



1 **Spatial relationship between hydrodynamic and physico-chemical parameters of surface**
2 **water for a basin with shale rock series as an indicator of intensity and direction of**
3 **chemical denudation in the Western Carpathians**

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23 **Abstract**

24 Hydrochemical evaluation of stream quality in the Western Carpathians requires a system
25 approach, gradually excluding factors less or more responsible for washing, mixing contaminants
26 and their farther transportation in the stream channel. In this work, the spatial autoregression
27 model was used to estimate the relationship between hydrodynamic and physico-chemical
28 parameters of surface water in various groups and variants of basin use. The highest mean shear
29 stress was 0.178 N·m⁻² in forest surface water. The highest mean Reynolds number (23654) was
30 recorded in the stream channel at permanent grassland, and the lowest number (0.426) at arable
31 lands. Analysis of spatial autoregression to a high degree showed space-time relations in various
32 measurement points. The turbulent diffusion coefficient should be regarded in the space-physical
33 model, constructed based on the influence of hydrodynamic indicators on the shaping of physico-
34 chemical parameters in the flysch basin. The autoregression confirmed that the turbulent
35 diffusion coefficient played a high role for ions K⁺ and P-PO₄³⁻ in surface water at arable lands



36 and for cation K^+ , as well as total iron for grassland ($p < 0.05$). A relation for physico-chemical
37 was not found for surface water in forests. The results, to a high degree, will be used to create an
38 erosion model concerning the alimentation of alluvial deposits from weathered Carpathian flysch
39 or surface wash depending on the material delivery in a basin.

40 1. Introduction

41 The factors influencing the physico-chemical state of surface water to the highest degree
42 are soil, relief (Shi et al., 2016), and soil plant cover occurring in the river-bank zone (Andersson
43 et al., 2015; Teixeira and Marques, 2016). In flysch basins, surface runoff is selective because
44 transported material often undergoes local accumulation on a slope during delivery to a stream
45 (Gil and Kotarba, 1977). Reliability of results is ensured when the picture of surface erosion is
46 being captured in mountain valleys, and proper frequency of measurement series is carried out,
47 together with specification of measurement error (Halecki et al., 2018a).

48 Investigations of water erosion can be carried out from a theoretical point of view and the
49 use of mathematical models (Panagos et al., 2015), for example, in the aspect of modelling the
50 shape of nitrates in surface runoff (Wang et al., 2014). Further, empirical equations and
51 theoretical methods are used for the settlement of river material quantity in various measurement-
52 control points (Mazur and Pałys, 1992) and intensity degree of suspended sediment transportation
53 in the shape of mechanical weathered rock mantle (Izmailow et al. 2008; Starkel 2011, Brynda et
54 al. 2014; Comino et al., 2016). It is important to remember that technical solutions simulating
55 digital sediment transportation are only approximate and do not reflect real data. The cause of
56 unreliable results is the short or irregular frequency of field measurements (Halecki et al., 2018 b).

57 Material transported in the Beskid region basins is connected with land use and the
58 occurrence of shale rock series (Halecki et al., 2018 c). Material delivered from slopes determines
59 the concentration of particular ions transferred in the stream channel. Ions can originate from the
60 dissolution of mineral fertilisers from adjacent areas of agricultural use (Hall et al., 2014; Gernez et
61 al., 2015). The fundamental biogenic indicators in the hydro-chemical evaluation of surface water
62 are phosphates and nitrates. Nitrogen most often appears as an ammonium ion
63 ($N-NH_4^+$), testifying point contamination of surface water. Nitrate ($N-NO_3^-$) and nitrite nitrogen
64 ($N-NO_2^-$) are indicators of the long-term influence of pollutants, particularly in the proximity of
65 crops. In summer and early autumn, $N-NO_3^-$ decreases, which is connected with the plant's



66 demand for nutrients (Ulénet al., 2012). The concentration of calcium (Ca^{2+}), magnesium (Mg^{2+}),
67 sodium (Na^+), potassium (K^+), total iron (Fe_{tot}), and nitrite nitrogen (N-NO_2) is differentiated in
68 spring. This shows that plants accumulate more nutrients during the warmer season (Padmalalet al.,
69 2012). For estimation of the physico-chemical quality of surface water, the good indicators
70 regarding transformation and accumulation are various forms of phosphorus. They pose a high
71 threat to the purity of surface water as factors that favour eutrophication and excessive rise of
72 biomass of particular algae (Smoroń, 2012). Changes in biogen concentration in surface streams
73 are influenced by the spatial distribution of arable land (Arienzo et al., 2012; Tasdighiet al., 2017).

74 Sedimentary rock undergoes an intensive leaching process. In surface water, sodium
75 cations can originate from industrial waste and potassium from agricultural cultivations, where
76 potassium fertilisers are used (Oster et al., 2016). Hydrolytic decay of minerals containing sodium
77 salts (aluminosilicates) and weathering of sedimentary rocks contribute to the creation of
78 alimentation (source) for sodium and potassium cations in surface water (Zhang et al., 2017).
79 Magnesium salts occur in all natural waters, both surface and underground. Leached Ca^{2+} and
80 Mg^{2+} cations penetrate the basin by infiltrating rainfall from fertilised cultivated areas and
81 supplying surface water (Grochowska, 2016). Increased concentration of Ca^{2+} and Mg^{2+} cations in
82 surface water testify to the occurrence of calcareous rocks, marls, and dolomites and depends on
83 the degree of basin management (Halecki, 2015). In addition, the concentration of Ca^{2+} cations
84 and SO_4^{2-} anions in small stream channels is attributed to geology, mainly the calcium carbonate
85 and sulphate concentration in the clastic rock mantle. Moreover, certain water quality factors
86 require specialistic investigation and hydrochemical evaluation according to the World Health
87 Organisation (WHO) standards. These include Ca^{2+} and Mg^{2+} cations, N-NO_2^- , N-NO_3^- , SO_4^{2-} , and
88 phosphate phosphorus (P-PO_4^{3-}) anions, and salinity indicators - specifically the electric
89 conductivity of water and the concentration of dissolved substances (sum of determined mineral
90 substances in a shape of fine clastic material) with total suspension (Shigutet al., 2017).

91 Dissolved oxygen is essential for the evaluation of water conditions. The activity of
92 microorganisms responsible for oxidation of organic compounds is examined using two
93 indicators, namely biochemical oxygen demand (BOD) and chemical oxygen demand (COD). In
94 evaluating surface water quality, higher oxygen use indicates contamination (Bo et al., 2017;
95 Effendi et al., 2018). On the other hand, a lower dissolved oxygen concentration can show
96 biochemical decay of accumulated organic substances, respiration of water organisms, or



97 oxidation of inorganic substances (Matta et al., 2017). Furthermore, water temperature controls
98 the solubility of solid substances and the concentration of suspension (Saito et al., 2005a). So,
99 low temperature weakens the chemical and biological activity of water and increases viscosity,
100 enabling the transportation of material (Jarocki, 1957). Higher values of dissolved oxygen and
101 BOD are connected with their high sensibility to changes in water caused by the activity of water
102 engineering (Parmar and Keshari, 2012). Essential differences between BOD and COD of
103 anthropogenic origin have been recorded in a seasonal hydro-chemical evaluation period
104 (Bellver-Domingo and Hernández-Sancho, 2018). Seasonal evaluation of temperature, pH, COD,
105 BOD and heavy metals: Fe, Mn, Ni, Cd, Cr, Co, Cu, Pb and Zn, is needed either for
106 determination of the recommended level and permissible concentration of pollutants for drinking
107 water (Vincent-Akpuandin, 2015). The anthropogenic activity contributes to delivering heavy
108 metals to surface water, mainly by using industrial objects (Juahiret al., 2010; Weber et al., 2014;
109 Assouline et al., 2015; Hu et al., 2015; Vaddeet al., 2018).

110 Contamination of surface water is proven by the decay of living organisms, which mineral
111 composition is related to the increase of sulphates. Chlorine is also present in the form of chloride
112 anions (Cl⁻). In surface water, it is transported as a result of weathering of rock-forming minerals.
113 Dissolution of evaporates (sedimentary rock) enriched in anhydrite constitutes a natural source of
114 SO₄²⁻ anions. Further, sulphate ions do not undergo sorption in an underground environment, and
115 similar to Cl⁻, it is the indicator of pollutant penetration to underground water (Sapek, 2008;
116 Geurtset al., 2009).

117 The first detailed aim of this study was to arrange useful values of the basin and,
118 secondly, to conduct a hydro-chemical evaluation of the surface water in various land
119 configurations of the basin lithology. The third aim was to determine hydraulic conditions in the
120 flysch stream channel. A final goal was to determine the main factors intensifying short-term
121 relations between hydrodynamic parameters of mountain streams and chemical compounds
122 leached from the slope to the water of the flysch stream as a result of interrill erosion. For such
123 detailed aims, hypotheses were formulated, methods and schemes of frequency measurement
124 series were set, and specifications of measurement errors were elaborated.

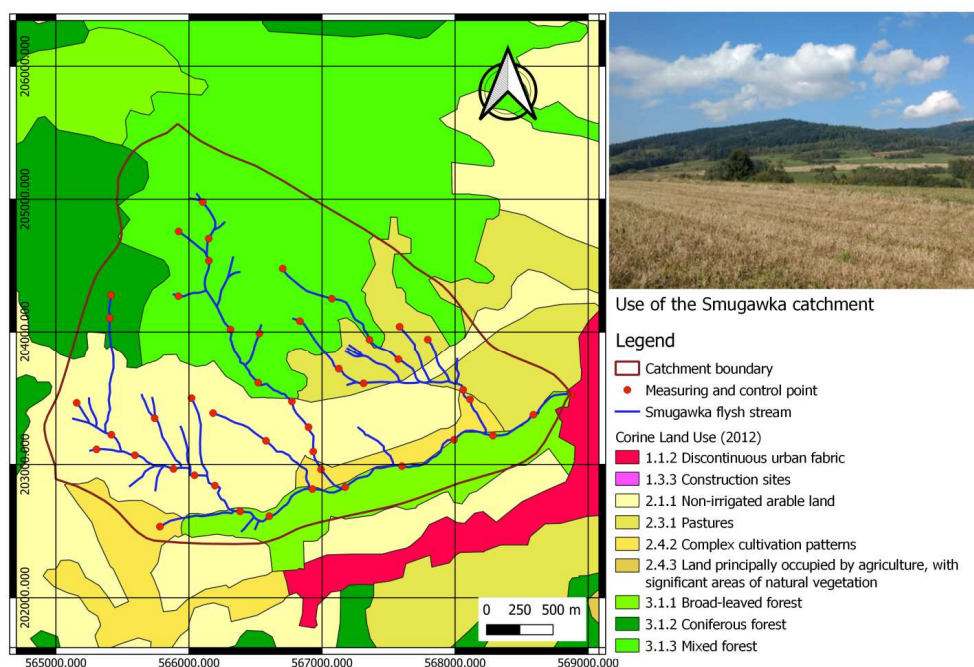
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126 2. Methods



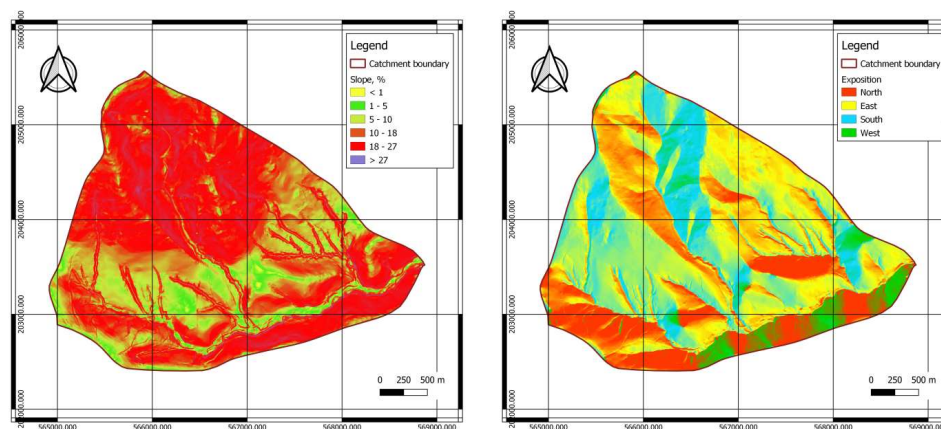
127 2.1. Investigated site and uptake and examination of surface water samples

128 The measurements were carried out in the Smugawka stream in the Beskid Wyspowy, the
129 Western Carpathians (Poland). The localisation of sampling points and land use is presented in
130 Figure 1. Slopes and exposition are shown in Figures 2. Water samples were taken in 1 dm³
131 volume containers once a month to determine the physico-chemical parameters of surface water.
132 The sampling period occurred in the spring-summer season (from March) and fall-winter (to
133 November) from 2014 to 2018. Sampling was carried out punctually, in cross-section levels, in
134 source and outlet places for the dimension of samples to be multiplied. After filling, each
135 container was emptied and then filled with flowing water for 120 seconds.



136

137 *Figure 1. Localisation of sampling points and land use*



138 *Figure 2.* Slope and exposition of the Smugawka stream catchment

139 One water sample was enough to determine the mean concentration of dissolved material
140 (Brański, 1968; Dalbianco et al., 2017) because it represents all hydrometric sections of the
141 channel in which strong longitudinal and transversal dispersion occurs. Moreover, the
142 relationship between the concentration of dissolved material is directly proportional to water flow
143 and in accordance with the hydrological system of a stream (Wojtasik and Szatten, 2014). Six
144 samples were taken during each field visit. Measurement of convection intensity was carried out
145 with flow intensity. At <0.3 m, water was taken directly to the container. After determining the
146 relationship between the mean concentration of total suspension in the section and the
147 concentration of punctual sediment in a cross-section, the number of samples was limited to 2-4.
148 The database contained 24 variables (physico-chemical features) in 25 replications. In summary,
149 600 samples were taken for laboratory analysis. To assure comparability of results, investigations
150 into physico-chemical surface water quality during field visits (additionally there were carried out
151 250 measurements) were conducted, recording actual results at low water levels and during
152 freshets.

153 **2.2. Hydro-chemical indicators**

154 Geodetic measurements were made using the tachymeter TOPCON GTS-226. Five cross-
155 sections of the bed channel slopes, from 0.79 to 2.18%, were localised along all stream lengths
156 (from sources to outlet). The average distance between them was 875 m. Granular measurements
157 of bottom sediment were carried out using the Wolman method (Wolman, 1954). In every



158 measurement point, 15 cobbles were taken, and their mean axis „b” was measured. The grain size
159 distribution curves were drafted, and their characteristic diameters: d_{\min} , d_{30} , d_{50} , d_{70} , and d_{\max} and
160 competent diameter: d_m were determined. Measuring points were located crosswise channel in a
161 distance of 0.5 m.

162 Hydrodynamic measurements were carried out at the same points as the granular ones.
163 Measurements were made using a hydrometric current meter VALEPORT Model 801 Flat EM
164 flow meter. This device allows the measurement of mean flow velocity in assigned time intervals.
165 It also measures instantaneous velocity and filling in the stream channel. The measurement scope
166 of the device is $5.0 \text{ m} \cdot \text{s}^{-1}$, and exactness $\pm 0.001 \text{ m} \cdot \text{s}^{-1}$. Through this device, mean velocity V_{mean} on
167 the height 0.4 m above the bottom, maximum velocity V_{\max} , and instantaneous velocity V
168 measured just above the bottom, allowed calculation of dynamic velocity based on charts of
169 velocity distribution above the bottom in a semi-logarithmic system (Gordon et al., 2007):

$$V_* = \frac{a}{5.75} [\text{m} \cdot \text{s}^{-1}]$$

170 where:

171 a = coefficient of straight line slope

172 $V = f(h)$, taking an equation form of $y = ax + b$ (x = height above bottom, the measurement of
173 velocity was carried out, b = free term of the equation).

174 Knowing dynamic velocity, shear stress was calculated as follows:

$$\tau = \rho \cdot (V_*)^2 [\text{N} \cdot \text{m}^{-2}]$$

175 where:

176 ρ = water density ($\text{kg} \cdot \text{m}^{-3}$)

177

178 The Reynolds number (mean, maximum, and grain) was used to determine the type of
179 flow pattern as laminar or turbulent. Further, the Freud number (mean and maximum) is a
180 measurement of bulk flow characteristics. These were respectfully calculated using the following
181 equations:



$$\begin{aligned}Re_{sr} &= \frac{V_{sr} \cdot h}{\nu} \\Re_{max} &= \frac{V_{max} \cdot h}{\nu} \\Re_{dm} &= \frac{V_* \cdot d_m}{\nu} \\Fr_{sr} &= \frac{V_{sr}}{\sqrt{gh}} \\Fr_{max} &= \frac{V_{max}}{\sqrt{gh}}\end{aligned}$$

182 where:

183 h = filling (m),

184 g = acceleration of gravity ($m \cdot s^{-2}$),

185 ν = kinematic coefficient of water viscosity, calculated from the equation:

$$\nu = \frac{0.00178}{\rho(1 + 0.0337t + 0.000221t^2)} [m^2 \cdot s^{-1}]$$

186 where:

187 t = water temperature ($^{\circ}C$)

188

189 In this study, hydrodynamic indicators allowed the calculation of the distance of full mix
190 in surface water, used in the analysis of spatial autoregression. Dynamic velocity (u) in the stream
191 was calculated using the following equation (Loga, 2016) :

$$u = \sqrt{h \cdot g \cdot S} [m \cdot s^{-1}],$$

192 where:

193 h = mean height of water (m),

194 g = acceleration of gravity ($m \cdot s^{-2}$),

195 S = slope of channel bottom (%).

196

197 There is a risk connected to the negative influence of the frequency of measurements.
198 In short periods of freshets, some part of the yearly sediment load cannot be observed. Therefore,



199 the following equation was used to present a moment in which the mix of material in water
200 occurs (a load of total suspended sediment and physico-chemical composition). That is some
201 quantitative representativeness of transported material. In addition, the turbulent diffusion
202 coefficient ($D_{t,y}$) was calculated for transverse mix in the particular years in a so-called natural
203 channel (Loga 2016):

$$D_{t,y} = 0,6 h \cdot u \quad [\text{m}^2 \cdot \text{s}^{-1}],$$

204 where:

205 h = mean height (m),

206 u = dynamic velocity ($\text{m} \cdot \text{s}^{-1}$).

207

208 **2.3. Determination of physico-chemical properties of surface water**

209 In the experimental area, the physico-chemical properties of surface water were measured.
210 The reaction pH of water was examined using the potentiometric method with a CP-104
211 ELMETRON equipped with a combined electrode. Specific electrical conductivity, expressed in
212 $\text{dS} \cdot \text{m}^{-1}$, was measured with a conductometer Elmetron CC-101. The concentration of dissolved
213 oxygen in ($\text{mg} \cdot \text{dm}^{-3}$) and the degree of water saturation by oxygen were determined using the
214 electrochemical method using an oxygen-meter Elmetron CO-411. Further, the water temperature
215 ($^{\circ}\text{C}$) was measured by a digital thermometer built in waterproof oxygen-meter CO-411. The total
216 content of dissolved substances in ppm was calculated by means of a digital TDS device. In the
217 laboratory, the following physico-chemical properties of water were determined using the
218 UNICAM SOLAAR 969 atomic absorption spectrophotometer: Ca^{2+} , Na^{+} , K^{+} , Mg^{2+} , Fe_{tot} , and
219 manganese Mn^{2+} .

220 Moreover, N-NH_4^{+} , N-NO_3^{-} , N-NO_2^{-} , P-PO_4^{3-} , and Cl^{-} were determined by the colorimetric
221 flow-injection method on the computer-controlled FIA Star 5000 apparatus of the FOSS
222 firm. Sulphates (SO_4^{2-}) were determined gravimetrically. Biological Oxygen Demand (BOD_5) was
223 analysed by the Winkler method: water samples were treated with manganese sulfate MnSO_4 ,
224 alkaline potassium iodide KI, and sulfuric acid H_2SO_4 , followed by titration with sodium
225 thiosulfate $\text{Na}_2\text{S}_2\text{O}_3$. The chemical oxygen demand (ChZT-Mn) was calculated by the



226 permanganate method by heating the water sample with KMnO_4 potassium permanganate in an
227 acidic environment, and the amount of consumed oxygen was determined by titration.

228 During subsequent readings of water levels, samples were taken from the stream by the
229 bathometric method (containers with a volume of 1 dm^3). The concentration of total suspended
230 solids was determined by the gravimetric method after drying using tared filters with an accuracy
231 of $\pm 0.0005 \text{ g}$. The dry residue (sum of mineral particles from the transported samples) from the
232 collected material was filtered to determine the different composite concentrations. The ionic
233 forms of zinc (Zn^{2+}), lead (Pb^{2+}), cadmium (Cd^{2+}) and copper (Cu^{2+}) were determined once per
234 quarter with the coulometric method using an electrochemical analyser for the determination of
235 trace amounts of heavy metals (EcaFlow 150 GLP by PolEko). The nephelometric-laser method
236 was also used to assess the degree of water turbidity using a Hach Lange 2100QS nephelometer
237 (turbidimeter) in the range of 0–2000 FNU, per the international standard (Nephelometric
238 Turbidity Unit; NTU).

239 **2.4. Data reduction to a spatial model and statistical analysis**

240 Before the analyses were performed, the asymptotic distribution for the χ^2 Jarque-Ber (JB)
241 test statistic was studied, which takes kurtosis and skewness into account (Jarque and Bera,
242 1987). The data was grouped in ascending order by the Anderson-Darling concordance test,
243 determined by the empirical distribution, and the normalisation for the mean and standard
244 deviation was calculated (Stephens, 1986). The test determines a weighted Cramér-von Mises
245 distance between the empirical (F_n) and theoretical (F) cumulative distribution factors, with
246 weights corresponding to the reciprocal of the empirical cumulative distribution (Anderson and
247 Darling, 1954). The Shapiro-Wilk test, based on positional statistics, was also used. The keynote
248 of this analysis is collinearity between the empirical quantile (i.e., the ordinal statistic) and the
249 theoretical quantile (i.e., the expected value of the ordinal statistic) along the $y = x$ straight line
250 (Shapiro and Wilk, 1965).

251 The homogeneity of variance was analysed by Levene's test (Levene, 1960). To reduce the
252 number of variables, the Principal Component Analysis (PCA) was used, presenting the
253 relationships between significant physico-chemical properties in the surface water of the stream
254 running near arable land, grasslands, and forests. Statistical significance for the matrix of
255 correlation coefficients and the strength of the relationship between the variables were assessed



256 with the Barlett sphericity test (Bartlett, 1950; Williams et al., 2010). Partial correlations with
257 bivariate correlation coefficients were performed using the Keiser-Mayer-Olkin coefficient
258 (KMO), determining the accuracy of selecting variables for the tested model. Adequate variables
259 have a KMO coefficient > 0.5 (MacCallum, 1983; Hair et al., 1995; Tabachnick and Fidell, 2007;
260 Szüle 2016).

261 In order to show the relationship between the tested physico-chemical properties of
262 surface water, values for eigenvectors were given, approximating the influence of primary
263 variables on the main component. The most significant variables were interpreted through factor
264 loadings, which also reflect the influence of individual variables on a given principal
265 component. The study used a correlation matrix, and the factor loadings were interpreted as
266 correlation coefficients between the original variables and the next analysed main component.
267 Following the PCA method, the variance size for each computed component and the so-called
268 primary variables were indicated. The most important variables were selected for further data
269 processing based on the value of factor loadings. The analysis was performed in PQ Stat
270 Software, version 1.6.6.

271 Spatial autoregression is a tool for predicting features in a linear system. The analysis
272 aimed to determine the factors that influence the temporal changes in the physico-
273 chemical composition of surface water. Therefore, the study used a spatial autoregression model
274 with a lagged response, facilitating the study of the relationship between the distance of complete
275 mixing and physico-chemical features randomly distributed in the water column. The general
276 form of spatial autoregression is given by the formula (Rangel et al., 2010):

$$Y = \rho \cdot W_y + X\beta + \varepsilon$$

277 where:

278 Y = vector ($N \cdot 1$) of the explained spatial process through a responsive (explained) variable,

279 W_y = weight matrix of the vicinity of the examined variables in different locations ($N \cdot N$ matrix
280 of spatial weights),

281 ρ = spatial autoregression parameter reflecting the strength of the relationship between the
282 variables (spatial interaction parameter),

283 X = matrix ($N \cdot K$) of explanatory processes,

284 β = vector ($K \cdot 1$) of structural indices,



285 ε = vector ($N \cdot 1$) of the random term (estimation error and spatial noise not described by the
286 model; random effect).

287

288 The spatial autoregression model was used to show the relationship between the variables
289 distributed in space over the course of a measurement series for homogeneous areas and to
290 facilitate the analysis of the variables, especially with water flows greater than average. The
291 transverse turbulence diffusion coefficient was assumed as the dependent variable, and surface
292 water's physico-chemical features were considered independent variables. The simultaneous
293 autoregressive model (SAR), with a system of lagged predictors (Rangel et al., 2010), was used
294 to calculate the spatial dependencies:

$$C = \sigma^2[(I - \rho \cdot W)^T]^{-1}[I - \rho \cdot W]^{-1}$$

295 where:

296 C = vector of the explained spatial process through a responsive (dependent) variable,

297 W = weight matrix of the vicinity of the examined variables in different locations matrix ($N \cdot N$)
298 of spatial weights,

299 σ^2 = residual variance between observations,

300 ρ = spatial autoregression parameter reflecting the strength of the relationship between the
301 variables (spatial interaction parameter),

302 I = linear transformation (matrix of $N \cdot N$ -type) of the dependent variable.

303 The autoregression analysis documented the overarching variables for each type of land
304 use, combining the physico-chemical characteristics of the surface water. The calculations were
305 made in the SAM 1.6.6 software.

306

307

308

309

310



311 **3. Results**

312 **3.1. Physico-chemical and hydrodynamic features of the stream bed**

313 The range of concentrations of the salinity indicators is presented in Table 1, and the
 314 range of concentrations of metals, including heavy metals, for the Smugawka stream throughout
 315 the study period is provided in Table 2. In surface water flowing out of arable land, a strong
 316 positive correlation was found for Ca^{2+} and P-PO_4^{3-} for the first two factors of principal
 317 component analysis (Table 3). In grassland, the water temperature played an important role,
 318 distinguishing strong factor loadings. In contrast, the concentration of Fe_{og} showed a positive
 319 correlation with the first and a high negative correlation with the second factor of the PCA (Table
 320 4). Water temperature and concentration of ChZT-Mn were also characterised by the highest
 321 variance for the factor axes in all tested surface water samples and the highest factor loadings
 322 (Table 5 and Figure 3). The transport intensity increased with the following gradient: forests >
 323 grassland > arable land (Table 6).

324 Table 1. The range of concentrations of salinity indicators for the Smugawka stream during the
 325 study period

326

Salinity indicators	Unit	Measuring point		
		Surface water runoff from arable land	Surface water runoff from grassland	Surface water runoff from forests
Na^+	[$\text{mg} \cdot \text{dm}^{-3}$]	4,80–88,31	3,91–68,72	8,92–37,54
K^+		1,78–5,20	0,68–10,40	3,20–5,40
Mg^{2+}		6,64–10,92	3,74–19,35	10,34–22,59
Ca^{2+}		39,53–133,08	31,22–147,23	58,63–80,82
SO_4^{2-}		43,87–204,24	23,56–193,36	45,05–120,93
Cl^-		23,94–134,45	35,05–195,36	24,59–145,48
*CZSR		101–147	78–199	148–178
Electrical conductivity	[$\mu\text{S} \cdot \text{cm}^{-1}$]	156–240	247–350	148–256

327 *CZSR – Total content of dissolved substances

328

329 Table 2. The range of concentrations of metals, including heavy metals, for the Smugawka stream
 330 during the study period

331

Indicators of metal concentrations	Unit	Measuring point		
		Surface water runoff from arable land	Surface water runoff from grassland	Surface water runoff from forests
Zn^{2+}	[$\mu\text{g} \cdot \text{dm}^{-3}$]	12,53–50,10	2,93–41,90	19,94–47,80
Pb^{2+}		1,61–6,80	0,01–7,60	0,01–1,34
Cd^{2+}		0,01–1,10	0,01–1,90	0,01–1,50
Cu^{2+}		1,51–3,04	0,91–1,50	0,01–0,70



Fe _{og}	[mg·dm ⁻³]	0,53–0,91	0,03–1,04	0,1–0,81
Mn ²⁺		0,12–0,18	0,06–0,09	0,09–0,10

332

333 Table 3. Factor loadings obtained from principal component analysis (PCA) for the studied
 334 physicochemical characteristics from surface water samples taken at arable land

335

Variable	Factor 1	Factor 2	Factor 3	Factor 4	Factor 5	Factor 6
K ⁺	-0,72	0,60	-0,10	0,05	-0,06	0,46
pH	-0,13	0,35	-0,78	-0,06	0,49	-0,11
N-NO ₂	0,31	0,17	0,19	0,90	0,15	0,01
Ca ²⁺	0,88	0,92	-0,03	-0,07	-0,01	0,14
Mg ²⁺	0,84	-0,86	0,02	-0,16	0,07	0,13
P-PO ₄ ³⁻	0,93	0,87	-0,01	-0,15	0,04	0,18
Na ⁺	0,87	-0,78	-0,31	0,13	0,19	0,23
Total suspended solids	-0,15	0,24	0,78	-0,21	0,52	0,04

336

337 Table 4. Factor loadings obtained from principal component analysis (PCA) for the studies
 338 characteristics physicochemical characteristics from surface water samples taken in grasslands

339

Variable	Factor 1	Factor 2	Factor 3	Factor 4	Factor 5	Factor 6
Temperature of water	0,82	0,76	-0,60	0,10	-0,10	0,12
Total suspended solids	-0,63	0,58	-0,62	-0,17	0,08	-0,14
Fe _{og}	0,75	-0,63	-0,15	-0,22	0,22	0,19
K ⁺	-0,59	0,51	-0,19	-0,53	-0,20	-0,13
Ca ²⁺	-0,61	0,54	0,17	-0,26	0,43	0,001
Na ⁺	-0,85	-0,73	0,04	-0,03	-0,13	0,11

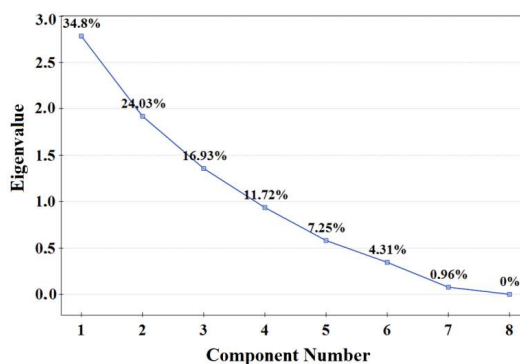
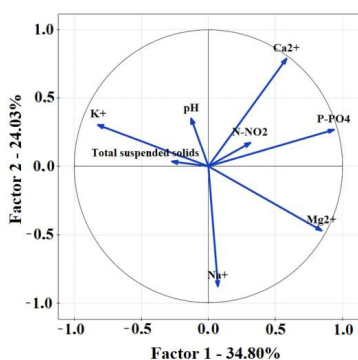
340

341 Table 5. Factor loadings obtained from principal component analysis (PCA) for the studied
 342 physicochemical characteristics from surface water samples taken in forests

343

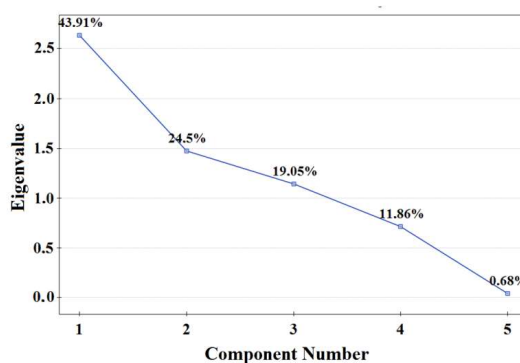
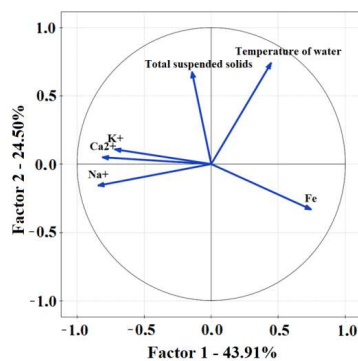
Variable	Factor 1	Factor 2	Factor 3	Factor 4	Factor 5	Factor 6
Temperature of water	0,84	0,86	-0,28	0,41	-0,03	0,32
ChZT-Mn	0,92	0,87	-0,51	0,21	0,33	0,08
Na ⁺	0,97	-0,78	-0,48	-0,44	0,12	0,23
Conductance	0,23	-0,35	-0,30	0,21	0,53	-0,59
BZT ₅	-0,16	-0,53	0,08	-0,64	0,46	0,22
Mg ²⁺	-0,38	0,13	0,55	0,13	0,50	-0,07
Dissolved oxygen	-0,56	-0,81	0,15	0,35	0,33	0,65
Fe _{og}	-0,51	0,84	-0,13	-0,01	-0,17	0,10

344



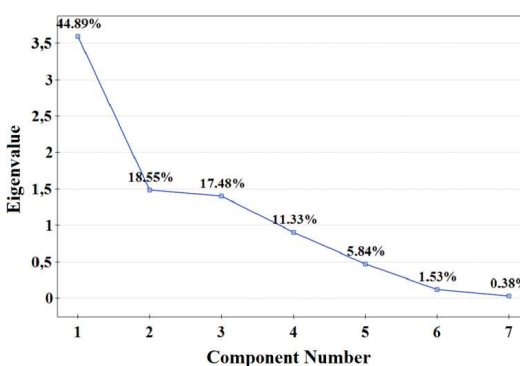
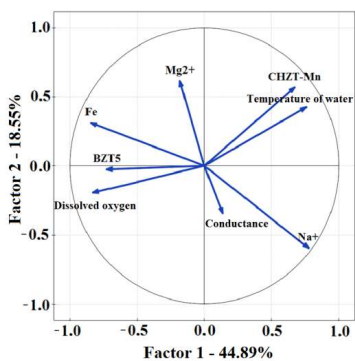
a) PCA analysis for physicochemical characteristics for samples collected in surface water near arable land

b) Scree plot for samples taken in surface water near arable land



c) PCA analysis for physicochemical characteristics for samples collected in surface water from grasslands

d) Scree plot for sample taken in surface water from grasslands



e) PCA analysis for physicochemical characteristics for samples collected in surface water near forests

f) Scree plot for samples taken in surface water near forests

345
 346
 347

Figure 3. PCA analysis and scree plots for physicochemical characteristics for samples taken in surface water: arable land (a, b), grassland (c, d), forest (e, f)



348 Table 6. Range of average values of soil material loss in the form of total suspended solids
 349 throughout the study period depending on the land use of the Smugawka catchment area
 350

Index of eroded soil material	Unit	Arable land	Grasslands	Forests
Circadian concentration of total suspended solids	[mg·dm ⁻³]	4,94–18,34	0,93–9,92	0,71–5,92
Total suspended solids load	[Mg]	0,05–0,52	0,05–0,34	0,02–0,23
Annual total suspended solids load	[Mg·rok ⁻¹]	18,69–190,27	19,65–126,95	7,38–53,78

351
 352

353 3.2. Data reduction and spatial model

354 Spatial autoregression was used to reveal the relationships between the studied variables.
 355 Relationships with some physico-chemical features of surface water were detected in the model.
 356 The SAR equation for the entire catchment area showed that P-PO₄³⁻, N-NO₂⁻, TDS, and the
 357 concentrations Na⁺, Mg²⁺, Ca²⁺ and Mn²⁺ (Table 7) are appropriate predictors for judging the
 358 distance of complete mixing. In general, the autoregressive modelling can be summed up by the
 359 fact that the few relationships for the physico-chemical characteristics of the surface water with
 360 the transverse turbulence diffusion coefficient of the flysch stream were significant (p < 0.05). In
 361 the spatial autoregression model divided into variants of use, statistically significant relationships
 362 were found for K⁺ and P-PO₄³⁻ in surface water of arable land and for Fe_{og} in surface water at
 363 grassland. No significant relationships were found at the measurement points in surface water
 364 near forests (Table 8). In the spatial autoregression model for the whole with the explanatory
 365 variable (the distance of complete mixing) detected the spatial dependencies for P-PO₄³⁻, N-NO₂⁻,
 366 Mn²⁺, TDS, Na⁺, Mg²⁺, Ca²⁺, and Mn²⁺ (Table 9).

367 Table 7. The results of the spatial autoregression model of the SAR type for the entire catchment.
 368 The - distance of full mixing was chosen as the explanatory variable,
 369 and physicochemical characteristics of surface water were taken as predictors of the model
 370

Physicochemical characteristics	OLS ratio	SAR ratio	Standard error
Physical characteristics			
Total suspended solids	-1,22	0,36	0,67
Temperature	10,41	16,28	5,90
Oxygen indicators			
Oxygen concentration	-0,46	-0,23	0,19
Degree of oxygen saturation	-0,03	-0,04	0,01
BZT ₅	-1,19	-1,77	0,37
ChZT-Mn	6,46	4,55	1,09
Biogenic indicators			
P-PO ₄ ³⁻	12,95	11,72	1,47



N-NH ₄ ⁺	11,40	16,2	1,78
N-NO ₃ ⁻	-0,03	4,88	0,36
N-NO ₂ ⁻	0,64	14,83	2,71
Salinity indicators			
*CZSR	5,70	8,53	0,27
SO ₄ ²⁻	-0,31	0,18	0,14
Cl ⁻	0,06	0,71	0,02
Na ⁺	0,49	-0,61	0,17
K ⁺	2,92	10,95	1,06
Mg ²⁺	-0,46	6,32	0,82
Ca ²⁺	-2,93	-4,41	1,41
Electrical conductivity	-1,19	4,83	-1,72
Metals, including heavy metals			
Zn ²⁺	-0,14	0,14	0,02
Pb ²⁺	-0,62	3,12	0,09
Cd ²⁺	-0,03	4,89	0,83
Cu ²⁺	0,98	3,12	0,56
Fe _{og}	6,46	12,81	1,37
Mn ²⁺	17,46	11,27	1,45

371 *CZSR – Total content of dissolved substances

372

373 Table 8. Results of the spatial autoregression model of the SAR type (The transverse turbulence
 374 diffusion coefficient was used as the dependent variable, and the most important
 375 physicochemical characteristics of surface water selected on the basis of principal component
 376 analysis for each use variant were used as predictors)

377

Variable	SAR ratio	Standard factor	Standard error	t	p
Survey points in the watercourse next to arable land					
The constant of the equation	9,46	-	5,35	0,23	3,50
K ⁺	-0,85	0,82	0,83	-4,06	0,004
pH	-2,50	-0,39	5,34	-2,25	0,32
N-NO ₂	0,009	0,05	0,02	0,36	0,78
Ca ²⁺	-0,83	8,14	1,67	3,97	0,15
Mg ²⁺	-0,37	0,04	0,01	0,01	0,04
P-PO ₄ ³⁻	-2,96	0,02	0,02	-2,05	0,006
Na ⁺	25,75	4,77	1,34	-4,06	4,95
Total suspended solids	2,56	0,32	0,34	3,72	10,03
Measurement points in the watercourse next to grasslands					
The constant of the equation	1,19	-	1,79	0,66	1,45
Temperature of water	0,04	0,01	0,42	0,10	0,51
The constant of the equation	0,12	0,02	1,27	0,09	0,92
Fe _{og}	-0,02	-0,45	0,006	-4,00	0,01
K ⁺	0,35	0,003	12,10	3,97	0,004
Ca ²⁺	-7,82	-0,09	0,30	0,01	0,99
Na ⁺	-1,95	-0,11	0,17	-2,05	0,07
Temperature of	-2,96	-0,81	0,93	-1,58	0,15



water					
Measurementpoints in the watercourse near the forests					
The constant of the equation	-	-0,90	0,24	-3,49	0,01
Temperature of water	-0,03	-0,03	0,07	-0,45	0,66
ChZT-Mn	-0,11	0,10	0,15	-0,72	0,49
Na ⁺	-0,04	-0,04	0,29	-0,51	0,88
Conductance	-0,03	0,09	0,27	0,53	0,56
BZT ₅	0,008	0,009	0,22	0,03	0,97
Mg ²⁺	-0,20	-0,13	0,09	-2,12	0,07
Dissolved oxygen	-0,46	0,003	0,004	0,21	0,99
Fe _{org.}	-0,03	0,19	0,06	0,15	0,24

378

379 4. Discussion

380 An innovative solution proposed in this dissertation is the combination
 381 of multidimensional statistics methods and a spatial model. During the preparation of the results,
 382 a close relationship was determined between selected physico-chemical characteristics of surface
 383 water quality using the spatial autoregression equation.

384

385 4.1. Influence of hydrodynamic parameters on the ionic composition of surface water

386 Determining the level of micro-pollutants, especially heavy metals, is essential for
 387 monitoring areas exposed to contamination in long-term studies. The ionic composition of
 388 surface water determines the intensity and direction of the chemical denudation process
 389 conditioned by lithology (Kim et al., 2012; Ziadat et al., 2015). It also determines the quality
 390 of the surface watercourses inventory and the hydrographic network's operation (Trivedi, 2010).
 391 Through PCA, the parent variables for each type of land use were documented. For further spatial
 392 analysis, the physico-chemical features of surface water with the highest factor loadings were
 393 selected. A high negative correlation was only calculated in the first PCA factor for the
 394 concentration of K⁺ and in the second PCA factor for concentrations of Mg²⁺ and Na⁺. This may
 395 be because these ions runoff from the catchment area.

396 Water self-purification processes are essential in the assessment and protection process
 397 (Moore and Langner, 2012; Halecki, 2015). Identifying the sources of nitrate pollution is one of
 398 the priority actions (Bu et al., 2017) because anthropogenic pollutants are different in surface
 399 water and may change seasonally (Xu et al., 2014; Hu et al., 2015). As for arable land, K⁺



400 showed a negative relationship with the first factor of the PCA. A negative relationship was also
401 found for Na^+ . This may mean that their concentration has been systematically decreasing. In
402 water samples collected near forested areas, a high factor load was calculated for the water
403 temperature and concentration of ChZT-Mn. They were also characterised by the highest
404 variance for the factor axes in all tested surface water samples and the highest factor loadings.
405 The water temperature could play a significant role in the concentration of ions dissolved in the
406 water of the flysch stream.

407 The state when the material was mixed in the water was considered in processing the
408 results. For this purpose, the transverse turbulence diffusion index was calculated to determine
409 transport intensity. The highest value of transport intensity, 1.02, was recorded in the riverbed of
410 the stream adjacent to arable land. According to PCA, the most important chemical parameter
411 associated with this land use type is Ca^{2+} and P-PO_4^{3-} . This is an interesting result because, when
412 assessing the impact of arable lands on the physico-chemical quality of surface water, particular
413 attention should be paid to the supply of phosphates and nitrates (Halecki et al., 2017). The
414 quantitative share of ions depends on the type of rocks and the physical properties of the
415 substrate, especially the infiltration coefficient, and for clastic weathering, on their mechanical
416 composition (Szostakiewicz-Hołownia, 2012). This aspect is of particular concern for flysch
417 catchments built of the Magura and Sub-Magura layers (Starkel, 2006; Starkelet. al., 2007).

418 The diversified rock bottom relief impacts high water turbulence and, thus, the transport of
419 weathered loads (Priess et al., 2015; Szatten, 2016). Stationary tests with a high measurement
420 frequency are then performed to determine the relationship between flooding and suspension
421 concentration in particular years (Huan, 2011). On the other hand, in the Beskidy catchments,
422 which include the variability of the transported clastic weathered water, quantitative relationships
423 between hydrological parameters and their supply against the background of water erosion have
424 been shown (Brański, 1975; Brański, 1990). The most significant distance of full mixing was
425 calculated for points in surface water near arable land at 52.26 m.

426 **4.2. Relationships between hydrological parameters and their supply against the** 427 **background of water erosion**

428 The intensity of transport of chemical compounds depends on turbulence and dispersion
429 due to spatial differentiation of the flow velocity. Turbulent vortices create a locally



430 inhomogeneous and non-stationary velocity field, accelerating mixing (Loga, 2016). This
431 hydrodynamic parameter was calculated for each cross-section on each section of the stream to
432 approximate the transport mechanism of the investigated ion forms and other materials dissolved
433 in water. The Reynolds number reached the highest values in the stream bed adjacent to the
434 permanent grassland, similar to the tangential stress. The bottom sediment ranged from 32 to 41
435 mm in the entire catchment area. It was observed that material with more significant fractions is
436 collected at the flysch bed surrounded by permanent grasslands.

437 **4.3. Model of developed techniques of spatial autoregression**

438 Ions in water may be related to land use, and the detection of heavy metals in surface
439 water determines its economic value (Fu et al., 2014; Liao et al., 2016) and the health of the
440 inhabitants of the surrounding areas, especially in terms of therapeutic properties of Carpathian
441 rivers (Operacz et al., 2018). The spatial pattern and correlations between the variables in the
442 entire catchment area have been shown based on the spatial autoregression model, empirical
443 equations, and multivariate analysis. Furthermore, the research confirms that in a small flysch
444 catchment, it is possible to balance the material leached from the soil because the chemical
445 denudation process shows a relationship with the concentration of dissolved electrolytes in the
446 flysch stream (Halecki et al., 2019).

447 The model of developed techniques of spatial autoregression also enabled the inclusion of
448 spatial analysis in the extrapolation of hydrochemical data and the intensity of erosion processes
449 during multiple measurement series. The technique of spatial autoregression was chosen for
450 its more effective presentation of these relationships and was used to verify the analysed data in a
451 spatial system. The applied analysis helped to check whether the variables are affected by
452 different land use modifications. The independent variables were the physico-chemical features of
453 surface water at different land use variants. Spatial SAR model results showed that P-PO_4^{3-} and
454 K^+ ions are leached mainly from soil in arable lands. Moreover, a statistically significant
455 relationship between K^+ and Fe_{og} cations was observed in surface water flowing from grasslands.
456 However, the SAR model of spatial autoregression did not detect statistically significant
457 differences in the ionic composition of surface water flowing out of the forests. The land use
458 could have contributed to the differences generated by the model. Therefore, the methods used to



459 evaluate the surface washout and soil erosion studies on used slopes, especially for agricultural
460 purposes, should be improved (Bakker et al., 2008; Smolska et al., 2010).

461 The spatial autoregression model for the whole with the explanatory variable (the distance
462 of full mixing) detected the spatial dependencies for P-PO_4^{3-} , Mn^{2+} , N-NO_2^- , TDS, Na^+ , Mg^{2+} ,
463 Ca^{2+} and Mn^{2+} . Since different hydraulic conditions were detected for each variant of land use, it
464 should be passed that the direction of the catchment development impacts the physico-chemical
465 quality of surface water.

466 4.4. Recommendations for hydraulic tests and physico-chemical assessment of surface water

467 Forecasting the physico-chemical quality of surface water and determining the
468 relationships between the concentration of metals, including heavy metals, requires the
469 determination of decisive indicators. Spatial autoregression is currently one of the most effective
470 methods of water quality forecasting for various sources of pollution (Yang et al., 2017). The
471 analysis of spatial autoregression primarily showed the spatio-temporal relationships at various
472 measuring points. Furthermore, including spatial analysis to determine many pollutants flowing
473 into the waters allows for solving technical problems related to surface washout and proper use of
474 arable land in the catchment area.

475 For spatial studies, the appropriate number of samples plays a crucial role in building a
476 spatial model (Vallejos and Osorio 2014), especially in work related to the physico-chemical
477 quality of surface water (Mahjouri and Kerachian, 2011). This means that the assessment of
478 surface water quality features depends on the spatial scale. Therefore, the evaluation of surface
479 water quality should be combined with a spatial analysis (Dash et al., 2015). The amount of
480 leached total suspended solids delivered during water erosion is a value recorded during short-
481 term flooding rainfalls. So, based on the hydrochemical properties, it is possible to determine
482 surface water quality (Misaghi et al., 2017). Only relevant parameters should be considered in
483 assessing the physico-chemical quality of surface water in agricultural areas and grasslands. It is
484 vital to choose a method consistent with the sample size (Griffith, 2005) in order to formulate a
485 dynamic model (Finley et al., 2012) or to calculate water quality indices (Yadav et al., 2015;
486 Naubi et al., 2016). Thus, the optimisation of spatial sampling is crucial (Hu and Wang, 2011) in
487 monitoring surface water pollution (Wu et al., 2005). The conclusions from this dissertation will
488 find practical application in isolating the primary factors favouring water erosion in streams with



489 larger catchment areas. In addition, recognising the effects of water erosion will be necessary for
490 evaluating the surface delivery of the material in further scientific research. The continuation of
491 scientific work in this field should focus on investigating the relationship between individual
492 washings and the concentration of material in the valleys of streams and rivers.

493 5. Conclusions

494 The effect of this work is important for determining the sources of diffuse pollution
495 flowing to the stream from agricultural areas in the periodic assessment of the physico-chemical
496 quality of surface water. The research thread undertaken in the dissertation will help determine
497 the intensity of water erosion and assess the sources of surface water supply. The assessment of
498 this phenomenon should aim at explaining the size of the load calculated into the delivery zones
499 of the transported material to the flysch stream and erosive factors, including susceptibility to
500 leaching. Some results will undoubtedly be used to create models of erosive feeding of alluvial
501 deposits from weathered flysch or surface wash, depending on the material delivery in the
502 catchment area. The results related to land use and its impact on the transport of chemical
503 compounds may be helpful in environmental management in mountain areas and chemical
504 monitoring in determining the course of erosion processes. Over the years, assessment of land use
505 changes may be of value for local residents involved in agriculture. Moreover, the assessment of
506 the short-term change in the concentration of the physico-chemical composition in the flysch
507 riverbed will be valuable for scientific purposes to develop a plan for the management and
508 elimination of pollutants resulting from anthropogenic pressure.

509 **Highlights:**

- 510 1. The physico-chemical composition of the flysch stream depends on the hydraulic
511 conditions.
- 512 2. The catchment use determines the intensity and direction of the chemical denudation
513 process.
- 514 3. The turbulent diffusion coefficient may influence the dissolved electrolytes in the
515 stream.
- 516 4. The spatial autoregression model detected relationships for nutrients leached from
517 arable land.



518 5. The assessment of the physico-chemical properties of surface water depends on the
519 spatial scale.

520

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524

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