A gridded multi-site precipitation generator for complex terrain: An evaluation in the Austrian Alps

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Abstract.

For climate change impact assessment, many applications require very high-resolution data of precipitation consistent both in space and time for current or future climate. In this regard, stochastic weather generators are designed as a statistical down-scaling tool that can provide such data. Here, we adopt the framework of a precipitation generator of Kleiber et al. (2012), which is based on latent and transformed Gaussian processes, and propose an extension to that for a mountainous region with complex topography by allowing elevation dependence in the model. The model is used to generate two-dimensional fields of precipitation with 1 km spatial and daily temporal resolution in a small region with highly complex terrain in the Austrian Alps. This study aims at evaluating the model for its ability to simulate realistic precipitation fields over the region using historical observations from a network of 29 meteorological stations as an input, discusses its added value over the original set-up and its limitations. Results show that the model generates realistic fields of precipitation with good spatial and temporal variability. The model is able to generate some of the difficult areal statistics useful for impact assessment such as areal dry and wet spells of different lengths and areal monthly mean of precipitation with great accuracy. The model also captures the inter-seasonal and intra-seasonal variability very well while the inter-annual variability is well captured in summer but largely underestimated in autumn and winter. The proposed model adds substantial value over the original modeling framework, specifically for the precipitation amount. The model is not able to reproduce realistic spatio-temporal characteristics of precipitation in autumn. We conclude that with further development, the model is a promising tool for downscaling precipitation in complex terrain for a wide range of applications in impact assessment studies.

1 Introduction

Precipitation is a major component of the hydrological cycle. With global warming, the hydrological cycle is expected to intensify and the risk associated with extreme events will increase (Tabari (2020); Pfahl et al. (2017)). The resulting changes in the precipitation will be unequally distributed around the world. There are many hydrologic responses to climate change and the potential impacts of these are likely to affect availability of fresh water, agriculture, timing and severity of wildfires and habitat sustainability (Bates et al. (2008); Kundzewicz et al. (2008)). With the increasing awareness about climate change and its global impact on the ecosystem and human societies (Konapala et al. (2020); Haddeland et al. (2014); Schewe et al. (2014)),

there is also an increasing need to understand the effects and impacts that would occur at local scale. Knowledge on how the local hydrological cycle and water resources will be affected by climate change is essential for planning reliable adaptation strategies and water policy.

30 In Austria, where a large part of the country is covered by mountains, the local hydrological cycle depends widely on temporal and spatial variations in precipitation. Tourism and agriculture are among the main drivers of Austria's economy. Accessibility of water resources for human consumption and ecosystems largely depends on the spatio-temporal distribution of precipitation. In the Austrian Alps, studies on the observed as well as projected impact of climate change show the changes on the availability of snow cover as well as water flux (e.g. Abermann et al. (2009); Wijngaard et al. (2016)). This ultimately will have impact on

35 the economy, the ecosystem, the environment and the society. To assess the impacts of climate change at local scales, precise climate information is critical which can serve the needs of the decision makers. Often such information should be consistent in space and time for present as well as future climate. Many applications in hydrology require very high resolution data of precipitation, typically of 1 km for spatial and daily for temporal scale. However, obtaining such high-resolution precipitation data is still a challenging task especially in the mountainous region (Henn et al. (2018)). Most importantly for the complex

40 topography such as the Austrian Alps, even 1 km resolution cannot include the impact of topography on climate correctly. For such regions, many applications in hydrology and ecology require even higher resolution data – in the spatial scale of 100 m and hourly for temporal scale. Climate models with higher resolutions like Regional Climate Models are also unable to provide such data. For that reason, various downscaling methods have been in use in the past few decades. Among all the downscaling methods, statistical downscaling using stochastic weather generators (WGs) has become very popular mainly because WGs

45 are computationally parsimonious.

A vast variety of WGs have been developed based on different approaches. Most widely used WGs are based on a rather simplistic approach in which the sites are mutually independent in space and time. Such WGs are generally referred to as single-site WGs. Among the single-site WGs, the most popular WGs are the parametric models based on Richardson (1981),

50 who used a Markov chain to simulate time series of precipitation occurrence (wet/dry days) and amount and other variables were generated upon the condition of whether the generated day is wet or dry (e.g. Dabhi et al. (2021), Caron et al. (2008); Zhang et al. (2004); Dubrovský et al. (2004); Wilks (1992)). The details on the available WGs can be found in the review articles by Ailliot et al. (2015); Maraun et al. (2010); Wilks and Wilby (1999) where Maraun et al. (2010) focused solely on precipitation.

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The major drawback of single-site WGs is that they are focused on a single location only, which can generate realistic data at a location but lacks the spatially correlated structure in the generated data. Getting a spatially and temporally consistent dataset – which is more realistic – from single-site models is impossible. For that reason, in the past two decades the focus has moved towards development of spatio-temporal WGs also known as multi-site WGs. As for precipitation, given the uneven nature of

60 its occurrence and intensity, it becomes more challenging to model it by keeping the spatio-temporal structure. In particular, in

complex topography like in the Alps, this task is even more challenging.

Numerous approaches have been proposed to generate spatially and temporally correlated precipitation data. Wilks (1998) did one of the early works on multi-site generation of daily precipitation where single-site parametric WGs at sites were forced with correlated random numbers to generate the occurrence of precipitation and the amount of precipitation was generated using a mixture of two exponential distributions. Other approaches to spatio-temporal modeling of precipitation are hidden Markov models (e.g. Ghamghami et al. (2016); Charles et al. (1999); Ailliot et al. (2009)), copula-based approaches (Serinaldi (2009); Bárdossy and Pegram (2009)), resampling based on k-nearest neighbours (Apipattanavis et al. (2007); Buishand and Brandsma (2001); Rajagopalan and Lall (1999)), Poisson cluster models (Ramesh et al. (2012); Cowpertwait (1995)), artificial neural networks (Harpham and Wilby (2005)) and Generalised Linear Model (GLM) based approaches (Kleiber et al. (2012); Verdin et al. (2018)). Baxevani and Lennartsson (2015) proposed a spatio-temporal model using a censored latent Gaussian field for precipitation generation. Olson and Kleiber (2017) used the Approximate Bayesian Computation method. Gao et al. (2021) developed a multi-site stochastic daily rainfall model by coupling a univariate Markov chain with a multi-site rainfall event model. There also exist more sophisticated WGs that can provide high-resolution spatio-temporal fields combining physical and stochastic approaches (e.g. Peleg et al. (2017); Paschalis et al. (2013)).

However, most of the aforementioned approaches simulate precipitation only at the locations where observations are available and such multi-site WGs have been implemented for the Alps. For example, Keller et al. (2015) and Keller et al. (2017) used a Wilks-type WG for precipitation simulation and downscaling, respectively, for a mountainous catchment in the Swiss Alps. Breinl et al. (2013) used a semi-parametric multi-site precipitation generator for the mountains in the Austrian-German Alps. However, high resolution data in space and time are needed to provide more realistic input for local climate impact assessment. To achieve that, a gridded multi-site model is required. Sparks et al. (2017) proposed a multi-site multivariate WG based on the use of periodically extended empirical orthogonal functions (EOFs), in which they modeled precipitation as a censored latent Gaussian process. They generated gridded data of precipitation and temperature over Europe using gridded input data, but their WG can not provide gridded data without gridded observations. Peleg et al. (2017) developed a WG called AWE-GEN-2d and used it in the Swiss Alps, which can generate two-dimensional fields of various meteorological variables where precipitation is generated at 2 km spatial and 5 min temporal resolution. Although their approach is sophisticated as it is a hybrid approach combining dynamical and statistical approaches, it requires spatially distributed data for calibration of the WG and it cannot generate data on a region outside the calibration region. Such WGs are of limited use if the observed gridded data are not available which is often the case. To our knowledge, not much work has been done for complex terrain like the European Alps using multi-site gridded WG without gridded observations.

Wilks (2009) developed one of the first ever WGs which can provide gridded data of precipitation and also temperature at locations with no observations. Kleiber et al. (2012) also gave an approach using the GLM based model which uses Gaussian processes to generate gridded data. Their approach is appealing as it generates the readily available field of precipitation using

kriging for interpolation of the model parameters. The advantage of their approach is that in the GLM framework, one could include various covariates such as large-scale climate indices, local climate information, topographical information etc. which makes the model more flexible. Another advantage is that it is a probabilistic approach which allows one to quantify the uncertainties in the parameter estimation. However, Kleiber et al. (2012) tested the model only for the multi-site precipitation generation, i.e. at locations with observation and not for the generated gridded data of precipitation. Since many applications for impact studies require gridded data as input, before producing the gridded data for such applications, one must evaluate the model for its ability to reproduce gridded fields. Another important point to be noted is, since the model can provide data at locations without historical observations, one can obtain historical time series of daily precipitation at those locations. In order to use the model for this purpose, it is necessary to assess the model performance for the gridded fields. Verdin et al. (2018) modified the framework of Kleiber et al. (2012) by including seasonal precipitation as an additional covariate and evaluated the model for gridded data, but it was implemented in flat terrain. Also, their modified model and the original model, both used an isotropic and stationary covariance structure with ordinary kriging (OK) for the interpolation of the model parameters which may not be suitable for the complex topographical terrain. Bennett et al. (2018) generated precipitation fields using a latent-variable approach that provides a parsimonious method to jointly generate rainfall occurrence and amount. They used an isotropic powered-exponential function for including spatial correlations and kriging for the interpolation of the parameters. However, they also implemented their model in relatively flat terrain in South Australia. To our knowledge, no space-time gridded precipitation generator has been evaluated for its ability to generate two-dimensional fields of precipitation in the highly complex terrain without needing gridded input data.

Here, we propose an extension to the framework of Kleiber et al. (2012) for complex terrain and evaluate the model for its capabilities to generate realistic two-dimensional fields of precipitation for a mountainous region in the Austrian Alps. In addition, we examine the added value of our model to the original isotropic set-up and discuss the limitations of the model.

This article is organized in a following way: Section 2 describes the extension to the isotropic framework for the implementation in a mountainous region. Section 3 details the study area, the data and the model evaluation strategy used in the study. Section 4 presents and analyses the results, Section 5 comprises a discussion of the results and Section 6 summarizes the study.

2 Model Description

2.1 Precipitation Occurrence

At a location s and on a day t, precipitation occurrence $O(s,t)$ is 0 (dry day) if no precipitation and 1 (wet day) if precipitation occurs. A wet day is defined when the precipitation amount exceeds 0.1 mm.

For a set of locations $\mathbf{s}$ and on a day t, a latent Gaussian process $W_O(\mathbf{s},t)$ is defined with mean function $\mu_O(\mathbf{s},t)$ and covariance function $C_O(\mathbf{h},\mathbf{v},t)$, where $h = |s_i - s_j|$ is the horizontal (Euclidean) distance between two locations denoted by i and j

and $v = |v_i - v_j|$ is the elevation difference between the two locations. The suffix 'O' stands for occurrence. Since we are implementing the model in a mountainous region, we define the covariances among the sites as a function of the difference in the elevation, too. In comparison with the original model where an isotropic covariance structure is used as a function of horizontal distances among the sites, this allows us to include anisotropy in the model. The precipitation occurrence then is defined as

$$O(\mathbf{s},t) = 0 \qquad if \qquad W_O(\mathbf{s},t) < 0$$
$$O(\mathbf{s},t) = 1 \qquad if \qquad W_O(\mathbf{s},t) \geq 0 \tag{1}$$

where the mean function is

$$\mu_O(\mathbf{s},t) = \beta_O(\mathbf{s})' X_O(\mathbf{s},t) \tag{2}$$

'X_O' is a vector of covariates and 'β_O' is a vector of regression parameters as in Eq. 5.

Kleiber et al. (2012) used a stationary and isotropic exponential covariance function of the form

$$C(\mathbf{h},t) = exp\left(\frac{-|\mathbf{h}|}{A(t)}\right) \tag{3}$$

where $A(t)$ is the time dependent scale parameter.

Since our goal is to use the model in complex topography, we introduce anisotropy in the covariance function (Eq. 3) by taking the difference in elevation between two locations. Thus, our stationary and anisotropic covariance function $C(\mathbf{h},\mathbf{v},t)$ takes the following form

$$C_O(\mathbf{h},\mathbf{v},t) = exp\left(-\frac{|\mathbf{h}|}{A(t)} - \frac{|\mathbf{v}|}{B(t)}\right) \tag{4}$$

where $A(t)$ and $B(t)$ are the time dependent scale parameters in the horizontal and vertical direction, respectively.

Elevation dependence in the covariance structure is the natural assumption in complex terrain. In the literature, it has been used for precipitation simulation in the mountains, e.g. Wilks (1999) and Wilks (2009).

At the base of this model is the single-site precipitation generator based on the GLM framework (e.g. Stern and Coe (1984); Chandler and Wheater (2002); Furrer and Katz (2007)) which is similar to a Richardson-type precipitation generator (Richardson (1981)), where daily precipitation occurrence is modeled as a first-order Markov chain and daily precipitation amount is

modeled using a gamma distribution. The GLM-based approach allows more flexibility as one may include as many covariates as desirable, through which the seasonality or the influence of large scale circulation on the local precipitation can be included. In the GLM approach, taking the previous day's occurrence as a covariate, forms a first-order Markov chain. Thus, at individual sites, the model reduces to a logit model which is given by

$$log\left(\frac{p_t}{1-p_t}\right) = \beta'_O(\mathbf{s})X_O(\mathbf{s},t) \tag{5}$$

where, p_t is the probability of occurrence on a day t.

Note that we use a logit link function instead of a probit link function, as was the choice in the original model. The parameters β_O are estimated at each location using the maximum likelihood approach.

$C_O(\mathbf{h},\mathbf{v},t)$ is estimated as the covariance matrix of the residuals in the logit model. We use the method of moments approach as suggested by Kleiber et al. (2012) to estimate the parameters of the covariance function. The parameters are estimated separately for each month to allow for seasonality in the generated data. The nugget term (in the kriging, see below) is considered zero.

Note that the Gaussian process modelling is considered a non-parametric method. In a parametric model, the number of parameters remains fixed with respect to the amount of the data available (i.e. number of stations in our case), while with non-parametric methods, the number of parameters grows with the number of data points.

The Gaussian process allows for a spatial interpolation method called 'kriging' which allows one to interpolate the model parameters β_O associated with each covariate, which are estimated at observation locations, to any location of interest. These interpolated coefficients are then used to obtain the mean function (Eq. 2). Here, we use kriging with external drift (KED) to interpolate the regression parameters. Since precipitation in the mountains is unequally distributed across the terrain, we allow elevation as an external drift in kriging such that the predicted values of precipitation (through the interpolated parameters) reflect the elevation dependence of precipitation. Again, inclusion of the elevation in kriging interpolation is natural in complex terrain. It has been used in the literature for precipitation interpolation in the mountains and proven to outperform OK, e.g. Tobin et al. (2011) and Rata et al. (2020). Also, we have found linear dependence in the model parameters with elevation (plots not shown). In KED, an auxiliary variable is assumed, which is elevation here, that is linearly related to the variable of interest which is the β parameter associated with each covariate in the model.

We also compare the results of our model with a simulation using ordinary kriging (OK) instead of KED in our model and also with the original isotropic model using OK and KED. This will be discussed in the Section 4.3.

2.2 Precipitation Amount

To simulate spatially correlated fields of precipitation, another Gaussian process $W_A(\mathbf{s}, t)$ is defined with mean function $\mu_A(\mathbf{s}, t)$ and covariance function $C_A(\mathbf{h}, \mathbf{v}, t)$, such that

$$Y(\mathbf{s}, t) = G_{\mathbf{s}, t}^{-1}(\Phi(W_A(\mathbf{s}, t)) \tag{6}$$

where $G_{\mathbf{s}, t}$ is the cumulative distribution function (CDF) of the gamma distribution at the locations $\mathbf{s}$ and time t, and Φ is the CDF of a standard normal distribution. The suffix 'A' stands for amount.

At an individual location, the amount model is the gamma GLM with logarithmic link function as given by Furrer and Katz (2007). The shape parameter α of the gamma distribution varies with space but not with time, while the scale parameter γ varies with both space and time. This way each location has its own distinct value of shape and scale parameters, with the scale parameter varying with time. Thus, we have

$$log(\gamma(\mathbf{s}, t)\alpha(\mathbf{s})) = \beta_A'(\mathbf{s})X_A(\mathbf{s}, t) \tag{7}$$

with the mean of the gamma distribution being the product of the scale and shape parameters, i.e. $\gamma\alpha$. X_A is the vector of covariates possibly different from those selected in the occurrence model.

The scale and shape parameters, γ and α respectively, and the model parameters β_A parameters are estimated at each individual observation site using the maximum likelihood approach and then are interpolated using KED. We allow the scale parameter to vary with every month to include seasonal variations in precipitation at each location.

The mean function $\mu_A(\mathbf{s}, t)$ of the Gaussian process $W_A(s, t)$ is again obtained from a regression on covariates. The covariance function $C_A(\mathbf{h}, \mathbf{v}, t)$ is the same as given in the occurrence model (Eq. 4), but with different parameters. The parameters of the covariance function are estimated for each month separately using the method of moments approach to allow the seasonality in spatio-temporal pattern of precipitation amount.

3 Implementation

3.1 Study area and data

The model is implemented in a small region comprising highly complex terrain (ranging from 256 m a.s.l. to over 3500 m a.s.l.) in the Austrian Alps. The area is surrounding the catchment of River Oetz, mainly in the federal state of Tyrol but also includes a part of the Autonomous Province of South Tyrol in northern Italy. The reason for selecting this region is that the catchment of river Oetz is a widely researched area (e.g. Wijngaard et al. (2016); Abermann et al. (2009)). To include more stations in the study, we allow stations from the surrounding region, also including northern Italy. The study region is comprised of several
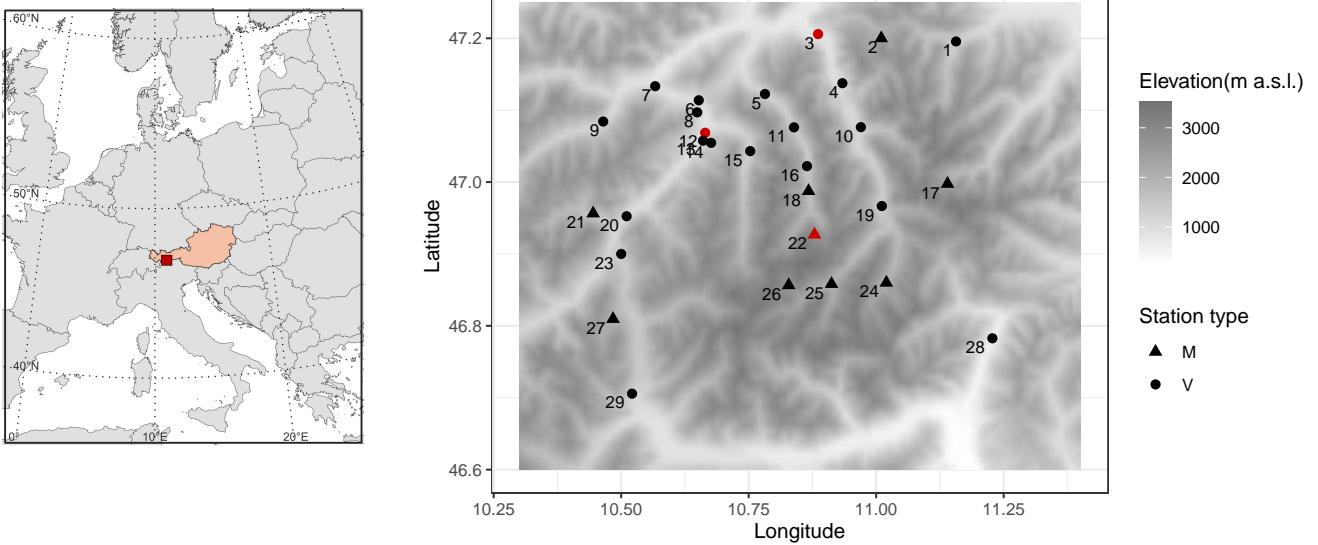
Figure 1. Study area showing the locations of the 29 meteorological observation stations (right) whose data are used in the study and the location of the region in the central Alps (left). Longitude and latitude are given in degrees east and north, respectively. Gray shading denotes the elevation (m a.s.l.). Stations with elevation higher than 1500 m a.s.l., usually high-mountain stations, are shown by the symbol type M and the stations with elevation less than 1500 m a.s.l., typically the valley stations, are shown by symbol type V. The stations shown in red colour are selected as example stations to illustrate the results in the article.

220   valleys including Oetz and Pitz Valley in Austria and Passeier Valley in South Tyrol. Daily observations from 29 meteorological stations (Figure 1) based on the availability of homogeneous time series of a period of 30 years from 1981 to 2010 are selected. Typically, for hydrological applications the input data required are hourly on a temporal scale for the spatial scale of terrain considered, but due to very few hourly data sets available over climate time scales, we selected daily data for the study. The dataset comprises data provided by the Austrian National Weather Service (ZAMG – Zentralanstalt für Meteorologie und

225   Geodynamik), the Hydrographic Service of Austria, the Hydrographic Service of the Autonomous Province of South Tyrol, the Institute of Atmospheric and Cryospheric Sciences – University of Innsbruck and TIWAG (Tiroler Wasserkraft AG). The highest station is on a glacier (Hintereisferner) at the elevation of 2860 m a.s.l., while the lowest is at 588 m a.s.l in northern Italy. A few stations have missing values for single days or a short period. Our model ignores such values while computing the observed statistics. The data at all the stations are thoroughly quality controlled by the respective service providers.

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In the northern part of the region, we have a dense network of stations, while the southern part has relatively fewer stations. The average inter-station distance between two locations is 28.15 km. The maximum inter-station distance is 72.84 km and the minimum inter-station distance is 1.25 km. The average altitude difference between two stations is 0.605 km, while the maximum altitude difference is 2.272 km. The locations of the 29 stations are shown in Figure 1 and further details about the

stations are given in Table 1.

The mean annual precipitation observed in the lowlands is approximately 780 mm on an average of 150 wet days in a year, while the highest mean annual precipitation is observed at the high-mountain station Dresdner Huette which is 1320 mm with an average of 176 wet days in a year. The highest number of mean annual wet days is 220 at St. Martin in Passeier Valley in South Tyrol with 887 mm mean annual precipitation.

Due to strongly different topography, a large variability both in space and in time is observed in the dataset. Out of all the 29 stations, Prutz has the most distinct climatological characteristics. For example, Prutz has the largest variability in almost all the months. Also, the most extreme precipitation (156.5 mm) in a day is recorded at Prutz (in July 2009) whereas amongst the remaining 28 stations, the second highest amount of precipitation recorded on the same day was at Dresdner Huette (35.1 mm). Apart from Prutz, only Dresdner Huette recorded a daily precipitation amount as high as 120.4 mm in the 30 years of record. At the location St. Leonhard im Pitztal, there are two stations operated by two different service providers. One is by the hydrographic service of Austria (St. Leonhard im Pitztal-1, see Table 1) in the northern part of the Valley and the other by ZAMG (St. Leonhard im Pitztal-2, see Table 1) in the southern part of the Pitz Valley. St. Leonhard im Pitztal-2 has somewhat different climatological characteristics than the nearby stations. Another station is St. Martin which has quite different climatologies compared to the Austrian stations. Note that this station is in the south of the Alpine crest (i.e. in northern Italy) and has the lowest elevation in the observed data. Thus, there are high variations in the observed climatologies of precipitation from valley to valley and also for stations within the same valley. This adds a particular challenge to the simulation.

To reduce the sampling uncertainty and increase the robustness of the observations, we increase the sample size of the observed data by considering a 7-days window centred at the day of interest. Thus, the chance that a particular date had, e.g., 30 dry days by random choice, is minimized – what avoids a probability of dry day being 1.0 (rather than 0.98, say), which is a problematic model setting. We generate N=30 stochastic realisations, each 30 years long (30 realisations×30 years = 900 years), of daily two-dimensional fields of precipitation on a 1 km grid over the region using the aforementioned observed daily data of 30 years. The Shuttle Radar Topography Mission (SRTM) 1 km (30 arc second) resolution dataset (Becker et al. (2009)) is used as a simulation grid. We select 1 km spatial resolution to reduce the simulation time as well as data storage requirement.

For the Northern Atlantic Oscillation Index (NAOI) (see Section 3.2), a daily time series from 1981 to 2010 is obtained from the National Oceanic and Atmospheric Administration (NOAA) website[1].

Note that the observed data are from different service providers, therefore the time of the data collection may differ which may affect the results.

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[1]https://www.cpc.ncep.noaa.gov/products/precip/CWlink/pna/nao.shtml

Table 1. List of the 29 meteorological stations whose data are used in the study. The names in **bold** letters are the three representative stations to illustrate the results.

Id no.	Name of the station	Longitude	Latitude	Altitude (m a.s.l.)
1	Gries im Sellrain	11.16	47.20	1200
2	Kuehtai	11.01	47.20	1970
3	**Oetz**	**10.89**	**47.21**	**760**
4	Umhausen	10.93	47.14	1040
5	Jerzens-Ritzenried	10.78	47.12	1120
6	Fliess	10.65	47.11	860
7	Landeck	10.57	47.13	800
8	Ladis-Neuegg	10.65	47.10	1350
9	See im Paznaun	10.46	47.08	1040
10	Laengenfeld	10.97	47.08	1180
11	St. Leonhard im Pitztal-1	10.84	47.08	1329
12	**Prutz**	**10.66**	**47.07**	**871**
13	Ried im Oberinntal	10.66	47.06	895
14	Fendels	10.68	47.05	1343
15	Kaunertal-Vergoetschen	10.75	47.04	1269
16	St. Leonhard im Pitztal-2	10.86	47.02	1460
17	Dresdner Huette	11.14	47.00	2290
18	Plangeross	10.87	46.99	1605
19	Soelden Schmiedhof	11.01	46.97	1380
20	Pfunds	10.51	46.95	992
21	Spiss	10.45	46.96	1540
22	**Pitztaler Gletscher (Pitztal Glacier)**	**10.88**	**46.93**	**2860**
23	Nauders	10.50	46.90	1360
24	Obergurgl	11.02	46.86	1940
25	Vent	10.91	46.86	1890
26	Vernagtbach	10.83	46.86	2640
27	Ausserrojen	10.48	46.81	1833
28	St. Martin im Passeier Beobachter	11.23	46.78	588
29	Marienberg	10.52	46.71	1310

3.2 Selection of covariates in the model

270   We allow several covariates in the model such that the model can capture a realistic structure of precipitation patterns over the region. This includes the day-to-day time dependence, seasonality as well as the influence of large scale circulation. As the

first covariate, we select previous day's occurrence (Occ_{t-1}) as possible covariate such that day-to-day temporal dependency in occurrence at a location is captured. To include seasonality, time dependent 1st and 2nd order harmonics of sine and cosine (see Table 2) are considered as the possible covariates. To allow the influence of large scale circulation over Europe, the NAOI is considered as a possible covariate. Studies show that there are links between the NAOI and precipitation characteristics (Casty et al. (2005); Beniston (1997)). A strongly positive NAOI is associated with persistent high pressure over the alpine region, resulting in warmer than average temperatures and lower than average precipitation. In general, winter NAOI correlates negatively with precipitation. Along with the described covariates, we also take into account their interaction terms as possible covariates.

For the selection of the covariates in the model, we use both the Akaike Information Criterion (AIC) (Akaike (1974)) and the Bayesian Information Criterion (BIC) (Schwarz (1978)). None of the two criteria selects the same set of covariates at all the stations. BIC has a tendency to select the simplest model whereas AIC has a tendency to select more complex models. However, it turns out that BIC helps in identifying the most important covariates at all the stations.

The three most important covariates for the occurrence model at the majority of stations are Occ_{t-1}, $\mathrm{Cos}(2\pi t/n)$ and $\mathrm{Cos}(4\pi t/n)$, where 't' is the day of year. We select those covariates which are selected by both AIC and BIC at the majority of the stations. The selected covariates are listed in Table 2. BIC selected this set of covariates at 18 stations out of total 29 stations (see Section 3.1), while AIC selected the same set of covariates at 11 stations. Thus, the vector of covariates in the model is

$$X_O(s,t) = (1, \mathrm{Occ_{t\text{-}1}}, \mathrm{Cos}\left(\frac{2\pi t}{n}\right), \mathrm{Sin}\left(\frac{2\pi t}{n}\right), \mathrm{Cos}\left(\frac{4\pi t}{n}\right), \mathrm{Cos}\left(\frac{4\pi t}{n}\right) * \mathrm{Sin}\left(\frac{4\pi t}{n}\right),$$
$$\mathrm{Occ_{t\text{-}1}} * \mathrm{Cos}\left(\frac{2\pi t}{n}\right), NAOI) \tag{8}$$

where 'n' is 365 or 366 in case of a leap year. The first term is associated with the intercept in the model.

For the precipitation amount also, we take into account all the possible covariates as described for the occurrence model. We select the covariates using both AIC and BIC for the amount model also. Additionally, selecting the same 7 covariates as in the occurrence model at the majority of stations, the 2nd harmonic of sine is also selected by both AIC and BIC (17 stations by BIC and 16 stations by AIC). Thus, we allow a total of 8 covariates in the amount model (see Table 2) and the vector of covariates for the amount model is

$$X_A(s,t) = (1, \mathrm{Occ_{t\text{-}1}}, \mathrm{Cos}\left(\frac{2\pi t}{n}\right), \mathrm{Sin}\left(\frac{2\pi t}{n}\right), \mathrm{Cos}\left(\frac{4\pi t}{n}\right), \mathrm{Sin}\left(\frac{4\pi t}{n}\right), \mathrm{Cos}\left(\frac{4\pi t}{n}\right) * \mathrm{Sin}\left(\frac{4\pi t}{n}\right),$$
$$\mathrm{Occ_{t\text{-}1}} * \mathrm{Cos}\left(\frac{2\pi t}{n}\right), NAOI) \tag{9}$$

Table 2. List of the covariates included in occurrence and amount model (Eq. 8 and Eq. 9)

No.	Name of the Covariate	Description	Occurrence model	Amount model
1	Occ_{t-1}	Previous day's occurrence	✓	✓
2	$Cos(2\pi t/n)$	1st harmonic of Cosine	✓	✓
3	$Sin(2\pi t/n)$	1st harmonic of Sine	✓	✓
4	$Cos(4\pi t/n)$	2nd Harmonic of Cosine	✓	✓
5	$Sin(4\pi t/n)$	2nd harmonic of Sine	✗	✓
6	$Cos(4\pi t/n) * Sin(4\pi t/n)$	Interaction of $Cos(4\pi t/n)$ and $Sin(4\pi t/n)$	✓	✓
7	$Occ_{t-1} * Cos(2\pi t/n)$	Interaction of Occ_{t-1} and $Cos(2\pi t/n)$	✓	✓
8	NAOI	Northern Atlantic Oscillation Index	✓	✓

'n' is 365 (366 in case of leap year)

300 The correlations for precipitation amount in the model are computed only for days when the precipitation was observed.

3.3 Model evaluation strategy

Although the model produces daily fields of precipitation, before evaluating the model for gridded data, we first evaluate it at the individual locations where the observations are available. This is common practice for validation of WGs that the input statistics must be reproduced. From the simulated gridded data, the 30 years time series of daily precipitation at the nearest grid
305 point to the observation locations is extracted from each of the N=30 realisations. The mean of the simulated statistics in each realisation is compared with the observed statistics. The validation is carried out for daily and long-term statistics along with considering more difficult statistics to be reproduced by the model. For the illustration of the results at individual locations, out of 29 stations, three example stations are selected which are: i) Oetz, ii) Pitztal Glacier, and iii) Prutz. These three stations are highlighted by red colour in the study area (Figure 1). The three stations are selected such that Oetz represents the results at val-
310 ley stations, Pitztal Glacier represents the results at the high-mountain stations and the third station Prutz is one of the stations with different climatic characteristics and has climatic characteristics most distinct from those at the surrounding stations (and is, therefore, most challenging). Note that Pitztal Glacier is the highest station amongst the 29 observation stations (see Table 1).

In the next step, we evaluate the model for its ability to reproduce spatial statistics. For that, gridded observed data are required.
315 We use the Alpine Precipitation Grid Dataset (APGD) (Isotta et al. (2014)) from the Swiss Federal Office for Meteorology and Climatology (MeteoSwiss) which has a 5 km spatial and daily temporal resolution. The dataset is based on measurements at high-resolution rain-gauge networks, incorporating more than 8500 stations from seven Alpine countries with more than 5500 rain-gauge measurements on average per day. With 10-15 km station spacing, the dataset is one of the densest in-situ observation networks over a high-alpine topography worldwide. These data are available from 1971 to 2008. Note that this dataset is

320   not a perfect reference. To obtain 30 years of gridded observations, we select the period from 1979 to 2008 for the validation of the simulated gridded data.

In order to assess the interpolation accuracy, we perform a hold-out cross validation in which one or more stations are withheld from the model fitting process. We withhold the same three stations which are selected for the illustration of the results,

325   i.e. Oetz, Pitztal Glacier, and Prutz. The model should be able to reproduce the observed statistics accurately at the withheld stations. For cross-validation also, we generate N=30 realisations of 30 years i.e. 900 years of data.

For the uncertainty estimation in the N=30 realisations, we use a tolerance interval (TI) (Patel (1986), Krishnamoorthy and Mathew (2009) and Young (2010)) instead of the conventional way of using a confidence interval for the sampling error, which

330   is sensitive to sample size. As opposed to confidence intervals, which would give expected bounds on the means of the simulated data, the tolerance interval gives bounds on the future individual observations. In our view, TIs provide an appropriate visualization of the expected variability of the simulated data, as well as a means of comparison with the original data. Here, we use a parametric two-sided TI with a normal distribution. The TIs are computed for each of the statistics considered in this study, obtained from the simulated 30 realisations at each station. As uncertainty criteria, we select a confidence interval of

335   95 % and a 99 % proportion of the population for the T.I, i.e. the TIs indicate the 99% range of the simulated values (with 95% confidence). The TIs are shown in each figure as a shaded area around the curve and denoted as TI_{99}^{95} throughout the article.

To quantify the model performance, along with various error metrics, we also take into account correlation coefficients (CC) and coefficients of determination (R^2). All together, we employ the following metrics: i) mean bias error (MBE), ii) mean ab-

340   solute error (MAE), iii) root mean square error (RMSE), iv) CC, and v) R^2. Each of the performance metrics serves a different purpose. MBE measures the overall bias in the model performance, MAE and RMSE both provide information on the mean magnitude of the error regardless of the direction of the error. However, the greater the difference between them, the greater the variance in the individual errors in the sample. Similarly, correlation shows the association between the two variables (which are observed and synthetic statistics here) but R^2 shows the proportion of data variation explained by the model. All the perfor-

345   mance metrics for the spatial statistics (corresponding to Figure 13) are obtained between the observed and synthetic statistics derived at all the grid points, e.g. for the spatial occurrence probabilities, the metrics are derived between the observed and synthetic "spatial series" of the occurrence probabilities.

4   Results

350   An extensive evaluation of the model generated data is carried out here.

4.1 Evaluation at individual stations

4.1.1 Daily occurrence probabilities at individual stations

First, we assess the performance of the occurrence model for the daily conditional probabilities on which the model is trained. These are important statistics as they are associated with the ability of the model to reproduce dry and wet spells. Figure 2 illustrates the annual cycle of the empirical and simulated daily conditional probability of a dry day following a dry day (P_{dd}) at the three selected stations along with the results of cross-validation for the same variable, whereas Figure 3 illustrates the daily conditional probability of a wet day following a wet day (P_{ww}). The solid lines are the fitted curve using the Locally Weigthed Scatterplot Smoothing (LOESS) method (Cleveland (1979)) to each of the three probabilities — observed, simulated and cross-validated, respectively. The annual cycle of P_{dd} is accurately simulated at Oetz and Pitztal Glacier while at Prutz, the model largely underestimates throughout the year. However, the seasonality in P_{dd} is well captured at Prutz. The annual cycle of P_{ww} is well captured at Oetz and Pitztal Glacier, but slightly underestimated at Oetz during the entire year whereas at Pitztal Glacier, the model accurately reproduces the probabilities throughout the year.

At Prutz, the model performs badly for P_{ww} (Figure 3 (c)). The observed seasonality in P_{ww} at Prutz is completely different compared to the seasonality at other stations which is not reproduced by the model at all. Similar to Prutz, the other two stations (St. Leonhard im Pitztal-2 and St. Martin), which have very different climatic characteristics, also exhibit marked differences (not shown) between the simulated and the observed daily values of P_{dd} and P_{ww}. What is noteworthy here is that the magnitude and the seasonality in both P_{dd} and P_{ww} at Prutz are close to the magnitude and seasonality of P_{dd} and P_{ww} at the valley station (Oetz). Similar behaviour is observed at the other two stations with poor performance. This is partly due to the fact that the covariates, which were selected at the majority of stations as optimal, did not optimally represent the statistics at Prutz (and neither at the other stations with distinct climatologies). This indicates that the selected set of covariates is not able to reproduce seasonality at all the stations because a small subset of stations has distinctly different seasonality. Furthermore, the performance of the model at Prutz is influenced by the climate characteristics at the nearby stations with regard to magnitude and seasonality and not the other way around. At all other stations, P_{dd} and P_{ww} are well simulated. This suggests that the selected harmonics are capable to capture the seasonality in daily occurrence probabilities. Moreover, the temporal dependency in the occurrence is well reproduced by the covariate Occ_{t-1}. In general, we find that the performance of the model at valley stations is similar to that at 'Oetz', and at the high-mountain stations similar to Pitztal Glacier.

Cross-validation produces similar results at all the three sites for P_{dd}. Compared to the results with the simulated probabilities, Prutz and Oetz show slightly more underestimation throughout the year, while Piztal Glacier has mostly the same results with only minor deviations. For P_{ww}, Oetz has similar results throughout the year, while Pitztal Glacier has deviations in spring and Prutz in autumn and winter. The seasonal differences are likely due to very different precipitation characteristics at Prutz which was removed from the model fitting along with Oetz and Pitztal Glacier in cross-validation.

385   To further examine the ability of the model to reproduce the observed climatology of wet days, we next consider the uncon-
ditional daily occurrence probability of wet days (P_w) (Figure 4). Again, at both Oetz and Pitztal Glacier, P_w is very well
simulated and the annual cycle of P_w is well captured by the model. At Prutz, the generator not only largely overestimates the
probabilities but is also not able to accurately capture the seasonality. Again, both the seasonality and magnitude at Prutz in
simulated probabilities are closer to those at valley stations (such as Oetz). Cross-validation produces similar results at all the
390   three sites for P_w also with small deviations compared to the simulated results.

The performance metrics for daily occurrence probabilities are shown in Figure 8 (a) (note that the performance metrics are
presented for the simulated data and not for the cross-validation). The metrics are computed for each of the 29 stations and
plotted as a boxplot. All the error metrics, CC and R^2 suggest the best performance is for P_{dd} followed by P_w and then P_{ww}.
395   The high values of CC and R^2 for P_{dd} and P_w demonstrate the overall very good performance for these two statistics. Con-
versely, the small values of CC and R^2 for P_{ww} indicate a relatively poor performance of the model. This is – at least partly –
due to the fact that the model performs very poorly in generating precipitation series at the stations with distinct climatologies
(see Figure 3 (c)).
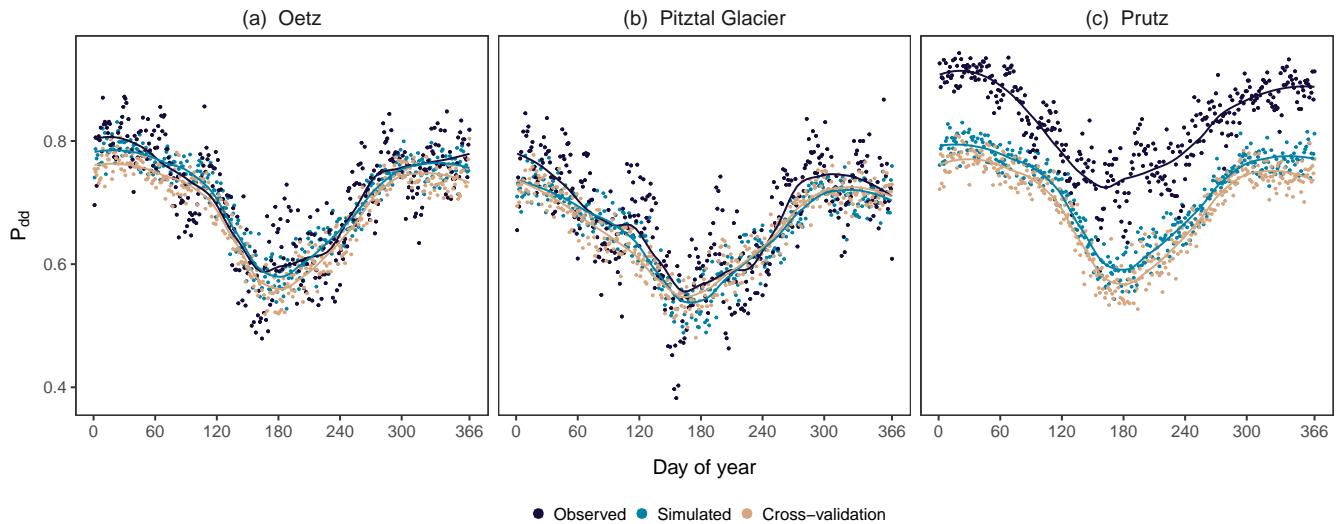
Figure 2. Daily conditional probability of a dry day following a dry (P_{dd}) day at three selected stations: (a) Oetz, (b) Pitztal Glacier, and (c)
Prutz, (Figure 1 and Table 1). The observed probabilities (navy blue) are obtained from the observed 30 years (1981–2010). The simulated
probabilities are the mean of the 30 realizations (skyblue) and from the hold-out cross validation simulation (brown). The solid lines are the
fitted curves using the LOESS method to the observed and simulated probabilities, respectively.
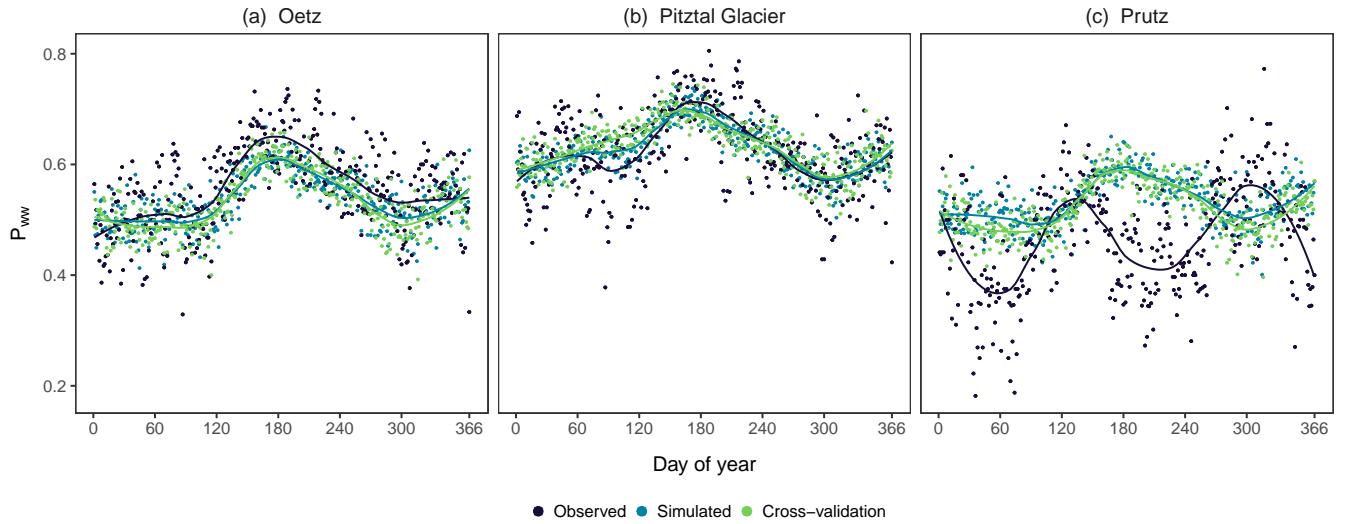
Figure 3. The same as Figure 2, but for daily conditional probability of a wet day following a wet day (P_{ww}).
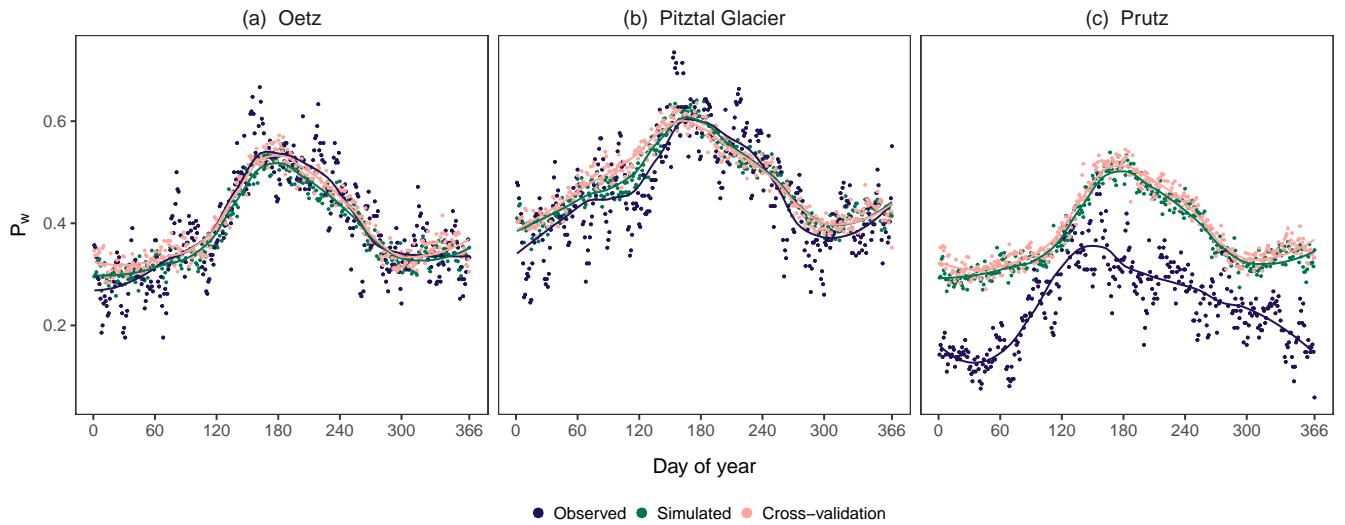
Figure 4. The same as Figure 2, but for daily occurrence probability of wet days (P_w).

400 4.1.2 Frequency of spells of different lengths at individual stations

Another important feature of the model is its ability to simulate long sequences of wet and dry days. Here, we examine the ability of the model to simulate wet and dry spells of different length at individual stations. Figure 5 displays the frequency of wet spells at the selected stations for the length 2 to 15 days. The precipitation generator is able to reproduce the wet spells of
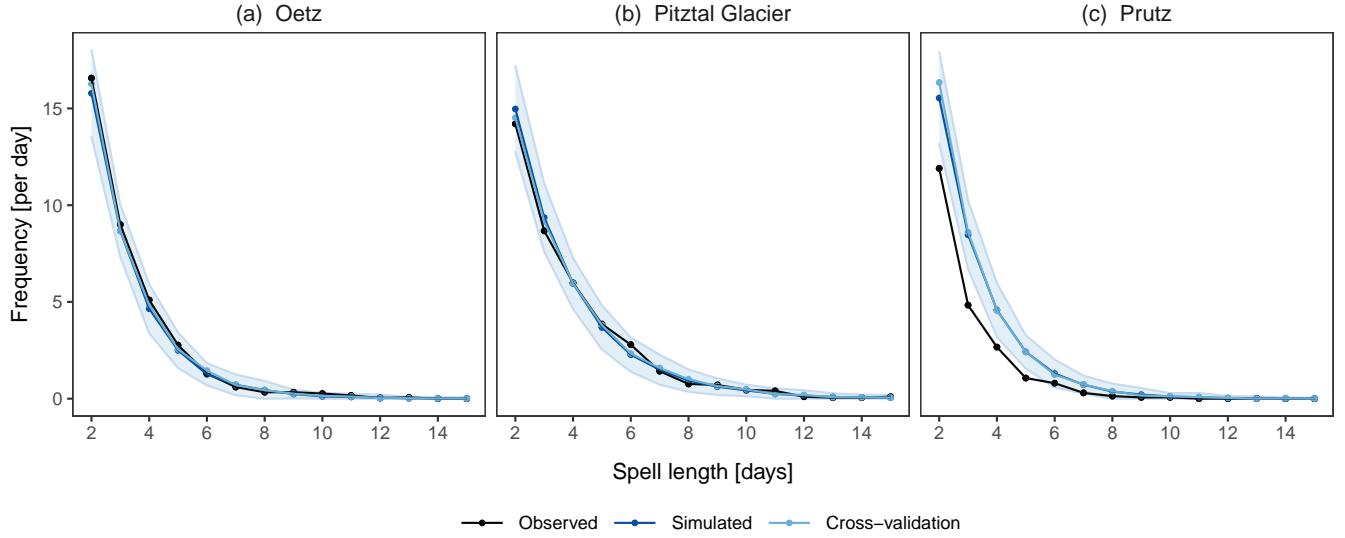
Figure 5. Frequency of wet spells of different length at three selected stations: (a) Oetz, (b) Pitztal Glacier, and (c) Prutz. The frequency of the observed spells is obtained from the observed 30 years (1981–2010). The frequency of the simulated spells is determined for each of the 30-year simulations separately and averaged over the 30 realizations. The shaded region is the selected tolerance interval TI^{95}_{99} for the simulated values in the 30 realisations. The two-sided tolerance interval is specified for 99 % of the population and with 95 % confidence level. Note that the TIs are not presented for cross-validation.

different length at the valley and high-mountain station accurately. The shaded region is the tolerance interval TI^{95}_{99} in the 30
405 realisations. The observed values of the spells of different length are within the tolerance interval which shows that the model does an excellent job in reproducing even the longer spells. It is noteworthy that the model captures the spells very accurately even if it is not trained on these statistics. There are very few occurrences of wet spells longer than 10 days at the majority of stations in the observed data which are also well reproduced. At Prutz, the model overestimates the wet spells of all different lengths. The observed values of the spells are not within the TI^{95}_{99} which shows that the overestimation is consistent in all the 30
410 realisations. The overestimation occurs because the model is not able to reproduce the conditional probabilities P_{ww} at Prutz even reasonably well. In fact, the large overestimation in P_w at Prutz (Figure 4) ultimately contributes to the overestimation in wet spells. Again, cross-validation produces similar results.

Figure 6 demonstrates the frequency of dry spells of length 2 to 25 days at three selected stations. The model is able to simulate
415 the dry spells of all different lengths at the valley station as well as high-mountain station with great accuracy. There are very few occurrences of extreme spells of length longer than 15 days in the observed data which are also well reproduced by the model. At Prutz, the model overestimates the spells of shorter length of up to 4 days but does a very good job for the longer spells. Again, the shaded regions are the statistical tolerance level TI^{95}_{99} in the 30 ensembles. The observed values of the spells
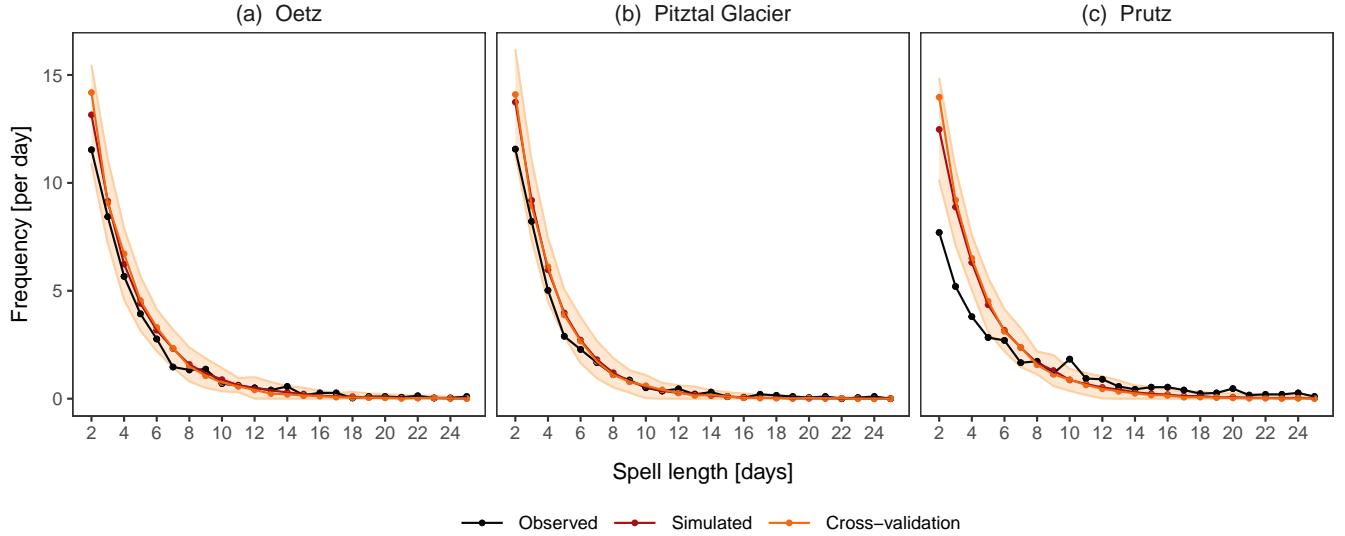
Figure 6. As Figure 5, but for the frequency of dry spells of different length.

are within the tolerance region except for the shorter spells even at Prutz. The worst performance for both wet and dry spells has
been observed at Prutz. In general, for the stations in the valley, the model does a better job than at the high-mountain stations
(figure not shown). The previous day's occurrence as a covariate (Occ_{t-1}) is able to simulate the long sequences of both wet
and dry days very well at the majority of the stations with a great accuracy. The occurrence model satisfactorily reproduces the
occurrence pattern at the majority of the stations, except at the few distinct stations with peculiar climatologies. The results of
the cross-validation are similar again.

As for dry and wet spells, the model performance is mostly similar for both the spells (Figure 8 (b)) except that the RMSE
is, in general, larger for wet spells suggesting overall slightly worse performance for wet spells. The CC and R^2 are nearly 1
which indicates an excellent agreement of the model with the observations for both dry and wet spells.

4.1.3 Monthly mean precipitation at individual stations

An important aspect of the precipitation generator is its ability to reproduce the amount of precipitation observed at the stations.
As the model for amount is the gamma distribution at the observed locations, the mean, which is the product of the shape and
scale parameters of the gamma distribution, should be well reproduced. Figure 7 displays the monthly mean of precipitation
in the observed and simulated data at the selected stations along with the results of cross-validation. At both Oetz and Pitztal
Glacier, the model is able to reproduce the mean very well as the observed values are within the TI_{99}^{95} (the shaded region). At
Prutz, the model underestimates the mean in April, May, October and November while in other months, the model is able to
reproduce the mean reasonably well as the observed values are within the TI_{99}^{95}. The results for cross-validation are similar at all
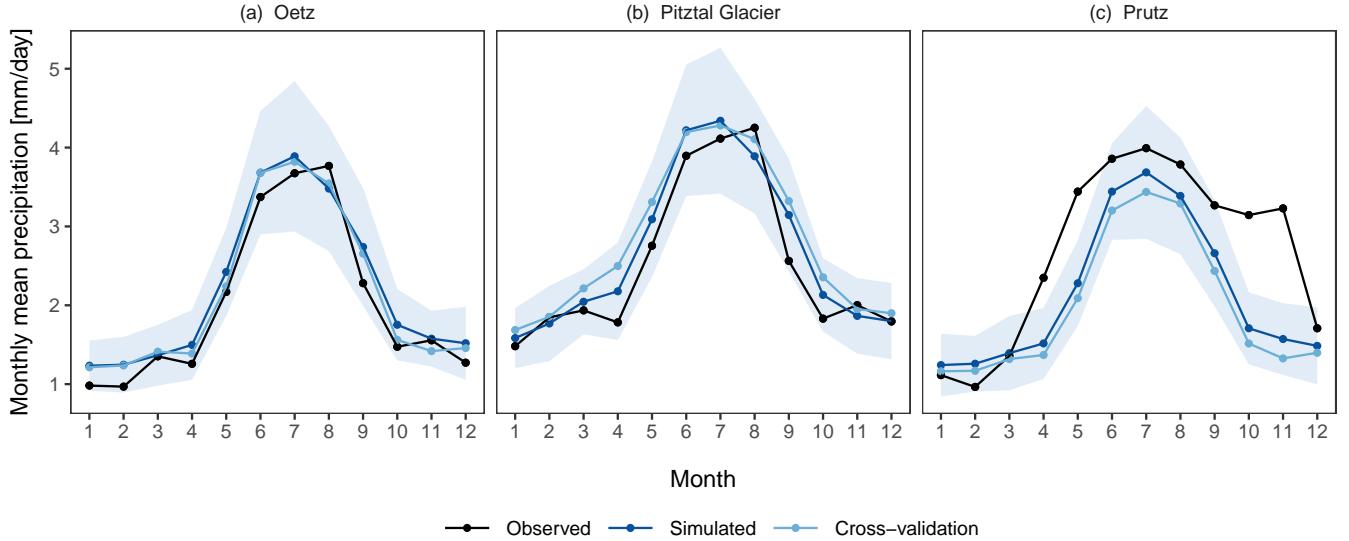
Figure 7. Monthly mean precipitation [mm/day] at the three selected stations: (a) Oetz, (b) Pitztal Glacier, and (c) Prutz. The observed values are obtained from the observed 30 years (1981–2010) and the simulated values are the mean of the 30 realisations. The shaded region is the selected tolerance interval TI_{99}^{95} for the simulated 30 realisations.

the three stations, but compared to simulated data, Pitztal Glaciers and Prutz show slight deviations. The performance metrics are displayed in Figure 8 (b). For the monthly mean precipitation, the model performs very well which can be seen by the small magnitudes of the errors – typically less than 0.5 mm/day, and strong values of CC and R^2.

440
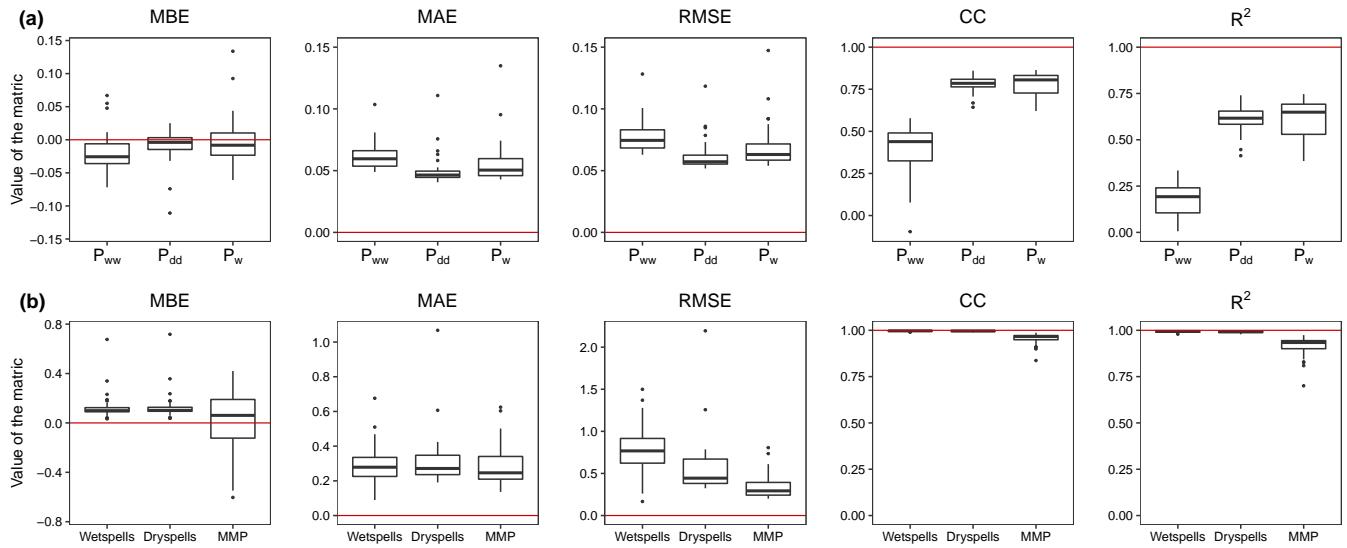
Figure 8. Performance metrics for the model at individual stations: (a) for daily occurrence probabilities (P_{dd}, P_{ww} and P_w), and (b) for the frequency of dry spells and wet spells of different length and monthly mean precipitation (MMP). The unit for the MBE, MAE and RMSE for dry/wet spells are [per year] and for MMP is [mm/day]. The red horizontal line in each panel plot corresponds to the optimal performance for the corresponding metric.

4.1.4 QQ plot at individual stations

Here, we examine the QQ plot at each individual station where the observed data are available (Figure 9). It can be seen that the distribution of the precipitation is accurately simulated at a majority of the stations. Largest discrepancy between the observed and simulated distribution is found at Prutz which has the largest spread in the observed data which is not well reproduced by the model. Apart from that, there are some stations including Dresdner Huette where the higher quantiles are not well reproduced. These are the stations which have longer tails in the distribution of precipitation in the observed data. This is a commonly reported problem in the literature that the gamma distribution is not adequate to simulate extremes.

We further examine the distributions of the generated data for each month at each of the 29 stations using the Kolmogorov-Smirnov test and the Wilkoxon-Mann-Whitney test. The results are shown in the Supplement (Figures S1–S2). As revealed by the QQ plots, the worst performance is observed at Prutz, St. Martin and St. Leonhard im Pitztal-2 which have the distinct climatologies.
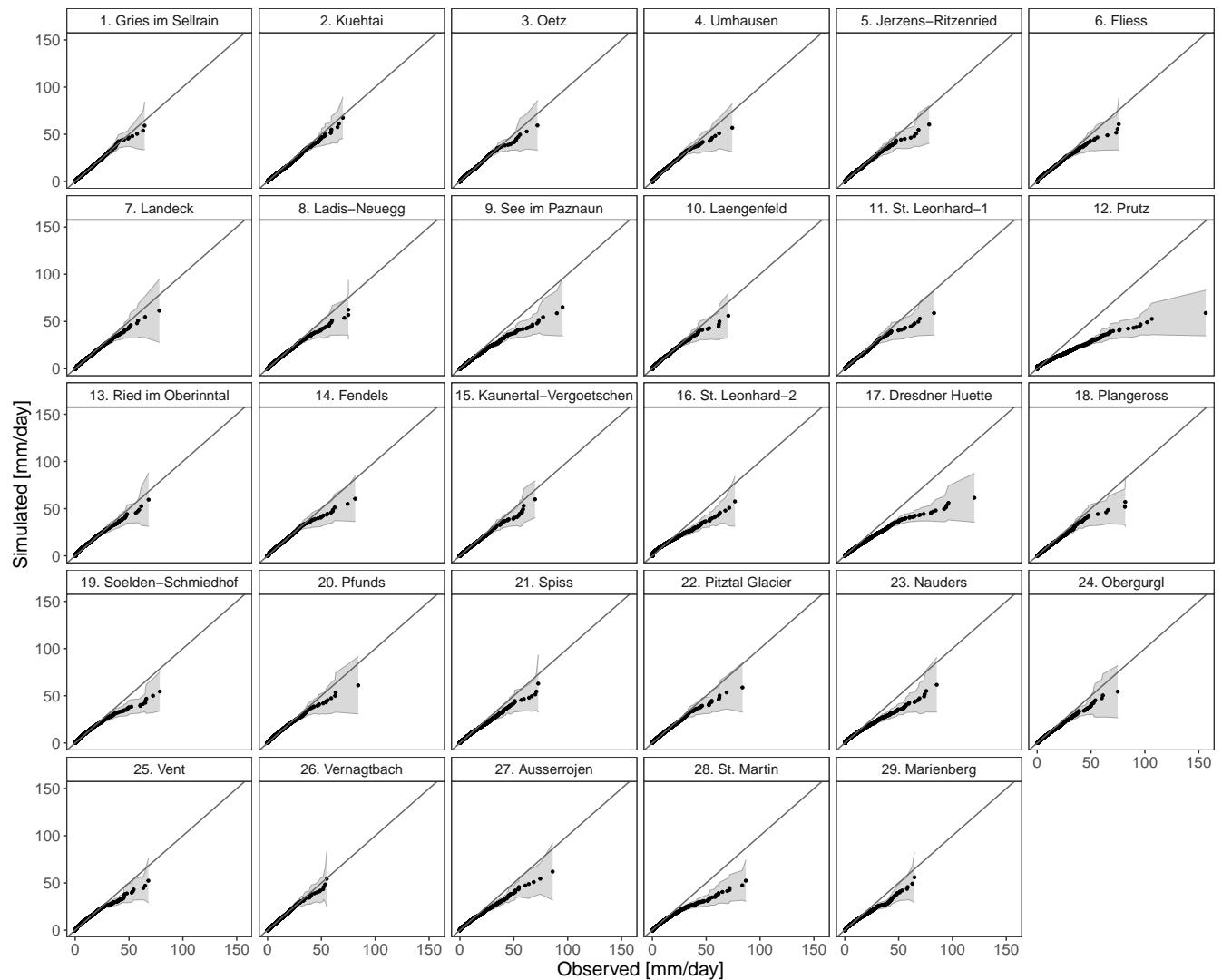
Figure 9. QQ plot showing percentiles of daily precipitation at each of the 29 individual stations (Table 1). The shaded region is the selected two-sided tolerance interval TI_{99}^{95} for the simulated 30 realisations. The black solid line in each panel plot corresponds to the 1:1 line.

4.2 Evaluation of simulated gridded data

In this section, we focus on evaluating the spatial dependence structure in the simulated data which is an important aspect for many hydrological applications. We present here the results of the validation — first for the spatial statistics related to the occurrence and then for the occurrence and amount together. More details on this aspect of validation is given in the discussion Section 5.

4.2.1 Frequency of areal spells of different length

One of the most challenging features of the gridded precipitation generator is its ability to reproduce the areal spells of wet and dry days of different length. This is one of the sought-after features in agricultural and hydrological applications. We define an areal wet (dry) spell as the number of consecutive days when 95 % of the study area is wet (dry). Figure 10 illustrates the areal spells of dry and wet days of different length in the observed and simulated data. It can be seen that the areal dry spells are better simulated than the areal wet spells. The areal dry spells of length greater than 2 days are accurately simulated, whereas the areal dry spells of 2 days are overestimated. For areal wet spells, the model underestimates the spells of all lengths. Larger discrepancies are found for shorter spells.

For the areal spells, the performance metrics are shown in Table 3. For areal dry spells, the error statistics suggest that the model has a tendency to overestimate the spells. This is because the model largely overestimates the areal dry spells of shorter lengths. The CC is perfect and R^2 is also nearly 1 which suggest an excellent performance of the model. For areal wet spells, the error metrics MBE, MAE and RMSE are small where a negative value of MBE indicates overall underestimation in the wet spells and strong CC and R^2 suggest a very good agreement with the observed values.

Table 3. Performance metrics for the gridded model for reproducing areal statistics

Statistics	MBE	MAE	RMSE	CC	R^2
Frequency of areal dry spells of different length	1.02 [per year]	1.38 [per year]	3.34 [per year]	1.00	0.99
Frequency of areal wet spells of different length	-0.41 [per year]	0.62 [per year]	1.30 [per year]	0.99	0.97
Monthly mean areal precipitation	-0.23 [mm/day]	0.26 [mm/day]	0.33 [mm/day]	0.96	0.93

4.2.2 Spatial distribution of occurrence probabilities

The spatial distribution of the simulated probabilities of wet days (P_w) for each season (Figure 11 (left column)) is compared to that of the (gridded) observations (Figure 11 (middle column)). It is noteworthy that the model generated data have 1 km spatial resolution, while the observed gridded data have 5 km resolution. The model is able to generate the seasonality in the spatial distribution of P_w very well. In particular, the higher probabilities in spring and summer are well reproduced. Also, at 1 km high resolution, the influence of topography on the spatial distribution of P_w is clearly visible. This is due to the inclusion
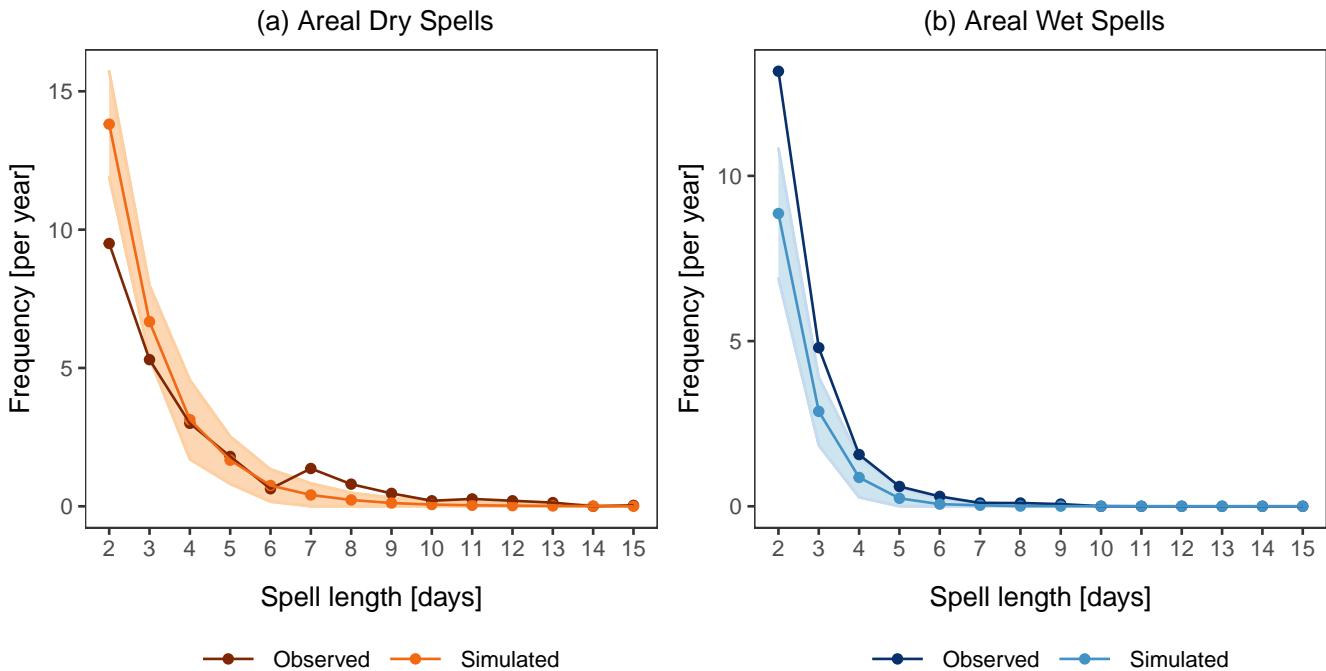
Figure 10. Frequency of the areal spells of different length (see Section 4.2.1): (a) Dry spells, (b) Wet spells. The observed spells are obtained from the APGD gridded data of 30 years (1979–2008). The simulated spells are the mean of the 30 realizations. The shaded region is the selected tolerance interval TI_{99}^{95} for the simulated data. The two-sided tolerance interval is specified for 99 % of the population and with 95 % confidence level.

480   of elevation in the kriging interpolation.

To test the model performance at each grid point, we upscale the simulated data from 1 km to 5 km and the difference between the observed and the simulated data is shown in the right column in Figure 11. It can be seen that in spring and summer the biases are small and mostly similar throughout the domain, while in winter there is an overestimation in the south-eastern
485   part of the region and a slight underestimation in the north-western part. This is because in the south-eastern part of the study area the density of stations is very low. There is one station (St. Martin) that is heavily influencing that area which has the highest probability of precipitation occurrence and that is manifested through the spatial interpolation of the model parameters through kriging (see Section 2) to generate the gridded fields. Except for summer, this station has a much higher probability of precipitation occurrence relative to the other stations. It is also reflected in the simulation that in summer in the south-eastern
490   part the bias is mostly zero.
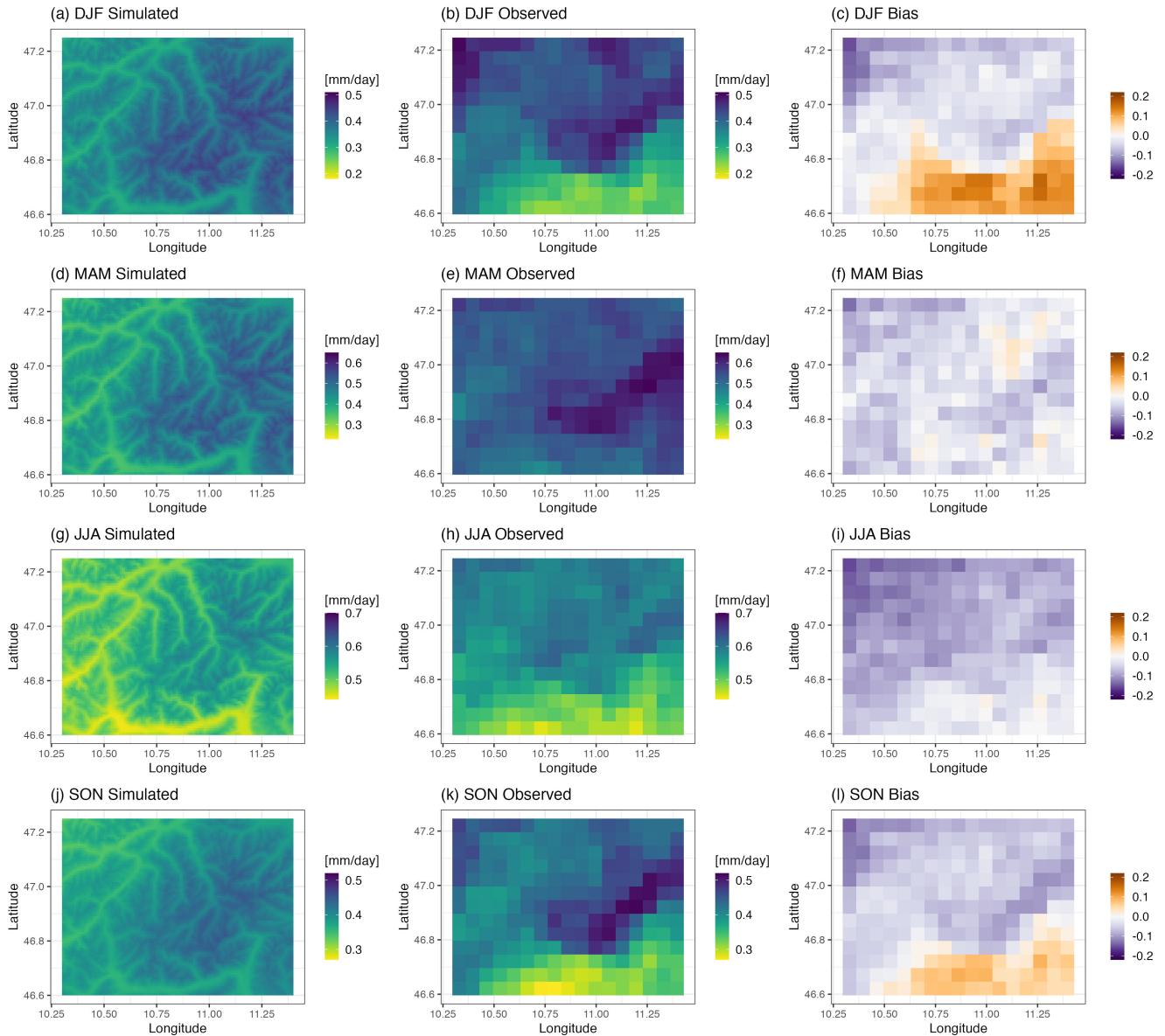
Figure 11. Spatial distribution of occurrence probability (P_w) in four seasons in the simulated 30 realisations (left column) and 30 years (1979–2008) APGD observed gridded data (middle column) and the bias (simulated-observed) in the simulated data (the last column). The spatial resolution for APGD data is 5×5 km. Note the different scales in the color coding for the different seasons.

Figure 13 (a) illustrates the performance metrics for the spatial distribution of P_w for each month. All the three error statistics are within 0.1 which suggests a very good performance of the model. The largest errors are for August where negative MBE indicates overall underestimation in P_w, while the largest overestimation is found in February. The lowest overall bias is found

495   in September. The highest value of CC and R^2 is observed in May. In colder months, i.e. in winter and autumn the correlations between the observed and simulated probabilities are weak and also the R^2 is small which suggests poor performance for these months. The worst performance is observed in January where the CC and R^2 are almost zero, while the best is found in May.

4.2.3   Spatial distribution of mean wet-day daily precipitation

500   Next, we consider the mean wet-day daily precipitation for assessing the ability of the model to reproduce the observed climatology of precipitation amount over the region. We consider the mean of daily precipitation on wet days in each season over the 30 years of simulated 30 ensembles, i.e. 900 years of data at each grid point (Figure 12 left column) and compare it with the observations (Figure 12 middle column). The right column in Figure 12 depicts the bias in the simulated and observed values at each grid point. The model is able to simulate the spatial seasonal variability in precipitation very well. In particular, the

505   summer precipitation due to the convective processes and thunderstorms, which account for high amounts of precipitation, is well reproduced. In the colder seasons (DJF and SON), the precipitation amount in the simulated data has visibly less elevation dependence and hence the generator has a tendency to generate less spatial variability of daily precipitation amount across the region. In contrast with the largely overestimated spatial probability of occurrence in the south-eastern part of the study area (Figure 11), the model underestimates the precipitation amount. St. Martin does not only have a high frequency of wet days but

510   also the highest precipitation amount in all the seasons among the 29 stations. The gamma distribution is not able to reproduce the large amount at this station which results in underestimating the amount in the surrounding area also.

Another important aspect to notice is that in the observed data APGD, the probabilities of wet days in the south-eastern part of the region in the colder seasons are low (Figure 11), whereas in the same part of the region, the precipitation amount is high –

515   particularly in SON (Figure 12). The model fails to capture this behaviour. This is also because in this part of the study area the model has to extrapolate the parameters as there is no observed station present beyond St. Martin. Hofstra et al. (2010) found that the density of station network used for interpolation influences the distribution of precipitation and the areal mean amount of precipitation. They found that when fewer stations are used, precipitation is over-smoothed which leads to a strong tendency for interpolated values to be less than the "true" value, and the effect was largest for higher percentiles. Dresdner Huette also

520   has a larger amount in observed data in autumn (and also in winter) compared to other stations. The model underestimates the amount in the region surrounding this station also, which is mainly in SON (see the north-eastern part in SON in Figure 12). Another reason for the underestimation in autumn could be the inability of the model to simulate orographic precipitation particularly related to föhn events. This altogether leads to large underestimation in autumn in this region.
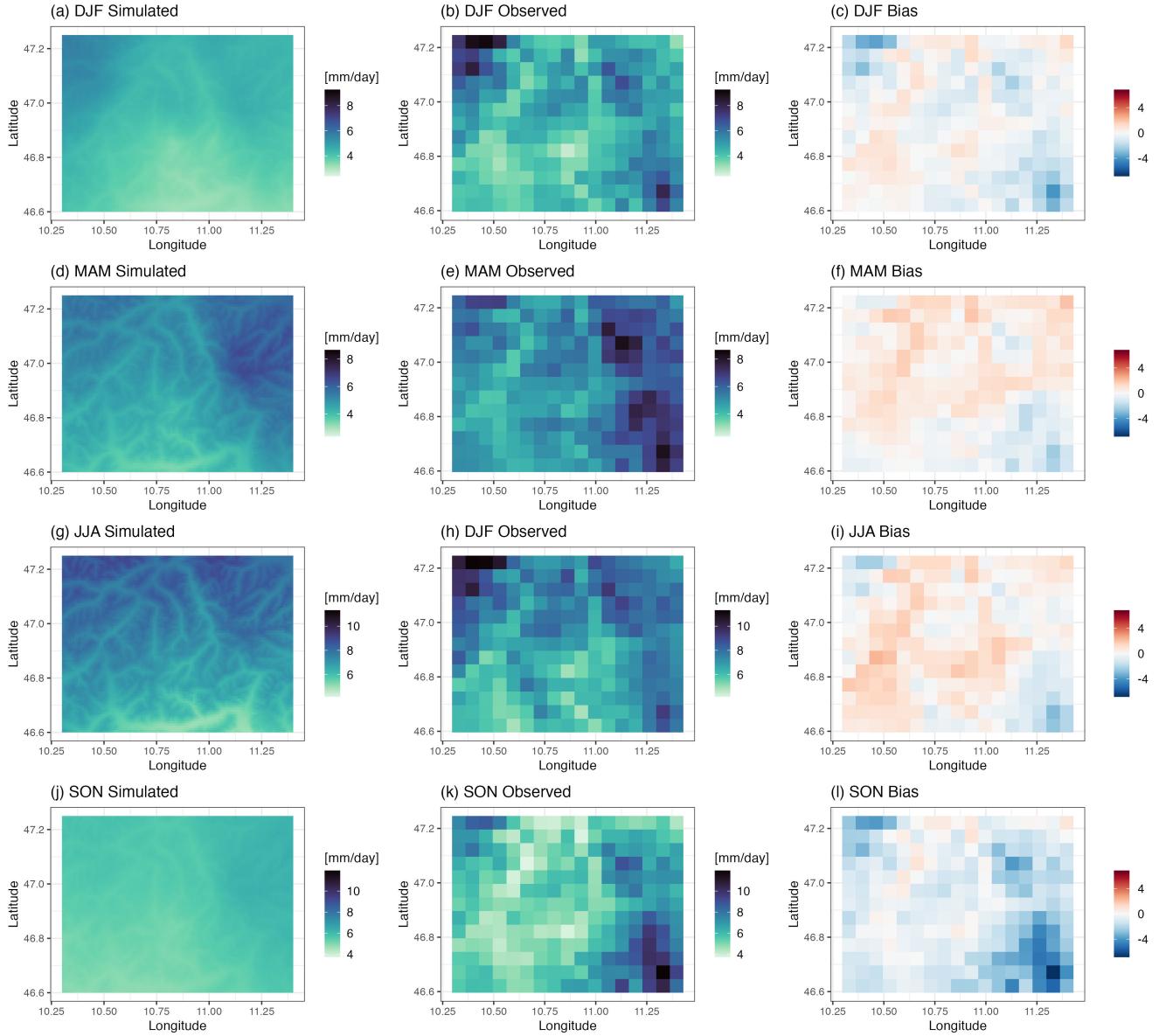
Figure 12. Spatial distribution of mean wet-day daily precipitation amount for four seasons [mm/day] for simulated (left column), observed (middle column) and bias (simulated-observed) in the simulated data (right column). The mean wet-day daily precipitation is obtained at each grid point for each season in each of the 30 realisations and the average of the 30 realisations is shown. The spatial resolution for simulated data is 1×1 km and for observed data is 5×5 km. Note the different scales in the color coding for the different seasons.
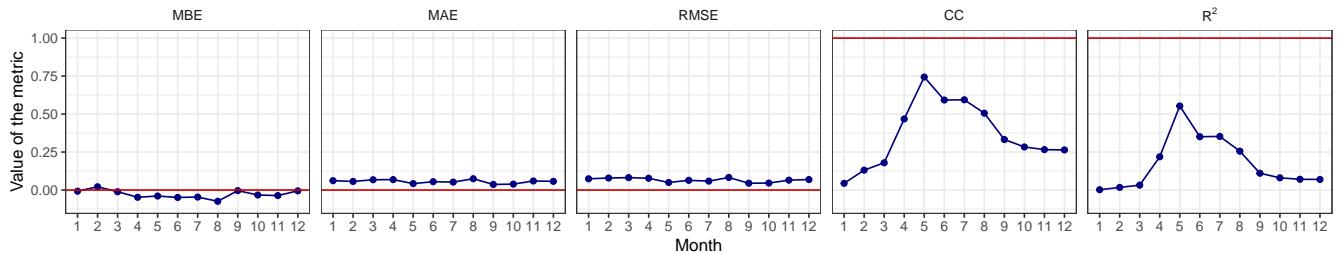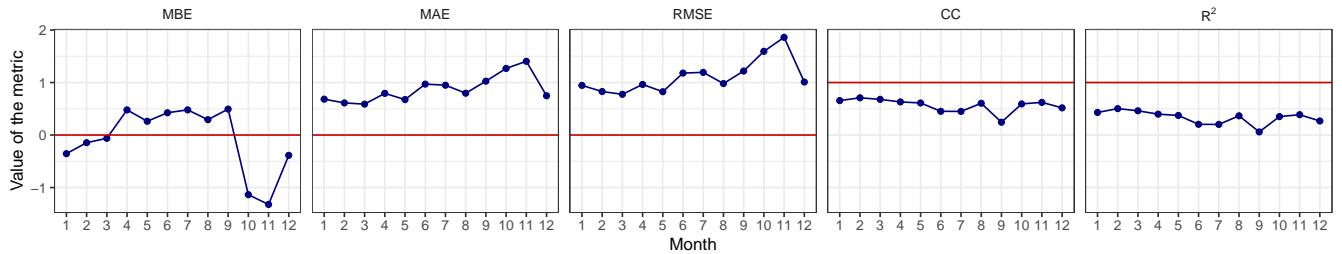
Figure 13. Performance metrics for the gridded model: (a) for spatial distribution of probabilities of wet days (P_w) in each month, and (b) for spatial distribution of mean wet-day daily precipitation where the unit for the MBE, MAE and RMSE is [mm/day]. The red horizontal line in each panel plot corresponds to the optimal performance for the corresponding metric.

525 Figure 13 (b) depicts the performance metrics for the spatial distribution of mean wet-day daily precipitation for each month. The MBE, MAE and RMSE are the largest in November and the negative value of MBE of approximately -2 mm/day suggests large underestimation in that month. For October also, large underestimation is found, while the smallest error metrics are found in March followed by February. The high values of CC and R^2 are in February and March which suggest the best performance for these months, while the worst performance is found in September. Contrary to the bad performance for spatial 530 distribution of P_w, the spatial distribution of mean precipitation is better reproduced in winter (compare the right columns of Figures 11 and 12).

4.2.4 Monthly mean areal precipitation

Next, we assess the ability of the precipitation generator to provide an areal climatology of the precipitation amount. This is 535 also one of the desired characteristics for impact modelling. Figure 14 displays the areal precipitation mean for each month in the observed as well as simulated gridded data. The precipitation generator, except in autumn, simulates the mean areal precipitation in all seasons with good accuracy as the observed values are within the TI_{99}^{95}. As seen in the spatial distribution of the amount of precipitation (Figure 12), in autumn the model underestimates mainly in October and November. The statistics are shown in Table 3. The small negative value of MBE indicates overall slight underestimation, while MAE and RMSE are
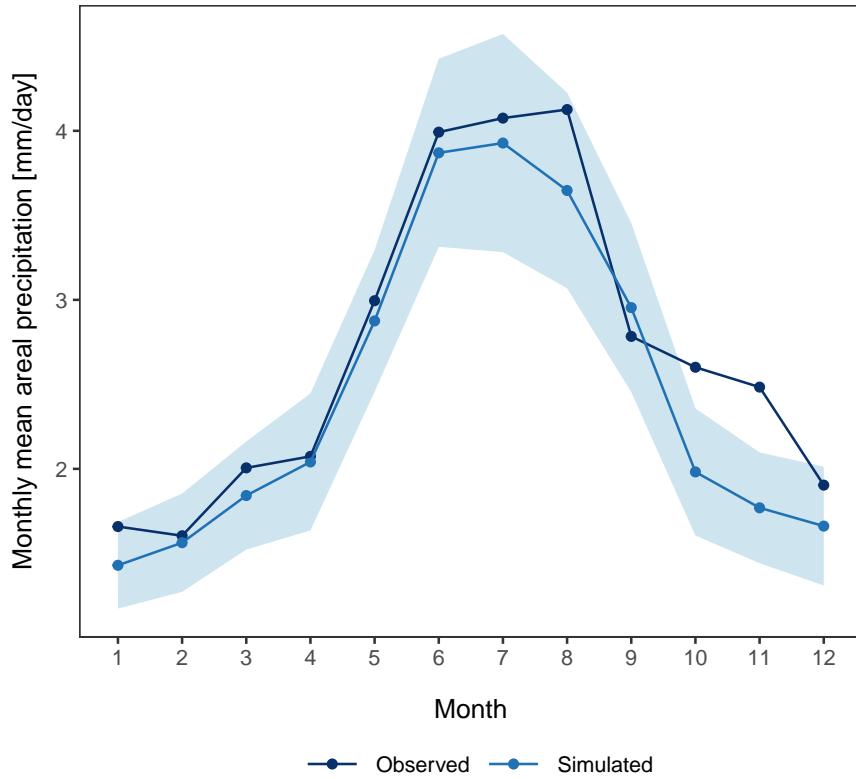
Figure 14. Monthly mean areal precipitation [mm/day]. The observed values are obtained from the APGD observed 30 years (1979–2008) and the simulated values are the mean of the 30 realisations. The shaded region is the selected two-sided tolerance interval TI_{99}^{95} for the simulated data.

540 also small, about less than 0.4 mm/day. The high values of CC and R^2 show that the model estimated means are in very good agreement with the observed.

4.2.5 Annual maximum precipitation sums

Here, we assess the ability of the model to reproduce extreme precipitation amount. We consider the annual maximum daily
545 precipitation at 29 sites. The daily sum of the observed daily precipitation at 29 sites is obtained and the maximum of the daily sum in every year is shown as a boxplot in Figure 15. For the simulated data, the time series of daily precipitation for 30 years at the nearest grid point to the 29 observation stations are extracted from the gridded data in each of the 30 realisations. The sum of the daily precipitation at the 29 sites in the simulated data is obtained and the maximum in each year is presented as a boxplot in Figure 15. There is an underestimation in the simulated maximum precipitation sums. The median is underestimated
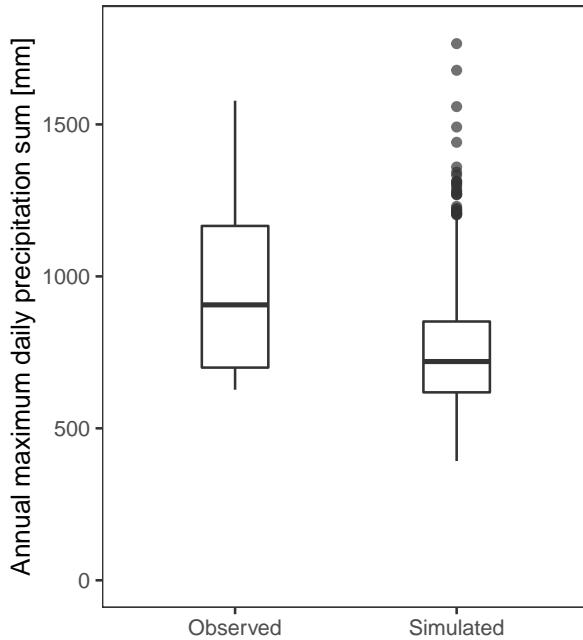
Figure 15. Annual maximum daily precipitation sum at 29 sites for observed and simulated data. The sum of daily precipitation at 29 sites is obtained and the maximum in each year is presented as a boxplot. Thus, the boxplot for observed is based on 30 values and for simulated is based on 900 (30 realisations of 30 years) values.

550 approximately by 20%. The interannual variability in the maximum precipitation sums is also underestimated as can be seen from the interquartile range (IQR). However, the range of the boxplot of the model simulated values is larger than that of the observed where the extreme outliers suggest the model generates higher extreme precipitation sums than those in the observed. In general, the model is able to generate the extreme precipitation reasonably well.

4.3 Comparison between anisotropic and isotropic model using KED and OK

555 Here, we compare the results of our simulations i.e. using KED in the anisotropic model (Aniso-KED) with three different model set-ups: i) by considering OK in the interpolation of the parameters of the anisotropic model (Aniso-OK), ii) using the original isotropic model which uses OK for the interpolation of the parameters (Iso-OK), and iii) using the isotropic model with KED (Iso-KED). We examine the results for the monthly sum of areal mean daily precipitation in these four cases of simulations (experiments) with the observed gridded data APGD. Figure 16 displays the distributions of the monthly sum of 560 areal mean daily precipitation for the simulated 900 years for each of the four experiments and the observed 30 years gridded data.

The model performance for Aniso-KED and Iso-KED is mostly similar in all the months, whereas the model performance for Iso-OK and Aniso-OK is similar in all the months. The model performance varies greatly from month to month for all four experiments, but both the experiments involving KED (Aniso-KED and Iso-KED) outperform those using OK in all the months. It is evident that by allowing the elevation as a covariate in the kriging interpolation for prediction at each grid point, the amount of precipitation is considerably improved in most months.

The median for both the KED experiments is overestimated in the months February, March, May and September and in the rest of the months, the median is underestimated. The IQR in Figure 16 shows the inter-annual variability (IAV) but the figure also shows the intra-seasonal and inter-seasonal variability. The model reproduces the intra-seasonal and inter-seasonal variability very well in all the four experiments but in general better for the experiments when KED is employed. The IAV is better simulated in summer than in colder months. The best performance is found in July, when the variability is larger in the simulated data for all the four experimental set-ups compared to the observed data. For both the simulations using KED, the median in July is furthermore very close to the observed value. This is a remarkably good performance as WGs are typically criticized to have a tendency to underestimate the low frequency variability, as it is the case for most of the months in our precipitation generator. Conversely, our model – for all the four experiments – is not able to reproduce the larger IAVs in other months (particularly October). This also could be one of the reasons for the large underestimation in the precipitation amount over the study area in the autumn season as discussed in the previous sections. It is noteworthy that from October to December, the model performance is essentially similar in all the four experiments. This shows that regardless of the type of correlation structure and the interpolation method, the model is unable to capture the spatial distribution of precipitation in those months. In general, the model performance is better in warmer months than in colder months.

The differences in the observed and simulated median and IQR in each month for each of the four experiments are listed in Table 4. Overall, the experiments with KED outperform those using ordinary kriging and the Iso-KED combination is slightly superior to the fully anisotropic combination. Apparently, even for such a small region in complex terrain as the present study area, an isotropic covariance is adequate to reproduce the precipitation fields. In the observed data, anisotropy is indeed present but the difference between the variation in the correlations with horizontal distance compared to that in the vertical is very small. Hence, there is almost no difference in the performance of the model in isotropic and anisotropic formulation. The two experiments with OK not just underestimate the amount of precipitation but also lack the topographical influence on the simulated precipitation amount and rather produce smooth precipitation fields over the region (see Figure S3 in Supplement). However, the influence of topography must be included in the model, to realistically simulate the precipitation fields, as we showed here by considering KED interpolation.

As for the occurrence model, the covariance structure has a slight influence on the model performance. Figure 17 compares the areal dry and wet spells for all the four experiments with those observed. The performance of the anisotropic model is better for dry spells where Aniso-KED performs sightly better than Aniso-OK. Contrarily, the performance of the isotropic model is
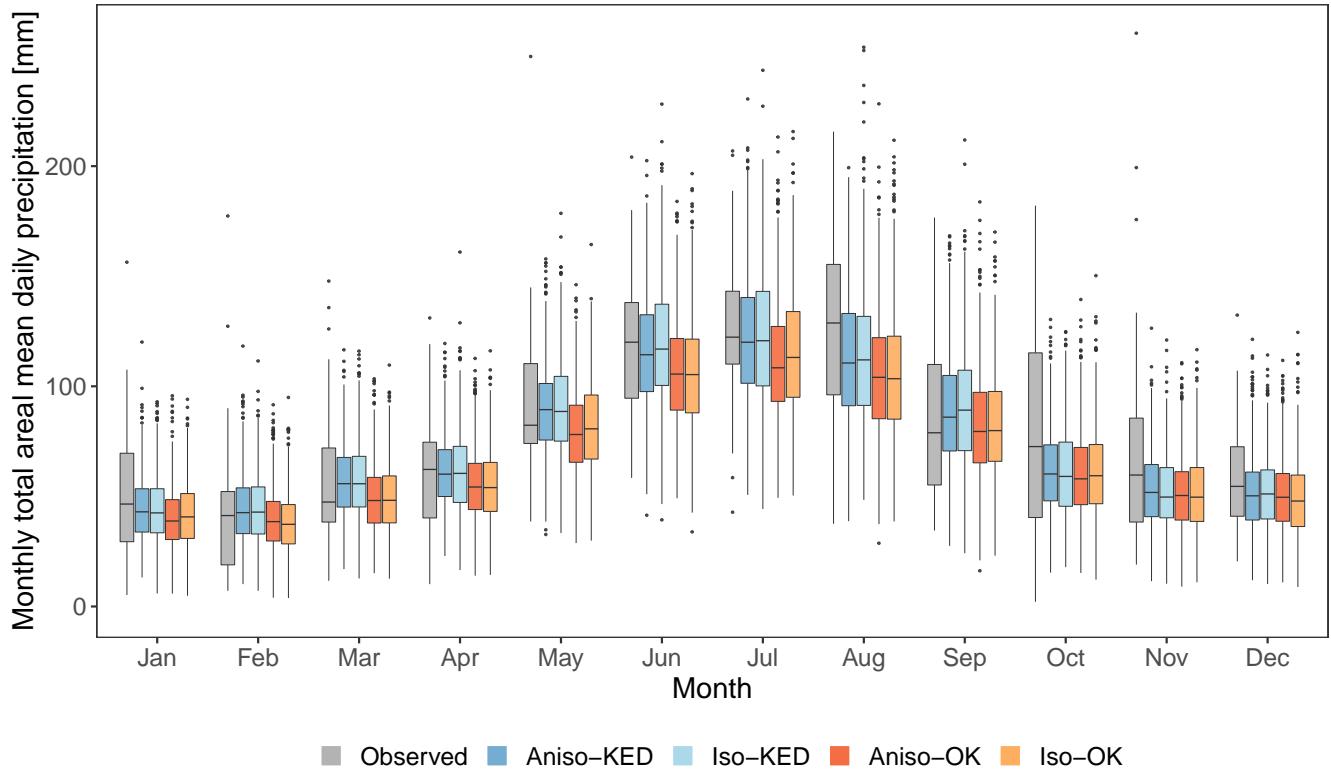
Figure 16. Monthly sum of areal mean daily precipitation [mm] for each month in APGD observed 30 years (1979–2008) and simulated 900 years (30 realisations of 30 years) of data for the four experiments: Aniso-KED, Iso-KED, Aniso-OK and Iso-OK (see text for specifications).

better for wet spells where Iso-KED performs slightly better than Iso-OK. For both areal dry and wet spells, KED interpolation adds little value to the simulation over using OK.

600

5   Discussion

In this study, we have tested two extensions of the original isotropic framework of Kleiber et al. (2012) for the simulation of precipitation fields. First, to include anisotropy in the covariance function and second, to apply KED in the interpolation of the parameters instead of OK. Our model adds substantial value over the original model in the highly complex mountainous region in the Austrian Alps. As a statistical downscaling tool, this model does not require any large computational resources. Obviously, compared to a single-site WG, it requires additional computing resources, but it is still very parsimonious and fast compared to dynamical downscaling models. It can be easily run on any high-end personal computer. In terms of model complexity, our model has only two parameters more than the original model which are the two range parameters in the verti-

Table 4. Difference (Simulated-Observed) in the median and interquartile range (IQR) of the observed and simulated monthly sum of areal mean daily precipitation [mm] in each month and each of the four experiments. The bold values indicate the best performance of the model in each month.

Month	Aniso-KED	Iso-KED	Aniso-OK	Iso-OK
	Difference in Median (IQR) [mm]			
Jan	**-3.54** (-20.52)	-4.01 (-20.16)	-7.70 (-22.06)	-5.84 (**-19.79**)
Feb	**1.33** (-12.52)	1.57 (**-11.93**)	-2.76 (-15.34)	-3.99 (-15.44)
Mar	8.35 (-11.10)	8.31 (**-10.66**)	**0.70** (-12.92)	0.78 (-12.32)
Apr	-2.15 (-13.13)	**-1.80** (**-8.94**)	-7.98 (-13.42)	-8.26 (-12.15)
May	7.08 (-10.58)	6.26 (**-6.86**)	-4.20 (-10.33)	**-1.63** (-7.19)
Jun	-5.69 (-8.56)	**-3.13** (**-6.55**)	-14.47 (-10.95)	-14.71 (-9.98)
Jul	-2.30 (5.87)	**-1.61**(9.83)	-13.91 (**0.95**)	-9.22 (5.84)
Aug	-18.18 (**-17.34**)	**-16.76** (-18.80)	-24.67 (-22.44)	-25.30 (-21.56)
Sep	7.10 (-20.44)	10.28 (**-18.20**)	**0.60** (-22.71)	1.00 (-23.03)
Oct	**-12.42** (-49.34)	-13.56 (**-45.64**)	-14.64 (-48.79)	-13.24 (-47.90)
Nov	**-7.90** (-23.54)	-10.00 (-24.40)	-9.30 (-25.25)	-10.04 (**-22.74**)
Dec	-4.29 (-9.79)	**-3.43** (-9.26)	-4.94 (-9.89)	-6.65 (**-8.19**)

cal direction due to inclusion of elevation in the covariance functions in the occurrence and amount model, respectively. And hence, in terms of computational time, the proposed model takes approximately 8% longer compared to the original model to complete the run which includes model fitting and simulation.

The major improvement in the results from our model comes from the KED interpolation, rather than the included anisotropy in the covariance structure. This suggests that there is no strong directional dependency in the precipitation simulation. Although there are minor differences in the model performance using the isotropic and anisotropic covariance functions, it can be concluded that an isotropic covariance function is sufficient even for small-scale topographic variability as in the present study in the European Alps. However, the topographical details must be included in the interpolation of the parameters of the model. Similar results can be expected for complex terrain in other mountainous regions.

At individual locations with observations, the model satisfactorily reproduces various observed statistics and the overall distribution of precipitation. The model is also able to capture spatial and temporal variability over the entire region reasonably well. It is capable to simulate the dry day statistics over the whole region very well, while for the wet day statistics, there is an underestimation observed. The frequency of areal dry spells of one or two days is strongly overestimated. The model uses previous day's occurrence as a covariate, which at an individual location creates a first-order two-state Markov chain. Dabhi et al. (2021) applied a first-order two-state Markov chain model at stations covering different climate zones in Europe and
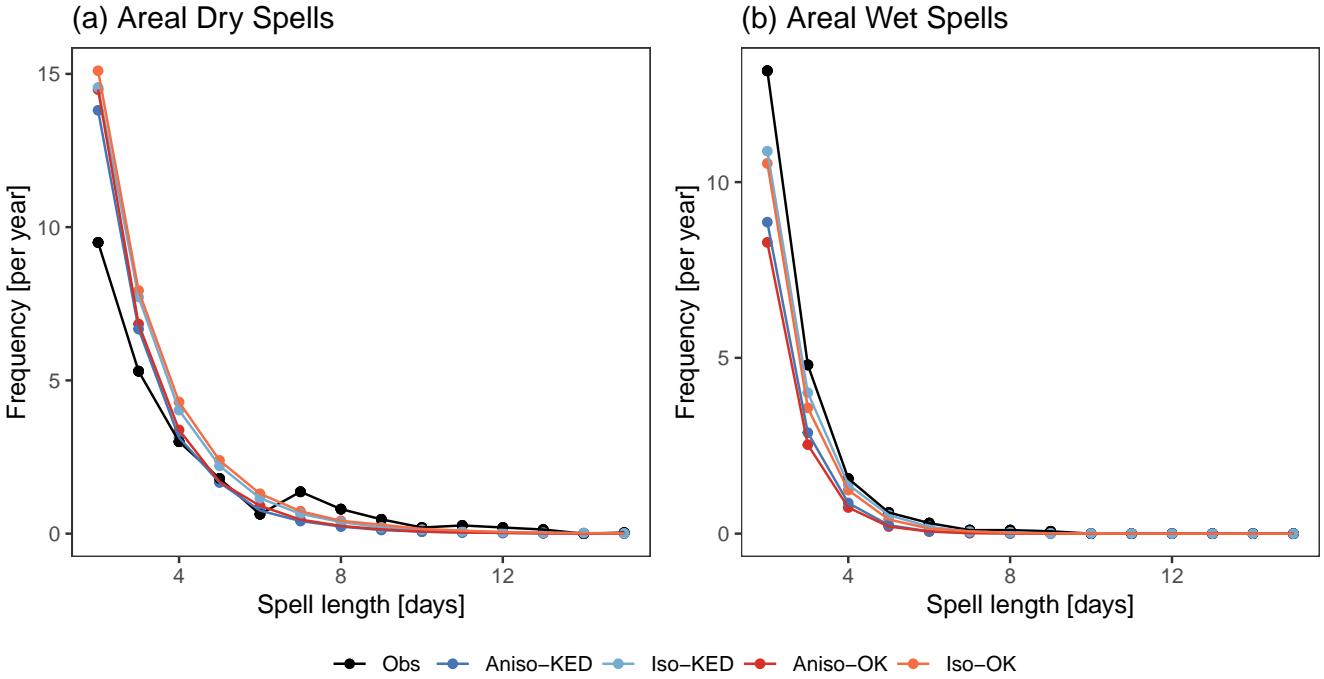
Figure 17. Frequency of areal spells [per year] of different length in each of the four experiments with 900 years of simulated data and in APGD observed 30 years (1979–2008) data: for (a) areal dry spells, and (b) areal wet spells. For the simulated dry and wet spells, their frequency is determined for each of the 30-year simulations separately and averaged over the 30 realizations.

found that it has a tendency to overestimate dry spells. The first-order Markov chain uses occurrence from only one day in the past, there may, however, be longer lasting correlations present in the data. Considering the occurrence from two or more days in the past, i.e. forming a second-order or higher order Markov chain at individual locations may potentially improve the results for both dry and wet spells. For example, Wilson Kemsley et al. (2021) studied the order of Markov chains in different climate regimes across the world where they showed that the third-order model reproduces observed dry spell distributions the best. Alternatively, by allowing other meteorological variables such as windspeed, humidity as covariates in the GLM can also improve the results (Ataharul Islam and Chowdhury (2006)).

The model captures the month-to-month variability in the monthly sum of precipitation very well which is due to the time-dependent harmonics of sine and cosine as covariates in the modeled spatial covariance structure. However, the inter-annual variability is largely underestimated mainly for the colder months. Even if we adopt the NAOI as a covariate to alleviate the often-discussed problem of overdispersion in this type of models, the overdispersion remains an issue. One reason is the tendency of the model to underestimate the large daily precipitation amounts. This is because the model generates the precipitation amount using a transformed Gaussian process which reduces to a gamma distribution at individual locations. The gamma dis-

640   tribution is not a heavy-tailed distribution and is therefore not well suited to reproduce the heavy precipitation. However, Wilks (2009) also found the underestimation of heavy precipitation in his study using a mixture of two exponential distributions to account for both smaller and larger precipitation amounts. This is another commonly reported problem for precipitation generators (e.g. Wilks (1999), Furrer and Katz (2008)). Allowing heavy-tailed distributions alleviates the overdispersion (Serinaldi and Kilsby (2014)), but simulating spatial extremes along with successfully capturing the smaller amounts with such a simple

645   model is even more challenging. Another way to reduce overdispersion is by allowing a suitable covariate such as seasonal total precipitation in the GLM as shown by Kim et al. (2012) or by including seasonal dry/wet indicators as in Kim and Lee (2017). Verdin et al. (2018) modified the model of Kleiber et al. (2012) by allowing the domain averaged seasonal total precipitation as a covariate and showed that the inclusion of this covariate improved the simulation. However, their study was focused on flat terrain where it is promising to take the areal precipitation as a covariate for the whole domain. This may not be suitable

650   for mountainous terrain and rather location specific climate information might be more promising. In this study, we wanted to evaluate the model for its ability to reproduce the observed statistics at locations where no observations are available and for that reason we avoided allowing any location specific information so that the model is not conditioned upon the availability of the information at each grid point. We believe that, by allowing such gridded information as a covariate, the model performance could be improved.

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    In complex mountainous terrain individual stations can exhibit precipitation characteristics quite distinct from those of neighbouring (or more distant) stations with more typical characteristics. The station-by-station evaluation (Section 4.1) has revealed that the model cannot reproduce precipitation at these distinct stations (e.g., Prutz in Figures 2–7). For the gridded simulations we would have expected that these distinct stations might negatively influence their neighbouring stations. In contrast, how-

660   ever, the nearby stations have influenced the distinct stations by generating too strong correlations. Figure S4 (see Supplement) depicts the inter-station correlations in the precipitation occurrence among the 29 stations in the observed against the simulated data. The cloud of points having weak correlations in the observed data belong to the distinct stations, but in the generated data the correlations are very strong. The reason is the good density of the network of 29 stations in such a small region. This is also the reason that the largest discrepancy in the performance of both the occurrence and amount models amongst the 29

665   stations is found at those distinct stations, because at those stations the model generates the statistics of the nearby stations rather than reproducing their own observations. Thus, a distinct station in a region with dense observational network cannot be reproduced, but doesn't strongly deteriorate the overall performance in the reproduction of the spatial information. This finding is also conformed by the hold-out cross-validation. As discussed in Section 3.2, this is due to the choice of the same set of covariates. Using the same set of covariates is, however, a necessary restriction if the model is aimed at gridded output fields.

670   Reproducing the spatial statistics over the whole region and especially in mountainous region is indeed a challenging task. And with such a simple model, capturing the realistic spatio-temporal fields of precipitation is even more challenging. Despite of that, our model successfully captures many difficult statistics useful for climate change impact applications such as long spells of dry and wet days and areal monthly mean precipitation.

675 However, if a 'distinct station' is located in a data sparse area (such as St. Martin in our study area), it dominates the entire neighbouring region and destroys the spatial structure. Thus, for a spatial precipitation generator in complex terrain the stations should not only be selected according to data availability (and quality) but also based on their precipitation characteristics. If they have distinctly different precipitation characteristics from the majority of the stations in the region, they should not be included in the training data set and if one is explicitly interested in such a station, one should use a single-site approach.

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Another limitation of this model is its inability to realistically simulate autumn and winter precipitation. This is because there are systematic differences in the characteristics of weather types between various seasons. In autumn and winter, westerly currents are stronger and the associated precipitation patterns are more pronounced than during spring and summer. The precipitation pattern in winter is associated with dynamically active synoptic-scale weather systems (fronts and low-pressure systems

685 specifically from the North Atlantic Ocean and the Adriatic Sea) in combination with orographic enhancement whereas the one in summer is related to convective activity which is either embedded in frontal systems or generated locally. Our model does not account for the wind-influence on the precipitation. This could be the reason for the model being not capable to capture the spatio-temporal patterns in autumn and winter. The convective season in Austria usually starts in May and lasts till September and during these months the model successfully captures the spatio-temporal patterns.

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The covariance function used in the model is assumed to be stationary which may not be a realistic assumption. Detecting spatial non-stationarity and modeling it is beyond the scope of this article and will be explored in future research. However, it is possible that by considering a non-stationary covariance function (e.g. Paciorek (2003)), the model performance may be improved. Since the precipitation-topography relationship is dominant in mountainous regions, the results can also be improved

695 by including not just elevation but other variables like slope, aspect, latitude and longitude in kriging (Wotling et al. (2000)). Considering east-west and north-south gradients of the topography along with elevation in the KED also improves the results (Hiebl and Frei (2017)). Since the inclusion of elevation at 1 km resolution has improved the results, we believe that considering even higher spatial resolution would provide even better simulation of precipitation.

700 **6 Conclusion**

A multi-site gridded precipitation generator that provides high resolution two-dimensional fields of precipitation in complex terrain using historical observations from a network of meteorological stations is developed, implemented and evaluated. The precipitation generator is an extension of the original framework of Kleiber et al. (2012) which uses a stationary isotropic covariance structure. The original framework is based on a latent Gaussian process for the occurrence, and a transformed

705 Gaussian process for the amount of precipitation where gamma distributed random numbers are transformed to normally distributed random numbers. This framework considers the parameters of a Generalized Linear Model (GLM) as a realisation of a spatial Gaussian process which allows one to spatially interpolate the parameters using kriging. In this article, two extensions

to the original framework are proposed: i) by allowing anisotropy in the covariance structure, and ii) by allowing elevation as an external drift in kriging. The anisotropy is included in the model by taking into account the elevation difference in the stationary covariance function of both, the occurrence and amount models. Along with that, elevation is allowed as an auxiliary variable in the kriging equations for the interpolation of the parameters of both the occurrence and amount models.

The model is tested in a small region (about 100×100 km^2) with highly complex terrain in the European Alps where 29 observational stations with 30 years of data (1981–2010) are available. The test region comprises stations with elevation differences of about 2300 m. Thirty realizations of 30 years of synthetic gridded data at 1×1 km^2 resolution are generated to allow for a robust statistical assessment.

The main findings from this study can be summarized as follows:

– At individual stations where observations are available, the model reproduces the observed statistics realistically well, including annual cycles of daily probabilities of precipitation occurrence and monthly means of precipitation, dry and wet spells of different length and the overall distribution of precipitation amount.

– The model has a great capability to capture the spatio-temporal statistics in the complex terrain, which includes the spatial distribution of occurrence probabilities and amount, areal dry and wet spells of different length, monthly mean areal precipitation and monthly sum of areal daily mean precipitation.

– The proposed extension considerably improves the simulation of spatio-temporal fields of precipitation – mainly due to the incorporation of elevation in kriging.

– The use of an isotropic or anisotropic covariance function in the mountainous region is equally good with marginal trade-offs for some of the statistics.

– The performance of the model varies greatly from month to month – being best in summer and worst in autumn.

– Intra-seasonal and inter-seasonal variabilities are well reproduced, while inter-annual variability is largely underestimated in autumn and winter.

– At a few of the 29 stations, where the observed precipitation statistics, and in particular their seasonality were distinctly different from all the other stations, the model performance is markedly compromised.

– The underestimation of large amounts of precipitation remains a problem.

Reproducing the spatio-temporal fields of precipitation in a region characterized by complex terrain like the Alps is a challenging task especially at locations where no observations are available. However, this is an essential requirement for hydrological modeling as hydrological models are driven by spatially and temporally coherent precipitation data. The proposed model can

respond to this need to some extent, nevertheless further improvement is required, as discussed in the article, to employ the
740   model for downscaling purpose.

Data availability. The data used in this study are available from the corresponding author upon request.

Author contributions. All the co-authors contributed to developing the idea. HD developed the code, prepared the input data, performed the simulations, did the analysis, prepared the visualizations and interpreted the results. MR supervised the work and MO provided vital suggestions. The manuscript was initially prepared by HD and finalized by all the co-authors.

745   *Competing interests.* The authors declare that they have no competing interest.

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