



1 Ensemble streamflow prediction considering the influence of 2 reservoirs in India

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9 Abstract

10 Developing an ensemble hydrologic prediction system is essential for reservoir operations and flood early warning.
11 However, efforts to build hydrologic ensemble prediction systems considering the influence of reservoirs have
12 been lacking in India. We examine the potential of the Extended Range Forecast System (ERFS, 16 ensemble
13 members) and Global Ensemble Forecast System (GEFS, 21 ensemble members) forecast for streamflow
14 prediction in India using the Narmada River basin as a testbed. We use the Variable Infiltration Capacity (VIC)
15 with reservoir operations (VIC-Res) scheme to simulate the daily river flow at four locations in the Narmada basin.
16 We examined the streamflow forecast skills of the ERFS forecast for the period 2003-2018 at 1-32 day lead. We
17 compared the streamflow forecast skills of raw meteorological forecasts from ERFS and GEFS at a 1-10 day lead
18 for the summer monsoon (June-September) 2019-2020. The ERFS forecast underestimated extreme precipitation
19 against the observations compared to the GEFS during the summer monsoon of 2019-2020. However, both the
20 forecast products showed better skills for minimum and maximum temperatures than precipitation. Ensemble
21 streamflow forecast from the GEFS performed better than the ERFS during 2019-2020. The performance of the
22 GEFS based ensemble streamflow forecast declines after five days lead. Overall, the GEFS ensemble streamflow
23 forecast can provide reliable skills at a 1-5 day lead. Our findings provide directions for developing a flood early
24 warning system based on ensemble streamflow prediction considering the influence of reservoirs in India.

25 1. Introduction

26 Floods are one of India's most destructive and frequently occurring natural disasters. Floods accounted for about
27 47% of natural disasters in India during the last 100 years (Tripathi, 2016). Riverine floods have been the most
28 common in India, where approximately five million people are affected annually (Luo et al., 2015). In India, the
29 frequency of floods has increased in the past (Singh and Kumar, 2013), with about 20% of the total flood-prone
30 area gets affected every year (Ray et al., 2019). Floods in 2018 caused an economic loss of more than twelve



31 billion dollars (USD) and resulted in the loss of 1808 lives (Joshi, 2020). In addition, climate warming is projected
32 to increase the frequency and intensity of riverine floods (Field et al., 2011; Luo et al., 2015).

33

34 Preparedness for disasters like floods can help mitigate economic loss and human lives (Jain et al., 2018). While
35 financial loss due to floods is projected to rise under the warming climate, human mortality can be reduced with
36 the flood early warning systems and effective communication (Dipti, 2017). Developing a flood prediction system
37 is necessary for early warning and preparedness. Streamflow prediction is an essential component of flood
38 forecasting, which helps in planning and decision-making (Georgakakos et al., 2012; Alfieri et al., 2013). Most of
39 the streamflow prediction systems in India are based on the deterministic approach (Harsha, 2020a; Todini, 2017),
40 which does not account for perturbations in initial conditions to quantify the uncertainty (Bowler et al., 2008).
41 Uncertainty quantification in streamflow prediction can reduce the risk of false alarms (Todini, 2017). In addition,
42 ensemble streamflow prediction is essential for the probabilistic flood forecast. The probabilistic approach
43 performs better than the deterministic approach by quantifying uncertainties associated with flood prediction and
44 early warning system (Krzysztofowicz, 2001). Previous studies used ensemble streamflow prediction in flood
45 forecasting (Cloke and Pappenberger, 2009; Nanditha and Mishra, 2021; Wu et al., 2020) using ensemble
46 meteorological forecast and hydrologic models (Zhang et al., 2020). Ensemble weather forecast provides multiple
47 members at the same location and time that can be used for probabilistic hydrologic prediction. However, several
48 challenges are associated with the operational ensemble streamflow forecast, including computational limitations,
49 explanation of ensemble forecasts to non-experts, and up-gradation in the policy to use the forecast for decision
50 making (Demeritt et al., 2010; Arnal et al., 2020). Despite these challenges, the advantages of ensemble flood
51 forecasts have been reported in previous studies (Pappenberger et al., 2012; Cloke and Pappenberger, 2009).

52

53 The Central Water Commission (CWC) manages flood forecast systems in India. The flood forecast network
54 monitors 325 stations covering low lying areas and towns close to reservoirs. CWC observes real-time water level
55 and discharge along the major rivers of India during the designated flood period. The flood forecast is performed
56 using statistical correlation methods from gauge to gauge. Moreover, Quantitative Precipitation Forecast (QPF)
57 from the India Meteorological Department (IMD) is used to forecast floods at a 3-day lead time (Teja and
58 Umamahesh, 2020). The current flood forecast approach used by CWC is deterministic, which lacks incorporating
59 uncertainties in the forecast and early warning system. An ensemble forecast system can help in the flood early
60 warning and decision making (Harsha, 2020b; Nanditha and Mishra, 2021). Moreover, river basins in India are
61 considerably influenced by reservoirs' presence, and incorporating the influence of reservoirs in streamflow
62 prediction remains a challenge. Incorporating reservoir influence in hydrologic models is essential as reservoirs
63 significantly affect the magnitude and timing of streamflow (Zajac et al., 2017; Yassin et al., 2019; Dang et al.,
64 2019a). However, most of the previous studies on flood forecasts and early warnings in India did not consider the
65 influence of reservoirs (Goswami et al., 2018; Sikder and Hossain, 2019).

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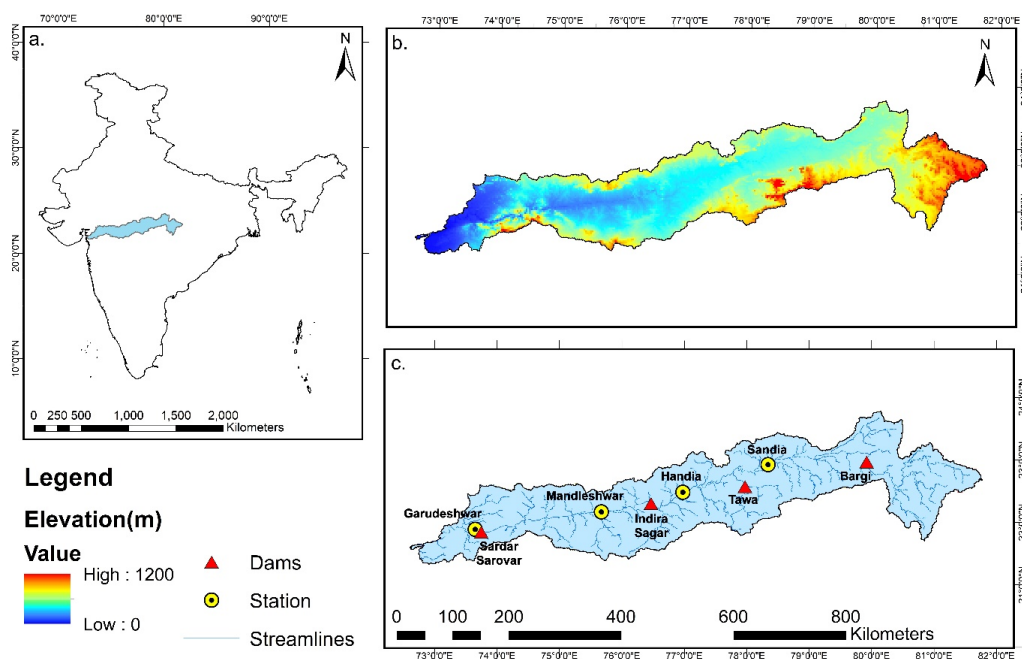
67 Various ensemble forecast products are available from the India Meteorological Department (IMD) and the Indian
68 Institute of Tropical Meteorology (IITM). However, the utility of these forecast products for streamflow prediction
69 and flood early warning at the river basin scale has not been examined. In addition, despite the advantages of
70 ensemble hydrologic prediction, India's current hydrologic forecast systems are mainly deterministic. Given the
71 increasing flood damage in India, the overarching aim of this work is to explore the utility of ensemble forecast
72 products for streamflow prediction in India. We considered the Narmada River basin as a testbed to examine the
73 potential of ensemble hydrologic prediction. We used the Variable Infiltration Capacity (VIC) with reservoir
74 operations (VIC-Res) scheme, which incorporates the effect of reservoirs (Dang et al., 2019). Extended Range
75 Forecast System (ERFS) and Global Ensemble Forecast System (GEFS) ensemble forecasts developed by IITM
76 were used to examine the hydrologic prediction skills at the selected gauge stations in the Narmada basin.

77

78 **2. Data and methods**

79 **2.1 Study region and datasets**

80 Narmada is the fifth biggest and the largest west flowing river in India. The Narmada river basin falls in two states,
81 Gujarat and Madhya Pradesh. Many tributaries contribute to the river through its way to the Arabian Sea, with the
82 Tawa river being its longest tributary. The catchment area of the river basin at the outlet is approximately 98,796
83 km². The upper portion of the basin falls in Madhya Pradesh. The mean annual rainfall in the Narmada basin is
84 1064 mm. Most of the total annual precipitation occurs during the summer monsoon season (June-September).
85 We used observed daily streamflow at four stations: Sandia, Handia, Mandleshwar, and Garudeshwar (Fig. 1).
86 There are several ongoing hydropower and irrigation projects in the Narmada basin. Our hydrologic modelling
87 framework has considered four dams: Bargi, Tawa, Indira Sagar, and Sardar Sarovar. Bargi and Tawa reservoirs
88 were primarily constructed for irrigation purposes. At the same time, Indira Sagar (0.975 Billion Cubic Meters
89 (BCM)) and Sardar Sarovar (5.8 BCM) are the two largest reservoirs that are used for multi-purpose.



90

91 **Figure 1. Basic information about (a) location in India, (b) topography, (c) streamlines, location of streamflow gauge**
92 **stations and reservoirs**

93 We used 0.25° (spatial resolution; ~25 x 25 km) gridded daily precipitation from IMD for the 1951-2020 period
94 (Pai et al., 2014). The daily gridded precipitation product is developed using daily rainfall from 6955 rain gauge
95 stations (Pai et al., 2015). Pai et al. (2015) examined daily rainfall trends, long-term climatology, and variability
96 over the central Indian region. We obtained daily 1° gridded maximum and minimum temperatures from IMD
97 (Srivastava et al., 2009). Srivastava et al. (2009) developed the gridded temperature dataset using observations
98 from 395 quality-controlled stations. We used bilinear interpolation to convert the 1° gridded temperature to 0.25°
99 resolution to make it consistent with the gridded precipitation. The VIC model also requires daily wind speed as
100 an input. We obtained the wind speed from the National Centers for Environmental Prediction (NCEP)-National
101 Centers for Atmospheric Research (NCAR)
102 (<https://psl.noaa.gov/data/gridded/data.ncep.reanalysis.pressure.html>). The wind speed at a coarser (1.875° x
103 1.905°) resolution was interpolated using bilinear interpolation to 0.25° to make it consistent with the other
104 meteorological datasets. The VIC model's vegetation parameters were obtained from the Advanced Very High-
105 Resolution Radiometer (AVHRR) global land cover, available at 1-km spatial resolution (Sheffield and Wood,
106 2007). Soil parameters at 0.25° were developed using the Harmonized World Soil Database (HWSD version 1.2)
107 [Gao et al., 2009]. We used digital elevation model data from Shuttle Radar Topography Mission (SRTM) at 90
108 m spatial resolution (Jarvis, 2008).



109 We observed daily streamflow, reservoir water level, and reservoir live storage data from the India -Water
110 Resources Information System (IWRIS; <http://www.indiawris.gov.in>), which is a joint venture of the Central
111 Water Commission, the Ministry of Jal Shakti, and the Indian Space Research Organization (ISRO). Streamflow
112 and reservoir levels are monitored at various stations within the Narmada basin by CWC. We selected the gauge
113 stations (Sandia, Handia, Mandleshwar, and Garudeshwar) that have observed flow data for at least 15 years. The
114 reservoir storage and water level data were obtained for different periods depending on the data availability.

115 We obtained the Extended Range Forecast System (ERFS) meteorological forecast for the 2003-2020 period. In
116 addition, the Global Ensemble Forecast System (GEFS) meteorological forecast was obtained for the summer
117 monsoon season (July-September) of 2019-2020 from the IITM. Both the ERFS and GEFS forecast products are
118 developed at IITM and are currently being used for the operational weather forecast by the IMD. In June 2018,
119 the high-resolution GEFS forecast was developed and then transferred to the IMD for operational forecasting
120 (Mukhopadhyay et al., 2018). The GEFS dataset has a horizontal resolution of T1534 (~12.5 km) and consists of
121 21 ensemble members (one control and twenty perturbed). The GEFS is being run operationally for the ten-day
122 lead forecast using daily Initial Conditions (ICs) during the summer monsoon period. The GEFS forecast
123 successfully predicted the 2018 Kerala extreme rainfall at 2-3 days lead and showed reasonable forecast skills at
124 5-7 days lead (Mukhopadhyay et al., 2018).

125 The ERFS multi-model system consists of four (CFSv2T382, CFSv2T126, GFSbcT382 and GFSbcT126) suites,
126 each having four ensemble members (one control and three perturbed). Therefore, sixteen ensemble members are
127 available for the ERFS forecast. The model is being run operationally for 32 days lead based on the initial
128 conditions of every Wednesday. We used the sixteen ensemble meteorological forecasts to simulate the daily
129 streamflow at 1-32 days leads at selected stations in the Narmada river basin. Shah et al. (2017) reported that ERFS
130 performed better than the Global Ensemble Forecast System v2 (GEFSv2) and Climate Forecast System v2
131 (CFSv2) in precipitation forecast during the summer monsoon season over India.

132 2.2 VIC-Res hydrologic model

133 We used the VIC-Res hydrologic model, a novel variant of the Variable Infiltration Capacity (VIC) model (Liang
134 et al., 1994a), to simulate streamflow. A combination of the VIC model and the routing model developed by Dang
135 et al. (2019) was used to simulate streamflow at the selected locations in the basin. Dang et al. (2019) incorporated
136 the effect of reservoirs by considering the reservoir storage dynamics and operating rules within the streamflow
137 routing model in the VIC-Res model. Therefore, the VIC-Res model consists of the rainfall-runoff model and a
138 routing model. The rainfall-runoff model generates water and energy flux within each grid using climate forcing,
139 soil parameters, land use/land cover, and the digital elevation model to simulate water and energy fluxes. The
140 model uses vegetation cover for each tile and three soil layers for each grid cell. The upper two soil layers control
141 runoff, infiltration, and evaporation, while the bottom layer governs baseflow. The routing model uses water fluxes
142 (runoff and baseflow) from each grid to simulate streamflow at selected gauge stations using the Saint-Venant



143 equations. The routing model uses flow direction, fraction, and station location as input to generate streamflow.
144 In addition, the VIC-Res model requires reservoir parameters and reservoir location as inputs. Dang et al. (2019)
145 reported that even the model without a reservoir exhibits almost the same accuracy. However, it leads to poor
146 representation of vital hydrological processes, such as infiltration or baseflow. The VIC-Res model also simulates
147 reservoir inflow, outflow, live storage, and water level daily.

148 We used observed daily precipitation, maximum and minimum temperatures from IMD, and wind speed from
149 NCEP-NCAR reanalysis as meteorological forcing. We used historical reservoir storage observations to input the
150 seasonal cycle for each reservoir into the model. An autocalibration module developed by Dang et al. (2019) was
151 used to calibrate soil parameters of the VIC-Res model for the Narmada River basin. The autocalibration module
152 uses the ϵ -NSGAI multi-objective evolutionary algorithm (Reed et al., 2013) to adjust the values of sensitive soil
153 parameters within the soil parameter file. The VIC-Res model simulations were conducted at 0.25° spatial
154 resolution in the Narmada River basin. We used five soil parameters (B_{inf} , D_s , D_{smax} , W_s and depth of three soil
155 layers as described in Mishra et al. (2010) to calibrate daily streamflow at the selected gauge stations in the basin.
156 B_{inf} is a Variable infiltration curve parameter. D_{smax} is the maximum velocity of baseflow. D_s is a fraction of D_{smax}
157 where non-linear baseflow begins. W_s is a fraction of maximum soil moisture non-linear baseflow occurs (Liang
158 et al., 1994b). Further details of the calibration parameters can be obtained from Mishra et al. (2010). The
159 autocalibration module optimizes the model's performance in simulating streamflow at selected stations
160 considering reservoir dynamics. We set our objective to maximize Nash-Sutcliffe Efficiency (NSE) (Dawson et
161 al., 2007; Nash and Sutcliffe, 1970). The model performance was evaluated for daily streamflow, the water level
162 of reservoirs, and the live storage of reservoirs using NSE and coefficient of determination (R^2). Daily streamflow
163 was calibrated and evaluated at Sandia, Handia, Mandleshwar, and Garudeshwar. We selected different periods
164 for the calibration and evaluation of the VIC-Res model based on the availability of observed streamflow. For
165 instance, we selected the years 1986-2000, 1986-2000, 1998-2005, 1998-2005 as the calibration period, while the
166 years 2001-2018, 2001-2018, 2015-2018, 2015-2018 as the evaluation period for stations Sandia, Handia,
167 Mandleshwar, and Garudeshwar, respectively. The VIC-Res model performance was also evaluated against water
168 level and live storage for Bargi, Tawa, Indira Sagar, and Sardar Sarovar reservoirs.

169 We first generated daily meteorological forcing of both ERFs and GEFS forecasts. The ERFs forecast is available
170 for the extended range (1-32 day lead), while the GEFS forecast is available at 1-10 day lead. We developed
171 observed initial conditions for each forecast date by forcing the observed meteorological forcing from IMD into
172 the calibrated VIC-Res model. We simulated a daily streamflow forecast at all the four selected gauge stations
173 using the meteorological forcing and initial conditions. The VIC-Res simulations were run for all the ensemble
174 members for ERFs and GEFS forecasts. The ensemble streamflow forecasts were simulated for 1-32 days lead
175 and ten days lead for ERFs and GEFS datasets. The ERFs forecast simulations were run for 1-32 days lead with
176 the initial conditions of every Wednesday generated from VIC-Res model using the observed forcings. Similarly,



177 GEFS streamflow forecast simulations were performed for 1-10 days lead with initial conditions one day before
178 the forecast.

179 2.3 Forecast skill evaluation

180 We evaluated the skills of the streamflow forecast generated using the ERF5 and GEFS meteorological forecast
181 by comparing the simulated streamflow forecast to the observed daily streamflow at each of the four locations.
182 The model simulated streamflow forecast was evaluated against the VIC-Res model simulated daily streamflow
183 using the observed forcing due to the unavailability of the observed streamflow for the years 2019-2020. The
184 ERF5 meteorological forcing was used to run the VIC-Res model for 1-32 days from each forecast date using the
185 initial condition generated using the observed forcing from IMD. Similarly, we ran the GEFS ensemble members
186 for a 1-10 days lead for each forecast date. We used bias and Normalized Root Mean Square Error (NRMSE) to
187 evaluate the performance of individual ensemble forecast members, which can be estimated as follows:

$$Bias = \sum_{i=1}^n (P_{sim,i} - P_{obs,i}) \quad (1)$$

$$NRMSE = \frac{RMSE}{\bar{O}} \quad (2)$$

where, \bar{O} = mean of observations.

$$RMSE = \sqrt{\frac{\sum_{i=1}^n (P_{sim,i} - P_{obs,i})^2}{n}} \quad (3)$$

188 where $P_{obs,i}$ and $P_{sim,i}$ are observed and simulated streamflow, respectively. Bias provides a measure of
189 correspondence between the mean of observations and the mean of the VIC-Res model simulations, while NRMSE
190 represents the relative magnitude of the squared error.

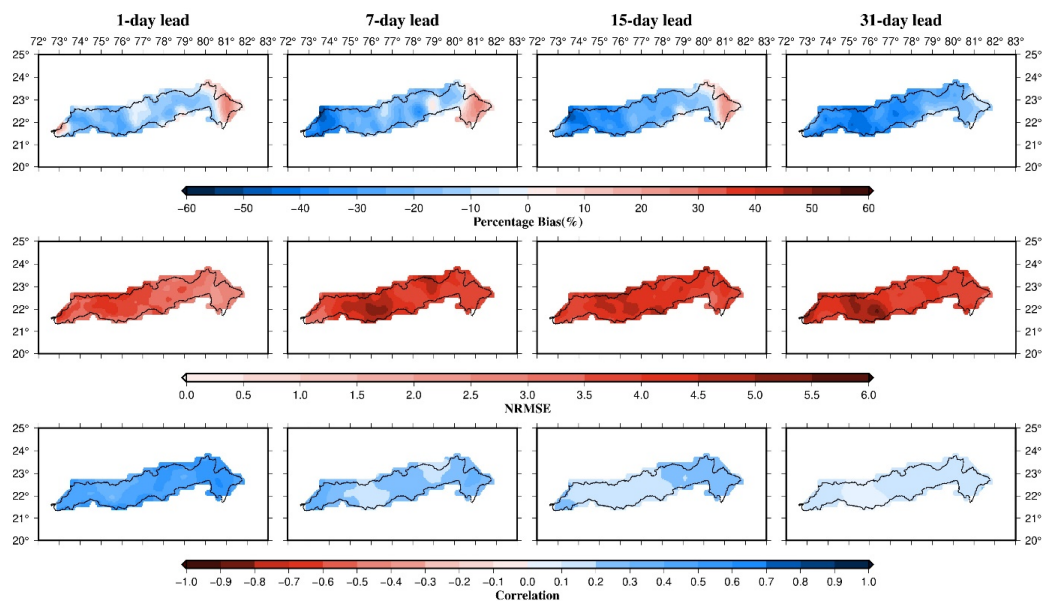
191 3 Results

192 3.1 Skill evaluation of raw meteorological forecasts

193 First, we evaluated ERF5 precipitation and temperature forecast skills for 1-, 7-, 15-, and 31-day leads. We used
194 bias, NRMSE, and correlation coefficient (r) to estimate the forecast skills. The forecast skill was evaluated for
195 the period 2003-2018. We estimated the forecast skill for each ensemble member and then calculated the median

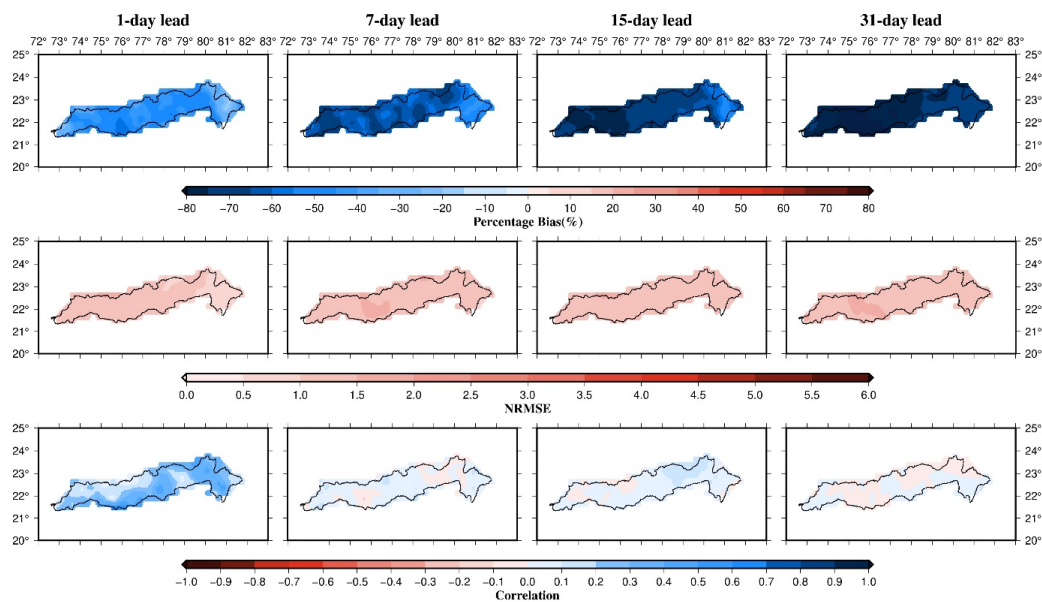


196 of the forecast skill of all the sixteen members for each grid in the Narmada river basin. Precipitation forecast from
197 ERFS showed a negative bias indicating an underestimation compared to observed rainfall. The dry bias in
198 precipitation forecast increases with the lead time (Fig. 2). For the 1-day lead, precipitation forecast from ERFS
199 showed a moderate positive correlation (median ~ 0.49), which declines with the lead time. Similarly, NRMSE in
200 precipitation forecast is large (>2.0) over the river basin. We also estimated bias in the precipitation forecast
201 exceeding the 90th percentile (Fig. 3). The extreme rainfall in the raw ERFS forecast dataset exhibited a weaker
202 correlation with the observed extreme precipitation. Moreover, a considerable dry bias in the extreme precipitation
203 forecast was found. We also evaluated forecast skills for maximum and minimum temperature against the observed
204 temperatures from IMD for the 2003-2018 period (Fig. S1 and S2). The daily temperature forecast showed a
205 relatively higher positive correlation with the observed temperatures from IMD. Moreover, lower NRMSE was
206 noted for the temperature forecast than the observed maximum and minimum temperatures. However, a positive
207 bias of ~ 1.5 °C (median of all grids in the basin) was found in minimum temperature forecast at all the lead times.



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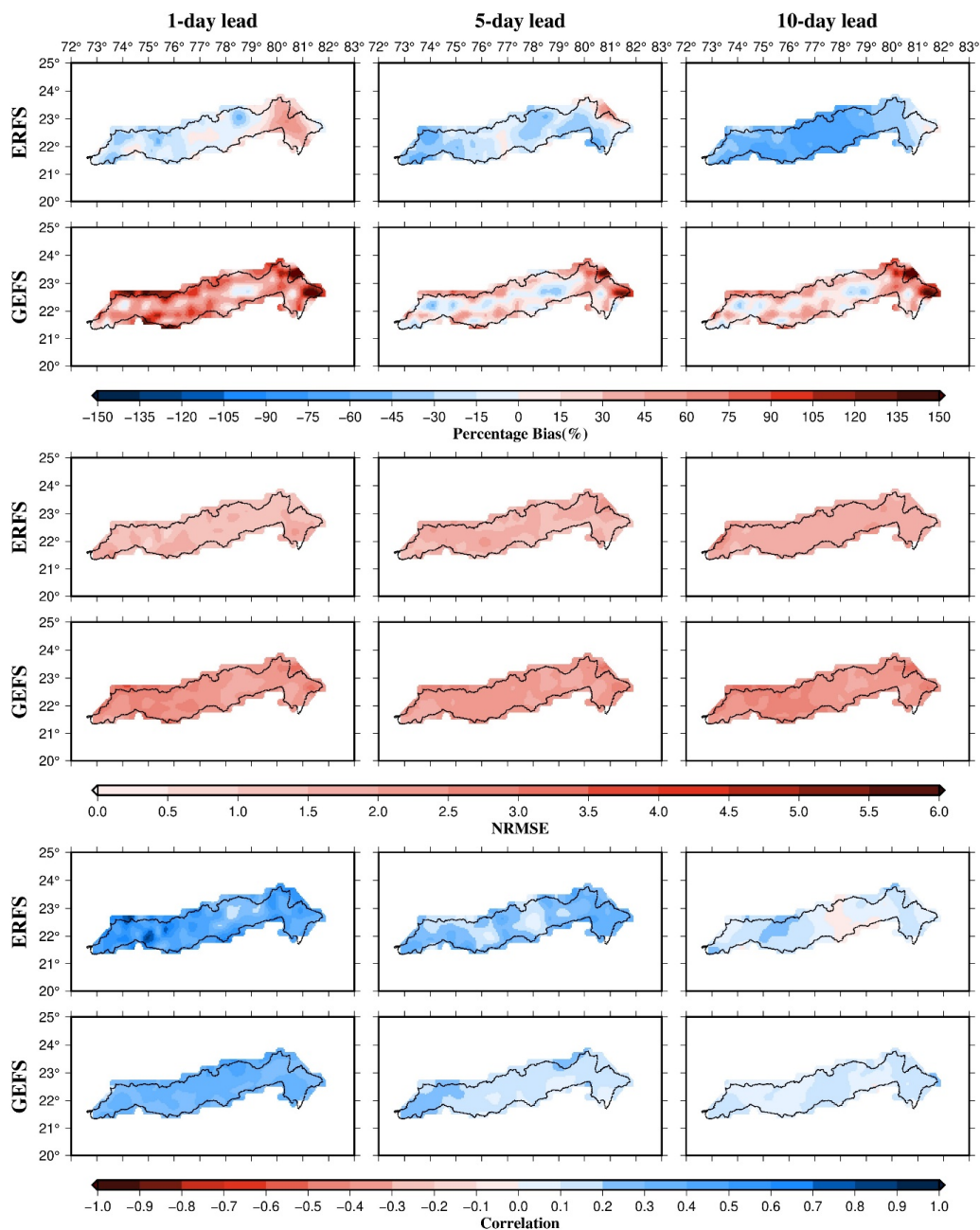
209 **Figure 2. Evaluation of ERFS precipitation forecast against observations for the 2003-2018 period. Forecast skills**
210 **were evaluated using bias, NRMSE, and correlation for each ensemble member and the median skill is presented.**



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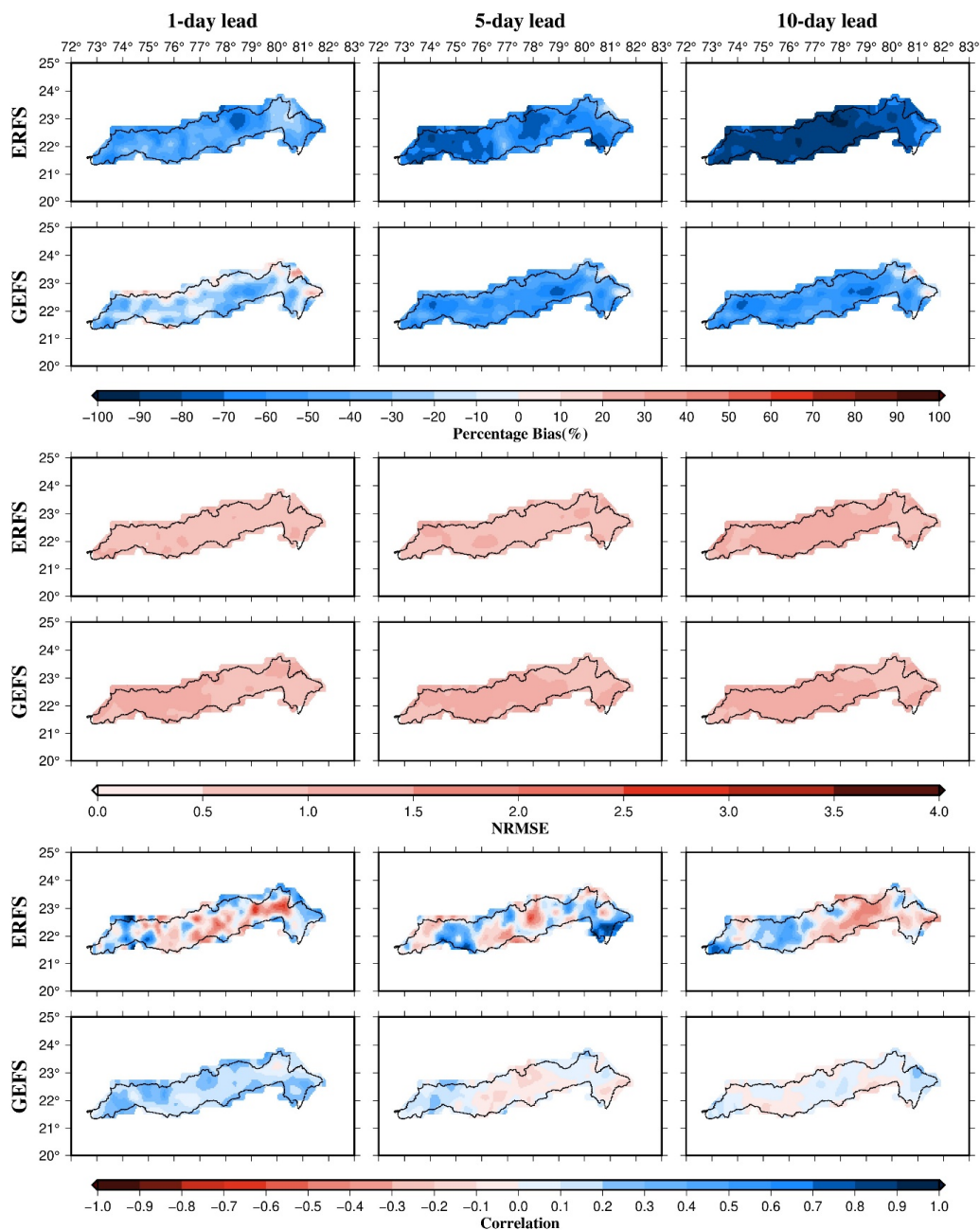
212 **Figure 3, Evaluation extreme precipitation (>90th percentile) forecast skill from ERF5 for the 2003-2018 period.**
213 **Forecast skills were evaluated using bias, NRMSE, and correlation for each ensemble member and the median skill is**
214 **presented.**

215 Next, we compared the ERF5 and GEFS ensemble forecast skills for the summer monsoon (June-September) of
216 the 2019-2020 period. We limit the comparison to the two years as the GEFS ensemble forecast is available only
217 for 2019-2020. We evaluated forecast skills for 1-, 5-, and 10-day lead (Fig. 4). Our results show that the ERF5
218 precipitation forecast has a dry bias across the river basin and all the leads (Fig 4). The GEFS precipitation forecast
219 showed a positive (wet) bias in the majority of the Narmada river basin. The forecast products (ERF5 and GEFS)
220 underestimate extreme rainfall in the Narmada basin (Fig 5). The dry bias in extreme rainfall increases with lead
221 time in the ERF5 and GEFS forecasts (Fig. 5). The forecast products showed a poor correlation with the observed
222 extreme precipitation in the Narmada river basin (Fig. 5). However, both the forecast products demonstrated
223 relatively better skills for maximum and minimum temperatures than precipitation (Fig. S3 and S4).



224

225 **Figure 4.** Comparison of the precipitation forecast skills from ERFs and GEFS for the summer monsoon period
226 during 2019-2020. Forecast skills were evaluated using bias, NRMSE, and correlation for each ensemble member of
227 ERFs and GEFS and the median skill is presented.



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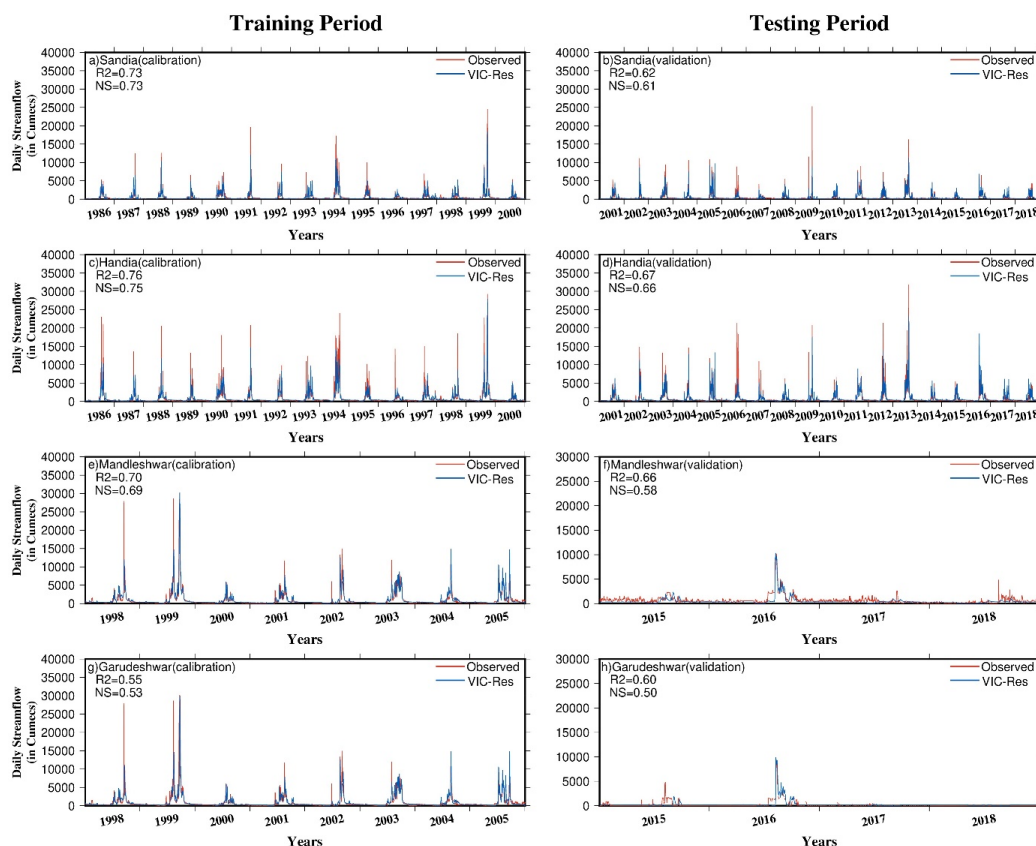
229 **Figure 5.** Comparison of the extreme precipitation (exceeding 75th percentile) forecast skills from ERFs and GEFS for
 230 the summer monsoon period during 2019-2020. Forecast skills were evaluated using bias, NRMSE, and correlation
 231 for each ensemble member of ERFs and GEFS and the median skill is presented.



232

233 3.2 Calibration and evaluation of the VIC-Res model

234 We performed calibration of reservoir level and storage and calibration of daily streamflow. Daily storage and
235 water level calibrated the VIC-Res model for four major reservoirs (Bargi, Tawa, Indira Sagar and Sardar Sarovar)
236 in the Narmada basin. The upstream catchment area of all the gauge locations and calibration parameters are shown
237 in supplementary Figure S5. We evaluated the VIC-Res model's performance using the coefficient of
238 determination (R^2) and Nash Sutcliffe Efficiency (NSE) (Fig. 6). The VIC-Res model simulates daily streamflow
239 at the selected stations in the basin. R^2 and NSE values were above 0.65 at Sandia, Handia, and Mandleshwar
240 stations for the calibration period. While at Garudeshwar, the VIC-Res model performed comparatively weaker
241 ($R^2 = 0.55$ & $NSE = 0.53$) for the calibration period.



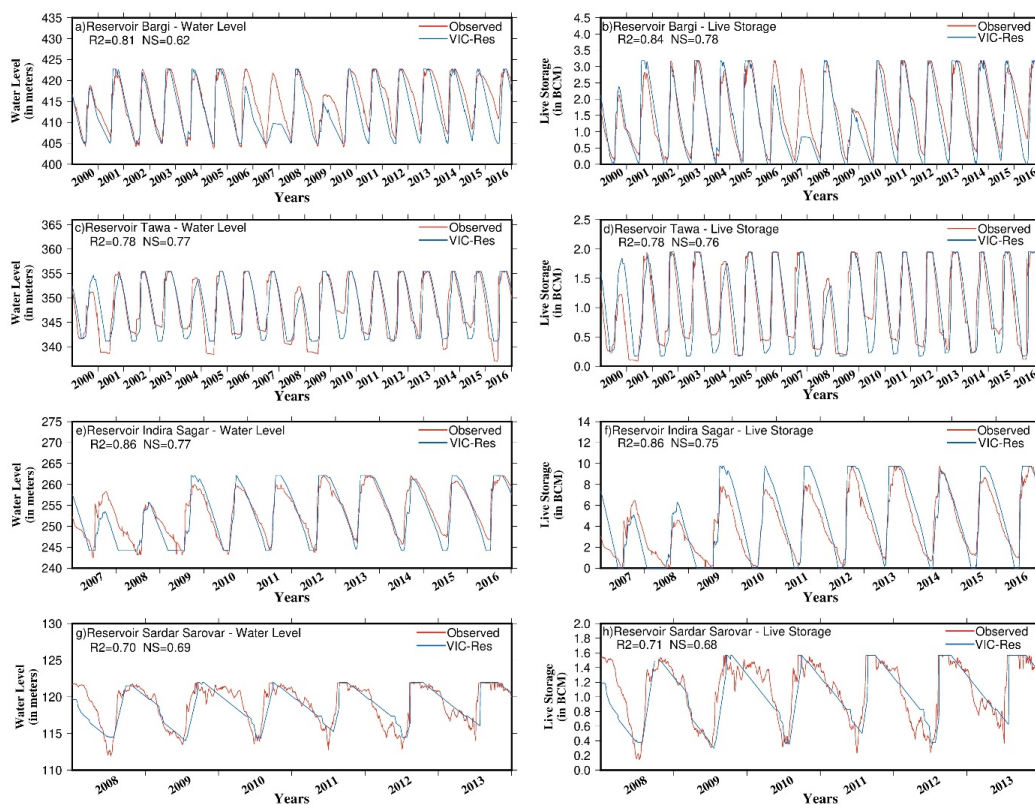
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243 **Figure 6. Calibration and evaluation of the VIC-Res model against observed daily streamflow at gauge stations at**
244 **Sandia, Handia, Mandleshwar and Garudeshwar. The performance of the VIC-Res model in simulating daily**
245 **streamflow was evaluated using the R^2 and NSE.**



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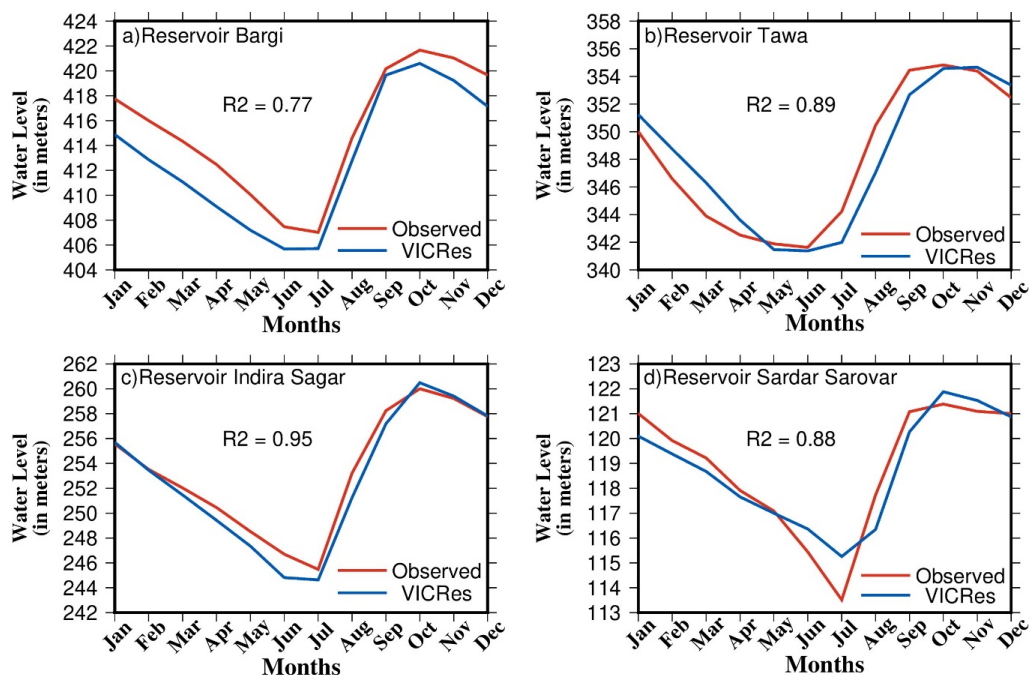
247 We considered the influence of major reservoirs on the simulated daily streamflow. Therefore, the VIC-Res
248 model's performance in simulating daily reservoir storage and the water level was evaluated against the streamflow
249 observations. We selected 2000-2016, 2000-2016, 2007-2016, and 2008-2013 as evaluation periods for Bargi,
250 Tawa, Indira Sagar, and Sardar Sarovar reservoirs, respectively, based on the availability of observations. We
251 estimated R^2 and NSE to evaluate the model's performance (Fig. 7). The model performed well in simulating all
252 the reservoirs' water levels and storage ($R^2 > 0.78$ and $NSE > 0.62$). We also compared the seasonal cycle of the
253 observed and simulated reservoir storage for all the four major reservoirs (Fig. 8). The model simulated monthly
254 seasonal cycle of reservoir storage compares well with the observed storage for all the dams with R^2 of more than
255 0.77. Overall, we find that the VIC-Res model can evaluate the ensemble streamflow forecast in the Narmada river
256 basin.



257

258 **Figure 7. Evaluation of the VIC-Res model in simulating daily water level and daily live storage at four major**
259 **reservoirs Bargi, Tawa, Indira Sagar and Sardar Sarovar.**

260



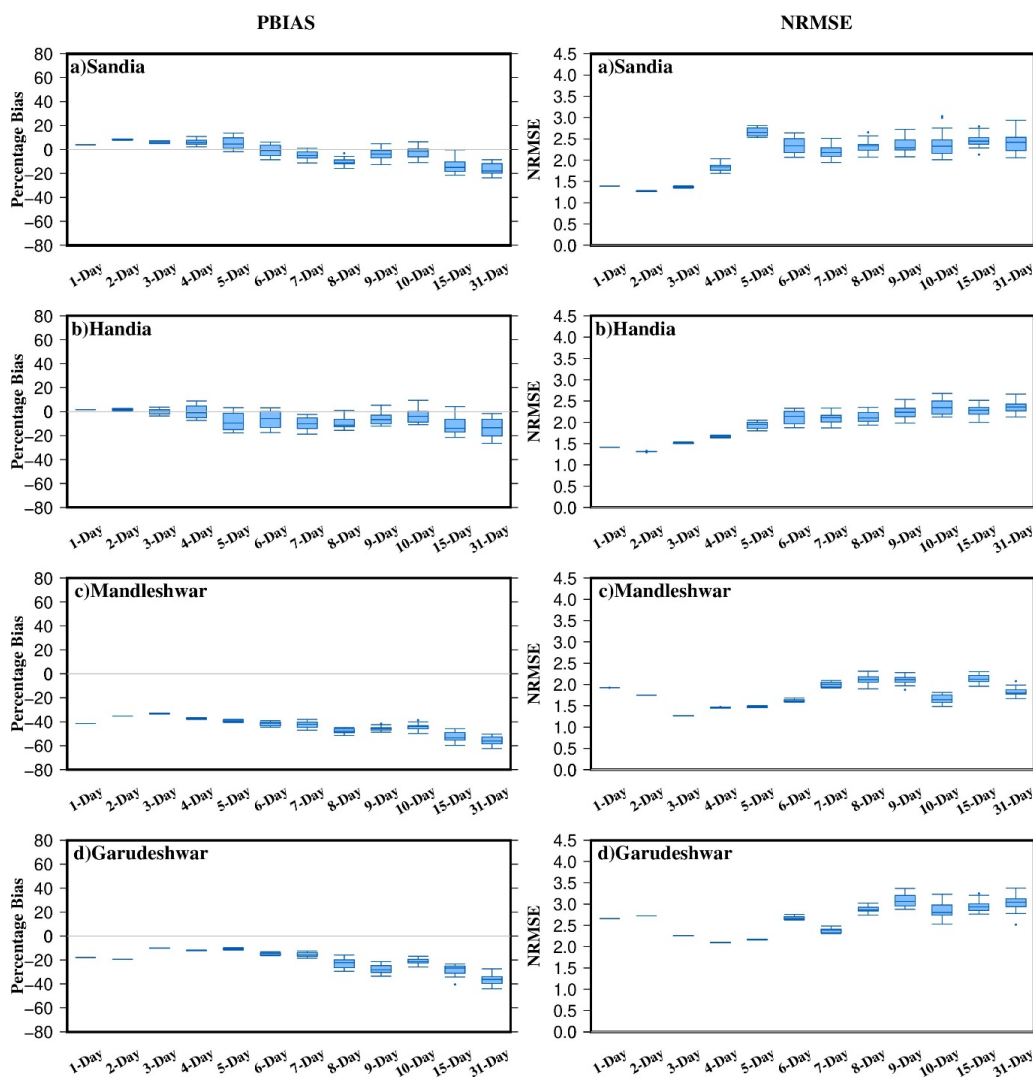
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262 **Figure 8. Comparison of observed and the VIC-Res model simulated reservoir water levels for four reservoirs in**
263 **Narmada river basin.**

264

265 3.3 Evaluation of ensemble streamflow forecast skills of ERFs

266 We estimated forecast skills of daily streamflow for 2003-2018 generated from each ensemble member of ERFs
267 for the twelve lead times (1-day to 10-day, 15-day, and 31-day). We selected a 1-10 day lead as GEFS forecast is
268 also available with the same lead. In addition, two other lead times (15 and 31 days) were selected to evaluate the
269 forecast skill of streamflow forecast from all the sixteen members of ERFs (Fig. 9). Both bias and NRMSE showed
270 a relatively lesser spread for the shorter lead (1-3 day) streamflow forecast from all the ensemble members of
271 ERFs (Fig. 9). However, uncertainty in streamflow forecast due to different ensemble members increases with the
272 lead time. NRMSE of streamflow forecast from ERFs also rises with the lead at all the stations. Ensemble
273 streamflow forecast from ERFs showed a positive bias for Sandia, Handia, and Garudeshwar, while a negative
274 bias was found for Mandleshwar station (Fig. 9).



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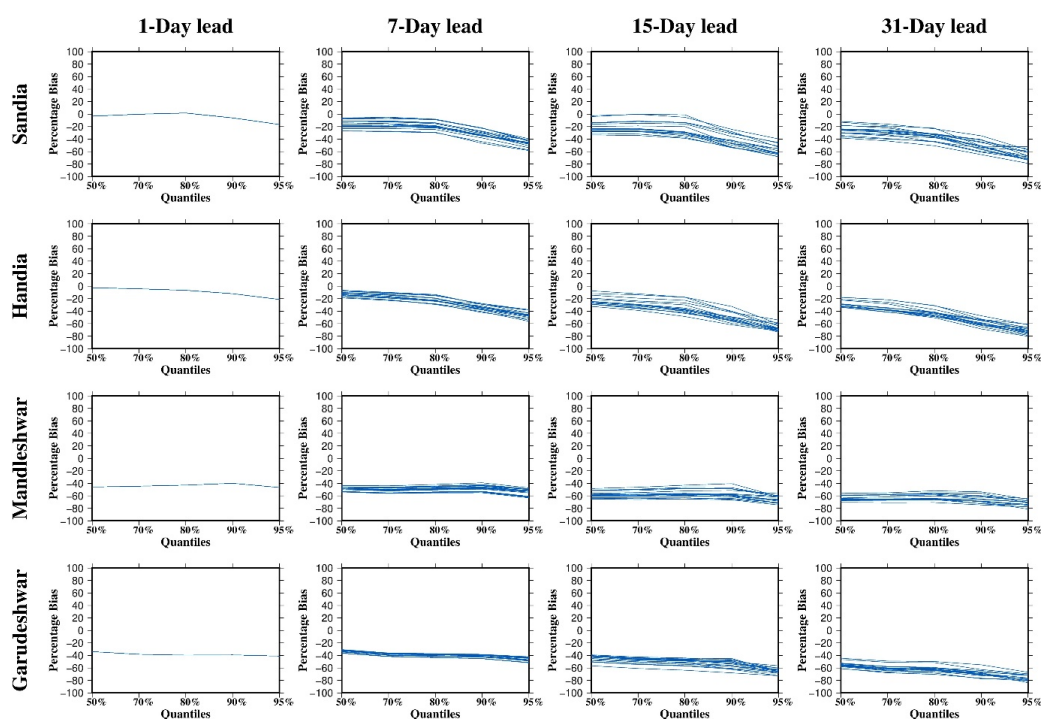
276 **Figure 9. Ensemble streamflow forecast skill based on the ERFs forecast for 2003-2018. The forecast was evaluated**
 277 **using bias (%) and NRMSE. Box and whisker plots show the skill for all 16 ensemble members at lead 1-10 day, 15**
 278 **day and 31 days at four gauge stations.**

279

280 We estimated the forecast skill in streamflow exceeding certain thresholds (50,70,80,90, and 95th percentiles) [Fig.
 281 10]. We find less spread in bias among different ensemble members for 1-day lead streamflow forecast from ERFs.
 282 However, the spread of bias in streamflow forecast due to different ensemble members increases with the lead



283 time (Fig. 10). Moreover, bias in streamflow forecast remains stable for all the selected percentile thresholds at a
284 1-day lead at all the four-gauge stations. On the other hand, bias in streamflow forecast increases for higher
285 percentiles at longer lead times. For instance, dry bias in streamflow forecast in all the ensemble members is higher
286 for the 95th percentile than for the 50th percentile. Therefore, our results show that regardless of the spread among
287 the ensemble members from ERFs, almost all the ensemble members underestimate the high flow at all the gauge
288 stations in the Narmada river basin (Fig. 10).



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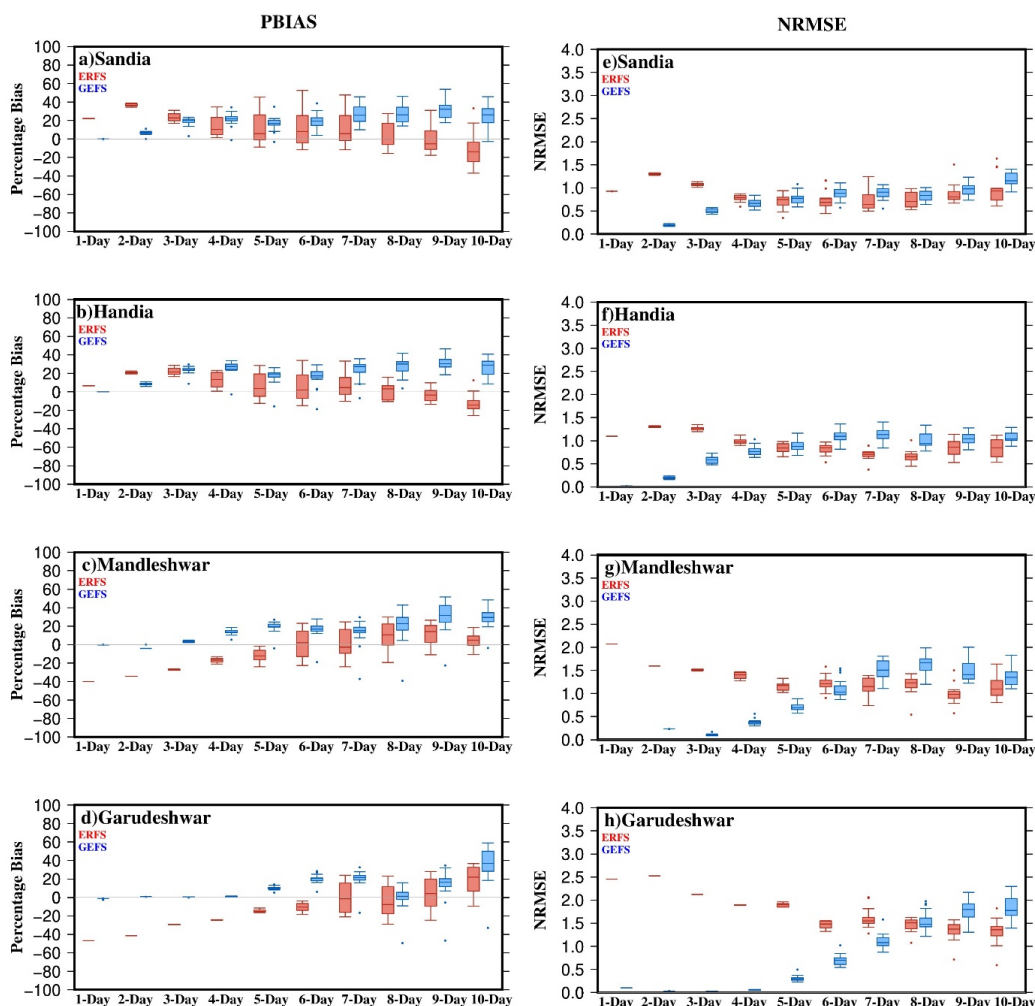
290 **Figure 10.** Bias in ensemble streamflow forecast estimated using ERFs for 2003-2018 for streamflow percentiles
291 exceeding 50th, 70th, 80th, 90th, and 95th thresholds. Bias in ensemble streamflow forecast was evaluated at 1, 7, 15, and
292 31 day lead.

293 3.4 Comparison of ensemble streamflow forecast skills ERFs and GEFS

294 We compared the streamflow forecast skills of 16 ensemble members from ERFs and 21 ensemble members from
295 GEFS. Since GEFS meteorological forecast is available only for 2019-2020, we compared the summer monsoon
296 season of these two years. ERFs forecast is available weekly for 1-32 days, while the GEFS forecast is generated
297 every day. Therefore, we compared the daily streamflow forecast from both the products for the weeks for which
298 the ERFs forecast was available for the summer monsoon of the 2019-2020 period. We compared the streamflow
299 forecast skills for all the ensemble members at 1 to 10 day leads at Sandia, Handia, Mandleshwar, and Garudeshwar



300 (Fig. 11). We find that the GEFS forecast has a better skill for the short lead time (~1-5 days) with less bias and
301 NRMSE. On the other hand, the ERFS ensemble forecast showed higher bias and NRMSE at shorter leads for
302 most of the selected locations in the Narmada basin. Streamflow forecast skill of GEFS declines rapidly after the
303 3-4 day lead time for most of the locations in the Narmada basin. The forecast products showed a larger spread
304 among the streamflow forecast ensemble members after five days lead. For short to medium range (~1 to 5 days),
305 the streamflow forecast from GEFS performed better with low NRMSE and bias for streamflow exceeding the
306 75th percentile of the summer monsoon period (Fig. S6). Moreover, streamflow forecast skill from the ERFS was
307 considerably lower than the GEFS at most of the locations for flow exceeding 75th percentiles (Fig. S6).

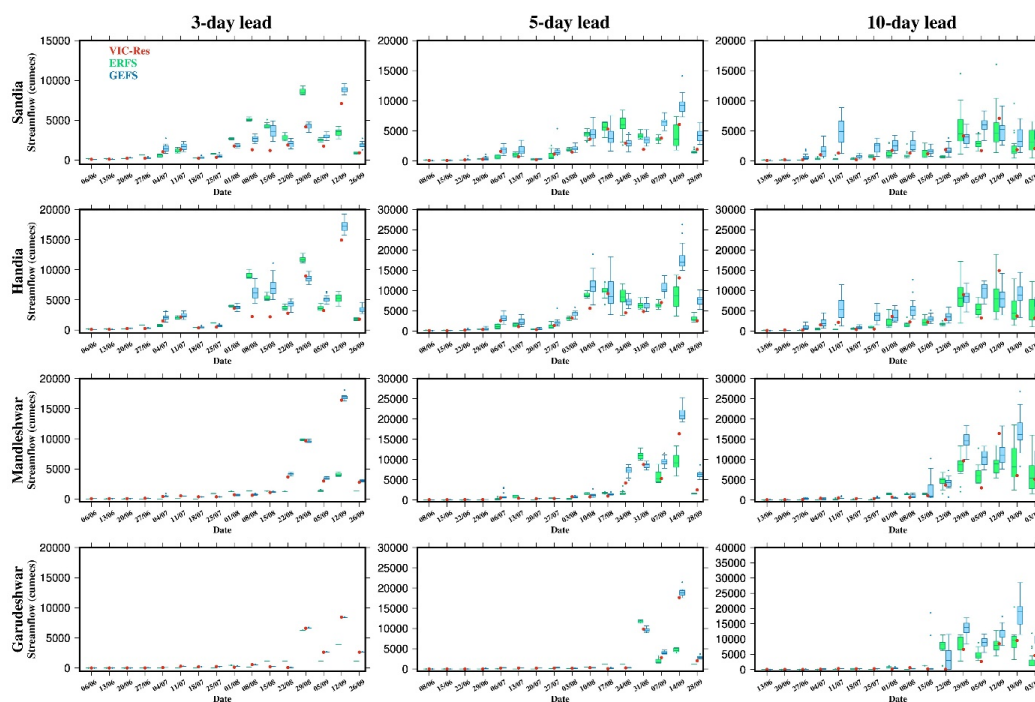


308



309 **Figure 11. Comparison of ensemble streamflow forecast skills from ERFs and GEFS for 2019-2020. The forecast skill**
310 **was evaluated considering the VIC-Res simulated streamflow with the observed forcing from IMD due to**
311 **unavailability of observed flow.**

312 We examined the daily streamflow forecast skill at 3-day, 5-day, and 10-leads from ERFs and GEFS forecasts for
313 the summer monsoon season of 2019 & 2020 against VIC-Res simulated streamflow using the observed
314 meteorological forcing at all the four gauge stations (Fig. 12 and Fig. S7). Since observed daily streamflow was
315 unavailable for skill assessment, the comparison was made against the VIC model simulated flow with the
316 observed meteorological forcing (Fig. 12 and Fig. S7). The GEFS forecast successfully captured streamflow peaks
317 in both 2019 and 2020 at a 3-day lead. In 2019, GEFS forecasts overestimated streamflow peaks at 3-day and 5-
318 day leads during the summer monsoon. On the other hand, the ensemble streamflow forecast developed using the
319 ERFs meteorological forecast showed a higher spread than GEFS (Fig. 12, Fig. S7). The spread in ensemble
320 streamflow forecast increases for both ERFs and GEFS forecast at a 10-day lead. However, the ERFs's streamflow
321 forecast showed a better skill at the 10-day lead. Despite having fewer ensemble members than the GEFS, the
322 ERFs forecast showed a broader spread in streamflow prediction, highlighting a higher uncertainty in prediction.
323 We find that GEFS overestimate streamflow the ERFs underestimates most of the locations and lead times.



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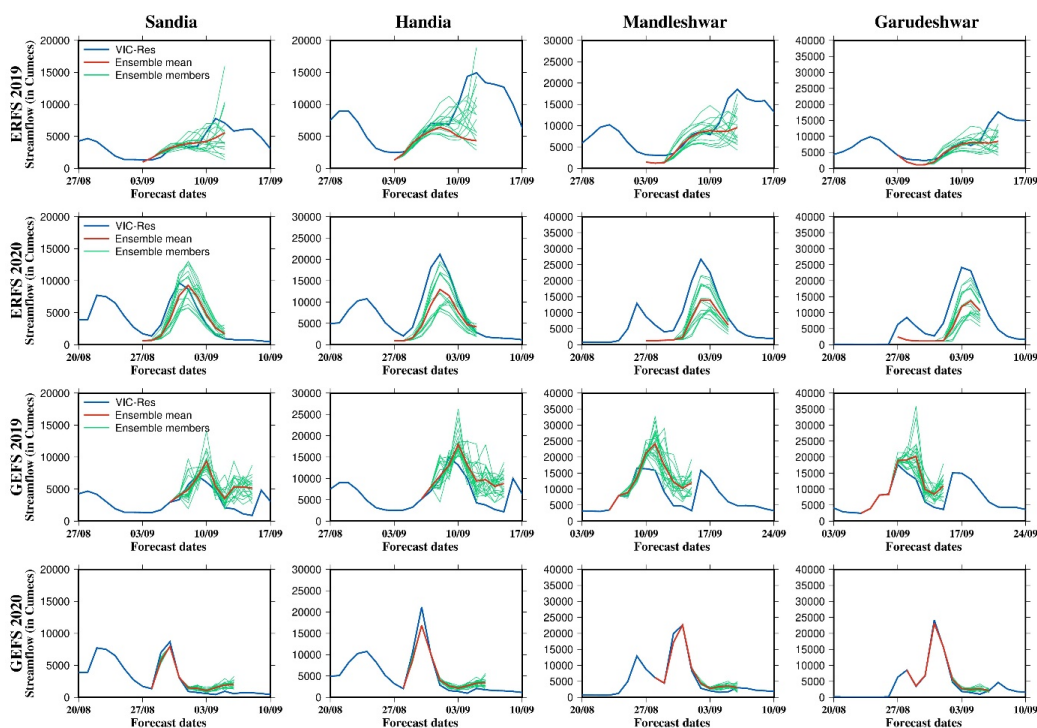
325



326 **Figure 12. Comparison of ensemble streamflow simulated using the VIC-Res model with ERFs and GEFS forecast**
327 **products during the summer monsoon of 2019. The forecast skill was evaluated considering the VIC-Res simulated**
328 **streamflow with the observed forcing from IMD due to unavailability of observed flow.**

329

330 We examined the streamflow forecast generated by all the ensemble members of ERFs and GEFS for a few events
331 using the VIC-Res model (Fig. 13). The ensemble streamflow prediction was compared considering the model
332 simulated streamflow with the observed forcing from IMD. In 2019, the ensemble mean streamflow from all the
333 ensemble members of ERFs considerably underestimated the peak flow (Fig. 13). However, a few ensemble
334 members of the ERFs forecast captured the peak flow at the four locations of the Narmada river basin (Fig. 13).
335 At Handia station, 1 out of 16 ensemble members exceeds the observed streamflow. Moreover, GEFS forecasts at
336 short leads (3-5 days) performed well in capturing peaks (Fig. 13). However, GEFS forecasts showed a smaller
337 spread in ensemble streamflow at the short lead time (1-5 days). Overall, we find that ensemble forecasts can be
338 used for probabilistic streamflow prediction.



339

340 **Figure 13. Ensemble streamflow simulations using the ERFs forecast at 5-11 day lead and GEFS forecast at 3-5 day**
341 **lead against the VIC-Res simulated streamflow with the observed meteorological forcing for 2019 and 2020.**



342

343 **4 Discussion and conclusions**

344 Streamflow forecast plays an essential role in efficient reservoir operations and flood mitigation (Chen et al., 2016;
345 Mediero et al., 2007). A reliable streamflow forecast can reduce uncertainty in reservoir operations and enhance
346 the development of a flood early warning system. Notwithstanding the considerable progress in an operational
347 meteorological forecast from different agencies, efforts to establish an ensemble streamflow forecast system at
348 river basin scales have been limited. Moreover, it remains unclear if other meteorological forecast products have
349 different streamflow forecast skills. We used the two meteorological ensemble forecast products available from
350 IMD to examine streamflow forecast skills in the Narmada river basin. We used the ensemble streamflow
351 prediction approach to generate the daily streamflow simulations considering the influence of reservoirs in the
352 Narmada river basin. We compared the performance of ERFs and GEFS ensembles for the summer monsoon
353 period of 2019-20. We also assessed the skills of the ERFs dataset solely for a more extended period from 2003
354 to 2018.

355 The ERFs ensemble forecast is available once a week at 1-32 days lead time. On the other hand, GEFS ensemble
356 forecasts are available daily at 1-10 days lead for the summer monsoon period of 2019-2020. Hagedorn et al.
357 (2005) reported that bias-correction of the raw forecast does not necessarily increase the forecast skill. Moreover,
358 statistical correction of the raw forecast is inappropriate, which can lose its effect propagating through the
359 hydrologic model (Zalachori et al., 2012; Crochemore et al., 2016; Benninga et al., 2017; Hagedorn et al., 2005).
360 Therefore, we did not bias-correct the raw meteorological ensemble forecasts from ERFs and GEFS. The skills of
361 ERFs and GEFS precipitation and temperature (minimum and maximum) forecasts were estimated for 1-, 5- and
362 10-day lead. The GEFS raw forecast showed better skills than the ERFs forecast for mean and extreme
363 precipitation. As precipitation plays a vital role in streamflow forecast (Meaurio et al., 2017; Demargne et al.,
364 2014; Pappenberger et al., 2005), our results showed that GEFS forecast provides better skills for streamflow
365 prediction in the Narmada River basin. The post-processing of streamflow data can significantly improve
366 performance (Tiwari et al., 2021; Muhammad et al., 2018), which can be used in the future to examine the
367 improvements in streamflow prediction. Moreover, a multi-model approach can be used to reduce the errors and
368 uncertainty in streamflow forecasts that could arise due to the parameterization of hydrological models (Velázquez
369 et al., 2011; Zarzar et al., 2018; Muhammad et al., 2018).

370

371 The skills of ERFs and GEFS ensemble forecasts were estimated for 1, 5 and 10-day leads. GEFS raw forecasts
372 illustrated better skills than ERFs forecasts for overall rainfall and extreme precipitation. As studies show that rain
373 plays a vital role in streamflow forecast (Demargne et al., 2014; Meaurio et al., 2017; Pappenberger et al., 2005),
374 we also observed the same results. The ensemble forecast with better skills performed well in predicting daily
375 streamflow. Correcting the bias of the input forecast may shrink the variability range of the result. However,



376 ensemble forecasts aim to capture uncertainties. Studies suggest that the post-processing of streamflow data can
377 significantly improve performance (Muhammad et al., 2018; Tiwari et al., 2021). A multi-model approach, where
378 more than one hydrologic model is used, can generalize the uncertainty introduced by the hydrologic model.
379 Various studies have reported improved forecast skills using the multi-model approach (Muhammad et al., 2018;
380 Velázquez et al., 2011; Zarzar et al., 2018).

381 Flood forecasting using the available meteorological forecast products can help in mitigating the losses through
382 early warnings. To account for the uncertainty arising from initial state and model parameterization, the individual
383 members of the ensemble weather forecast can provide better information than their ensemble mean (Saleh et al.,
384 2019). The probabilistic approach over the deterministic method provides the range of variability, which can help
385 determine the probability of exceeding a specific threshold of streamflow (Hsiao et al., 2013). The shift from the
386 existing 'flood forecast system' to the 'ensemble-based probabilistic forecast' requires modifications in the current
387 flood forecast practice. The transition is expected to change various aspects of the existing decision-making
388 process. The forecasters need to train the on-duty officers adequately and the authorities on probabilistic forecasts.
389 We evaluated the streamflow forecast skills at 1-32 day lead in the Narmada river basin. The increased lead time
390 in streamflow forecast can assist in developing efficient communication methods of information (Arnal et al.,
391 2020; Ramos et al., 2010). Moreover, ensemble streamflow forecast at longer leads can be effectively used in
392 optimizing reservoir operations (Alemu et al., 2011). Our results show that, while the mean of the ensemble
393 members failed to capture the high flows, a few individual ensemble members performed better in capturing peak
394 flow, which can be used to develop probabilistic early warnings.

395 Based on our findings, the following conclusions can be made:

- 396 1) The raw precipitation forecast from both GEFS and ERFS datasets showed moderate skills (bias, NRMSE
397 and correlation) against observations from IMD at 1-day, 5-day and 10-day lead times. While both (ERFS
398 and GEFS) forecast products underestimated extreme precipitation, dry bias in the ERFS forecast was
399 more prominent than the GEFS forecast. For instance, raw precipitation forecast from ERFS showed
400 negative bias across the Narmada river basin. On the other hand, the raw precipitation forecast from GEFS
401 exhibited both negative and positive bias. Both the forecast products showed better skills for maximum
402 and minimum temperatures than precipitation.
- 403 2) We calibrated and evaluated the VIC-Res model to simulate streamflow, considering the influence of
404 reservoirs at four gauge stations in the Narmada River Basin. The model reproduced daily streamflow,
405 reservoir water level, and storage reasonably well against the observations.
- 406 3) Comparing the streamflow forecast skills of both the ensemble forecasts showed that GEFS forecasts
407 performed better than the ERFS at all the locations in the basin. However, both the forecast products
408 underestimated the extremes, which can be due to dry bias in extreme precipitation. The spread in
409 streamflow due to different ensemble members increased with the forecast lead time. Overall, an
410 ensemble forecast can be used to develop a probabilistic forecast based flood early warning system.



- 411 **Data availability:** All the datasets used in this study can be obtained from the corresponding author.
412
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