Does peatland rewetting mitigate flooding from extreme rainfall events?

3 Shirin Karimi^a, Virginia Mosquera^{a*}, Eliza Maher Hasselquist^a, Järvi Järveoja^a and Hjalmar Laudon^a

^a Swedish University of Agricultural Sciences, Department of Forest Ecology and Management, Umeå,

5 Sweden

6 *Corresponding author: virginia.mosquera@slu.se

7 Abstract

8 Pristine peatlands are believed to play an important role in regulating hydrological extremes because 9 they can act as reservoirs for rainwater and release it gradually during dry periods. Rewetting of drained 10 peatlands has therefore been considered an important strategy to reduce the catastrophic effects of 11 flooding. With the anticipation of more frequent extreme rainfall events in the future due to a changing 12 global climate, the importance of peatland rewetting in flood mitigation becomes even more important. 13 To date, however, empirical data showing that rewetting of drained peatlands actually restores their 14 hydrological function similar to pristine peatlands is largely lacking, particularly for boreal fens. To 15 assess whether peatland rewetting can mitigate flooding from extreme rainfall events and ensure water security in a future climate, we measured event-based runoff responses before and after rewetting 16 17 using a BACI approach (before-after and control-impact) within a replicated, catchment scale study at the Trollberget Experimental Area in northern Sweden. High-resolution hydrological field observations, 18 19 including groundwater level (GWL), discharge, and rainfall data were collected over four years, allowing 20 us to detect and analyze 17 rainfall-runoff events before and 30 events after rewetting. We found that 21 the rewetted sites experienced an increase in the GWL following rewetting and that this was 22 consistently observed across all distances from the blocked ditch within the peatland. Our rainfall-23 runoff analysis revealed that rewetting significantly decreased peak flow, runoff coefficient, and reduced the overall flashiness of hydrographs, making the rewetted sites function more like the pristine 24 25 control peatland. However, "lag time" which was already similar to pristine conditions was pushed farther away from pristine conditions following rewetting. Yet, our results showed that the 26 27 effectiveness of ditch-blocking in flood moderation was strongly influenced by the initial condition and 28 catchment percent of restoration, as one of our two rewetted peatlands did not show significant

- change attributed to it being already similar to the pristine site, suggesting less treatment effect; and
- 30 the other catchment, with higher restoration percentage, had a better response to treatment. In
- summary, our findings suggest that peatland rewetting has the potential to mitigate flood responses,
- 32 however, further research over a longer time period is needed as peat properties and the peatland
- 33 vegetation will develop and change over time.

34 Keywords

Boreal landscape, peatland hydrology, rewetting, flood mitigation, extreme events.

37 Introduction

38 Pristine peatlands are the predominant wetland type in the boreal biome. They encompass 15% of the 39 boreal land surface area and serve as significant carbon sinks and methane sources, playing a crucial 40 role in regulating the global climate (Helbig et al., 2020). In recent years, there has been an increased 41 recognition of the importance of peatlands in carbon capture, flood management, water quality, and biodiversity (Holden et al., 2017). Regrettably, these valuable ecosystems have undergone substantial 42 43 human-induced damages, with more than half of the total pristine peatlands in Europe estimated to have been impacted by drainage for agriculture, forestry, or peat extraction (Andersen et al., 2017). 44 45 Drained peatlands cannot sustain critical ecosystem services, such as climate regulation through long-46 term carbon sequestration or buffering extreme hydrological events, imposing a significant cost on 47 society—a burden that could be alleviated through appropriate rewetting measures (Loisel & Gallego-48 Sala, 2022). Additionally, there are growing concerns surrounding climate change projections for the Northern Hemisphere, indicating an expected increase in more frequent extreme precipitation events, 49 50 along with extended dry periods (Aghakouchak et al., 2020; Hawcroft et al., 2018).

51 Pristine peatlands function as significant water reservoirs, efficiently storing substantial amounts of 52 water during periods of high rainfall (Acreman & Holden, 2013). As extreme rainfall events are 53 anticipated to become more frequent in the evolving global climate, understanding the role of peatland 54 rewetting in flood mitigation is increasingly vital. Rewetting projects typically involve physical interventions such as ditch-blocking or infilling, aiming to increase groundwater level (GWL). Moreover, 55 the blocking of ditches cuts off preferential pathways along open drains, and when combined with 56 57 pooling behind dams, has the potential to act as a buffer during peak flow events, slowing water release 58 and mitigating the flashiness of the discharge response (Holden, 2006; Holden & Burt, 2003). Therefore, 59 by reducing peak flows, peatland rewetting can also contribute to natural flood management by attenuating downstream flow and diminishing flood risk. Furthermore, the reduction of peak flows 60 61 could play an important role in mitigating further erosion of peatlands and minimizing sediment 62 production, as well as carbon loss (Shuttleworth et al., 2015).

In Sweden, peatlands cover approximately 65,600 km² (16% of the Swedish land area) and are predominantly located within boreal regions (Franzen et al., 2012; Montanarella et al., 2006). The historical practice of draining peatlands began in the early 18th century for agricultural purposes and later in the 19th century for forestry, resulting in the excavation of over 1 million km of ditches, primarily dug by hand (Laudon et al., 2022). Consequently, the rewetting of degraded peatlands in Sweden has become a pressing priority to enhance the hydrological functioning of these ecosystems (Bring et al., 2022). As a response, several national programs for peatland rewetting have emerged, with a primary emphasis on reintroducing essential ecosystem services, notably flood control. In a significant move, in
2018, 27 million euros was allocated to facilitate peatland rewetting in Sweden. However, the empirical
underpinning supporting the desired outcomes of peatland rewetting is still largely lacking.

73 The effect of peatland rewetting on hydrological responses during rainfall events has received scientific 74 attention over the past decades (Gatis et al., 2023; Goudarzi et al., 2021; Ketcheson & Price, 2011; 75 Menberu et al., 2018; Shuttleworth et al., 2019). Event-based analysis of stream hydrographs by 76 employing various metrics related to hydrograph magnitude and timing, is a common approach for 77 investigating dominant runoff generation processes in catchments and understanding how quickly 78 water is mobilized from the landscape (Haque et al., 2022; Ketcheson & Price, 2011; Kirchner et al., 79 2023). These response metrics provide valuable insights into catchment storage and release 80 mechanisms (Blume et al., 2007). One widely acknowledged aspect is the impact of rewetting on the 81 event runoff coefficient, which represents the ratio of event runoff depth to event rainfall depth (Evans 82 et al., 1999; Shuttleworth et al., 2019). Therefore, comparing event characteristics before and after 83 rewetting offers a means to understand hydrological processes and runoff generation mechanisms at 84 the catchment scale, thereby improving our understanding of flood estimation during extreme events.

85 A common limitation in the current literature is the predominant focus on event characteristics in 86 natural or relatively un-impacted catchments, with few studies addressing rewetted peatlands. 87 Therefore, the extent of hydrological changes due to rewetting is not well understood. Some studies highlight the positive impact of peatland rewetting on flood moderation with a reduction of peak storm 88 flow (Gatis et al., 2023; Javaheri & Babbar-Sebens, 2014; Lane et al., 2003; Shuttleworth et al., 2019; 89 90 Wilson et al., 2011), however there are differences in the extent of the flood moderation. For example, 91 Gatis et al. (2023) reported a 49% reduction in peak storm flow after rewetting, while Shuttleworth et 92 al. (2019) found a 24% reduction in peak storm flows and a 94% extension in lag times without a change 93 in runoff coefficients. The challenges in understanding the effects of rewetting at the catchment scale 94 are further underscored by the inherent high spatial variability of peatland hydrology and physical 95 characteristics (Evans et al., 1999). The apparent discrepancies in study outcomes, coupled with 96 significant variations among different research sites, highlight the importance of addressing this 97 through further in-depth investigations for developing effective strategies in peatland management, 98 especially given the evolving trend in climate.

Moreover, a recent meta-analysis conducted by Bring et al. (2020) has brought further attention to a noteworthy knowledge gap in understanding the impact of rewetting on GWL changes at different distances from the intervention (i.e. blocked ditches). While existing studies have contributed valuable data on the overall hydrological effects of peatland rewetting, a comprehensive spatial analysis of 103 groundwater changes following rewetting remains inadequately explored. Despite this shortage, some 104 studies suggest that the impact of rewetting, especially through ditch blocking, is localized, resulting in 105 more pronounced GWL rise in close proximity to the ditch (Armstrong et al., 2010; D'Acunha et al., 106 2018; Haapalehto et al., 2014; Wilson et al., 2010). In a prior study (Karimi et al., 2024) in the same 107 catchments as reported on below the overall effect of rewetting on hydrological functioning was found, 108 including a significant rise in GWL post-rewetting. However, a thorough examination of groundwater 109 changes at varying distances from the ditch, considering its important role in discharge regulation, is 110 essential to enhance our mechanistic understanding of flow generation after rewetting. Without such 111 monitoring, the estimation and extrapolation of discharge responses across the landscape become 112 more uncertain. Therefore, a more detailed spatial analysis of GWL changes is crucial for those involved 113 in managing these peatlands.

114 Given that there have been inconsistent reports in the literature on the extent to which rewetted peatlands will affect hydrological functioning, particularly with regards to natural flood management, 115 116 we build on methods used to examine the effect of pristine peatlands on flood attenuation (Karimi et 117 al., 2024) to that of rewetting's impact on hydrological functioning. We used a hydro-climate data set comprised of one-year pre- and three years post-rewetting and incorporate two control catchments to 118 ensure the robustness of our findings. The primary objective of this paper was to test whether peatland 119 120 rewetting has any natural flood management effect. We hypothesized that rewetting leads to a 121 reduction in peak flow, runoff coefficient, Hydrograph Shape Index (HSI), and an increase in lag time, 122 resulting in a generally less flashy hydrograph. Moreover, as GWL is an important indicator of the amount of water stored in the peatland and the effect of the rewetting, we asked how far from the 123 124 ditch GWL was increased by the ditch blocking. We hypothesized that the areas closest to the ditch 125 would increase more than the areas further away from the blocked ditch.

126

128 Materials and methods

129 Study sites

This study took place in the Trollberget Experimental Area (TEA), situated approximately 50 km 130 northwest of Umeå (TEA; 64.181550N, 19.835378E) (Figure 1). The TEA's peatland is an oligotrophic 131 132 minerogenic fen. Prior to rewetting the peatland was dominated by Sphagnum spp., complemented by sparse sedges, dwarf shrubs, and sparse tree canopy (basal area = $2.6 \text{ m}^2 \text{ ha}^{-1}$) of slow-growing Scots 133 pine (Pinus sylvestris). The underlying soils consist mainly of humic podzol, with some drier areas 134 135 featuring Humu-ferric podzol and wetter regions comprising Histosols. Peat depth is on average 2.41 m (Laudon et al., 2023). The bulk density of the drained peatland varied between 0.05 and 0.13 g cm⁻³ 136 137 within the top 55 cm of the peat profile. The bulk density generally increased with distance from the central ditch and with peat depth (Casselgård, 2020). The climate of the area is classified as cold 138 139 temperate humid, characterized by a mean average temperature of 2.4°C and annual precipitation of 140 623 mm (approximately 30% as snow), based on data collected from 1980 to 2020 at the nearby 141 Svartberget Climate Station (Laudon et al., 2021).

142 The peatland at TEA was drained by manual ditch-digging in the early 1920s primarily for forestry 143 purposes, however, because of nutrient limitation the peatland remained unproductive with sparse 144 tree cover (Laudon et al., 2023). The peatland is divided into two catchments draining in two directions, 145 referred to here as R1 and R2, with drainage areas of 33 and 60 ha, respectively (Figure 1). Both 146 catchments are similar in topography and vegetation, however, R1 had an open-water pond shown on 147 historic maps pre-drainage. In the 1930s, the uplands of the peatland were also drained leading to 1152 m of ditches in R1 and 5189 m of ditches in R2 (Laudon et al., 2021). In November 2020, the peatland 148 149 was rewetted by filling and blocking all the ditches in the peatland, whereas ditches in the surrounding 150 none-peat areas were left unmanaged. As a result of these efforts, 59% of the ditches of R1 and 16% 151 of the ditches of R2 were blocked. The ditches were filled using peat from the site with additional dams 152 built at regular intervals using the tree logs harvested from the site. The logs were placed horizontally, 153 except at the two outlet locations where the logs were inserted vertically into the peat and layered 154 additionally with geotextile. To protect the soil characteristics as much as possible, the heavy machinery 155 (i.e., excavators) used moveable log mats while moving on the mire. Additionally, the sparse tree cover 156 that grew on the peatland was cut to reduce evapotranspiration and complement the ditch blocking 157 (Karimi et al., 2024). Finally, after restoration, the open-water pond at R1 re-appeared ca 100 m 158 upstream of the sampling location (Figure 1b).

159 The Degerö Stormyr

160 This study leveraged available data from a nearby natural fen, Degerö Stormyr (273-ha catchment), located approximately 24 km from the TEA, at the Kulbäcksliden Research Infrastructure (KRI, 161 162 Noumonvi et al., 2023) (64.182029N, 19.556543E) to serve as the control for the rewetted peatland R1 163 and R2 catchments. Degerö Stormyr is characterized as an acidic, oligotrophic, minerogenic, mixed mire system. This intensively studied peatland complex exhibits varying vegetation compositions, 164 predominantly featuring Sphagnum moss and sedges. The depth of the peat has an average thickness 165 166 of ~3 m (Noumonvi et al., 2023). The bulk density of the peatland varied between 0.02 to 0.06 g/cm3 167 within the top 34 cm of the peat profile (Fig. 2 in Casselgård, 2020).

168 C4 (Kallkälsmyren)

The second control catchment, C4 (Kallkälsmyren), situated within the Krycklan Catchment Study (KCS) (64.260722N, 19.770339E). C4 is a nutrient-poor, minerogenic fen located approximately 10 km from the rewetted catchment. It encompasses a catchment area of 18 ha, with 40% covered by peatlands and the remainder by forest (Laudon et al., 2021). Similar to TEA, the climate is characterized as a cold temperate humid type with persistent snow cover during the winter season. The peat vegetation cover is dominated by *Sphagnum* spp.

175 Data collection

176 At the TEA, GWL were measured between 2019 and 2023 at an hourly resolution using 30 dipwells. Half 177 of these dipwells were continuously monitored for GWL using data loggers (Levelogger 5, Solinst, Canada), while the remaining were manually measured every two weeks during the snow-free season. 178 179 Dipwells were distributed along five transects. Each transect consisted of six wells with increasing 180 distances of approximately 10, 50 and 100 m from the main ditch (Figure 1). For the Degerö Stormyr 181 control site, GWL data for the corresponding period were obtained from four wells that are part of the 182 ICOS database (www.icos-sweden.se/data). Due to technical issues with the groundwater loggers, no groundwater data for recent years was available for the C4 control catchment. 183

The discharge data at two TEA mire outlets was collected between 2019 and 2023 at an hourly resolution using 90 degree sharp-crested V-notches with connected data loggers for continuous water level measurements (TruTrack WT-HR, Intech Instruments, Australia). Automatic observations were not possible year-round as there was no heating in place, which limited data collection during the winter low flow periods. Manual water level measurements were taken twice a month to calibrate automatic water level data, and stage-discharge relationships were defined using manual flow gauging. Specific discharge (discharge per unit catchment area) was calculated using catchment areas derived from the Deterministic 8 (D8) algorithm based on a 2×2 m resolution DEM in which we first burned the ditches into the digital elevation model (DEM) to the depth of 0.5 m (Whitebox GAT 3.3) (Laudon et al., 2021). For this study, we only utilized discharge data from the C4 control site due to its proximity to the rewetted site. At C4, the outlet is equipped with a similar V-notch weir situated within a heated dam house, facilitating continuous stage height monitoring year-round. Discharge measurements and calibrations followed the same protocol and interval as those implemented at TEA (Laudon et al., 2021).

197 Rainfall data were acquired from a reference climate station at Svartberget Research Station 198 (64.244376N, 19.766378E, 225 m.a.s.l) (Laudon et al., 2021). Rainfall measurements were logged every 199 10 minutes using a tipping-bucket rain gauge (ARG 100, EML, UK). The climate station is integral to the 200 reference climate monitoring program at Vindeln experimental forests, adhering to the WMO standard 201 for meteorological measurements (Karlsen et al., 2019).

302 GWL analysis

203 First, the hourly groundwater data were examined for outliers, and any gaps were filled using the 204 Generalized Extreme Studentized Deviate (ESD) filter (Rosner, 1983). The algorithm processes a time-205 series dataset by calculating a rolling mean and standard deviation with a window size of 6 hours. 206 Outliers were identified by comparing each data point to the moving average, and values exceeding the 207 3-standard deviation threshold were identified as outliers and subsequently removed from the dataset. Subsequently, the data were gap-filled using the Spline interpolation method, an advanced form of 208 209 interpolation that utilizes piecewise polynomial functions to estimate data between two known points. 210 The data were aggregated to daily time scales. For our analysis we used the GWL data from 1st of June 211 to the end of October as our study focused on rainfall events; outside this period, precipitation often 212 occurs as snow and dipwells could be frozen. For each catchment R1 and R2, the GWL data were averaged, and pairwise comparisons test were conducted to assess if there were any significant 213 214 differences between pre-rewetting and multiple post-rewetting years. As the data were not normally 215 distributed and we were interested in the distribution of the data and not the means, the non-216 parametric Wilcoxon tests were used. Then, a Bonferroni-Holm correction was applied to adjust for 217 multiple comparisons. The differences were considered significant when p<0.05. Moreover, to examine 218 the impact of rewetting on GWL at all distances from the main ditch, data were disaggregated based on distances of 10, 50, and 100 m to the main ditch. It is noteworthy that additional side ditches existed 219 220 in the proximity of some of the dipwells, indicating a potential limitation in the study design.

221 Rainfall-runoff events detection

222 As a first step, we segmented the 2020–2023 summer-autumn precipitation record into distinct rainfall events using the inter-event time definition (IETD) via the IETD R package (Duque, 2022). The IETD 223 224 establishes a minimum dry period between independent rainfall events as a criterion for grouping them. To distinguish independent rainfall events from continuous precipitation, we set a minimum 225 threshold of 0.1 mm h⁻¹ at the start of an event. Events were considered distinct if they were separated 226 by at least 12 hours without rainfall. The methodology for identifying runoff events was based on the 227 228 framework outlined by Luscombe (2014) and was further adapted to the specific characteristics of our 229 study area. Runoff events were defined as periods during which the observed discharge exhibited significant deviations from the baseflow. Rainfall events were matched with the runoff events that 230 231 followed within a specified time window (12.5 h). We calculated rolling quantiles for this time window 232 (12.5 h) at the 30th and 95th percentile (Q30th and Q95th respectively). A rolling quantile for the 70th percentile for a one month period was also calculated (MQ90). Where (Q95th – Q30th) > MQ90, the 233 234 flow was considered to be elevated and any fluctuation in flow was driven by precipitation; therefore 235 measured discharge was used (Gatis et al., 2023; Puttock et al., 2021). A final, visual inspection of the 236 time series with detected events was used to quality control these data and ensure that all significant 237 rainfall and flow events were extracted from the dataset.

238 Flood mitigation effects

239 To evaluate the flood mitigation effect of peatland rewetting and determine its impact, we employed a 240 set of response metrics to characterize hydrologic responses during events following the rewetting process. These response metrics included event duration, rainfall volume, peak flow, runoff coefficient, 241 lag time, and HSI. We calculated these response metrics for both the rewetted and control sites. The 242 selection of these response metrics was based on their widespread use in hydrological comparison 243 studies (Edokpa et al., 2022; Wilson et al., 2011). Peak flow response was computed as the maximum 244 discharge observed during each event. Runoff coefficient was determined as the ratio of total event 245 runoff to total event rainfall. Lag time calculated as the time between peak rainfall and peak discharge 246 247 in each event. HSI, defined as the ratio of peak storm discharge to total storm discharge, was used to provide a straightforward measure of the overall hydrograph shape (Shuttleworth et al., 2019). The 248 response metrics for the rewetted catchment R2 and the control site were derived using the start and 249 250 end times of rainfall-runoff events identified at R1 catchment.

251 Statistical analyses

252 The statistical design used in this study focuses on the BACI approach (before-after and control-impact) 253 as used previously in hydrological studies (Holden et al., 2017; Laudon et al., 2023; Menberu et al., 254 2018; Shuttleworth et al., 2019). We standardized the response metrics derived from the two 255 catchments (R1 and R2) of the rewetted site against the control catchment (treatment minus control) 256 to distinguish responses resulting from rewetting treatment from natural variation, changes over time 257 and seasons. Due to variations in the frequency of events between the pre- and post-rewetting periods, 258 and the non-normal distribution of response metrics, a non-parametric test was employed. Specifically, 259 the Wilcoxon test was conducted to investigate statistically significant changes in the distribution of 260 data for each catchment (R1 and R2) of the rewetted site before and after rewetting, with a focus on 261 understanding the extremes, rather than solely examining means (Shuttleworth et al., 2019). 262 Significance was determined at p < 0.05. Additionally, we aggregated all years post-rewetting together due to the highly variable number of events occurring during each year post-rewetting. Statistical 263 264 analysis was undertaken in R version 4.1.2. (R Core Team, 2021) with data processing, summary 265 statistics and plotting undertaken using the R package Tidyverse (Wickham, 2016).

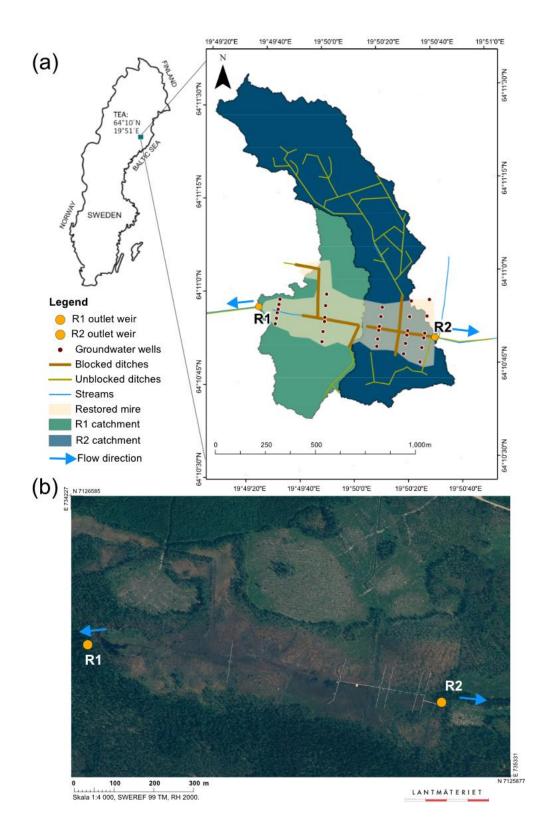


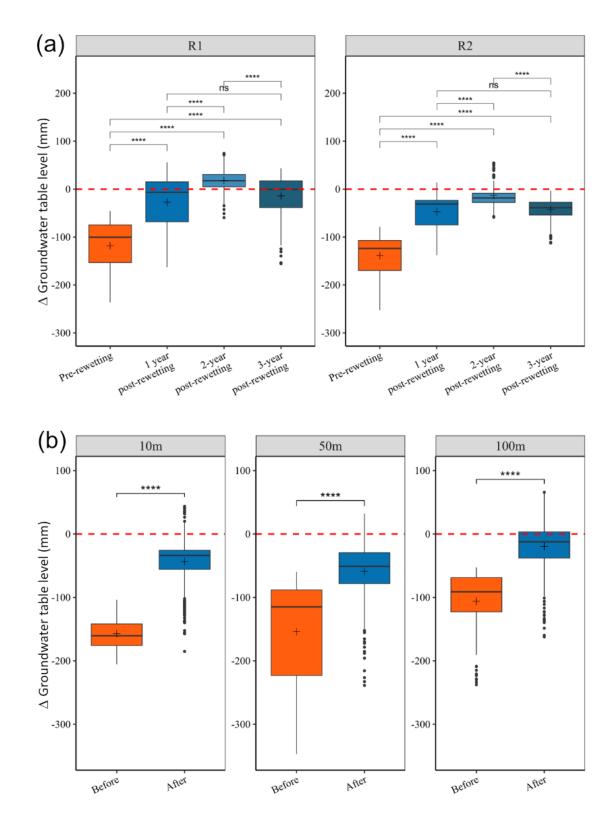


Figure 1. Trollberget Experimental Area (TEA) catchments with monitoring locations (A). Yellow circles show the locations of the outlets of the catchment areas for R1 and R2 (weir locations) of the rewetted peatland. Red small circles designate groundwater dipwells. Note that in R1, 59% of the ditches were blocked, while in R2, only 16% of the ditches were blocked. Aerial view of the rewetted peatland with boardwalks visible as white lines (B). (Aerial map from Lantmäteriet)

271 **Results**

272 The impact of rewetting on GWL variation

273 Peatland rewetting led to a significant increase in GWL at the two rewetted catchments (R1 and R2) 274 compared to the control site (Figure 2a). The relative difference in GWL between the rewetted and 275 control sites (treatment minus control) at varying distances to the ditch also showed a significant 276 decrease after rewetting (Figure 2b). Interestingly, this impact demonstrated variability depending on 277 the distance from the ditch, with wells located closest to the ditch showing a more pronounced 278 response compared to those farther away. Prior to rewetting, the median GWL was significantly 279 (p<0.05) lower next to the ditch (-228 mm) compared to the furthest distance away (-174 mm). 280 Furthermore, GWL exhibited greater variability in the middle of the transect (50 m from the ditch), 281 reaching a minimum of 507 mm from the ground. After rewetting, the largest median GWL change was 282 observed at a distance of 10 meters, with an increase of 119 mm. This was followed by a median 91 283 mm increase at a distance of 100 meters and a median 62 mm increase at a distance of 50 meters. The 284 median GWL at the control sites was similar during the pre- and post-rewetting periods (-79 and -78 285 mm, respectively). Finally, GWL in the three different distances from ditch significantly increased in the 286 first year after rewetting, significantly increased further in the second year and significantly decreased 287 in the third year, however, still being significantly higher than the pre-rewetting and first year post-288 rewetting GWL.



290 Figure 2.a) Difference (treatment-control) in groundwater table level (GWL) at the two rewetted catchments (R1 and R2) based 291 on daily data gathered between June to October in the years 2020 (pre-rewetting) and 2021, 2022 and 2023 (3 years post-292 rewetting) regardless of distance to ditch. b) Relative difference in GWL based on varying distances to the main ditch (i.e. 10, 293 50 and 100 m); all years post-rewetting are combined (sample sizes for pre-rewetting and post-rewetting were 153 and 428, 294 respectively). The red dashed line indicates the GWL of the control site; positive values indicate that the GWL is higher at the 295 rewetted site than at the control, while negative GWL indicate the opposite. The box plots show the minimum, first quartile, 296 median, third quartile, and maximum, with outliers as dots. The stars indicate the levels of significant difference between the 297 marked comparisons as determined using a Wilcoxon test (**** $p \le 0.0001$).

Table 1. Median, minimum (min), maximum (max) and 5th-95th quantile of groundwater level (GWL) change pre- and postrewetting for different distances to the ditch (i.e. 10, 50 and 100 m) and the control site.

	Distance	Median (mm)	Min (mm)	Max (mm)	5th-95th quantile (mm)
PRE-REWETTING	10 m	-228	-364	-120	194
	50 m	-190	-507	-60	370
	100 m	-174	-416	-44	304
	Control	-79	-186	9	156
POST-REWETTING	10 m	-108	-272	-33	197
	50 m	-127	-366	-30	233
	100 m	-83	-341	5	240
	Control	-78	-234	3	171

300 The impact of rewetting on runoff responses

Based on the response at R1, 17 rainfall-runoff events before and 30 events after rewetting were 301 302 extracted and analyzed (Figure 3). The impact of rewetting on runoff responses during rainfall-runoff events is depicted through examples of event-scale hydrographs (Figure 4, Table S 1). The illustrations 303 304 display the variation in discharge response across control and the two rewetted catchments (R1 and R2) for different event sizes and antecedent GWL conditions, during both pre-and post-rewetting 305 periods. In the pre-rewetting period, despite the control site having the shallowest GWL at -15 mm, 306 307 exhibited the smallest peak flow of 0.29 mm h⁻¹. In contrast, the rewetted site R1, with an antecedent GWL of -82 mm h⁻¹, reached a peak of 0.93 mm h⁻¹. One and two years after rewetting, R1 still had the 308 highest peak flows at 0.71 mm h⁻¹ and 0.61 mm h⁻¹, respectively, while the rewetted catchment R2 309 310 showed similarities to the control site (Figure 2). However, three years after rewetting, although R1 had the shallowest antecedent GWL at -5.15 mm, the peak flow was almost half of the peak in the control 311 catchment (0.14 mm h^{-1} and 0.26 mm h^{-1} , respectively). 312

313

314

315

316

317

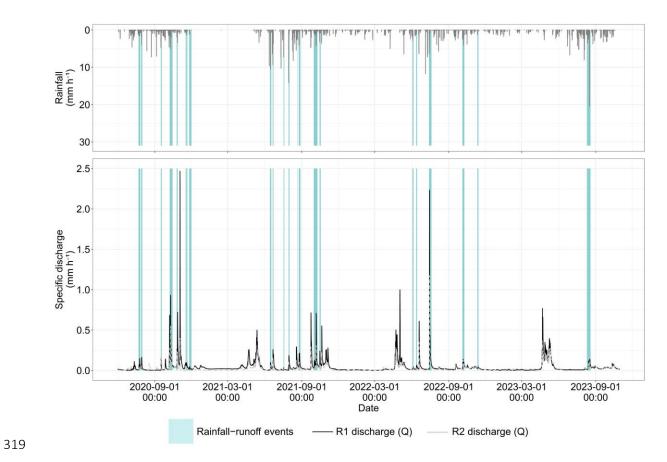
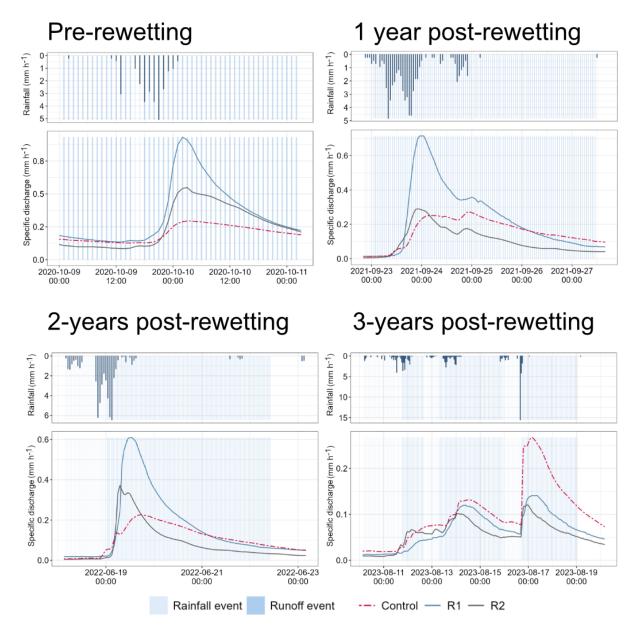


Figure 3. Identified rainfall-runoff events using discharge measured at the rewetted catchment R1 across the entire study period
 (black line). Same rainfall-runoff events identified using R1 discharge are shown for R2 (grey line).



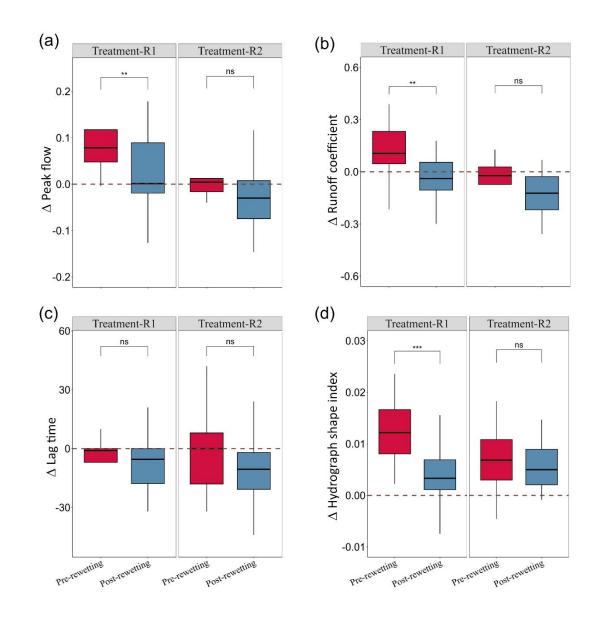
330 Figure 4. Examples of runoff responses of control and the two catchments (R1 and R2) of the rewetted site during rainfall-runoff

events for each of the four pre- and post-rewetting years. Note that the scales for the y-axes show different magnitudes of specific discharge.

334 Flood mitigation effects of rewetting

The magnitude of the effects of peatland rewetting was investigated for 47 rainfall-runoff events (17 events before rewetting and 30 events after rewetting) to test if the rewetting's effects were significant under a larger number of events. Storm magnitudes ranged between 5 and 50 mm in total precipitation before rewetting, and 2 and 63 mm after rewetting. The relative differences between the two catchments (R1 and R2) of the rewetted site and control sites (rewetted minus control) for each metric are shown in Figure 5.

- The analysis of rainfall-runoff events revealed a reduction in relative peak flow at the two catchments (R1 and R2) of the rewetted site following rewetting (Figure 5a). However, the reduction was significant only at R1. Specifically, the median peak flow at R1 decreased from 0.14 to 0.10 mm h⁻¹ post-rewetting. In contrast, at R2, there was an increase from 0.04 to 0.08 mm/h post-rewetting. Interestingly, the control site experienced a rise in median peak flow from 0.05 to 0.12 mm h⁻¹ during the post-rewetting period.
- The median runoff coefficient in the two catchments of the rewetted site showed an increase from 0.36 to 0.40 and from 0.14 to 0.20 at R1 and R2, respectively, after rewetting. The runoff coefficient at the control site increased from 0.17 before rewetting to 0.40 after rewetting. Relative to the control site, both restored sites, R1 and R2, experienced a decline in runoff coefficients during the post-rewetting phase. Notably, this reduction was statistically significant solely at R1 (p<0.01) (Figure 5b).
- After rewetting, the median lag time in the two catchments (R1 and R2) of the rewetted site decreased by 0.5 and 7 hours, reaching 15 and 10 hours for R1 and R2, respectively. In contrast, the control catchment exhibited an increase in median lag time from 14 to 23 hours during the post-rewetting period. However, pairwise test results indicated that there was no statistically significant change at both rewetted catchments (R1 and R2) following rewetting (Figure 5c).
- The median HSI values for both catchments (R1 and R2) of the rewetted site and control sites decreased after the rewetting period, shifting from 0.023 to 0.021, 0.034 to 0.025, and 0.027 to 0.026 at control, R1, and R2, respectively (Fig. 5d). The effect of rewetting in reducing HSI was significant only at R1 (p < 0.0001). Prior to rewetting, the relative HSI at R1 was 0.012, and after rewetting, it decreased to 0.003. The relative HSI also experienced a decline at R2, dropping from 0.006 pre-rewetting to 0.004 after rewetting. However, this decrease was not statistically significant.



363

Figure 5.Differences between the rewetted and control sites pre- and the combined three years of post- rewetting period for (a) peak flow, (b) runoff coefficient, (c) lag time, and (d) Hydrograph shape index. The relative difference was computed as treatment minus control and the red dashed line indicates the value of the control site; thus, positive values indicate that the hydrological response is greater at the treatment site than at the control site, while negative values indicate the opposite. The box plots show the minimum, first quartile, median, third quartile, and maximum, with outliers as points. The stars indicate the

370 levels of significance in Wilcoxon test (** $p \le 0.01$; *** $p \le 0.001$; "ns" denotes not significant.).

372 Discussion

373 Despite significant interest in peatland rewetting, there has been limited research on its effects on 374 hydrological functioning and the scale of these impacts. We found that peatland rewetting on nutrient-375 poor minerogenic fens, one of the most common peatland types in Fennoscandia, was generally 376 positive for use as natural flood management. Rewetting has begun to influence GWL, runoff responses during rainstorms, and flood mitigation (though the latter was observed in only one of the two study 377 378 catchments), while also shifting these hydrological characteristics closer to pristine conditions by increasing water storage in the peatland. However, special attention should be given to the diverse 379 380 characteristics of peatlands in the boreal biome before generalizing the effect of peatland rewetting on 381 hydrological functioning.

382 The impact of rewetting on GWL

383 Using the BACI experimental approach, we found that the mean GWL of the rewetted sites rose rather rapidly after ditch-blocking at both R1 and R2 to the near pristine levels of our control site. Our results 384 385 align broadly with several other studies mainly from Finland, Canada and UK that found that peatland 386 rewetting raised GWL to near pristine levels (Armstrong et al., 2022; Dixon et al., 2014; Haapalehto et 387 al., 2014; Howie et al., 2009; Menberu et al., 2016; Shuttleworth et al., 2019; Soomets et al., 2023). 388 Notably, GWL rose in all studies, despite variation in the extent of recovery, climate and drainage 389 conditions, such as age, depth, and pattern of ditching. Nevertheless, peatland type is likely to have a 390 major impact on the time taken to increase GWL to pristine-like levels after ditch blocking. For example, 391 fen peatlands, such as our system, with a rather flat or slightly domed surface and slow lateral 392 movement of water could have a faster response to ditch blocking compared to blanket bogs which 393 may exist even on slopes of 20 to 25 degrees (Laine et al., 2011). Furthermore, our results also revealed 394 that the median GWL at R1 closely resembled that of the control site after rewetting. However, at R2, 395 the median GWL remained slightly lower post-rewetting. This difference may be attributed to the 396 presence of shrubs and sparse tree cover (higher water uptake) on the mire at R2, as well as a lower 397 proportion of blocked ditches within the catchment. However, as found by Howie et al. (2009) in Southwestern British Columbia, the difference between the effect of rewetting on GWL between the 398 399 two catchments could also be attributed to differences in the drying of the peatland. This, coupled with shrinkage and subsidence of the peat, could lead to a reduction in hydraulic conductivity, possibly 400 401 hindering the effectiveness of restoration efforts in reversing the impacts of drainage.

402 Our results addressed a gap in the existing literature by examining the spatial variability of GWL 403 recovery at different distances from the ditch, a factor largely neglected in prior research, particularly 404 within the context of boreal ecosystems (Bring et al., 2022). After rewetting, our results show a 405 significant increase in GWL at all distances from the ditch, however with variation spatially. These 406 detailed spatial results of GWL increase at different distances to the ditches, show that all of the 407 locations in the mire had undergone rewetting and any observed differences in event runoff responses 408 could be attributed to changes in GWL and water storage within the full extent of the peatland. 409 Furthermore, we found that after rewetting, significant differences persisted between locations, with the highest GWL found at the furthest distance (i.e., 100 m), contrary to most other published studies 410 411 that have found that the impact of rewetting on GWL diminishes with increasing distance from the main 412 ditch (Bring et al., 2022). This difference in GWL between distances from ditch might be a topographic 413 effect, where peat surfaces could be at higher elevations at the further distance due to long-term 414 consolidation of the peat near ditches and therefore the absolute GWL was also higher at the further 415 distance (Holden et al., 2017). Either way, our results showed that rewetting was successful in returning 416 the pristine-like patterns of higher GWL, however, restoration must also intend to reach an even GWL 417 throughout the mire which we expect to see with time (Haapalehto et al., 2014; Laine et al., 2011). Furthermore, our data serves as a valuable resource for peatland managers, especially in boreal 418 ecosystems, helping to gain a better understanding of site-specific changes in hydrology and the 419 420 associated ecosystem services that result from the rewetting of peatlands.

421 The impact of rewetting on runoff responses

Event-based analysis of discharge responses provides important information on treatment effects on 422 423 the hydrological functionality. Relying solely on daily discharge analysis does not offer much insight into 424 discharge responses to precipitation, including the lag time to peak flow. For instance, examining the 425 hourly hydrograph revealed that, discharge responses at R1 exhibited flashier characteristics with higher peaks compared to those at R2, yet R2 did not show a significant difference in the overall 426 427 decrease in relative peak flow after rewetting. This discrepancy could possibly be attributed to the 428 smaller increase of GWL of R2 compared to R1, likely influenced by a lower proportion of blocked 429 ditches in the catchment, the re-creation of the open-water pond in R1, or the fact that the relative 430 peak flow at R2 before rewetting was already similar to the control peatland. Specifically, the re-431 creation of the old open-water pond ca. 100 meters from R1 outlet after rewetting, could potentially be functioning as a "natural peatland pool", hence influencing runoff responses (Arsenault et al., 2019). 432 433 Nonetheless, three years after ditch blocking, both catchments showed attenuated hydrograph shapes 434 during most storm events. These changes in the hydrograph characteristics following rewetting indicate that hydrological restoration positively affects overall flow regimes influencing flow pathways and 435 436 water storage within the peat, leading to reduced peak flow and increased hydrological residence time

in the peatland (Gatis et al., 2023). However, the scarcity of continuous, prolonged datasets from
rewetted peatlands, particularly for boreal minerogenic fens, poses a significant challenge in
conducting comprehensive comparisons across various peatland sizes, types, and rewetting durations,
as most rewetting projects have only recently commenced. Therefore, a more extended period of postrewetting monitoring is necessary to fully understand how the discharge patterns of drained peatlands
evolve after rewetting.

443 Flood mitigation effects of rewetting

444 Rewetting resulted in a significant reduction in event peak flow response at R1. It is noteworthy that, 445 although there was a significant decrease, the median relative peak flow at R1 after rewetting was still 446 higher than the control. This is in contrast to R2, where the decrease in peak flow was not significant, 447 but where the median peak flow started out similar to the control mire, but decreased to be somewhat 448 lower than the control after rewetting. Our findings align with the results observed by Wilson et al. (2011), where they showed peak flow hydrographs from ditches with considerable change after 449 rewetting, with lower peak flow rates, less runoff and rainwater being released during events. In 450 451 contrast, Shantz and Price (2006) evaluated the hydrological characteristics of a restored peatland in 452 Quebec, Canada, and observed higher discharge peaks during summer at the restored site compared 453 to the control site. They attributed this to wetter antecedent conditions and faster drainage response 454 following rainfall. Conversely, our research reveals that despite observing a rise in GWL after rewetting, 455 rewetted peatlands can exhibit less intense flood responses and offer improved retention of rainfall. 456 Moreover, although reduced runoff rate following restoration inevitably increases the chances of 457 overland flow, this pathway is considerably slower than flow through the drainage networks and 458 therefore restoration may lead to reduction in peak flow magnitude (Grand-Clement et al., 2013). Here, 459 the old-open water pond, functioning as a peatland pool, could potentially increase runoff detention and therefore reduced peak flow (Arsenault et al., 2019; Kløve, 2000). Finally, it is noteworthy that even 460 461 before rewetting, our system was already dominated by *Sphagnum* spp., therefore compared to other 462 degraded (especially bare) peatlands, our system had a higher potential for reducing sheet erosion and 463 downstream flood peaks after rewetting (Holden et al., 2008). Either way, our results suggest that 464 contrary to conclusions drawn in many previous studies (Holden et al., 2004; Holden & Burt, 2003), the 465 rewetted peatlands in our study exhibited more controlled and resilient hydrological behavior 466 delivering "natural flood management" by attenuating downstream flow and reducing flood risk. Yet, 467 our results could be affected by the short time period since rewetting. With time, the peatland might 468 respond differently depending on the type of newly established vegetation and initial moisture

469 conditions, where rainfall events could trigger rapid and concentrated runoff and discharge (Holden &470 Burt, 2003).

471 Runoff coefficient is another key indicator for flood mitigation and corresponds to catchment storage 472 capacity. Specifically, a reduction in runoff coefficient indicates a gain in storage capacity either behind the dam in the blocked ditches or due to increased surface roughness which then reduces the drainage 473 efficiency, thereby increasing water storage capacity (Menberu et al., 2016). Our results showed that 474 475 reduction in runoff coefficient was significant at R1, showing less runoff being exported with rainfall 476 events after rewetting. Yet again, this reduction was not significant at R2, and the same patterns 477 follows, where the initial condition of the runoff coefficient of R2 was lower than the control mire 478 before rewetting and after rewetting the coefficient lowered even further. In line with our results, the 479 reduction in runoff coefficient after peatland rewetting has been reported in many studies (Ketcheson 480 & Price, 2011; Shantz & Price, 2006; Wilson et al., 2011), in some, to the extent of being the most 481 significant hydrological effect of peatland rewetting (Ketcheson & Price, 2011). However, caution is 482 needed in interpreting some of these results due to the potential influence of the relatively short time series during which the peatland could have been still undergoing filling (Ketcheson & Price, 2011) or 483 484 an eventual increase in the runoff coefficient due to a declining efficiency of the ditch blocking 485 (Menberu et al., 2018).

486 An increasing lag time traditionally serves as a positive indicator for flood modification, as downstream 487 flow becomes less "flashy". Contrary to expectations, the lag time between the initiation of a rainfall 488 event and the peak discharge decreased after rewetting in both rewetted catchments. However, it is 489 important to note that this decrease, while observed, was not statistically significant for either 490 catchment. A shorter lag time may support the interpretation that flow peaks observed after rewetting 491 originated from the near vicinity of the monitoring site, while water from upstream areas was 492 attenuated before reaching the outlet (Gatis et al., 2023). These explanations could be further 493 supported by coupling these results with the positive effect on base flow that has been seen in our 494 study site (Karimi et al., 2024). Finally, HSI, which serves as a direct indicator of system flashiness, 495 exhibited a notable decrease at the R1 catchment following rewetting. Unlike the pattern observed in 496 R2's response to other hydrological factors, the HSI at R2 was higher than that of the control site. 497 Although R2 showed a non-significant reduction after rewetting, it remained consistently higher than the control. This reduction seems to be a combined effect of ditch blocking and presence of vegetation 498 499 as other studies have shown that ditch blocking by itself does not appear to alter the "flashiness" of 500 stormflow (Shuttleworth et al., 2019).

501 The significant decreases in peak flow, runoff coefficient and HSI observed at R1, compared to the non-502 significant changes at R2, can be attributed to several factors. Firstly, the BACI analysis indicated that, 503 prior to rewetting, R1 had flashier hydrological responses. Moreover, R2's responses were already more 504 similar to the control site, suggesting less potential to observe a significant post-rewetting effect. 505 Additionally, a smaller portion of R2 catchment was restored compared to R1, which could mean that 506 the overall water storage at R2 remained lower than at R1. Consequently, water may still drain more 507 quickly at R2, leading to less noticeable impacts from the rewetting efforts. Furthermore, as mentioned 508 above, the differences in responses could in part be due to the recreation of the peatland pool that 509 likely influences runoff detention. All in all, the diverse responses observed in flood response 510 characteristics, both in our study and other investigations, raises questions regarding the overall 511 effectiveness of peatland rewetting, at least in the short term. While it appears successful in reducing 512 peak flow, runoff coefficient, and overall flashiness of hydrographs (as shown by HSI), our results 513 suggests it might not be as effective in increasing lag time from peak rainfall to peak flow occurrence. 514 This limitation could potentially be attributed to the need for new peat formation before becoming fully 515 hydrologically restored. Hence, a crucial question regarding the duration of these effects and the time 516 necessary for lag time recovery remains unanswered.

517 The effectiveness of ditch-blocking in flood moderation is influenced by various factors, including the 518 initial condition of a drained peatland, the extent of peat degradation and restoration, and changes in 519 its properties (Menberu et al., 2016). Furthermore, there may be a delayed effect in the peatland's 520 response to ditch-blocking, and the corresponding flood mitigation may progressively change over time 521 in the years following the blocking of ditches due to changes in peat properties and vegetation cover. 522 Overall, our rewetted sites, having been drained for a century, still may not function as a natural 523 peatland and a full hydrological recovery will take substantially longer than the recovery period we 524 measured here. Several factors linked to prolonged drainage could contribute to a long recovery period. 525 For instance, peat oxidation and compaction may lead to increased bulk density, which in turn, affects 526 the ability of the site to effectively retain and release water (Liu & Lennartz, 2019). Hence, our three-527 year post-rewetting monitoring period, while longer than many other studies, still offers rather limited 528 insight into the impact of rewetting on flood moderation under extreme storm events, especially in 529 more severe future climatic conditions. Therefore, further monitoring is required to understand the 530 influence of restoration practices on peatland hydrological functioning.

531 Conclusion

532 Our results showed that the effect of rewetting on flow moderation from rainfall events is not as simple 533 as restoring GWL. This gradual and evolving process of peatland hydrological functioning due to a long 534 history of peat compaction and decomposition, then the subsequent re-establishment of peat-forming 535 vegetation after rewetting emphasizes the importance of sustained long-term monitoring to fully understand the outcomes of rewetting. Moreover, our findings indicate that peatland rewetting has 536 537 the potential for flood mitigation and, in some cases, the ability to mitigate runoff from rainfall events better than pristine sites. This was supported by reductions in peak flow, runoff coefficient, and less 538 539 flashy hydrograph responses (HSI). However, the results showed that peatland rewetting would not 540 necessarily increase the lag time between the peak of a rainfall event and peak discharge. Significant 541 changes were only observed at one of the two restored peatlands. These differences seem to be 542 attributed to (1) a higher percentage of the ditches in the catchment being restored and (2) that the none-effected site already being similar to the pristine site, suggesting less potential changes post-543 544 rewetting. Nevertheless, uncertainties remain in our understanding of the contribution of peatland rewetting to natural flood management over longer timescales or during large historical flood events. 545 Therefore, we emphasize the significance of long-term monitoring combined with hydrological 546 547 modeling to determine whether peatlands will consistently mitigate floods as climate change 548 intensifies.

549

551 Code and data availability

All data used in this study are freely available. The discharge data can be obtained from https://data.fieldsites.se/portal/. The GWL data up to October 2023 are available from the corresponding author. The original R codes for extracting rainfall-runoff events are available from Gatis et al. (2023) at https://ore.exeter.ac.uk/repository/handle/10871/134028.

556 Financial support

557 The TEA infrastructure was initiated and co-funded by the European Union GRIP on LIFE IP project 558 (LIFE16IPE SE009 GRIP) led by the Västerbotten Administration Board and Swedish Forest Agency, with 559 additional financial infrastructure and research support from The Kempe Foundation and the Swedish 560 Research Council Formas grants (2018-00723 (to EMH), 2018-02780 (to HL), 2020-01372 (to HL), 2021-561 02114 (to HL), as well as by the Knut and Alice Wallenberg (Grants 2018.0259 and 2023.0245). The 562 KCS/KFI infrastructure and long-term data collection have been funded by the Swedish Research 563 Council VR (SITES, grant number 2021-00164).

564 Acknowledgements

565 We would like to thank all the skilled and dedicated field personnel at the Svartberget research station.

566 Competing interests

567 The authors declare that they have no conflict of interest.

569 **References**

- 570 Acreman, M., & Holden, J. (2013). How wetlands affect floods. *Wetlands*, *33*(5), 773–786.
 571 https://doi.org/10.1007/s13157-013-0473-2
- Aghakouchak, A., Chiang, F., Huning, L. S., Love, C. A., Mallakpour, I., Mazdiyasni, O., Moftakhari, H.,
 Papalexiou, S. M., Ragno, E., & Sadegh, M. (2020). Climate Extremes and Compound Hazards in a
 Warming World. *Annual Review of Earth and Planetary Sciences*, 48, 519–548.
 https://doi.org/10.1146/annurev-earth-071719-055228
- Andersen, R., Farrell, C., Graf, M., Muller, F., Calvar, E., Frankard, P., Caporn, S., & Anderson, P. (2017).
 An overview of the progress and challenges of peatland restoration in Western Europe. *Restoration Ecology*, 25(2), 271–282. https://doi.org/10.1111/rec.12415
- Armstrong, A., Holden, J., Kay, P., Francis, B., Foulger, M., Gledhill, S., McDonald, A. T., & Walker, A.
 (2010). The impact of peatland drain-blocking on dissolved organic carbon loss and discolouration
 of water; results from a national survey. *Journal of Hydrology*, *381*(1–2), 112–120.
 https://doi.org/10.1016/j.jhydrol.2009.11.031
- Armstrong, L., Peralta, A., Krauss, K. W., Cormier, N., Moss, R. F., Soderholm, E., McCall, A., Pickens, C.,
 & Ardón, M. (2022). Hydrologic Restoration Decreases Greenhouse Gas Emissions from Shrub Bog
 Peatlands in Southeastern US. *Wetlands*, 42(7), 81. https://doi.org/10.1007/s13157-022-01605-y
- Arsenault, J., Talbot, J., Moore, T. R., Beauvais, M. P., Franssen, J., & Roulet, N. T. (2019). The Spatial
 Heterogeneity of Vegetation, Hydrology and Water Chemistry in a Peatland with Open-Water
 Pools. *Ecosystems*, 22(6), 1352–1367. https://doi.org/10.1007/s10021-019-00342-4
- Blume, T., Zehe, E., & Bronstert, A. (2007). Rainfall-runoff response, event-based runoff coefficients and
 hydrograph separation. *Hydrological Sciences Journal*, 52(5), 843–862.
 https://doi.org/10.1623/hysj.52.5.843
- Bring, A., Rosén, L., Thorslund, J., Tonderski, K., Åberg, C., Envall, I., & Laudon, H. (2020). Groundwater
 storage effects from restoring, constructing or draining wetlands in temperate and boreal
 climates: a systematic review protocol. *Environmental Evidence*, 9(1), 1–11.
 https://doi.org/10.1186/s13750-020-00209-5
- Bring, A., Thorslund, J., Rosén, L., Tonderski, K., Åberg, C., Envall, I., & Laudon, H. (2022). Effects on
 groundwater storage of restoring, constructing or draining wetlands in temperate and boreal
 climates: a systematic review. *Environmental Evidence*, *11*(1), 38.

https://doi.org/10.1186/s13750-022-00289-5

- Casselgård, M. (2020). Effects of 100 years of drainage on peat properties in a drained peatland forests
 in northern Sweden Examensarbeten 2020:4 Fakulteten för skogsvetenskap Institutionen för
 skogens ekologi och skötsel. https://stud.epsilon.slu.se
- D'Acunha, B., Lee, S. C., & Johnson, M. S. (2018). Ecohydrological responses to rewetting of a highly
 impacted raised bog ecosystem. *Ecohydrology*, *11*(1), 1–12. https://doi.org/10.1002/eco.1922
- Dixon, S. D., Qassim, S. M., Rowson, J. G., Worrall, F., Evans, M. G., Boothroyd, I. M., & Bonn, A. (2014).
 Restoration effects on water table depths and CO2 fluxes from climatically marginal blanket bog.
 Biogeochemistry, *118*(1–3), 159–176. https://doi.org/10.1007/s10533-013-9915-4
- 608 Duque, M. L. F. (2022). *R package IETD* (pp. 1–9). https://doi.org/10.3390/w6010045>.Restrepo-Posada
- Edokpa, D., Milledge, D., Allott, T., Holden, J., Shuttleworth, E., Kay, M., Johnston, A., Millin-Chalabi, G.,
 Scott-Campbell, M., Chandler, D., Freestone, J., & Evans, M. (2022). Rainfall intensity and
 catchment size control storm runoff in a gullied blanket peatland. *Journal of Hydrology*, *609*(October 2021), 127688. https://doi.org/10.1016/j.jhydrol.2022.127688
- Evans, M. G., Burt, T. P., Holden, J., & Adamson, J. K. (1999). Runoff generation and water table
 fluctuations in blanket peat: Evidence from UK data spanning the dry summer of 1995. *Journal of Hydrology*, 221(3–4), 141–160. https://doi.org/10.1016/S0022-1694(99)00085-2
- Franzen, L. G., Lindberg, F., Viklander, V., & Walther, A. (2012). The potential peatland extent and
 carbon sink in Sweden, as related to the Peatland/Ice Age Hypothesis. *Mines and Peat*, *10*(8), 1–
 19.
- Gatis, N., Benaud, P., Anderson, K., Ashe, J., Grand-Clement, E., Luscombe, D. J., Puttock, A., & Brazier,
 R. E. (2023). Peatland restoration increases water storage and attenuates downstream stormflow
 but does not guarantee an immediate reversal of long-term ecohydrological degradation. *Scientific Reports*, 13(1), 1–14. https://doi.org/10.1038/s41598-023-40285-4
- Goudarzi, S., Milledge, D. G., Holden, J., Evans, M. G., Allott, T. E. H., Shuttleworth, E. L., Pilkington, M.,
 & Walker, J. (2021). Blanket Peat Restoration: Numerical Study of the Underlying Processes
 Delivering Natural Flood Management Benefits. *Water Resources Research*, 57(4).
 https://doi.org/10.1029/2020WR029209
- Grand-Clement, E., Anderson, K., Smith, D., Luscombe, D., Gatis, N., Ross, M., & Brazier, R. E. (2013).
 Evaluating ecosystem goods and services after restoration of marginal upland peatlands in South-

west England. Journal of Applied Ecology, 50(2), 324–334. https://doi.org/10.1111/13652664.12039

- Haapalehto, T., Kotiaho, J. S., Matilainen, R., & Tahvanainen, T. (2014). The effects of long-term drainage
 and subsequent restoration on water table level and pore water chemistry in boreal peatlands. *Journal of Hydrology*, *519*(PB), 1493–1505. https://doi.org/10.1016/j.jhydrol.2014.09.013
- Haque, A., Ali, G., & Badiou, P. (2022). Event-based analysis of wetland hydrologic response in the
 Prairie Pothole Region. *Journal of Hydrology*, 604(September 2021), 127237.
 https://doi.org/10.1016/j.jhydrol.2021.127237
- Hawcroft, M., Walsh, E., Hodges, K., & Zappa, G. (2018). Significantly increased extreme precipitation
 expected in Europe and North America from extratropical cyclones. *Environmental Research Letters*, 13(12). https://doi.org/10.1088/1748-9326/aaed59
- Helbig, M., Waddington, J. M., Alekseychik, P., Amiro, B. D., Aurela, M., Barr, A. G., Black, T. A., Blanken,
 P. D., Carey, S. K., Chen, J., Chi, J., Desai, A. R., Dunn, A., Euskirchen, E. S., Flanagan, L. B., Forbrich,
 I., Friborg, T., Grelle, A., Harder, S., ... Zyrianov, V. (2020). Increasing contribution of peatlands to
 boreal evapotranspiration in a warming climate. *Nature Climate Change*, *10*(6), 555–560.
 https://doi.org/10.1038/s41558-020-0763-7
- Holden, J. (2006). Chapter 14 Peatland hydrology. In I. P. Martini, A. M. Cortizas, & W. Chesworth (Eds.), *Developments in Earth Surface Processes* (Vol. 9, Number C, pp. 319–346).
 https://doi.org/10.1016/S0928-2025(06)09014-6
- Holden, J., & Burt, T. P. (2003). Runoff production in blanket peat covered catchments. *Water Resources Research*, *39*(7). https://doi.org/10.1029/2002WR001956
- Holden, J., Chapman, P. J., & Labadz, J. C. (2004). Artificial drainage of peatlands: Hydrological and
 hydrochemical process and wetland restoration. *Progress in Physical Geography*, 28(1), 95–123.
 https://doi.org/10.1191/0309133304pp403ra
- Holden, J., Green, S. M., Baird, A. J., Grayson, R. P., Dooling, G. P., Chapman, P. J., Evans, C. D., Peacock,
 M., & Swindles, G. (2017). The impact of ditch blocking on the hydrological functioning of blanket
 peatlands. *Hydrological Processes*, *31*(3), 525–539. https://doi.org/10.1002/hyp.11031
- Holden, J., Kirkby, M. J., Lane, S. N., Milledge, D. G., Brookes, C. J., Holden, V., & McDonald, A. T. (2008).

657 Overland flow velocity and roughness properties in peatlands. *Water Resources Research*, 44(6),

658 1–11. https://doi.org/10.1029/2007WR006052

- Howie, S. A., Whitfield, P. H., Hebda, R. J., Munson, T. G., Dakin, R. A., & Jeglum, J. K. (2009). Water table
 and vegetation response to ditch blocking: Restoration of a raised bog in southwestern british
 columbia. *Canadian Water Resources Journal*, 34(4), 381–392.
 https://doi.org/10.4296/cwrj3404381
- Javaheri, A., & Babbar-Sebens, M. (2014). On comparison of peak flow reductions, flood inundation
 maps, and velocity maps in evaluating effects of restored wetlands on channel flooding. *Ecological Engineering*, *73*, 132–145. https://doi.org/10.1016/j.ecoleng.2014.09.021
- Karimi, S., Hasselquist, E., Salimi, S., Järveoja, J., & Laudon, H. (2024). Rewetting impact on the
 hydrological function of a drained peatland in the boreal landscape. *Journal of Hydrology*,
 668 641(April). https://doi.org/10.1016/j.jhydrol.2024.131729
- Karlsen, R. H., Bishop, K., Grabs, T., Ottosson-Löfvenius, M., Laudon, H., & Seibert, J. (2019). The role of
 landscape properties, storage and evapotranspiration on variability in streamflow recessions in a
 boreal catchment. *Journal of Hydrology*, *570*(December 2018), 315–328.
 https://doi.org/10.1016/j.jhydrol.2018.12.065
- Ketcheson, S. J., & Price, J. S. (2011). The impact of peatland restoration on the site hydrology of an
 abandoned block-cut bog. *Wetlands*, *31*(6), 1263–1274. https://doi.org/10.1007/s13157-0110241-0
- Kirchner, J. W., Benettin, P., & Van Meerveld, I. (2023). Instructive Surprises in the Hydrological
 Functioning of Landscapes. *Annual Review of Earth and Planetary Sciences*, *51*, 277–299.
 https://doi.org/10.1146/annurev-earth-071822-100356
- Kløve, B. (2000). Retention of suspended solids and sediment bound nutrients from peat harvesting
 sites with peak runoff control, constructed floodplains and sedimentation ponds. *Boreal Environment Research*, 5(1), 81–94.
- Laine, A. M., Leppälä, M., Tarvainen, O., Päätalo, M. L., Seppänen, R., & Tolvanen, A. (2011). Restoration
 of managed pine fens: Effect on hydrology and vegetation. *Applied Vegetation Science*, *14*(3),
 340–349. https://doi.org/10.1111/j.1654-109X.2011.01123.x
- Lane, S., Brookes, C., Hardy, R., Holden, J., James, T., Kirkby, M., Mcdonald, A., Tayefi, V., & Yu, D. (2003).
 Land Management, Flooding and Environmental Risk: New Approaches To a Very Old Question. *CIWEM National Conference, September*, 20.
- 688 Laudon, H., Hasselquist, E. M., Peichl, M., Lindgren, K., Sponseller, R., Lidman, F., Kuglerová, L.,

- Hasselquist, N. J., Bishop, K., Nilsson, M. B., & Ågren, A. (2021). Northern landscapes in transition:
 Evidence, approach and ways forward using the Krycklan Catchment Study. *Hydrological Processes*, *35*(4), 1–15. https://doi.org/10.1002/hyp.14170
- Laudon, H., Lidberg, W., Sponseller, R. A., Maher Hasselquist, E., Westphal, F., Östlund, L., Sandström, 692 C., Järveoja, J., Peichl, M., & Ågren, A. (2022). Emerging technology can guide ecosystem 693 security. 694 restoration for future water Hydrological Processes, 36(10), 1-5. 695 https://doi.org/10.1002/hyp.14729
- Laudon, H., Mosquera, V., Eklöf, K., Järveoja, J., Karimi, S., Krasnova, A., Peichl, M., Pinkwart, A., Tong,
 C. H. M., Wallin, M. B., Zannella, A., & Hasselquist, E. M. (2023). Consequences of rewetting and
 ditch cleaning on hydrology, water quality and greenhouse gas balance in a drained northern
 landscape. *Scientific Reports*, *13*(1), 20218. https://doi.org/10.1038/s41598-023-47528-4
- Liu, H., & Lennartz, B. (2019). Hydraulic properties of peat soils along a bulk density gradient—A meta
 study. *Hydrological Processes*, *33*(1), 101–114. https://doi.org/10.1002/hyp.13314
- Loisel, J., & Gallego-Sala, A. (2022). Ecological resilience of restored peatlands to climate change.
 Communications Earth and Environment, *3*(1), 1–8. https://doi.org/10.1038/s43247-022-00547-x
- Luscombe, D. J. (2014). Understanding the ecohydrology of shallow, drained and marginal blanket
 peatlands. *Department of Geography, PhD,* 180.
 https://ore.exeter.ac.uk/repository/handle/10871/15967
- Menberu, M. W., Haghighi, A. T., Ronkanen, A. K., Marttila, H., & Kløve, B. (2018). Effects of Drainage
 and Subsequent Restoration on Peatland Hydrological Processes at Catchment Scale. *Water Resources Research*, 54(7), 4479–4497. https://doi.org/10.1029/2017WR022362
- Menberu, M. W., Tahvanainen, T., Marttila, H., Irannezhad, M., Ronkanen, A.-K. A., Penttinen, J., &
 Kløve, B. (2016). Water-table-dependent hydrological changes following peatland forestry
 drainage and restoration: Analysis of restoration success. *Water Resources Research*, *52*(5), 3742–
- 713 3760. https://doi.org/10.1002/2015WR018578
- Montanarella, L., Jones, R. J. A., & Hiederer, R. (2006). The distribution of peatland in Europe. *Mires and Peat*, 1(1), 1–10. http://hdl.handle.net/1826/3415
- Noumonvi, K. D., Ågren, A., Ratcliffe, J. L., Öquist, M. G., Ericson, L., Tong, C. H. M., Järveoja, J., Zhu, W.,
 Osterwalder, S., Peng, H., Erefur, C., Bishop, K., Laudon, H., Nilsson, M. B., & Peichl, M. (2023). The
 Kulbäcksliden Research Infrastructure: a unique setting for northern peatland studies. *Frontiers in*
 - 30

- 719 *Earth Science, 11*(May). https://doi.org/10.3389/feart.2023.1194749
- Puttock, A., Graham, H. A., Ashe, J., Luscombe, D. J., & Brazier, R. E. (2021). Beaver dams attenuate
 flow: A multi-site study. *Hydrological Processes*, *35*(2). https://doi.org/10.1002/hyp.14017
- Rosner, B. (1983). Percentage points for a generalized esd many-outlier procedure. *Technometrics*,
 25(2), 165–172. https://doi.org/10.1080/00401706.1983.10487848
- Shantz, M. A., & Price, J. S. (2006). Characterization of surface storage and runoff patterns following
 peatland restoration, Quebec, Canada. *Hydrological Processes*, 20(18), 3799–3814.
 https://doi.org/10.1002/hyp.6140
- Shuttleworth, E. L., Evans, M. G., Hutchinson, S. M., & Rothwell, J. J. (2015). Peatland restoration:
 Controls on sediment production and reductions in carbon and pollutant export. *Earth Surface Processes and Landforms*, 40(4), 459–472. https://doi.org/10.1002/esp.3645
- Shuttleworth, E. L., Evans, M. G., Pilkington, M., Spencer, T., Walker, J., Milledge, D., & Allott, T. E. H.
 (2019). Restoration of blanket peat moorland delays stormflow from hillslopes and reduces peak
 discharge. *Journal of Hydrology X*, *2*, 100006. https://doi.org/10.1016/j.hydroa.2018.100006
- Soomets, E., Lõhmus, A., & Rannap, R. (2023). Restoring functional forested peatlands by combining
 ditch-blocking and partial cutting: An amphibian perspective. *Ecological Engineering*, *192*(May
 2021). https://doi.org/10.1016/j.ecoleng.2023.106968
- Wickham, H. (2016). ggplot: Elegant Graphics for Data Analysis (Version 3.4.2) [R package]. Springer Verlag. https://cran.r-project.org/web/packages/ggplot2/index.html
- Wilson, L., Wilson, J., Holden, J., Johnstone, I., Armstrong, A., & Morris, M. (2010). Recovery of water
 tables in Welsh blanket bog after drain blocking: Discharge rates, time scales and the influence of
 local conditions. *Journal of Hydrology*, *391*(3–4), 377–386.
 https://doi.org/10.1016/j.jhydrol.2010.07.042
- Wilson, L., Wilson, J., Holden, J., Johnstone, I., Armstrong, A., & Morris, M. (2011). The impact of drain
 blocking on an upland blanket bog during storm and drought events, and the importance of
 sampling-scale. *Journal of Hydrology*, 404(3–4), 198–208.
 https://doi.org/10.1016/j.jhydrol.2011.04.030
- 746

748 Supplementary information

750 Table S 1. Characteristics of the four rainfall-runoff events shown in Figure 4 for the rewetted (R1 and R2) and control sites 751 during the pre- and post-rewetting years.

	Site	Total rain (mm)	Peak flow (mm h ⁻¹)	752 Antecedent GWI (mm) 753
Pre-rewetting		37		<u> </u>
	Control		0.29	- ¹⁵ 754
	R1		0.93	-82
	R2		0.54	-102 755
1 year post- rewetting		63		756
	Control		0.27	-35
	R1		0.71	-24 757
	R2		0.29	-85
2-years post- rewetting		53		758
	Control		0.22	⁻⁴⁷ 759
	R1		0.61	-34
	R2		0.37	-57 760
3-years post- rewetting		69		761
	Control		0.26	-19
	R1		0.14	-5.1 762
	R2		0.12	-40