



Estimating flood discharge at river bridges using the entropy theory. Insights from Computational Fluid Dynamics flow fields

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Abstract. Estimating the flow velocity and discharge in rivers is of particular interest for monitoring, modelling, and
10 research purposes. Instruments for measuring water level and surface velocity are generally mounted on bridge decks, and
this poses a challenge because the bridge structure (e.g., piers and abutments) can lead to perturbed flow fields. The current
research aims to investigate the applicability of the entropy theory to estimate the velocity distribution and the discharge in
the vicinity of river bridges. To this purpose, a Computational Fluid Dynamics (CFD) model is used to obtain three-
15 dimensional flow fields along a stretch of the Paglia River (central Italy), where a historical multi-arch bridge strongly
affects flood flows. The input data for the entropy model include the cross-sectional bathymetry and the surface velocity
provided by the numerical simulations. Different flow conditions and cross-sections, either upstream and downstream of the
bridge, are considered. It is found that the entropy model can be applied safely upstream of the bridge, also when forced with
a single (i.e., the maximum) value of the surface velocity, with errors on total discharge below 13% in the considered case.
On the contrary, downstream the bridge, the wakes generated by the bridge piers strongly affect the velocity distribution,
20 both in the spanwise and in the vertical directions, and for very long distances. Here, notwithstanding the complex and
multimodal spanwise distribution of flow velocity, the entropy model estimates the discharge with error lower than 8% if
forced with the river-wide distribution of the surface velocity. The present study has important implications for the optimal
positioning of sensors and suggest the potential of using CFD modelling and entropy theory jointly to foster the knowledge
of river systems.

25 **1 Introduction**

Velocity and discharge measurements in rivers are fundamental for monitoring, modelling, and research purposes (Depetris,
2021; Di Baldassarre and Montanari, 2009; Dottori et al., 2013; Gore and Banning, 2017; Herschy, 2009). Unfortunately,
measuring river discharge can be very challenging due to different reasons, for example in the case of intermittent rivers
typical of semi-arid regions, of flash floods in mountain areas, of flood flows involving wide floodplains, of freshwater flows
30 affected by saline tidal intrusions in estuaries, etc. Traditional methods of measuring flow characteristics are generally



expensive, time-consuming, and risky for operators, mainly during severe flow conditions, and they are not feasible in remote and inaccessible locations.

In natural rivers with large cross-sections, the streamwise velocity typically shows a logarithmic vertical distribution due to the bottom roughness. According to field data, the maximum velocity is found just below the free surface and gradually
35 decreases towards the bed (Franca et al., 2008; Guo, 2014). However, plenty of factors contributes in making the velocity distribution irregular. For instance, channel bends and deformed bathymetry produce large-scale secondary currents (Constantinescu et al., 2011; Lazzarin and Viero, 2023; Yang et al., 2012), and the presence of banks and of discontinuities of bed elevation in the spanwise directions can generate secondary currents of the second kind (Nikora and Roy, 2011; Proust and Nikora, 2020), which all increase the three-dimensionality of the flow field and alter the vertical and spanwise
40 distribution of the flow velocity.

The presence of in-stream structures such as bridges, that are characterized by the presence of piers and/or of lateral abutments, can induce significant alterations on the flow field (Laursen, 1963, 1960), producing complex and rapidly varying flow patterns, with the formation of strong three dimensional flow structures (Ataie-Ashtiani and Aslani-Kordkandi, 2012; Chang et al., 2013; Salaheldin et al., 2004). Secondary currents in the cross-section transport low momentum fluid from
45 lateral region to the center of the channel, and high-momentum fluids from the free surface toward the bed (Bonakdari et al., 2008; Nezu and Nakagawa, 1993; Yang et al., 2004). This creates systems of vortices with horizontal (horseshoe vortex) or vertical axes (wake vortex) that in turn modify the velocity distribution (Kirkil and Constantinescu, 2015; Sumer et al., 1997). The wakes generated by in-stream obstacles and contractions can propagate downstream of bridges for quite long distances (Briaud et al., 2009; Yang et al., 2021).

50 The cross-sectional velocity distribution at bridges is of particular interest in hydrology and hydraulics. Indeed, measuring instruments such as hydrometers, as well as radar sensors or cameras for estimating the surface velocity, are generally mounted on bridge decks for convenience reasons. Notwithstanding the recommendation of installing height gage at the upstream side of bridges (Meals and Dressing, 2008), measuring instruments are often located downstream of bridges, which is expected to complicate the discharge estimate (Kästner et al., 2018).

55 Besides the measurement of the flow discharge, the knowledge of flow field nearby bridges has additional practical implications; the flow velocity is the dominant parameter to study the local scour at a bridge pier, which may result in being responsible for the bridge collapse in some extreme conditions (Barbetta et al., 2017; Federico et al., 2003; Lu et al., 2022). The formation of scours at piers and abutments can be attributed to a significant extent to the flow patterns produced at their immediate vicinity, such as the flow contraction and the large-scale turbulent structures (Cheng et al., 2018; Khosronejad et
60 al., 2012).

While monitoring river discharge on the ground has definite advantages (Fekete et al., 2012), the use of traditional methods (e.g., current meters, ADCPs) is not straightforward in case of high-flow conditions. Alternative methods have been proposed to estimate the velocity distributions and the flow discharge through indirect approaches (Bogning et al., 2018; Fekete and Vörösmarty, 2002; Spada et al., 2017; Vandaele et al., 2023; Zhang et al., 2019). These methods typically make



65 use of some field data, whose measurements is generally relatively easy, to reconstruct the entire cross-sectional
distributions. Some of this data include the water level and the velocity at the free surface, which can be collected by
permanent measurement stations mounted on bridge decks and based on a number of diverse techniques (e.g., Eltner et al.,
2020; Herschy, 2009; Jodeau et al., 2008; Schweitzer and Cowen, 2021).

One of the most promising methods to exploit joint measures of water level and surface velocity is based on the entropy
70 concept. Researchers have widely applied this concept to predict velocity distributions as well as other relevant parameters in
open channels (Bonakdari et al., 2015; Chiu, 1989; Chiu and Said, 1995; Chiu et al., 2005; Ebtehaj et al., 2018; Moramarco
and Singh, 2010; Singh et al., 2017; Sterling and Knight, 2002; Termini and Moramarco, 2017). Termini and Moramarco
(2020) identified the cross-sectional velocity by applying the entropy concept and by testing the spatial distribution of the
entropic parameter, M , along with the location of the velocity dip. Using a one-dimensional flood routing model and an
75 entropy velocity profile, Abdolvandi et al. (2021) developed a novel conversion factor estimation procedure. Using two
different assumptions, they estimated discharge at weak gauging sites; constant velocity factor and aspect ratio-related
velocity factor. Based on a picture-based technique, Chahrour et al. (2021) measured the discharge of the Isère River at
Grenoble University Campus using the entropy-based approach. A particle tracking velocimetry (PTV) approach was used to
estimate surface velocities from video images, which were then used to calculate entropy. Using point velocity data collected
80 at $0.6 D$ depth from the surface of the water (being D the total water depth), Vyas et al. (2021) examined the entropy
concept. A correlation was established between the maximum point-velocity based on the sectional mean flow velocity and
the maximum point-velocity estimated at $0.6 D$ depth. Using ADCP surface velocity measurements at different cross-
sections, Bahmanpouri et al. (2022a) applied the entropy approach at the confluence of large rivers (Negro and Solimões
rivers); despite the complex hydrodynamic settings, the entropy method well predicted the velocity field and the flow
85 discharge. Ammari et al. (2022) applied the entropy concept to derive the discharge rate based on the estimate of a single
parameter for the drainage network of a wide region of the central and east part of Algeria. Their outputs suggested a smooth
spatial distribution of the entropy parameter along channels, associated to reliable discharge estimations. Bahmanpouri et al.
(2022b), based on only one surface velocity measurement provided by an unmanned aerial vehicle (UAV), estimated the
cross-sectional velocity and the discharge (with errors lower than 13%) along two European rivers.

90 All these studies demonstrated the accuracy of the entropy method referring to undisturbed flow conditions, or to cases like
confluences or low curvature bends characterized by large-scale three-dimensional effects and secondary currents. However,
in practical applications, water levels and free-surface velocities are often measured by instruments mounted on bridge
decks. In this case, the vicinity of structures such as bridge piers and abutments can produce flow contraction, separation,
and turbulent wakes, generally associated with abrupt variations in the local flow field and deformed cross-sectional
95 distribution of the flow velocity.

Accordingly, the present research is meant to investigate the predictive ability of the entropy theory in estimating the
velocity distribution, and hence the streamflow discharge, in the case of complex flow fields generated by the presence of a
bridge. We choose a reach of the Paglia River, in the central Italy, as a relevant case study; here, a level gauge and a radar for

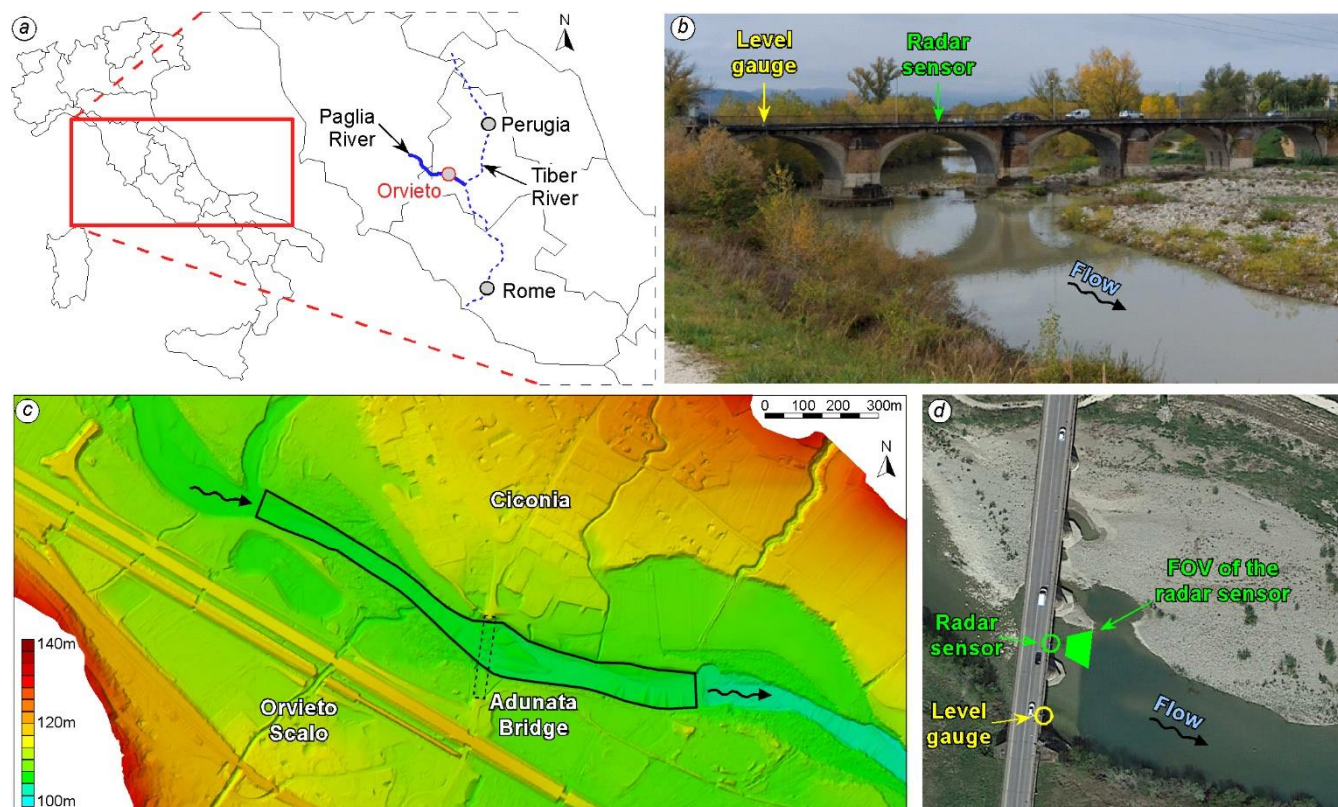


measuring the surface velocity are mounted on a historical multi-arch bridge, which produces strong flow-structure
100 interactions. A physics-based, three-dimensional Computational Fluid Dynamics model is used to provide complete and
accurate velocity distributions, both upstream and downstream of the bridge section, which are compared with those
obtained using the entropy theory forced with different input data. Guidelines for the proper application of the entropy
theory, and the optimal choice and positioning of measuring instruments are finally given.

2 Material and Methods

105 2.1 Field Site

The Paglia river, in the central part of Italy (Figure 1a), is a tributary of Tiber River, subjected to severe flooding and high
sediment transport. The reach of interest is across the Adunata bridge (Figure 1b), near the town of Orvieto, where the Paglia
river subtends a basin area of about 1'200 km². Here, the average discharge is of 10 m³/s; however, flood discharge can reach
estimated values of up to 2'500 m³/s.



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Figure 1. a) Location of the field site; b) downstream view of the Adunata bridge on the Paglia river during normal flow condition (11.11.2021); c) Digital Terrain Model (DTM) nearby the Adunata bridge (dotted line), with the domain of the 3D CFD model (black line); d) location of the level gauge and of the radar sensor with the field of view (FOV) on aerial image (© Google Earth).



The Adunata bridge connects the settlements of Orvieto Scalo and Ciconia, as part of the Italian State Road n.71 (Figure 1c).
115 It is a masonry multi-arch bridge, with 5 arches ending at four piers on the river bed. On the right-hand side, an abutment sustains the bridge and separates it from the floodplain; on the left-hand side, the bridge deck is supported by the main levee. Close to the bottom, the piers have a roughly elliptical shape with the mayor axis, aligned with the flow, 15 m long, and the minor axis, orthogonal to the flow, 5.7 m wide. At the bottom, each pier is sustained by an elliptical plinth whose profile is 2.0 m larger than the pier. The center-distance between the piers is 23.2 m. The piers width increases approaching the deck
120 because of the arches; the deck width is approximatively 10 m.

The main thread of the flow is at the right-hand side of the river, and a large depositional area forms on the left-hand side just downstream of the bridge (Figure 1b). The main channel axis is characterized by a significant curvature, bending to the left at the bridge section (Figure 1c).

2.2 Available Data

125 In the downstream side of the Adunata Bridge, a water level gauge is placed at the center of the first arch, and a radar sensor for measuring the water surface velocity is located at the second arch (Figure 1d). The field of view (FOV) of the radar sensor is also shown in Figure 1d. The time resolution of both the sensors is 10 min.

In addition, a number of flowrate measures and cross-sectional velocity distributions are provided by the Umbria Region Hydrological Service. The flowrate data were collected using a current meter a few tens of meters downstream of the
130 Adunata Bridge, by wading in the period 2009-2011 (flowrate ranging between 3.3 and 14.3 m³/s), and from the bridge in the period 1995-2010 (flowrate ranging between 16.8 and 147 m³/s); additional flowrate data were collected using an Acoustic-Doppler Current Profiler (ADCP) some hundreds of meters upstream of the bridge in the period 2014-2019 (flowrate ranging between 0.37 and 45 m³/s). The official rating curve for the Adunata Bridge, provided by ARPA Lazio, is based on these measures.

135 As detailed in the following sections, the rating curve, the water levels, and the free-surface velocity data collected by the two sensors, are used to validate the numerical models; the cross-sectional velocity distributions measured with the ADCP are used to further assess the entropy-based velocity distributions.

2.3 Numerical Model

The commercial CFD software STAR-CCM+ (Siemens) is used for the numerical simulations. It implements the Finite
140 Volume method to compute the flow field on unstructured, Cartesian computational grids. In the present application, the two-phase Volume of Fluid (VoF) method is used to distinguish water and air in the computational domain (Hirt and Nichols, 1981). This method was shown to well capture the water surface in complex open channel flows (Horna-Munoz and Constantinescu, 2018; Lazzarin et al., 2023b; Li and Zhang, 2022; Luo et al., 2018; Yoshimura and Fujita, 2020).

In the used setup, the model solves the Reynolds-Averaged Navier-Stokes (RANS) equations, in which the stress tensor in
145 the momentum equations is related to the mean flow quantities by adopting the Boussinesq approximation. The eddy



viscosity, μ_T , is determined by solving transport equations for the turbulent kinetic energy, k , and dissipation rate, ε , according to the realizable k - ε turbulence model (Shih et al., 1995), which was shown to provide reliable predictions for large-scale complex flows in natural rivers (e.g., Horna-Munoz and Constantinescu, 2018).

The simulations are advanced in time with an implicit, 1st order discretization, until reaching steady state conditions. The computational domain reproduces a $\sim 1'100$ m long reach of the Paglia River (Figure 1c), centered at the Adunata bridge. The average size of the grid elements is of 1.0 m. Starting 100 m upstream of the bridge and up to 300 m downstream of the bridge, the grid is refined using elements with average length 0.5 m. To well capture the near-wall boundary layer, a prism layer refinement with three layers is used to reduce the wall-normal thickness of the grid cells close to solid boundaries (i.e., the riverbed and bridge structure). The final computational grid is made of ~ 4 million elements.

A rough-wall, no-slip condition is imposed at the solid boundaries by means of a wall function (roughness height of 0.1 m at the bottom, and of 0.01 m at the bridge surfaces). The upper boundary of the computational grid is treated as a symmetry plane (i.e., slip-condition) for the air-flow. The water elevation at the outlet (i.e., downstream section) is kept fixed in time by imposing a suitable hydrostatic-pressure distribution. The value of the downstream level, for each of the simulated scenarios, is derived from an auxiliary two-dimensional (2D), depth-averaged hydrodynamic model calibrated on available data; the 2DEF model has been used to this purpose (see Appendix A for details on the model and its calibration/verification). A constant-in-time, logarithmic velocity distribution is imposed as upstream boundary condition for the water fraction. For the air fraction (upper part of the numerical domain), zero velocity and zero pressure are imposed at the inlet and at the outlet, respectively.

The 3D-CFD model is validated by comparing the surface velocity computed by the model with those measured by the radar sensor located downstream of the bridge (see the yellow bullets in Fig. A2c,d, in the Appendix A).

2.4 Flood events considered in the study

Three different steady flow conditions are simulated with the 3D-CFD model STAR-CCM+, which correspond to the peak flow conditions of flood events occurred in 2012, 2019, and 2022 (Table 1). During the most severe flood of 2012, water flowed also through the floodplains adjacent to the main river channel. The preliminary simulation carried out with the 2DEF depth-averaged model showed that 700 m³/s flowed through the floodplain, overflowing the bridge access roads, and 1800 m³/s flowed within the main channel; this last value is used in the 3D-CFD simulation, which considers only the main channel of the river. The flood events of 2019 and 2022, although being quite ordinary, were the largest floods occurred after the installation of the radar sensor for the surface velocity data (surface velocity data are not available for the 2012 flood).

Event	Discharge [m ³ /s]
2012	1800 (2500)
2019	450
2022	160

Table 1. Simulations performed in the present work. The value in brackets indicate the total discharge considering also the flow over floodplains, not considered in the 3D simulations.



2.5 Entropy theory

Chiu (1989) developed an estimation of cross-sectional velocity distribution, $U(x,y)$, using the entropy probability density function. Later, Moramarco et al. (2004) simplified the estimation. Using this approach, one can determine the entropy-based velocity profile along the verticals as follows:

$$U(x_i, y) = \frac{U_{\max v}(x_i)}{M} \ln \left[1 + (e^M - 1) \frac{y}{D(x_i) - h(x_i)} \exp \left(1 - \frac{y}{D(x_i) - h(x_i)} \right) \right] \quad i = 1 \dots N_v \quad (1)$$

180 where U is the time-averaged velocity, $U_{\max v}(x_i)$ is the maximum value of U along the i^{th} vertical, x_i is the distance of the i^{th} sampled vertical from the left bank, $h(x_i)$ is the dip, i.e., the depth of $U_{\max v}(x_i)$ below the water surface, $D(x_i)$ the flow depth, y is the distance of the velocity point from the bed, and N_v is the number of verticals sampled across the river section. M can be estimated using the linear entropic relation using the mean and the maximum flow velocity, U_m and U_{\max} , measured within the entire cross-section (Chiu, 1989):

$$U_m = \left(\frac{e^M}{e^M - 1} - \frac{1}{M} \right) U_{\max} = \phi(M) U_{\max} \quad (2)$$

185 In general, for a given river site, $\phi(M)$ is assumed to be constant for all flow conditions, while for ungauged sites $\phi(M)$ can be estimated as (Moramarco and Singh, 2010):

$$\phi(M) = \frac{\frac{1}{n} R^{1/6}}{\sqrt{g} \frac{1}{k} \left[\ln \left(\frac{y_{\max}}{y_o} \right) + \frac{h}{y_{\max}} \ln \left(\frac{h}{D} \right) \right]} \quad (3)$$

where y_{\max} is the location of U_{\max} from the bottom and y_o is the datum where the velocity is equal to zero, k is the von Karman constant, R is the hydraulic radius, n is the Manning roughness and D is the maximum flow depth.

Whether at a river site only the surface velocities, $U_{\text{surf}}(x_i, D(x_i))$ are available, then $U_{\max v}(x_i)$ can be estimated as (Fulton and
 190 Ostrowski, 2008):

$$U_{\max v}(x_i) = \frac{U_{\text{surf}}(x_i, D(x_i))}{\frac{1}{M} \ln [1 + (e^M - 1) \delta(x_i) e^{1 - \delta(x_i)}]} \quad (4)$$

where $\delta(x_i) = D(x_i) / [D(x_i) - h(x_i)]$. Specifically, if $h(x_i) = 0$, it follows that $\delta(x_i) = 1$ and, hence, $U_{\max v}(x_i) = U_{\text{surf}}(x_i, D(x_i))$. The magnitude of $\delta(x_i)$ can be obtained based on the iterative procedure proposed by Moramarco et al. (2017). The procedure can be applied for sites with a given $\phi(M)$. The procedure is based on assigning an initial dip, $h(x_{i,p=1})$, where the maximum surface velocity occurs (p is the iteration).

195 For the current research, according to the initial value of dip, a laboratory distribution law of dip suggested by Yang et al. (2004) is implemented, and the $U_{\max v}(x_{i,p=1})$ is assessed by Eq. (4) for all the considered verticals. $U_{\max(p=1)}$ is identified as the maximum of $U_{\max v}(x_{i,p=1})$. Therefore, once $U_{\max v}(x_{i,p=1})$ is replaced in Eq.(1), it enables estimation of the depth-averaged velocities in each cross-section. For the first iteration, the mean flow velocity, $U_{m(p=1)}$, can be estimated using the velocity-area method. As a consequence, $\phi(M_{\text{com},p=1})$ can be computed by Eq.(2), using $U_{m(p=1)}$ and $U_{\max(p=1)}$. The iteration continues



200 until the error of $\phi(M_{com,p}) - \phi(M)$ becomes lower than 0.01. For more details, the reader is referred to Moramarco et al.
(2017).

3 Results and discussions

The comparison between the entropy-based and the CFD-derived velocity distributions is performed considering four cross-
sections (Figure 2), at a distance of 50 m upstream and 50, 100, and 200 m downstream of the bridge, and the three flood
205 events of 2012, 2019, and 2022 (see Table 1).

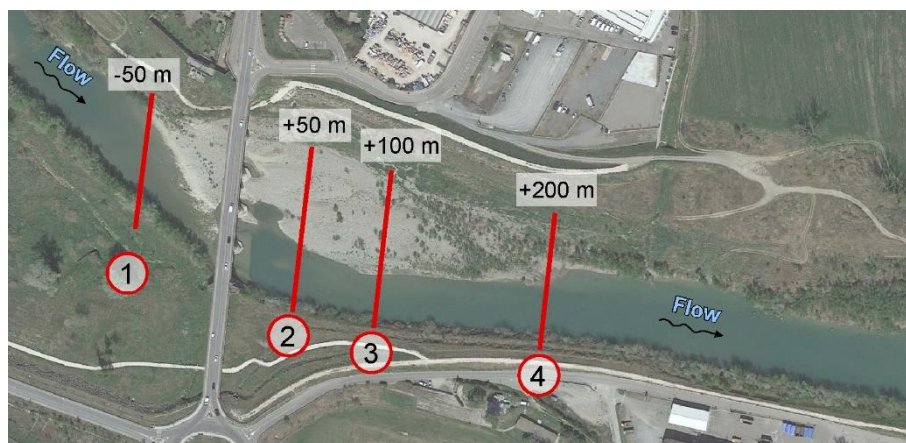


Figure 2. Location of the Adunata Bridge and of the four selected cross-sections (aerial image from © Google Earth).

First, the study analyzes the variability of the entropy parameter, $\phi(M)$, at the four cross-sections, as derived from the cross-
sectional velocity distributions provided by both the 3D-CFD model and the ADCP measures (Sect. 3.1). Then, in applying
210 the entropy model to estimate the cross-sectional velocity, two different procedures are considered. In the first one, the
entropy model is forced with the river-wide distribution of the surface velocities computed by the 3D-CFD model (this is
described in the following Sect. 3.2); in the second one, only the maximum value of the surface velocity computed by the
3D-CFD model is considered as input for the entropy model (Sect. 3.3). The first procedure is applied to all the four cross-
sections, whereas the latter is only applied to cross-sections 1 and 4, i.e., where the effects of the bridge piers is minimal so
215 that the spanwise velocity distribution is unimodal.

3.1 Variability of the entropy parameter

Some relevant parameters that characterize the flow field (e.g., aspect ratio, average and maximum velocity) at the selected
cross-sections are presented in Table 2 for the peak flow condition of the three flood events.



Year	Distance from the bridge (m)	Channel aspect ratio (width/depth)	Average Velocity (m/s)	Maximum Velocity (m/s)	$\phi(M)$	M
2012	-50	9.26	4.43	6.82	0.650	1.91
	+50	13.78	2.91	7.01	0.415	-1.03
	+100	11.05	3.61	6.68	0.541	0.50
	+200	8.5	4.06	5.48	0.740	3.4
2019	-50	16.3	3.0	4.21	0.711	2.87
	+50	18.45	1.93	3.74	0.515	0.18
	+100	14.75	2.08	3.26	0.639	1.75
	+200	12.84	2.40	3.47	0.690	2.51
2022	-50	27.7	2.57	3.40	0.755	3.71
	+50	27.3	1.33	2.58	0.514	0.17
	+100	20.9	1.55	2.19	0.711	2.86
	+200	13.23	1.97	2.56	0.767	3.96

220 **Table 2. Entropy parameter, $\phi(M)$, and other relevant flow data for the cross-section of Figure 2 and the three considered flood events of Table 1.**

Since the entropy parameter is typically assumed to be constant for all flow conditions at a given cross-section, it is of interest to analyze its actual variation by exploiting the flow fields provided by the 3D-CFD model. The values of $\phi(M)$ reported in Table 2 are plotted in Figure 3 as a function of the downstream distance from the bridge. At the first cross-section downstream of the bridge (i.e., cross-section 2), although referring to different flow conditions, the entropy parameters computed with the 3D-CFD and the current meter velocity distributions show the same magnitude, further confirming the reliability of the 3D-CFD model.

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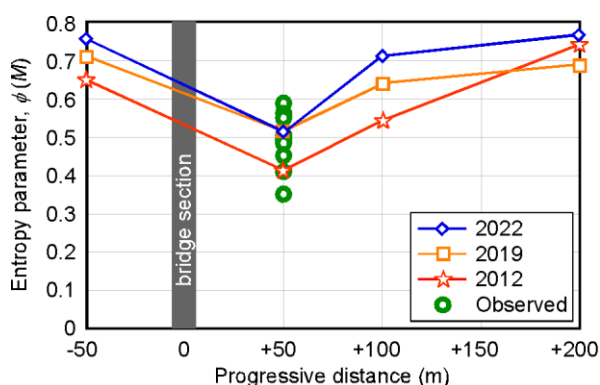


Figure 3. Entropic parameter $\phi(M)$ as a function of the distance from the bridge (positive downstream), computed from the 3D-CFD flow fields of the different simulated scenarios; green dots refers to data derived from ADCP field measures.

230 For each flood event, at cross-sections 1 and 4, i.e., where the flow field is not characterized by the wakes generated by the bridge piers, the entropic parameter assumes similar values, which can be identified as “undisturbed” values. The variability of such undisturbed values of $\phi(M)$ with the flowrate is relatively small, as all the values fall in the range



0.65 < $\phi(M)$ < 0.75, which is in agreement with the range found by Bahmanpouri et al. (2022b) for similar European rivers. On the contrary, at cross-sections 2 and 3, just downstream of the bridge, the values of $\phi(M)$ are consistently reduced due to the effect of the bridge. For the moderate peak flow of the 2022 event, the entropy parameter recovers undisturbed values already at cross-section 3, i.e., 100 m downstream of the bridge. In the largest flood event of 2012, which produced near pressure-flow conditions at the bridge with marked localized increasing of the flow velocity, $\phi(M)$ decreases from 0.64 to 0.42, and a sensible reduction is still present 100 m downstream of the bridge (cross-section 3).

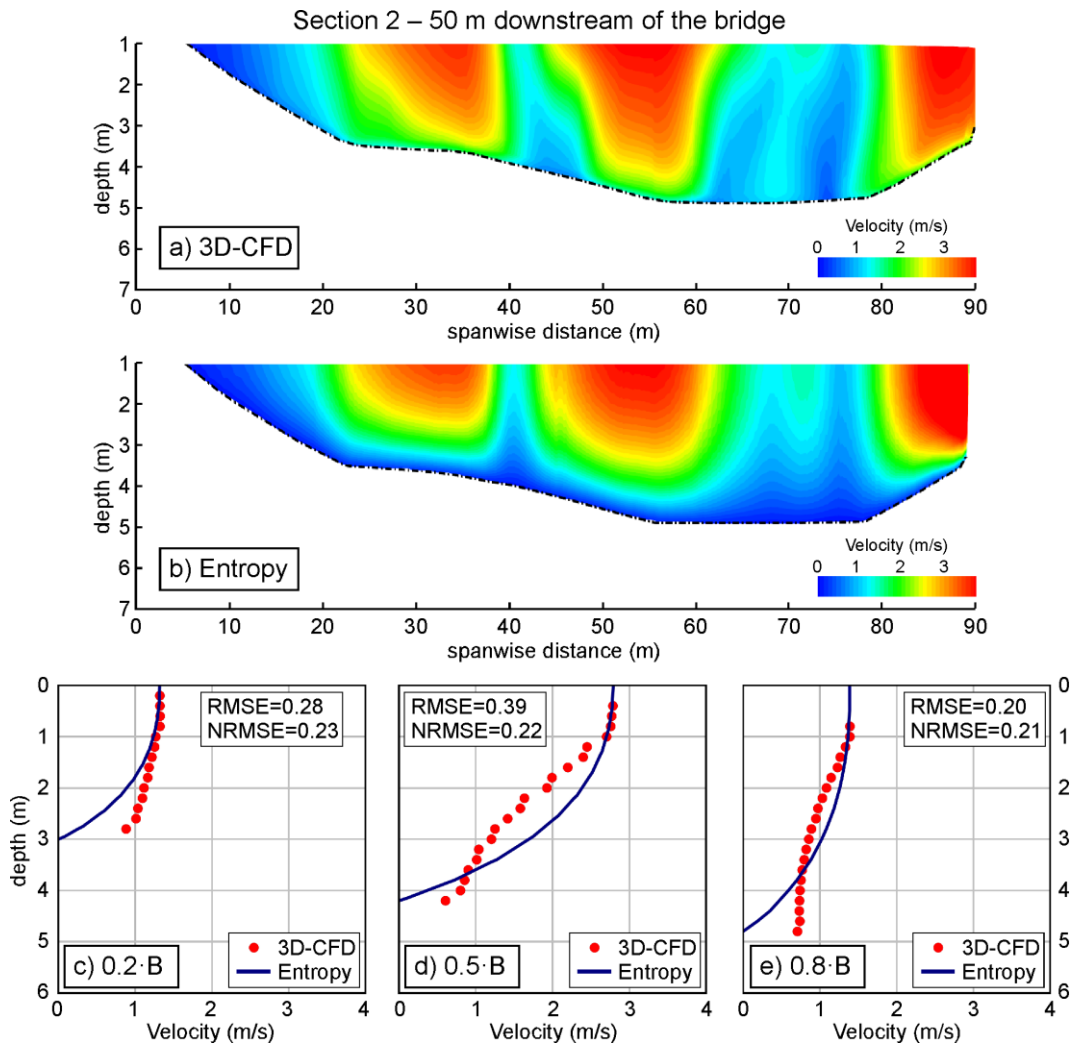
This first analysis suggests that assuming constant values of $\phi(M)$ can be reasonable in undisturbed river reaches; however, in case of irregular flow fields induced by the interactions with in-stream structures, the entropy parameter $\phi(M)$ can vary with respect to undisturbed values and, in addition, it can show substantial variations with the flowrate.

3.2 Entropy model forced with the river-wide profile of free-surface velocity

The efficacy of the entropy model is here tested for the case in which the surface velocity is known for all the width of the cross-section. This could be the case in which the river-wide surface velocity is estimated from imaging techniques (e.g., Eltner et al., 2020; Schweitzer and Cowen, 2021). The results, in terms of cross-sectional velocity distributions, are presented for brevity only for the intermediate peak flow of the 2019 flood event, and for the most challenging cross-sections just downstream of the bridge, where the flow field is disturbed by the pier wakes. The same results, for the peak flows of 2012 and 2022 events, are provided as supplementary material.

Figure 4 presents the cross-sectional velocity distribution 50 m downstream of the bridge (cross-section 2). As shown by the 3D-CFD flow field (Fig. 4a) and reflected in the low value of $\phi(M)$ for this cross-section (Table 2 and Figure 3), the effect of the piers is very strong, such that there is a clearly uneven distribution of the cross-sectional velocity because of the wakes developing downstream the piers. Despite that, using as input the river-wide distribution of the surface velocity provided by the CFD simulation, the entropy model can reliably capture the salient features of the cross-sectional velocity distribution.

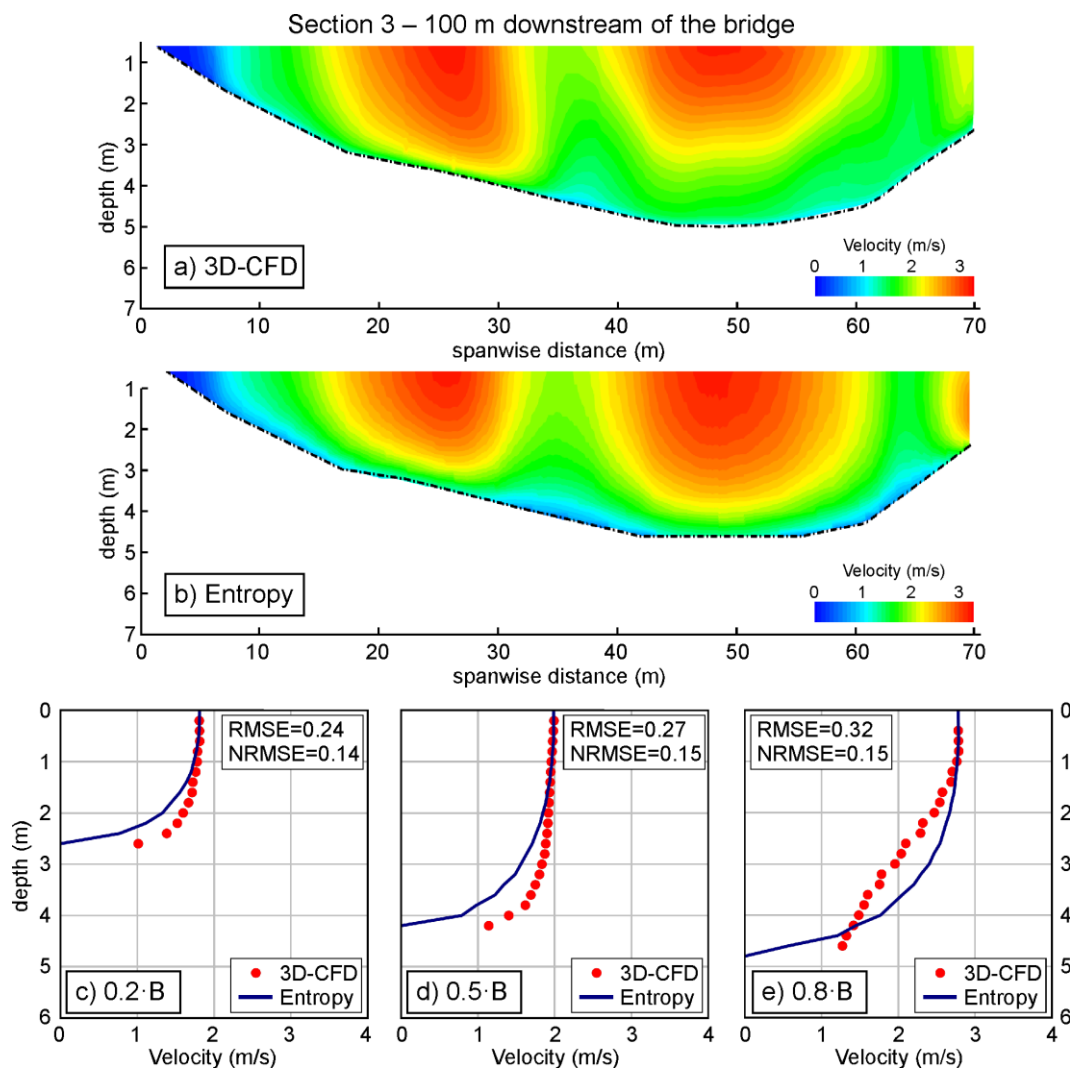
Figure 4(c-e) highlights the comparison of 3D-CFD and entropy flow velocities along three verticals located at 0.2 B , 0.5 B , and 0.8 B (where B is the channel width). Compared to the results of the 3D-CFD model, the entropy approach underestimates the velocity close to the bed. Just downstream of the bridge, due to the presence of the bridge arches, the flow field provided by the 3D-CFD model is configured as a sort of partial orifice flow that increases the vertical uniformity of the velocity distribution compared to a uniform shear flow. Of course, the entropy model cannot capture such a localized flow features, which entails some difference in the patchiness of the physics-based and the entropy velocity distributions (Figure 4a-e). Since the velocities and the volumetric fluxes are still relatively small near the bed, these discrepancies marginally affect the estimation of the section-averaged velocity and, consequently, of the total discharge (Table 3). The percentage error is quite larger (7.6%) for the very high-flow condition of the 2012 event (see Supplementary Materials), due to the accentuation of orifice-flow conditions associated to the higher water levels.



265 **Figure 4.** Flood event of 2019, cross-section 2 (50 m downstream of the bridge). Velocity distributions provided by (a) the 3D-CFD model, and (b) the entropy model forced with the river-wide distribution of the free-surface velocity. Comparison of vertical distributions of velocity at $0.2B$ (c), $0.5B$ (d), and $0.8B$ (e), where B is the width of the cross-section.

Figure 5 depicts the cross-sectional velocity distributions at a larger distance from the bridge, i.e. at cross-section 3, placed 100 m downstream the bridge. The visual comparison with Figure 4 suggests that the effects of the piers on the flow field are reduced because of the increased distance, and the cross-sectional distribution provided by the 3D-CFD model (Figure 5a) appears more regular. The statistical analysis confirms that in this case the entropy model (Figure 5b) is able to simulate the velocity profiles with a higher accuracy.

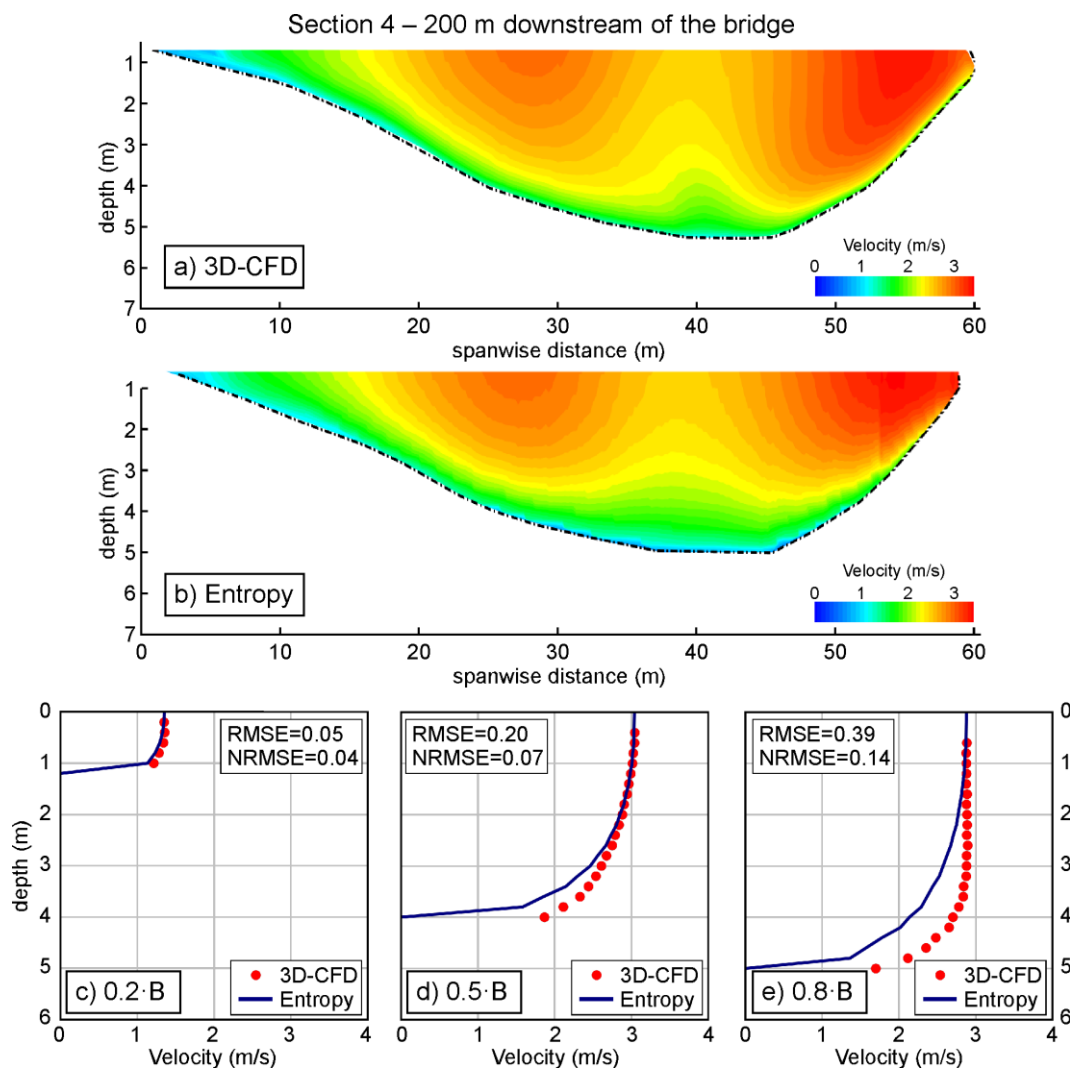
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275 **Figure 5. Flood event of 2019, cross-section 3 (100 m downstream of the bridge). Velocity distributions provided by (a) the 3D-CFD model, and (b) the entropy model forced with the river-wide distribution of the free-surface velocity. Comparison of vertical distributions of velocity at $0.2B$ (c), $0.5B$ (d), and $0.8B$ (e), where B is the width of the cross-section.**

Figure 6 shows the cross-sectional velocity distributions of 3D-CFD and entropy models for the cross-section 4, i.e. 200 m downstream of the bridge. Compared to cross-section 3, the effect of the bridge piers is further reduced, because of both the distance and the more compact shape of the cross-section. Since the effect of the bridge piers is minimum, the statistical analysis shows a better agreement of the entropy model results with the CFD-based data. Though areas with high velocities are still visible in simulations with higher values of the discharge (i.e., events of 2012 and 2019), for the high-flow conditions of 2022, the effect of the bridge pier has completely vanished. Therefore, the lower the flow discharge the lower the distance from the bridge to reach the normal flow condition without the bridge effect.

280



285 **Figure 6. Flood event of 2019, cross-section 4 (200 m downstream of the bridge). Velocity distributions provided by (a) the 3D-CFD model, and (b) the entropy model forced with the river-wide distribution of the free-surface velocity. Comparison of vertical distributions of velocity at $0.2B$ (c), $0.5B$ (d), and $0.8B$ (e), where B is the width of the cross-section.**

The results here presented show that, when the river-wide distribution of the free-surface velocity is provided, the entropy method allows to provide good estimations of the cross-sectional velocity distribution even when the influence of bridge
 290 piers, and thus the unevenness of the flow field, is relevant. The main discrepancies are observed in the regions of flow with low values of velocity, which slightly affect the estimation of the flow discharge. Table 3 lists some statistics and error percentages for the depth-averaged velocity and discharge estimates for all cross-section and the three events considered. The estimation provided by the entropy method are in good agreement with results of CFD model, both upstream and downstream of the Adunata bridge. Though the precision is slightly reduced downstream of the bridge, the results are
 295 reliable also in the vicinity of the structure (i.e., at cross-section 2), suggesting the applicability of the entropy model to



estimate the flow discharges even in case of irregular distributions of the cross-sectional velocity, provided that the river-wide distribution of the surface velocity is used as input data.

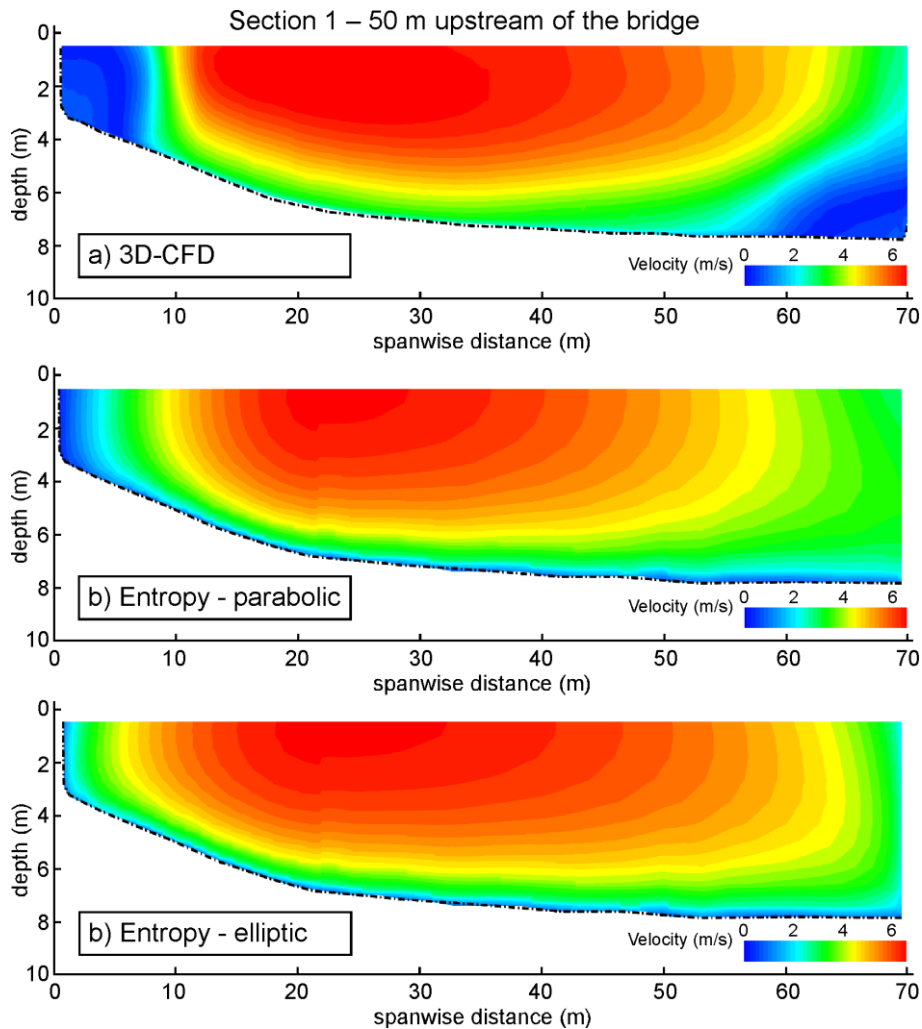
Flood event	Cross-section	Distance from the bridge (m)	Average velocity (m/s)		Discharge (m ³ /s)		Error percentage (%)
			3D-CFD	Entropy	3D-CFD	Entropy	
2012	1	-50	4.43	4.64	1'800	1'885	+4.7
	2	+50	2.91	2.69		1'664	-7.6
	3	+100	3.61	3.54		1'765	-2.0
	4	+200	4.06	4.30		1'906	+5.9
2019	1	-50	3.0	3.0	450	450	+0.1
	2	+50	1.93	1.90		443	-1.5
	3	+100	2.08	2.12		459	+2.0
	4	+200	2.40	2.54		476	+5.8
2022	1	-50	2.55	2.66	160	166	+3.7
	2	+50	1.32	1.24		150	-6.3
	3	+100	1.55	1.51		157	-1.9
	4	+200	1.97	1.98		161	+0.6

Table 3. Flood event of 2019. Comparison between 3D-CFD outputs and entropy-based estimations forced with the river-wide distribution of the free-surface velocity.

3.3 Entropy model forced with a single value of free-surface velocity

In this section, the results are presented considering only a single value of the surface velocity as input for the entropy model, which corresponds to the maximum surface velocity predicted by the 3D-CFD model. Two different spanwise velocity distributions are enforced in the entropic model, namely a parabolic spanwise distribution (PSD) and an elliptic spanwise distribution (ESD). Of course, applying the entropy model using a unique value of the velocity is particularly sensitive of this value and suppose a unimodal velocity distribution in the spanwise direction. For this reason, this kind of approach cannot be used in the cross-sections immediately downstream the bridge, where velocities show large spatial variations (see e.g., Figure 4). Herein, the results are presented for cross-section 1, i.e. 50 m upstream of the bridge for the high-flow condition of the 2012 event, and for cross-section 4, i.e. 200 m downstream of the bridge for the modest peak flow condition of the 2022 event, where the effect of bridge piers on the velocity distribution wears off in a shorter distance.

Figure 7 shows the distribution of the surface velocity based on CFD outputs and both the PSD and ESD entropy models. The agreement of both the PSD and the ESD is generally good in the central and the right parts of the channel, and less good in the left part of the channel. Here, due to the irregular bathymetry (i.e., gravel deposit), the 3D-CFD model predicts localized stagnation zones that cannot be captured by the entropy model based on a single value of the surface velocity. This is confirmed by Figure 8a, which shows the cross-sectional distribution of the depth-average velocity and three vertical profiles. In the perspective of estimating the flow discharge, the lateral discrepancies represent a minor limit, as the central part of the cross-sections conveys the largest part of the total discharge.



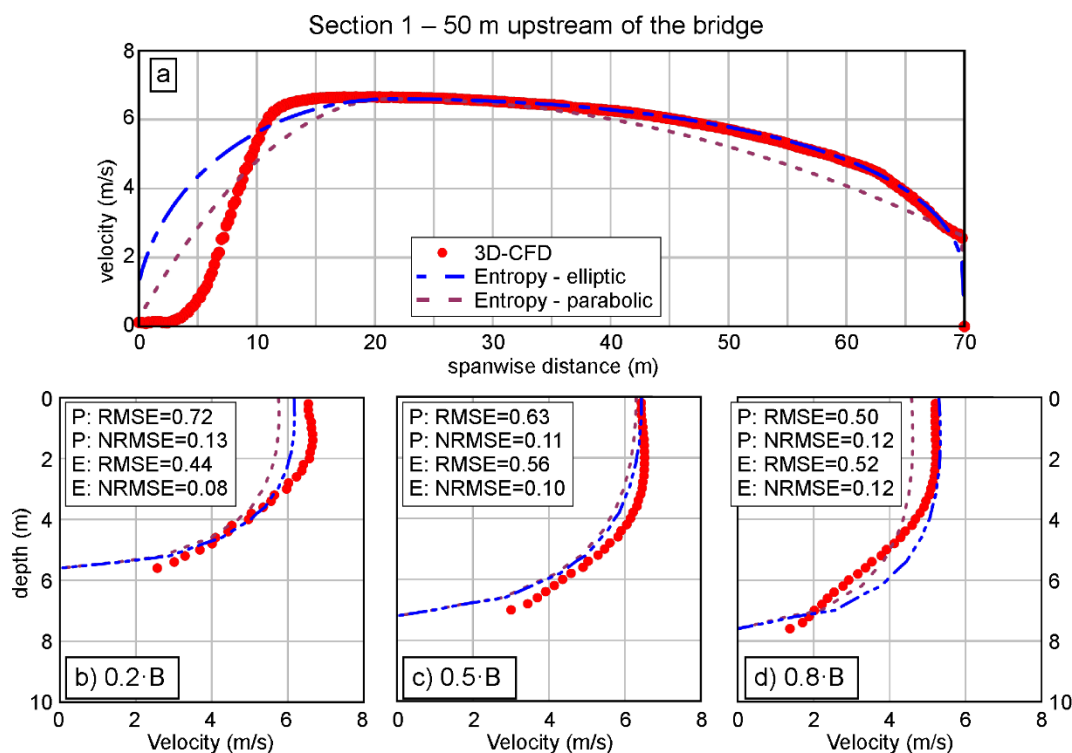
320 **Figure 7. Flood event of 2012, cross-section 1 (50 m upstream of the bridge). Cross-sectional velocity distribution computed with the 3D-CFD model (a), the entropy theory with parabolic (b) and elliptic (c) spanwise velocity distribution.**

Overall, the results based on ESD are more accurate than results based on the PSD: they provide similar results at the center of the channel, but the parabolic distribution generally underestimates the flow velocity close to the banks. Both cross-sectional and vertical distributions of the velocity profiles (Figure 7a and Figure 8c) highlight the existence of a velocity dip, i.e. the maximum velocity is below the water surface, particularly at the center of the channel. This is generally the consequence of secondary currents superposed to the main flow (Termini and Moramarco, 2020). Yang et al. (2004) and Moramarco et al. (2017) reported that for large aspect ratios of channel flow, B/D , the dip phenomenon appears primarily near the sidewall region, whereas for relatively low aspect ratios ($B/D = 9.26$ for cross-section 1) the velocity dip is generally located at the center of the channel (Bahmanpouri et al., 2022b, a; Kundu and Ghoshal, 2018; Moramarco et al., 2017; Termini and Moramarco, 2020). In this case, the 3D flow field from the CFD simulation shows that the dip depends on the

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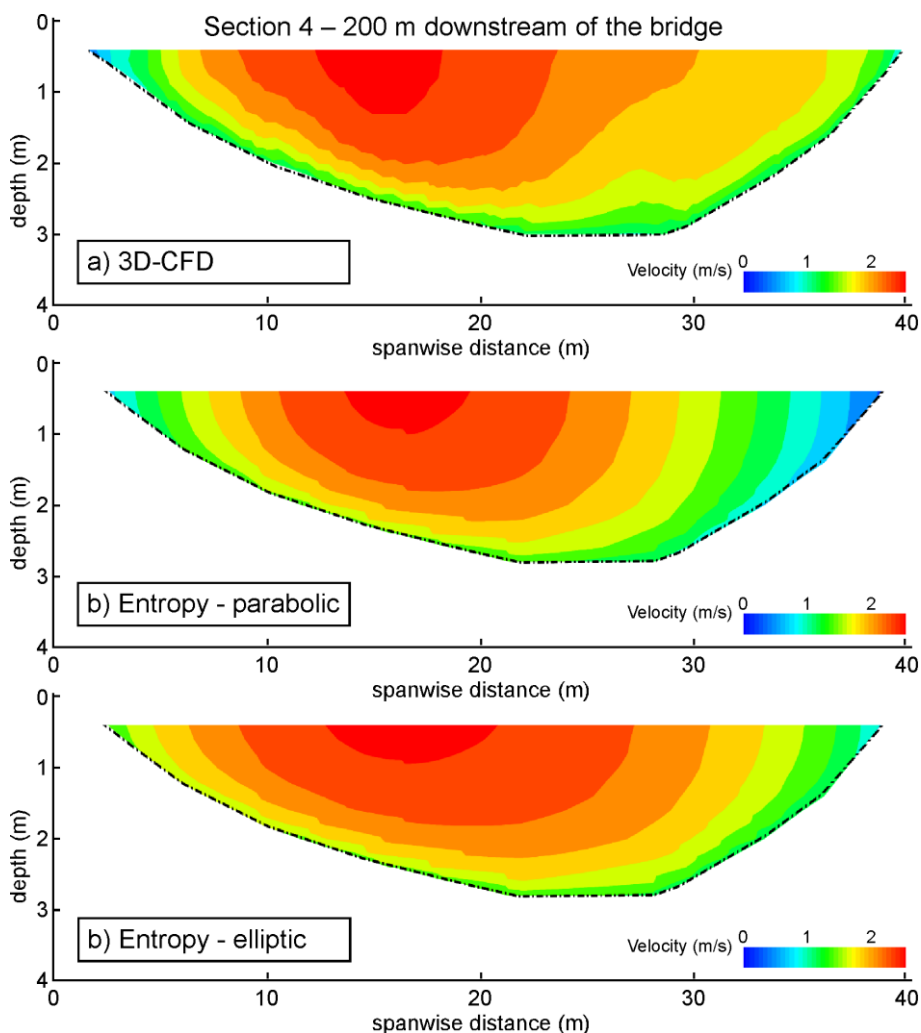
330 counter-clockwise rotating secondary current generated by the upstream right-handed bend. Indeed, rotational inertia makes these curvature-induced helical flow structures to propagate downstream for relatively long distances (Dominguez Ruben et al., 2021; Lazzarin and Viero, 2023; Thorne et al., 1985).



335 **Figure 8.** Flood event of 2012, cross-section 1 (50 m upstream of the bridge). Spanwise distribution of the surface velocity (a); comparison of vertical distributions of velocity at 0.2*B* (b), 0.5*B* (c), and 0.8*B* (d).

The distribution of the velocity at the free surface for the cross-section 4 (200 m downstream of the bridge) is presented in Figure 9 for the moderate peak flow condition of the 2022 event. For this cross-section, in the 3D-CFD results (Figure 9a), the maximum surface velocity is located on the left side of the channel, rather than at its center (this aspect is discussed in the following). Forced with the maximum water surface velocity, the entropy model well reproduces the velocity field in the central part of the riverbed. Larger discrepancies are instead observed in the lateral part of the cross-section, with the elliptic spanwise distribution (ESD) that performs slightly better than the parabolic (PSD), particularly in the right side. Figure 10 shows the cross-sectional distribution of the depth-averaged velocity and the velocity distribution along three verticals. In terms of cross-sectional average velocity and flow discharge, both the PSD and ESD produce error that are lower than 10% (Table 4), then quite larger than those obtained using the river-wide surface velocity as input for the entropy model.

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Figure 9. Flood event of 2022, cross-section 4 (200 m downstream of the bridge). Cross-sectional velocity distribution computed with the 3D-CFD model (a), the entropy theory with parabolic (b) and elliptic (c) spanwise velocity distribution.

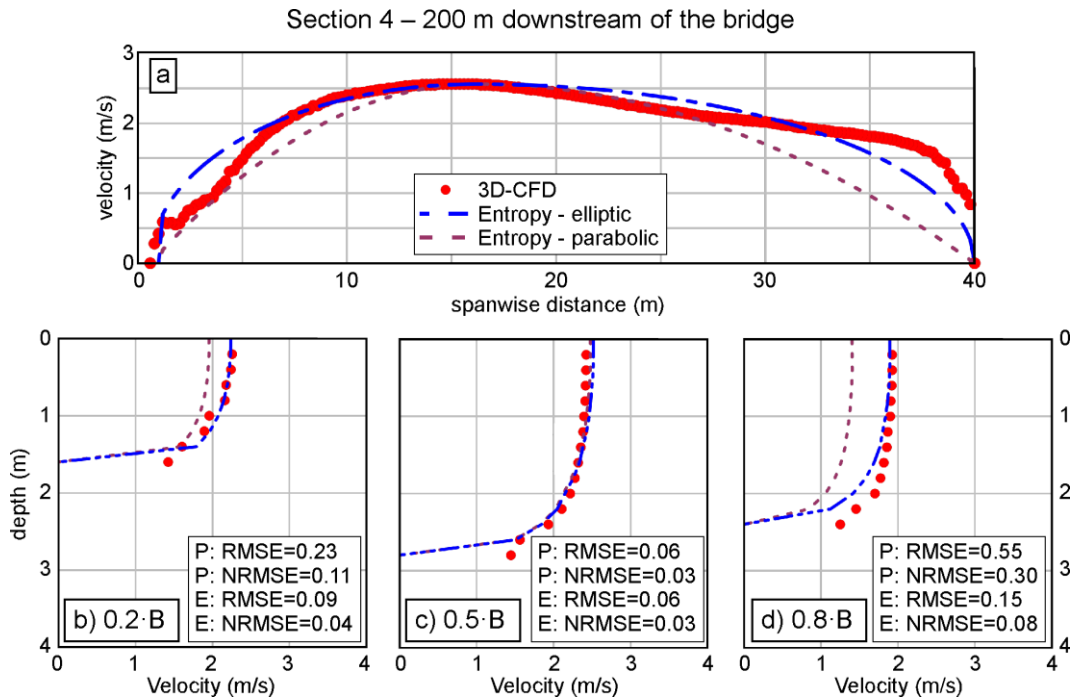
A last point worth of discussing regards the unusual cross-sectional distribution of flow velocity at section 4 (Figure 9a). The reason why the 3D-CFD model locates the maximum velocity at the left of the talweg is the alternate vortex shedding occurring downstream of the bridge piers, which propagates beyond the last considered cross-section. This is evident in the map of instantaneous surface-velocity of Figure 11. This particular occurrence poses interesting questions on the application of the entropy model to estimate the flow discharge downstream of in-stream structures. First, the spanwise location of the maximum surface velocity is subject to a periodical shift, which prevents its correct detection by means of a fixed sensor with a small-size field of view. Secondly, marked time-varying flow fields, which occasionally (or periodically) deviate from nearly uniform flow conditions, can hardly be captured by any preset velocity distribution. To alleviate the problem, the

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periodic signal of surface velocity can be filtered, which is equivalent to look at time-averaged modelled flow fields, which requires knowing the frequency of vortex shedding.



360 **Figure 10.** Flood event of 2022, cross-section 4 (200 m downstream of the bridge). Spanwise distribution of the surface velocity (a); comparison of vertical distributions of velocity at 0.2*B* (b), 0.5*B* (c), and 0.8*B* (d).

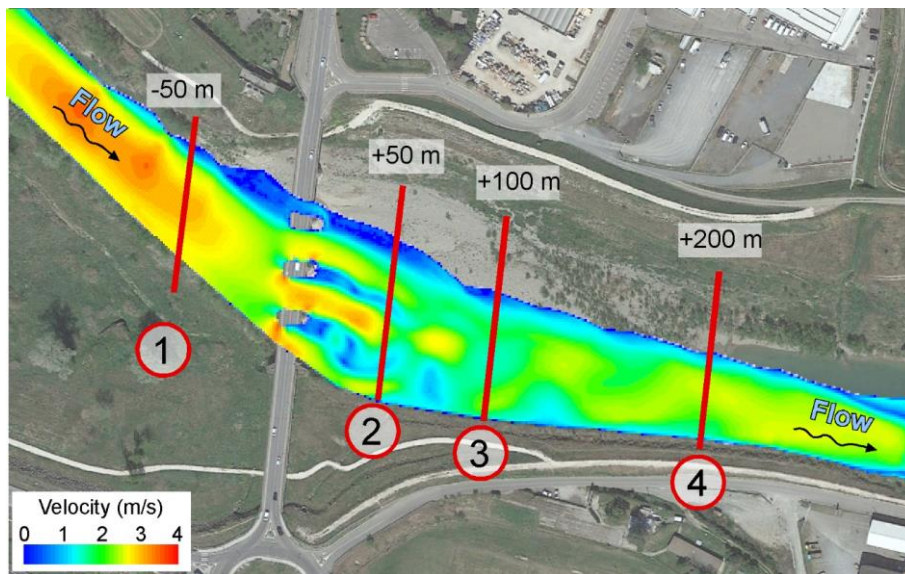


Figure 11. Flood event of 2022. Colormap of the instantaneous surface velocities computed with the 3D-CFD model for the Paglia River at the Adunata bridge (aerial image from © Google Earth).



The results shown in this Section confirm the general accuracy of the entropy model in predicting the velocity distributions. As expected, when using a single value of velocity in place of the river-wide distribution of surface velocity, the precision of the method decreases. Provided that using a single velocity is beyond the scope of the method when the velocity distribution is markedly irregular, the entropy approach can still be forced with a single surface velocity, and produce accurate results, when there are no evidences of strong disturbances of the flow. Indeed, such an approach cannot capture marked unevenness in the flow field, as shown in the case of the lateral low-velocity regions at cross-section 1 for the 2012 event (Figure 7), and in the time-varying flow field of cross-section 4 for the 2022 event (Figure 9).

Distance from the bridge (m) and year	Average velocity (m/s)			Discharge (m ³ /s)			Error percentage (%)	
	3D-CFD	Entropy		3D-CFD	Entropy		Parabolic	Elliptic
		Parabolic	Elliptic		Parabolic	Elliptic		
-50 (2012)	4.43	4.44	4.83	1'800	1'804	1'962	+0.2	+9.0
-50 (2019)	3.00	3.24	3.40	450	486	510	+8.0	+13.3
-50 (2022)	2.55	2.65	2.74	160	166	172	+3.8	+7.5
+200 (2022)	1.97	1.81	2.02	160	147	164	-8.0	+2.0

Table 4. Comparison between 3D-CFD and entropy-based outputs considering a single surface velocity.

4 Conclusions

The present study investigates the entropy-based cross-sectional distributions of velocity, as well as the associated river discharge, for different flow conditions in a representative case study. As sensors for continuous monitoring of water level and discharge are often mounted on existing bridges, a stretch of the Paglia River is analyzed where a multi-arch bridge with thick piers, which hosts a level gauge and a radar sensor, strongly affects the flow field.

With the goal of assessing the applicability of the entropy model in case of flows disturbed by the presence of in-stream structures, a 3D-CFD model is set up to obtain reliable, physics-based velocity distributions at relevant cross-sections, both upstream and downstream of the bridge. The entropy model is then applied to reproduce this set of velocity distributions, using the modelled surface velocity and the bathymetric data as input.

As a first point, the study highlighted the potential of using accurate, physics-based, 3D-CFD models to deepen the knowledge of river and, specifically, of theoretical methods for discharge estimation. Indeed, 3D-CFD models allows providing pictures of complex flow fields that are more complete than, e.g., ADCP measures, in terms of spatial and temporal distribution and, above all, valid for high-flow regimes which typically prevent any direct measurement of the flow field beneath the free-surface. This entails unexplored chances of outlining best-practices in the use of simplified methods for continuous discharge monitoring, and, as a consequence, to improve their accuracy.

According to the present analysis, the entropy model revealed remarkable skills in reproducing also disturbed and uneven flow fields when the river-wide distribution of the surface velocity is used as input data. This occurred also just downstream of the bridge, where the pier-induced wakes made the velocity distribution multimodal and extremely irregular, with error on



discharge estimates lower than 8%. The availability of innovative measuring techniques, able to collect river-wide surface velocity data at a relatively low cost, adds value to the present findings.

On the other side, the accuracy of the entropy model is reduced when only the maximum surface velocity is used as input data, so that the spanwise velocity distribution has to be assumed on a theoretical basis (e.g., parabolic or elliptical). While
395 such a method is absolutely discouraged in case of disturbed flow fields (e.g., downstream of in-stream structures), it still provides accurate estimates where the velocity field is sufficiently regular.

As a final recommendation, measuring instruments and sensor for surface velocity become more effective when placed upstream of in-stream structures, i.e., where the flow field is only marginally affected by the structure and the velocity distribution is far more regular.

400 Future research will include the analysis of stage-dependent variations of cross-sectional velocity distribution, particularly in case of compound cross-sections that are typical of natural rivers. More complex scenarios that still need a comprehensive assessment, and which could largely benefit from physics-based numerical modelling, include the case of mobile beds, in which the geometrical variability occurring at the passage of floods adds uncertainty to the discharge estimation.

Appendix A

405 To impose the boundary conditions to the 3D-CFD model, a 2D depth-averaged model of a longer stretch of the Paglia River has been setup. We used the 2DEF Finite Element model (Defina, 2003; Lazzarin et al., 2023a; Viero, 2019; Viero et al., 2014, 2013), which solves a modified version of the shallow water equations (SWEs) that allow for a robust treatment of wetting and drying over irregular topographies (D'Alpaos and Defina, 2007; Defina, 2000). The SWEs are written as:

$$\eta(h_s) \frac{\partial h_s}{\partial t} + \nabla \cdot \mathbf{q} = 0$$
$$g \nabla h_s + \frac{D}{Dt} \left(\frac{\mathbf{q}}{Y} \right) + \frac{\boldsymbol{\tau}}{\rho Y} - \nabla \cdot \mathbf{Re} = 0$$
(A1)

in which h_s is the free surface elevation, t is the time, ∇ and $\nabla \cdot$ denote the 2D gradient and divergence operators, respectively, $\mathbf{q} = (q_x; q_y)$ is the depth-integrated velocity (i.e., the unit-width discharge), Y is the equivalent water depth (i.e., the volume of water per unit area), and $\eta(h_s)$ a storativity coefficient to account for the wetted fraction of the domain, $\boldsymbol{\tau} = (\tau_x; \tau_y)$ is the bed shear stress, evaluated using the Gauckler-Strickler formula, ρ is the water density, and \mathbf{Re} the horizontal components of the Reynolds stresses, modelled according to the Boussinesq approximation. A mixed Eulerian-Lagrangian approach allows evaluating the total derivative of the flow velocity in the momentum equations using finite differences and a
415 backward tracing technique based on the method of characteristics (Defina, 2003; Giraldo, 2003; Walters and Casulli, 1998). Then, the SWEs are solved with a finite element method, based on triangular, unstructured grids. The model also allows to couple 2D triangular elements with 1D elements (either open- or closed-sections) to model the minor hydraulic network efficiently; other 1D elements are used to model particular devices, such as pumps, weirs, etc. (Martini et al., 2004). The



420 model has been successfully used to simulate flows in various rivers (e.g., Mel et al., 2020a, 2020b; Viero et al., 2019); its effectiveness have been demonstrated also in different research field, such as lagoon and marine environments (e.g., Carniello et al., 2012; Pivato et al., 2020; Tognin et al., 2022; Viero and Defina, 2016).

In the present case, the computational mesh covered a stretch of the Paglia river about 7 km long, from 600 m upstream of the Adunata Bridge to the confluence with the Tiber river, including floodable floodplains (Fig. A1). The average mesh size ranges from 10 m in the riverbed near the Adunata bridge, to 30 m in the floodplains and far downstream of the Adunata
425 bridge. The computational mesh included 61'000 triangular elements, 16 1D elements to simulate underpasses, and 4 1D weir elements to simulate the sill located 500 m downstream of the Adunata Bridge.

The inflow hydrographs, prescribed at the upstream mesh inlet, were derived from water levels measured at the Adunata Bridge using the associated rating curve. At the outlet, an arbitrary rating curve was applied as downstream boundary condition; a sensitivity analysis showed that, because of the distance from the Adunata Bridge, this boundary condition did
430 not produce any perceivable effect in the water levels simulated at the study site.

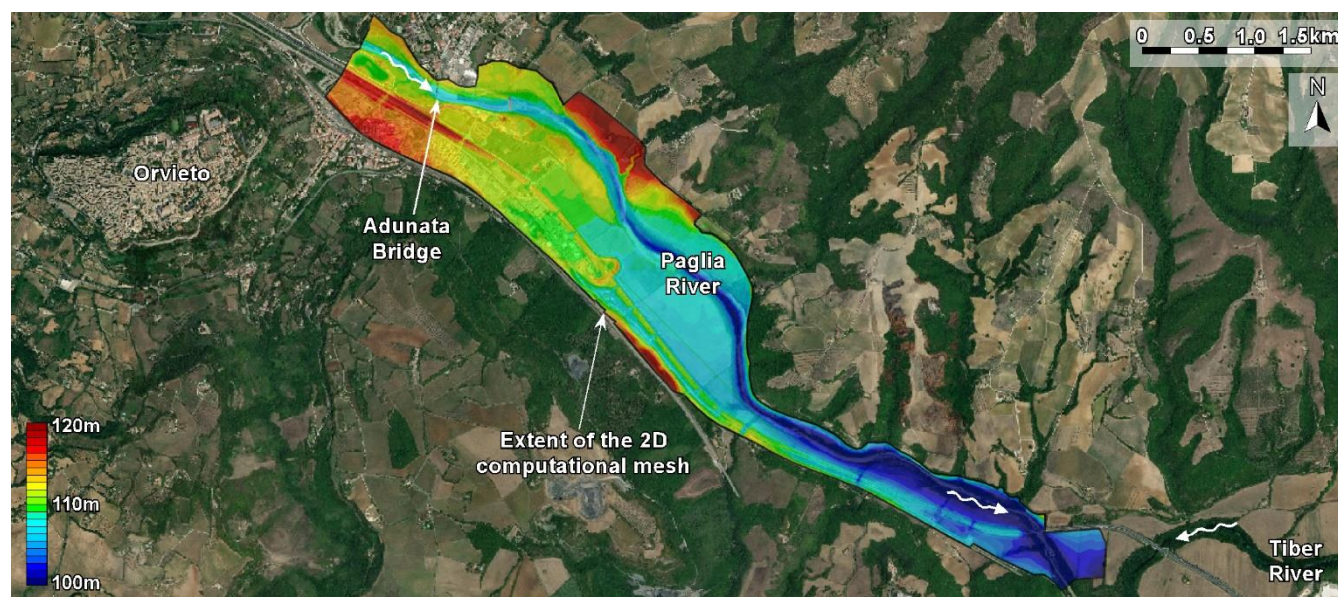


Figure A1. Spatial extent of the 2D computational mesh (aerial image from World Imagery). The color map shows the bottom elevation of the grid elements derived from the LiDAR-based DTM.

Different Gauckler-Strickler coefficients were assigned to the different parts of the domain (e.g., floodplains, densely
435 vegetated areas, etc.) based on the soil cover. The value assigned to the main riverbed were calibrated to match the time series of the water levels measured at the Adunata bridge gauging station for the 2019 flood event (Fig. A2a) and, for the most sever flood event occurred in 2012, the model results were also checked in terms of extent of flooded areas. The minor flood of 2022 was used to verify the model (Fig. A2b). Finally, the depth-averaged velocity just downstream of the Adunata Bridge was compared with the free-surface velocity measured by the radar sensor. Due to the use of a coarse grid and to the



440 depth-average assumption, the 2D model underpredicts the measured water surface systematically (Fig. A2c,d); however, using an amplification factor of 1.7 (gray dots in Fig. A2c,d), the predicted values are quite similar to the measured ones.

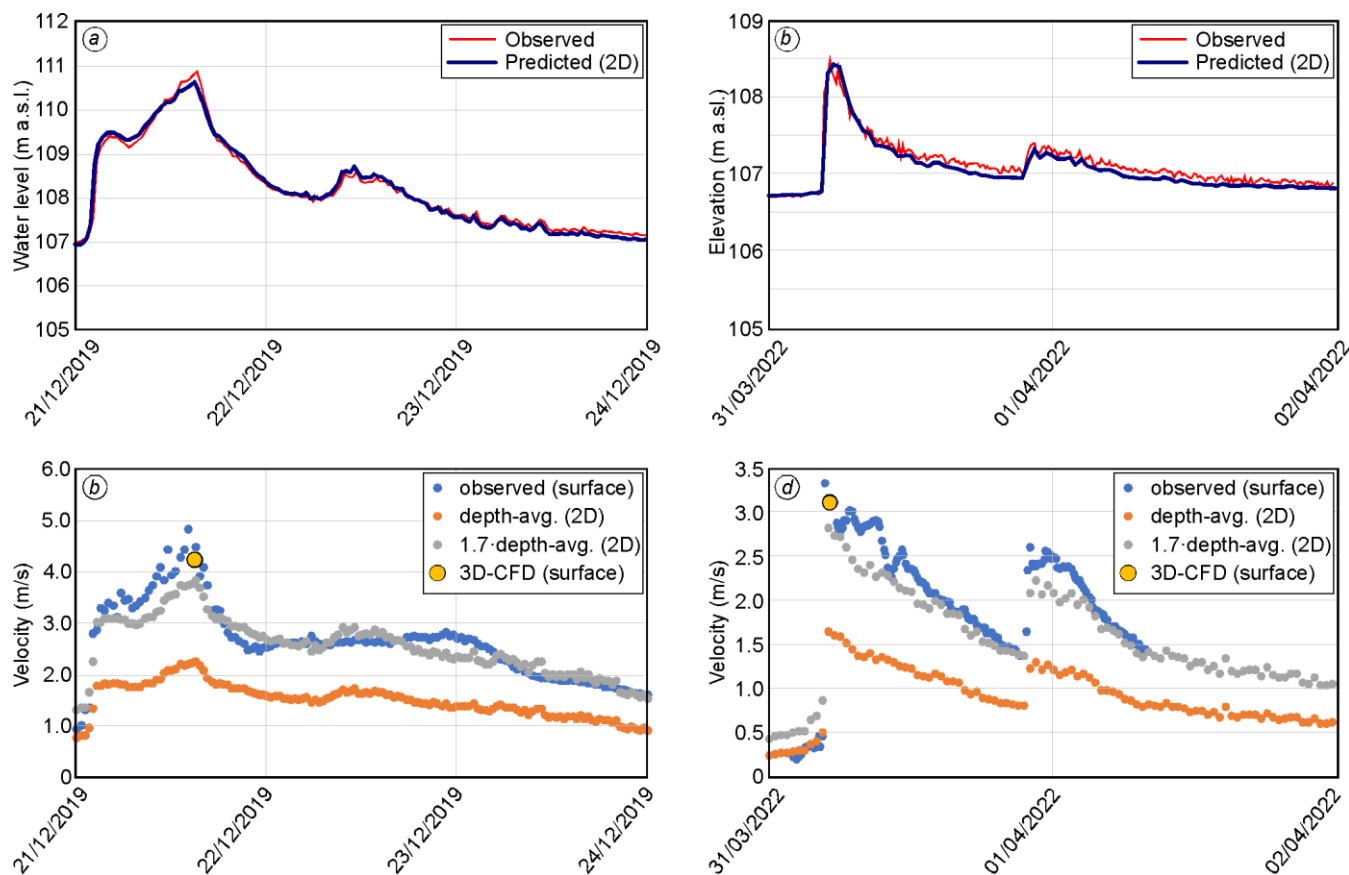


Figure A2. Observed (red) and predicted (blue) water levels at the Adunata Bridge gauging station for the flood events of 2019 (a) and 2022 (b). Observed and predicted water velocity for the flood events of 2019 (c) and 2022 (d).

445 Data availability

Data available on request from the authors.

Author contribution

Conceptualization: FB, TL, SB, TM, DPV. Formal analysis: FB and TL. Funding acquisition: TM. Investigation: FB and TL. Methodology: FB, TL, SB, TM, DPV. Project administration: SB, TM, DPV. Software: FB, TL, TM, DPV. Supervision: SB, TM, DPV. Visualization: FB, TL, DPV. Writing – original draft preparation: FB. Writing – review & editing: TL, SB, TM, DPV.



Competing interests

The authors declare that they have no conflict of interest.

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