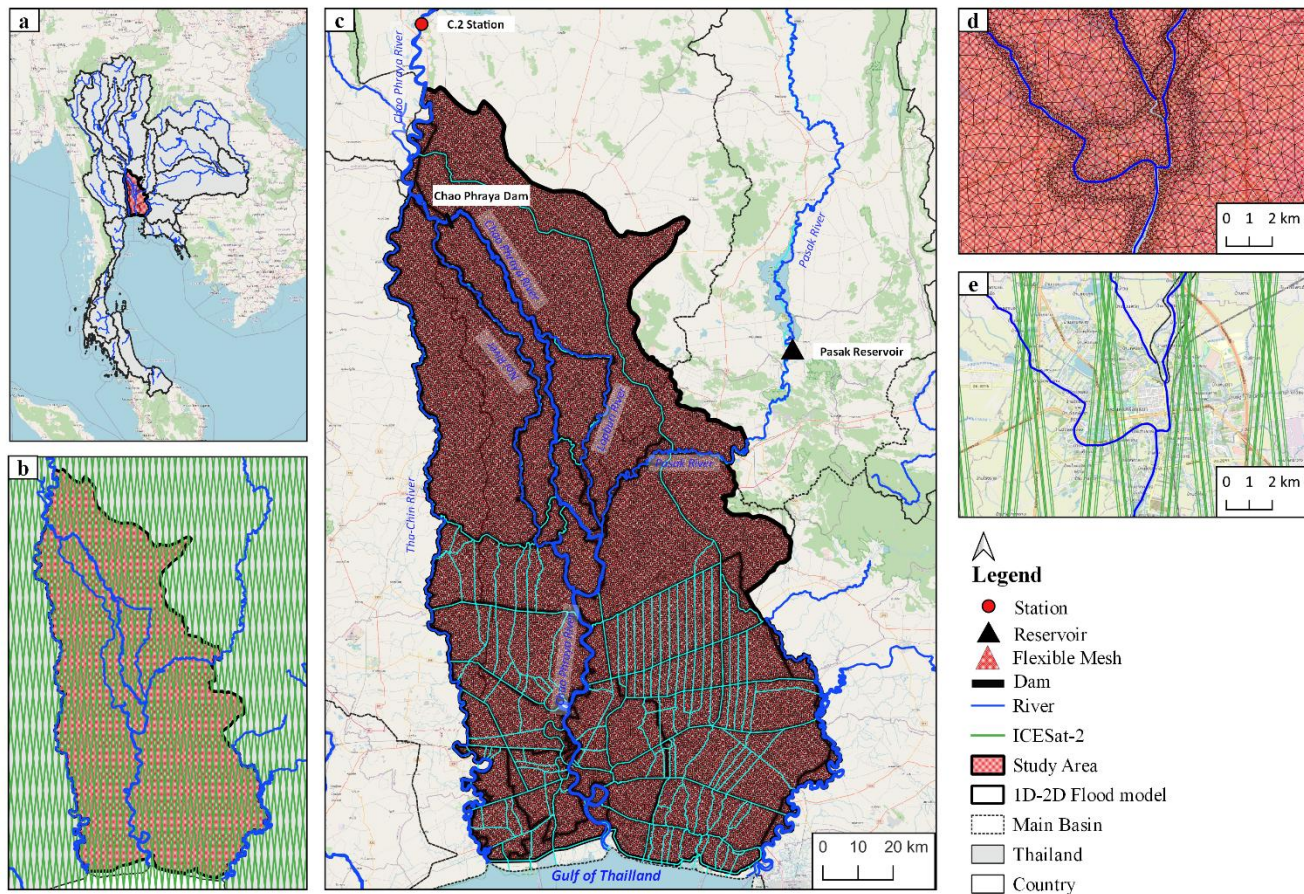


Figure 1: (a) the location of the study area, (b) the ICESat-2 orbit, (c) the study area/1D-2D flood model, (d) the flexible mesh in the flood model, and (e) ICESat-2 beam pairs. © OpenStreetMap contributors 2015. Distributed under the Open Data Commons Open Database License (ODbL) v1.0.

3 Materials

3.1 1D-2D Flood modelling

In this study, we used the flood model from the decision support system for flood forecasting and water management in the CPY River basin, developed in collaboration with HII and DHI A/S since 2012 (Sisomphon et al., 2013) and updated with new information in 2016 (Charoensuk et al., 2018). The decision support system for flood forecasting and water management in the CPY basin continues to operate, supporting the Thai Government in managing flood risk and providing real-time flood forecasts.

The flood model uses the MIKE FLOOD software developed by DHI A/S. A MIKE FLOOD model (DHI Water and Environment, 2019) consists of coupled one-dimensional (1D) and two-dimensional (2D) models, namely MIKE11 and

MIKE21, respectively. The 1D hydraulic model (MIKE11) simulates unsteady flow in river networks solving the Saint-Venant equations with an implicit finite difference solver (DHI Water and Environment, 2021). The main branches of MIKE11 include the Chao Phraya, Tha-Chin, Lopburi, Noi and Pasak rivers. Cross-sections, rainfall-runoff, boundary conditions, hydrodynamic parameters, and control structures were implemented in MIKE11. The MIKE21 model is an overland flow model utilizing 2D shallow water equations (Danish Hydraulic Insitute, 2016). MIKE21 employs a 2D flexible mesh based on the digital elevation model (DEM) to assess flood depth and its propagation. The river network in MIKE11 is dynamically linked to floodplain bathymetry through lateral links. The lateral links connect the river to the floodplain along its length using the cell-to-cell method, allowing water to overflow to the floodplain in the MIKE21 overland flood model. The lateral link connection uses the weir equation to calculate overflow in MIKE FLOOD (DHI Water and Environment, 2019).

The 1D-2D flood model, documented in (Hydro-Informatics Institute, 2017), establishes the following boundary conditions: upstream boundary forcing with discharge from C.2 station and releases from the Pasak Reservoir from the Royal Irrigation Department (RID) in the CPY and Pasak rivers, respectively. Meanwhile, the downstream boundary connects to the Gulf of Thailand using sea level measurements from the Hydrographics Department, Royal Thai Navy (NAVY), as illustrated in Figure 1(c). The MIKE11 model was calibrated using water level observations presented in (Charoensuk et al., 2024). MIKE21 utilized a flexible mesh to simulate overland flow, as illustrated in Figure 1(d), and MIKE FLOOD was calibrated against flood maps and satellite data from 2011, as detailed by (Charoensuk et al., 2018).

3.2 Geoid Models

To measure elevations around the Earth, a vertical reference is needed, with mean sea level chosen as the reference. The geoid is the level (equipotential) surface of the Earth's gravity field that best coincides with mean sea level. This surface connects the oceans and extends through the continents. The geoid serves as the reference surface for levelled heights, commonly expressed as 'heights above sea level'. In order to compare heights from different data sources, all data has to be re-referenced to the same geoid model. A geoid model is a spatial representation of geoid height, encompassing both global and local scales. This study has collected three geoid models, summarized in Table 1. Thailand has its own local geoid model. The latest one, TGM2017, was released in 2018. This geoid is based on new gravity measurements taken around Thailand and has been shown to better match the expected geoid heights than the EGM2008 model (Dumrongchai et al., 2021). TGM2017 provides the best fit for Thailand, it was chosen as the primary geoid model, and all heights were re-referenced to TGM2017.

Table 1: Geoid model datasets

Geoid model	Scale	Download	References
EGM96: The Earth Gravitational Model 1996	Global	https://earth-info.nga.mil/	(Lemoine et al., 1998)
EGM2008: The Earth Gravitational Model 2008	Global	https://earth-info.nga.mil/	(Pavlis et al., 2012)
TGM2017: Thailand geoid model 2017	Local	On request	(Dumrongchai et al., 2021)

3.3 Digital Elevation Models (DEM)

155 A digital elevation model (DEM) is a quantitative representation of the Earth’s surface elevation. The term "DEM" encompasses both digital terrain models (DTM) and digital surface models (DSM). A DSM maps the heights of all features on the surface, such as vegetation and buildings, while a DTM only represents the actual height of the terrain (“bare earth”). Multiple digital elevation models are available, local DEMs are often preferred due to their higher spatial resolution and vertical accuracy (McClean et al., 2020). In this study, we have collected 10 DEM products, as shown in Figure 2. A summary of these products is presented in Table 2Error! Reference source not found. and detailed statistical analyses are provided in Table A1. The three local DEM products were obtained from the Thai agency, namely LDD DEM, JICA DEM, and merged LDD-JICA DEM. Additionally, seven global DEMs were collected, including ASTEM GDEM V3, SRTMv3 DEM, MERIT DEM, FABDEMv1-2 DEM, GLO30 DEM, TanDEM-X, and TanDEM-EDEM.

165 Table 2: Digital Elevation Model (DEM)

Dem Product	Spatial resolution	Data Collection (Year)	Datum Reference	Type	Scale	Acquisition technique
LDD DEM	5 m	2004	EGM96 geoid	DTM	Local	Aerial stereo photo
JICA DEM	2 m	2012	EGM2008 geoid	DTM	Local	Airborne LiDAR
merged LDD-JICA DEM	2 m	-	TGM2017 geoid	DTM	Local	Fusion of multisource data
ASTEM GDEM V3	1 arcsecond (~30 m)	2000-2010	EGM96 geoid	DSM	Global	Satellite stereo images
SRTM DEM V3	1 arcsecond (~30 m)	2000	EGM96 geoid	DSM	Global	SAR Interferometry
MERIT DEM	3 arcseconds (~90 m)	2000	EGM96 geoid	DSM	Global	Fusion of multisource data
GLO30 DEM	1 arcsecond (~30 m)	2011-2015	EGM2008 geoid	DSM	Global	Fusion of multisource data
FABDEM v1-2	1 arcsecond (~30 m)	2011-2015	EGM2008 geoid	Base on DSM remove building and forest	Global	Fusion of multisource data
TanDEM-X DEM	0.4 arcsecond (~12 m)	2011-2015	WGS84 ellipsoidal height	DSM	Global	SAR Interferometry
TanDEM-X EDEM	1 arcsecond (~30 m)	2011-2015	WGS84 ellipsoidal height	DSM	Global	Fusion of multisource data

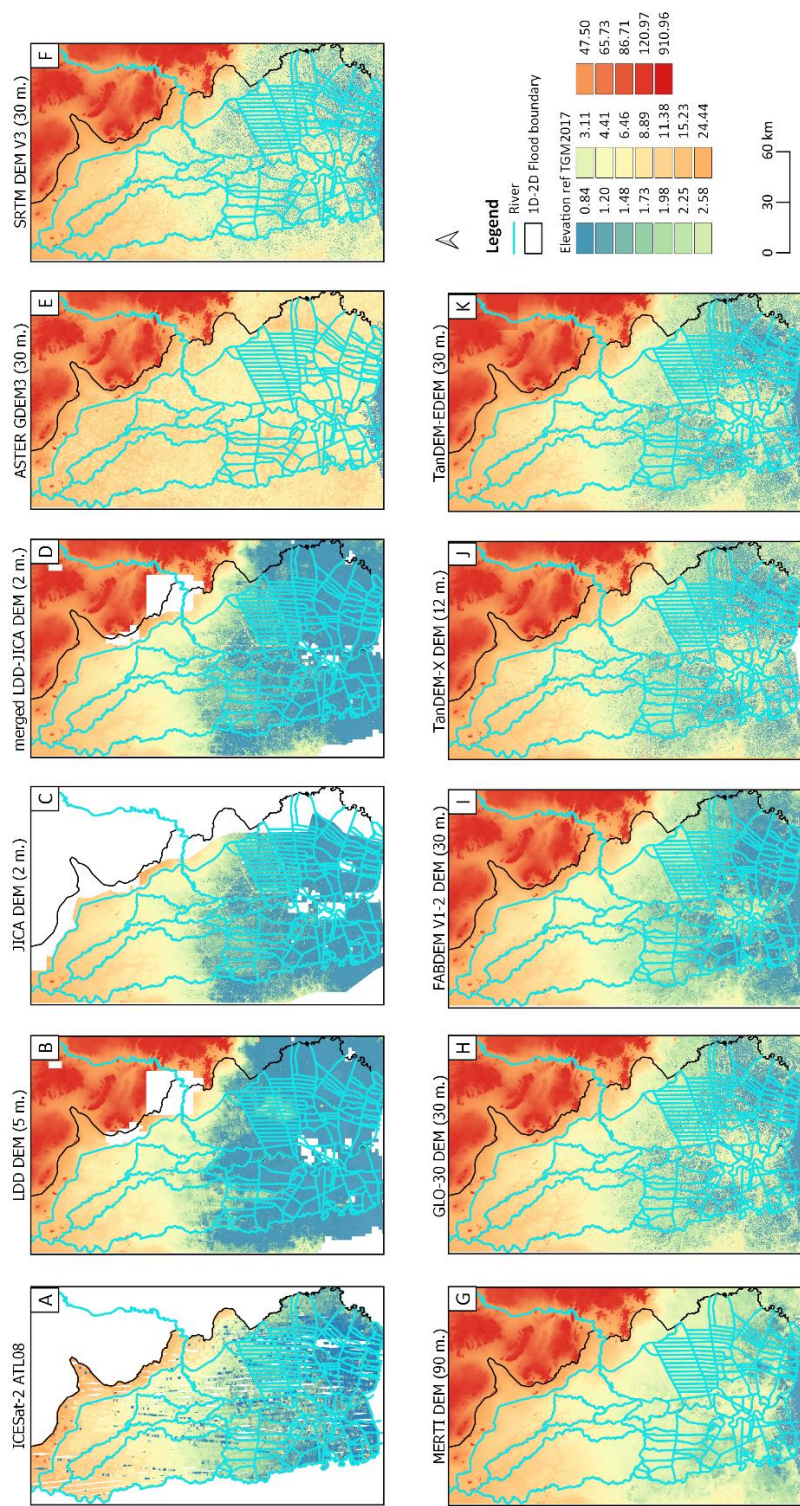


Figure 2: ICESat-2 ATL08 and DEM products, including: A) ICESat-2 ATL08 Surface elevation, B) Land Development Department (LDD) DEM, C) JICA DEM, D) Merged LDD-JICA DEM, E) ASTER GDEM Version 3, F) SRTM DEM Version 3, G) MERIT DEM, H) GLO-30 DEM, I) FABDEM v1.2 DEM, J) TanDEM-X DEM, and K) TanDEM-EDEM.

3.3.1 LDD DEM

The LDD DEM data is supplied by the Land Development Department of Thailand (LDD) in a grid format, with a resolution of 5x5 meters. This DEM was generated using photogrammetry using aerial stereo photo pairs with known scales (Paengwangthong & Sarapirome, 2012). This approach involves deducing distances between points from photos and determining object heights by identifying stereoscopic parallax from multiple pictures and rectifying with ground control points (GCPs) (Sholarin & Awange, 2015). Subsequently, orthorectification and interpolation are used to generate a DEM and mask off buildings and vegetation. Because buildings and vegetation are removed, the LDD DEM approximates a DTM (Sholarin & Awange, 2015).

3.3.2 JICA DEM

The JICA DEM was produced through a collaborative effort between the Royal Irrigation Department (RID) and the Japan International Cooperation Agency (JICA) at a resolution of 2x2 meters (Japan International Cooperation Agency (JICA), 2012). The JICA DEM was generated using Airborne Laser Scanning techniques with the LiDAR (Light Detection And Ranging) aerial technology. The LiDAR aerial survey employs a pulse laser to measure distances between the target and sensor, and it is applied on a large scale. The distance from the vehicle to the surface can be determined based on the travel time of the laser pulse (Argall & Sica, 2003). The JICA DEM was processed into a DTM filtering out features such as transportation facilities, buildings, and vegetation from the original data, as described in (Japan International Cooperation Agency (JICA), 2012).

3.3.3 Merged LDD-JICA DEM

The merged LDD-JICA DEM was generated by integrating the LDD DEM and JICA DEM as described by (Charoensuk et al., 2018). The JICA DEM served as the primary dataset, while the LDD DEM was utilized in areas with gaps within the 1D-2D Flood modeling boundary. To incorporate the LDD DEM into the merged LDD-JICA DEM within data gaps, we applied bias correction. The native LDD DEM and JICA DEM datasets were not referenced to the same vertical datum. The processing of the merged LDD-JICA DEM consists of two primary steps (Figure A2): 1) re-referencing both LDD DEM and JICA DEM to the TGM2017 reference, and 2) calculation of the correlation coefficient between the JICA and LDD DEM for 1000 random points, using linear regression to correct the bias in the LDD DEM, as shown in Figure A3. Following this, the JICA DEM and LDD DEM are combined to create the merged LDD-JICA DEM using linear regression. The resulting combined merged LDD-JICA DEM has a resolution of 2x2 meters.

3.3.4 ASTER GDEM3

The Advanced Spaceborne Thermal Emission and Reflection Radiometer (ASTER GDEM3), serving as a global DEM, was developed by the Ministry of Economy, Trade, and Industry (METI) of Japan in collaboration with The United States National

Aeronautics and Space Administration (NASA) and was published on 2019. The footprint of ASTER GDEM spans latitudes from 83°N to 83°S. The study area utilized ASTER GDEM3 (Abrams et al., 2020), which can be downloaded from the associated website: <https://gdemdl.aster.jspacesystems.or.jp/>. More information is shown as Table 2.

3.3.5 SRTMv3 DEM

The Shuttle Radar Topography Mission (SRTM) DEM, developed by NASA, was a collaborative effort involving the National Geospatial-Intelligence Agency (NGA) and the German and Italian Space Agencies. It was part of an international project aimed at acquiring radar data, which were used to create the first near-global set of land elevations (Werner, 2001). The DEM was launched in 2000 (Farr et al., 2007), and many improvements have been made since then. The SRTMv3 DEM, the latest version, was used for the study area and can be downloaded from the associated website: <https://search.earthdata.nasa.gov/search>.

3.3.6 MERIT DEM

The Multi-Error-Removed Improved-Terrain (MERIT) DEM, developed by (Yamazaki et al., 2017). MERIT DEM improves upon previous DEMs by systematically removing various error components such as absolute bias, stripe noise, speckle noise, and tree height biasn from SRTM3 DEM (Farr et al., 2007) , AW3D-30 m DEM (Tadono et al., 2015) and gap-filling with the Viewfinder Panoramas (VFP) DEM (<http://viewfinderpanoramas.org/dem3.html>). The MERIT DEM is a DSM with resolution of 3 arc seconds. It was utilized for the study area and is available for download from the dedicated website: http://hydro.iis.utokyo.ac.jp/~yamadai/MERIT_DEM/index.html/.

3.3.7 GLO30 DEM

The Copernicus DEM, published in 2019 by the European Space Agency (ESA) (AIRBUS, 2020), represents an upgraded iteration of the WorldDEM. The backbone of the Copernicus WorldDEM is the TanDEM-X mission data, yet void filling techniques and integration of other data sources are used to enhance data completeness and accuracy. The Copernicus DEM is provided in three different DSM instances: EEA-10, GLO-30, and GLO-90. For this study, GLO-30 was utilized, offering 1 arc-second resolution. It can be downloaded from the dedicated website: <https://spacedata.copernicus.eu/de/collections/copernicus-digital-elevation-model>.

3.3.8 FABDEMv1-2

Forest And Building removed Copernicus Digital Elevation Model (FABDEM) was developed in collaboration between Bristol-based flood modelling company Fathom and the University of Bristol FloodLab. The FABDEM V1-0, launched in 2021 (Laurence Hawker, 2021), is derived from the Copernicus GLO-30 (AIRBUS, 2020) DSM. FABDEM V1-2, released in 2023 (Hawker et al., 2023), has a 1 arc-second resolution and is based on a DSM which removes buildings and vegetation.

This dataset was employed for the study area and is available for download from <https://data.bris.ac.uk/data/dataset/s5hqmjcdj8yo2ibzi9b4-ew3sn>.

3.3.9 TanDEM-X DEM

230 TanDEM-X (TerraSAR-X add-on for Digital Elevation Measurement) is an innovative space borne-radar interferometer based on two TerraSAR-X radar satellites flying in close formation (Krieger et al., 2007). The TanDEM-X mission represents a collaborative effort between the German Aerospace Center (DLR) and AIRBUS (Wessel, 2016), with the aim of generating a globally consistent DEM. TanDEM-X, launched in 2016, is a DSM with resolutions of 0.4, 1, and 3 arcseconds. The 3-arcsecond TanDEM-X product is readily accessible and can be downloaded directly from [https://geoservice.dlr.de](https://geoservice.dlr.de/data-assets/ju28hc7pui09.html)
235 [/data-assets/ju28hc7pui09.html](https://geoservice.dlr.de/data-assets/ju28hc7pui09.html). However, the 0.4 and 1 arcsecond products are available from DLR upon request. It is important to note that the TanDEM-X product has not undergone full processing to eliminate artifacts, outliers, noisy regions, and data gaps. As a result, its adoption in flood modeling has been limited (McClean et al., 2020). In this study, we employed TanDEM-X with a 0.4 arcsecond resolution for our flood modeling purposes.

3.3.10 TanDEM-EDEM

240 The TanDEM-X Edited Digital Elevation model (TanDEM-EDM) is an edited version of the TanDEM-X Global with a 1-arcsec (~30 m) pixel resolution released in 2023 (Wessel, 2016). The main update in TanDEM-EDEM version 1 includes filling gaps with suitable alternative DEM data and improving representation of water bodies. The TanDEM-EDEM dataset, which is a DSM, was utilized for the study area and is readily available for download from: https://download.geoservice.dlr.de/TDM30_EDEM/. It has a resolution of 30 m.

245 3.4 ICESat-2 satellite laser altimetry

Ice, Cloud, and Land Elevation Satellite-2 (ICESat-2) is a laser altimetry satellite launched by the National Aeronautics and Space Administration (NASA) in 2018. As the follow-on satellite of ICESat, ICESat-2 continues elevation measurements of ice sheets, glaciers, sea ice, and various other land features with a 91-day exact repeat orbit. ICESat-2 carries the Advanced Topographic Laser Altimeter System (ATLAS), which works by transmitting 10,000 laser pulses per second using laser light
250 of 532 nm (Neumann et al., 2019). The pulse rate enables the satellite to capture a measurement every 70 cm along the ground track. The pulse divides into six beams, organized into three pairs. Each pair comprises one right-side beam and one left-side beam, striking the Earth at a 90 m distance from each other. The distance between each pair is 3.3 km, as depicted in Figure 1(E).

The National Snow and Ice Data Center (NSIDC) portal has developed various products that incorporate photon travel times
255 and locations determined using the built-in GPS from the ICESat-2 satellite. This mission generates 21 products, as detailed on their website: <https://nsidc.org/data/icesat-2/products>. The two data products used in this study are ATL03 and ATL08, as

summarized in Table 3, and the ground track pattern of ICESat-2 in the study area is shown in Figure 1(b). The ICESat-2 data were obtained from the NSIDC website via their data access tool (<https://nsidc.org/data/data-access-tool>).

Table 3: ICESat-2 product

ICESat-2 product		Data Collection (Year)	Datum Reference
ATL03	Global Geolocated Photon Data (DSM)	2018 - 2022	WGS84 ellipsoid
ATL08	Land/Water/Vegetation Elevation (DSM)	2018 - 2022	WGS84 ellipsoid

260 **3.4.1 ATL08**

The ATL08 product is derived from ICESat-2 ATL03 data, which provides detailed information on time, latitude, longitude, and height for each photon track. This dense photon dataset enables subsequent analyses and the creation of surface-specific products, such as land ice height and sea ice freeboard (Tom Neumann et al., 2021). The ATL08 product offers estimates of terrain heights, canopy heights, canopy cover, and other descriptive parameters at fine spatial scales in the along-track direction. A fixed segment size of 100 m was chosen to provide continuity of data parameters on the ATL08 data product. Height estimates from ATL08 can be compared with other geodetic data and serve as input for higher-level products like ATL13 (inland water-related heights) and ATL18 (terrain and canopy feature maps) (Neuenschwander et al., 2022). In this study, we used ATL08 land heights from ICESat-2 as the benchmark, to which various DEM products were compared.

3.5 Flood Map/Surface Water Extent (SWE) dataset

270 In this study, SWE and flood maps were collected from two sources: surface water extent (SWE) data from the WorldWater project (<https://worldwater.earth/>), funded by the European Space Agency (ESA) and the Geo-Informatics and Space Technology Development Agency (Public organization) (GISTDA) in Thailand. The flood map datasets are summarized in Table 4 and presented in Figure 3.

3.5.1 WorldWater Surface Water Extent (SWE)

275 We used SWE products from the WorldWater project, using data from the Sentinel-1 and Sentinel-2 imaging satellites, both integral parts of the ESA Copernicus program. The Sentinel-1 satellite, launched in 2014, is equipped with a SAR constellation consisting of two polar-orbiting satellites, with objectives on land and ocean monitoring. Sentinel-1 comprises a C-band SAR sensor with a 10-meter spatial resolution (Torres et al., 2012). The Sentinel-2 satellites consist of two satellites, namely Sentinel-2A and Sentinel-2B, launched in 2015 and 2017, respectively. The dual-satellite system operates in coordination with a 180° phase difference in the sun-synchronous orbit, supporting both land and ocean monitoring (European Space Agency (ESA), 2015). The WorldWater SWE mapping algorithm utilized Sentinel-1 and Sentinel-2 data from 2017 to 2021 to develop a SWE dataset. The details of Sentinel-1 and Sentinel-2 dataset are accessible at depicted in Copernicus Open Access Hub. This algorithm utilizes a fusion approach (Tottrup et al., 2022), combining optical and radar observations, to provide a more robust delineation of water surfaces. The SWE products provide information on water occurrence, monthly water presence,

285 water seasonality, maximum and minimum water extent, all accessible on the website: <https://swdap.worldwater.earth/>. The monthly water presence of the Worldwater SWE in November 2017 is illustrated in Figure 3(c). It is important to note that the WorldWater SWE dataset uses a median composite of all Sentinel-1 and Sentinel-2 acquisitions within a given month to predict monthly surface water presence. Consequently, it does not necessarily reflect the maximum extent of flooding within that month.

290 **3.5.2 GISTDA Flood map**

GISTDA is a Thai space agency and space research organization that utilizes satellites such as Cosmo-SkyMed, KOMPSAT, LANDSAT-5, RADARSAT-2, and THAICHOTE (Channumsin et al., 2020) to conduct research and development. GISTDA receives observations of the Earth through the use of Synthetic Aperture Radar (SAR) and optical sensor satellites (Nithirochananont et al., 2010). SAR satellite information is derived from two constellations: RADARSAT and the Advanced
295 Land Observing Satellite (ALOS). RADARSAT comprises two SAR satellites, while ALOS integrates a SAR satellite with an optical satellite. Both RADARSAT and ALOS possess SAR data processing systems. In flooded areas, the Earth's surface appears smooth in the wavelength of the SAR. This smooth surface causes microwaves to reflect in a specular way, resulting in low backscatter values. This characteristic allows for real-time flood imaging and identification. The SAR data undergoes processing, and image quality enhancement while eliminating any noise present in the data products (Auynirundronkool et al.,
300 2012).

To generate flood maps from satellite data, GISTDA employed several analysis methods, including supervised classification, visual analysis, and thresholding, which were combined with field images. Subsequently, GISTDA used the boundaries of natural and permanent water sources from the existing database and removed these areas from the flood map. Since 2005, GISTDA has annually published nowcasting flood maps and flood occurrence maps on <https://flood.gistda.or.th/>, which were
305 utilized in this study. The GISTDA flood occurrence map is shown in Figure 3(b).

From 2014 to 2023, HII analysedanalyzed flood frequency maps from GISTDA. The assessment focused on the frequency of flood occurrences, which were categorized into three levels: low, medium, and high-risk flood frequency. Low-risk flood frequency is defined as 1-3 occurrences within the 10-year span, medium risk as 4-7 occurrences, and high risk as 8-10 occurrences, as depicted Figure 3(a).

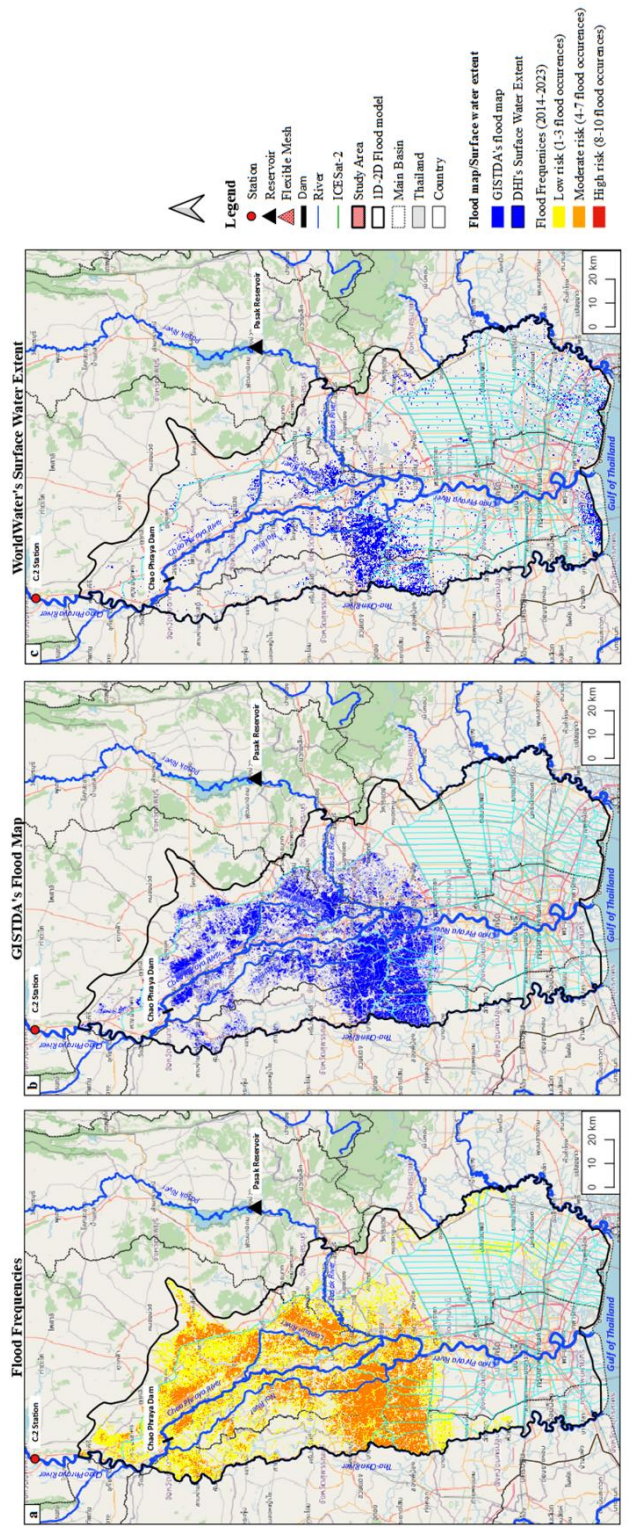


Figure 3: The flood map/surface water extent in study area, presenting (a) flood frequency from HIL, (b) GISTDA's flood map in Nov 2017, and (c) WorldWater's SWE in Nov 2017. © OpenStreetMap contributors 2015. Distributed under the Open Data Commons Open Database License (ODbL) v1.0.

Table 4: Flood map datasets

Product	Resolution (m.)	Period	Frequency	Type file	Download
GISTDA's flood map	-	2005-2021	On request, satellite track, and annual	Shape file	https://flood.gistda.or.th/
WorldWater's surface water extent	10	2017-2021	Monthly and annual	Raster file	https://swdap.worldwater.earth/

4 Methodology

The workflow used in this study, illustrated in Figure 4, comprises two primary components. The first component, namely DEM analysis, focuses on evaluating the DEMs (Sect. 3.3) with the ICESat-2 benchmark (Sect. 3.4) as a high-precision reference, which effectively serves as the "ground truth." The best DEM identified in the DEM analysis is then used as input to the flood map analysis. The flood map analysis, focuses on evaluating flood maps generated by the 1D-2D flood model (Sect. 3.1) against WorldWater SWE and GISTDA flood maps (Sect. 3.5).

4.1 DEM Analysis

The primary objective is to assess the accuracy and reliability of the DEMs by comparing them with elevation data obtained from ICESat-2 using statistical methods. In the study area, ICESat-2 ATL08 data were primarily used for evaluation, while ICESat-2 ATL03 data were employed in complex terrain. Figure 4 (a) illustrates the workflow involving processing and re-referencing steps. Subsequently, the evaluation of DEMs and ICESat-2 is conducted using statistical methods.

4.1.1 ICESat-2 ATL08 Data Processing

ATL08 provides estimates of terrain height, canopy height, and canopy cover at fine spatial scales in the along-track direction. For each parameter, terrain surface elevation and canopy heights were provided at a fixed along-track segment size of 100 meters (Neuenschwander et al., 2022). The ATL08 dataset comprises a total of 18 land parameters, such as Mean terrain height for segment (h_te_mean), Mode of terrain height for segment (h_te_mode), Number of ground photons in segment (n_te_photins), slope of terrain within segment (terrain_slope), Best fit terrain elevation at the 100 m segment mid-point location (h_te_best_fit), and others. We processed the ATL08 dataset, extracting the latitude and longitude of the photon signals along with the photon heights above the WGS84 ellipsoid. The terrain elevation parameter used for evaluation was h_te_best_fit.

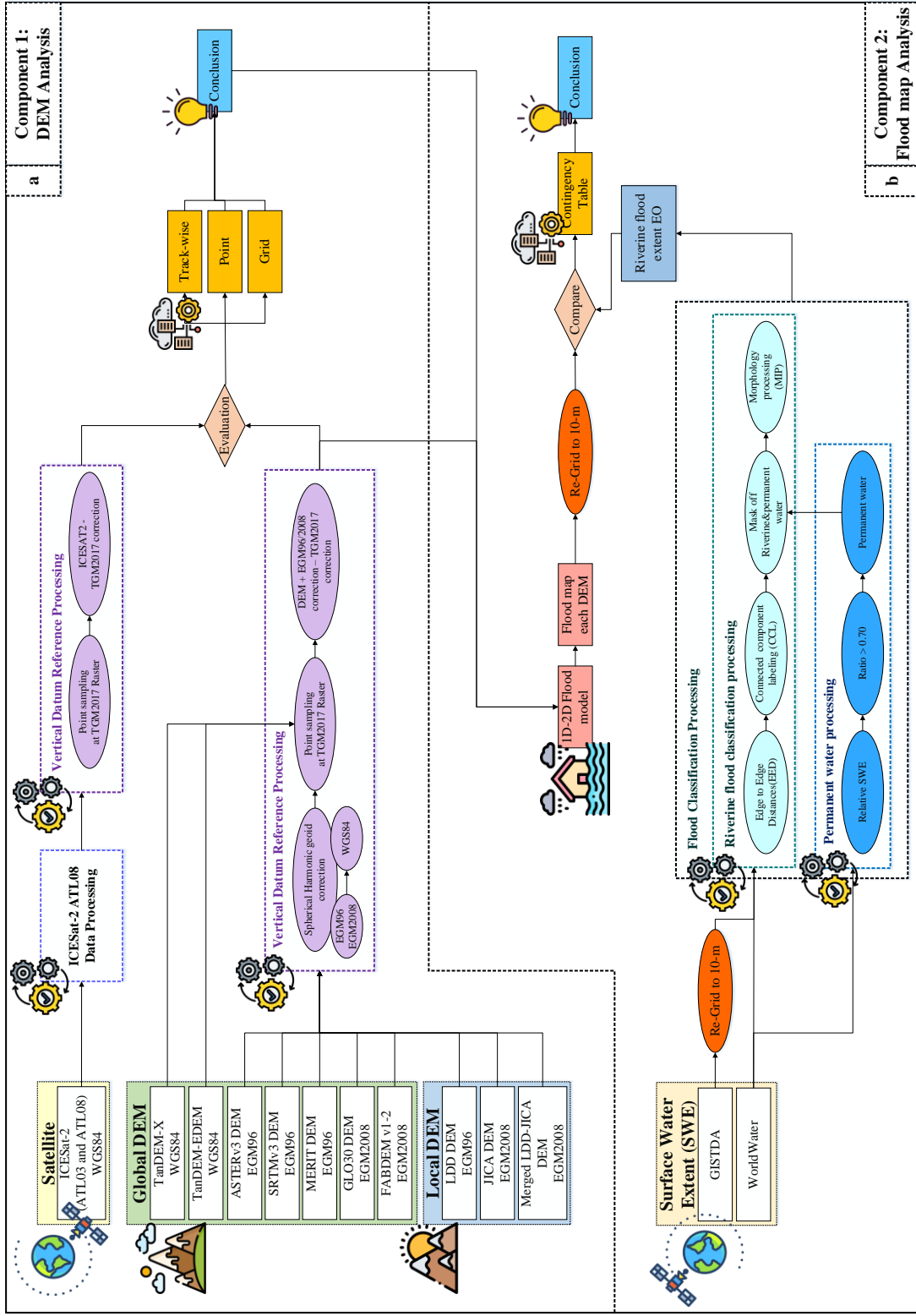


Figure 4: Overall Methodology: (a) Component 1: DEM Analysis – Involves processing ICESat-2 ATL08 data, applying vertical datum referencing, and evaluating DEMs against the ICESat-2 ATL08 benchmark through point, grid, and track-wise comparisons. (b) Component 2: Flood Map Analysis – Includes setting up the 1D-2D flood model, performing flood classification, and evaluating flood maps using appropriate methods.

335 4.1.2 Vertical Datum Reference Processing

To evaluate the DEMs with the ICESat-2 benchmark, it is necessary to use the same vertical datum reference. Vertical datum reference processing was employed to standardize the datum reference. In this study, the vertical datum reference was TGM2017, using Eq. (1) to establish accurate measurements of vertical elevation

$$H = h - N \quad (1)$$

340 Where H is ortometric height, h is ellipsoid height, and N is geoid height.

$$H_{DEM\ ref\ TGM2017} = h_{DEM} + N_{DEM} - N_{TGM2017} \quad (2)$$

Where $H_{DEM\ ref\ TGM2017}$ is the DEM referenced to TGM2017, h_{DEM} represents the original DEM, N_{DEM} is the geoid reference of the original DEM, $N_{TGM2017}$ is TGM2017 geoid model.

To obtain DEMs referenced to TGM2017, EGM96 and EGM2008 height corrections were added to the DEM heights, followed
345 by subtracting the TGM2017 geoid corrections, as shown in Eq. (2). The geoid model datasets are shown in Sect. 3.2 for reference. For ICESat-2 elevations referenced to TGM2017, the TGM2017 correction was subtracted from the ICESat-2 elevation data.

4.1.3 Evaluation of DEMs using ICESat-2 ATL08 Benchmark

The DEM products were estimated and evaluated using statistical methods, including bias (mean error, ME), mean absolute
350 error (MAE) (Willmott, 2005), mean square error (MSE), root mean square error (RMSE) (Chai & Draxler, 2014) and percent bias (PBIAS) (D. N. Moriasi et al., 2007). The overall purpose of implementing these statistical methods is to evaluate the ICESat-2 ATL08 data paired with the 10 DEM products covering the study area. Subsequently, the performance of the DEMs was systematically compared using statistical indices defined in Eqs. (3)–(6) (Samantaray & Sahoo, 2024).

$$ME = \frac{1}{n} \sum_{i=1}^n (Y_i - \hat{Y}_i) \quad (3)$$

$$355\ MAE = \frac{1}{n} \sum_{i=1}^n |Y_i - \hat{Y}_i| \quad (4)$$

$$MSE = \frac{1}{n} \sum_{i=1}^n (Y_i - \hat{Y}_i)^2 \quad (5)$$

$$RMSE = \sqrt{\sum_{i=1}^n \frac{(Y_i - \hat{Y}_i)^2}{n}} \quad (6)$$

$$PBIAS = 100 \times \left[\frac{\sum_{i=1}^n (Y_i - \hat{Y}_i)}{\sum_{i=1}^n \hat{Y}_i} \right] \quad (7)$$

Where \hat{Y}_i represents ICESat-2 ATL08 elevation, Y_i denotes the elevation for each DEM (i.e., LDD DEM, JICA, merged LDD-
360 JICA DEM, ASTEM GDEM V3, SRTM DEM, MERIT DEM, FABDEM v1-2 DEM, GLO30 DEM, TanDEM-X, and TanDEM-EDEM), and n is the number of observations. The ideal PBIAS value is 0: positive values indicate that the DEM products tend to overestimate compared to the ICESat-2 ATL08 benchmark, while negative values indicate a tendency toward underestimation.

We conducted three types of comparisons as follows:

365 **Point comparison**

Point comparison was performed for every segment of the ICESat-2 ATL08 pass over the study area. This approach aimed to provide a quantitative overview of the quality and identify potential discrepancies among the ten DEMs in comparison to ICESat-2 ATL08 data (Weifeng et al., 2024), using statistical methods. A total of 954,800 elevation points were extracted from the study area for point-to-point comparison.

370 **Grid comparison**

The grid comparison was conducted using a regular square grid over the study area. This comparison provides an overview of the spatial variation of the quality of the DEMs in comparison to ICESat-2 ATL08 benchmark. In this study, we employed a 5-km resolution for grid comparison, which involved calculating statistical measures for every segment within each grid cell and displaying the evaluation spatially on a map.

375 **Track-wise comparison**

The track-wise comparison was conducted using tracks of ICESat-2 over the study area. The distance between the ICESat-2 points was calculated using UTM x and y coordinates, as shown in Eq.(7). The track-wise comparison represents an overall elevation profile comparison between DEMs and ICESat-2 ATL08 data over the study area.

$$distance = \sqrt{(x_0 - x_i)^2 + (y_0 - y_i)^2} \quad (7)$$

380 Where x represents the x coordinates, and y denotes the y coordinates.

4.2 Flood Map Analysis

The purpose of the flood map analysis is to evaluate the performance of simulated flood maps from the 1D-2D flood model using various DEM products selected from the first component in comparison to the WorldWater SWE and GISTDA flood maps. This comparative analysis aims to assess the accuracy and effectiveness of the improved flood simulation model.

385 **4.2.1 1D-2D Flood Modelling Setup**

The setup of the 1D-2D flood model mirrored the original model, retaining the same parameters with only the DEM being modified to generate the flood map. The DEM products were selected based on the evaluation of DEMs against the ICESat-2 ATL08 benchmark. Flood maps in the lower CPY basin were simulated using the 1D-2D flood model for the years 2017, and 2021. The flood map simulation results from the 1D-2D flood model present flood extents that occurred during the simulation
390 period and at each daily time step (DHI, 2018). In this study, we employ simulated flood maps generated from a 1D-2D flood model using the merged LDD-JICA DEM and FABDEMv1-2 DEM products and compare them with WorldWater SWE and GISTDA flood maps.

4.2.2 Flood Classification Processing

The flood map and SWE dataset used for evaluation in this study (Sect. 3.5) had different resolutions, formats, and flood map definitions. To effectively assess the simulated flood map from the 1D-2D flood model, we compared it to the WorldWater SWE and GISTDA flood map. However, it is crucial to employ the same resolution, format, and flood definition. Common types of flooding include flood irrigation, pluvial flash floods, coastal floods, and riverine floods. The 1D-2D flood model only simulates riverine floods, caused by high water levels in the rivers, eventually overflowing onto the neighboring land due to high river discharge over an extended period. In order to compare the simulated flood map to the satellite EO products, we first have to extract riverine flooding patterns from the surface water extent maps provided by satellite EO. This is done in the following steps:

Permanent water processing

Permanent water bodies should be removed from the satellite EO SWE maps prior to comparison. The GISTDA datasets does not include permanent water bodies. The WorldWater product includes permanent water bodies, which must be removed prior to comparison with simulated flood maps. We use relative water frequency (Yamazaki et al., 2015), which measures the occurrence of surface water within a defined time period. The relative water frequency fr of pixel was defined by Eq.(8) and shown as Figure A4 (a).

$$f_r(t) = \frac{f_a(t)}{f_v(t)} \quad (8)$$

Where f_a depicts the frequency of surface water detections during a certain time period for each pixel, and f_v represents the frequency of valid observations during the same period for each pixel.

The relative water frequency ranges between 0.0 to 1.0. The permanent water designation indicates that there was observed water coverage in every single observation of the considered time period, which corresponds to a relative water frequency of 1.0 (Martinis et al., 2022). In many cases, lower thresholds of 0.9, 0.7, and 0.5 were applied ((Rao et al., 2018); (Yamazaki et al., 2015)). The permanent water map for each threshold is illustrated in Figure A4. In this study, the threshold for relative water frequency is set to 0.7, indicating that a pixel is considered permanent water if it is present in 70% or more of the valid observations over the specified time period. The output of the permanent water processing is utilized in riverine flood classification processing to remove permanent water from the WorldWater SWE.

Riverine flood classification processing

The WorldWater and GISTDA datasets contain both riverine floods and other inundated areas caused, for instance, by irrigation or pluvial floods. In order to separate riverine floods in the satellite EO flood maps, we used the following method Figure A5:

- Expand the wet area from WorldWater and GISTDA by 200 meters using expand segmentation labels (ESL) without overlap (Van Der Walt et al., 2014). The ESL method merges labels in a label image based on the distances between each pixel. Labels that are close by will be merged.

425 ▪ Subsequently, label each pixel using connected component labeling (CCL) ((Rosenfeld & Pfaltz, 1966) and (AbuBaker et al., 2007)). The CCL method is employed to detect connected regions in binary digital image. The assumption of riverine flood identification is based on the presence of wet connected pixels originating from the river. These are then masked off using ESL, and the riverine flood label is selected.

430 ▪ Subsequently, the SWE undergoes morphological image processing (MIP) using a closing algorithm (Van Der Walt et al., 2014). The structuring element, footprint, passed to the closing algorithm is a boolean array describing the neighborhood. We used a disk to create a circular structuring element with a radius of 2, implemented as the footprint. The output provides riverine flood maps, namely WorldWater and GISTDA flood map, for evaluation with other flood map products.

4.2.3 Flood map evaluation methods

435 This study evaluates the flood map of the lower CPY River basin using the contingency table (“Glossary of Terms,” 1998), comparing flood maps from two different dimensions, as shown in Table 5. We evaluated the flood maps produced by the 1D-2D flood model by comparing them with the monthly surface water presence maps from WorldWater and GISTDA for the years 2017 and 2021. We mainly use probability of detection (POD), false alarm ratio (FAR), and critical success index (CSI) (Forecast, 1995) to perform the evaluation. These statistics are based on the number of grid cells or pixels in the study area is

440 defined as:

$$POD = \frac{Hit}{Hit+Miss} \tag{9}$$

$$FAR = \frac{False\ alarm}{Hit+False\ alarms} \tag{10}$$

$$CSI = \frac{Hit}{Hit+False\ alarms+Miss} \tag{11}$$

Where, *Hit* represents the number of correctly detected flooded pixels from two different dimensions. *True negative* donates the number of correctly detected non-flooded or dry areas from two different dimensions. *Miss* indicates the number of floods from dimension 1 that are not detected by dimension 2, while *False Alarm* represents the number of floods from dimension 2 which did not occur floods in dimension 1. A perfect score for both POD and CSI is 1, while a value of 0 represents the best score for FAR.

445

Table 5: Contingency table

Modelled flood map		Observed flood map	
		Flood	Unflood
		Hits	False alarms
		Misses	True negative

5.1 1D-2D Flood model calibration results

The 1D river model was calibrated using in-situ water surface elevation data for the period 2012 to 2013. The calibration results of the main river in the study area are presented in (Charoensuk et al., 2024). The overall performance during the calibration period is generally satisfactory for all main rivers, with an average R^2 of 0.96, RMSE of 0.30 m, and NSE of 0.90.

455 The 1D-2D flood model has been calibrated for extreme floods in 2011, as presented in (Charoensuk et al., 2018). Normally, flooding in Thailand is influenced by meteorological conditions, river conveyance, and sea level rise. However, the primary cause of the 2011 flood was dike breaching along the Chao Phraya River, resulting in uncontrollable flood inundation. The simulated flood, when compared with the GISTDA's flood map, satisfactorily corresponds to flood depth, flood propagation direction, and duration.

460 **5.2 Results of DEM evaluation against the ICESat-2 ATL08 benchmark**

5.2.1 Point comparison evaluation results

Figure 5 illustrates point comparisons between the statistical metrics of 10 DEM products against ICESat-2 ATL08 benchmark. As depicted in the Figure 5(a), the average ME of the local DEM products was -0.88 m, whereas the average ME of global DEM products was +1.62 m. The results indicate that local DEM products tend to have negative bias, while global DEM

465 products tend to show positive bias when compared against ICESat-2 ATL08 benchmark. This tendency is attributed to the algorithms described in Sect. 3.3, which remove buildings and vegetation from the local DEM products. Moreover, the local DEM products have a finer grid resolution compared to the global DEM products. The average performance statistics of the local DEM and global DEM were 1.25 and 2.17 m for MAE, 4.23 m and 13.52 m for MSE, and 2.04 and 3.38 m. for RMSE, as shown in Figure 5(b), Figure 5(c) and Figure 5(d) respectively.

470 Table 6 presents the statistical results of point comparisons between 10 DEM products compared with ICESat-2 ATL08, indicating that the accuracy of JICA DEM and FABDEMv1-2 DEM was higher than other local and global DEM, respectively. The statistical results of JICA DEM were -0.65 m, 1.04 m, 3.51, 1.87 m, and -17.00% for ME, MAE, MSE, RMSE, and PBias, respectively. Specifically, the FABDEMv1-2 DEM showed the highest accuracy, with ME, MAE, MSE, RMSE, and PBias values of 0.25 m, 0.80 m, 3.79, 1.95 m, and 4.59% respectively.

475 Figure 6 presents the histogram distributions of ME for 10 DEM products relative to ICESat-2 ATL08 benchmark. The histogram distributions illustrate that the entire curve of both local and global DEMs shifts towards negative and positive biases, respectively. These shifts indicate that local DEMs, including LDD DEM, JICA DEM, and merged LDD-JICA DEM, exhibit a negative bias in elevation relative to the ICESat-2 ATL08 benchmark, with ME averages of -1.30 m, -0.65 m, and -0.68 m, respectively.

480 Conversely, the shifts observed in the histogram distribution of global DEMs, including ASTERV3 DEM, SRTMv3 DEM, Merit DEM, GLO30 DEM, FABDEMv1-2 DEM, TanDEM-X DEM, and TanDEM-EDEM DEM, indicate a positive bias of

the elevation of ICESat-2 ATL08 benchmark. The ME averages for these DEMs were +4.78 m, +2.03 m, +1.56 m, +0.84 m, +0.25 m, +0.94 m, and +0.91 m, respectively. Further details are provided in Figure A7 and Figure A8, illustrating the mean absolute error (MAE) and mean squared error (MSE), respectively.

Table 6: Table of statistical metrics, comparing 10 DEM products against the ICESat-2 benchmark. The resulting averages are computed across the datasets in study area.

DEM product	Scale	Statistical method				
		ME (m.)	MAE (m.)	MSE (m.)	RMSE (m.)	PBIAS (%)
LDD	Local	-1.30	1.64	5.45	2.33	-34.76
JICA	Local	-0.65	1.04	3.51	1.87	-17.00
merged LDD-JICA	Local	-0.68	1.08	3.74	1.93	-15.38
Average local DEMs		-0.88	1.25	4.23	2.04	-22.38
ASTER	Global	+4.77	5.57	44.28	6.65	47.71
SRTM	Global	+2.04	2.58	12.92	3.59	27.99
MERIT	Global	+1.56	1.79	6.76	2.6	22.99
GIO30	Global	+0.84	1.3	5.89	2.43	13.87
FABDEMv1-2	Global	+0.25	0.8	3.79	1.95	4.59
TanDEM-X	Global	+0.94	1.73	13.29	3.65	15.24
TanDEM-EDEM	Global	+0.91	1.43	7.74	2.78	14.84
Average global DEMs		+1.62	2.17	13.52	3.38	21.03

5.2.2 Grid comparison evaluation results

Figure 7 displays the ME spatial grid comparison of 10 DEM products against the ICESat-2 ATL08 benchmark, with a resolution of 5x5 km. As shown in the figure, the local DEMs indicated overall lower values than the benchmark, with LDD DEM showing the lowest ME. In contrast, the overall ME spatial grid comparison of global DEMs was higher than the benchmark and clearly reveals that most global DEMs exhibit poor performance in urban areas. Notably, in the lower middle of the study area lies Bangkok, the capital city of Thailand. However, the FABDEMv1-2 DEM performed better in urban areas compared to other global DEMs, which can be attributed to the fact that vegetation and buildings are eliminated in this DEM described in Sect. 3.3.8 and (Dandabathula et al., 2023).

5.2.3 Track-wise comparison evaluation results

The track-wise comparison involves comparing the land elevation profile over the study area between the 10 DEM products and ICESat-2 ATL08 benchmark (cf. Figure 8). As shown in Figure 8, it is evident that the local DEMs exhibit lower land elevation compared to the ICESat-2 ATL08 benchmark. For most of the tracks, the LDD DEM measures a lower elevation than benchmark, while the JICA and merged LDD-JICA DEM follow the ICESat-2 ATL08 measurements more closely. This trend is consistent along the majority of the tracks, indicating that the LDD DEM a negative bias in elevation when compared to ICESat-2. Additionally, both the JICA and merged LDD-JICA DEMs closely track the ICESat-2 measurements for most of the tracks. Moreover, local DEMs show lower elevations in urban areas, in agreement with ICESat-2 ATL08. However, we

expect that both local DEMs and ICESat-2 ATL08 still have residual positive bias compared to the true bare earth elevation in urban areas.

505 The overall result of the track-wise comparison of global DEMs shows a higher elevation than the benchmark, especially in urban areas, clearly indicating higher elevations in these urban areas, as illustrated in Figure 8. In most tracks, ASTERv3 and SRTMv3 DEMs exhibit a notable positive bias and fluctuations compared to the benchmark. Meanwhile, Merit, GLO30, TanDEM-X, and TanDEM-EDEM DEMs tend to follow a fluctuating pattern and measure slightly higher than the benchmark's track. FABDEMv1-2 closely aligns with the benchmark, indicating its strong performance. The track-wise comparison
510 provides more detailed information in Appendix A.

The summary results of the evaluation of the 10 DEM products are presented in the parallel plot shown Figure 9, which displays the 10 DEM products along with the results of statistical methods including MAE, RMSE, and DEM resolution. In the local DEM products, it is notable that the LDD DEM exhibits higher error and resolution compared to the JICA and merged LDD-JICA DEMs. Both the JICA and merged LDD-JICA DEMs demonstrate similar accuracy, but the JICA DEM does not cover
515 the entire study area (Figure 2). Therefore, we utilized the merged LDD-JICA DEM from the local DEM product to implement the 1D-2D flood model. For the global DEM product, the FABDEMv1-2 demonstrates the best performance compared to other global DEM products. Therefore, we selected the FABDEMv1-2 DEM to implement in the 1D-2D flood modelling, even though its spatial resolution is lower than TanDEM-X DEM.

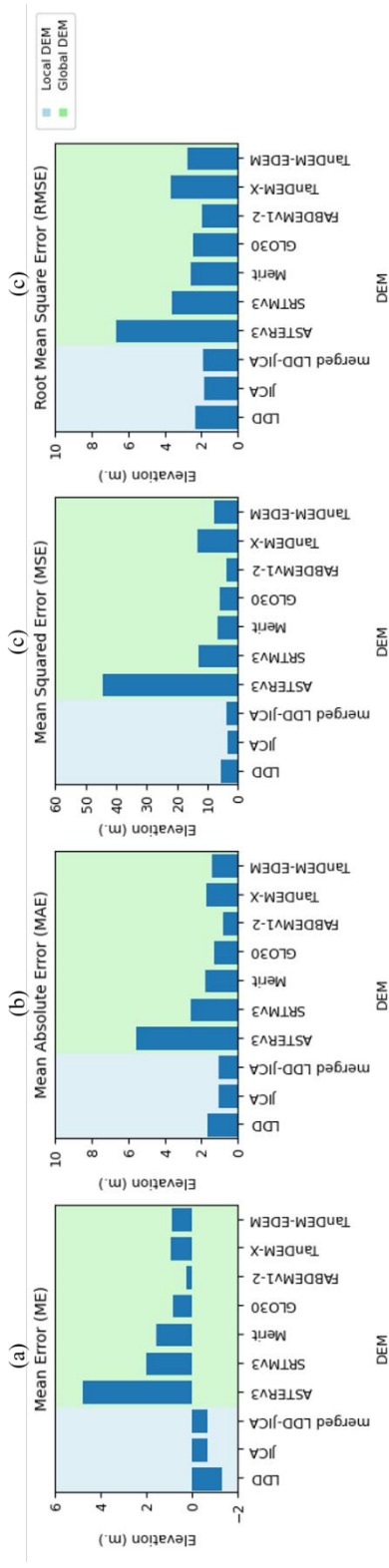


Figure 6: a bar chart of statistical metrics, comparing 10 DEM products against the ICESat-2 ATL08 benchmark, including (a) mean error (ME), (b) mean absolute error (MAE), (c) mean squared error (MSE), and (d) root mean square error (RMSE). The resulting averages are computed across the datasets in study area.

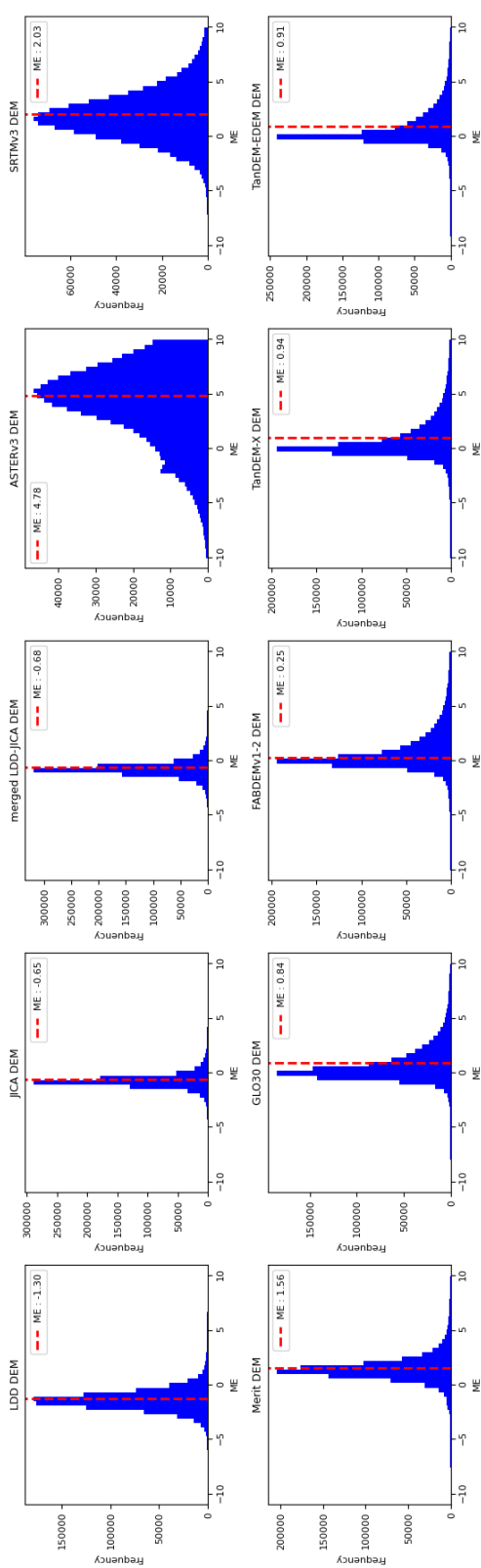


Figure 5: histogram distribution of the mean error (ME), comparing 10 DEM products against the ICESat-2 ATL08 benchmark.

Mean Error (ME) resolution 5x5 km.

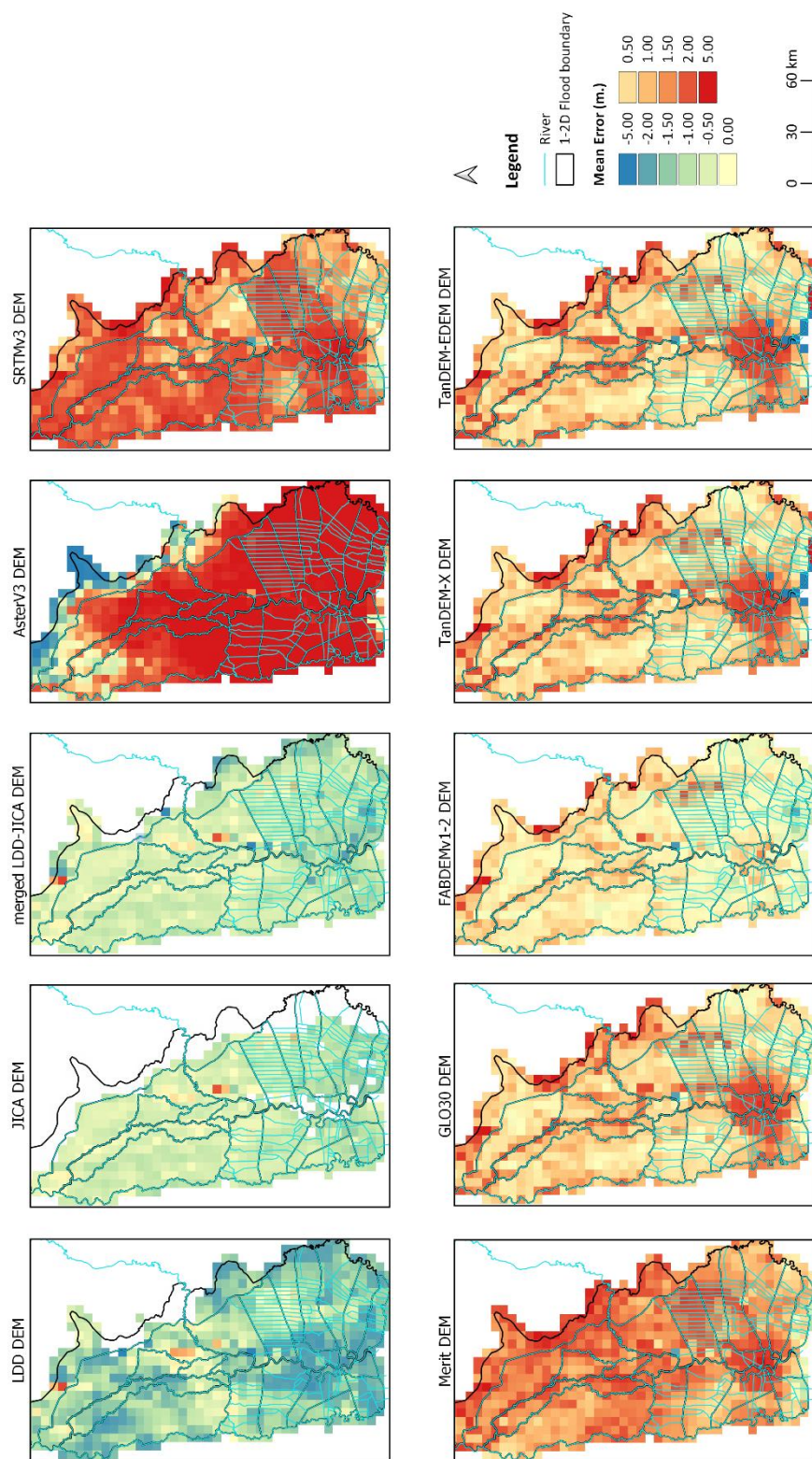


Figure 7: the mean error (ME) spatial grid comparison of 10 DEM products against the ICESat-2 ATL08 benchmark, with a resolution of 5x5 km.

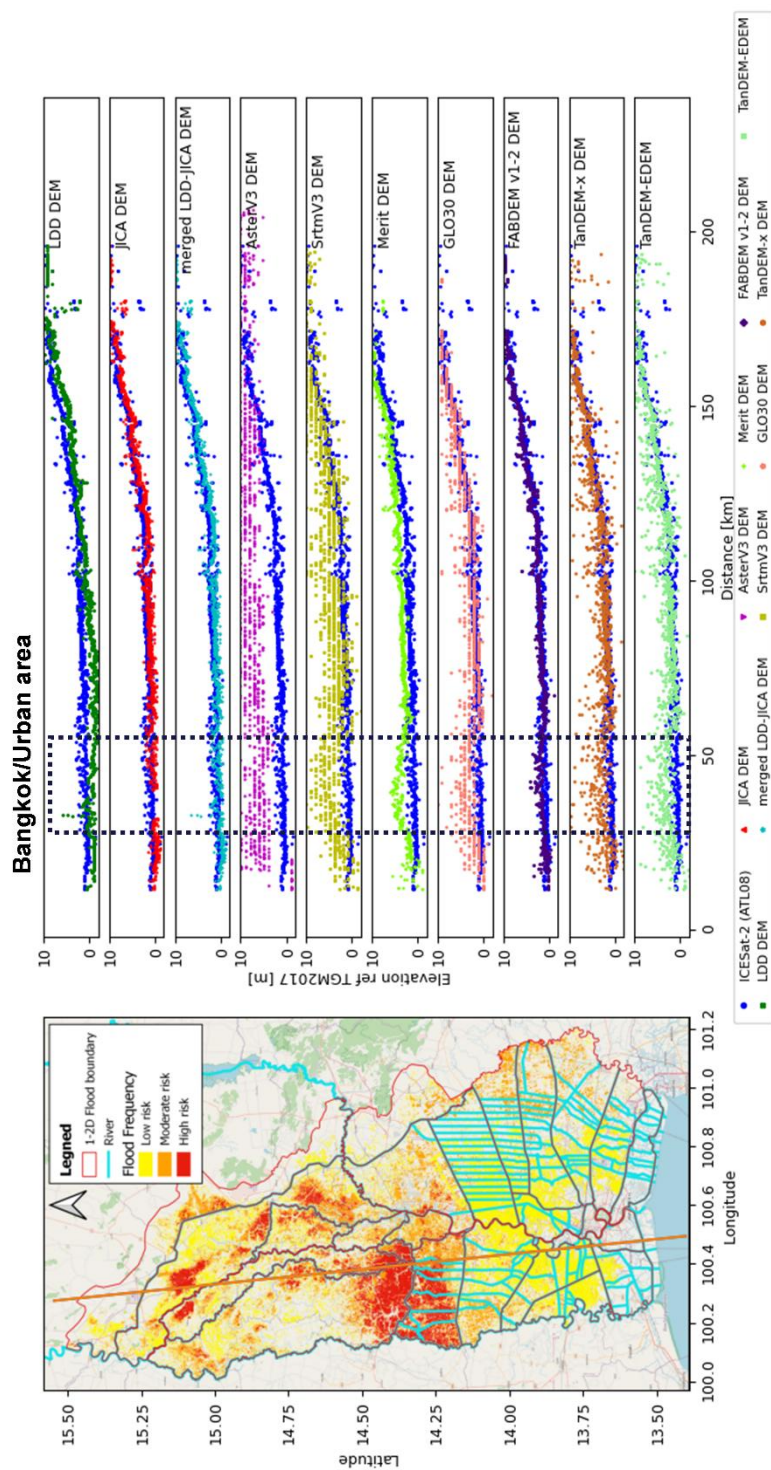
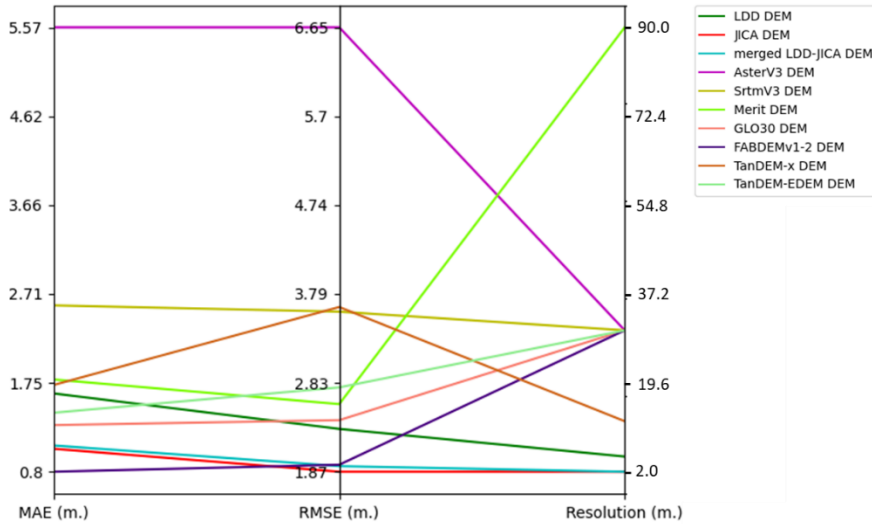


Figure 8: The track-wise comparison of 10 DEM products with ICESat-2 ATL08 benchmark. © OpenStreetMap contributors 2015, Distributed under the Open Data Commons Open Database License (ODbL) v1.0.



525 **Figure 9: The parallel plot of 10 DEMs evaluation with the ICESat-2 ATL08 benchmark.**

5.3 Results of the evaluation of flood inundation maps

We evaluated simulated flood maps produced using two DEMs: 1) the merged LDD-JICA DEM and 2) the FABDEMv1-2 DEM, as described in Section 5.2 from the 1D-2D flood model. The simulated flood map generated by the 1D-2D flood model, referred to as the Model flood map, were evaluated using flood maps from WorldWater and GISTDA for September, October, and November (flood season) in the years 2017 and 2021. The 1D-2D flood model generated daily simulated flood maps. To ensure accurate comparisons, we selected the dates of satellite passes over the study area according to WorldWater and GISTDA datasets. These dates were then combined to represent the flood areas that occurred in each month. The results of the flood map evaluation were categorized based on the DEM and compared to the flood maps from WorldWater and GISTDA. Table 7 provides a comparison of the POD, FAR, and CSI scores for the flood simulation using the merged LDD-JICA DEM, month, and year. Overall, the flood model using the merged LDD-JICA DEM tends to overestimate flooding, particularly in the eastern part of the study area. This overestimation in the eastern part of the study area was attributed to the boundary between the JICA and LDD DEMs in the merged LDD-JICA DEM. The average FAR values of 0.926 and 0.790, along with POD values of 0.713 and 0.585 compared to WorldWater and GISTDA flood maps, respectively, indicate that the Model flood map portrays a larger flood extent while still effectively detecting floods. The average CSI values of 0.072 and 0.183 indicate low model performance and a reflection of the larger flood extent simulation when compared to the flood maps by WorldWater and GISDTA. The overall flood map evaluation based on the FABDEMv1-2 DEM indicates that the Model flood map tends to overestimate, with average FAR values of 0.916 and 0.730 compared to WorldWater and GISTDA flood maps, respectively. Meanwhile, the average CSI values of 0.081 and 0.230 indicate low performance.

Figure 10 shows flood maps and contingency tables for September, October, and November in 2017 and 2021. Figure 10(a-1) presents contingency tables comparing WorldWater monthly SWE and Model flood maps based on the merged LDD-JICA

DEM in 2017. The results of the evaluation show low CSI values of 0.046, 0.071, and 0.076 for September, October, and November in 2017, respectively, indicating that the Model flood map based on the merged LDD-JICA DEM has low performance. Additionally, the number of False alarms was high, resulting in high FAR values of 0.952, 0.926, and 0.923 for September, October, and November in 2017, respectively. Figure 10(b-1) illustrates contingency tables comparing GISTDA and Model flood maps based on the merged LDD-JICA DEM in 2017. The POD values of 0.259, 0.567, and 0.642 are due to the high number of Misses, particularly in September in the upper part of the study area. Moreover, the results show more false alarms in the eastern part of the study area, attributed to the combination of LDD and JICA DEMs. The FAR values of 0.913, 0.727, and 0.699 for September, October, and November in 2017, respectively. The CSI values were low in September at 0.070 but increased to 0.226 and 0.258 for October and November, respectively. The detailed statistics are summarized in Table 7.

Figure 10(a-2) and Figure 10 (b-2) present contingency tables comparing WorldWater and Model, and GISTDA and Model for each month in 2021, respectively. The results of flood map evaluation in 2021 followed a similar trend to that of the 2017 flood. In Figure 10(a-2), low CSI values of 0.091, 0.071, and 0.075 are depicted for September, October, and December in 2021, respectively. Additionally, FAR values of 0.903, 0.928, and 0.923, and POD values of 0.593, 0.845, and 0.835, were high observed for September, October, and November in 2021, respectively. These values suggest that the WorldWater flood map indicates a smaller flood extent compared to the Model flood map based on the merged LDD-JICA DEM. Figure 10(b-2) illustrates an increase in CSI values to 0.133, 0.214, and 0.200 for September, October, and November in 2021, respectively, confirming that the Model flood map based on the merged LDD-JICA DEM fit the GISTDA flood map as well. However, the FAR values were high at 0.852, 0.760, and 0.790 for September, October, and November in 2021, respectively, indicating that the Model flood map based on the merged LDD-JICA DEM shows overestimated flood extents. Despite this, the POD values of 0.564, 0.667, and 0.810 suggest that the Model flood map based on the merged LDD-JICA DEM can effectively detect GISTDA flood map extents, particularly in October and November.

Figure 11 shows flood maps and contingency tables in 2017 and 2021. Figure 11(a-1) illustrates contingency tables comparing WorldWater and Model flood maps based on FABDEMv1-2 DEM for each month in 2017. The evaluation results clearly indicate that the Model flood tends to overestimate the extent of flooding, as evidenced by FAR values of 0.946, 0.913, and 0.914 and low CSI values of 0.052, 0.084, and 0.085 in September, October, and November in 2017, respectively. However, the POD values were high, with values of 0.625, 0.710, and 0.907 in September, October, and November, respectively, indicating that the Model flood map based on FABDEMv1-2 DEM can effectively corresponds to detect the WorldWater flood map as well, as shown in Table 7. Figure 11(b-1) presents contingency tables comparing GISTDA and Model floods for each month in 2017. The Figure 11(b-1) confirms the observations made in Figure 11(a-1), indicating that the Model flood map tends to overestimate the extent of flooding compared to the GISTDA flood map. However, the FAR values decrease slightly to 0.834, 0.612, and 0.591, and the POD values decrease to 0.331, 0.664, and 0.672 in September, October, and November in 2017, respectively. The decrease in POD is attributed to a higher number of Misses in the upper part of the study area, suggesting that the GISTDA flood map depicts more flooding than the Model flood map based on FABDEMv1-2 DEM. On the other hand, the CSI improved to 0.124, 0.325, and 0.341 in September, October, and November in 2017, respectively,

580 indicating that the model results are more accurate when compared with GISTDA flood map. Additionally, the figure illustrates that the GISTDA flood map shows a greater extent of flooding compared to the WorldWater flood map.

Figure 11(a-2) and Figure 11(b-2) depict contingency tables comparing WorldWater and Model, and GISTDA and Model for each month in 2021, respectively. The Model flood map based on FABDEMv1-2 DEM exhibits an overestimation of flooding, particularly noticeable in the eastern part of the study area. Figure 11(a-2) illustrates high FAR values of 0.887, 0.920, and 0.916 that indicating that there are more False alarms in September, October, and November, respectively. The POD was high values of 0.584, 0.885, and 0.850 and low CSI values of 0.105, 0.079, and 0.083 in September, October, and November in 2021, respectively. This figure illustrates that the Model and the WorldWater flood map indicates more and less flooding, respectively. Figure 11(b-2) reveals more Misses in the upper part of the study area, resulting in a decrease in the POD values to 0.502, 0.680, and 0.837 compared to Figure 11(a-2). Despite this, the FAR values remain high at 0.832, 0.738, and 0.776, particularly notable in the eastern part of the study area. However, the Model flood map effectively correspond to detects the GISTDA flood map as well. The CSI values of 0.144, 0.234, and 0.215 for September, October, and November in 2021, respectively, indicate that the Model flood map exhibits improved accuracy in comparison to the GISTDA flood map. The overall assessment of the Model flood map, based on both the merged LDD-JICA and FABDEMv1-2 DEMs, indicates an overestimation of flood extent compared to both WorldWater and GISTDA flood maps. When comparing the model flood map based on the merged LDD-JICA DEM and FABDEMv1-2 DEM with each of the WorldWater and GISTDA flood maps, the results consistently indicate a slight improvement in performance for the Model flood map based on FABDEMv1-2. The CSI of the Model flood map based on FABDEMv1-2 increases by 0.010 and 0.047 compared to the Model flood map based on the merged LDD-JICA DEM for WorldWater and GISTDA flood maps, respectively. Additionally, the FAR is reduced by approximately 0.010 and 0.060 for WorldWater and GISTDA flood maps, respectively. Although the study used flood classification processing to extract riverine flood maps from the SWE map for comparison, there are still limitations. Continuous improvement in flood classification process are necessary. The study results show that the overall assessment of flood simulation based on FABDEMv1-2 DEM reveals a slight improvement of 13.55-25.56% in terms of the CSI compared to flood simulation based on the merged LDD-JICA DEM. However, the DEM is one factor contributing to improved performance, many other factors still require further improvement.

605 Figure 12 illustrates the overall performance of the Model flood map, based on both the merged LDD-JICA and FABDEMv1-2 DEMs. The results are presented in three box plots, corresponding to the evaluation metrics: POD, FAR, and CSI. FABDEMv1-2 exhibits a slightly higher median POD than the merged LDD-JICA, indicating a better ability to correctly detect flood events. The interquartile range (IQR) for the merged LDD-JICA is wider, suggesting greater variability in detection performance compared to FABDEMv1-2, which shows more consistent POD values. Both DEMs show relatively high FAR values, with the merged LDD-JICA having a slightly higher median FAR, indicating it generates more false alarms. FABDEMv1-2 has a smaller IQR, reflecting more consistent performance in minimizing false alarms compared to the merged LDD-JICA. Additionally, FABDEMv1-2 demonstrates a significantly higher median CSI than the merged LDD-JICA,

reflecting superior overall performance in balancing correct detections and false alarms. The narrower IQR for FABDEMv1-2 suggests more consistent performance, while the merged LDD-JICA shows greater variability in CSI values.

Table 7: The statistical metrics of the contingency table, comparing flood map dimensions 1 and 2.

Time	Dimension 1: WorldWater Dimension 2: Model			Dimension 1: GISTDA Dimension 2: Model			DEM product
	POD	FAR	CSI	POD	FAR	CSI	
2017-09	0.549	0.952	0.046	0.259	0.913	0.070	merged LDD-JICA
2017-10	0.612	0.926	0.071	0.567	0.727	0.226	merged LDD-JICA
2017-11	0.842	0.923	0.076	0.642	0.699	0.258	merged LDD-JICA
2021-09	0.593	0.903	0.091	0.564	0.852	0.133	merged LDD-JICA
2021-10	0.845	0.928	0.071	0.667	0.760	0.214	merged LDD-JICA
2021-11	0.835	0.923	0.075	0.810	0.790	0.200	merged LDD-JICA
Total average	0.713	0.926	0.072	0.585	0.790	0.183	merged LDD-JICA
2017-09	0.625	0.946	0.052	0.331	0.834	0.124	FABDEMv1-2
2017-10	0.710	0.913	0.084	0.664	0.612	0.325	FABDEMv1-2
2017-11	0.907	0.914	0.085	0.672	0.591	0.341	FABDEMv1-2
2021-09	0.584	0.887	0.105	0.502	0.832	0.144	FABDEMv1-2
2021-10	0.885	0.920	0.079	0.680	0.738	0.234	FABDEMv1-2
2021-11	0.850	0.916	0.083	0.837	0.776	0.215	FABDEMv1-2
Total average	0.760	0.916	0.081	0.614	0.730	0.230	FABDEMv1-2

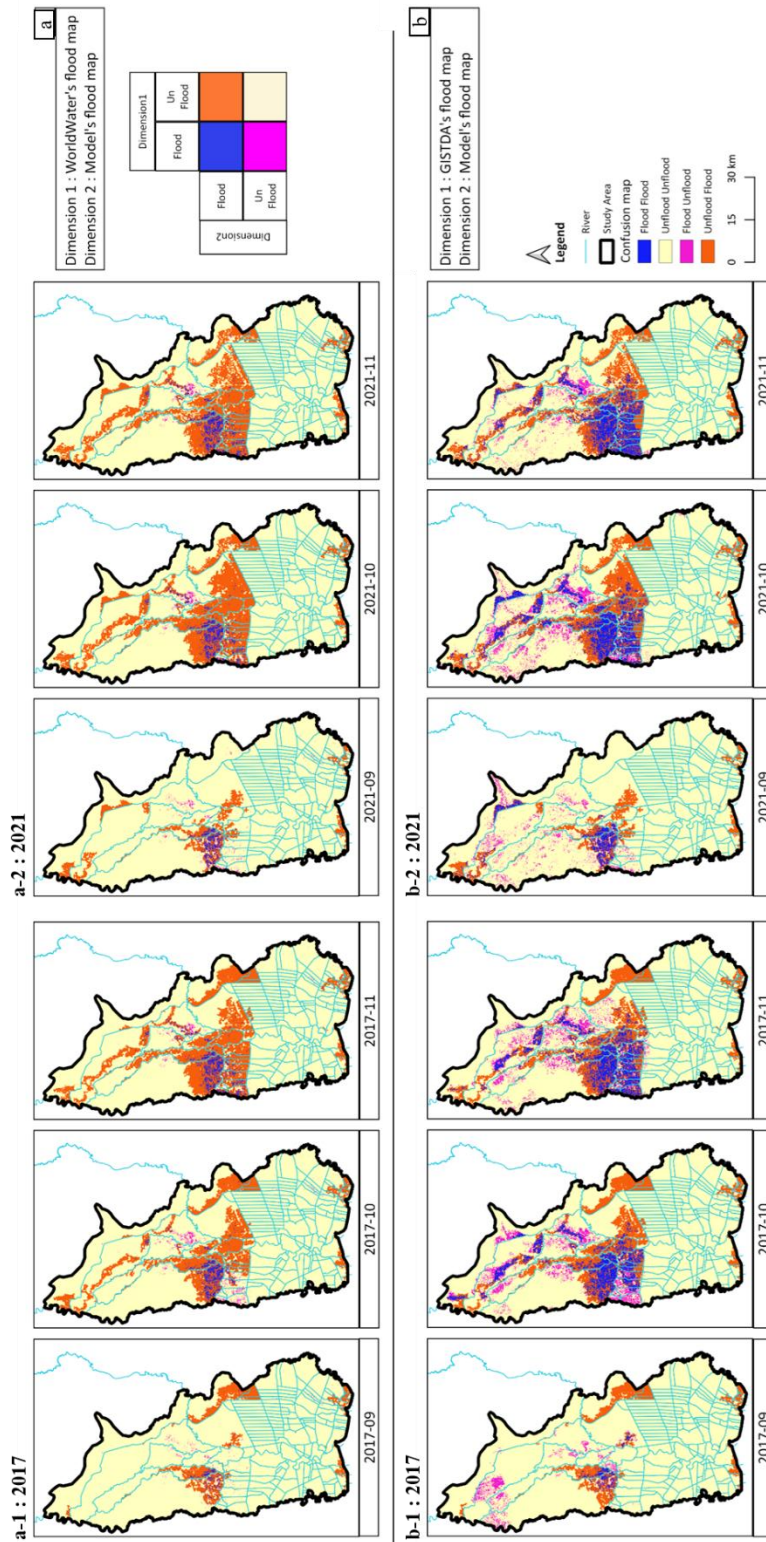


Figure 10: the contingency table of dimension 1 and dimension 2 flood maps on the spatial map: (a) Comparison between WorldWater and Model based on the merged LDD-JICA DEM, (a-1) in 2017 and (a-2) in 2021; (B) Comparison between GISTDA and Model based on the merged LDD-JICA DEM, (b-1) in 2017 and (b-2) in 2021.

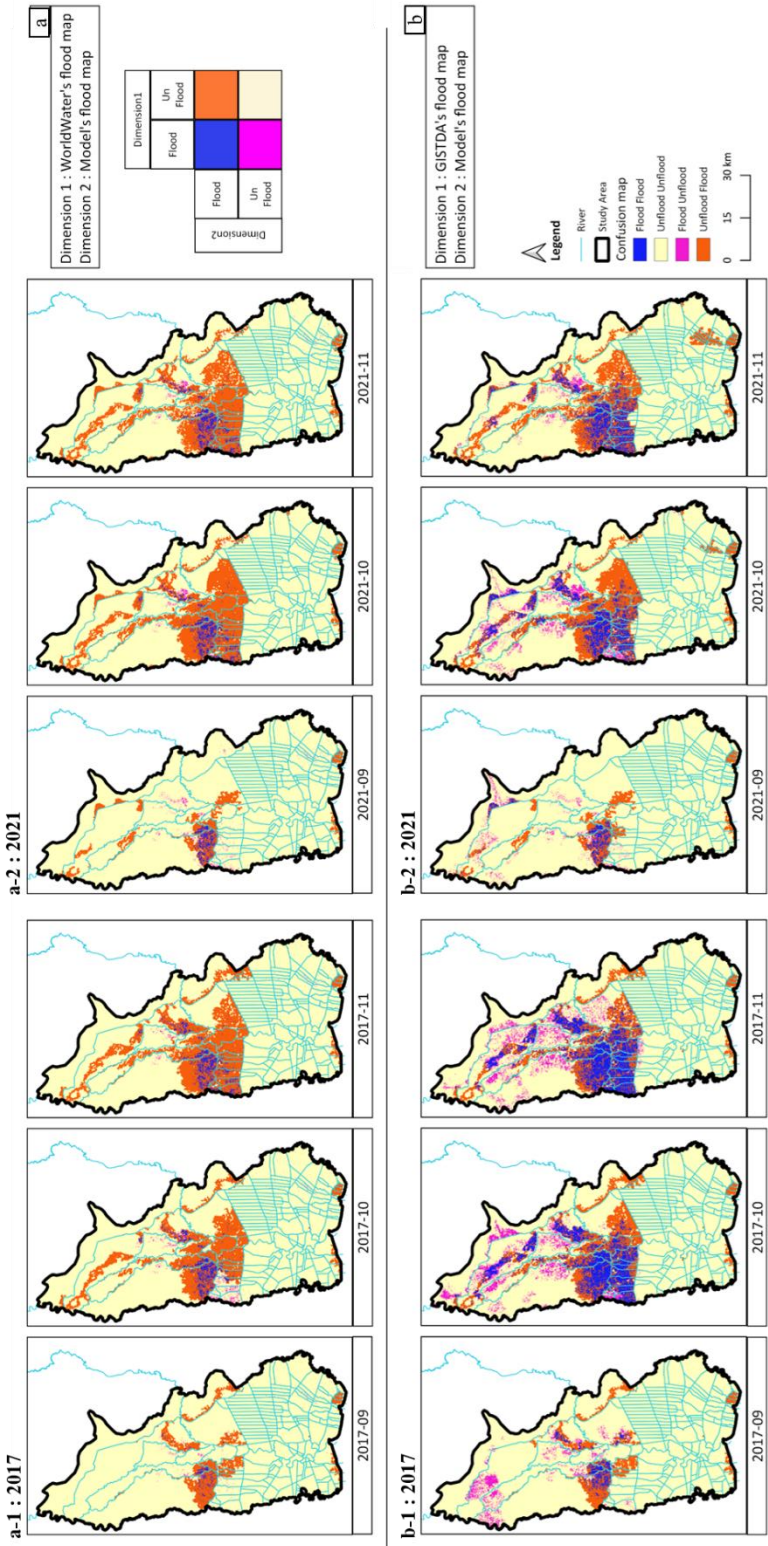


Figure 11: the contingency table of dimension 1 and dimension 2 flood maps on the spatial map: (a) Comparison between WorldWater and Model based on FABDEMv1-2 DEM, (a-1) in 2017 and (a-2) in 2021; (b) Comparison between GISTDA and Model based on FABDEMv1-2 DEM, (b-1) in 2017 and (b-2) in 2021.

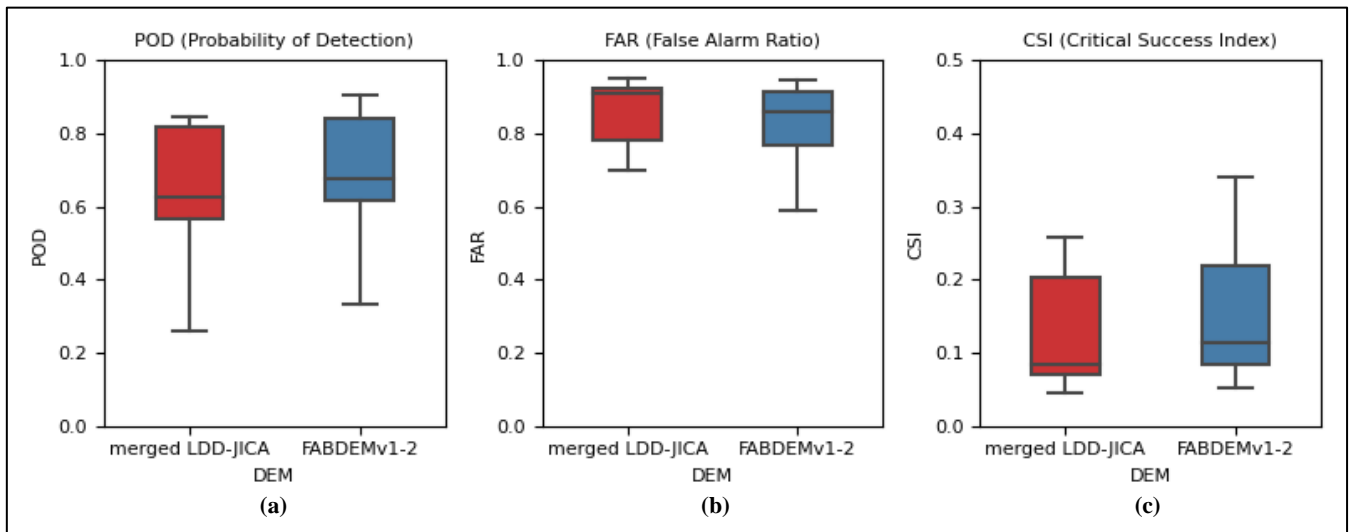


Figure 12: box plots illustrating the performance of the flood model based on the merged LDD-JICA and FABDEMv1-2 DEMs across three statistical metrics: (a) Probability of Detection (POD), (b) False Alarm Ratio (FAR), and (c) Critical Success Index (CSI).

625 6 Discussion

6.1 Overall result of DEM analysis workflow

The result of DEM analysis shows that ICESat-2 ATL08 data offer a unique advantage in verifying DEM accuracy (Carabajal & Boy, 2020). The overall precision of DEM products was evaluated using the ICESat-2 ATL08 benchmark, showing that JICA and FAMDEMv1-2 DEM were significantly better than the local and global DEM products in terms average of RMSE, with values of 1.87 m and 1.95 m, respectively (Figure 5Error! Reference source not found. and Table 6) in point comparison. The merged LDD-JICA DEM showed a slight difference of 0.06 m in average RMSE compared to the JICA DEM. This variance is primarily attributed to the combination of LDD and JICA DEMs, with JICA DEM being chosen as the primary DEM. However, it is noteworthy that the local DEM product exhibited a negative average bias (ME) ranging from -1.30 to -0.65 m, indicating that elevation of local DEM products is lower than the benchmark. Another study conducted in Spain, which verified Airborne LiDAR data with ICESat-2 ATL08, also reported a negative bias, with average ME values of -0.48 m (Zhu et al., 2022). On the other hand, the average ME of the global DEM products yielded positive values ranging from 0.25 to 4.77 m, indicating that the global DEM products overestimate the benchmark. This result has been previously confirmed in studies such as ASTERv3 (Weifeng et al., 2024), STRMv3 and TanDEM-X (Liu et al., 2020). The ASTERv3 DEM showed the lowest overall accuracy, with an average RMSE of 6.65 m. This is in line with other areas, such as the Qinghai-Tibet Plateau, where the RMSE reached 11.47 m (Weifeng et al., 2024). The TanDEM-EDEM is an updated version of TanDEM-X, which can reduce the error value from 3.65 to 2.78 m in terms of average RMSE.

In Figure 7, Figure A10 and Figure A11 illustrate the spatial grid comparison of 10 DEM products against the ICESat-2 ATL08 benchmark, with a resolution of 5x5 km for ME, MAE, and RMSE, respectively. The results clearly reveal that the global DEM tends to overestimate, particularly when compared to the ASTERv3 DEM. As shown in the figures, the error of the global DEM clearly clusters in urban areas, except for the FABDEMv1-2, which employs an algorithm to remove building discrepancies, as discussed in Sect. 3.3.8. Although, ICESat-2 ATL08 is capable of measuring land elevation very accurately, some urban areas still exhibit positive bias, particularly in high-rise dense areas (Liu et al., 2020), as shown in Figure 8. This suggests that the DEM analysis workflow can effectively utilize ICESat-2 ATL08 data for evaluation. In certain areas, the incorporation of ATL03 data may be necessary to enhance the evaluation process.

Despite the high spatial resolution of the local DEM (merged LDD-JICA DEM), which is derived from airborne LiDAR and expected to be highly accurate, the results of this study demonstrate that global DEM (FABDEMv1-2) can surpass it in specific ICESat-2-based evaluation metrics. Although this finding may appear counterintuitive, it is attributable to several underlying factors that influence DEM performance and consistency.

- ICESat-2 ATL08 benchmark: the ICESat-2 ATL08 dataset is derived from ICESat-2 ATL03 photon cloud data through a sequence of processing steps designed to extract accurate land surface elevation and canopy height information. The algorithm comprises several key stages, including noise filtering, surface detection, top-of-canopy identification, photon classification, photon label refinement, canopy height estimation, and link scale for data product as depicted (Neuenschwander & Pitts, 2019a). Notably, the algorithm is optimized to produce smoothed land surface elevation estimates over fixed segment lengths of 100 meters. This smoothing inherently aligns better with the spatial resolution of coarser global DEMs. Consequently, global DEMs—such as FABDEMv1-2—tend to yield terrain representations that are more consistent with the footprint-averaged elevations derived from ATL08.
- Quality of DEMs: the accuracy of local DEMs is highly dependent on data acquisition techniques, the acquisition time of the data, and post-processing workflows. Errors can arise from incomplete ground point classification, outdated surveys, or inconsistencies in vertical datum alignment. In the case of the merged LDD-JICA DEM, multiple datasets were combined—some of which were collected long ago (Sect 3.3)—potentially introducing temporal inconsistencies. Although the JICA DEM component is considered a reliable source, the merging process and age of the data may reduce overall accuracy. In contrast, the ICESat-2 ATL08 dataset benefits from continuous updates, offering more current elevation observations. This is particularly relevant in rapidly evolving urban areas such as Bangkok, where frequent land use changes can quickly render older DEMs obsolete.

6.2 Overall result of flood map analysis workflow

The flood classification processing aims to classify flood types from SWE map. This method is based on various assumptions and simplifications. The validity of the approach is hard to evaluate, given the lack of ground-truth flood extent observations. However, it is evident that in this study area, surface water extent is not only due to riverine flooding but also various other flooding mechanisms such as irrigation and pluvial flooding.

675 The Model flood map, based on both Model and FABDEMv1-2 DEMs, tends to overestimate flood extent compared to the
classified flood maps derived from SWE data provided by the GISTDA and WorldWater projects. Additionally, the flood map
based solely on FABDEMv1-2 performs slightly better than the one based on the merged LDD-JICA DEM, with an
improvement of approximately 13.55 – 25.56 percent according to the CSI. The overestimation of flood inundation from the
680 flood model occurs predominantly in the eastern part of the CPY River, indicating a clear need for improvement the 1D-2D
flood model. Although this study has incorporated high-quality DEM data implemented into the 1D-2D flood model, there are
still many factors affecting flood map generation. For instance, the 1D-2D flood model, developed long ago (Sect. 3.1), needs
to be updated and recalibrated due to continuous developments in water management plans, such as the Ayutthaya Bypass
channel (JICA, 2018) and ongoing land use changes in the lower CPY basin (Visessri & Ekkawatpanit, 2020), which impact
flood map simulations.

685 The results of the flood map comparisons demonstrate that the CSI value is relatively better when compared with GISTDA,
but lower when compared with WorldWater. It is observed that the overall WorldWater flood map shows relatively low
flooding compared to the GISTDA flood map. This is due to fundamental differences in the mapping approaches with
WorldWater aiming to provide long-time series of the typical distribution and persistence of monthly surface water presence
whereas GISTDA is targeting real time maps showing the extent of flooding at a specific moment in time. Additionally,
690 WorldWater uses only Sentinel-1 and Sentinel-2 data, whereas GISTDA combines data from multiple other satellites, as
described in Sect. 3.5.1. This can be further verified for accuracy with additional information from news sources and by cross-
referencing with ICESat-2 ATL13 data, extracted from ICESat-2 ATL03 (inland water surface heights), in main rivers ((Coppo
Frias et al., 2023) and (Dandabathula & Srinivasa Rao, 2020)). This suggests that the flood analysis workflow can effectively
verify the performance of flood simulation using satellite data. Although the flood simulation results in this study meet
695 acceptable standards and are sufficiently reliable for practical applications, the SWE data were generated using different
algorithms and satellite sources, resulting in variations in the datasets. These observed datasets were subsequently compared
to simulated flood maps derived from various DEM products.

6.3 Advantages and Limitations

This study proposes enhance the performance of 1D-2D flood models using satellite laser altimetry and multi-mission surface
700 water extent maps from EO data. The proposed workflows, encompassing comprehensive DEM analysis and flood map
analysis, are designed be adaptable, scalable, and standardized for the development of 1D-2D flood models. These workflows
enable their application across diverse spatial domains, ranging from local to national scales, and can be readily tailored to
address flood management challenges in other regions or countries.

Furthermore, the increasing availability of EO data has proven highly effective in improving the accuracy of 1D-2D flood
705 models, particularly in calibration and validation processes. ICESat-2 data, with its high precision and 91-day exact repeat
orbit, serves as a robust benchmark for evaluating DEM products. Its near-real-time capability is particularly beneficial in

areas undergoing rapid land-use changes. Freely available global DEM products, developed using advanced EO techniques, provide high-resolution and high-quality elevation data essential for accurate modeling.

710 Datasets such as WorldWater's SWE maps and GISTDA's flood maps are valuable resources for assessing the accuracy and reliability of simulated flood maps. By comparing observed flood extents with model outputs, these datasets help identify discrepancies, refine model parameters, and enhance the performance of flood models. This iterative process facilitates the development of more reliable and accurate tools for flood forecasting and management.

715 Despite these advantages, EO data is not without limitations. For instance, ICESat-2 offers an elevation accuracy of approximately 0.70 m ((Neuenschwander & Pitts, 2019b) and (Carabajal & Boy, 2020)). However, delays or gaps in EO data acquisition, as discussed in Sect. 3.4 and 3.5, can affect the evaluation of simulated flood maps. Furthermore, while the best available DEMs were selected for this study, elevation inaccuracies in certain areas may still compromise the precision of flood inundation maps. Periodic updates to the data, as explained in Sect.6.2, are necessary to address these limitations and maintain modeling accuracy.

6.4 Future applications

720 The workflows developed in this study represent a significant advancement in upgrading 1D-2D flood models by integrating satellite laser altimetry and multi-mission satellite surface water extent (SWE) maps. These workflows not only enhance the accuracy and reliability of flood inundation simulations but also offer scalable solutions for improving flood forecasting systems across multiple regions. Their success in the Chao Phraya River Basin sets the foundation for expanding these methodologies to other regions of Thailand, including the Eastern (Finn et al., 2018), Northeastern (Thanathanphon et al., 725 2014), Southern, and Western regions. These regions will benefit from improved model calibration, validation, and more accurate flood forecasts, thereby supporting better decision-making for flood mitigation, response, and water resource management.

Moreover, satellite technology offers new opportunities for measuring water surface elevation (WSE) such as ICESat-2 ATL13 (Jasinski et al., 2023), Surface Water and Ocean Topography (SWOT) (Biancamaria et al., 2016), CryoSat-2 ((Kittel et al., 730 2021);(Y. Shen et al., 2020)), Jason-2 and Envisat (Okeowo et al., 2017). These technologies enhance calibration, validation, error diagnosis, and monitoring of main river systems, especially in areas with limited ground-based instrumentation.

On a broader scale, the workflows could be adapted for use in other countries, particularly in regions facing similar challenges related to data scarcity, terrain complexity, and high flood risk. The integration of satellite EO data, combined with local hydrological models, could provide valuable insights for flood-prone regions across Southeast Asia and beyond, contributing 735 to global efforts in disaster risk reduction and climate resilience.

7 Conclusion

The present study enhanced the performance of 1D-2D flood models using satellite laser altimetry and multi-mission surface water extent maps from Earth Observation (EO) data. We demonstrated two workflows in the lower CPY basin.

- 740 ▪ **DEM analysis workflow:** This involved evaluating DEM accuracy using satellite laser altimetry data from ICESat-2 ATL08 before integrating the DEM products into the flood model. The assessment aimed to assess the overall performance of DEM products through vertical, spatial, track-wise analysis, and statistics measures to select the most suitable DEM for the study area. Furthermore, this workflow is transferable to other study areas, providing a method to reduce uncertainty before developing flood models. The results show that the merged LDD-JICA and FABDEMv1-2 DEMs are highly suitable in the study area, with RMSE values of 1.93 and 1.95 m., respectively.
- 745 ▪ **Flood map analysis workflow:** This workflow encompassed riverine flood classification and the evaluation of simulated flood maps generated by the 1D-2D flood model using multi-mission satellite SWE maps. While the flood classification algorithm still presents challenges, it is important to recognize that SWE maps derived from satellite EO cannot be directly compared with the output of flood models without further processing. The flood map evaluation method facilitated the assessment of flood simulation accuracy against satellite SWE maps, employing statistical and spatial analyses. These
750 evaluations contribute significantly to the calibration and validation of flood maps derived from the 1D-2D flood model. The results indicate that simulated flood maps based on FABDEMv1-2 DEM can improve the performance of the 1D-2D flood model by 13.55% to 25.56%, as determined by the CSI, when compared to simulated flood maps based on the merged LDD-JICA DEM.

Integrating these workflows will enhance the efficiency of the 1D-2D flood model and showcase the potential of utilizing EO
755 satellite data to enhance flood modelling capabilities for operational flood forecasting in Thailand and elsewhere.

Appendix A

Table A 1: Descriptive statistics of ten different DEM products.

DEM product	Statistical Parameters				
	Min (m.)	Max (m.)	Mean (m.)	Standard Deviation (m.)	Median
ICESat-2 ATL08	-7.00	218.42	5.29	6.81	2.49
LDD	-9.41	254.27	4.34	7.75	1.51
JICA	-22.97	239.31	4.20	5.48	1.95
merged LDD-JICA	-16.00	378.73	5.21	8.26	1.87
ASTER	-2.00	267.93	6.23	8.23	2.85
SRTM	-34.97	262.17	8.02	8.47	5.25
MERIT	-1.29	257.32	7.53	8.17	4.34
GIO30	-15.93	271.15	6.87	8.30	4.15
FABDEMv1-2	-14.99	267.93	6.22	8.23	2.85
TanDEM-X	-7.00	274.93	7.06	8.48	4.24
TanDEM-EDEM	-36.91	271.26	6.93	8.35	4.13

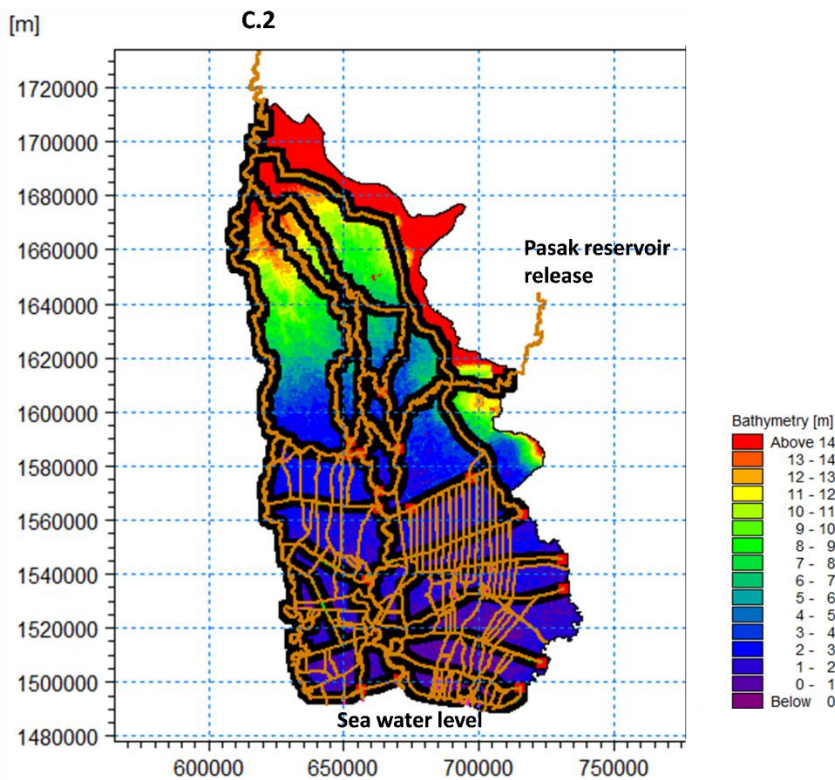


Figure A 1: 1D-2D Flood model

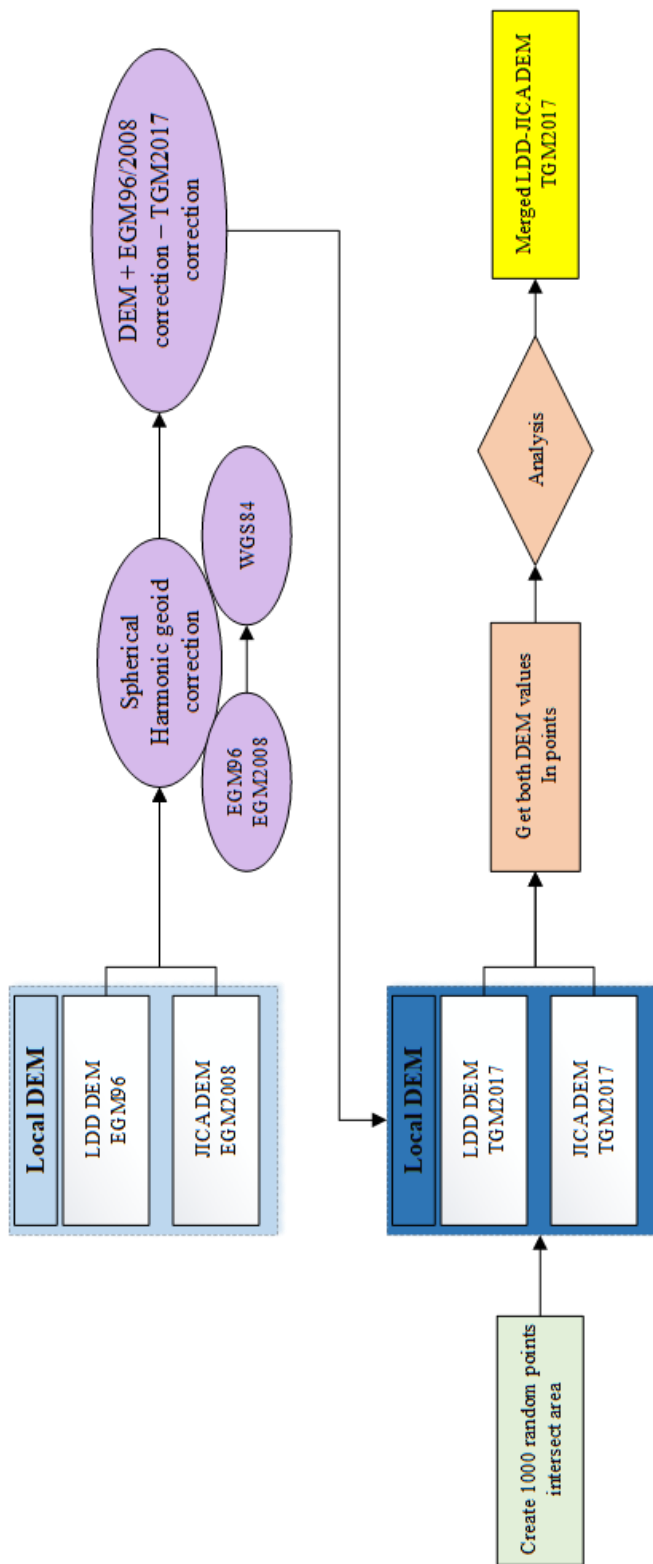


Figure A 2 : the merged LDD-JICA DEM's workflow

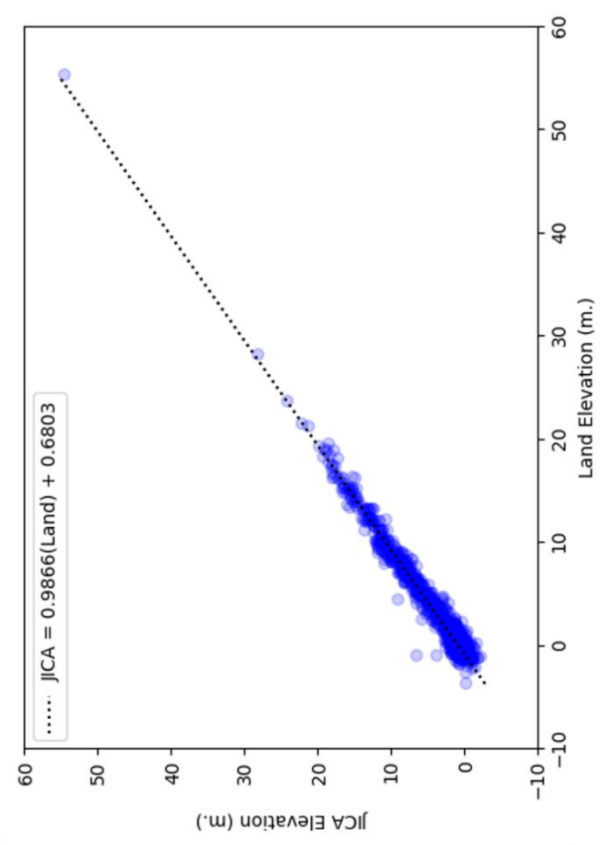
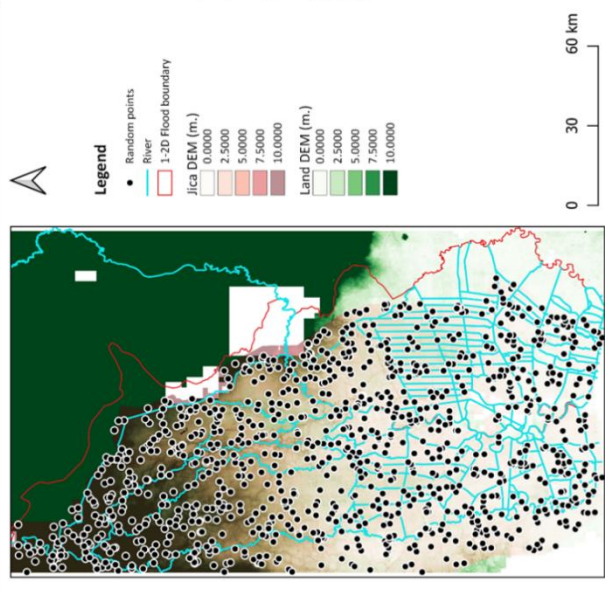


Figure A 4: the merged LDD-JICA DEM's processing

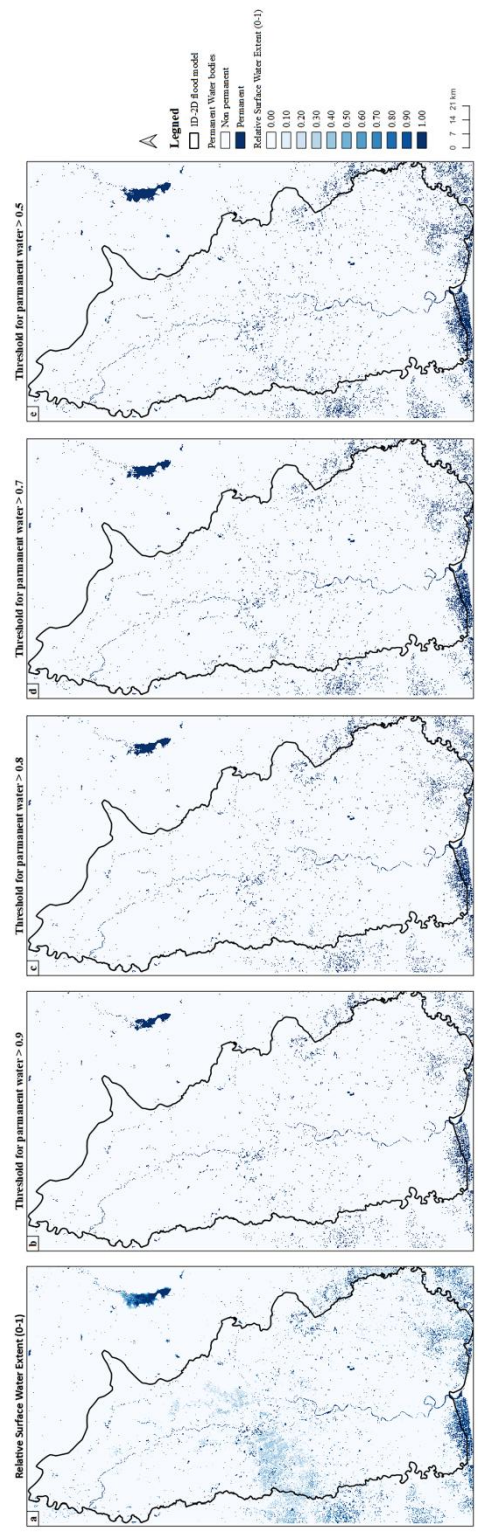


Figure A 3: Relative water frequency (a), threshold 0.9 (b), threshold 0.8 (c), threshold 0.7 (d) and threshold 0.5 (e).

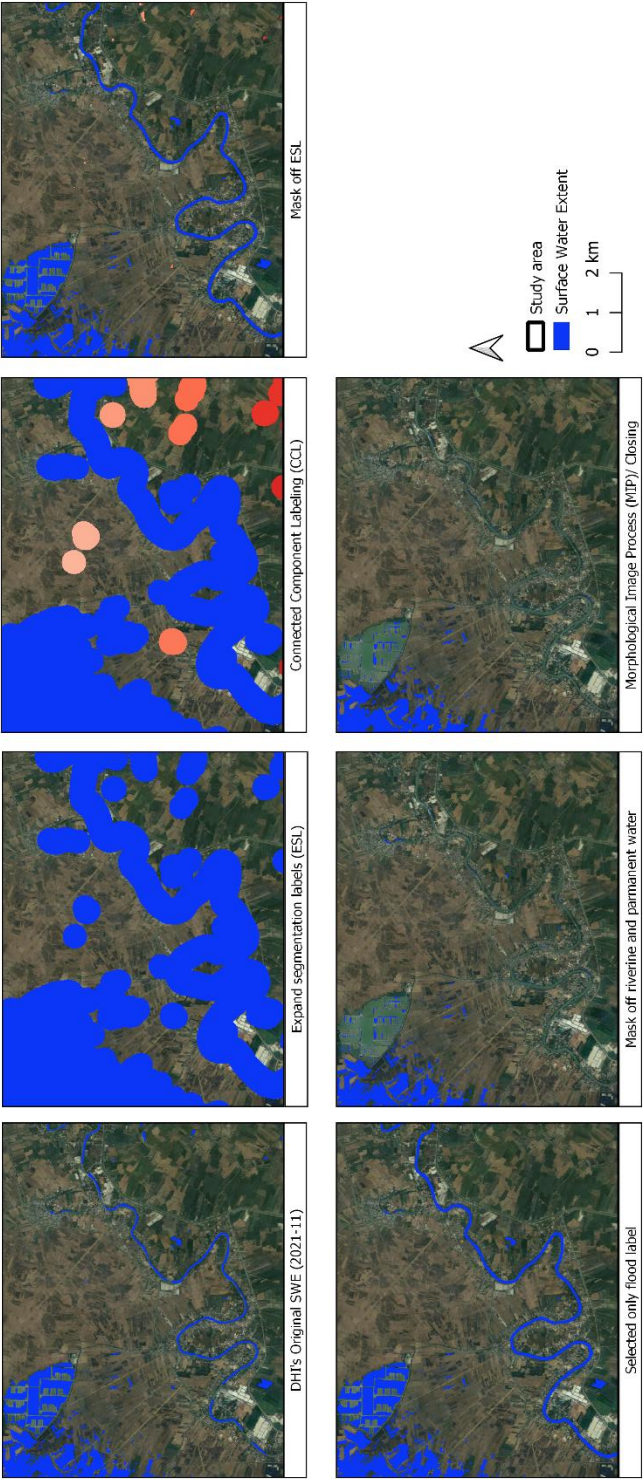


Figure A 6: Riverine flood classification processing. © Google Earth

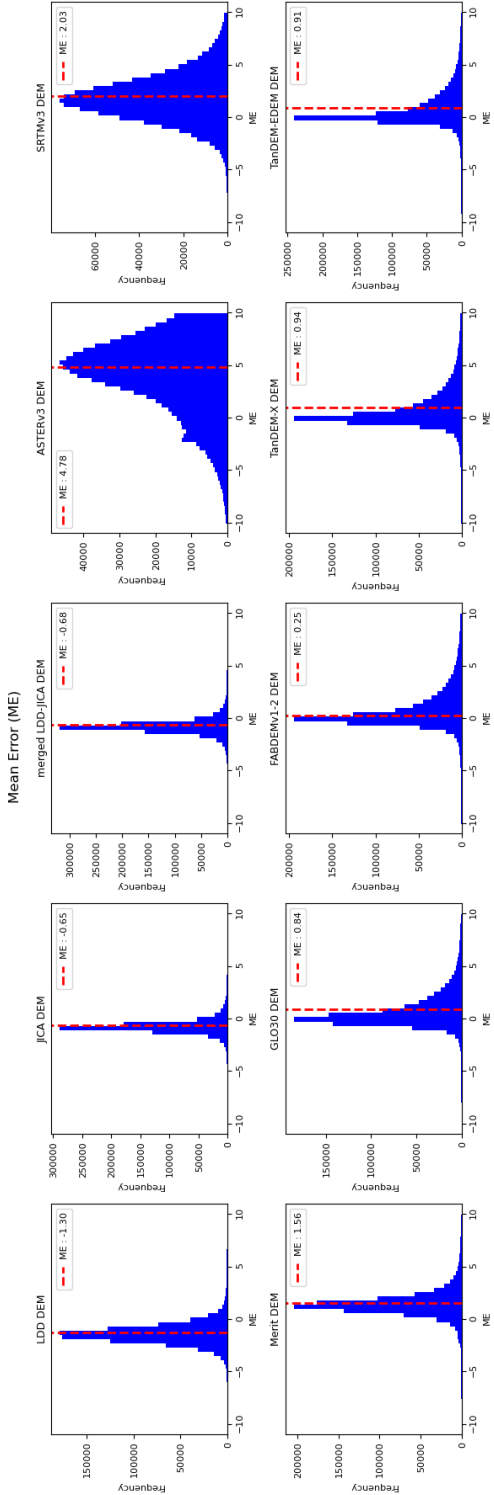


Figure A 5: The histogram distribution of the mean error (ME), comparing 10 DEM products against the ICESat-2 ATL08 benchmark.

Mean Absolute Error (MAE)

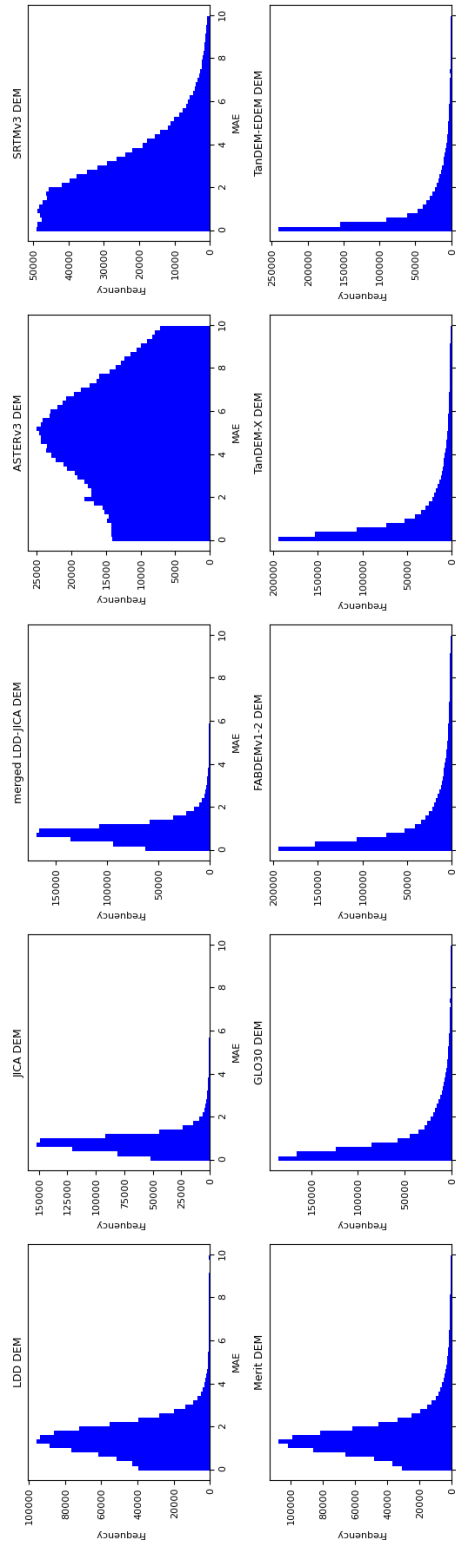


Figure A 8: The histogram distribution of the mean absolute error (MAE), comparing 10 DEM products against the ICESat-2 ATL08 benchmark.

Mean Squared Error (MSE)

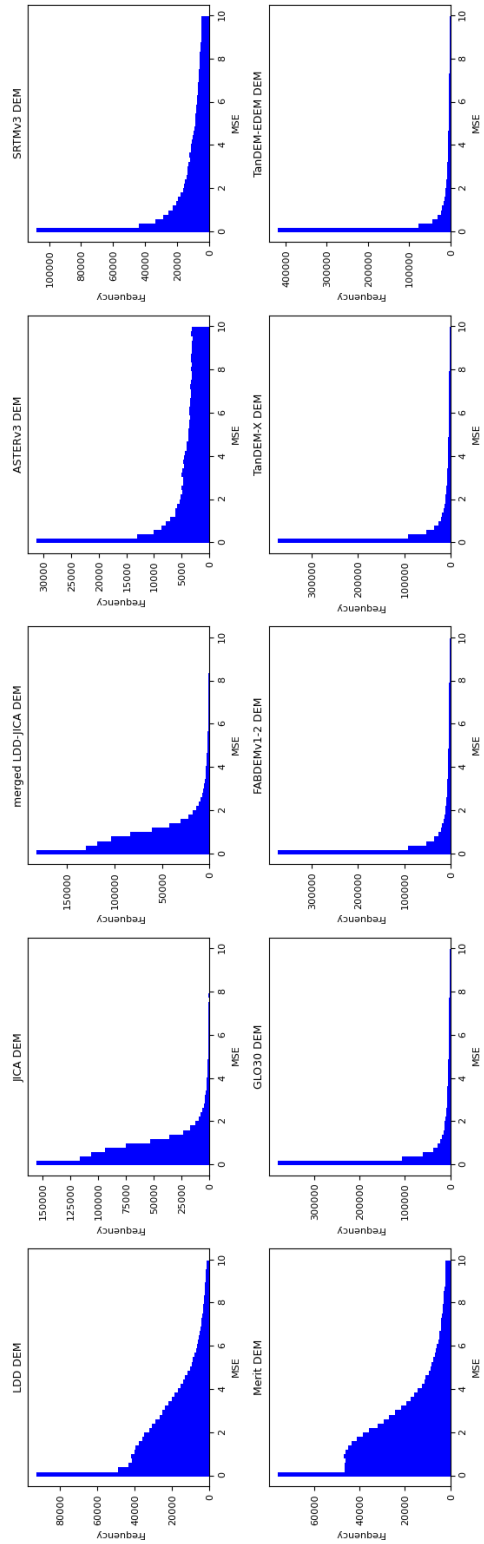


Figure A 7: The histogram distribution of the mean square error (MSE), comparing 10 DEM products against the ICESat-2 ATL08 benchmark.

Mean Error (ME) resolution 5x5 km.

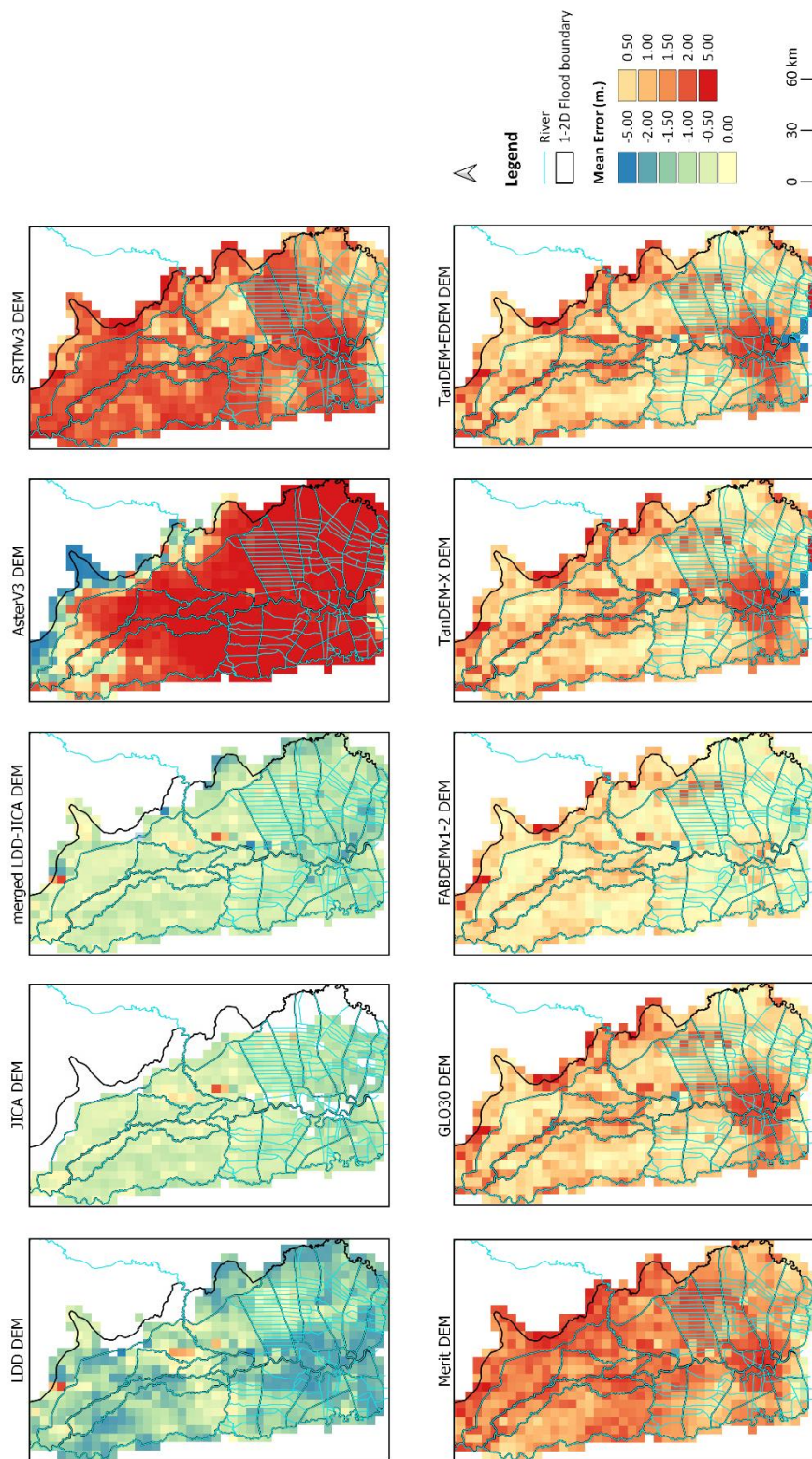


Figure A 9: The mean error (ME) spatial grid comparison of 10 DEM products against the ICESat-2 ATL08 benchmark, with a resolution of 5x5 km.

Mean Absolute Error (MAE) resolution 5x5 km.

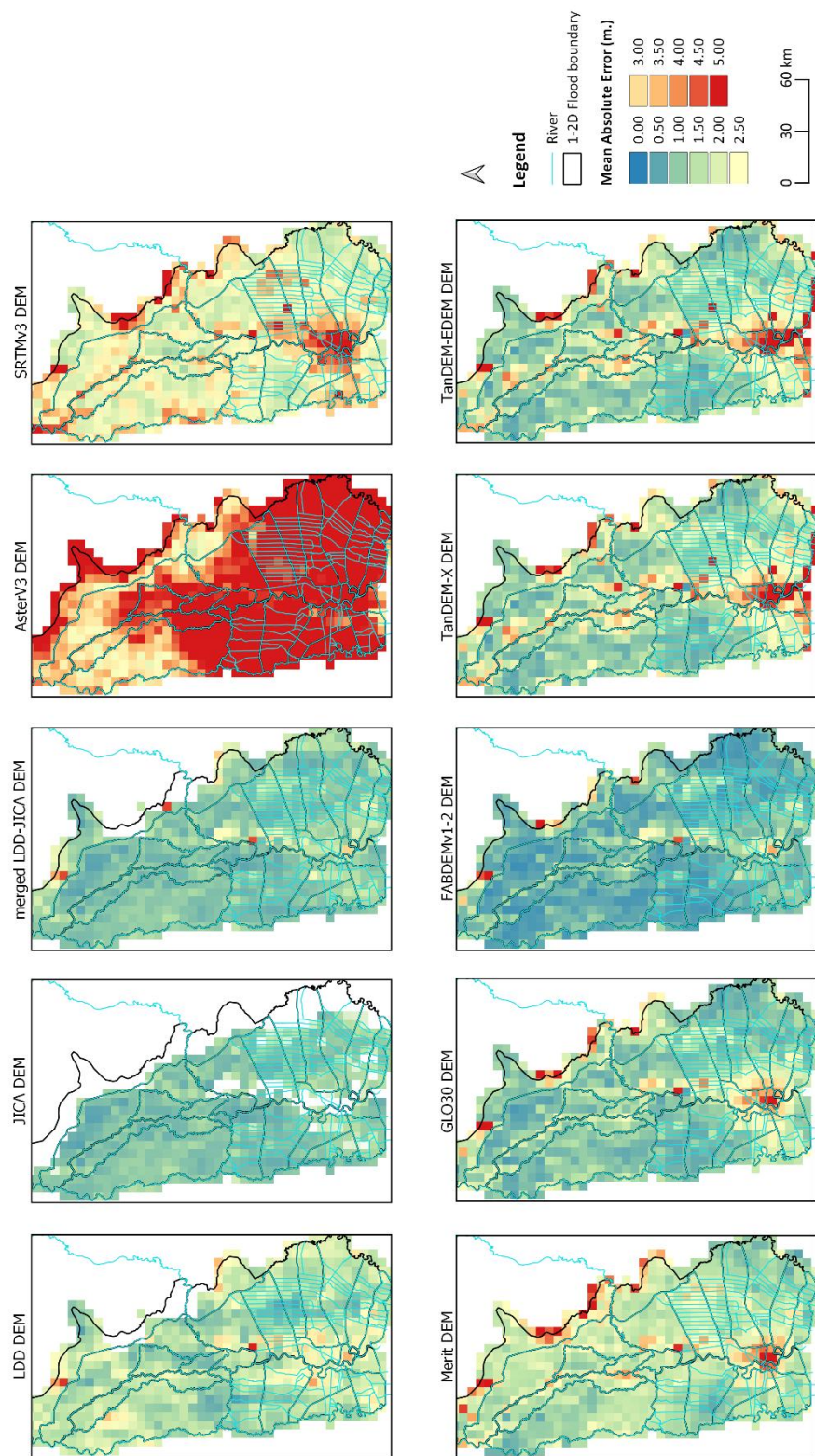


Figure A 10: The mean error (MAE) spatial grid comparison of 10 DEM products against the ICESat-2 ATL08 benchmark, with a resolution of 5x5 km.

Root Mean Square Error (RMSE) resolution 5x5 km.

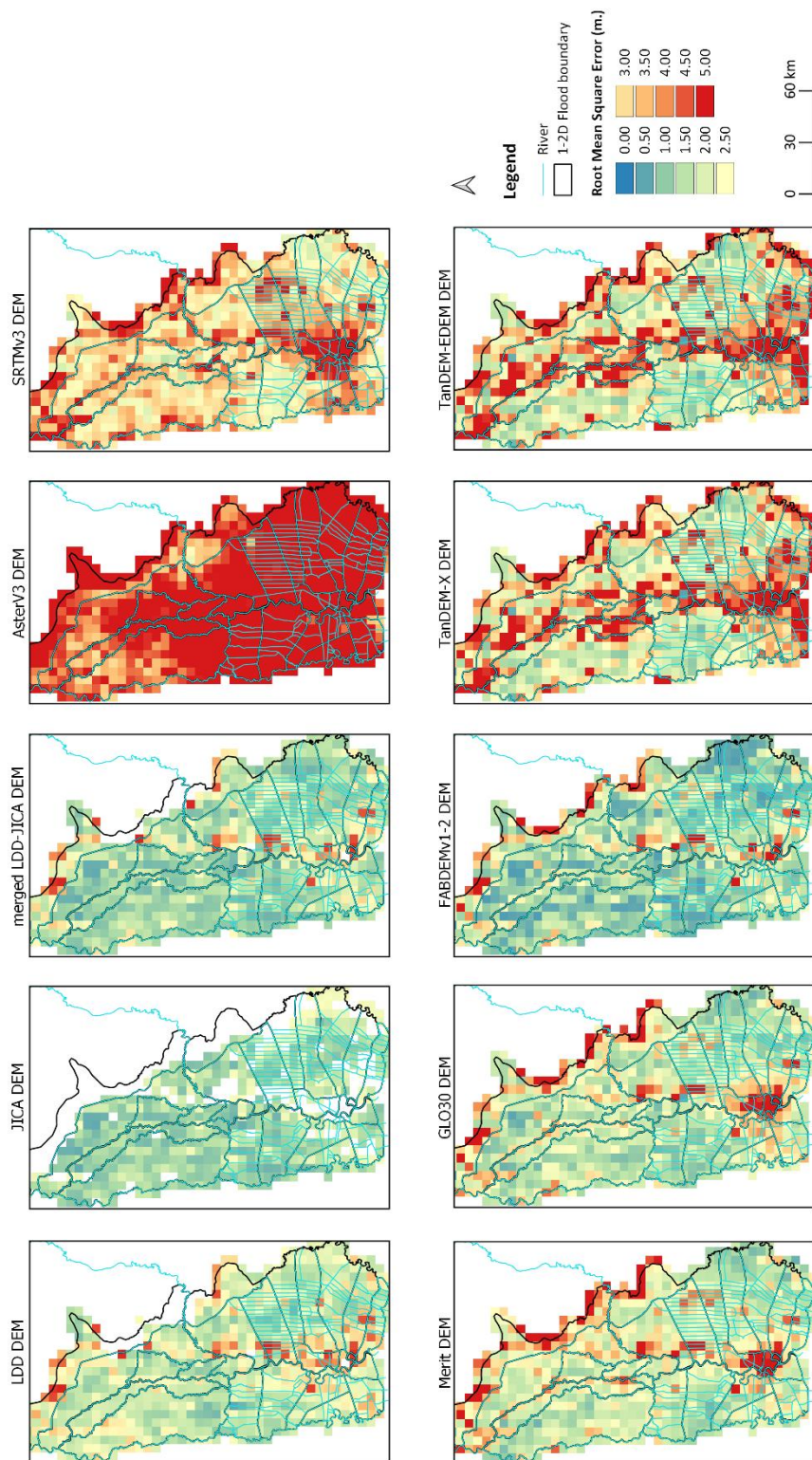


Figure A 11: The mean error (RMSE) spatial grid comparison of 10 DEM products against the ICESat-2 ATL08 benchmark, with a resolution of 5x5 km.

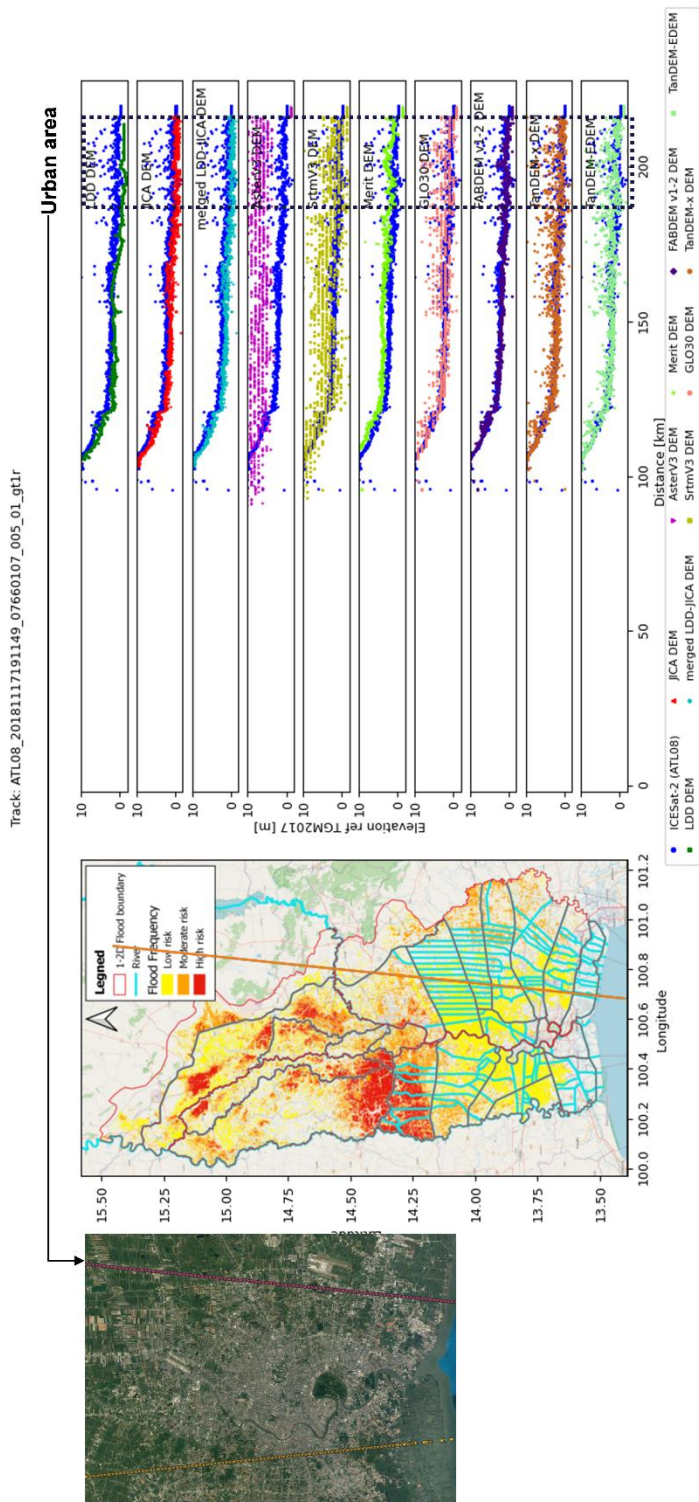


Figure A 12: The track-wise comparison of 10 DEM products with ICESat-2 ATL08 benchmark. © OpenStreetMap contributors 2015. Distributed under the Open Data Commons Open Database License (ODbL) v1.0. © Google Earth.

Author contributions

Theerapol Charoensuk: Conceptualization, Methodology, Validation, Formal analysis, Investigation, Data curation, Visualization, Writing – review & editing, Writing – original draft. Claudia Katrine Corvenius Lorentzen: Coding and Visualization. Anne Beukel Bak: Coding and Visualization. Jakob Luchner: Software, review. Christian Tøttrup: Data
775 curation, review. Peter Bauer-Gottwein: Methodology, Conceptualization, Resources, Supervision, Writing – review & editing.

Competing interests

The authors declare that they have no conflict of interest.

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