1 Does peatland rewetting mitigate flooding from extreme

2 rainfall events?

3 Shirin Karimi[a], Virginia Mosquera[a*], Eliza Maher Hasselquist[a], Järvi Järveoja[a] and Hjalmar Laudon[a]

4 [a] Swedish University of Agricultural Sciences, Department of Forest Ecology and Management, Umeå,

5 Sweden

6 *Corresponding author: virginia.mosquera@slu.se

7 Abstract

8 Pristine peatlands are believed to play an important role in regulating hydrological extremes because

9 they can act as reservoirs for rainwater and release it gradually during dry periods. Rewetting of drained

10 peatlands has therefore been considered an important strategy to reduce the catastrophic effects of

11 flooding. With the anticipation of more frequent extreme rainfall events in the future due to a changing

12 global climate, the importance of peatland rewetting in flood mitigation becomes even more important.

13 To date, however, empirical data showing that rewetting of drained peatlands actually restores their

14 hydrological function similar to pristine peatlands is largely lacking, particularly for boreal fens. To

15 assess whether peatland rewetting can mitigate flooding from extreme rainfall events and ensure water

16 security in a future climate, we measured event-based runoff responses before and after rewetting

17 using a BACI approach (before-after and control-impact) within a replicated, catchment scale study at

18 the Trollberget Experimental Area in northern Sweden. High-resolution hydrological field observations,

19 including groundwater level (GWL), discharge, and rainfall data were collected over four years, allowing

20 us to detect and analyze 17 rainfall-runoff events before and 30 events after rewetting. We found that

21 the rewetted sites experienced an increase in the GWL following rewetting and that this was

22 consistently observed across all distances from the blocked ditch within the peatland. Our rainfall-

23 runoff analysis revealed that rewetting significantly decreased peak flow, runoff coefficient, and

24 reduced the overall flashiness of hydrographs, making the rewetted sites function more like the pristine

25 control peatland. However, "lag time" which was already similar to pristine conditions was pushed

26 farther away from pristine conditions following rewetting. Yet, our results showed that the

27 effectiveness of ditch-blocking in flood moderation was strongly influenced by the initial condition and

28 catchment percent of restoration, as one of our two rewetted peatlands did not show significant

29   change attributed to it being already similar to the pristine site, suggesting less treatment effect; and

30   the other catchment, with higher restoration percentage, had a better response to treatment. In

31   summary, our findings suggest that peatland rewetting has the potential to mitigate flood responses,

32   however, further research over a longer time period is needed as peat properties and the peatland

33   vegetation will develop and change over time.

34   Keywords

35   Boreal landscape, peatland hydrology, rewetting, flood mitigation, extreme events.

36

Introduction

Pristine peatlands are the predominant wetland type in the boreal biome. They encompass 15% of the boreal land surface area and serve as significant carbon sinks and methane sources, playing a crucial role in regulating the global climate (Helbig et al., 2020). In recent years, there has been an increased recognition of the importance of peatlands in carbon capture, flood management, water quality, and biodiversity (Holden et al., 2017). Regrettably, these valuable ecosystems have undergone substantial human-induced damages, with more than half of the total pristine peatlands in Europe estimated to have been impacted by drainage for agriculture, forestry, or peat extraction (Andersen et al., 2017). Drained peatlands cannot sustain critical ecosystem services, such as climate regulation through long-term carbon sequestration or buffering extreme hydrological events, imposing a significant cost on society—a burden that could be alleviated through appropriate rewetting measures (Loisel & Gallego-Sala, 2022). Additionally, there are growing concerns surrounding climate change projections for the Northern Hemisphere, indicating an expected increase in more frequent extreme precipitation events, along with extended dry periods (Aghakouchak et al., 2020; Hawcroft et al., 2018).

Pristine peatlands function as significant water reservoirs, efficiently storing substantial amounts of water during periods of high rainfall (Acreman & Holden, 2013). As extreme rainfall events are anticipated to become more frequent in the evolving global climate, understanding the role of peatland rewetting in flood mitigation is increasingly vital. Rewetting projects typically involve physical interventions such as ditch-blocking or infilling, aiming to increase groundwater level (GWL). Moreover, the blocking of ditches cuts off preferential pathways along open drains, and when combined with pooling behind dams, has the potential to act as a buffer during peak flow events, slowing water release and mitigating the flashiness of the discharge response (Holden, 2006; Holden & Burt, 2003). Therefore, by reducing peak flows, peatland rewetting can also contribute to natural flood management by attenuating downstream flow and diminishing flood risk. Furthermore, the reduction of peak flows could play an important role in mitigating further erosion of peatlands and minimizing sediment production, as well as carbon loss (Shuttleworth et al., 2015).

In Sweden, peatlands cover approximately 65,600 km^2 (16% of the Swedish land area) and are predominantly located within boreal regions (Franzen et al., 2012; Montanarella et al., 2006). The historical practice of draining peatlands began in the early 18th century for agricultural purposes and later in the 19th century for forestry, resulting in the excavation of over 1 million km of ditches, primarily dug by hand (Laudon et al., 2022). Consequently, the rewetting of degraded peatlands in Sweden has become a pressing priority to enhance the hydrological functioning of these ecosystems (Bring et al., 2022). As a response, several national programs for peatland rewetting have emerged, with a primary

70   emphasis on reintroducing essential ecosystem services, notably flood control. In a significant move, in

71   2018, 27 million euros was allocated to facilitate peatland rewetting in Sweden. However, the empirical

72   underpinning supporting the desired outcomes of peatland rewetting is still largely lacking.

73   The effect of peatland rewetting on hydrological responses during rainfall events has received scientific

74   attention over the past decades (Gatis et al., 2023; Goudarzi et al., 2021; Ketcheson & Price, 2011;

75   Menberu et al., 2018; Shuttleworth et al., 2019). Event-based analysis of stream hydrographs by

76   employing various metrics related to hydrograph magnitude and timing, is a common approach for

77   investigating dominant runoff generation processes in catchments and understanding how quickly

78   water is mobilized from the landscape (Haque et al., 2022; Ketcheson & Price, 2011; Kirchner et al.,

79   2023). These response metrics provide valuable insights into catchment storage and release

80   mechanisms (Blume et al., 2007). One widely acknowledged aspect is the impact of rewetting on the

81   event runoff coefficient, which represents the ratio of event runoff depth to event rainfall depth (Evans

82   et al., 1999; Shuttleworth et al., 2019). Therefore, comparing event characteristics before and after

83   rewetting offers a means to understand hydrological processes and runoff generation mechanisms at

84   the catchment scale, thereby improving our understanding of flood estimation during extreme events.

85   A common limitation in the current literature is the predominant focus on event characteristics in

86   natural or relatively un-impacted catchments, with few studies addressing rewetted peatlands.

87   Therefore, the extent of hydrological changes due to rewetting is not well understood. Some studies

88   highlight the positive impact of peatland rewetting on flood moderation with a reduction of peak storm

89   flow (Gatis et al., 2023; Javaheri & Babbar-Sebens, 2014; Lane et al., 2003; Shuttleworth et al., 2019;

90   Wilson et al., 2011), however there are differences in the extent of the flood moderation. For example,

91   Gatis et al. (2023) reported a 49% reduction in peak storm flow after rewetting, while Shuttleworth et

92   al. (2019) found a 24% reduction in peak storm flows and a 94% extension in lag times without a change

93   in runoff coefficients. The challenges in understanding the effects of rewetting at the catchment scale

94   are further underscored by the inherent high spatial variability of peatland hydrology and physical

95   characteristics (Evans et al., 1999). The apparent discrepancies in study outcomes, coupled with

96   significant variations among different research sites, highlight the importance of addressing this

97   through further in-depth investigations for developing effective strategies in peatland management,

98   especially given the evolving trend in climate.

99   Moreover, a recent meta-analysis conducted by Bring et al. (2020) has brought further attention to a

100   noteworthy knowledge gap in understanding the impact of rewetting on GWL changes at different

101   distances from the intervention (i.e. blocked ditches). While existing studies have contributed valuable

102   data on the overall hydrological effects of peatland rewetting, a comprehensive spatial analysis of

groundwater changes following rewetting remains inadequately explored. Despite this shortage, some studies suggest that the impact of rewetting, especially through ditch blocking, is localized, resulting in more pronounced GWL rise in close proximity to the ditch (Armstrong et al., 2010; D'Acunha et al., 2018; Haapalehto et al., 2014; Wilson et al., 2010). In a prior study (Karimi et al., 2024) in the same catchments as reported on below the overall effect of rewetting on hydrological functioning was found, including a significant rise in GWL post-rewetting. However, a thorough examination of groundwater changes at varying distances from the ditch, considering its important role in discharge regulation, is essential to enhance our mechanistic understanding of flow generation after rewetting. Without such monitoring, the estimation and extrapolation of discharge responses across the landscape become more uncertain. Therefore, a more detailed spatial analysis of GWL changes is crucial for those involved in managing these peatlands.

Given that there have been inconsistent reports in the literature on the extent to which rewetted peatlands will affect hydrological functioning, particularly with regards to natural flood management, we build on methods used to examine the effect of pristine peatlands on flood attenuation (Karimi et al., 2024) to that of rewetting's impact on hydrological functioning. We used a hydro-climate data set comprised of one-year pre- and three years post-rewetting and incorporate two control catchments to ensure the robustness of our findings. The primary objective of this paper was to test whether peatland rewetting has any natural flood management effect. We hypothesized that rewetting leads to a reduction in peak flow, runoff coefficient, Hydrograph Shape Index (HSI), and an increase in lag time, resulting in a generally less flashy hydrograph. Moreover, as GWL is an important indicator of the amount of water stored in the peatland and the effect of the rewetting, we asked how far from the ditch GWL was increased by the ditch blocking. We hypothesized that the areas closest to the ditch would increase more than the areas further away from the blocked ditch.

Materials and methods

Study sites

This study took place in the Trollberget Experimental Area (TEA), situated approximately 50 km northwest of Umeå (TEA; 64.181550N, 19.835378E) (Figure 1). The TEA's peatland is an oligotrophic minerogenic fen. Prior to rewetting the peatland was dominated by *Sphagnum* spp., complemented by sparse sedges, dwarf shrubs, and sparse tree canopy (basal area = 2.6 m^2 ha^{-1}) of slow-growing Scots pine (*Pinus sylvestris)*. The underlying soils consist mainly of humic podzol, with some drier areas featuring Humu-ferric podzol and wetter regions comprising Histosols. Peat depth is on average 2.41 m (Laudon et al., 2023). The bulk density of the drained peatland varied between 0.05 and 0.13 g cm^{-3} within the top 55 cm of the peat profile. The bulk density generally increased with distance from the central ditch and with peat depth (Casselgård, 2020). The climate of the area is classified as cold temperate humid, characterized by a mean average temperature of 2.4°C and annual precipitation of 623 mm (approximately 30% as snow), based on data collected from 1980 to 2020 at the nearby Svartberget Climate Station (Laudon et al., 2021).

The peatland at TEA was drained by manual ditch-digging in the early 1920s primarily for forestry purposes, however, because of nutrient limitation the peatland remained unproductive with sparse tree cover (Laudon et al., 2023). The peatland is divided into two catchments draining in two directions, referred to here as R1 and R2, with drainage areas of 33 and 60 ha, respectively (Figure 1). Both catchments are similar in topography and vegetation, however, R1 had an open-water pond shown on historic maps pre-drainage. In the 1930s, the uplands of the peatland were also drained leading to 1152 m of ditches in R1 and 5189 m of ditches in R2 (Laudon et al., 2021). In November 2020, the peatland was rewetted by filling and blocking all the ditches in the peatland, whereas ditches in the surrounding none-peat areas were left unmanaged. As a result of these efforts, 59% of the ditches of R1 and 16% of the ditches of R2 were blocked. The ditches were filled using peat from the site with additional dams built at regular intervals using the tree logs harvested from the site. The logs were placed horizontally, except at the two outlet locations where the logs were inserted vertically into the peat and layered additionally with geotextile. To protect the soil characteristics as much as possible, the heavy machinery (i.e., excavators) used moveable log mats while moving on the mire. Additionally, the sparse tree cover that grew on the peatland was cut to reduce evapotranspiration and complement the ditch blocking (Karimi et al., 2024). Finally, after restoration, the open-water pond at R1 re-appeared ca 100 m upstream of the sampling location (Figure 1b).

The Degerö Stormyr

This study leveraged available data from a nearby natural fen, Degerö Stormyr (273-ha catchment), located approximately 24 km from the TEA, at the Kulbäcksliden Research Infrastructure (KRI, Noumonvi et al., 2023) (64.182029N, 19.556543E) to serve as the control for the rewetted peatland R1 and R2 catchments. Degerö Stormyr is characterized as an acidic, oligotrophic, minerogenic, mixed mire system. This intensively studied peatland complex exhibits varying vegetation compositions, predominantly featuring *Sphagnum* moss and sedges. The depth of the peat has an average thickness of ~3 m (Noumonvi et al., 2023). The bulk density of the peatland varied between 0.02 to 0.06 g/cm3 within the top 34 cm of the peat profile (Fig. 2 in Casselgård, 2020).

C4 (Kallkälsmyren)

The second control catchment, C4 (Kallkälsmyren), situated within the Krycklan Catchment Study (KCS) (64.260722N, 19.770339E). C4 is a nutrient-poor, minerogenic fen located approximately 10 km from the rewetted catchment. It encompasses a catchment area of 18 ha, with 40% covered by peatlands and the remainder by forest (Laudon et al., 2021). Similar to TEA, the climate is characterized as a cold temperate humid type with persistent snow cover during the winter season. The peat vegetation cover is dominated by *Sphagnum* spp.

Data collection

At the TEA, GWL were measured between 2019 and 2023 at an hourly resolution using 30 dipwells. Half of these dipwells were continuously monitored for GWL using data loggers (Levelogger 5, Solinst, Canada), while the remaining were manually measured every two weeks during the snow-free season. Dipwells were distributed along five transects. Each transect consisted of six wells with increasing distances of approximately 10, 50 and 100 m from the main ditch (Figure 1). For the Degerö Stormyr control site, GWL data for the corresponding period were obtained from four wells that are part of the ICOS database (www.icos-sweden.se/data). Due to technical issues with the groundwater loggers, no groundwater data for recent years was available for the C4 control catchment.

The discharge data at two TEA mire outlets was collected between 2019 and 2023 at an hourly resolution using 90 degree sharp-crested V-notches with connected data loggers for continuous water level measurements (TruTrack WT-HR, Intech Instruments, Australia). Automatic observations were not possible year-round as there was no heating in place, which limited data collection during the winter low flow periods. Manual water level measurements were taken twice a month to calibrate automatic water level data, and stage-discharge relationships were defined using manual flow gauging. Specific

discharge (discharge per unit catchment area) was calculated using catchment areas derived from the Deterministic 8 (D8) algorithm based on a 2×2 m resolution DEM in which we first burned the ditches into the digital elevation model (DEM) to the depth of 0.5 m (Whitebox GAT 3.3) (Laudon et al., 2021). For this study, we only utilized discharge data from the C4 control site due to its proximity to the rewetted site. At C4, the outlet is equipped with a similar V-notch weir situated within a heated dam house, facilitating continuous stage height monitoring year-round. Discharge measurements and calibrations followed the same protocol and interval as those implemented at TEA (Laudon et al., 2021).

Rainfall data were acquired from a reference climate station at Svartberget Research Station (64.244376N, 19.766378E, 225 m.a.s.l) (Laudon et al., 2021). Rainfall measurements were logged every 10 minutes using a tipping-bucket rain gauge (ARG 100, EML, UK). The climate station is integral to the reference climate monitoring program at Vindeln experimental forests, adhering to the WMO standard for meteorological measurements (Karlsen et al., 2019).

GWL analysis

First, the hourly groundwater data were examined for outliers, and any gaps were filled using the Generalized Extreme Studentized Deviate (ESD) filter (Rosner, 1983). The algorithm processes a time-series dataset by calculating a rolling mean and standard deviation with a window size of 6 hours. Outliers were identified by comparing each data point to the moving average, and values exceeding the 3-standard deviation threshold were identified as outliers and subsequently removed from the dataset. Subsequently, the data were gap-filled using the Spline interpolation method, an advanced form of interpolation that utilizes piecewise polynomial functions to estimate data between two known points. The data were aggregated to daily time scales. For our analysis we used the GWL data from 1st of June to the end of October as our study focused on rainfall events; outside this period, precipitation often occurs as snow and dipwells could be frozen. For each catchment R1 and R2, the GWL data were averaged, and pairwise comparisons test were conducted to assess if there were any significant differences between pre-rewetting and multiple post-rewetting years. As the data were not normally distributed and we were interested in the distribution of the data and not the means, the non-parametric Wilcoxon tests were used. Then, a Bonferroni-Holm correction was applied to adjust for multiple comparisons. The differences were considered significant when $p < 0.05$. Moreover, to examine the impact of rewetting on GWL at all distances from the main ditch, data were disaggregated based on distances of 10, 50, and 100 m to the main ditch. It is noteworthy that additional side ditches existed in the proximity of some of the dipwells, indicating a potential limitation in the study design.

Rainfall-runoff events detection

As a first step, we segmented the 2020–2023 summer–autumn precipitation record into distinct rainfall events using the inter-event time definition (IETD) via the IETD R package (Duque, 2022). The IETD establishes a minimum dry period between independent rainfall events as a criterion for grouping them. To distinguish independent rainfall events from continuous precipitation, we set a minimum threshold of 0.1 mm h^{-1} at the start of an event. Events were considered distinct if they were separated by at least 12 hours without rainfall. The methodology for identifying runoff events was based on the framework outlined by Luscombe (2014) and was further adapted to the specific characteristics of our study area. Runoff events were defined as periods during which the observed discharge exhibited significant deviations from the baseflow. Rainfall events were matched with the runoff events that followed within a specified time window (12.5 h). We calculated rolling quantiles for this time window (12.5 h) at the 30th and 95th percentile (Q30th and Q95th respectively). A rolling quantile for the 70th percentile for a one month period was also calculated (MQ90). Where (Q95th – Q30th) > MQ90, the flow was considered to be elevated and any fluctuation in flow was driven by precipitation; therefore measured discharge was used (Gatis et al., 2023; Puttock et al., 2021). A final, visual inspection of the time series with detected events was used to quality control these data and ensure that all significant rainfall and flow events were extracted from the dataset.

Flood mitigation effects

To evaluate the flood mitigation effect of peatland rewetting and determine its impact, we employed a set of response metrics to characterize hydrologic responses during events following the rewetting process. These response metrics included event duration, rainfall volume, peak flow, runoff coefficient, lag time, and HSI. We calculated these response metrics for both the rewetted and control sites. The selection of these response metrics was based on their widespread use in hydrological comparison studies (Edokpa et al., 2022; Wilson et al., 2011). Peak flow response was computed as the maximum discharge observed during each event. Runoff coefficient was determined as the ratio of total event runoff to total event rainfall. Lag time calculated as the time between peak rainfall and peak discharge in each event. HSI, defined as the ratio of peak storm discharge to total storm discharge, was used to provide a straightforward measure of the overall hydrograph shape (Shuttleworth et al., 2019). The response metrics for the rewetted catchment R2 and the control site were derived using the start and end times of rainfall-runoff events identified at R1 catchment.

Statistical analyses

The statistical design used in this study focuses on the BACI approach (before-after and control-impact) as used previously in hydrological studies (Holden et al., 2017; Laudon et al., 2023; Menberu et al., 2018; Shuttleworth et al., 2019). We standardized the response metrics derived from the two catchments (R1 and R2) of the rewetted site against the control catchment (treatment minus control) to distinguish responses resulting from rewetting treatment from natural variation, changes over time and seasons. Due to variations in the frequency of events between the pre- and post-rewetting periods, and the non-normal distribution of response metrics, a non-parametric test was employed. Specifically, the Wilcoxon test was conducted to investigate statistically significant changes in the distribution of data for each catchment (R1 and R2) of the rewetted site before and after rewetting, with a focus on understanding the extremes, rather than solely examining means (Shuttleworth et al., 2019). Significance was determined at $p < 0.05$. Additionally, we aggregated all years post-rewetting together due to the highly variable number of events occurring during each year post-rewetting. Statistical analysis was undertaken in R version 4.1.2. (R Core Team, 2021) with data processing, summary statistics and plotting undertaken using the R package Tidyverse (Wickham, 2016).

266

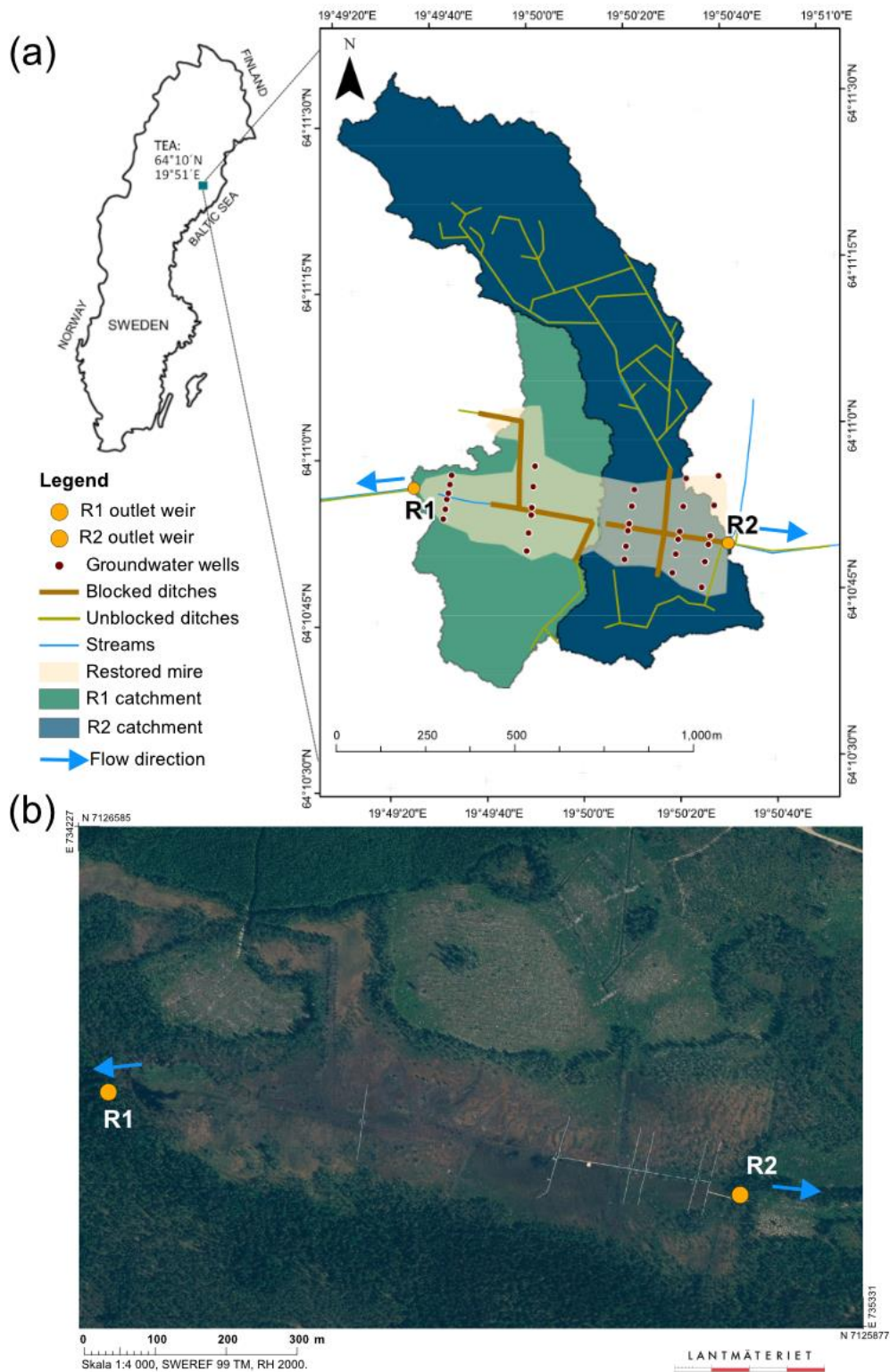



Figure 1. Trollberget Experimental Area (TEA) catchments with monitoring locations (A). Yellow circles show the locations of the outlets of the catchment areas for R1 and R2 (weir locations) of the rewetted peatland. Red small circles designate groundwater dipwells. Note that in R1, 59% of the ditches were blocked, while in R2, only 16% of the ditches were blocked. Aerial view of the rewetted peatland with boardwalks visible as white lines (B). (Aerial map from Lantmäteriet)

Results

The impact of rewetting on GWL variation

Peatland rewetting led to a significant increase in GWL at the two rewetted catchments (R1 and R2) compared to the control site (Figure 2a). The relative difference in GWL between the rewetted and control sites (treatment minus control) at varying distances to the ditch also showed a significant decrease after rewetting (Figure 2b). Interestingly, this impact demonstrated variability depending on the distance from the ditch, with wells located closest to the ditch showing a more pronounced response compared to those farther away. Prior to rewetting, the median GWL was significantly ($p<0.05$) lower next to the ditch (-228 mm) compared to the furthest distance away (-174 mm). Furthermore, GWL exhibited greater variability in the middle of the transect (50 m from the ditch), reaching a minimum of 507 mm from the ground. After rewetting, the largest median GWL change was observed at a distance of 10 meters, with an increase of 119 mm. This was followed by a median 91 mm increase at a distance of 100 meters and a median 62 mm increase at a distance of 50 meters. The median GWL at the control sites was similar during the pre- and post-rewetting periods (-79 and -78 mm, respectively). Finally, GWL in the three different distances from ditch significantly increased in the first year after rewetting, significantly increased further in the second year and significantly decreased in the third year, however, still being significantly higher than the pre-rewetting and first year post-rewetting GWL.

289

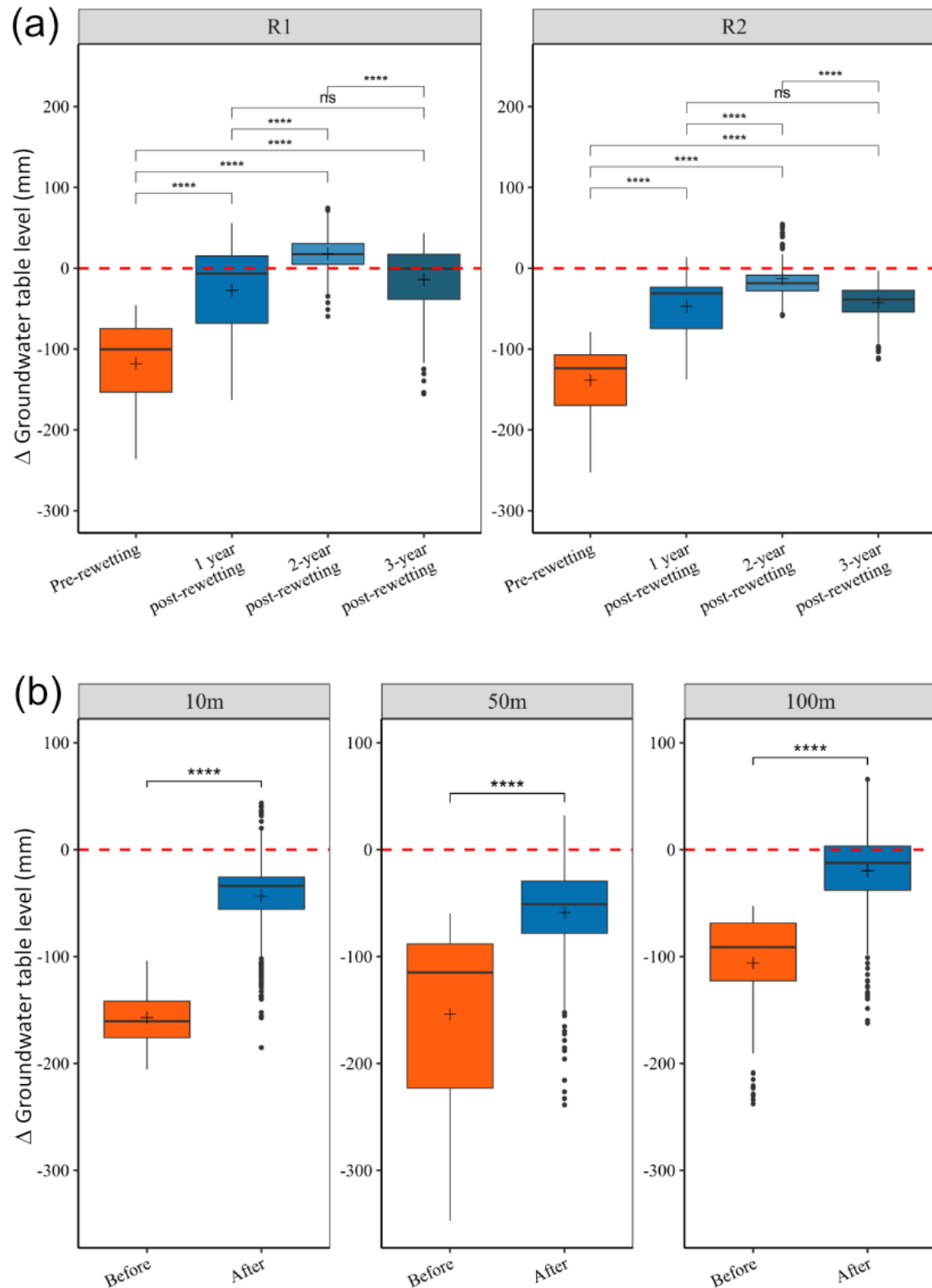



Figure 2.a) Difference (treatment-control) in groundwater table level (GWL) at the two rewetted catchments (R1 and R2) based
on daily data gathered between June to October in the years 2020 (pre-rewetting) and 2021, 2022 and 2023 (3 years post-
rewetting) regardless of distance to ditch. b) Relative difference in GWL based on varying distances to the main ditch (i.e. 10,
50 and 100 m); all years post-rewetting are combined (sample sizes for pre-rewetting and post-rewetting were 153 and 428,
respectively). The red dashed line indicates the GWL of the control site; positive values indicate that the GWL is higher at the
rewetted site than at the control, while negative GWL indicate the opposite. The box plots show the minimum, first quartile,
median, third quartile, and maximum, with outliers as dots. The stars indicate the levels of significant difference between the
marked comparisons as determined using a Wilcoxon test (****$p \le 0.0001$).

298 *Table 1. Median, minimum (min), maximum (max) and 5th-95th quantile of groundwater level (GWL) change pre- and post-*
299 *rewetting for different distances to the ditch (i.e. 10, 50 and 100 m) and the control site.*

	Distance	Median (mm)	Min (mm)	Max (mm)	5th- 95th quantile (mm)
PRE-REWETTING	10 m	-228	-364	-120	194
	50 m	-190	-507	-60	370
	100 m	-174	-416	-44	304
	Control	-79	-186	9	156
POST-REWETTING	10 m	-108	-272	-33	197
	50 m	-127	-366	-30	233
	100 m	-83	-341	5	240
	Control	-78	-234	3	171

The impact of rewetting on runoff responses

301 Based on the response at R1, 17 rainfall-runoff events before and 30 events after rewetting were
302 extracted and analyzed (Figure 3). The impact of rewetting on runoff responses during rainfall-runoff
303 events is depicted through examples of event-scale hydrographs (Figure 4, Table S 1). The illustrations
304 display the variation in discharge response across control and the two rewetted catchments (R1 and
305 R2) for different event sizes and antecedent GWL conditions, during both pre-and post-rewetting
306 periods. In the pre-rewetting period, despite the control site having the shallowest GWL at -15 mm,
307 exhibited the smallest peak flow of 0.29 mm h^{-1}. In contrast, the rewetted site R1, with an antecedent
308 GWL of -82 mm h^{-1}, reached a peak of 0.93 mm h^{-1}. One and two years after rewetting, R1 still had the
309 highest peak flows at 0.71 mm h^{-1} and 0.61 mm h^{-1}, respectively, while the rewetted catchment R2
310 showed similarities to the control site (Figure 2). However, three years after rewetting, although R1 had
311 the shallowest antecedent GWL at -5.15 mm, the peak flow was almost half of the peak in the control
312 catchment (0.14 mm h^{-1} and 0.26 mm h^{-1}, respectively).

313

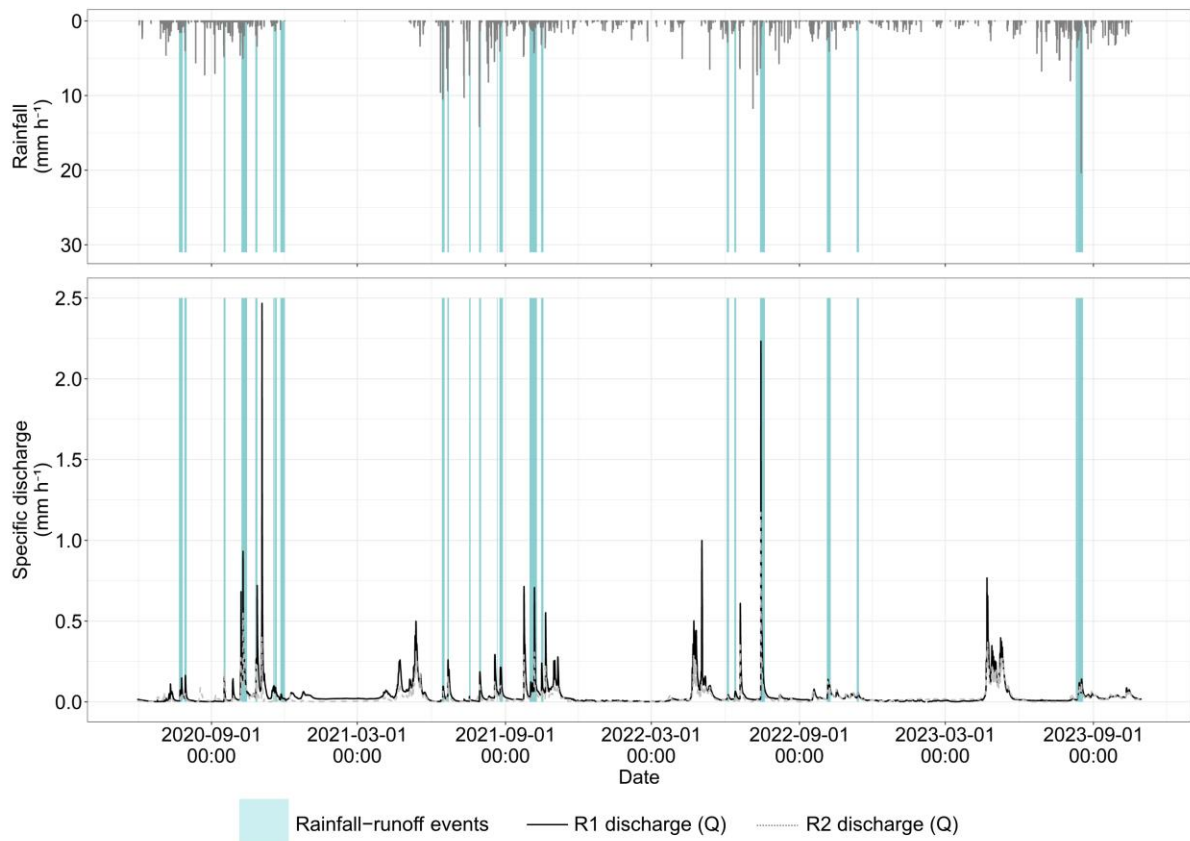
314

315

316

317

318

319

320   *Figure 3. Identified rainfall-runoff events using discharge measured at the rewetted catchment R1 across the entire study period*
321   *(black line). Same rainfall-runoff events identified using R1 discharge are shown for R2 (grey line).*

322

323

324

325

326

327

328

329

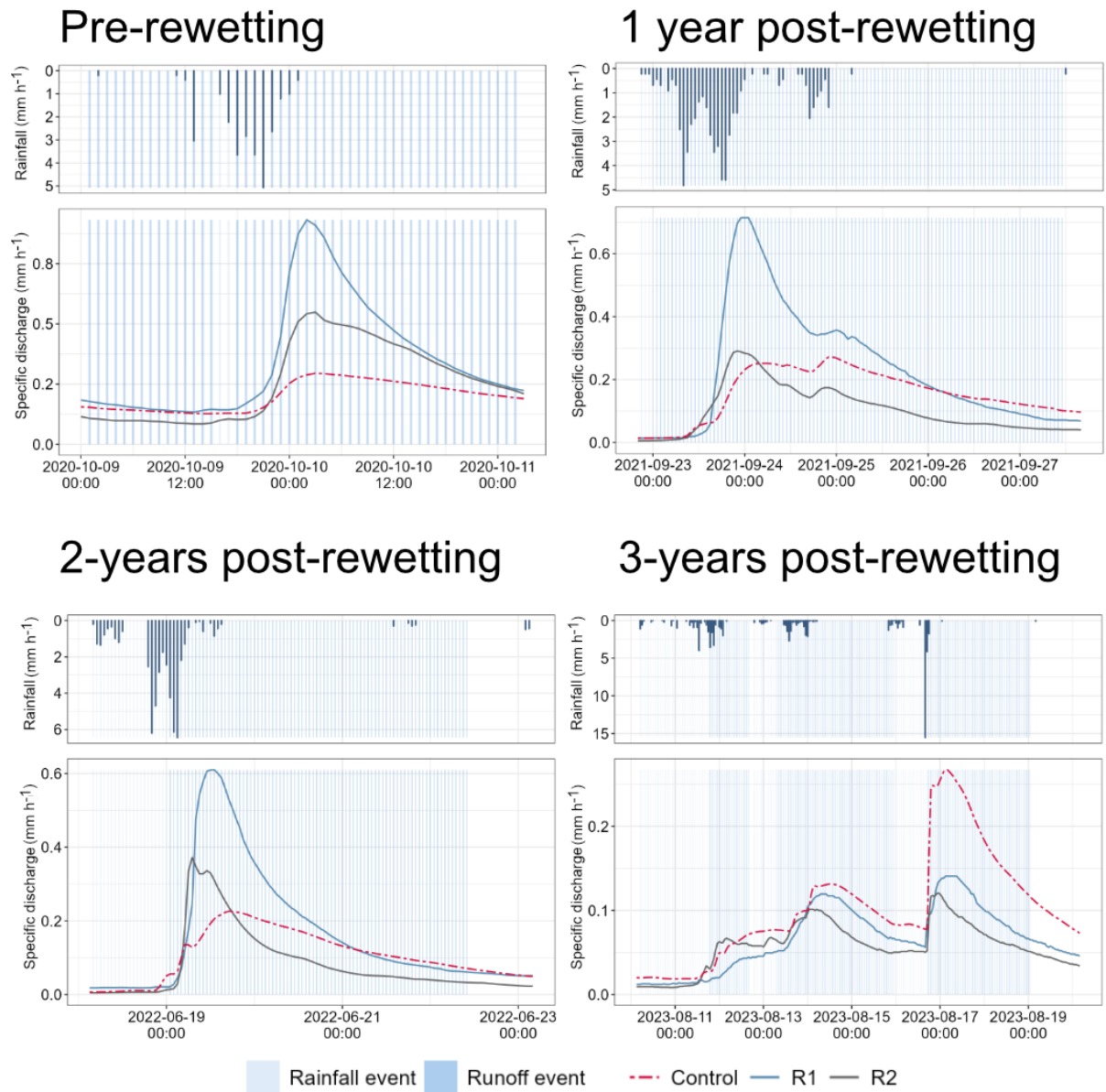



Figure 4. Examples of runoff responses of control and the two catchments (R1 and R2) of the rewetted site during rainfall-runoff events for each of the four pre- and post-rewetting years. Note that the scales for the y-axes show different magnitudes of specific discharge.

330
331
332

333

Flood mitigation effects of rewetting

The magnitude of the effects of peatland rewetting was investigated for 47 rainfall-runoff events (17 events before rewetting and 30 events after rewetting) to test if the rewetting's effects were significant under a larger number of events. Storm magnitudes ranged between 5 and 50 mm in total precipitation before rewetting, and 2 and 63 mm after rewetting. The relative differences between the two catchments (R1 and R2) of the rewetted site and control sites (rewetted minus control) for each metric are shown in Figure 5.

The analysis of rainfall-runoff events revealed a reduction in relative peak flow at the two catchments (R1 and R2) of the rewetted site following rewetting (Figure 5a). However, the reduction was significant only at R1. Specifically, the median peak flow at R1 decreased from 0.14 to 0.10 mm h^{-1} post-rewetting. In contrast, at R2, there was an increase from 0.04 to 0.08 mm/h post-rewetting. Interestingly, the control site experienced a rise in median peak flow from 0.05 to 0.12 mm h^{-1} during the post-rewetting period.

The median runoff coefficient in the two catchments of the rewetted site showed an increase from 0.36 to 0.40 and from 0.14 to 0.20 at R1 and R2, respectively, after rewetting. The runoff coefficient at the control site increased from 0.17 before rewetting to 0.40 after rewetting. Relative to the control site, both restored sites, R1 and R2, experienced a decline in runoff coefficients during the post-rewetting phase. Notably, this reduction was statistically significant solely at R1 ($p<0.01$) (Figure 5b).

After rewetting, the median lag time in the two catchments (R1 and R2) of the rewetted site decreased by 0.5 and 7 hours, reaching 15 and 10 hours for R1 and R2, respectively. In contrast, the control catchment exhibited an increase in median lag time from 14 to 23 hours during the post-rewetting period. However, pairwise test results indicated that there was no statistically significant change at both rewetted catchments (R1 and R2) following rewetting (Figure 5c).

The median HSI values for both catchments (R1 and R2) of the rewetted site and control sites decreased after the rewetting period, shifting from 0.023 to 0.021, 0.034 to 0.025, and 0.027 to 0.026 at control, R1, and R2, respectively (Fig. 5d). The effect of rewetting in reducing HSI was significant only at R1 ($p < 0.0001$). Prior to rewetting, the relative HSI at R1 was 0.012, and after rewetting, it decreased to 0.003. The relative HSI also experienced a decline at R2, dropping from 0.006 pre-rewetting to 0.004 after rewetting. However, this decrease was not statistically significant.

363

364

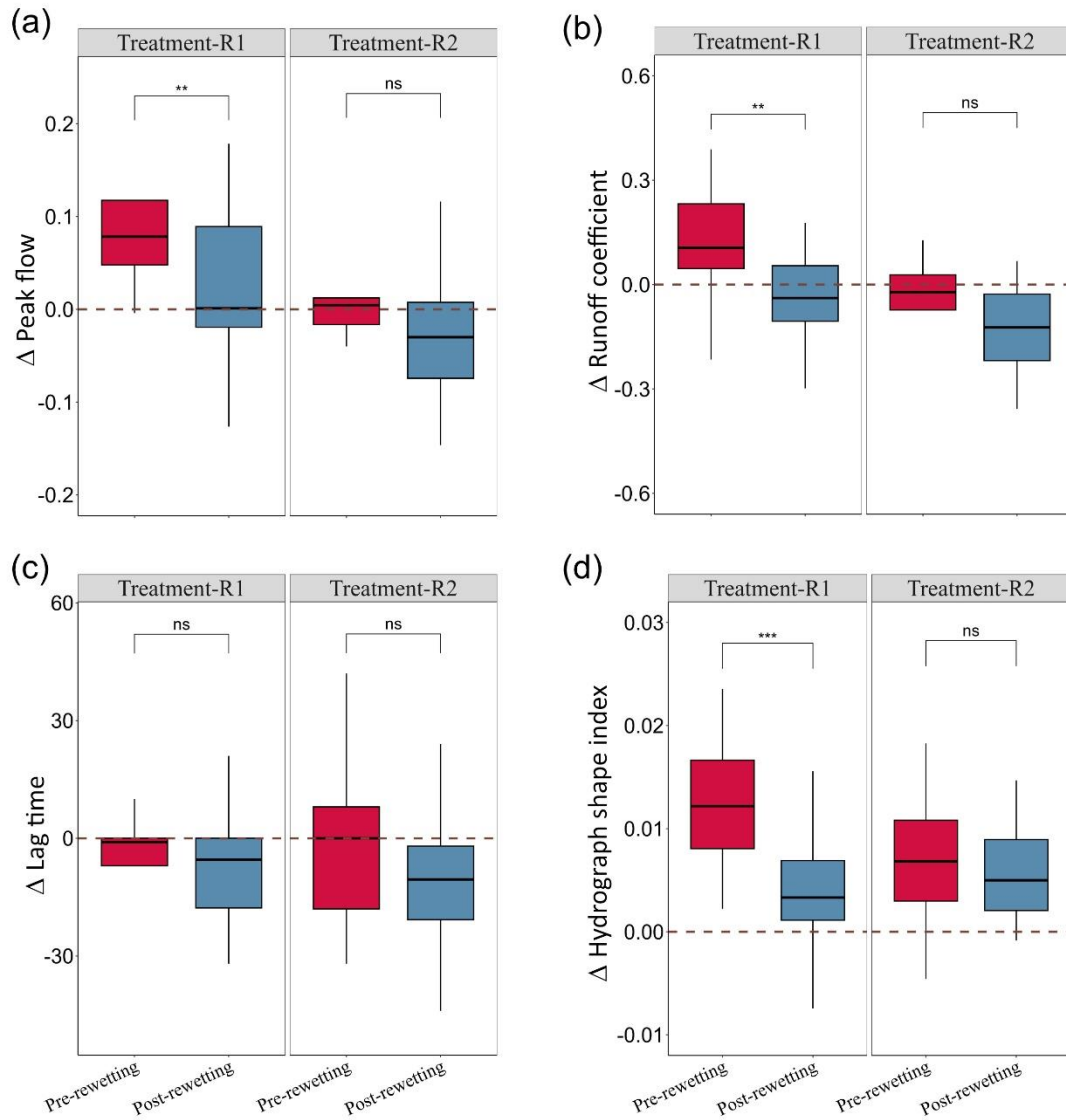



Figure 5.Differences between the rewetted and control sites pre- and the combined three years of post- rewetting period for (a) peak flow, (b) runoff coefficient, (c) lag time, and (d) Hydrograph shape index. The relative difference was computed as treatment minus control and the red dashed line indicates the value of the control site; thus, positive values indicate that the hydrological response is greater at the treatment site than at the control site, while negative values indicate the opposite. The box plots show the minimum, first quartile, median, third quartile, and maximum, with outliers as points. The stars indicate the levels of significance in Wilcoxon test (**p ≤ 0.01; ***p ≤ 0.001; "ns" denotes not significant.).

371

Discussion

Despite significant interest in peatland rewetting, there has been limited research on its effects on hydrological functioning and the scale of these impacts. We found that peatland rewetting on nutrient-poor minerogenic fens, one of the most common peatland types in Fennoscandia, was generally positive for use as natural flood management. Rewetting has begun to influence GWL, runoff responses during rainstorms, and flood mitigation (though the latter was observed in only one of the two study catchments), while also shifting these hydrological characteristics closer to pristine conditions by increasing water storage in the peatland. However, special attention should be given to the diverse characteristics of peatlands in the boreal biome before generalizing the effect of peatland rewetting on hydrological functioning.

The impact of rewetting on GWL

Using the BACI experimental approach, we found that the mean GWL of the rewetted sites rose rather rapidly after ditch-blocking at both R1 and R2 to the near pristine levels of our control site. Our results align broadly with several other studies mainly from Finland, Canada and UK that found that peatland rewetting raised GWL to near pristine levels (Armstrong et al., 2022; Dixon et al., 2014; Haapalehto et al., 2014; Howie et al., 2009; Menberu et al., 2016; Shuttleworth et al., 2019; Soomets et al., 2023). Notably, GWL rose in all studies, despite variation in the extent of recovery, climate and drainage conditions, such as age, depth, and pattern of ditching. Nevertheless, peatland type is likely to have a major impact on the time taken to increase GWL to pristine-like levels after ditch blocking. For example, fen peatlands, such as our system, with a rather flat or slightly domed surface and slow lateral movement of water could have a faster response to ditch blocking compared to blanket bogs which may exist even on slopes of 20 to 25 degrees (Laine et al., 2011). Furthermore, our results also revealed that the median GWL at R1 closely resembled that of the control site after rewetting. However, at R2, the median GWL remained slightly lower post-rewetting. This difference may be attributed to the presence of shrubs and sparse tree cover (higher water uptake) on the mire at R2, as well as a lower proportion of blocked ditches within the catchment. However, as found by Howie et al. (2009) in Southwestern British Columbia, the difference between the effect of rewetting on GWL between the two catchments could also be attributed to differences in the drying of the peatland. This, coupled with shrinkage and subsidence of the peat, could lead to a reduction in hydraulic conductivity, possibly hindering the effectiveness of restoration efforts in reversing the impacts of drainage.

Our results addressed a gap in the existing literature by examining the spatial variability of GWL recovery at different distances from the ditch, a factor largely neglected in prior research, particularly

404 within the context of boreal ecosystems (Bring et al., 2022). After rewetting, our results show a
405 significant increase in GWL at all distances from the ditch, however with variation spatially. These
406 detailed spatial results of GWL increase at different distances to the ditches, show that all of the
407 locations in the mire had undergone rewetting and any observed differences in event runoff responses
408 could be attributed to changes in GWL and water storage within the full extent of the peatland.
409 Furthermore, we found that after rewetting, significant differences persisted between locations, with
410 the highest GWL found at the furthest distance (i.e., 100 m), contrary to most other published studies
411 that have found that the impact of rewetting on GWL diminishes with increasing distance from the main
412 ditch (Bring et al., 2022). This difference in GWL between distances from ditch might be a topographic
413 effect, where peat surfaces could be at higher elevations at the further distance due to long-term
414 consolidation of the peat near ditches and therefore the absolute GWL was also higher at the further
415 distance (Holden et al., 2017). Either way, our results showed that rewetting was successful in returning
416 the pristine-like patterns of higher GWL, however, restoration must also intend to reach an even GWL
417 throughout the mire which we expect to see with time (Haapalehto et al., 2014; Laine et al., 2011).
418 Furthermore, our data serves as a valuable resource for peatland managers, especially in boreal
419 ecosystems, helping to gain a better understanding of site-specific changes in hydrology and the
420 associated ecosystem services that result from the rewetting of peatlands.

The impact of rewetting on runoff responses

422 Event-based analysis of discharge responses provides important information on treatment effects on
423 the hydrological functionality. Relying solely on daily discharge analysis does not offer much insight into
424 discharge responses to precipitation, including the lag time to peak flow. For instance, examining the
425 hourly hydrograph revealed that, discharge responses at R1 exhibited flashier characteristics with
426 higher peaks compared to those at R2, yet R2 did not show a significant difference in the overall
427 decrease in relative peak flow after rewetting. This discrepancy could possibly be attributed to the
428 smaller increase of GWL of R2 compared to R1, likely influenced by a lower proportion of blocked
429 ditches in the catchment, the re-creation of the open-water pond in R1, or the fact that the relative
430 peak flow at R2 before rewetting was already similar to the control peatland. Specifically, the re-
431 creation of the old open-water pond ca. 100 meters from R1 outlet after rewetting, could potentially
432 be functioning as a "natural peatland pool", hence influencing runoff responses (Arsenault et al., 2019).
433 Nonetheless, three years after ditch blocking, both catchments showed attenuated hydrograph shapes
434 during most storm events. These changes in the hydrograph characteristics following rewetting indicate
435 that hydrological restoration positively affects overall flow regimes influencing flow pathways and
436 water storage within the peat, leading to reduced peak flow and increased hydrological residence time

437   in the peatland (Gatis et al., 2023). However, the scarcity of continuous, prolonged datasets from

438   rewetted peatlands, particularly for boreal minerogenic fens, poses a significant challenge in

439   conducting comprehensive comparisons across various peatland sizes, types, and rewetting durations,

440   as most rewetting projects have only recently commenced. Therefore, a more extended period of post-

441   rewetting monitoring is necessary to fully understand how the discharge patterns of drained peatlands

442   evolve after rewetting.

443   ## Flood mitigation effects of rewetting

444   Rewetting resulted in a significant reduction in event peak flow response at R1. It is noteworthy that,

445   although there was a significant decrease, the median relative peak flow at R1 after rewetting was still

446   higher than the control. This is in contrast to R2, where the decrease in peak flow was not significant,

447   but where the median peak flow started out similar to the control mire, but decreased to be somewhat

448   lower than the control after rewetting. Our findings align with the results observed by Wilson et al.

449   (2011), where they showed peak flow hydrographs from ditches with considerable change after

450   rewetting, with lower peak flow rates, less runoff and rainwater being released during events. In

451   contrast, Shantz and Price (2006) evaluated the hydrological characteristics of a restored peatland in

452   Quebec, Canada, and observed higher discharge peaks during summer at the restored site compared

453   to the control site. They attributed this to wetter antecedent conditions and faster drainage response

454   following rainfall. Conversely, our research reveals that despite observing a rise in GWL after rewetting,

455   rewetted peatlands can exhibit less intense flood responses and offer improved retention of rainfall.

456   Moreover, although reduced runoff rate following restoration inevitably increases the chances of

457   overland flow, this pathway is considerably slower than flow through the drainage networks and

458   therefore restoration may lead to reduction in peak flow magnitude (Grand-Clement et al., 2013). Here,

459   the old-open water pond, functioning as a peatland pool, could potentially increase runoff detention

460   and therefore reduced peak flow (Arsenault et al., 2019; Kløve, 2000). Finally, it is noteworthy that even

461   before rewetting, our system was already dominated by *Sphagnum* spp., therefore compared to other

462   degraded (especially bare) peatlands, our system had a higher potential for reducing sheet erosion and

463   downstream flood peaks after rewetting (Holden et al., 2008). Either way, our results suggest that

464   contrary to conclusions drawn in many previous studies (Holden et al., 2004; Holden & Burt, 2003), the

465   rewetted peatlands in our study exhibited more controlled and resilient hydrological behavior

466   delivering "natural flood management" by attenuating downstream flow and reducing flood risk. Yet,

467   our results could be affected by the short time period since rewetting. With time, the peatland might

468   respond differently depending on the type of newly established vegetation and initial moisture

conditions, where rainfall events could trigger rapid and concentrated runoff and discharge (Holden & Burt, 2003).

Runoff coefficient is another key indicator for flood mitigation and corresponds to catchment storage capacity. Specifically, a reduction in runoff coefficient indicates a gain in storage capacity either behind the dam in the blocked ditches or due to increased surface roughness which then reduces the drainage efficiency, thereby increasing water storage capacity (Menberu et al., 2016).Our results showed that reduction in runoff coefficient was significant at R1, showing less runoff being exported with rainfall events after rewetting. Yet again, this reduction was not significant at R2, and the same patterns follows, where the initial condition of the runoff coefficient of R2 was lower than the control mire before rewetting and after rewetting the coefficient lowered even further. In line with our results, the reduction in runoff coefficient after peatland rewetting has been reported in many studies (Ketcheson & Price, 2011; Shantz & Price, 2006; Wilson et al., 2011), in some, to the extent of being the most significant hydrological effect of peatland rewetting (Ketcheson & Price, 2011). However, caution is needed in interpreting some of these results due to the potential influence of the relatively short time series during which the peatland could have been still undergoing filling (Ketcheson & Price, 2011) or an eventual increase in the runoff coefficient due to a declining efficiency of the ditch blocking (Menberu et al., 2018).

An increasing lag time traditionally serves as a positive indicator for flood modification, as downstream flow becomes less "flashy". Contrary to expectations, the lag time between the initiation of a rainfall event and the peak discharge decreased after rewetting in both rewetted catchments. However, it is important to note that this decrease, while observed, was not statistically significant for either catchment. A shorter lag time may support the interpretation that flow peaks observed after rewetting originated from the near vicinity of the monitoring site, while water from upstream areas was attenuated before reaching the outlet (Gatis et al., 2023). These explanations could be further supported by coupling these results with the positive effect on base flow that has been seen in our study site (Karimi et al., 2024). Finally, HSI, which serves as a direct indicator of system flashiness, exhibited a notable decrease at the R1 catchment following rewetting. Unlike the pattern observed in R2's response to other hydrological factors, the HSI at R2 was higher than that of the control site. Although R2 showed a non-significant reduction after rewetting, it remained consistently higher than the control. This reduction seems to be a combined effect of ditch blocking and presence of vegetation as other studies have shown that ditch blocking by itself does not appear to alter the "flashiness" of stormflow (Shuttleworth et al., 2019).

501 The significant decreases in peak flow, runoff coefficient and HSI observed at R1, compared to the non-
502 significant changes at R2, can be attributed to several factors. Firstly, the BACI analysis indicated that,
503 prior to rewetting, R1 had flashier hydrological responses. Moreover, R2's responses were already more
504 similar to the control site, suggesting less potential to observe a significant post-rewetting effect.
505 Additionally, a smaller portion of R2 catchment was restored compared to R1, which could mean that
506 the overall water storage at R2 remained lower than at R1. Consequently, water may still drain more
507 quickly at R2, leading to less noticeable impacts from the rewetting efforts. Furthermore, as mentioned
508 above, the differences in responses could in part be due to the recreation of the peatland pool that
509 likely influences runoff detention. All in all, the diverse responses observed in flood response
510 characteristics, both in our study and other investigations, raises questions regarding the overall
511 effectiveness of peatland rewetting, at least in the short term. While it appears successful in reducing
512 peak flow, runoff coefficient, and overall flashiness of hydrographs (as shown by HSI), our results
513 suggests it might not be as effective in increasing lag time from peak rainfall to peak flow occurrence.
514 This limitation could potentially be attributed to the need for new peat formation before becoming fully
515 hydrologically restored. Hence, a crucial question regarding the duration of these effects and the time
516 necessary for lag time recovery remains unanswered.

517 The effectiveness of ditch-blocking in flood moderation is influenced by various factors, including the
518 initial condition of a drained peatland, the extent of peat degradation and restoration, and changes in
519 its properties (Menberu et al., 2016). Furthermore, there may be a delayed effect in the peatland's
520 response to ditch-blocking, and the corresponding flood mitigation may progressively change over time
521 in the years following the blocking of ditches due to changes in peat properties and vegetation cover.
522 Overall, our rewetted sites, having been drained for a century, still may not function as a natural
523 peatland and a full hydrological recovery will take substantially longer than the recovery period we
524 measured here. Several factors linked to prolonged drainage could contribute to a long recovery period.
525 For instance, peat oxidation and compaction may lead to increased bulk density, which in turn, affects
526 the ability of the site to effectively retain and release water (Liu & Lennartz, 2019). Hence, our three-
527 year post-rewetting monitoring period, while longer than many other studies, still offers rather limited
528 insight into the impact of rewetting on flood moderation under extreme storm events, especially in
529 more severe future climatic conditions. Therefore, further monitoring is required to understand the
530 influence of restoration practices on peatland hydrological functioning.

Conclusion

531 Our results showed that the effect of rewetting on flow moderation from rainfall events is not as simple

533 as restoring GWL. This gradual and evolving process of peatland hydrological functioning due to a long

534 history of peat compaction and decomposition, then the subsequent re-establishment of peat-forming

535 vegetation after rewetting emphasizes the importance of sustained long-term monitoring to fully

536 understand the outcomes of rewetting. Moreover, our findings indicate that peatland rewetting has

537 the potential for flood mitigation and, in some cases, the ability to mitigate runoff from rainfall events

538 better than pristine sites. This was supported by reductions in peak flow, runoff coefficient, and less

539 flashy hydrograph responses (HSI). However, the results showed that peatland rewetting would not

540 necessarily increase the lag time between the peak of a rainfall event and peak discharge. Significant

541 changes were only observed at one of the two restored peatlands. These differences seem to be

542 attributed to (1) a higher percentage of the ditches in the catchment being restored and (2) that the

543 none-effected site already being similar to the pristine site, suggesting less potential changes post-

544 rewetting. Nevertheless, uncertainties remain in our understanding of the contribution of peatland

545 rewetting to natural flood management over longer timescales or during large historical flood events.

546 Therefore, we emphasize the significance of long-term monitoring combined with hydrological

547 modeling to determine whether peatlands will consistently mitigate floods as climate change

548 intensifies.

549

550

Code and data availability

All data used in this study are freely available. The discharge data can be obtained from https://data.fieldsites.se/portal/. The GWL data up to October 2023 are available from the corresponding author. The original R codes for extracting rainfall-runoff events are available from Gatis et al. (2023) at https://ore.exeter.ac.uk/repository/handle/10871/134028.

Financial support

The TEA infrastructure was initiated and co-funded by the European Union GRIP on LIFE IP project (LIFE16IPE SE009 GRIP) led by the Västerbotten Administration Board and Swedish Forest Agency, with additional financial infrastructure and research support from The Kempe Foundation and the Swedish Research Council Formas grants (2018-00723 (to EMH), 2018-02780 (to HL), 2020-01372 (to HL), 2021-02114 (to HL), as well as by the Knut and Alice Wallenberg (Grants 2018.0259 and 2023.0245). The KCS/KFI infrastructure and long-term data collection have been funded by the Swedish Research Council VR (SITES, grant number 2021-00164).

Acknowledgements

We would like to thank all the skilled and dedicated field personnel at the Svartberget research station.

Competing interests

The authors declare that they have no conflict of interest.

References

Acreman, M., & Holden, J. (2013). How wetlands affect floods. *Wetlands*, *33*(5), 773–786. https://doi.org/10.1007/s13157-013-0473-2

Aghakouchak, A., Chiang, F., Huning, L. S., Love, C. A., Mallakpour, I., Mazdiyasni, O., Moftakhari, H., Papalexiou, S. M., Ragno, E., & Sadegh, M. (2020). Climate Extremes and Compound Hazards in a Warming World. *Annual Review of Earth and Planetary Sciences*, *48*, 519–548. https://doi.org/10.1146/annurev-earth-071719-055228

Andersen, R., Farrell, C., Graf, M., Muller, F., Calvar, E., Frankard, P., Caporn, S., & Anderson, P. (2017). An overview of the progress and challenges of peatland restoration in Western Europe. *Restoration Ecology*, *25*(2), 271–282. https://doi.org/10.1111/rec.12415

Armstrong, A., Holden, J., Kay, P., Francis, B., Foulger, M., Gledhill, S., McDonald, A. T., & Walker, A. (2010). The impact of peatland drain-blocking on dissolved organic carbon loss and discolouration of water; results from a national survey. *Journal of Hydrology*, *381*(1–2), 112–120. https://doi.org/10.1016/j.jhydrol.2009.11.031

Armstrong, L., Peralta, A., Krauss, K. W., Cormier, N., Moss, R. F., Soderholm, E., McCall, A., Pickens, C., & Ardón, M. (2022). Hydrologic Restoration Decreases Greenhouse Gas Emissions from Shrub Bog Peatlands in Southeastern US. *Wetlands*, *42*(7), 81. https://doi.org/10.1007/s13157-022-01605-y

Arsenault, J., Talbot, J., Moore, T. R., Beauvais, M. P., Franssen, J., & Roulet, N. T. (2019). The Spatial Heterogeneity of Vegetation, Hydrology and Water Chemistry in a Peatland with Open-Water Pools. *Ecosystems*, *22*(6), 1352–1367. https://doi.org/10.1007/s10021-019-00342-4

Blume, T., Zehe, E., & Bronstert, A. (2007). Rainfall-runoff response, event-based runoff coefficients and hydrograph separation. *Hydrological Sciences Journal*, *52*(5), 843–862. https://doi.org/10.1623/hysj.52.5.843

Bring, A., Rosén, L., Thorslund, J., Tonderski, K., Åberg, C., Envall, I., & Laudon, H. (2020). Groundwater storage effects from restoring, constructing or draining wetlands in temperate and boreal climates: a systematic review protocol. *Environmental Evidence*, *9*(1), 1–11. https://doi.org/10.1186/s13750-020-00209-5

Bring, A., Thorslund, J., Rosén, L., Tonderski, K., Åberg, C., Envall, I., & Laudon, H. (2022). Effects on groundwater storage of restoring, constructing or draining wetlands in temperate and boreal climates: a systematic review. *Environmental Evidence*, *11*(1), 38.

599   https://doi.org/10.1186/s13750-022-00289-5

600   Casselgård, M. (2020). *Effects of 100 years of drainage on peat properties in a drained peatland forests*
601   *in northern Sweden Examensarbeten 2020:4 Fakulteten för skogsvetenskap Institutionen för*
602   *skogens ekologi och skötsel*. https://stud.epsilon.slu.se

603   D'Acunha, B., Lee, S. C., & Johnson, M. S. (2018). Ecohydrological responses to rewetting of a highly
604   impacted raised bog ecosystem. *Ecohydrology*, *11*(1), 1–12. https://doi.org/10.1002/eco.1922

605   Dixon, S. D., Qassim, S. M., Rowson, J. G., Worrall, F., Evans, M. G., Boothroyd, I. M., & Bonn, A. (2014).
606   Restoration effects on water table depths and CO2 fluxes from climatically marginal blanket bog.
607   *Biogeochemistry*, *118*(1–3), 159–176. https://doi.org/10.1007/s10533-013-9915-4

608   Duque, M. L. F. (2022). *R package IETD* (pp. 1–9). https://doi.org/10.3390/w6010045>.Restrepo-Posada

609   Edokpa, D., Milledge, D., Allott, T., Holden, J., Shuttleworth, E., Kay, M., Johnston, A., Millin-Chalabi, G.,
610   Scott-Campbell, M., Chandler, D., Freestone, J., & Evans, M. (2022). Rainfall intensity and
611   catchment size control storm runoff in a gullied blanket peatland. *Journal of Hydrology*,
612   *609*(October 2021), 127688. https://doi.org/10.1016/j.jhydrol.2022.127688

613   Evans, M. G., Burt, T. P., Holden, J., & Adamson, J. K. (1999). Runoff generation and water table
614   fluctuations in blanket peat: Evidence from UK data spanning the dry summer of 1995. *Journal of*
615   *Hydrology*, *221*(3–4), 141–160. https://doi.org/10.1016/S0022-1694(99)00085-2

616   Franzen, L. G., Lindberg, F., Viklander, V., & Walther, A. (2012). The potential peatland extent and
617   carbon sink in Sweden, as related to the Peatland/Ice Age Hypothesis. *Mines and Peat*, *10*(8), 1–
618   19.

619   Gatis, N., Benaud, P., Anderson, K., Ashe, J., Grand-Clement, E., Luscombe, D. J., Puttock, A., & Brazier,
620   R. E. (2023). Peatland restoration increases water storage and attenuates downstream stormflow
621   but does not guarantee an immediate reversal of long-term ecohydrological degradation.
622   *Scientific Reports*, *13*(1), 1–14. https://doi.org/10.1038/s41598-023-40285-4

623   Goudarzi, S., Milledge, D. G., Holden, J., Evans, M. G., Allott, T. E. H., Shuttleworth, E. L., Pilkington, M.,
624   & Walker, J. (2021). Blanket Peat Restoration: Numerical Study of the Underlying Processes
625   Delivering Natural Flood Management Benefits. *Water Resources Research*, *57*(4).
626   https://doi.org/10.1029/2020WR029209

627   Grand-Clement, E., Anderson, K., Smith, D., Luscombe, D., Gatis, N., Ross, M., & Brazier, R. E. (2013).
628   Evaluating ecosystem goods and services after restoration of marginal upland peatlands in South-

629   west England. *Journal of Applied Ecology*, *50*(2), 324–334. https://doi.org/10.1111/1365-
630   2664.12039

631   Haapalehto, T., Kotiaho, J. S., Matilainen, R., & Tahvanainen, T. (2014). The effects of long-term drainage
632   and subsequent restoration on water table level and pore water chemistry in boreal peatlands.
633   *Journal of Hydrology*, *519*(PB), 1493–1505. https://doi.org/10.1016/j.jhydrol.2014.09.013

634   Haque, A., Ali, G., & Badiou, P. (2022). Event-based analysis of wetland hydrologic response in the
635   Prairie Pothole Region. *Journal of Hydrology*, *604*(September 2021), 127237.
636   https://doi.org/10.1016/j.jhydrol.2021.127237

637   Hawcroft, M., Walsh, E., Hodges, K., & Zappa, G. (2018). Significantly increased extreme precipitation
638   expected in Europe and North America from extratropical cyclones. *Environmental Research*
639   *Letters*, *13*(12). https://doi.org/10.1088/1748-9326/aaed59

640   Helbig, M., Waddington, J. M., Alekseychik, P., Amiro, B. D., Aurela, M., Barr, A. G., Black, T. A., Blanken,
641   P. D., Carey, S. K., Chen, J., Chi, J., Desai, A. R., Dunn, A., Euskirchen, E. S., Flanagan, L. B., Forbrich,
642   I., Friborg, T., Grelle, A., Harder, S., … Zyrianov, V. (2020). Increasing contribution of peatlands to
643   boreal evapotranspiration in a warming climate. *Nature Climate Change*, *10*(6), 555–560.
644   https://doi.org/10.1038/s41558-020-0763-7

645   Holden, J. (2006). Chapter 14 Peatland hydrology. In I. P. Martini, A. M. Cortizas, & W. Chesworth (Eds.),
646   *Developments in Earth Surface Processes* (Vol. 9, Number C, pp. 319–346).
647   https://doi.org/10.1016/S0928-2025(06)09014-6

648   Holden, J., & Burt, T. P. (2003). Runoff production in blanket peat covered catchments. *Water Resources*
649   *Research*, *39*(7). https://doi.org/10.1029/2002WR001956

650   Holden, J., Chapman, P. J., & Labadz, J. C. (2004). Artificial drainage of peatlands: Hydrological and
651   hydrochemical process and wetland restoration. *Progress in Physical Geography*, *28*(1), 95–123.
652   https://doi.org/10.1191/0309133304pp403ra

653   Holden, J., Green, S. M., Baird, A. J., Grayson, R. P., Dooling, G. P., Chapman, P. J., Evans, C. D., Peacock,
654   M., & Swindles, G. (2017). The impact of ditch blocking on the hydrological functioning of blanket
655   peatlands. *Hydrological Processes*, *31*(3), 525–539. https://doi.org/10.1002/hyp.11031

656   Holden, J., Kirkby, M. J., Lane, S. N., Milledge, D. G., Brookes, C. J., Holden, V., & McDonald, A. T. (2008).
657   Overland flow velocity and roughness properties in peatlands. *Water Resources Research*, *44*(6),
658   1–11. https://doi.org/10.1029/2007WR006052

659   Howie, S. A., Whitfield, P. H., Hebda, R. J., Munson, T. G., Dakin, R. A., & Jeglum, J. K. (2009). Water table

660   and vegetation response to ditch blocking: Restoration of a raised bog in southwestern british

661   columbia. *Canadian Water Resources Journal*, *34*(4), 381–392.

662   https://doi.org/10.4296/cwrj3404381

663   Javaheri, A., & Babbar-Sebens, M. (2014). On comparison of peak flow reductions, flood inundation

664   maps, and velocity maps in evaluating effects of restored wetlands on channel flooding. *Ecological*

665   *Engineering*, *73*, 132–145. https://doi.org/10.1016/j.ecoleng.2014.09.021

666   Karimi, S., Hasselquist, E., Salimi, S., Järveoja, J., & Laudon, H. (2024). Rewetting impact on the

667   hydrological function of a drained peatland in the boreal landscape. *Journal of Hydrology*,

668   *641*(April). https://doi.org/10.1016/j.jhydrol.2024.131729

669   Karlsen, R. H., Bishop, K., Grabs, T., Ottosson-Löfvenius, M., Laudon, H., & Seibert, J. (2019). The role of

670   landscape properties, storage and evapotranspiration on variability in streamflow recessions in a

671   boreal catchment. *Journal of Hydrology*, *570*(December 2018), 315–328.

672   https://doi.org/10.1016/j.jhydrol.2018.12.065

673   Ketcheson, S. J., & Price, J. S. (2011). The impact of peatland restoration on the site hydrology of an

674   abandoned block-cut bog. *Wetlands*, *31*(6), 1263–1274. https://doi.org/10.1007/s13157-011-

675   0241-0

676   Kirchner, J. W., Benettin, P., & Van Meerveld, I. (2023). Instructive Surprises in the Hydrological

677   Functioning of Landscapes. *Annual Review of Earth and Planetary Sciences*, *51*, 277–299.

678   https://doi.org/10.1146/annurev-earth-071822-100356

679   Kløve, B. (2000). Retention of suspended solids and sediment bound nutrients from peat harvesting

680   sites with peak runoff control, constructed floodplains and sedimentation ponds. *Boreal*

681   *Environment Research*, *5*(1), 81–94.

682   Laine, A. M., Leppälä, M., Tarvainen, O., Päätalo, M. L., Seppänen, R., & Tolvanen, A. (2011). Restoration

683   of managed pine fens: Effect on hydrology and vegetation. *Applied Vegetation Science*, *14*(3),

684   340–349. https://doi.org/10.1111/j.1654-109X.2011.01123.x

685   Lane, S., Brookes, C., Hardy, R., Holden, J., James, T., Kirkby, M., Mcdonald, A., Tayefi, V., & Yu, D. (2003).

686   Land Management , Flooding and Environmental Risk : New Approaches To a Very Old Question.

687   *CIWEM National Conference*, *September*, 20.

688   Laudon, H., Hasselquist, E. M., Peichl, M., Lindgren, K., Sponseller, R., Lidman, F., Kuglerová, L.,

689   Hasselquist, N. J., Bishop, K., Nilsson, M. B., & Ågren, A. (2021). Northern landscapes in transition:
690   Evidence, approach and ways forward using the Krycklan Catchment Study. *Hydrological*
691   *Processes*, *35*(4), 1–15. https://doi.org/10.1002/hyp.14170

692   Laudon, H., Lidberg, W., Sponseller, R. A., Maher Hasselquist, E., Westphal, F., Östlund, L., Sandström,
693   C., Järveoja, J., Peichl, M., & Ågren, A. (2022). Emerging technology can guide ecosystem
694   restoration for future water security. *Hydrological Processes*, *36*(10), 1–5.
695   https://doi.org/10.1002/hyp.14729

696   Laudon, H., Mosquera, V., Eklöf, K., Järveoja, J., Karimi, S., Krasnova, A., Peichl, M., Pinkwart, A., Tong,
697   C. H. M., Wallin, M. B., Zannella, A., & Hasselquist, E. M. (2023). Consequences of rewetting and
698   ditch cleaning on hydrology, water quality and greenhouse gas balance in a drained northern
699   landscape. *Scientific Reports*, *13*(1), 20218. https://doi.org/10.1038/s41598-023-47528-4

700   Liu, H., & Lennartz, B. (2019). Hydraulic properties of peat soils along a bulk density gradient—A meta
701   study. *Hydrological Processes*, *33*(1), 101–114. https://doi.org/10.1002/hyp.13314

702   Loisel, J., & Gallego-Sala, A. (2022). Ecological resilience of restored peatlands to climate change.
703   *Communications Earth and Environment*, *3*(1), 1–8. https://doi.org/10.1038/s43247-022-00547-x

704   Luscombe, D. J. (2014). Understanding the ecohydrology of shallow, drained and marginal blanket
705   peatlands. *Department of Geography*, *PhD*, 180.
706   https://ore.exeter.ac.uk/repository/handle/10871/15967

707   Menberu, M. W., Haghighi, A. T., Ronkanen, A. K., Marttila, H., & Kløve, B. (2018). Effects of Drainage
708   and Subsequent Restoration on Peatland Hydrological Processes at Catchment Scale. *Water*
709   *Resources Research*, *54*(7), 4479–4497. https://doi.org/10.1029/2017WR022362

710   Menberu, M. W., Tahvanainen, T., Marttila, H., Irannezhad, M., Ronkanen, A.-K. A., Penttinen, J., &
711   Kløve, B. (2016). Water-table-dependent hydrological changes following peatland forestry
712   drainage and restoration: Analysis of restoration success. *Water Resources Research*, *52*(5), 3742–
713   3760. https://doi.org/10.1002/2015WR018578

714   Montanarella, L., Jones, R. J. A., & Hiederer, R. (2006). The distribution of peatland in Europe. *Mires and*
715   *Peat*, *1*(1), 1–10. http://hdl.handle.net/1826/3415

716   Noumonvi, K. D., Ågren, A., Ratcliffe, J. L., Öquist, M. G., Ericson, L., Tong, C. H. M., Järveoja, J., Zhu, W.,
717   Osterwalder, S., Peng, H., Erefur, C., Bishop, K., Laudon, H., Nilsson, M. B., & Peichl, M. (2023). The
718   Kulbäcksliden Research Infrastructure: a unique setting for northern peatland studies. *Frontiers in*

Earth Science, *11*(May). https://doi.org/10.3389/feart.2023.1194749

Puttock, A., Graham, H. A., Ashe, J., Luscombe, D. J., & Brazier, R. E. (2021). Beaver dams attenuate flow: A multi-site study. *Hydrological Processes*, *35*(2). https://doi.org/10.1002/hyp.14017

Rosner, B. (1983). Percentage points for a generalized esd many-outlier procedure. *Technometrics*, *25*(2), 165–172. https://doi.org/10.1080/00401706.1983.10487848

Shantz, M. A., & Price, J. S. (2006). Characterization of surface storage and runoff patterns following peatland restoration, Quebec, Canada. *Hydrological Processes*, *20*(18), 3799–3814. https://doi.org/10.1002/hyp.6140

Shuttleworth, E. L., Evans, M. G., Hutchinson, S. M., & Rothwell, J. J. (2015). Peatland restoration: Controls on sediment production and reductions in carbon and pollutant export. *Earth Surface Processes and Landforms*, *40*(4), 459–472. https://doi.org/10.1002/esp.3645

Shuttleworth, E. L., Evans, M. G., Pilkington, M., Spencer, T., Walker, J., Milledge, D., & Allott, T. E. H. (2019). Restoration of blanket peat moorland delays stormflow from hillslopes and reduces peak discharge. *Journal of Hydrology X*, *2*, 100006. https://doi.org/10.1016/j.hydroa.2018.100006

Soomets, E., Lõhmus, A., & Rannap, R. (2023). Restoring functional forested peatlands by combining ditch-blocking and partial cutting: An amphibian perspective. *Ecological Engineering*, *192*(May 2021). https://doi.org/10.1016/j.ecoleng.2023.106968

Wickham, H. (2016). *ggplot: Elegant Graphics for Data Analysis (Version 3.4.2) [R package]*. Springer-Verlag. https://cran.r-project.org/web/packages/ggplot2/index.html

Wilson, L., Wilson, J., Holden, J., Johnstone, I., Armstrong, A., & Morris, M. (2010). Recovery of water tables in Welsh blanket bog after drain blocking: Discharge rates, time scales and the influence of local conditions. *Journal of Hydrology*, *391*(3–4), 377–386. https://doi.org/10.1016/j.jhydrol.2010.07.042

Wilson, L., Wilson, J., Holden, J., Johnstone, I., Armstrong, A., & Morris, M. (2011). The impact of drain blocking on an upland blanket bog during storm and drought events, and the importance of sampling-scale. *Journal of Hydrology*, *404*(3–4), 198–208. https://doi.org/10.1016/j.jhydrol.2011.04.030