Quantify and reduce flood forecast uncertainty by the CHUP-BMA method

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Abstract. The Bayesian model averaging (BMA), hydrological uncertainty processor (HUP), and HUP-BMA methods have been widely used to quantify flood forecast uncertainty. This study proposed the CHUP-BMA method by introducing a copula-based HUP in the framework of BMA to bypass the need for normal quantile transformation of the HUP-BMA method. The proposed ensemble forecast scheme consists of 8 members (two forecast precipitation inputs, two advanced long short-term memory (LSTM) models, and two objective functions used to calibrate parameters) and is applied to the interval basin between Xiangjiaba and Three Gorges Reservoir (TGR) dam-site. The ensemble forecast performance of the HUP-BMA and CHUP-BMA methods is explored in the 6-168h forecast horizons. The TGR inflow forecasting results show that the two methods can improve the forecast accuracy over the selected member with the best forecast accuracy, and the CHUP-BMA performs much better than the HUP-BMA. Compared with the HUP-BMA method, the forecast interval width and continuous ranked probability score metrics of the CHUP-BMA method are reduced by a maximum of 28.42% and 17.86% within all forecast horizons, respectively. The probability forecast of the CHUP-BMA method has better reliability and sharpness and is more suitable for flood ensemble forecasts, providing reliable risk information for flood control decision-making.

1 Introduction

Accurate and reliable flood forecasting is one of the necessary measures to reduce flood disasters and improve water resource utilization (Zhou et al., 2019; Vegad and Mishra, 2022). With the development of hydrological theory and flood forecasting techniques, the flood forecasting accuracy and lead time have been significantly improved in recent years (Xu et al., 2022; Cui et al., 2023). However, neither physically-based and conceptual hydrological models nor data-driven models can guarantee to obtain perfect forecasting in real conditions. Because of the influence of the changing environment and the limitations of human perception of complex hydrological processes, the meteorological forcing and other inputs, hydrological model structure, and parameters, etc., contain significant uncertainties (Cloke et al., 2009), which leads to the simulation and forecast results of the model inevitably containing integrated uncertainties from multiple sources (Liu et al., 2022). Traditional flood forecasting schemes are mostly deterministic forecast results without considering forecast

30 uncertainty (Zhong et al., 2018a; Gelfan et al., 2018), which makes decision-makers unable to grasp useful risk information

31 beyond the forecast value. Excessive superstition on a single forecast value will likely lead to poor decision-making

32 (Krzysztofowicz et al., 1999). Therefore, it is essential to quantify and reduce flood forecast uncertainty in practical

33 applications.

34   Probabilistic flood forecasting is one of the effective methods to quantify integrated forecast uncertainty (Matthews et

35 al., 2022). It not only provides a deterministic forecast value, but also provides forecast uncertainty (or risk) information by

36 means of quantile, confidence interval, or density function (Biondi and Todini, 2018; Ferretti et al., 2020; Zhou et al., 2022),

37 which is more scientifically reasonable and practically useful compared with deterministic forecasts and helps decision-

38 makers consider forecast risk quantitatively (Todini, 2008). Various probabilistic forecasting methods based on statistical

39 post-processing of numerical forecast data have been developed in recent years. Among these methods, probabilistic

40 ensemble forecasting is considered to overcome the limitations of a single model or a simple average with fixed model

41 weights (Han and Coulibaly, 2017) and contains richer forecast information because it can consider the ensemble forecast

42 results of multiple models to quantify and reduce integrated uncertainty that contains uncertainties in the inputs, model

43 structure, and parameters (Li et al., 2017; Saleh et al., 2016). Bayesian model averaging (BMA), proposed by Raftery et al.

44 (2005), uses the Bayesian theory and a total probability formulation to transform ensemble forecasts into probabilistic

45 forecasts and is one of the most representative and reliable methods that has been widely used to supplement uncertainty

46 information beyond point estimates (Shu et al., 2022).

47   The BMA method has been applied to temperature, precipitation, and wind speed ensemble forecasts of meteorological

48 forcing (Raftery et al 2005; Sloughter et al, 2007; Sloughter et al, 2010). After confirming that the BMA method can

49 effectively quantify forecast uncertainty and obtain highly accurate deterministic forecasts, it is widely used in hydrological

50 forecasting to quantify forecast uncertainty from different sources, such as model inputs, structure, and parameters. The

51 standard BMA method assumes that each member's posterior probability distribution approximately obeys a normal

52 distribution (Huang et al., 2019; Guo et al., 2021). However, some variables, such as wind speed, rainfall, runoff, etc.,

53 usually obey skewed distributions and require methods such as Box-Cox to convert non-Gaussian variables to standard

54 normal variables that affect the accuracy of probability distribution estimation (Duan et al., 2007; Liu et al., 2018). Many

55 authors have investigated the applicability of BMA in flood ensemble forecasting and tried to overcome its limitations

56 (Madadgar and Moradkhani, 2014; Darbandsari and Coulibaly,2020). Sloughter et al. (2010) proposed an improved BMA

57 method by assuming that the posterior probability distribution of each member could obey a specific non-normal distribution

58 (e.g., Gamma distribution) and using the member forecast values to estimate the mean and variance of the distribution.

59 Madadgar and Moradkhani (2014) introduced the Copula function to solve the posterior probability distribution of members

60 in the BMA method and proposed the Copula-based BMA method, which avoids the assumption of the posterior probability

61 distribution and further reduces the application limitation of the BMA method. In order to ensure that the quantiles of

62 forecast distributions after Box-Cox transformation are within the actual physical range, Baran et al. (2019) introduced upper

63 and lower truncated normal distributions into the BMA, and found that the double truncated BMA had reliable forecasting

64 ability compared to ensemble model output statistics. The advantage was more obvious when rolling window training

65 periods are used. Hemri et al. (2013) introduced the principle of geostatistical output perturbation into the BMA method and

66 proposed a multivariate BMA, which extended the membership probability distribution into a multivariate normal

67 distribution function. Relative to the univariate BMA method, the multivariate BMA can not only consider the temporal

68 correlation between forecast flows, but also improve the forecast reliability when the forecast system was changing, i.e.,

69 fewer models were available due to dropping out at particular lead times. Meanwhile, the BMA method usually ensembles

70 the forecast results of multiple models to be as close to the actual values as possible. However, too many ensemble members

71 may generate redundant information. Darbandsari and Coulibaly (2020) introduced the Shannon entropy theory to select the

72 forecast members that satisfy the above conditions before applying BMA. Their results showed that the BMA method

73 incorporating entropy could improve the probabilistic forecasting performance for high flows over the standard BMA

74 method. In addition, some studies have developed various methods based on the BMA principle, such as the multi-model

75 ensemble forecasting method based on Vine Copula (Zhang et al., 2022) and the combination of BMA and data assimilation

76 techniques (Parrish et al., 2012).

77   However, most studies ignore an essential issue: the BMA does not consider the constraint of initial conditions (i.e.,

78 observed flow at the start of the forecast). It can be shown from Raftery et al. (2005) that the conditional distribution of the

79 member ($Q_{f,i}$) in the BMA is assumed to follow the normal distribution with expectation $\mu_i = a_i + b_i \cdot Q_{f,i}$ (a_i and b_i are the bias

80 correction coefficients) and variance σ_i, which implies that the conditional distribution is only related to the member's

81 forecasted flow and not affected by the observed flow at the forecast start time. It is unreasonable to produce the same

82 posterior distribution when the forecast results are the same at different moments.

83   The hydrological uncertainty processor (HUP) can obtain the posterior distribution function of the actual value under

84 the condition of the forecast value and the observed flow at the start time based on Bayesian principles and the assumption of

85 perfect rainfall forecasting (Krzysztofowicz and Kelly, 2000). Darbandsari and Coulibaly (2021) firstly utilized the HUP

86 method to derive the posterior distribution of each member considering the initial constraints, and then used the BMA

87 method to weight the conditional distribution of all members to obtain the final posterior distribution, which is called the

88 HUP-BMA method. Their results showed that the HUP-BMA method outperforms the HUP method and improves the BMA

89 method in short-term probabilistic forecasting. In addition, the derivability of the posterior distribution for the ensemble

90 members is theoretically enhanced, the heteroskedasticity of the ensemble members is considered, and the interpretability

91 and logical rationality of the BMA method are improved.

92   Although it has been demonstrated that considering initial conditions in the BMA method can improve ensemble

93 forecast performance, there are still issues to be explored. The HUP-BMA method requires a normal quantile conversion

94 method to convert the flow data series to Gaussian space to solve the posterior distribution. The process is not only tedious

95 and complicated, but also prone to bias in the inverse conversion. To this end, Liu et al. (2018) adopted the copula to derive

96 the conditional distribution of the observed flow under the conditions of the forecasted flow, which avoids the assumption

97 that the flow series obeys a normal distribution in the HUP and relaxes the application limitation. The study shows that the

98   CHUP can improve the probabilistic forecasting performance of the HUP method. It is anticipated that coupling CHUP to

99   the BMA may improve the HUP-BMA accuracy and applicability, which motivates the current study.

100   The main innovations and research steps are shown as follows: (1) A novel CHUP-BMA method is proposed for the

101   first time by coupling CHUP into BMA, which not only avoids the normal distribution assumption in HUP-BMA, but also

102   considers the constraints of the initial condition of the forecast. (2) An ensemble forecast containing eight members is

103   constructed by combining two types of forecast precipitation, two long short-term memory (LSTM) models, i.e., the

104   recursive encoder-decoder structure-based LSTM-RED model and the feature-temporal dual attention-based DA-LSTM-

105   RED model, and two objective functions of model calibration. (3) The ensemble forecast performance of the proposed

106   method is analysed and discussed in comparison to the HUP-BMA benchmark method in terms of the deterministic and

107   probabilistic forecasts. The interval basin between Xiangjiaba Dam and the Three Gorges Dam in the Yangtze River, China,

108   is selected as case study.

109   The rest of the paper is organized as follows. Section 2 introduces the case study and materials. The methods are

110   presented in Section 3. Section 4 evaluates the deterministic and ensemble forecast results. Conclusions and prospects are

111   given in Section 5.

112   **2 Case study and materials**

113   **2.1 Study basin**

114   Three Gorges Reservoir (TGR) is the largest hydraulic project in the world and plays a vital role in flood control, power

115   generation, and other water resource management issues (Zhong et al., 2020). The TGR controls a watershed area of about 1

116   million km^2. The total reservoir capacity is about 39.3 billion m^3, with a flood control capacity of about 22.15 billion m^3.

117   The TGR inflow is directly influenced by the runoff yield of the cascade reservoir interval-basin between Xiangjiaba

118   and TGR (Fig.1), with a basin area of about 127,400 km^2 (Zhou et al., 2019). The inflow of the TGR consists of the outflow

119   discharge from the Xiangjiaba Reservoir, the inflow of several tributaries such as Min, Tuo, Jialing, and Wu Rivers, and the

120   rainfall of the interval-basin. The flow sources are complex and have different effects on the TGR inflow. Moreover, TGR is

121   a river-type reservoir with a length of about 600 km at the normal storage level (175m) and an average width of only 1.1 km,

122   resulting in uncertainty in rainfall intensity and storm-centre positioning (Zhong et al., 2020). Therefore, there is significant

123   uncertainty in the flood forecast of TGR. It has been a major challenge to quantify and reduce forecast uncertainty.

124   Table 1 shows the flow propagation time from the hydrological control stations of the mainstream and tributaries to the

125   TGR dam. The outflow discharge of Xiangjiaba Reservoir, located on the Jinsha River, is observed at the Pingshan

126   hydrological station and represents the mainstream flow. The discharge values from large tributaries (Min, Jialing, Tuo, and

127   Wu Rivers) are observed at the Gaochang, Fushun, Beibei, and Wulong hydrological stations, respectively.

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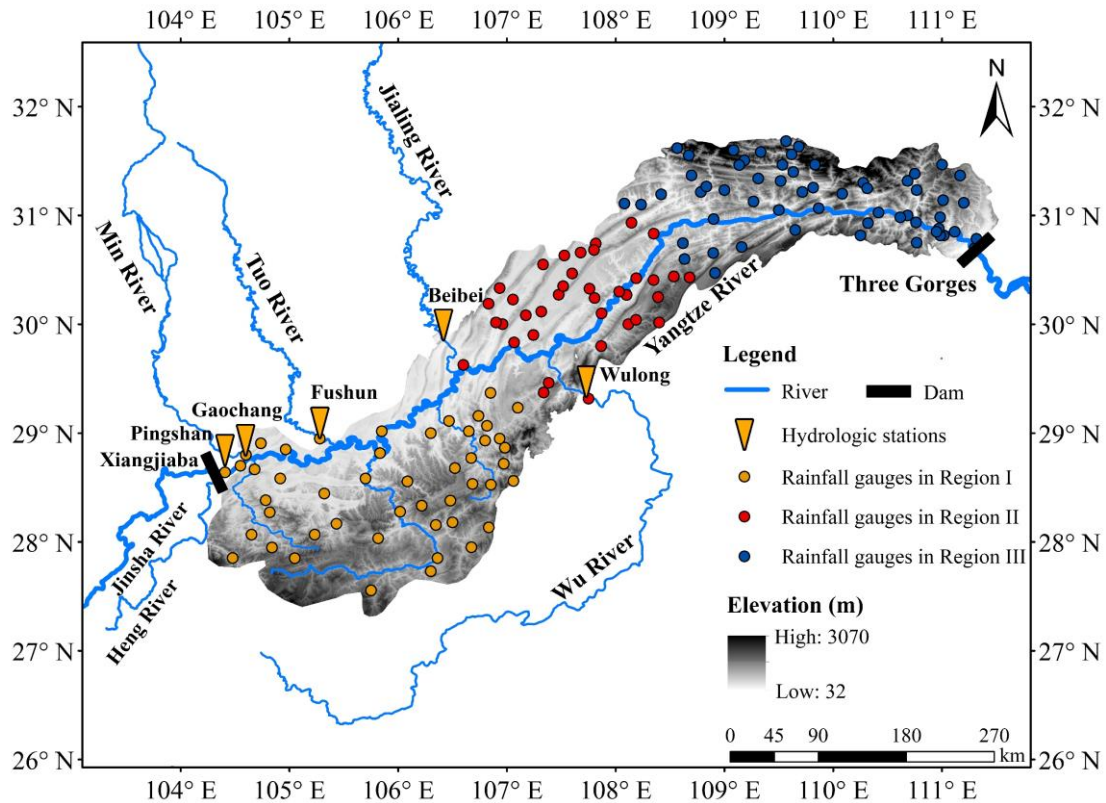

130 **Fig. 1** Schematic diagram of the interval-basin between Xiangjiaba and TGR dam-site which is divided into three sub-regions.

131

132 Considering the uneven distribution of rainfall intensity because of the narrow and long basin, the interval basin

133 between Xiangjiaba and TGR dam-site is divided into three sub-basins: Pingshan-Cuntan, Cuntan-Wanxian, and Wanxian-

134 TGR dam-site. Their watershed areas are 76,900, 22,900, and 27,600 km² respectively. Meanwhile, there are 45, 38, and 60

135 gauged rainfall stations in these three sub-regions, respectively.

136

137   **Table 1** List of flow propagation time for hydrological control stations to TGR dam-site

Rivers	Hydrological control stations	To TGR dam flow propagation time (h)
Jinsha	Pingshan	48-66
Min	Gaochang	48-66
Jialing	Beibei	24-42
Tuo	Fushun	42-60
Wu	Wulong	15-30

138

139 **2.2 Study materials**

140   This study collects 6h observed flow discharges at TGR dam-site and five hydrological stations (Table 1), and 6h

141 observed rainfall in the interval-basin during the 2010-2021 flood season (May-September). The Tyson polygon method is

142 used to calculate areal average rainfall using rainfall station data for each sub-basin area. Meanwhile, this study collects the

143 forecasted precipitation data issued by the European Centre for Medium-Range Weather Forecasts (ECMWF) and the

144 Hydrology Bureau of the Yangtze River Water Resources Commission (HBYRWRC) for the 2017-2021 flood season in the

145 three sub-basins. Their forecast time starts at 8:00, with the 6-168h forecast horizons and the 6h forecast interval. The spatial

146 resolution of each grid for the ECMWF forecasted precipitation is 0.125°×0.125°. The HBYRWRC forecasted precipitation

147 is the areal average forecasted precipitation data.

148   The training period is from 2010 to 2016, and the validation period is from 2017 to 2021. Since the precipitation

149 forecast starts at 8:00 a.m., the forecasted flow for the 6-168h forecast horizons is also calculated from the daily 8:00 a.m. in

150 the validation period.

151 **3 Methods**

152 **3.1 Proposed CHUP-BMA method**

153 **3.1.1 Bayesian model averaging (BMA)**

154   Bayesian model averaging (BMA) method's principle is as follows.

$$p(Q_o|Q_{f,1}, Q_{f,2}, \ldots, Q_{f,k}) = \sum_{i=1}^{k} w_i \cdot p(Q_o|Q_{f,i}) \tag{1}$$

155 where, where, $p(\cdot)$ denotes the probability density function. Q_o denotes the observed flow corresponding to the forecast

156 moment (target value). k is the number of ensemble members. Q_f denotes the forecasted flow of ensemble members. w_i

157 denotes the weight of the i-th model. $p(Q_o/Q_{f,i})$ denotes the conditional probability density of Q_o conditional on $Q_{f,i}$, which is

158 assumed to approximately obey a normal distribution with the expectation of $\mu_{i} = a_i + b_i \cdot Q_{f,i}$ and variance of σ_i. a_i and b_i are the

159 bias correction coefficients obtained by linear fitting of $Q_{f,i}$ to Q_o.

160   Therefore, Eq. (1) can be rewritten as follows.

$$p(Q_o|Q_{f,1}, Q_{f,2}, \ldots, Q_{f,k}) = \sum_{i=1}^{k} w_i \cdot N(Q_o|\mu_i, \sigma_i) \tag{2}$$

161   From Eq. (2), it can be seen that the BMA method does not consider the influence of the initial state (the actual

162 observed flow at the start of the forecast) on the posterior distribution. When the member forecasts at different times are the

163 same, the posterior probability distribution generated by the BMA is also the same, which lacks logical rationality.

3.1.2 Hydrological uncertainty processor (HUP)

Based on the assumption that the precipitation uncertainty is zero, under the condition that the i-th ensemble member forecasts ($Q_{f,i}$) and the observed flow at the start of the forecast (Q_b), the posterior distribution of Q_o derived by the HUP method is as follows.

$$p\left(Q_o|Q_{f,i}, Q_b\right) = \frac{p\left(Q_{f,i}|Q_o, Q_b\right) \cdot p\left(Q_o|Q_b\right)}{\int_{-\infty}^{+\infty} p\left(Q_{f,i}|Q_o, Q_b\right) \cdot p\left(Q_o|Q_b\right) dQ_o} \tag{3}$$

where, $p(Q_o/Q_b)$ is the prior density function, $p(Q_{f,i}/Q_o,Q_b)$ is the likelihood density function. $p(Q_o/Q_{f,i},Q_b)$ is the posterior density function.

The HUP method assumes that flow series transformed to normal space obey the Gaussian distribution. The cumulative distribution function is different for forecasted and observed flows. The common normal quantile transformation is key to the application of the HUP method, which its significance is to make the HUP method applicable to variables with any marginal distributions, heteroskedasticity, and nonlinear dependence structures (Krzysztofowicz and Kelly, 2000; Darbandsari and Coulibaly, 2021).

$$\hat{Q}_o = N^{-1}\left(P(Q_o)\right), \hat{Q}_{f,i} = N^{-1}\left(P(Q_{f,i})\right) \tag{4}$$

where, $P(\cdot)$ denotes the probability distribution function. $N^{-1}(\cdot)$ denotes the inverse function of the standard normal distribution. $\hat{Q}_o$ and $\hat{Q}_{f,i}$ are the observed and forecasted flow transformed to the normal space, respectively.

The HUP method also assumes that the observed flow obeys the strictly stationary first-order Markov process (Krzysztofowicz and Kelly, 2000), i.e., the flows between adjacent forecast horizons obey the linear constraint after the normal transformation.

$$\hat{Q}_{o,t} = c_t \times \hat{Q}_{o,t-1} + \varepsilon_t \tag{5}$$

where, $\hat{Q}_{o,t}$ is the observed flow corresponding to the t-th forecast horizon. c is the regression coefficient. ε is the residual, obeying $N(0,1-c_t^2)$.

The prior density function expressions are as follows.

$$p\left(\hat{Q}_{o,t}|\hat{Q}_b\right) = \frac{1}{(1-C_t^2)^{0.5}} n\left\{\frac{\hat{Q}_{o,t} - C_t \times \hat{Q}_b}{(1-C_t^2)^{0.5}}\right\}, C_t = \prod_{i=1}^{t} c_i \tag{6}$$

where, $n(\cdot)$ denotes standard normal density function; $\hat{Q}_b$ is the observed flow at the start of the forecast transformed to the normal space.

$\hat{Q}_b$, $\hat{Q}_o$, and $\hat{Q}_{f,i}$ are assumed to obey a linear relationship. The expression of the likelihood function in normal space is as follows.

$$\hat{Q}_{f,i,t} = a_t \times \hat{Q}_{o,t} + d_t \times \hat{Q}_b + b_t + \theta_t$$

$$p\left(\hat{Q}_{f,i,t}|\hat{Q}_{o,t}, \hat{Q}_b\right) = \frac{1}{\sigma_t} n\left\{\frac{\hat{Q}_{f,i,t} - (a_t \times \hat{Q}_{o,t} + d_t \times \hat{Q}_b + b_t)}{\sigma_t}\right\} \tag{7}$$

187 where, θ_t is an independent variable obeying N(0,σ_t^2). a_t, d_t, and b_t are regression coefficients.

188   The posterior density function under normal space can be derived by substituting Eqs. (6) and (7) into Eq. (3).

$$p\left(\hat{Q}_{o,t}\middle|\hat{Q}_{f,i,t}, \hat{Q}_b\right) = \frac{1}{Y_t} n\left\{\frac{\hat{Q}_{o,t} - \left(A_t \times \hat{Q}_{f,i,t} + D_t \times \hat{Q}_b + B_t\right)}{Y_t}\right\},$$

$$A_t = \frac{a_t y_t^2}{a_t^2 y_t^2 + \sigma_t^2}, B_t = \frac{-a_t b_t y_t^2}{a_t^2 y_t^2 + \sigma_t^2}, D_t = \frac{C_t \sigma_t^2 - a_t d_t y_t^2}{a_t^2 y_t^2 + \sigma_t^2}, Y_t = \left(\frac{y_t^2 \sigma_t^2}{a_t^2 y_t^2 + \sigma_t^2}\right)^{0.5}, y_t^2 = 1 - C_t^2 \tag{8}$$

189   The posterior distribution function under the normal space can be converted to the original space by Jacobian
190 transformation (Liu et al., 2016). The posterior density function of $Q_{o,t}$ under $Q_{f,i,t}$ and Q_b conditions is as follows.

$$p\left(Q_{o,t}\middle|Q_{f,i,t}, Q_b\right) = \frac{J(Q_{o,t})}{Y_t} n\left\{\frac{N^{-1}\left(P(Q_{o,t})\right) - A_t N^{-1}\left(P(Q_{f,i,t})\right) - D_t N^{-1}(P(Q_b)) - B_t}{Y_t}\right\},$$

$$J(Q_{o,t}) = \frac{p(Q_{o,t})}{n\left(N^{-1}\left(P(Q_{o,t})\right)\right)} \tag{9}$$

191 where, $J(\cdot)$ is the Jacobian transformation function.

3.1.3 HUP-BMA method

193   Darbandsari and Coulibaly et al. (2021) applied the hydrological uncertainty processor (HUP) to the ensemble forecast
194 members, substituted the posterior density function obtained by the HUP method (Eq. (9)) into the BMA framework (Eq.
195 (2)), and then obtained the posterior distribution function of the target flow based on the initial state and the forecasted flow
196 of the ensemble member. Therefore, the expression of the HUP-BMA method is as follows.

$$p\left(Q_o\middle|Q_{f,1}, Q_{f,2}, \ldots, Q_{f,k}, Q_b\right) = \sum_{i=1}^{k} w_i \cdot \frac{J(Q_{o,t})}{Y_t} n\left\{\frac{N^{-1}\left(P(Q_{o,t})\right) - A_t N^{-1}\left(P(Q_{f,i,t})\right) - D_t N^{-1}(P(Q_b)) - B_t}{Y_t}\right\} \tag{10}$$

3.1.4 Copula-based HUP-BMA (CHUP-BMA) method

(1) Copula-based HUP

199 According to Sklar's theorem (Sklar, 1959), the joint distribution of *m* variables is as follows.

$$P(x_1, x_2, \ldots \ldots, x_m) = C_m(P(x_1), P(x_2), \ldots \ldots, P(x_m)) \tag{11}$$

200 where, $C_m(\cdot)$ denotes the m-dimensional copula distribution.

201   The copula-based HUP method (CHUP) was proposed by Liu et al. (2018), which can avoid the normal quantile
202 transformation process of the flow series in the standard HUP method. With the help of the copula function, the prior density
203 function in Eq. (3) can be derived as follows.

$$p(Q_o|Q_b) = \frac{\partial^2 C_2(P(Q_o), P(Q_b))}{\partial P(Q_o)\partial P(Q_b)} \cdot \frac{dP(Q_o)}{dQ_o} = c_2(P(Q_o), P(Q_b)) \cdot p(Q_o) \tag{12}$$

where, $c_m(\cdot)$ denotes the m-dimensional copula density function. m denotes the dimension.

The likelihood density function in Eq. (3) can be derived as follows.

$$p(Q_{f,i}|Q_o,Q_b) = \frac{\dfrac{\partial^3 C_3(P(Q_o),P(Q_{f,i}),P(Q_b))}{\partial P(Q_o)\cdot\partial P(Q_{f,i})\cdot\partial P(Q_b)}}{\dfrac{\partial^2 C_2(P(Q_o),P(Q_b))}{\partial P(Q_o)\cdot\partial P(Q_b)}}\cdot\frac{dP(Q_{f,i})}{dQ_{f,i}} = \frac{c_3(P(Q_o),P(Q_{f,i}),P(Q_b))}{c_2(P(Q_o),P(Q_b))}\cdot p(Q_{f,i}) \tag{13}$$

The posterior density function in Eq. (3) can be derived as follows.

$$p(Q_o|Q_{f,i},Q_b) = \frac{c_3(P(Q_o),P(Q_{f,i}),P(Q_b))}{\int_0^1 c_3(P(Q_o),P(Q_{f,i}),P(Q_b))dP(Q_o)}\cdot p(Q_o) \tag{14}$$

(2) Copula-based HUP-BMA method

Applying CHUP to the i-th ensemble member, the posterior probability distribution function $p(Q_o|Q_{f,i},Q_b)$ of Q_o based on $Q_{f,i}$ and Q_b can be obtained. Coupling $p(Q_o|Q_{f,i},Q_b)$ into the BMA framework, the copula-based HUP-BMA (CHUP-BMA) method can be constructed, and Eq. (2) can become as follows.

$$p(Q_o|Q_{f,1},Q_{f,2},\dots,Q_{f,k},Q_b) = \sum_{i=1}^{k} w_i \cdot \frac{c_3(P(Q_o),P(Q_{f,i}),P(Q_b))}{\int_0^1 c_3(P(Q_o),P(Q_{f,i}),P(Q_b))dP(Q_o)}\cdot p(Q_o) \tag{15}$$

The forecast uncertainty is quantified by the forecast interval with a 90% confidence level. Before constructing the copula, selecting the marginal distribution and the copula type is usually necessary. This study intends to select the appropriate marginal distribution and copula function from five common distribution functions, such as Pearson type III (P-III), Gamma, Normal, Lognormal, and Weibull, and five common copula functions, such as Gumbel-Hougaard, Frank, Clayton, Student-t (Student) and Gaussian copula, according to the root mean square error (RMSE) minimization criterion, respectively. The definition and mathematical expressions of copula functions can be referred to Liu et al. (2018) and Chen and Guo (2019).

Darbandsari and Coulibaly (2021) demonstrated that the HUP-BMA method could improve the probabilistic forecasting performance of the HUP and BMA methods in the short forecast horizons. Therefore, this paper focuses on analyzing and comparing the performance of the HUP-BMA and CHUP-BMA methods. The HUP-BMA and CHUP-BMA methods only calibrate the ensemble members' weights through the Expectation-Maximization (EM) algorithm (Darbandsari and Coulibaly, 2021). Meanwhile, since the forecast accuracy of ensemble members may change with time due to seasonality and other factors (Zhong et al., 2020), the sliding window approach is used to update the weighting parameters. Parrish et al. (2012) and Darbandsari and Coulibaly (2019) have shown that the BMA method with the sliding window can obtain better probabilistic forecast performance compared to the method without the sliding window.

3.2 Ensemble forecasting scheme

An ensemble forecast scheme containing multi-source uncertainties in the model input, the model structure, and the parameter is constructed using a multi-member approach consisting of two forecasted precipitation, two models, and two objective functions used to calibrate parameters, as shown in Fig. 2.

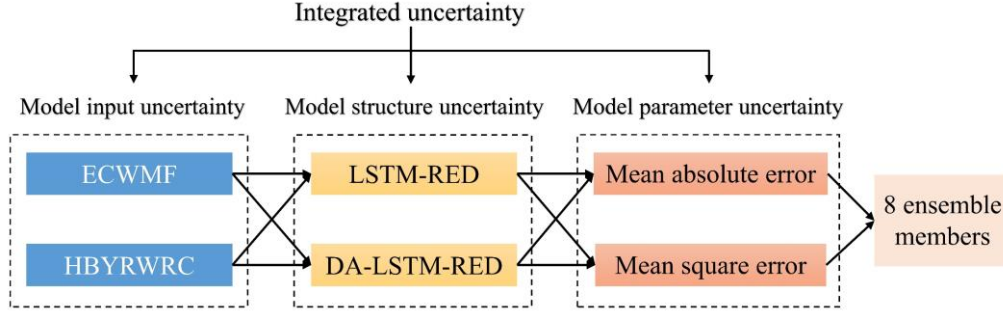



Fig. 2 The TGR's flood ensemble forecast scheme

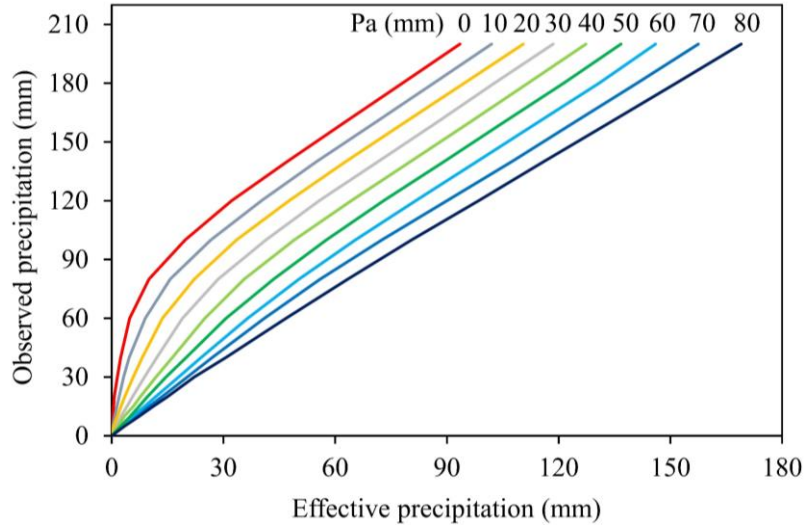
3.2.1 Model input uncertainty

There are five flow discharge inputs from five large tributaries (Jinsha, Min, Jialing, Tuo, and Wu Rivers) in our case study. The flow discharges are observed at the Pingshan, Gaochang, Fushun, Beibei, and Wulong hydrological controlled stations, respectively. Since these observed (or forecasted) flows are respectively regulated by their upstream cascade reservoirs, these flow data inputs are more accurate than the rainfall inputs. This study collected the forecasted precipitation data from the European Centre for Medium-Range Weather Forecasts (ECMWF) and HBYRWRC in these three sub-basins. Since the rainfall data is more diverse and has relatively large uncertainty, the forecast rainfall input variable is used to explore the impact of forecast rainfall uncertainty on the TGR inflow forecasts. The TGR is a river-type reservoir, so building a river confluence model for flood forecasting is necessary. The observed and forecasted precipitations are converted into the effective precipitation in the three sub-basin areas, which accounts for the losses of plant reception, infiltration, evaporation, etc. The rainfall-runoff relationship (Fedora and Beschta, 1989) commonly used in the Yangtze River basin to calculate the effective precipitation. The antecedent precipitation index, which is the key variable of the method, can be calculated by the following equation to represent the soil moisture content (Zhong et al., 2018b).

$$P_{a,t+1} = k(P_{a,t} + P_t) \tag{16}$$

$$P_{a,t+1} \leq I_m \tag{17}$$

where $P_{a,t}$ denotes the antecedent precipitation index on the t-th day, P_t is the daily precipitation, I_m is the water storage capacity of the basin, k denotes evaporation reduction index.

249   The values of k and I_m for these three sub-basins are listed in Table 2, which are obtained from the HBYRWRC. Since
250 the rainfall-runoff relationship graph method have been widely used for runoff generation calculation in the Yangtze River
251 basin, the rainfall-runoff relationship between Xiangjiaba and Three Gorges dam-site uncontrolled interval basin are
252 established and plotted in Fig. 3, which is used to calculate the effective precipitation based on the antecedent precipitation
253 index (P_a) and observed (or forecasted) precipitation for these three sub-basins.
254

255
256 **Fig. 3** Rainfall-runoff relationship between Xiangjiaba and Three Gorges dam-site uncontrolled interval basin
257
258   **Table 2** The k and I_m values for the three sub-basin areas

Sub-basin	k	Im
Pingshan-Cuntan	0.90	50
Cuntan-Wanxian	0.95	80
Wanxian-TGR dam-site	0.95	80

259
260   After obtaining the daily antecedent precipitation index at 8:00, the antecedent precipitation index for the 6-h time scale
261 is calculated as follows.

$$P_{a,t,m} = \left(P_{a,t} + \sum P_{t,n}\right) \times k^{\frac{h}{24}} \tag{18}$$

262 where, $P_{a,t,m}$ denotes the antecedent precipitation index at m:00 on the t-th day. $\sum P_{t,n}$ denotes the cumulative observed
263 precipitation from 8:00 to m:00 on the t-th day. h denotes the time gap from 8:00 to m:00 on the t-th day.

3.2.2 Model structure uncertainty

The TGR inflow forecasting is influenced by the upstream mainstream and tributary reservoir scheduling decisions, the rainfall intensity and distribution in the interval basin, and the changes in the subsurface characteristics, which is challenging to establish complex and physical-based hydrological models (Yang et al., 2019; Cho et al., 2022; Hauswirth et al., 2023). The simulation or forecast accuracy in this interval-basin needs to be improved to meet the needs of the work. Therefore, two advanced data-driven models for obtaining multi-step-ahead flood processes forecasting, namely the long short-term memory (LSTM-RED) model based on an encoder-decoder structure and the coupled dual attention LSTM-RED (DA-LSTM-RED) model, are used for confluence calculations as a way to consider the uncertainty in the model structure. Since the forecast data series at the outlets of tributaries are inconsistent, the observed flow at the outlets of five large tributaries are used to train and validate the proposed models.

(1) Long short-term memory model based on encoder-decoder structure

The structure of LSTM neural network includes forgetting gate, input gate, updating the state of the memory unit, and output gate (Hochreiter and Schmidhuber, 1997). The forgetting gate can select the relatively important information in the previous memory unit. The input gate can select useful information from the input variables at the current moment. The memory unit state can store relatively important information extracted from historical moments, which is updated under the control of the forgetting gate and the input gate. The output gate selects and outputs useful information from the memory cell state. More detailed procedures of the LSTM neural network formulation have been described by Kratzert et al. (2018).

This study nests an LSTM neural network into a recursive encoder-decoder (RED) structure that can be obtained for forecasting flood processes to build an LSTM-RED model. Among them, the RED structure is similar to that of Kao et al. (2020). The description of the LSTM neural network can be found in Cui et al. (2022). The encoding process of the RED structure is used to extract the critical information (C_t) of the input (Xiang et al., 2020). In the decoding process, forecast information of the same category as the encoding process is input to the neural network of the latter moment, besides the C_t and the output of the hidden layer at the previous moment.

(2) LSTM-RED neural network coupled dual attention mechanism

The LSTM-RED model based on the dual attention mechanism (DA-LSTM-RED) is established by adding the feature-temporal dual attention mechanism to the LSTM-RED model, which can enable the model to highlight effective information in different types and moments of the input. The DA mechanism (Fig. 4) consists of the feature attention module, the temporal attention module, and the forecast input processing module.

The feature attention module can adaptively highlight the critical input types by assigning feature weights to the input of the encoding process (Qin et al., 2017). The temporal attention module can highlight the information (hidden layer states) extracted at critical time step by assigning temporal weights to the information extracted at all time step in the encoding process (Ding et al., 2020). Meanwhile, the feature weights are averaged based on temporal weights and applied them to the

296   forecast information inputted in the decoding process, thus highlighting the key forecast input variables. The principle of the

297   DA-LSTM-RED model can be found in Cui et al. (2023).

298

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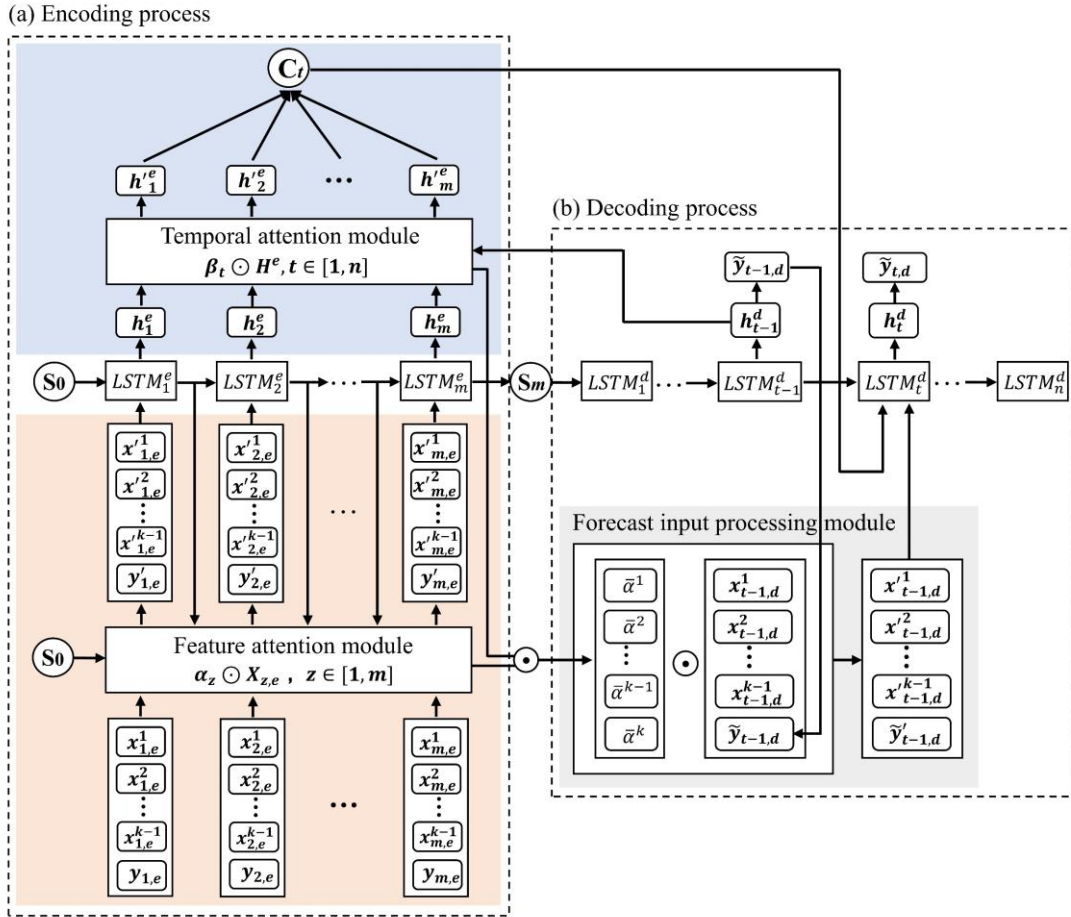
Fig. 4 Schematic diagram of the DA-LSTM-RED model. *e* and *d* are the encoding and decoding processes, respectively. *k* is the number of input types. $X_{z,e}$ is the input variable of encoding process, $X_{z,e}=\{x^1_{z,e}, x^2_{z,e}, ..., x^{k-1}_{z,e}, y_{z,e}\}$. α_z denotes the weights of the input variables, $\alpha_z=\{\alpha^1_z, \alpha^2_z, ..., \alpha^k_z\}$. *m* is the input time-steps in the encoding process. *S* is the hidden layer output. *n* is the maximum forecast horizon. H^e is the hidden layer state, $H^e=\{h^e_1, h^e_2, ..., h^e_m\}$. β_t denotes the weights of the hidden layer states of the encoding process, $\beta_t=\{\beta^1_t, \beta^2_t, ..., \beta^m_t\}$. *C* denotes the key information highlighted by the temporal attention. $\bar{\alpha}$ denotes the forecast input weights.

305

306   **(3) Model input and hyperparameter selection**

307   In this study, the input types for the encoding process include effective precipitation in the three sub-basins, flow

308   discharge in the mainstream and tributaries (i.e., five hydrological stations in Table 1), and previously observed inflow to the

309   TGR for a total of nine data types. In order to make the model learn comprehensive information, input variables with the last

310   11-time steps (66h) are inputted to the encoding process according to the flow propagation times from the hydrological

311   stations to the TGR dam site in Table 1.

312   The forecasted effective precipitation, the forecasted flow of the mainstream and tributaries, and the forecasted flow for

313   the previous forecast horizon are used as inputs of the decoding process. Among them, the forecasted effective precipitation

314   is calculated by the observed precipitation during the training period and by the forecast precipitation during the validation

315   period. The forecasted flow of the upstream mainstream and tributaries is replaced by the observed flow during the training

316   and validation periods. The TGR's observed inflow for the 6-168h forecast horizons is the target output, needed for practical

317   forecasting.

318   The input and output data are handled by the normalization method. Moreover, the trial-and-error method is used for

319   debugging the network hyperparameters. The model is trained by the Adam method (Kingma& Ba, 2014).

320   **3.2.3 Model parameter uncertainty**

321   Different parameter-optimization objective functions may focus on different forecast results (Zhong et al., 2020). For

322   example, the mean absolute error function focuses on the magnitude of the error mean. The mean square error function

323   usually is sensitive to outliers with large errors, which may make the model parameters with different objective functions

324   produce forecast results with different focus points (Duan et al., 2007). Therefore, it is necessary to consider the uncertainty

325   of the model parameters. Neural network models usually train model parameters (such as model internal weights and bias

326   values, etc.) based on loss functions, so this paper uses two common loss functions, namely the mean absolute error and the

327   mean square error, to train the model as a way to consider the uncertainty of model parameters.

328   **3.3 Evaluation metrics**

329   **3.3.1 Deterministic forecast evaluation metrics**

330   The accuracy of deterministic forecast is evaluated by three metrics: the Nash-Sutcliffe efficiency (Nash and Sutcliffe,

331   1970) (NSE), the mean absolute error (MAE) and the relative error of total runoff (RE).

$$NSE = 1 - \frac{\sum_{i=1}^{N}(Q_{o,i} - Q_{f,i})^2}{\sum_{i=1}^{N}(Q_{o,i} - \overline{Q_o})^2} \tag{19}$$

$$RE = \frac{\sum_{i=1}^{N} Q_{f,i} - \sum_{i=1}^{N} Q_{o,i}}{\sum_{i=1}^{N} Q_{o,i}} \times 100\% \tag{20}$$

$$MAE = \frac{1}{N}\sum_{i=1}^{N}|Q_{o,i} - Q_{f,i}| \tag{21}$$

332   where, N is the sample number. $\overline{Q_o}$ and $\overline{Q_f}$ are the average of the observed and forecasted flow, respectively.

333   The Nash-Sutcliffe efficiency (NSE) is one of the most important metrics in flood forecasting, reflecting the degree of

334   fit between forecasted and observed flows (Nash & Sutcliffe, 1970). Since the accurate runoff volume predictions is more

important than peak discharge for the operation of a large reservoir (Cui et al., 2023), the relative error for total runoff volume (RE) is also chosen. The mean absolute error (MAE) can reflect the forecast error for each moment, and compared with the continuous ranked probability score (CRPS) of the ensemble forecast (Raftery et al., 2005), which can reflect the effectiveness of the ensemble forecast correction.

3.3.2 Probabilistic forecast evaluation metrics

(1) Forecast interval evaluation metrics

The forecast interval is evaluated by three metrics: the average coverage rate (CR), average interval width (IW), and the percentage of observations bracketed by the unit confidence Interval (PUCI) (Li et al., 2011).

$$CR = \frac{n_c}{N} \tag{22}$$

$$IW = \frac{1}{N}\sum_{i=1}^{N}(Q_{u,i} - Q_{l,i}) \tag{23}$$

$$PUCI = \frac{CR}{\frac{1}{N}\sum_{i=1}^{N}(\frac{Q_{u,i} - Q_{l,i}}{Q_{o,i}})} \tag{24}$$

where, n_c denotes the number of Q_o located in the forecast interval. Q_u and Q_l are the upper and lower boundaries of the forecast interval with a 90% confidence level, respectively.

The average coverage rate (CR) is one of the most necessary metrics for evaluating the reliability of forecast intervals (Li et al., 2021). The average interval width (IW) is the metric that directly reflects the level of forecast uncertainty, which is an important metric for evaluating the effectiveness of the proposed methods. The percentage of observations bracketed by the unit confidence Interval (PUCI) is a comprehensive metric for evaluating the performance of forecast intervals in quantifying uncertainty (Xiong et al., 2009). Therefore, the CR, RB, and PUCI metrics are selected to evaluate the forecast intervals performance.

(2) Probabilistic forecast evaluation metrics

The probabilistic forecast is evaluated by three metrics: the α_index (Renard et al., 2010), the ignorance score (IGS) (Gneiting et al., 2005), and continuous ranked probability score (CRPS) (Raftery et al., 2005).

$$\alpha\_index = 1 - \frac{2}{N}\sum_{i=1}^{N}|q_{e,i} - q_{th,i}| \tag{25}$$

$$IGS = -\frac{1}{N}\sum_{i=1}^{N}ln(p(Q_{o,i})) \tag{26}$$

$$CRPS = \frac{1}{N}\sum_{i=1}^{N}\int_{0}^{+\infty}\left(P_i(r) - I(r - Q_{o,i})\right)^2 dr, \tag{27}$$

$$I\left(r - Q_{o,i}\right) = \begin{cases} 1 & r \geq Q_{o,i} \\ 0 & r < Q_{o,i} \end{cases}$$

354 where, $q_{e,i}$ and $q_{th,i}$ denote observed and theoretical p-values of $Q_{o,i}$, respectively. p-value denotes the posterior probability

355 distribution value of the $Q_{o,i}$ (Renard et al., 2010). $I(\cdot)$ denotes the indicator function. r denotes the flow variable.

356   The α_index metric can quantitatively assess the reliability of ensemble probabilistic forecasts from the perspective of

357 distribution function values (Renard et al., 2010). The closer the α_index value is to 1, the more reliable the probabilistic

358 forecast is. The IGS and CRPS metrics can reflect the reliability and sharpness of the probabilistic forecast. The former can

359 quantify the forecast probability density at the observation, while the latter can indicate the fit performance between the

360 posterior probabilistic distribution and the actual probabilistic distribution of Q_o (Raftery et al., 2005). Both CRPS and IGS

361 are negative scores, i.e., the smaller the value, the better. The IGS imposes severe penalties for particularly poor probabilistic

362 predictions and may be extremely sensitive to outliers and extreme events, yet also lacks robustness (Raftery et al., 2005).

363 **4 Results evaluation**

364 **4.1 Deterministic forecast results of ensemble member**

365   Since the study focuses on the differences in ensemble forecast performance between the HUP-BMA and CHUP-BMA

366 methods, the overall forecast accuracy of members is analysed (Fig. 5), and the differences in forecast accuracy between

367 members are not explicitly analysed. As shown in Fig. 5, using the observed values as input during the training period, high

368 forecast accuracy can be acquired in different forecast horizons, with the NSE values exceeding 0.95 and the MAE values

369 below 1400 m³/s, and the absolute value of RE within 4%.

370   After combining the forecasted precipitation during the validation period, the NSE values show a decreasing trend, and

371 the MAE and RE values show an increasing trend with the increase of the forecast horizon. Taking the NSE metrics of the 1-

372 7d forecast horizons as an example (Table 3), the average value of the NSE metric decreases from 0.97 to 0.89, which

373 indicates that the forecast accuracy gradually decreases. Meanwhile, the range of evaluation metrics gradually increases with

374 the increase of the forecast horizon. It can be seen from Table 3 that the difference between the maximum and minimum

375 values of NSE indicators for the 1d forecast horizon is only 0.01. In contrast, the difference for the 7d forecast horizon is as

376 high as 0.05, which indicates that the difference in forecast accuracy of members is also more significant, and the forecast

377 uncertainty gradually increases. Overall, the NSE values of the forecast members in the 6-168h forecast horizons are higher

378 than 0.88, and the absolute values of the RE metrics are within 7%. Hence, the forecast accuracy of members is relatively

379 high, and the forecast error is low, which can be used for flood ensemble forecasting.

380

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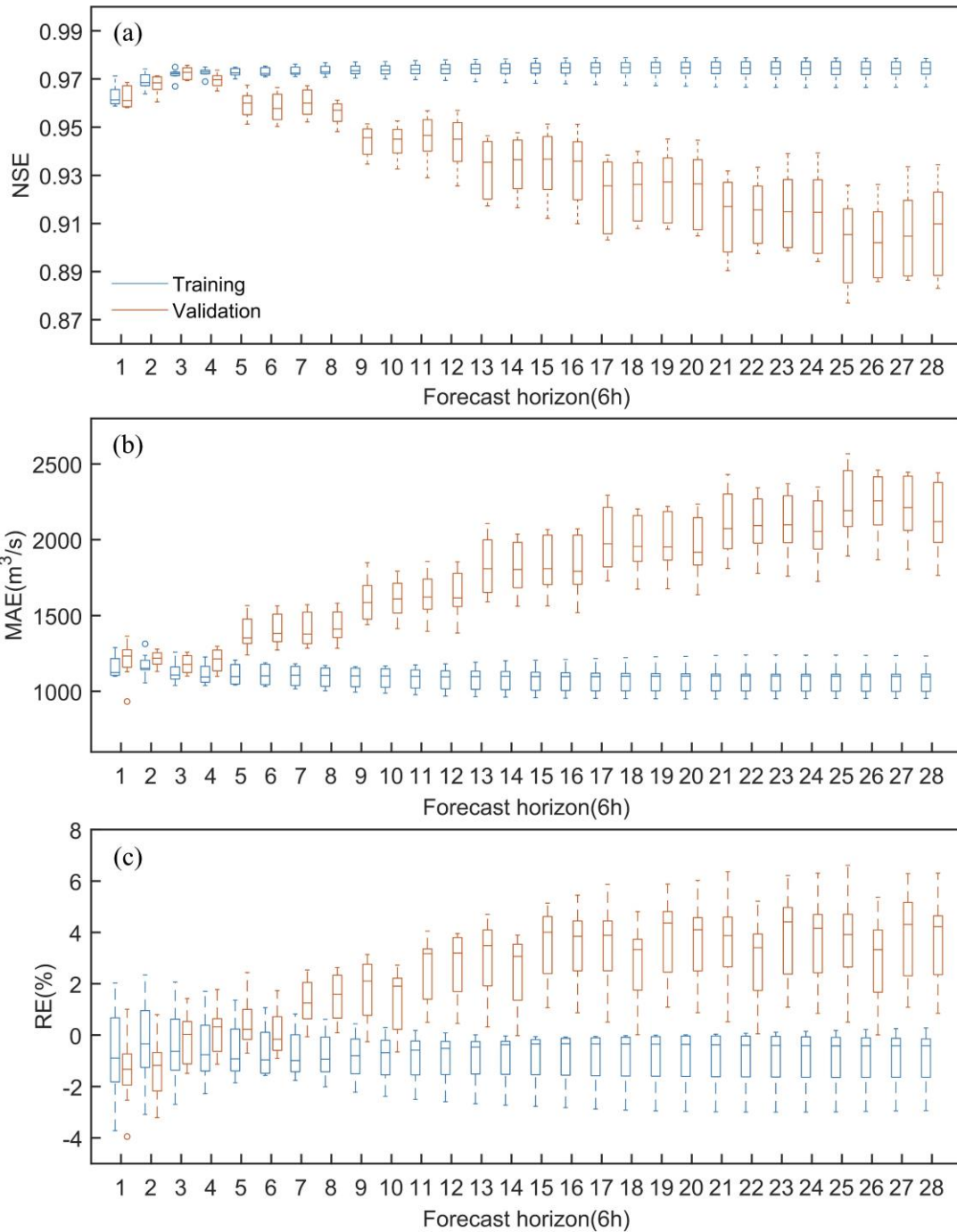


Fig. 5 Statistical chart of evaluation metrics of 8 ensemble members

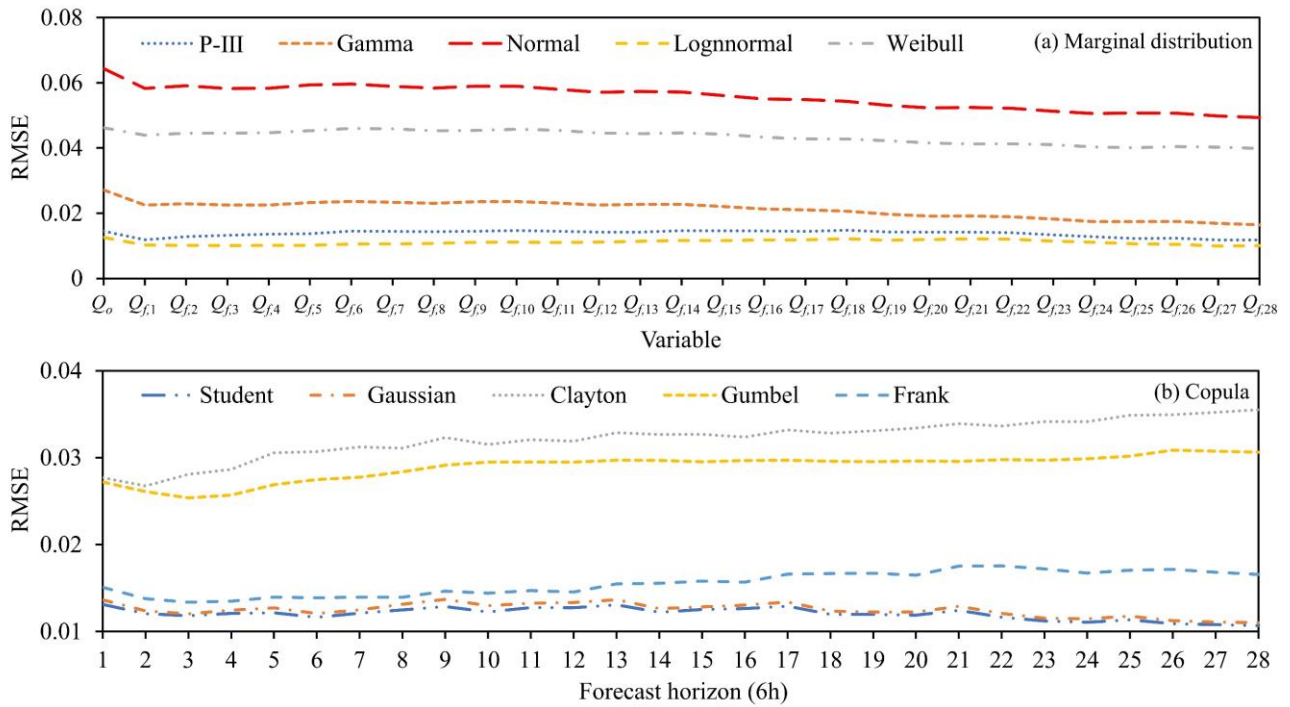
Table 3 Mean, minimum, and maximum values of NSE metrics for 8 ensemble members in the validation period

Forecast horizon (h)	Mean	Max	Min	Forecast horizon (h)	Mean	Max	Min
6	0.96	0.97	0.96	90	0.93	0.95	0.91
12	0.97	0.97	0.96	96	0.93	0.95	0.91
18	0.97	0.98	0.97	102	0.92	0.94	0.90
24	0.97	0.97	0.97	108	0.92	0.94	0.91
30	0.96	0.97	0.95	114	0.93	0.95	0.91
36	0.96	0.97	0.95	120	0.92	0.94	0.90
42	0.96	0.97	0.95	126	0.91	0.93	0.89
48	0.96	0.96	0.95	132	0.91	0.93	0.90
54	0.94	0.95	0.93	138	0.92	0.94	0.90
60	0.94	0.95	0.93	144	0.91	0.94	0.89
66	0.95	0.96	0.93	150	0.90	0.93	0.88
72	0.94	0.96	0.93	156	0.90	0.93	0.89
78	0.93	0.95	0.92	162	0.91	0.93	0.89
84	0.93	0.95	0.92	168	0.91	0.93	0.88

4.2 Ensemble forecast results

4.2.1 Marginal distribution and copula function selection

It is necessary first to fit the marginal distributions of the observed flow and the forecasted flow of the 6~168h forecast horizons. The Q_o and Q_b obey the same distribution. The RMSE criterion is used to select the marginal distribution type. In each forecast horizon, the RMSE values of the 8 members are averaged to obtain the marginal distribution suitable for the forecasted flow intuitively. Meanwhile, according to Eq.(14), the three-dimensional joint distribution of Q_o, Q_b, and Q_f needs to be constructed. The RMSE criterion is used to select the copula function. Similarly, the RMSE values for the 8 members of each forecast horizon were averaged.

Figs.6(a) and (b) show the RMSE values generated by fitting the marginal distribution and copula function, respectively. It can be seen from Fig. 6(a) that the Lognormal distribution has the lowest RMSE value among the five alternative marginal distributions and is chosen as the sequence marginal distribution type. As shown in Fig.6(b), the student copula has the lowest RMSE value in the 6-168h forecast horizons and is chosen to construct the three-dimensional joint distribution function of Q_o, Q_b, and Q_f.

402

403 **Fig. 6** The RMSE values of Q_o, Q_b and Q_f sequence marginal distributions and copula functions. 1, 2, … ,28 denote 6h, 12h, …, 168h

404 forecast horizons, respectively.

4.2.2 Sliding window length selection

406   Since there is no specific method or rule to calculate the sliding window length, this study adopts the CRPS metric as

407 the objective function and the trial-and-error method to select the sliding window length. The range of window lengths is [40,

408 200].

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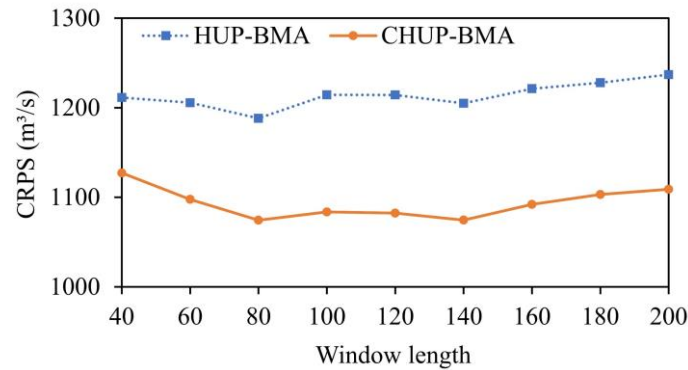
410

411   **Fig. 7** The average CRPS values of the CHUP-BMA and HUP-BMA methods with different window lengths

412

413   To facilitate the selection of the sliding window lengths, Fig. 7 shows the average CRPS values of the HUP-BMA and
414 CHUP-BMA methods for all forecast horizons with different window lengths. It can be seen from Fig. 7 that the HUP-BMA
415 and CHUP-BMA methods all have the lowest CRPS values at the sliding window length of 80. Therefore, 80 is the optimal
416 window length for the ensemble forecasting study.

4.2.3 Deterministic forecast results of ensemble forecast

417

418   The HUP-BMA and CHUP-BMA methods use expected values of ensemble forecasts as deterministic forecast results.
419 In order to analyze the deterministic forecast performance of ensemble forecasts, one member with the best forecast accuracy
420 is selected for comparative analysis based on the criteria of the relatively low RE and MAE values and relatively high NSE
421 values, which is composed of the forecast rainfall from ECWMF, the DA-LSTM-RED model, and the objective
422 function with mean square error to optimize the parameters.

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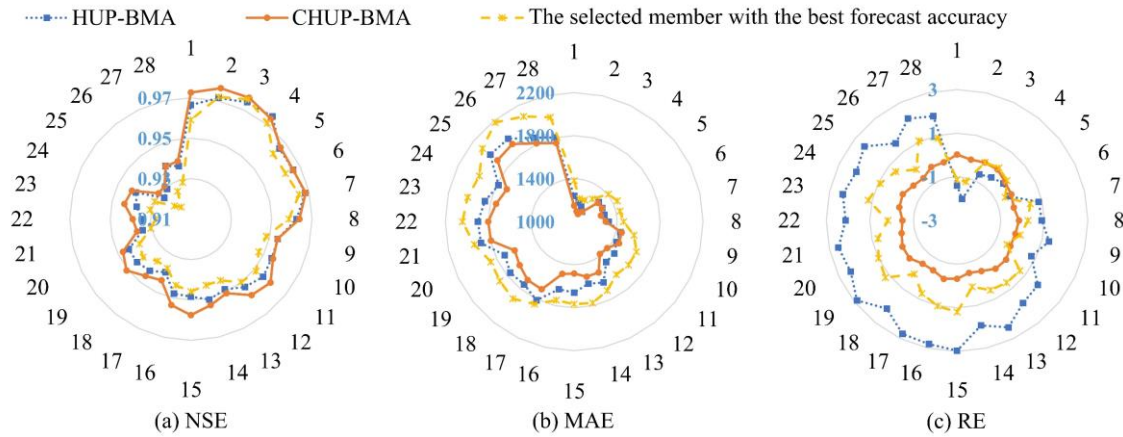

424
425 **Fig. 8** Deterministic forecast evaluation metrics for the HUP-BMA, the CHUP-BMA, and the selected member with the best forecast
426 accuracy

427

428   Fig.8(a), (b), and (c) show the NSE, MAE, and RE metrics of three deterministic forecast results, respectively. It can be
429 seen that the NSE metrics show a decreasing trend, and the MAE metrics show an increasing trend as the forecast horizon
430 increases, indicating a gradual decrease in forecast accuracy.

431   As shown in Fig. 8(a), the NSE metrics of three forecast results are at least 0.92 during the 6-168h forecast horizons.
432 The difference between the two is small, not more than 0.02. Among them, the CHUP-BMA method has the best NSE
433 metrics. However, the advantage value gradually decreases as the forecast horizon increases. The NSE metrics of the HUP-
434 BMA method are better than those of the selected forecast member in most forecast horizons. From Fig.8(b), the maximum
435 and mean values of MAE are 1923 and 1513 m³/s for the CHUP-BMA method, 1999 and 1582 m³/s for the HUP-BMA
436 method, and 2179 and 1719 m³/s for the selected forecast member, respectively. The CHUP-BMA method has the best MAE

metric, with the maximum and average reduction of 10.69% and 4.36% relative to the HUP-BMA method, respectively. Meanwhile, the MAE values of two ensemble forecasting methods are lower than those of the selected forecast members. As shown in Fig.8(c), the maximum and mean of the RE metric are 0.02% and -0.27% for the CHUP-BMA method, 2.97% and 1.36% for the HUP-BMA method, and 1.20% and 0.34% for the selected forecast member, respectively. The CHUP-BMA method can reduce the RE metrics of the selected forecast member in most forecast horizons, while the HUP-BMA method has no advantage in the RE metric. Overall, ensemble forecast methods can somewhat improve the selected best member forecast accuracy. The CHUP-BMA method's expectation forecast has the best accuracy, which indicates that the copula-based CHUP-BMA method can improve the performance of the HUP-BMA method in correcting errors.

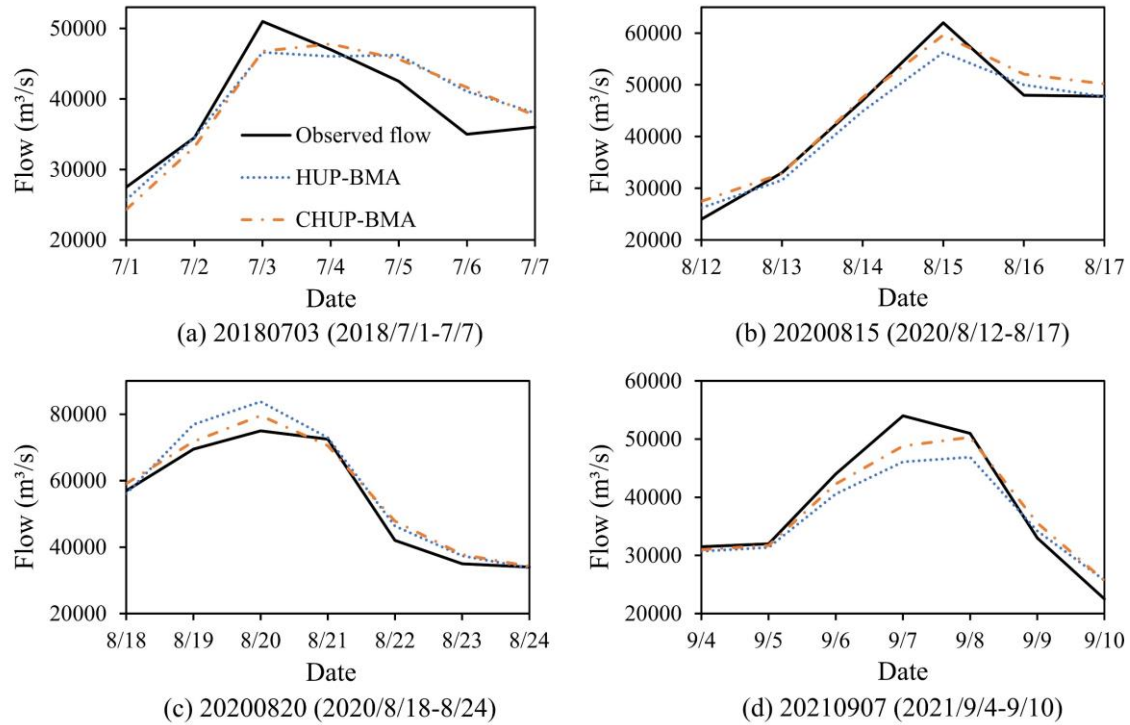
To further analyze the accuracy of ensemble forecast methods, seven floods with peaks exceeding 50,000 m³/s during the 24 and 168h forecast horizons in the validation period (2017-2021) are selected for analyzing. The average relative error metric of peak (PRE) (Cui et al., 2022) is added to analyze the forecasting performance for flood peaks. Table 4 demonstrates the forecast evaluation metrics for the seven flood events. With the increase in the forecast horizon, the NSE metric shows a decreasing trend, and the RE and MAE metrics show an increasing trend, indicating a gradual decrease in forecasting performance. It can be seen from Table 4 that (1) in the 24h forecast horizon, the forecast accuracy of the two methods is similar for most flood events and quality metrics, (2) in the 168h forecast horizon, the forecast accuracy of the CHUP-BMA method is better than HUP-BMA method in most flood events and quality metrics. The average values of the NSE, RE, MAE, and PRE are 0.88, -0.63%, 2980m³/s, and -4.55% for CHUP-BMA, and 0.84, -2.38%, 3188m³/s, and -6.46% for HUP-BMA, respectively, indicating an overall improvement of CHUP-BMA over HUP-BMA in forecasting accuracy.

Table 4 Evaluation metrics for forecast flood events for 24 and 168h forecast horizons

Flood event	Method	Forecast horizon (h)	Evaluation metric			
			NSE	RE (%)	MAE (m³/s)	PRE (%)
20180703 (2018/7/1-7/7)	HUP-BMA	24	0.93	1.95	1697	-3.29
		168	0.80	1.69	2709	-8.60
	CHUP-BMA	24	0.94	3.63	1667	1.64
		168	0.78	1.30	2988	-6.26
20180714 (2018/7/11-7/17)	HUP-BMA	24	0.85	-1.38	2768	-8.04
		168	0.97	0.11	1101	0.88
	CHUP-BMA	24	0.84	-1.97	2874	-7.70
		168	0.95	-2.37	1587	-1.23
20200717 (2020/7/14-7/20)	HUP-BMA	24	0.91	-7.02	3094	-10.02
		168	0.64	-11.67	5965	-19.00
	CHUP-BMA	24	0.91	-4.75	3211	-8.80
		168	0.75	-7.45	5255	-13.58

20200727 (2020/7/25-7/31)	HUP-BMA	24	0.97	-0.22	1371	0.02
		168	0.84	-4.73	3044	-13.47
	CHUP-BMA	24	0.94	4.40	1819	3.62
		168	0.88	0.04	3155	-7.79
20200815 (2020/8/12-8/17)	HUP-BMA	24	0.93	-1.31	2714	-8.21
		168	0.94	-1.96	2259	-9.25
	CHUP-BMA	24	0.96	2.06	2062	-3.53
		168	0.95	3.05	2167	-3.82
20200820 (2020/8/18-8/24)	HUP-BMA	24	0.95	-0.79	2772	0.22
		168	0.92	5.74	3509	11.72
	CHUP-BMA	24	0.96	2.58	2125	2.60
		168	0.96	4.08	2816	6.06
20210907 (2021/9/4-9/10)	HUP-BMA	24	0.94	-3.26	2231	-7.43
		168	0.87	-4.66	3042	-13.15
	CHUP-BMA	24	0.97	-0.64	1722	-4.07
		168	0.94	-0.99	2016	-6.82

457

458   To further demonstrate the accuracy of flood process forecasting and applicability of the two methods, four relatively
459 large flood events are selected for comparative analysis for 168h forecast horizon (Fig.9). In the 20180703-flood event (Fig.
460 9a), the two methods have similar forecast performance, underestimating the peak and rising water processes and
461 overestimating the receding water process. The CHUP-BMA method has relatively low PRE values and total runoff error.
462 The HUP-BMA method accurately forecasts the peak present time. In the 20200815-flood event (Fig.9b), two methods
463 underestimate the flood peak and overestimate the receding water process. The HUP-BMA method has a larger flood peak
464 error, and the CHUP-BMA method has a better fitting performance. In the 20200820-flood event (Fig.9c), two methods
465 overestimate the observed flood process, with the CHUP-BMA method having the lower peak and total runoff error than the
466 HUP-BMA method. In the 20210907-flood event (Fig.9d), the CHUP-BMA and HUP-BMA methods underestimate the
467 flood peak and delay the forecast peak occurring time. The former has smaller peak and water volume error.
468

470 **Fig. 9** Forecasted flood events during 168h forecast horizon for the HUP-BMA and the CHUP-BMA methods

471

472 **4.2.4 Probabilistic forecast results of ensemble forecast**

473   **(1) Evaluation of forecast interval**

474   Figs. 10(a), (b), and (c) show the CR, IW, and PUCI metrics for the forecast interval with a 90% confidence level,

475 respectively. Fig.10(a) shows that during the 6-168h forecasting period, the maximum, minimum, and mean of the CR metric

476 for the forecast interval of the CHUP-BMA method are 0.92, 0.88, and 0.89, respectively, and 0.93, 0.88, and 0.91 for the

477 HUP-BMA method, respectively. The CR values of the two methods' forecast intervals are close to or exceed the 90%

478 confidence level, indicating that the forecast intervals are reliable.

479   It is obvious from Fig.10(b) that the forecast interval width tends to increase with the increase of the forecast horizon,

480 indicating that the forecast uncertainty gradually increases. The maximum, minimum, and mean of the IW metrics for the

481 forecast interval of the CHUP-BMA method are 7820, 3337, and 6257 m³/s, respectively, and 8888, 4662, and 7345 m³/s for

482 the HUP-BMA method, respectively. The forecast intervals of the CHUP-BMA method are significantly narrower than those

483 of the HUP-BMA method, with the maximum and average reduction of 28.42% and 15.32%, respectively, which indicates

484 that the CHUP-BMA method can effectively reduce the interval width and forecast uncertainty.

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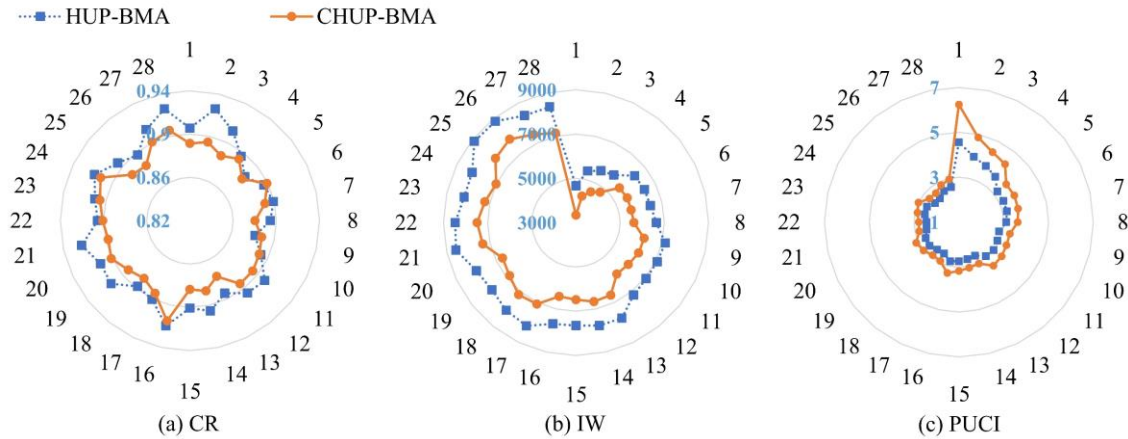
Fig. 10 Evaluation metrics of forecast intervals with the 90% confidence level of the HUP-BMA and CHUP-BMA methods

From Fig.10(c), the maximum, minimum, and mean of the PUCI metric for the forecast interval of the CHUP-BMA method are 6.24, 2.65, and 3.48, respectively, and 4.55, 2.35, and 2.95 for the HUP-BMA method, respectively. The CHUP-BMA method has the higher PUCI values, indicating that the forecast interval of the CHUP-BMA method reflects the forecast uncertainty relatively well.

In summary, the CHUP-BMA outperforms the HUP-BMA method under the premise that the CR values are close to or exceed the 90% confidence level. The CHUP-BMA method has narrower forecast intervals and better performance in quantifying forecast uncertainty. Although the HUP-BMA method has a higher CR value, its IW value is larger, and the PUCI value is smaller for the long forecast horizon, indicating that the forecast interval is too conservative to estimate the uncertainty range reasonably.

In order to visually analyze the ability of the CHUP-BMA method to quantify forecast uncertainty, the forecast intervals with a 90% confidence level of the HUP-BMA and CHUP-BMA methods for 168h forecast horizon in the 2020 flood season are compared. It can be seen from Fig.11 that the forecast intervals of the two ensemble forecasts can cover most of the observed flows and always cover the annual maximum flood peak, indicating that the forecast intervals are reliable. Meanwhile, the forecast intervals of the CHUP-BMA method are remarkably narrower than those of the HUP-BMA method, indicating that the forecast uncertainty of the former is relatively low, which can provide more reasonable risk information for TGR flood control decisions.

(2) Evaluation of overall probabilistic forecast

Fig.12 shows the PIT histograms of the HUP-BMA and CHUP-BMA methods for 24, 96, and 168h forecast horizons. It can be significantly observed that the PIT plots of the HUP-BMA method show a ∩-shaped distribution, which indicates that the forecast distribution is over-dispersed and overestimates the forecast uncertainty, explaining the phenomenon of wide intervals. Meanwhile, the PIT plot of CHUP-BMA is more uniformly distributed than that of the HUP-BMA method, which can obtain a better calibration performance.

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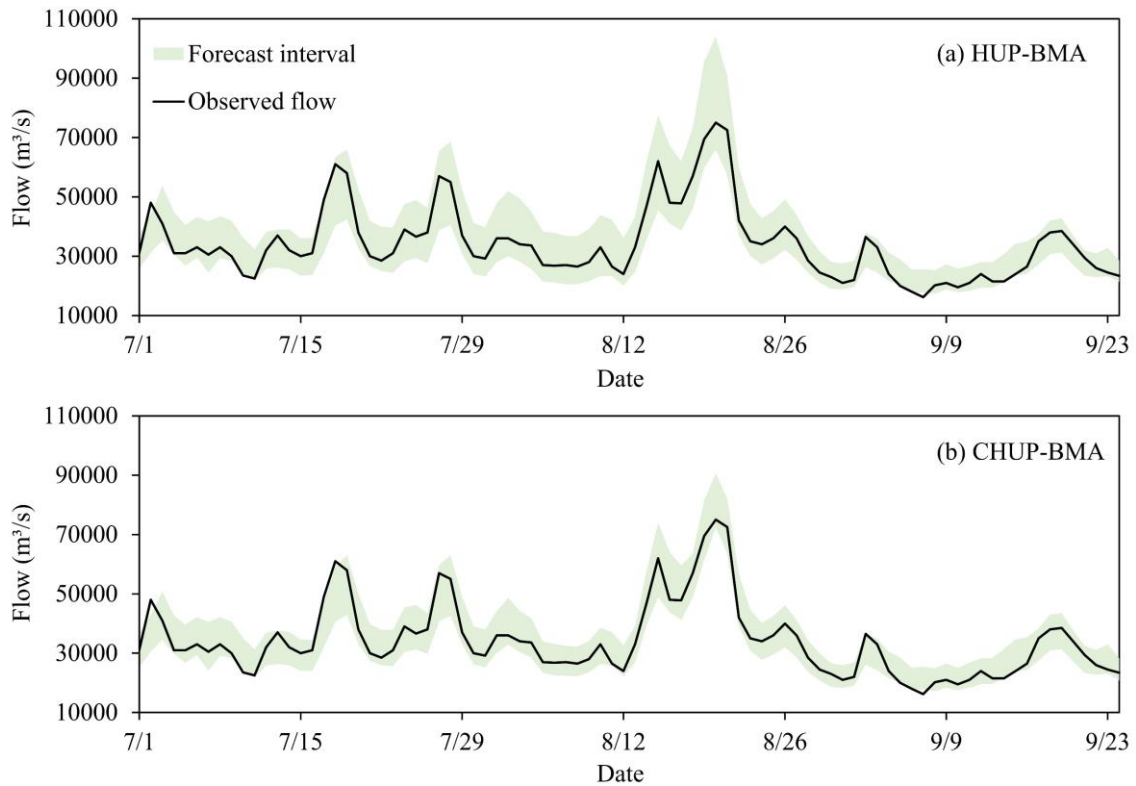
Fig. 11 Forecast intervals with the 90% confidence level for the HUP-BMA and CHUP-BMA methods from 2020/7/1 8:00 to 9/24 8:00

514

Meanwhile, Fig.13(a), (b), and (c) show the evaluation metrics of α_index, IGS, and CRPS metrics for the two ensemble probabilistic forecasts, respectively. It can be seen from Fig. 13(a) that the α_index metrics of the CHUP-BMA method-based probabilistic forecasts are significantly higher than those of the HUP-BMA method in the 6-168h forecast horizons. Among them, the maximum, minimum, and mean of the α_index metric for CHUP-BMA method-based probabilistic forecasts are 0.98, 0.93, and 0.97, respectively, and 0.95, 0.88, and 0.93 for the HUP-BMA method, respectively. The α_index metric of the CHUP-BMA method-based probabilistic forecast is closer to the perfect value of 1, indicating that its probability forecast is the more reliable.

It can be seen from Fig.13(b) that the IGS values of the two methods gradually increase with the increase of the forecast horizon, indicating that the forecast uncertainty gradually increases. The maximum, minimum, and mean of the IGS metric for the CHUP-BMA method are 9.10, 8.33, and 8.87, respectively, and 9.16, 8.59, and 8.98 for the HUP-BMA method, respectively. It can be seen that the IGS metrics of the CHUP-BMA method are consistently lower than those of the HUP-BMA method, which indicates that the CHUP-BMA method has better ensemble forecast performance relative to the HUP-BMA method by assigning a higher probability density around the actual values.

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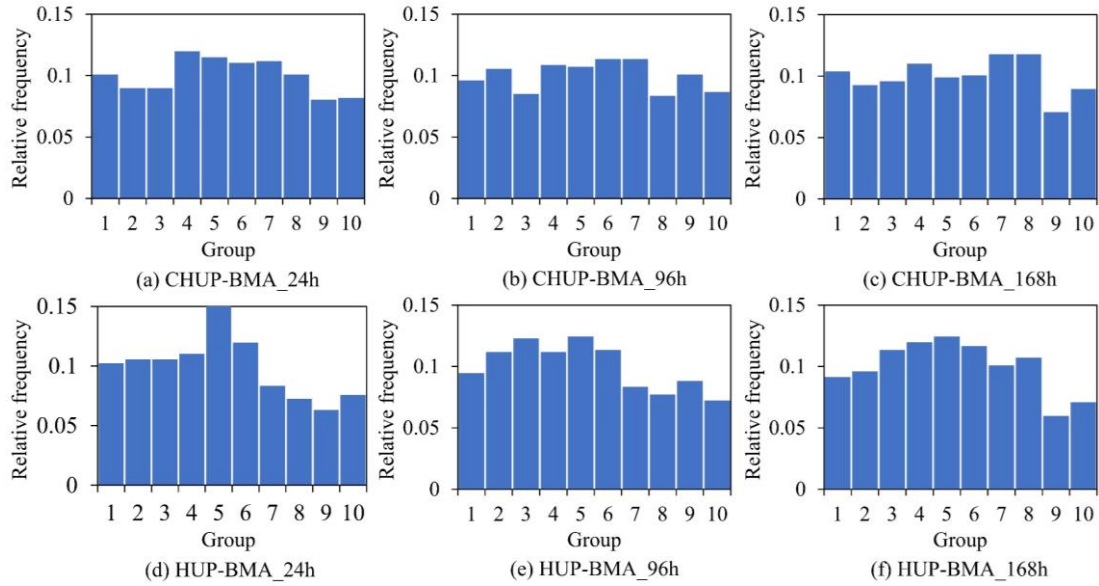
529

530 **Fig. 12** The probability integral transform (PIT) histograms of the HUP-BMA and CHUP-BMA methods for the ensemble forecasts of the
531 24, 96, and 168h forecast horizons.

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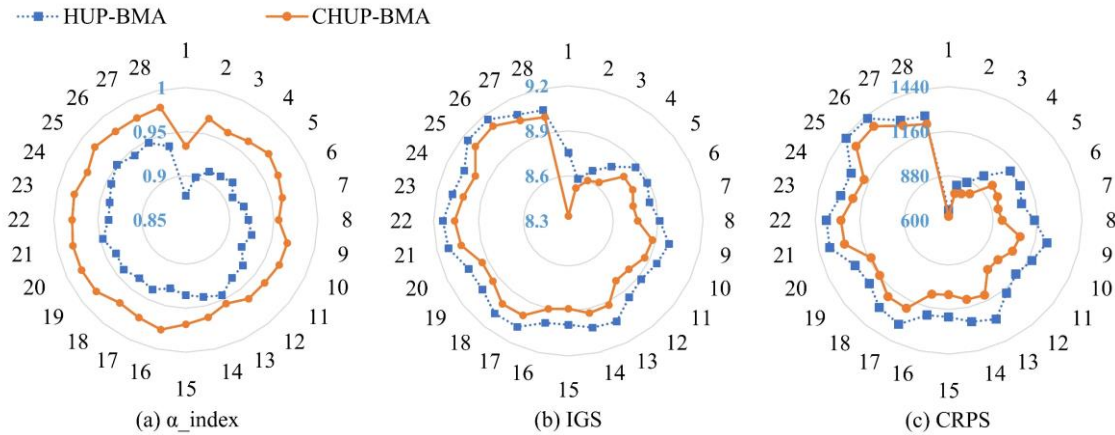
533

534 **Fig. 13** Evaluation metrics of α_index, IGS, and CRPS metrics of two ensemble forecasts. The α_index metric can assess the reliability of
535 ensemble forecasts, while the IGS and CRPS metrics can reflect the reliability and sharpness of the ensemble forecast. The closer the
536 α_index metric is to 1, and the smaller the IGS and CRPS metrics are, the better the performance of the ensemble forecast.

537
538   As shown in Fig. 13(c), the CRPS values of the two methods are lower than the MAE values of the selected member
539 (Fig. 8(b)), indicating that the probabilistic forecasts are effective and can fit the probabilistic distribution of the target values
540 well. Meanwhile, during the 6-168h forecast horizons, the maximum, minimum, and mean of the CRPS metric for the
541 CHUP-BMA method are 1356, 625, and 1074 m³/s, respectively, and 1425, 662, and 1188 m³/s for the HUP-BMA method,

respectively. It can be seen that the CRPS values of the CHUP-BMA method are lower than those of the HUP-BMA method, with a maximum and average reduction of 17.86% and 9.71%, respectively. It can be seen that the CHUP-BMA method can better fit the posterior distribution of the actual values and effectively improve the probabilistic forecast performance of the HUP-BMA method.

From the Table 5, it can be seen that the T-statistics at the 0.05 significance level for all three metrics are higher than the threshold value, indicating that there is a significant difference between the scores of the CHUP-BMA and HUP-BMA methods, i.e., the CHUP-BMA method is significantly better than the HUP-BMA method for ensemble forecasting metrics and performance.

Table 5 T-test results of ensemble forecast metrics at 0.05 significance level

Metric	α_index		IGS		CRPS	
	HUP-BMA	CHUP-BMA	HUP-BMA	CHUP-BMA	HUP-BMA	CHUP-BMA
Mean	0.93	0.97	8.98	8.87	1188	1074
Variance	0.0003	0.0001	0.02	0.03	32247	33716
Degree of freedom	46.00		52.00		54.00	
T-statistic	-10.76		2.36		2.34	
T-threshold	1.68		1.67		1.67	
Difference significance analysis	Significant		Significant		Significant	

In summary, the CHUP-BMA method considers the influence of the initial state on the ensemble forecast, bypasses the normal quantile transformation of the HUP-BMA method, derives the posterior distribution of the target flow without restrictions, and improves the probabilistic forecast performance of the HUP-BMA method. Therefore, the ensemble forecasting by CHUP-BMA method can provide more reasonable and reliable risk information for the TGR.

5 Conclusion and prospects

In this study, we proposed a novel CHUP-BMA method, which not only can consider the influence of the initial state on the ensemble forecast, but also can avoid the assumption of normal distribution in the HUP-BMA method and derive the posterior distribution function more accurately. An ensemble forecast scheme that consists of two forecasted precipitation, two hydrological models, and two objective functions of parameter calibration was established. The ensemble forecasting performance of the HUP-BMA and CHUP-BMA methods was discussed from the perspective of deterministic and probabilistic forecasts. The flood ensemble forecasting experiment with 6-168h forecast horizons was conducted in the Xiangjiaba-TGR dam-site interval basin. The main conclusions were summarized as follows.

565   (1) The two ensemble forecasting methods can improve the members' forecast accuracy. The proposed CHUP-BMA
566 method performs better than the HUP-BMA method, and the MAE metric is reduced by a maximum of 10.69% within 6-
567 168h forecast horizons.

568   (2) The coverage rate of the forecast interval of the CHUP-BMA method is close to or exceeds the specified 90%
569 confidence level, and the forecast interval is significantly narrower than that of the HUP-BMA method, with a maximum
570 reduction of 28.42% during 6-168h forecast horizons, which can effectively reduce the forecast uncertainty.

571   (3) The probabilistic forecast of the CHUP-BMA method has better reliability and sharpness, and its CRPS values are
572 reduced by a maximum of 17.86% relative to the HUP-BMA method, which indicates that the CHUP-BMA method can
573 better fit the posterior distribution of the actual values.

574   (4) The CHUP-BMA method can derive the posterior distribution of the target flow without restriction under the
575 condition of considering the initial constraint, which makes the BMA method more towards perfection. Therefore, it is more
576 suitable for the flood forecasting in the 6-168h forecast horizons and provides reliable risk information for reservoir
577 scheduling decision-making.

578   The present study focuses on flood ensemble forecasting for the TGR's 6-168h forecast horizons. Future studies can
579 explore the ensemble forecasting performance of the proposed CHUP-BMA method for longer forecast horizons and further
580 validate the effectiveness of the proposed method in global basins. Meanwhile, the vine copula, which facilitates multivariate
581 joint distribution modelling, can be considered for constructing the CHUP-BMA method and exploring its advantages and
582 effectiveness in ensemble flood forecasting. And the effective way or method of guiding reservoir scheduling based on
583 ensemble forecasts can be further explored so that ensemble forecasts can be widely used in decision-making.

584 **Appendix A:The model parameters for ensemble membership**

585   We set the number of neural network layers and neurons to be the same for the encoding and decoding processes, with
586 trial-and-error preferences for the number of hidden layers, neurons, and dropout. Meanwhile, the batch size, epoch, and
587 learning rate are set to 100, 500, and 0.001, respectively. The different model parameters are shown in Table A1.

588

589   **Table A1** The model parameters for ensemble membership

Ensemble member type	Neuron	Hidden layers	Dropout
ECMWF&DA-LSTM-RED&MSE	64	1	0.001
ECMWF&LSTM-RED&MSE	64	1	0.001
ECMWF&DA-LSTM-RED&MAE	32	1	0.01
ECMWF&LSTM-RED&MAE	64	1	0.1
HBYRWRC &DA-LSTM-RED&MSE	32	1	0.1

HBYRWRC &LSTM-RED&MSE	32	1	0.001
HBYRWRC &DA-LSTM-RED&MAE	64	1	0.001
HBYRWRC &LSTM-RED&MAE	48	1	0.01

590

Code availability

The code used to support the findings of this study are available from the corresponding author upon request.

Data availability

The data generated and/or analyzed during the current study are not publicly available for legal/ethical reasons but are available from the corresponding author on reasonable request.

Author contributions

Zhen Cui and Shenglian Guo conceived and designed the experiments; Zhen Cui performed the experiments and wrote the manuscript draft; Zhen Cui, Shenglian Guo, Chong-Yu Xu, Hua Chen, Dedi Liu, and Yanlai Zhou reviewed and edited the manuscript.

Competing interests

The authors declare that they have no conflict of interest.

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