# A Bayesian updating framework for calibrating <u>the</u>hydrological parameters of road network<u>s</u> using taxi GPS data

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Abstract. Hydrological parameters should pass through a careful calibration procedure before aiding decision\_-making. However, great-significant difficulties difficulty is are encountered when applying existing calibration methods to regions in whichwhere runoff data are inadequate. To achieve accurate fill the gap of hydrological calibration for the ungauged road networks, we proposed a Bayesian updating framework thatto calibrates hydrological parameters based on taxi GPS data. Hydrological parameters wereare calibrated by adjusting their values such that the runoff generated by the acceptable parameter sets corresponds to could yield the road disruption periods during which no taxi points are observed. The proposed method wasis validated on through 10 flood-prone roads in Shenzhen, and the results revealeds that the trends of runoff could be correctly predicted for 8 out of 10 roads. This study shows-demonstrates that the integration of hydrological models and

20 taxi GPS data <u>can provide suggests</u>-viable alternative measures for the-model calibration to derive, and provides actionable insights for flood hazard mitigation.

## **1** Introduction

<u>In the context Under the background</u> of climate change and increased urbanization, flooding <u>poses is posing</u> farreaching threats to <u>the</u> urban road network<u>s</u> of coastal <u>metropolises metropolis</u> (Balistrocchi et al., 2020). In Australia,

- 25 around-approximately 53% of flood-related drowning deaths were the result of <u>vehicles</u> driving into flooded waters. <u>Additionally, i</u>Indirect losses caused by flooding, such as cancelled commutes, mandatory detours, and travel time delays, <u>even-often</u> outweigh direct losses (Kasmalkar et al., 2020). Quantifying the impact of flooding exposure requires the prediction of surface runoff over the roads and computation of road disruptions induced by the runoff, which are critical for the implementation of to flooding mitigation, traffic resilience improvement, and <u>risk-early warning systems</u>.
- 30 Public concerns about regarding road flooding hazards <u>have</u> created <u>the pressing needpressure</u> to develop fine-grained and accurate models for hydrological simulation. <u>HThe hydrological modelling is</u>, as a <u>quite-relatively</u> well-established

theory <u>that can</u>, provides an approximations of the real-world hydrological system, and has been widely used in many roadrelated studies (Versini et al., 2010; Yin et al., 2016; Safaei-Moghadam et al., 2022). As the Because hydrological modelling is subject to uncertainty, which that arises from the over-simplifiedfalse reflection of the hydrological systems, the initial

- 35 and boundary conditions, and the-lack of true knowledge, parameters forof hydrological models should must be carefully calibrated prior to their before applying application to solve practical problems, so that the models can is capable of closely matching the historical evidences-trends (Gupta et al., 1998). Asn un-calibrated models are is indefensible and sterile, so-very few models documented in the literature have been applied without any calibrationsa calibration procedure (Beven, 2012).
- 40 During-Over the last-past four decades, numerous studies have been <u>conducted on devoted to</u> the development of calibration methods. Methodologies <u>forof</u> model calibration range from <u>the</u>-simple trial-and-error <u>methods</u>, <u>which-that</u> adjusts one parameter value <u>in</u> each <u>turn-iteration</u> until the differences between predicted and observed values <u>areis</u> satisfactory, to <u>the</u> Bayesian updating framework, <u>which-that</u> rejected the <u>concept of a idea that there is</u> single correct solution. To a great extent, <u>No matter what kinds of methods</u>, <u>hydrological models are basically calibrated based on the</u>
- 45 runoff data alone (Dembélé et al., 2020), so the success of model calibration is\_, to a great extent, dominated by the availability of field-observed runoff data. However, runoff data are generally only gathered at <u>a only</u> few sites, and some cities even never measured runoff data in the built-up regions (Gebremedhin et al., 2020). <u>AlEven</u> though the runoff data could can be effectively collected by administration departments in some cities, they haved no motivations to share the se data to with the public. For example, among China's top 10 largest cities<sup>1</sup>, only Shenzhen has shared runoff-related data on
- 50 the <u>an</u> open data platform. <u>FAs</u> for the model calibration for <u>at</u> the road scale, the runoff data are even more difficult to acquire, because a road networks <u>are is</u> far denser than a river networks and flood gauges are only <u>located installed</u> in a few flood-prone roads <u>considering based on</u> their high measurement cost, leaving most of roads <u>ungagedungauged</u>. As pointed out by Beven (2012, p:55), "the ungauged catchment problem is one of the real challenges for hydrological model ers."

Thise lack of hydrological data <u>has</u> prompted researchers to seek <u>extra additional</u> data sources to support flood-related decision\_-making. <u>In response to this need, big data, owing toBased on</u> the advance<u>ment</u> of mobile telecommunication technologies, <u>big data</u> are emerging as alternative sources of information for coping with flood risks (Paul et al., 2018; Li et al., 2018; Gebremedhin et al., 2020). Citizens <u>can</u> voluntarily or passively <u>acting act</u> as human sensors <u>to</u> generate georeferenced data to improve flood monitoring. <u>Typical Many</u> studies <u>have leveraged</u> involve the use of crowdsourceding social media data-(Brouwer and Eilander, 2017; Li et al., 2018) (Brouwer and Eilander, 2017; Sadler et al., 2018; Zahura et

60 al., 2020), mobile phone data (Yabe et al., 2018; Balistrocchi et al., 2020), and taxi GPS data (She et al., 2019; Kong et al., 2022). However, most of-previous works <u>have</u> concentrated on using big data either for flood mapping or mining spatiotemporal patterns (Restrepo-Estrada et al., 2018), and-<u>parameter calibration for ungauged roads based on big data for ungauged roads.</u>

<sup>&</sup>lt;sup>1</sup> Rank by the resident population in 2021.

This study differentiates extends from our previous study (Kong et al., 2022) by going one a step further than simply

- 65 recognizing the floodeding roads. We propose a calibration method for road-related hydrological parameters using based on the taxi GPS data. Many studies have shown that vehicle-related information during the rainfall, such asincluding vehicle volume, speed, and trajectory information, is usefuleritical forto floodeding road detection (Zhang et al., 2019; Qi et al., 2020; Yao et al., 2020). When a road segment is inundated by the heavy rainfall, the vehicle volume may present exhibit a sharp or gradual drop; depending on the intensity of the rainfall event. Conversely, anthe abnormal drop inof vehicle volume
- 70 during the rainfall <u>may</u> implyies that the <u>a</u> road <u>may has</u> experienced <u>some</u> rainfall-induced inundations. This motivatesd us to use <u>a</u> traffic-related data sources to calibrate hydrological parameters. In this study, we developed a transformation process <u>which that</u> converts the rainfall time series <u>data in</u> to the <u>a</u> time series of the probabilities that no taxis <u>will</u> drive through on the <u>a</u> road (<u>no-taxi-passing probability hereafter</u>) for every <u>a given</u> hydrological parameter set., <u>and-We</u> then assigned a probability to every parameter set by integrating the no-taxi-passing probability with the observed taxi GPS data.
- 75 We not only-outlined a generalized taxi-data-driven calibration framework <u>and but also realize implemented</u> the <u>a</u> framework with specific hydrological and transportation models.

## 2 Methodology

## 2.1 A-Bayesian updating procedure

Observed data are not always as informative as expected and may be inconsistent with other data sources; so hydrologists <u>usually-typically</u> adopt the Bayesian framework to update hydrological parameters, which provides a generalized formalism that integrates prior probability representing the prior knowledge with the likelihood that reflects how well-accurately the presumeda model can reproduce the observations to form the a posterior probability. Suppose we have several hydrological models, each with different sets of parameters. Then, the purpose of the Bayesian updating procedure adopted in this study is to assign a posterior probability to every hydrological parameter set as new taxi data become 85 available.

Two components are critical for the this Bayesian updating procedure. One is the prior probability, and the other is the likelihood function. RegardingFor the prior probability, for their famous calibration model called generalized likelihood uncertainty estimation, Beven and Binley (1992) stated in their famous calibration model, generalized likelihood uncertainty estimation (GLUE), that all the parameter combinations are considered equally probable before extra additional information

90 is introduced. After the first updat<u>e</u>ing, the prior probability of each updating <u>run-iteration\_could\_can</u> be replaced by the posterior probability of the latest updating <u>iteration</u>. <u>L</u>The likelihood, <u>which isas</u> a measurement of how well the <u>a</u> given model conforms to the observed taxi behaviour, is not as easy to compute as the prior probability, because the parameter set to be estimated is hydrology\_-related, wh<u>ereasile</u> the observed evidence is taxi-<u>based</u>related. The<u>refore</u>, <u>question</u> then arises <u>as-we must determine</u> how to construct a taxi-based proxy whose probability <u>equalsis equal to</u> to that of the associated

95 hydrological parameter and construct a function enabling the transformation from the hydrological parameters to the taxirelated proxiesy.

The <u>selected</u>-proxy <u>selected</u> in this study <u>wasis</u> the time series of <u>probability that no taxi drives through the road in a</u> given time interval (short for the no-taxi-passing probability). Figure 1 <u>illustrates presents</u> a generalized procedure <u>for of</u> converting the <u>a precipitation rainfall process time series in</u> to the <u>a</u> time series of no-taxi-passing probabilit<u>ies</u> for each

- 100 hydrological parameter. Thise procedure consists of three steps. First, a hydrological model is used to convert the <u>a</u> rainfall time series inprocess to the <u>a</u> hydrograph, which is a graph showing runoff with respect to time past a specific point. Second, a runoff-disruption function that, which relates the runoff to the probability that the <u>a</u> road is blocked, is used to transition transform the hydrograph into the <u>a</u> time series of road disruption probabilitiesy. Third, the taxi arriving arrival rate is combined with the time series of road disruption probabilitiesy to yield derive the <u>a</u> time series of no-taxi-passing
- 105 probabilit<u>iesy</u>. Note that tThe hydrological model and the taxi <u>arrivalarriving</u> rate are considered to be unique for every road and <u>are invariable within a short period</u>, while whereas the runoff-disruption function is identical for all roads.



Figure 1 A <u>gG</u>eneralized procedure <u>for</u>of converting <u>the a</u>rainfall time series <u>in</u>to <u>the a</u> time series of no-taxi-passing probabilit<u>iesy.</u>

Integrating thise three-step process with the Bayesian equation enables us to compute the posterior probability of the a parameter set based on the taxi data. For a specific road, suppose there are *N* hydrological parameter sets to be estimated. As <u>Because</u> the runoff-disruption function and the taxi <u>arrivalarriving</u> rate are assumed to be <u>invariant fixed</u> for the road, we can construct a composite function converting the *i*th parameter set, <u>which is</u> denoted as  $\theta^{(i)} - \theta^{(t)}$ , <u>in</u>to the <u>a</u> time series of no-taxipassing probabilitiesy, which is denoted as  $\Omega^{(i)}$ . Therefore, the probability of  $\theta^{(i)}$  to be being optimal is equals to the probability of  $\Omega^{(i)}$  beingto be true, which can be expressed as follows:

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$$P(\boldsymbol{\theta}^{(i)}) = P(\boldsymbol{\varOmega}^{(i)}) \tag{1}$$

where  $P(\boldsymbol{\theta}^{(i)})$  and  $P(\boldsymbol{\Omega}^{(i)})$  are the prior probabilities of  $\boldsymbol{\theta}^{(i)}$  and  $\boldsymbol{\Omega}^{(i)}$  respectively. As taxi observations become available,  $P(\boldsymbol{\theta}^{(i)})$  (or  $P(\boldsymbol{\Omega}^{(i)})$ ) can be updated using the Bayes <u>Ftheorem as</u>:

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$$-P(\theta^{(i)}|X) = P(\Omega^{(i)}|X) \propto P(\Omega^{(i)})\mathcal{L}(X|\Omega^{(i)}) (2)$$
$$P(\theta^{(i)}|X) = P(\Omega^{(i)}|X) \propto P(\theta^{(i)})\mathcal{L}(X|\theta^{(i)})$$
(2)

where X is the taxi observation, and  $P(\boldsymbol{\theta}^{(i)}|X)$  and  $P(\boldsymbol{\Omega}^{(i)}|X)$  are the posterior probabilities of  $\boldsymbol{\theta}^{(i)}$  and  $\boldsymbol{\Omega}^{(i)}, \boldsymbol{\Omega}^{(i)}$  respectively under the conditional on the of taxi observation, X. The  $\mathscr{L}(X|\boldsymbol{\Omega}^{(i)})$   $\mathcal{L}(X|\boldsymbol{\theta}^{(i)})$  is the likelihood of X given  $\boldsymbol{\Omega}^{(i)}, \boldsymbol{\theta}^{(i)}$ . The optimal parameter set is the one that which derives yields the  $\boldsymbol{\Omega}^{(i)}$  that most closely best fits the observed taxi data.

125 The solution of <u>Solving</u> Eq. (2) involves the calculation of  $P(\Omega^{(i)}) P(\theta^{(i)})$  and  $\mathscr{L}(X|\Omega^{(i)}) \mathcal{L}(X|\theta^{(i)})$ . According to Eq.(1),  $P(\Omega^{(i)})$  can be replaced with  $P(\theta^{(i)})$ , which is the prior probability of parameter sets. The derivation of  $P(\theta^{(i)})$ 

depends on prior knowledge regarding the parameter distribution, which is typically obtained using traditional hydrological methods. However, this prerequisite knowledge may not always be readily accessible based on limited data availability. In such cases, Beven and Binley (1992) suggested that prior to the introduction of any quantitative and qualitative information, any parameter set combination should could be considered to be equally likely. This impliesed that the parameter set is

drawn from a uniform distribution <u>as follows</u>:

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 $-P(\Omega^{(i)}) = P(\theta^{(i)}) = 1/N - (3)$ 

$$P(\boldsymbol{\theta}^{(i)}) = 1/N \tag{3}$$

In this study, we compared the effects of two types of prior parameter distributions, namely a uniform distribution and a distribution derived from digital elevation model (DEM) data, on the resulting posterior distributions.

Next,  $\mathcal{L}(\mathbf{X}|\boldsymbol{\theta}^{(i)}) \mathscr{L}(\mathbf{X}|\boldsymbol{\Omega}^{(i)})$ , which is as a likelihood function, describes the joint probability of the observed taxi data,  $\mathbf{X}$ , as a function of the chosen  $\mathcal{Q}^{(i)} \boldsymbol{\theta}^{(i)}$ . Consider a rainfall event that is broken divided into T 5 min intervals. From the taxi data, we obtained can obtain a sequence of taxi-related observations, which are denoted as  $\mathbf{X} = \{\mathbf{x}X_1, \mathbf{x}X_2, \dots, \mathbf{x}X_T\}$ , where  $\mathbf{x}X_t = 1$  if the observed road has is observed with at least one taxi passing during by in the tth 5 min interval, and  $\mathbf{x}X_t = 0$  otherwise. The  $\mathbf{\Omega}^{(i)} = \{\omega_1^{(i)}, \omega_2^{(i)}, \dots, \omega_T^{(i)}\}$  is also a T-dimensional vector, where  $\omega_t^{(i)}$  is the no-taxi-passing probability at in the tth 5 min interval taking with  $\boldsymbol{\theta}^{(i)}$  as the parameter set. Note that  $\mathbf{\Omega}^{(i)}$  is only determined by the chosen hydrological parameter and the rainfall process time series, and is not measured from the observed data. Considering that the arrival of taxis is independent of with respect to time,  $\mathcal{L}(\mathbf{X}|\boldsymbol{\theta}^{(i)}) \mathscr{L}(\mathbf{X}|\boldsymbol{\Omega}^{(j)})$  can be formulated as:

$$-\mathcal{L}(X|\Omega^{(i)}) = \prod_{t=1}^{T} (1-\omega_t^{(i)})^{X_t} (\omega_t^{(i)})^{1-X_t} - (4)$$

$$\mathcal{L}(\boldsymbol{X}|\boldsymbol{\theta}^{(i)}) = \mathcal{L}(\boldsymbol{X}|\boldsymbol{\Omega}^{(i)}) = \prod_{t=1}^{T} (1 - \omega_t^{(i)})^{x_t} (\omega_t^{(i)})^{1 - x_t}$$
(4)

By substituting Eq.(3) and Eq.(4) back-into Eq.(2), the Eq.(5) is following equation can be obtained:

$$- P(\boldsymbol{\theta}^{(i)}|\boldsymbol{X}) \propto \frac{1}{N} \prod_{t=1}^{T} (1 - \omega_{t}^{(i)})^{X_{t}} (\omega_{t}^{(i)})^{1 - X_{t}} (5)$$

$$P(\boldsymbol{\theta}^{(i)}|\boldsymbol{X}) \propto P(\boldsymbol{\theta}^{(i)}) \prod_{t=1}^{T} (1 - \omega_{t}^{(i)})^{X_{t}} (\omega_{t}^{(i)})^{1 - X_{t}} (5)$$

- Equation (5) is the proposed Bayesian updating model <u>forto</u> calibratinge the hydrological parameters based on the taxi data, where X <u>could-can</u> be directly measured and  $\Omega^{(i)} \omega_t^{(i)}$  is calculated through the three-step process shown-<u>illustrated</u> in Fig. 1, which will be discussed in detail in the <u>next-following</u> section. Having <u>chosen-selected</u> an updating model, the optim<u>alum</u> parameter for one period of observations may not be optimal for another period. <u>As-Because</u> the model may have a-continuing inputs of new taxi observations, the posterior probability for  $\theta^{(i)}$  should be updated as new evidence becomes
- available. For the second update, the posterior <u>probability</u> from the first observation becomes the prior <u>probability</u> for the second observation, and the posterior probability for  $\theta^{(i)}$  is recursively updated as:

$$P(\boldsymbol{\theta}^{(i)}|\boldsymbol{X}_2) \propto \mathcal{L}(\boldsymbol{X}_2|\boldsymbol{\theta}^{(i)}) P(\boldsymbol{\theta}^{(i)}|\boldsymbol{X}_1)$$
(6)

where  $X_1$  and  $X_2$  are the first and the second taxi observations.

### 2.2 Instantiation Instantization of the three-step procedure

160 Section 2.1 presenteds a generalized three-step procedure which for convertings the a rainfall time series into the a time series of no-taxi-passing probabilitiesy. In this section, we specialize thise process by integrating existing theories with our model. The tThree conceptualized steps shown illustrated in Fig. 1 are were replaced substituted with three more concrete sub-models. Firstly, a Soil Conservation Service (SCS) unit hydrograph wasis used to turn-convert the rainfall excess into the a hydrograph of the target road. Secondly, an empirical runoff-disruption function based on , whose data extracted from various experimental, observational, and modelling studies, is was applied to convert the hydrograph into the a time series of the road disruption probabilitiesy. Thirdly, a Poisson distribution, representing the distribution of taxi arrivalarriving rate, is was combined with the road disruption probability time series to vield derive the a no-taxi-passing probability time series.

# Step 1: Converting rainfall into runoff based on the SCS unit hydrograph

Not all rainfall will-produces runoff because storage from soils storage can absorb a certain amount of rainlight shower.
 While-However, in the-urbanized areas, only a small proportion of rainfall infiltrates into the soil or is retained on the land surface, leaving most of themrain to flow across the urban surfaces and becomes become the direct runoff. The rainfall that yields becomes the direct runoff is termed referred to as rainfall excess. The Natural Resources Conservation Service

 $(NRCS)^2$  developed a method to estimate the rainfall excess based on the soil types and land uses using the <u>following</u> curve number equation:

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$$P_e = \begin{cases} (P_a - 0.2S)/(P_a + 0.8S) & P_a > 0.2S \\ 0 & P_a \le 0.2S \end{cases}$$
(7)

where  $P_e$  is the accumulated rainfall excess in cm,  $P_a$  is the accumulated rainfall in cm, and S is the potential retention after runoff begins, which is <u>defined as</u> a function of the curve number <u>as follows</u>:

$$S = 2.54 \times (1000/CN - 10) \tag{8}$$

where *CN* is the curve number. For urban and residential land, the curve number varies from 65-40 to 8595, depending on
 the impervious areas (Natural Resources Conservation Service, 2010a). Because prior knowledge on the *CN* is unavailable, it was considered as a calibrated parameter in this study. For sake of brevity, the curve number will not be regarded as a parameter to be calibrated in this study but as a given parameter with the value of 85.

Next, tThe rainfall excess derived usingby Eq. (7) wasis inputted into the unit hydrograph to produce-derive the runoff. The unit hydrograph is a commonly used rainfall-runoff model that converts rainfall excess into a temporal distribution of
 direct runoffconverts rainfall excess to direct runoff. First proposed by Sherman in 1932, the unit hydrograph is defined as the hydrograph resulting from one unit of rainfall excess distributed uniformly over a catchment area. It assumes the that rainfall is uniform over the catchment area and that runoff increases linearly with the rainfall excess. Although under most conditions these assumptions cannot be perfectly satisfied under most conditions, the results obtained from the unit hydrograph are generally acceptable for most practical usescases. The model, was originally designed for larger watersheds,

190 but it has been found to be applicable to some catchments areas less than 5,000 m<sup>2</sup> in size (Chow et al., 1988).

The unit hydrograph applies is only applicable to for the watershed areas where the runoff data arewere measured. The paucity of the runoff data motivated sparkled the idea development of the synthetic unit hydrograph (SUH) concept. The term "synthetic" in SUH denotes the refers to a unit hydrograph derived from watershed characteristics, rather than empirical rainfall-runoff relationships. In this study, we utilized the SCS unit hydrograph, which is a dimensionless SUH proposed by the NRCS. For the dimensionless SUH, the discharge (i.e., *y*-axis) is expressed as the ratio of discharge *q* to the peak discharge  $q_p$  and the time (i.e., *x*-axis) is expressed as the ratio of time *t* to the peak time  $t_p$ . Therefore, the SCS unit hydrograph, rigorously speaking, is not exactly an SUH itself, but is a useful tool for constructing an SUH.

The shape of <u>an</u> SCS unit hydrograph is totally <u>entirely</u> determined by the peak rate factor. A standard value of 2.08 for the peak rate factor is recommended and commonly used by the NRCS (Fig.\_2). To construct <u>an</u> SUH from the <u>an</u> SCS unit hydrograph, the *x*-axis of the SCS unit hydrograph is multiplied by  $t_p$  and the *y*-axis <u>is multiplied</u> by  $q_p$ . The values of  $q_p$ and  $t_p$  are functions of the catchment area and the time of concentration <u>as follows</u>:

$$t_p = 0.6t_c + D/2$$
(9)

$$q_p = 2.08A/t_p \tag{10}$$

<sup>&</sup>lt;sup>2</sup> The NRCS used to be was originally called the US Soil Conservation Service (SCS).

where  $t_c$  is the time of concentration in hours, A is the catchment area in km<sup>2</sup>, and D is the duration of unit rainfall excess in

- 205 hours, which iwas set to 1/12 h one-twelfth of an hour (i.e., 5 min) in this study. Notably As can be seen, the catchment area and time of concentration are required to construct an SUH, and they are the other two hydrological parameters we would that should be calibrated based on the taxi data. Although numerous tools and theories have been developed for estimating catchment area and time of concentration, these two parameters are still prone to significant errors, particularly in urban areas, because of challenges in accurately delineating urban catchments (Huang and Jin, 2019; Li et al., 2020). Urban
- 210 catchment delineation is more complex than natural catchment delineation. Urban catchments have spatially heterogeneous surface cover types, which change with city development and construction, and modify runoff parameters (Goodwin et al., 2009). High densities of residential and commercial buildings obstruct flow paths and alter flow directions of storm water runoff, complicating rainfall-runoff and overland flow processes in urban areas (Ji and Qiuwen, 2015). Additionally, accurate urban catchment delineation necessitates high-resolution DEMs, which are not always available. In many Chinese
- 215 <u>cities, high-resolution DEMs are considered confidential data and are generally inaccessible to non-governmental</u> organizations. Based on these challenges, deriving accurate catchment area and time of concentration data in urban areas is <u>difficult in Shenzhen.</u>

For <u>the</u> sake of simplicity, the peak rate factor <u>would was</u> not <u>be</u> calibrated and <u>be was fixed as at</u> 2.08, although some studies have <u>showed-indicated</u> that it has a <u>much-wide</u> range from 0.43 for steep terrain to 2.58 for very flat terrain (Chow et

al., 1988). After  $t_c$  and A wereare chosen, an SUH could can be constructed, and then we used it to convert the rainfall excess into the runoff by applying the discrete convolution equation. The detailed computation process of the discrete convolution equation can be found in most hydrological textbooks (e.g., see Chow et al., (1988)-pp: 211-213), and will not be discussed here. To be clear, a The graphical workflow in Fig. 3 shows-illustrates the transformation of how the rainfall time series is data transformed into the a hydrograph for every parameter set.





The goal of Step 2 is to convert the hydrograph generated in Step 1 into the a time series of road disruption probabilities, or more specifically, the probability that a taxi driver chooses to turn their car when arriving at a flooded road. Most models in the literature assume that a road is either open or closed, which usually does not correspond to the empirical evidence that many drivers may take risks to drive on inundated roadsalong the rood even though it is inundated. In order

<sup>t</sup>To transition from a binary view of a flooded road being considered "open" or "closed." Pregnolato et al. (2017) proposed 235 to the use of a curve that relates the depth of floodwater to a reduction in vehicle speed to show-indicate the probability of road disruption.- and such This idea wasis soon followed adopted by Contreras-Jara et al. (2018) and Nieto et al. (2021).

A driver will turn around when he believes that the flow rate is too risky for their overcomes the vehicle configuration. From this perspective, the road disruption probability is equals to equal to the probability that the vehicle performance is

- 240 lower-less than the flow rate perceived by a driver. However, it is a difficult task to quantify the factors that influence<del>common belief of what guide people's</del> willingness of people to drive through a flooded waterway roadway, and is also impossible difficult to obtain the precise knowledge of regarding all taxi-flood intersections. Alternatively, to ensure the vehicle stability in flood flows, guidelines are usually typically recommended based on the limiting values of depth times velocity., and mMany studies researchers have carried outconducted laboratory testing on the stability of different kinds
- 245 types of vehicle models exposed to different combinations of depth and velocity (Merz and Thieken, 2009; Shah et al., 2018). As sSuggested by Pregnolato et al. (2017), we constructed the our runoff-disruption function by integrating data from reviewed the literatures and some authoritative guidelines. In this study, the road disruption probability wasis defined as the probability that the product of flow velocity and flow depth is higher was greater than the stability limits extracteding from the literatureexisting studies, which are shown-listed in Table 1 and plotted in Fig. 4. The expression of the fitting curve is: y

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$$v = [1 + \exp(-16.6(x - 0.48)^2)]^{-1}$$
(11)

where x is the product of flow velocity and flow depth, and y is the disruption probability. According to Eq. (11), a road has a disruption probability of 50% when the product of flow velocity and flow depth is 0.47 m<sup>2</sup> s<sup>-1</sup>, and is totally disrupted when the product is higher-greater than 0.80 m<sup>2</sup> s<sup>-1</sup>. By Aapplying the fitting curve, we can easily convert the flood runoff into the disruption probability as follows:

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$$P(Disrupt)_{t}^{(i)} = \left[1 + \exp\left(-16.6\left(q_{t}^{(i)}/W - 0.48\right)^{2}\right)\right]^{-1}$$
(12)

where  $P(Disrupt)_{t}^{(i)}$  and  $q_{t}^{(i)}$  are the road disruption probability and discharge at in the *t*th interval 5-min\_derived by from the hydrological model with the parameter set  $\theta^{(i)}$ , respectively, and W is the road width.

_		<b>X7 1 • 1</b>	П (	Recommended limits for vehicle stability		
	Reference	Vehicle type	Feature	( <b>m</b> <sup>2</sup> <b>s</b> <sup>1</sup> )		
_	Shah et al. (2018)	Volkswagen Scirocco	Flow direction = $0^{\circ}$	velocity×depth<0.014		
	Al-Qadami et al. (2022)	Perodua Viva	Ground clearance =0.18	velocity×depth<0.39		

# Table 1 Guidelines recommended by in the existing literatures.





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Figure 4 Empirical runoff-disruption function derived from the existing literature.s

#### Step 3: Derivatione of the time series of no-taxi-passing probabilitiesy

A road has is considered to have no taxis passing by in a fixed time step-interval if the road has no taxis visiting arriving or if every taxi that arrives at the road turns around. So Therefore, the no-taxi-passing probability can be inferred calculated by using the following equation:

$$- \omega_t^{(i)} = \sum_{n=0}^{\infty} P(Arrival\_taxi=n) \times (P(Disrupt)_t^{(i)})^n - (13)$$
$$\omega_t^{(i)} = \sum_{n=0}^{\infty} P(Arrival\_taxi=n)_t \times (P(Disrupt)_t^{(i)})^n$$
(13)

where  $\omega_t^{(i)}$  is the no-taxi-passing probability in the *t*th <u>interval 5 min</u>, and <u> $P(Arrival_taxi = n)P(Arrival_taxi = </u>$ 

n) $(Arrival_t axi = n)_t$  is the probability that *n* taxis arrives at the road segment <u>during</u> the *t*th <u>5 minimterval</u>. Equation (13)

270 indicates that if every taxi <u>that</u> arrives at the road segment makes a turn because of the flooded <u>waterwayroadway</u>, <u>then the</u> taxi volume <u>of on</u> the road will be zero. In this study,  $P(Arrival_taxi = n)(Arrival_taxi = n)_t P(Arrival_taxi = n)$  wasis assumed to follow the Poisson distribution:

 $-\frac{P(Arrival taxi=n)}{e^{-\lambda_t}\lambda^n/n!}$ (14)

$$P(Arrival_taxi = n)_t = e^{-\lambda_t} \lambda_t^n / n!$$
(14)

where  $\lambda \lambda_t$  is the average number of taxis arriving at the road <u>during the *t*th interval</u>. By <u>Ss</u>ubstitut<u>ing</u> Eq. (14) into Eq. (13), we <u>deriveobtain</u>:

$$-\underbrace{\omega_{t}^{(i)} = \sum_{n=0}^{\infty} (e^{-\lambda} \lambda^{n} / n!) \times (P(Disrupt)_{t}^{(i)})^{n} - (15)}_{\omega_{t}^{(i)} = \sum_{n=0}^{\infty} (e^{-\lambda_{t}} \lambda_{t}^{n} / n!) \times (P(Disrupt)_{t}^{(i)})^{n}}$$

$$\underbrace{By}_{n=0} x^{n} / n!, \text{ Eq.}_{(15)} \text{ can be further converted into:}$$

$$(15)$$

$$280 \qquad \qquad \omega_t^{(i)} = e^{-\lambda} \sum_{n=0}^{\infty} \left( P(Disrupt)_t^{(i)} \lambda \right)^n / n! = \exp\left(\lambda \left( P(Disrupt)_t^{(i)} - 1 \right) \right) \qquad (16)$$

$$\omega_t^{(i)} = e^{-\lambda_t} \sum_{n=0}^{\infty} \left( P(Disrupt)_t^{(i)} \lambda_t \right)^n / n! = \exp\left(\lambda_t \left( P(Disrupt)_t^{(i)} - 1 \right) \right) \qquad (16)$$
Equation (16) indicates that  $\omega_t^{(i)}$  is totally entirely determined by  $\lambda_t$  and  $P(Disrupt)_t^{(i)}$ . Since Because  $P(Disrupt)_t^{(i)}$ 

Equation (16) indicates that  $\omega_t^{(0)}$  is totally <u>entirely</u> determined by  $\neq \lambda_t$  and  $P(Disrupt)_t^{(0)}$ . Since Because  $P(Disrupt)_t^{(0)}$  is <u>given-obtained through-from</u> Step 2, what is left to determine is the value of  $_{-\lambda}\lambda_t$ . The value of  $\lambda_t$  fluctuates according to the time of day, indicating higher taxi volume during congested periods and lower volume during non-congested periods. Therefore, we calculate  $\lambda_t$  by averaging the taxi volume during the *t*th interval to account for time-of-day variations. It should be noted that Aas the intensity of rain increases gets heavier, experienced taxi drivers will avoid flood-prone roads in advance, which meanings that strictly speaking,  $\neg \lambda \lambda_t$ , strictly speaking, is a decreasing function of rainfall intensity. However, fitting the rainfall- $\frac{1}{\lambda}\lambda_t$  curve requires substantial many taxi GPS trajectories to inspect the route choices of taxi drivers under heavy rain, which is <u>currently unfeasible inoutside the scope of</u> this study. We assume that  $\lambda$  is a constant quantity which keeps unchanged with respect to rainfall. The value of  $\lambda$  can be calculated by averaging all 5 min taxi volume using the historical taxi GPS data. Therefore, we assumed that  $\lambda_t$  was independent of rainfall.

Finally, Table 2 lists all the sub\_models and parameters <u>used in of</u> the three-step process. The core principle of the threestep process <u>iwas</u> to calculate the time series of no-taxi-passing probabilities  $\Omega^{(i)}$ , <u>givenfor</u> each parameter set,  $\theta^{(i)}$ .

As<u>Because</u> the best choice of a-model is often data\_specific, it is <u>probable\_likely</u> that the model combination proposed in this study is not optimal for other <u>studiesscenarios</u>. To apply the <u>proposed</u> calibration method in practice<u>al use</u>, one must specify the sub\_models for in the three-step process must be specified according to the available data, prior knowledge, and accuracy requirements.

Purpose of the each step	Specific model Parameter		Source of parameters	
	Curve number equation 1. Curve number		Parameters to be calibrated Existing literature	
Step 1: Convert <del>the </del> rainfall <u>data</u> <u>in</u> to <del>the a</del> hydrograph	SCS unit hydrograph	2. peak rate factor         2. Catchment           area         3. Time of concentration	Existing literatureParameters to be calibrated	
	SCS unit hydrograph	3. Catchment area           4. Time of concentration4. Peak           rate factor	Parameters to be calibrated <u>Existing literature</u>	
Step 2: Convert the hydrograph into the a time series of disruption probabilitiesy	Empirical runoff-disruption function	5. Limit of product of flow velocity and depth	Existing literature <del>s</del>	
Step 3: Convert the time series of disruption probabilitiesy into the a time series of no-taxi probabilitiesy	Taxi arrival rate followings the Poisson distribution	6. Average taxi volume <u>s</u> in 5 min <u>periods</u>	Taxi GPS data	

Table 2 Specific sub\_models and parameters used inof the three-step process.

## 3 <u>A wW</u>orking example

300 The method outlined above was tested on an-the intersection of located in-Xinzhou Rroad and Hongli Rroad in Shenzhen, which is recognized as a waterlogging-flood-prone point by the Water Authority of Shenzhen Municipality. Recall that the parameters to be calibrated are the curve number *CN*, catchment area, *A*, and time of concentration, *t<sub>c</sub>*. The parameter spaces for *CN*, *A*, and *t<sub>c</sub>*. are determined by DEMs and other prior knowledge, which will be discussed in Section 4. The range of parameters should be wide enough to encompass most possible values. After several rounds of testing, the maximum value for *A* is set as 0.5 km<sup>2</sup>, and the maximum value for *t<sub>c</sub>* is 5 h. Optimal parameter sets for most roads would fall into the region enclosed by the maximum parameter sets. Table 3 shows presents the details information of the parameter sets to be calibrated, which totally-form 3,0008 × 20 × 30 = 4,800 possible combinations. For ease of exposition, we assume that all parameters are uniformly distributed.

Donomotor	Annotation Minimum		Marimum	Incrementel	Number of possible	
r ai ametei	Annotation	Winninum	waxiniuni	merementai	parameter values	
Curve number	<u>CN</u>	<u>40</u>	<u>75</u>	<u>5</u>	<u>8</u>	
Catchment area	Α	$0.1 \text{ km}^2 0.01 \text{ km}^2$	<u>0.29 km<sup>2</sup>0.5 km<sup>2</sup></u>	0.01 km <sup>2</sup>	<u>20</u> 50	
Time of concentration	$t_{C}$	<u>0.75 h</u> 1/12 h	<u>3.2 h</u> 5-h	1/12 h	<u>30</u> 60	

Table 3 Detailed information of on parameter sets to be calibrated.

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The t<u>T</u>axi GPS data collected during two storm events <u>that</u> occurr<u>eding</u> on <u>9</u>-May <u>9</u>, 2015 and <u>23</u>-May <u>23</u>, 2015 were are used to calibrate the parameter sets <u>of for</u> the <u>target</u> intersection. Rainfall time series <u>data</u> and taxi observations <del>under</del> during two-these two storms are shown-presented in Fig. 5. Each taxi observation contains two time series: <u>One is</u> the time

series of <u>5 min</u> taxi volumes at <u>5 min intervals</u>, and the other is the <u>5 min time series of road statuses at 5 min intervals</u>, which is <u>These were</u> derived from the taxi volumes, with the <u>a</u> value <u>of one to be 1</u> if the taxi volume <u>wasis higher greater</u> than <u>0-zero</u> and <u>a value of zero0</u> if the taxi volume <u>is-was 0 zero</u>.



**Figure 5** Rainfall and taxi observations used to calibrate the hydrological parameters: (a)  $\frac{5 \min rR}{a}$  ainfall time series  $\frac{in 5}{2}$ . (b)  $\frac{5 \min rr}{a}$  ainfall time series  $\frac{in 5}{2}$ . (c)  $\frac{7}{2}$ . May  $\frac{23}{2}$ ,  $2015_{a^{-}}$  (c)  $\frac{7}{2}$ . Ttaxi observations on  $\frac{9}{2}$ . May  $\frac{9}{2}$ ,  $2015_{a^{-}}$  (c)  $\frac{1}{2}$ .

\_\_Given the rainfall on 9-May 9, 2015, we should-must calculate the time series of no-taxi-passing probabilities possibility for each parameter combination. Because there are 4,800 parameter sets, we can derive 4,800 possible time series of no-taxi-passing probabilities. For simplicity, we only present the 3,120th parameter set (i.e., *CN* = 65, *A* = 0.2 km<sup>2</sup>, and *t<sub>c</sub>* = 2.75 h) as an example to demonstrate the working of the proposed method. According to the three-step process, the first step is to convert the original rainfall into the rainfall excess using the curve number method given *CN* = 65 (Fig. 6(a)). Then, for each combination of *A* and *t<sub>c</sub>*, we construct a SUH. As there are 3,000 parameter sets, we can construct 3,000 different SUHs. For simplicity, we only chose the 1,170th parameter set, i.e. *A*=0.2 km<sup>2</sup> and *t<sub>c</sub>*=2.75 h, as examples to show the calibration works. Using Eq.(9) and Eq.(10), the we calculated the peak discharge *q<sub>p</sub>* and peak time *t<sub>p</sub>* using Eqs. (9) and (10) ean be calculated as:

$$t_p = 0.6 \times 2.75 + \frac{1}{2 \times 12} \approx 1.69 \ h$$
$$q_p = 2.08 \times \frac{0.2}{1.69} \approx 0.24 \ \text{m}^3 \text{s}^{-1}$$

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The SUH <u>iwas</u> derived <u>by-through multiplication</u><u>multiplied</u>\_by  $t_p$  on the *x*-\_axis and <u>by-</u> $q_p$  on the *y*-\_axis of the standard SCS unit hydrograph (Fig.\_6(b)). Next, the rainfall excess <u>shown-presented</u> in Fig.\_6(a) <u>iwas</u> combined with the derived SUH to <u>yield-obtain the-a</u> hydrograph through <u>the-convolution</u> (Fig.\_6(c)).

In the second step, the runoff <u>iwas</u> transformed <u>into the a</u> time series of road disruption probabilit<u>ies</u> based on the 335 runoff-disruption function (Fig.\_6(d)).<u>Note that t</u> he runoff-disruption function takes the product<del>ion</del> of water depth and velocity (in <u>the units</u> of m<sup>2</sup> s<sup>--1</sup>) as input<u>s</u>. Therefore, the original runoff (in <u>the units</u> of m<sup>3</sup> s<sup>--1</sup>) <u>produced byderived in</u> the first step should be divided by the road width before inputting<u>it in</u> to the runoff-disruption function.

In the third step, the time series of road disruption probabilit<u>iesy</u> (Fig.6(e)) <u>iwas</u> converted to that of no-taxi-passing probabilit<u>iesy</u> using Eq. (16) (Fig. 6(f)). According to the historical taxi GPS data, the average number of taxis arriving at the road, λ, is 10.0 taxi per 5 min. The average number of taxis during the flooding period is presented in Fig. (6f), and tThe

derived time series of no-taxi-passing probabilitiespossibility\_is shown-presented in Fig. 6(g).





Figure 6 An eExample transformation of how the a rainfall time series is transformed into the no-taxi-passing probabilitiesy using the three-step procedure for the 1170th 3,120th parameter set: (a) Ttime series of rainfall and rainfall excess. (b) SUH constructed using the 3,120th1170th parameter set. (c) Dderived runoff. (d) Eempirical runoff-disruption function. (e) Dderived time series of disruption probabilitiesy. (f) Disruption no taxi passing probability function average taxi volume in 5 min intervals. and (g) Dderived no-taxi-passing probabilitiesy.

After the time series of no-taxi-passing probabilities for every parameter set were set were

degree of belief that a given parameter set is optimal <u>was calculated</u> by integrating it with the taxi observations on <u>9-May 9</u>,
 2015. According to Eq. (5), the posterior probability of the <u>1,1703,120th</u> parameter set is calculated as:

$$P(\theta^{(1170)}|X) \propto P(\Omega^{(1170)}) \mathscr{L}(X \mid \Omega^{(1170)}) = \frac{1}{3000} \prod_{t=1}^{T} (1 - \omega_t^{(1170)})^{X_t} (\omega_t^{(1170)})^{1 - X_t}$$
$$P(\theta^{(3120)}|X) \propto L(X|\theta^{(3120)}) P(\theta^{(3120)}) = 1/4800 \times \prod_{t=1}^{T} (1 - \omega_t^{(3120)})^{X_t} (\omega_t^{(3120)})^{1 - X_t}$$

where  $\mathscr{L}(\theta^{(1170)}|X) \mathcal{L}(\theta^{(3120)}|X)$  is the likelihood-posterior distribution of probabilities that the 1,1703,120th parameter set is optimal conditioning conditional on X, which represents is the taxi observations on 9-May 9, 2015 shown-presented in Fig. 5(c). The  $\mathcal{P}(\Omega^{(1170)}) \mathcal{P}(\theta^{(3120)})$  is the prior probability of the 1,1703,120th parameter set to being optimal, and its values is 1/30001 / 4,800 because there are 3,0004,800 possible combinations.

<u>By</u> <u>Ff</u>ollowing the <u>abovethis</u> process, we can calculate the posterior probabilit<u>iesy</u> for every parameter set. <u>FurthermoreAdditionally</u>, the posterior probability distribution of <u>a</u> parameter set <u>could\_can</u> be updated using the taxi 360 observations and rainfall data on <u>23</u>-May <u>23</u>, 2015 <u>as</u>:

$$P(\boldsymbol{\theta}^{(i)}|\boldsymbol{X}_2) \propto \mathcal{L}(\boldsymbol{X}_2|\boldsymbol{\theta}^{(i)}) P(\boldsymbol{\theta}^{(i)}|\boldsymbol{X}_1)$$

where  $P(\theta^{(i)}|X_1)$  is the original posterior probability distribution <u>calculated calibrated</u> based on the storm on 9-May 9, 2015, and  $P(\theta^{(i)}|X_2)$  is the updated posterior distribution after the data of storm <u>from 23</u> May 23, 2015 are added. Fig. 7 illustrates<u>ed</u> the evolution of the probability distribution with the availability of <u>more additional</u> taxi data. <u>The first row in Fig.</u>

365 <u>7 represents the prior joint distribution of hydrological parameter sets, and the second and third rows represent the posterior distribution after each round of updating.</u> The posterior distribution dominates the uniform prior distribution after the first updateing, and the distribution is refined slightly a little bit after the second updateing.



370 Figure 7 Evolution of the posterior probability distribution of hydrological parameter sets: (a) Prior distribution before updating. (b) Posterior distribution after the first updating. (c) Posterior distribution after the second updating.

## 4 Method Vyalidation and result

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## 4.1 Method validation Data description

The proposed method <u>iwas</u> validated <u>upon on</u> flood-prone roads located in Shenzhen, China, which is a coastal city frequently hit by extreme storms <u>in during</u> summer. Another reason that Shenzhen is chosen is that only Shenzhen, as far as we known, To the best of our knowledge, Shenzhen is the only city that has shared the runoff-related data to with the public in China. Three data sources, which namely are taxi GPS data, rainfall data, and authoritative water level data, were

are used to validate the our parameter calibration method. Hydrological parameters are-were calibrated using the first two data sources; and the water level data acteds as the ground truth to validate the proposed method. Taxi GPS data wereare anonymized and aggregated to in the road every 5 min intervals. The rRainfall data, which are were also collected every in 5 min intervals, are were measured at 115 gauging stations citywide; and are mapped to the road network throughout Shenzhen using the Oordinary Kriging spatial interpolation algorithm. The water level data are were only measured at some certain waterlogging flood-prone points; with a dynamic sampling interval ranging from 5 min when the weather was rainy to 1 h when the weather was clear rainless. The proposed calibration method was validated by checking analyzing the hydrographs derived from the calibrated hydrological models and against the authoritative water levels for 10 selected roads. Detailed information of on the three data sources is provided are listed in Table 4.

Table 4 Detailed information	n <del>of <u>on the</u> three data sources.</del>
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Item	Taxi GPS data <sup>1</sup>	Rainfall data <sup>1</sup>	Water level data <sup>2</sup>	
	Transport Commission of	Mateorological Bureau of	Shenzhen Municipal	
Source	Shenzhen Municipality	Shonzhon Municipality	Government Data Open	
		Shenzhen Municipanty	Platform <sup>1</sup>	
Record	Taxi volume of each road	5 min accumulative rainfall	Water level	
Data collection period	May 2015	2015 and 2019	2019	
Data collection interval	5 min	5 min	5 min -1 h	
Location	Citywide	115 rainfall gaging stations	171 flooding gaging sites	

<sup>1</sup> The complete taxi GPS data and rainfall data are not openly accessible <u>due-owing</u> to the requirements of data policy. To validate <u>the-our</u> research findings, we have uploaded the necessary data <u>in-to</u> Zenodo (Kong, 2022).

<sup>2</sup> Openly available at the site: https://opendata.sz.gov.cn/data/dataSet/toDataDetails/29200\_01403147.

<u>The Tt</u>wo storm events, occurred on 9-May 9, 2015 and 23-May 23, 2015 are-were treated as calibration events and the a storm occurred on 11-June 11, 2019, wasis retained for testing. ObviouslyClearly, there is a four4 year span-gap between the calibration data and validation data due tobased on the data availability. The hydrological environments of flood-prone roads may have changed during these years, which could render the parameters calibrated based on data from 2015 inaccurate for

395 <u>analysis in 2019.</u> To reduce the validation error caused by <u>thisthe</u> time <u>differencegap</u>, <u>the</u> roads to be validated should <u>have</u> be<u>en</u> vulnerable to flooding <u>on-in</u> both 2015 and 2019 so that <u>the</u> hydrological parameters of these roads <u>would</u> have <u>a</u> higher chance <u>to-of</u> remaining unchanged. Therefore, <u>in-a</u> total of 10 flood-prone roads, <u>which-that</u> were labelled as <u>flood prone</u> roadssuch in <u>on</u> both the List of 2015 Flood-prone Roads in Shenzhen (Water Authority of Shenzhen Municipality, 2015) and the List of 2019 Flood-prone Roads in Shenzhen (Water Authority of Shenzhen Municipality, 2019), were carefully selected (Fig. 8).



Figure 8 Spatial distribution of 10 flood-prone roads in Shenzhen.

Next, the posterior probability of parameter sets after calibration for the 10 roads are illustrated in Fig.9. As shown in Fig.9, the posterior probability distribution of parameter sets for most flood prone roads are clustered around the optimal parameter set after two runs of updating, indicating that the uncertainty of parameter sets is refined to a much smaller area when taxi observations are added. It should also be noted that the posterior probability of parameter set for the Jinlian Road (Road ID=10) is evenly distributed on a triangular region (Fig.9j). By examining the taxi data of the road, we found that the taxi volume was greater than 0 for most 5 min intervals during two storms, indicating that the road was not disrupted during two storms. As hydrological parameters are calibrated by adjusting their values such that the runoff generated by the acceptable parameter sets could yield the disruption period during which no taxi points are observed, the lack of no taxi passing period would provide less information for calibration compared with when no taxi passing period is observed. This explains why the posterior probability is not refined to a small area domain. However, we can still get some valuable information from Fig.9j. First, the catchment area for the Jinlian Road should not be too large to generate the runoff which may cause the road disruption.

Second, the catchment area is highly intercorrelated with time of concentration. As the catchment area gets larger, the time of

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concentration is more likely to increase so that the high runoff volume could not converge in a short time.



**Figure 9** Posterior probability distribution of hydrological parameter sets after the first updating for 10 flood prone roads. Subplots (a) (j) represent the probability distribution for Road 1–10.

# 4.2 Prior distributions of calibrated parameters

- 420 We introduced two types of prior distributions to demonstrate the effects of prior distributions on calibrated parameters. The first prior distribution was determined based on prior knowledge and DEMs from Shenzhen, which were obtained from ASTER GDEM V3, which is a product of NASA and Japan's Ministry of Economy, Trade, and Industry (METI) (ASTER Global Digital Elevation Map, 2023). This global DEM covers the entire land surface of the earth with a 30 m resolution, exhibiting notable improvements in horizontal and vertical accuracy while reducing anomalies compared to previous
- 425 versions. We inputted the DEMs from Shenzhen into the hydrological software PCSWMM to delineate catchments and calculate the catchment area. Subsequently, we computed the time of concentration using the watershed lag method (Natural Resources Conservation Service, 2010b). As suggested by Zhang and Huang (2018), we used the average curve number for Shenzhen in 2015, which was assessed to be 60, as the estimated curve number for each road under validation.
- We then constructed a discretized parameter space for the three parameters for each road as follows. For the curve number, we examined eight possible values centered on 60 with steps of five. For the catchment area, we considered 20 possible values centered on the estimated value with steps of 0.01 km<sup>2</sup>. For the time of concentration, we considered 30 possible values centered on the estimated value with steps of 5 min. After constructing the parameter space for the parameters, we assigned a triangular prior distribution to each, which assumed the maximum probability at the estimated value and linearly decreased to zero at the parameter space boundaries, as depicted in Fig. 9.



Curve number	<u>8</u>	<u>5</u>	Maximum probability at 60 and linearly reduces to zero at the parameter space boundaries	<u>8</u>	<u>5</u>	<u>1/8 for each</u> possible value
Catchment	20	0.01	Maximum probability at the	20	0.01	<u>1/20 for each</u>
area	<u>20</u>	0.01	estimated value and linearly reduces	20	0.01	possible value
Time of	20	1/12	to zero at the parameter space	<u>30</u>	1/12	<u>1/30 for each</u>
concentration	<u>.50</u>	<u>1/12</u>	boundaries		1/12	possible value

### **4.3 Posterior distributions after calibration**

- We first calibrated the parameters based on the prior distributions calculated according to the DEMs and other prior
   knowledge. The resulting posterior distributions are presented in Fig. 10. Each row in Fig. 10 represents a different road, and each column represents a curve number. Each subplot presents the joint probability distribution of the catchment area and time of concentration for a given curve number. The color intensity in Fig. 10 represents the magnitude of the probabilities. Following two iterations of updating, the posterior probability distributions for both the catchment area and time of concentration converge around the optimal parameter sets for most flood-prone roads. This demonstrates that incorporating
- 450 taxi observations significantly reduces the uncertainty associated with catchment area and time of concentration. The probability typically achieves its maximum value when the curve number is either 55 or 60. Furthermore, each subplot contains a salient cluster with higher probability than other regions, suggesting that there may be multiple acceptable parameter sets.

Furthermore, the optimal catchment area under a given curve number decreases as the curve number increases, whereas the optimal time of concentration under a given curve number increases with the curve number. This is logical, because a higher curve number corresponds to increased rainfall excess under identical rainfall conditions, requiring a reduction in catchment area to maintain the runoff that best aligns with the taxi observations. Similarly, an increase in the time of concentration diminishes the peak runoff produced by the additional runoff generated by a higher curve number, thereby preserving the optimal runoff status.





Figure 11 Marginal prior and posterior probability distributions of the curve number for 10 flood-prone roads. For example, for road ID = 6, the optimal parameter set consists of a catchment area of 0.19 km<sup>2</sup>, time of concentration of 0.9 h, and curve number of 55. To investigate the effects of these parameters on the hydrograph and time series of no-taxipassing probabilities, we held two parameters constant at their optimal values and observed the impact of changing the third parameter. Our findings are presented in Fig. 12. One can see that when the catchment area varies from 0.04 to 0.23 km<sup>2</sup>, the maximum no-taxi-passing probability increases from 20% to 100% and the duration for which the no-taxi-passing probability varies from 0.5 to 1.8 h. In contrast, when the curve number varies from 40
480 to 75, the maximum no-taxi-passing probability is fixed at 100%, the duration for which the no-taxi-passing probability exceeds 0.5 extends from 1.1 to 1.3 h, and the peak time of the no-taxi-passing probability remains fixed at the 1.1 h. These

small fluctuations in the time series of no-taxi-passing probabilities are representative of why the distribution of curve numbers remains relatively stable after calibration compared to the catchment area and time of concentration.



- 485 **Figure 12** Impacts of three parameters on the variation of the time series of runoff and no-taxi-passing probabilities: (a) catchment area conditional on runoff, (b) catchment area conditional on the no-taxi-passing probability, (c) time of concentration conditional on runoff, (d) time of concentration conditional on the no-taxi-passing probability, (e) curve number conditional on runoff, and (f) curve number conditional on the no-taxi-passing probability.
- The posterior distributions calibrated based on the uniform prior distribution are presented in Fig. 13. When comparing two posterior distributions derived from two prior distributions, it is clear that the posterior distributions of the catchment area and time of concentration are very similar, indicating that the impact of prior distributions on these parameters rapidly diminishes after taxi-related knowledge is added. As stated by Beven and Binley (1992 pp: 286), "as soon as information is added in terms of comparisons between observed and predicted responses then, if this information has value, the distribution of calculated likelihood values should dominate the uniform prior distribution when uncertainty estimates are recalculated."



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#### **4.4 Validation results**

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After the parameter sets were calibrated, they were combined with the an\_SCS unit hydrograph to construct the an\_SUH, which were was further combined with the rainfall data occurring from on 11-June 11, 2019 to produce the predicted hydrograph. As Because the posterior probability associated with each parameter set can be regarded as a fuzzy measure reflecting the degree of belief that the parameter set is true, the weighted runoffs values for each parameter set were summed to produce calculate the final predicted runoff:

$$Q = \sum_{i=1}^{N} P(\theta^{(i)} | X) Q^{(i)}$$
(17)

505 where <u>Here</u>, Q is the final predicted runoff,  $Q^{(i)}$  is the simulated runoff derived from the *i*th parameter set, and  $P(\theta^{(i)}|X)$  is the posterior probability of the *i*th parameter set, <u>which acting acts</u> as the <u>a</u> weight.

The output of the calibrated hydrological model is runoff (with the-units of m<sup>3</sup> s<sup>--1</sup>), whereas the validation data <u>areis</u> water level <u>data</u> (with the-units of *m*). <u>BecauseAs</u> the calibration data and validation data <u>arise-came</u> from multiple sources and have different units, conventional error-based statistics such as the mean absolute error (<u>MAE</u>)wereare not suitable <u>forin</u> this study. <u>Most often, tT</u>he discharge of <u>a</u> stream is rarely measured directly. Instead, streamflow is typically determined by converting measured water depth (i.e., <u>water</u> stage) into discharge through a rating curve, which provides a functional relationship between the water stage and discharge at a specified point (Le Coz et al., 2014). Inspired by the application of the rating curve, we validated the <u>our</u> method by <u>developing the rating curve for every road, and then estimate estimating</u> the goodness\_-of\_-fit between the water level which was measured in the field and the corresponding runoff which was predicted

515 <u>based on the proposed calibration method.of those rating curves.</u> A higher goodness of fit indicates synchronous trends between the runoff and water level, which indirectly demonstrates the feasibility of the proposed method.

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Because the posterior distributions derived from the two types of prior distributions were very similar, we only considered the posterior distribution calibrated based on prior distributions derived from DEMs and other prior knowledge for validation. Comparisons between the observed water depth and the simulated runoff for 10 selected roads are shown presented in Fig. 140, and rating-corresponding curvesscatter plots constructed by fitting the runoff-stage scatter plot are shown-presented in Fig. 154. We use the Pearson correlation coefficient, which measures the linear correlation between two variables, as the <u>a</u> goodness\_-of\_-fit indicator. The result shows<u>One can see</u> that 8 of 10 roads have rating curves with <u>are</u> characterized by significant positive Pearson coefficients, indicating that the runoff and water have similar and consistent variation processtrends.





**Figure**  $\frac{10}{14}$  Comparisons between the observed water depth and the simulated runoff for <u>rRoads</u> 1 to-10. The maximum value is 30 m<sup>3</sup> s<sup>-1</sup> of on the left *y*-axis (i.e., runoff) and 0.6 m of on the right *y*-axis (i.e., stage) for every each subplot.





It is worth-not<u>eworthying</u> that goodness\_-of\_-fit <u>solely-simply</u> describes the degree of correlation between the observed and simulated data, and may contain validation bias. As suggested by Legates and McCabe (1999), correlation-based statistics <u>areis</u> insensitive to additive and proportional differences between the simulations and observations. Therefore, the fitting of a

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5 rating curve <u>is only a partial reveals part of the validation truth</u>, and the usefulness of the <u>proposed</u> calibration method <del>needs</del> <u>requires</u> further <u>inspectionanalysis</u>.

# 4.2 Application of the method to plot flooding maps in Shenzhen

from 1.67 h to 3.15 h as the return period increases from 2 year to 50 year.

Based on the proposed calibration method, we simulated how the road network experiences flooding for different rainfall return periods. Three storm events of different return periods (*T* = 2, 10, and 50 years) were designed according to the Rainfall
Intensity Formula of Shenzhen (Meteorological Bureau of Shenzhen Municipality, 2015). Each storm lasts 3 hours, with an accumulative rainfall amount of 159 mm, 230 mm, and 283 mm for the 2, 10, and 50 year return period. Hydrological parameters of high level flood prone roads, including expressway, main road, and secondary road, are calibrated using the taxi data on 9 May 2015 and 23 May 2015. The flood prone roads are identified based on the algorithm proposed in our previous studies (Kong et al., 2022). The road discharges under different rainfall return periods are simulated by inputting the designed of parts of the road network, which locate in Baoan District, Shenzhen, for different return periods. With the return periods rising from 2 year to 50 year, the average peak discharge for flood prone roads increases by 80.6%, with the value from 13.9 m<sup>3</sup> s<sup>4</sup> to 25.1 m<sup>3</sup> s<sup>4</sup>. Inputting the simulated runoff to the empirical runoff-disruption function, expressed in Eq.(12), the time series of disruption probability for every road could be derived<sup>3</sup>. To facilitate discussion, we temporarily define the disruption

550 period as the time when disruption probability is higher than 0.5. The average disruption period for flood prone roads increases

<sup>&</sup>lt;sup>3</sup>-Original runoff should be divided by the road width before inputting to the empirical runoff-disruption function.



Figure 12 Spatio temporal evolution of simulated runoff for different return periods in Baoan District, Shenzhen.

#### **5** Discussion

- 555 Three main points are worth discussing about the proposed calibration method are worthy of further discussion. The first is that; although the presented validation results support the use of taxi GPS data to calibrate hydrological parameters for poorly gauged road networks, the proposed method is more applicable to the roads which that are is frequently visited by taxis. Uncertainty increases as the taxi volume of on a road decreases. A road is considered to be passable when at least one taxi GPS points is are observed during the a time interval, while but we cannot assert that the a road is disrupted when the 560 taxi volume is zero. When a road with frequent taxis traffic frequently passing by is observed with no taxi GPS points during the a storm, it is highly probable that the road is disrupted by the flooding, which provides relatively reliable information for parameter calibration. Conversely, when a road with few little taxis visiting traffic has no taxi points during the a storm, there is a great relatively chance high likelihood that the road remains passable and is simply exhibiting its typical trend ofjust has no taxis as usual. Therefore, Thethe proposed calibration method thus becomes relatively unreliable considering that when the a "no-taxi-passing period" is no longer a good proxy fore the "disruption period" for the taxi data sparse road
  - on a road with sparse taxi data. To compensate for the <u>a</u> shortage of the taxi GPS data, extra <u>additional</u> data sources, such as ride-hailing data and bus data, should be incorporated in the future work.

Secondly, the disruption of one road may cause cascading failures, so that where the disruption is may be \_rapidly propagateding from the inundated road to the adjacent non-inundated roads under the constraints of the road connectivity. For

- 570 a road which that is disrupted, but not inundated by the a storm, the implementation of the proposed calibration method may be subject to structural errors. Assume Consider there are two connected roads called, namely Road 1 and Road 2, which that are both disrupted during a storm, and have taxi volumes of two roads are therefore of zero (Fig.13\_16). The difference lies in that In this case, Road 1 is disrupted by the flooding, while whereas Road 2 is only disrupted due to because it is connected ling to the disrupted road, i.e. Road 1. If taxi data areis the only data source used for calibration, then the posterior distributions of
- 575 <u>the hydrological parameters for Road 1 and Road 2 should-will be identical after calibration, because the sequences of taxi volume are identical for both roads. ClearlyHowever, we know that the hydrological parameters for of these two roads could are not be the same, because only one road is flooded otherwise Road 1 and Road 2 should be both flooded. Just like we cannot simply treat the "no-taxi-passing period" as the "disruption period", we cannot confuse the "disruption period" with the "flooded period." In the future work, an algorithm that facilitates enabling to distinguishing the flooding-induced disruption should be developed.</u>



Figure <u>13-16</u> <u>A gG</u>raphical representation to show<u>highlighting</u> the difference between the "disruption period" and <u>flooded</u> <u>period</u> the "no taxi passing period."

Thirdly, the specific-proposed three-step process, which consists of the an SCS unit hydrograph, the empirical runoff-585 disruption function, and the Poisson distribution, isperforms as a realization of the generalized framework shown-presented in Fig. 1. The sSub-models used inof the three-step process are not deterministic, and can be flexibly substituted replaced with by other sub-models according to the needed complexity requirements and data availability. For example, an alternative to the SCS unit hydrograph is the distributed hydrological model. UnlikeCompared with the SCS unit hydrograph, the distributed hydrological model partitions a watershed into physically homogeneous units and captures the complex spatial variation 590 induced by human activity in high resolution, which may be more applicable to the-urbanized environments, such as the-road networks. However, considering that some critical data such as including the road drainage data and land use data are missing, as well as the calibration procedure will become extremely computationally intensive cost associated with the distributed hydrological model, we did not use the distributed hydrological model adopt this model in this study. Another assumption we made in this study is that the number of taxis arriving at a road follows a Poisson distribution. By conducting the Chi-square 595 goodness of fit test, we found that the frequency distribution of taxi volumes adheres to a Poisson distribution for more than 50% of 5 min intervals for 7 of the 10 roads presented in Fig. 8, indicating that the Poisson model appears to be a suitable assumption. However, this hypothesis may not be universally applicable, particularly in different urban contexts, where alternate distributions such as the Weibull distribution may provide a more accurate representation.

#### **6** Conclusion

- An urban flooding model requires various types of data for calibration. In this study, we proposed a Bayesian calibration framework for the hydrological parameters of the <u>a</u> road network based on the taxi GPS data. A three-step procedure, consisting of a rainfall-runoff model, <u>a</u>-runoff-disruption <u>modelfunction</u>, and <u>a disruption</u> no-taxi-passing probability model, <u>enables</u> <u>enabled</u> us to transform the <u>a</u> given rainfall time series <u>into the a</u> time series of no-taxi-passing probabilit<u>ies</u>, which acting-acted set, which is key to <u>the taxi-data-driven</u> model calibration. The calculated no-taxi-passing probabilit<u>ies</u>, <u>which acting-acted</u> as a proxy <u>forof</u> the associated hydrological parameter set<u>s</u>, <u>is further were</u> compared <u>with theto</u> observed taxi data through <u>based on</u> the Bayes equation to assess the posterior probability <u>distributions</u> of the hydrological parameter set<u>s</u>. <u>Three</u> <u>parameters</u>, <u>namely the curve number</u>, catchment area, and time of concentration, were calibrated. The proposed calibration method <u>wasis</u> instantiated by combining <del>some</del>-classical hydrological <u>models</u> and <u>with</u> traffic flow models; and <u>is was</u> validated on 10 flood-prone roads in Shenzhen. The validation results <del>show indicate</del> that the trends of runoff could be correctly predicted for <u>eight</u> roads, which demonstrates the potential of calibrating hydrological parameters based on taxi GPS data-indicating a
  - good performance for hydrological parameter calibration.

This study <u>illustrates highlights</u> the <u>great</u>-potential of integrating transportation-related data with hydrological theory <u>for</u> <u>thein</u> transportation resilience improvement and flood risk management <u>for of the</u>-road networks. We hope that our study <u>can</u> provides a flexible calibration framework for countries <u>which that are short of have little</u> runoff data, but rich <u>of taxi</u> data. We <u>accept acknowledge</u> that the application of the <u>proposed</u> method is currently limited by the heterogeneous spatial distributions of taxis citywide and <u>the</u>-cascading effects of road inundation, but expect this to change with the increasing availability of vehicle data and continuously optimization of modelling <u>approaches</u>.

#### Code and data availability

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The data and code used to validate the <u>proposed</u> method are available <u>onat</u> Zenodo 620 (<u>https://doi.org/10.5281/zenodo.7894921</u><u>https://doi.org/10.5281/zenodo.7294849</u>).

#### Author contributions

JY <u>conceptualised\_conceptualized</u> the article and collected the field data. XK designed the methodology and was responsible for the code compilation. KX plotted the figures and revised the manuscript. BD <u>manged\_managed</u> the

implementation of research activities. SJ discussed the results and contributed to the method validation. XK wrote the final

625 version of the article with contributions from all co-authors.

#### **Competing interests**

The contact corresponding author has declared declares that none of the authors haves any competing interests.

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