

Response to Reviewer #1

R1. This article looks at methods for partitioning changes in the rainfall-runoff relationship between vegetation changes, climatic variation, and non-stationarity in runoff generation.

Response: Thanks very much for your great efforts to assess our manuscript. We have studied your and reviewers' comments carefully and have made corrections/revisions as suggested. The point-to-point responses to the comments and revision are detailed below. In the following, we have detailed how these comments (in black) are raised and our responses (in deep sky blue).

R2. The three acronyms PCM, TTM and SBM are introduced and explained in parentheses by the second sentence of the Abstract, but not until the third paragraph of the Introduction and then after the acronyms are already used. Please use and explain them as soon as they are mentioned in the main text of the paper.

Response: We apologize for these oversights. Following your suggestion, we have added the introduction of three acronyms PCM, TTM and SBM in the main text of the paper (Line 52-53; Line 102; Line 334; Line 549-550).

Relevant text reads (Line 52-53): "Three commonly used methods for separating the impacts of vegetation change on catchment water yield are the paired-catchment method (PCM), the time-trend method (TTM), and the sensitivity-based method (SBM)."

(Line 102): "The monthly data were used for the paired-catchment method (PCM), time-trend method (TTM), ..."

(Line 334): "By using the paired-catchment method (PCM), time-trend method (TTM), and sensitivity-based method (SBM), ..."

(Line 549-550): "The paired-catchment method (PCM) is not invalidated by the non-stationarity induced by multiyear drought because of the role of the control catchment. However, the essence of the time-trend method (TTM) and the sensitivity-based method (SBM) is to separate runoff changes caused by non-stationary..."

R3. Is the Pettitt (1975) method used in this work? It is mentioned once (p4 L101) where the authors state that the Mann-Kendall test is used for ranking tests of non-stationarity then never again.

Response: Thanks for bringing this to our attention. In the revised manuscript, what kinds of time series had been analysed using these two methods and the purposes of the applications have been mentioned in section 3.1, section 4.1 and Table 1 (Line 119-126; Line 256; Line 321).

The Pettitt (1979) method is mainly used to solve two problems in this study. One is to analyse the abrupt change points of annual rainfall, runoff and potential evapotranspiration (refer to Table 1). The abrupt change point of annual runoff is used to divide the calibration period and the prediction period. And the other one is to analyse the abrupt change points of the slope of the rainfall-runoff cumulative curve (refer to Fig. 4). The existence of the abrupt change points of the slope of the rainfall-runoff cumulative curve means that the relationship between rainfall and runoff become non-stationary.

The Mann-Kendall test (Kendall, 1975; Mann, 1945) is used to analyse the change trend (increase or decrease) of annual rainfall, runoff, and potential evapotranspiration (refer to Table 1).

Relevant text reads (Line 119-126): “The Mann-Kendall test (Kendall, 1975; Mann, 1945) was used to detect the long-term trend of annual rainfall, runoff, and potential evapotranspiration and the Sen’s slope estimator (Sen, 1968) was used to obtain the degrees of above changes. If the Z value estimated by the Mann-Kendall test method is less than zero, it indicates a downward trend; on the contrary, if the Z value is greater than zero, it indicates an upward trend. β estimated by the Sen’s method represents the slope of the change trend. The Pettitt method (Pettitt, 1979) is a rank-based nonparametric statistical test method and is used to detect abrupt change points of annual rainfall, runoff, potential evapotranspiration records and the rainfall and runoff cumulative curves. The abrupt change point of annual runoff is used to divide the calibration period and the prediction period. The Mann-Kendall and Pettitt methods are the most frequently used statistical methods for identification of changes in hydro-meteorological data (Peng et al., 2020).”

(Line 256): “Two change points estimated by the Pettitt method occurred in December 1996 and January 2010 in the Red Hill catchment,…”

(Line 321): “^athe change point year estimated by the Pettitt method.”

R4. Question of equilibration between rainfall-runoff process within catchment and between paired catchments? The catchments must be small enough and the changes of a suitable scale that either both equilibrate quickly, or at the same rate so that cumulative fluxes still appear as a straight line.

Response: Thanks for your constructive comment. In the revised manuscript, the principle and the purposes of the applications of double mass curve have been described in more detail in section 3.1 (Line 128-133).

The double mass curve is the simplest, but is most intuitive and most widely used method to analyse the stationarity or multiyear evolution trend of hydrometeorological variables. It is true that cumulative fluxes can still appear as a straight line when both hydrometeorological variables equilibrate quickly, or at the same rate under the condition of stationary changes. It can be seen from the straight line of the period before October 2001 in Fig. 4 (b). However, when catchments are affected by non-stationary changes (such as multiyear drought, vegetation change, etc.), the amount and rate of changes to reach new equilibrium of both hydrometeorological variables (rainfall-runoff process within catchment or runoff-runoff process between paired catchments) will be different. For example, the same rainfall will lead to less runoff after multiyear drought and afforestation (refer to Fig. 4 (a)). These different changes will lead to the abrupt change points of the double mass curve, and there are different equilibrium states before and after the abrupt change points (the slope of curve during different periods is changed). The double mass curve in this study is used to explore the impact of multiyear drought and vegetation change on the non-stationary changes of rainfall-runoff relationship of paired catchments.

Relevant text reads (Line 128-133): “The DMCs plot the accumulated values of one variable against the accumulated values of another related variable for a concurrent period (Searcy and Hardison, 1960; Wang et al., 2013). It can still appear as a straight line when both hydrometeorological variables (rainfall and runoff) equilibrate quickly, or at the same rate under the condition of stationary changes. A break in the slope of the DMCs detected by the Pettitt method means that a change in the constant of proportionality between rainfall and runoff has occurred. The difference in the slope of the lines

indicates the shift in the rainfall-runoff relationship and the degree of change in the relation.”

R5. Do the authors think that having a single effect occur with reasonable gaps is necessary to analyse the data? In the case of Red Hill/Kileys Run the catchments were paired (hydrologically) well then had the afforestation and years of data, then the drought with years of data, then the post-drought conditions and again with years of data. Is there a risk if changing climate conditions inducing a non-stationary response would interfere with the land-use response if they occurred closely chronologically? Could a method determine the changes and separate them?

Response: Thanks for your constructive comments. We have added discussions regarding to your concerns in the revised manuscript (Line 470-483).

We agree that one of investigated effects (i.e. multiyear drought) was occurred part of the study period (2000 – 2009). We think it is still necessary to analysis the impacts of drought as we have demonstrated that multiyear drought has induced rainfall-runoff relationship changes in the control catchment. Negligence of such effect can result in significant differences in estimated effects of vegetation change as shown in Fig. 6, which is also reported by Zhao et al. (2010) using about 16 years data.

Only two years of rainfall was above the average after 2000 and drought lasted from 2000 to 2009 in the Red Hill paired catchments. After 2009 (post drought), the slope of double mass curve is still very close to that during the multiyear drought period. Thus the single effects of drought are considered to have similar effects as vegetation treatment and are evaluated between two periods (i.e. pre- and post-change periods).

We agree the risks of interferences between drought and land-use response are existed. Several studies have reported that not only land use types but also soil and catchment properties may lead to different effects of drought on runoff (Saft et al., 2015; van Dijk et al., 2013). Here, one of the assumptions of the new framework is that the effects of three factors (vegetation change, hydroclimatic non-stationarity and climate variability) are independent of each other. We have to make this assumption to enable us to separate three effects with the help of paired catchments. The sum of the contribution of three factors to runoff change is 111.4% in the Red Hill catchment and it is close to 100%, which proves that the assumption is basically reasonable and valid. However, considering the complexity of the interaction amongst different factors and the values that are difficult to quantify, the new framework cannot separate the interaction of three factors under the current experimental conditions and data. The research about how to estimate the interactions amongst different factors need to be carefully observed and investigated in the future.

Relevant text reads (Line 470-483): “Negligence of non-stationarity induced by multiyear drought can result in significant differences in estimated effects of vegetation change as shown in Fig. 6, which has also been reported by Zhao et al. (2010) using about 16 years data at the same site. In the new framework, effect of multiyear drought is estimated between pre- and post-change periods as that of vegetation change, although two years of rainfall is above the average after 2000. Because the slope of DMCs is still very close to that during the period after 2009 (post-drought) (see Fig. 4). Interactions between the impact of prolonged drought and that of land-use change may be existed. Several studies have reported that not only land use types but also soil and catchment properties may lead to different effects of drought on runoff (Saft et al., 2015; van Dijk et al., 2013). Here, one of the assumptions of the new framework is that the effects of three factors (vegetation change, hydroclimatic non-stationarity and climate variability) are independent of each other. We have to make this assumption to enable us

to separate three effects with the help of paired catchments. The sum of the contribution of three factors to runoff changes is 111.4% in the Red Hill catchment, which is close to 100% and proves that the assumption is basically reasonable and valid. Considering these complex and secondary interactions amongst different factors, the new framework cannot separate them under the current experimental design and available data. How to estimate the interactions amongst different factors need to be carefully observed and investigated in the future.”

References:

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Peng, T., Tian, H., Singh, V. P., Chen, M., Liu, J., Ma, H., and Wang, J.: Quantitative assessment of drivers of sediment load reduction in the Yangtze River basin, China, *J. Hydrol.*, 580, 124242, <https://doi.org/https://doi.org/10.1016/j.jhydrol.2019.124242>, 2020.

Pettitt, A. N.: A non-parametric approach to the change-point problem, *Journal of the Royal Statistical Society: Series C (Applied Statistics)*, 28, 126-135, <https://doi.org/10.2307/2346729>, 1979.

Saft, M., Western, A. W., Zhang, L., Peel, M. C., and Potter, N. J.: The influence of multiyear drought on the annual rainfall-runoff relationship: An Australian perspective, *Water Resour. Res.*, 51, 2444-2463, <https://doi.org/10.1002/2014WR015348>, 2015.

Searcy, J. K., and Hardison, C. H.: Double-mass Curves, United states government printing office, Washington, 65 pp., 1960.

Sen, P. K.: Estimates of the Regression Coefficient Based on Kendall's Tau, *J. Am. Stat. Assoc.*, 63, 1379-1389, <https://doi.org/10.1080/01621459.1968.10480934>, 1968.

van Dijk, A. I. J. M., Beck, H. E., Crosbie, R. S., de Jeu, R. A. M., Liu, Y. Y., Podger, G. M., Timbal, B., and Viney, N. R.: The Millennium Drought in southeast Australia (2001-2009): Natural and human causes and implications for water resources, ecosystems, economy, and society, *Water Resour. Res.*, 49, 1040-1057, <https://doi.org/10.1002/wrcr.20123>, 2013.

Wang, W., Shao, Q., Yang, T., Peng, S., Xing, W., Sun, F., and Luo, Y.: Quantitative assessment of the impact of climate variability and human activities on runoff changes: a case study in four catchments of the Haihe River basin, China, *Hydrol. Process.*, 27, 1158-1174, <https://doi.org/https://doi.org/10.1002/hyp.9299>, 2013.

Zhao, F., Zhang, L., Xu, Z., and Scott, D. F.: Evaluation of methods for estimating the effects of vegetation change and climate variability on streamflow, *Water Resour. Res.*, 46, <https://doi.org/10.1029/2009WR007702>, 2010.

Response to Reviewer #2

General comments

R1. In general it's clear that the authors improved the original manuscript (<https://doi.org/10.5194/hess-2021-5>). They did improved the flow of the storyline and the analysis of the control catchment. The manuscript has appropriate objectives and gives insight in the change in hydrological processes caused by the combination of land use and climate changes.

Response: Thanks very much for your great efforts to assess our manuscript. We have studied your and reviewers' comments carefully and have made corrections/revisions as suggested. The point-to-point responses to the comments and our plans for revision are listed below. In the following, we have detailed how these comments (in black) are raised and our responses (in deep sky blue).

R2. Although the assumptions and following conclusions are constructively fine and according to the applied methods, the arguments (especially in the abstract and conclusion) are not very clear. Although some comparative values per method are presented, for a reader it is not clear where the conclusions are based on.

Response: Thanks for bringing this to our attention. More specific descriptions about where the conclusions are based on have been mentioned in the new manuscript (Line 22-24; Line 289-290; Line 347-351; Line 355-357; Line 547-548; Line 553; Line 555-556; Line 556-557).

Relevant text reads (Line 22-24): "Based on the new framework, impacts of multiyear drought and climate variability on runoff of the control catchment (Kileys Run) were 87.2% and 12.8%, respectively. Impacts of afforestation, multiyear drought and climate variability on runoff of the treated catchment (Red Hill) were 32.8%, 54.7% and 23.9%, respectively."

(Line 289-290): "The runoff coefficient decreased by 87.8% and 63.3% during the drought period in the Red Hill and Kileys Run catchments, respectively."

(Line 347-351): "For the Kileys Run catchment, runoff change induced by multiyear drought ($\overline{\Delta Q_c^d}$) is -110.2 mm, and runoff change caused by climate variability ($\overline{\Delta Q_c^{clm}}$) by subtracting runoff change caused by multiyear drought from the total runoff changes ($\overline{\Delta Q_c^{total}} = -126.4$ mm) is -16.2 mm. Impacts of afforestation, multiyear drought and climate variability on runoff of the Red Hill catchment are 32.8%, 54.7% and 23.9%, respectively, and impacts of multiyear drought and climate variability on runoff of the Kileys Run catchment are 87.2% and 12.8%, respectively."

(Line 355-357): "Estimated impacts of afforestation on runoff decreased greatly from 93.5% to 38.8% (=93.5%-54.7%) calculated by the TTM and decreased greatly from 76.1% to 21.4% (=76.1%-54.7%) by the SBM."

(Line 547-548): "The runoff coefficient decreased by 87.8% and 63.3% during the drought period in the Red Hill and Kileys Run catchments, respectively."

(Line 553): "Estimated afforestation impacts were 32.8%, 93.5%, and 76.1% of total runoff changes by the PCM, TTM and SBM, respectively."

(Line 555-556): "Impacts of afforestation, multiyear drought, and climate variability on runoff of the

treated catchment (Red Hill) were 32.8%, 54.7% and 23.9%, respectively.”

(Line 556-557): “The contribution of vegetation change to runoff reduction using the three methods under the new framework become consistent (32.8%, 38.8% and 21.4%).”

Specific comments

R3. The abstract ends with “paired-catchment method is proven to be still the most reliable method even the control catchment experienced climate-induced shift in rainfall-runoff relationship”, but within the abstract no information on the control catchments are presented. I would suggest to give actual numbers so that the reader is able to draw this conclusion himself.

Response: Thanks for your suggestions. We have added the results about the control catchment in abstract section (Line 22-23).

Relevant text reads (Line 22-23): “Impacts of multiyear drought and climate variability on runoff of the control catchment (Kileys Run) were 87.2% and 12.8%, respectively.”

R4. In my opinion it’s a missed opportunity that the majority of input by the previous reviewers and especially the replies by the authors are not processed within this manuscript.

e. Response to reviewer 1 (<https://doi.org/10.5194/hess-2021-5-AC1>): “During 1988 the uppermost 50 ha of Red Hill was planted top. radiata, with the remaining 145 ha planted to p. radiata in April 1989. During 2003 the plantations in Red Hill catchment were thinned to remove pulpwood and to promote the growth of sawlogs in the remaining stands. During the prolonged drought period, no trees died in the treated catchment. The change of annual PET can be seen in Figure 8 (b) (Page 42). PET showed an insignificant ($p\text{-value}>0.1$) increasing trend of 3.5 mm year⁻¹. PET initially decreased before 1996 and then increased after 1996. The mean annual PET during the period of 1990-1996, 1997-2009, 2010-2015 and 1997-2015 are 1168 mm, 1262 mm, 1186 mm and 1238 mm, respectively. Compared with the period of 1990-1996 and 2010-2015, the mean annual PET during the period of 1997-2009 (the prolonged drought occurred) increased by 94 mm and 76 mm, respectively. Compared with the period of 1990-1996, the mean annual PET during the period of 1997-2015 increased by 70 mm. It is consistent with the cognition that afforestation and drought can make PET increase.”

e. reviewer 2 (<https://doi.org/10.5194/hess-2021-5-AC2>), which you may use in the introduction and/or discussion: “Considering the influence of prolonged drought on the rainfall-runoff relationship in the treated catchment, the result of the paired catchment method is closest to the real runoff change caused by vegetation change. Because the control catchment indirect eliminate the influence of prolonged drought and climate variability on the treated catchment under the assumption that the response of the two catchments to prolonged drought is similar. Interannual changes in watershed storage occur primarily in soil water and shallow groundwater, pools that are often hydrologically active at time scales shorter than 1 year (Sayama et al., 2011). Rice and Emanuel (2019) indicates that down-regulation of transpiration and inhibition of hydrologic connectivity by forest vegetation represent two important negative feedback processes that can avert large losses in soil water or plant-accessible groundwater during dry periods. In doing so, this feedback mechanism has the potential to reinforce steady-state (or near-steady-state) conditions in dry conditions. So, the sensitivity-based method may

be less affected by the prolonged drought because it is used in the forest and the time scale of PET and P data used in this method is annual scale. Runoff changes calculated by the sensitivity-based method are induced by climate variability.”

Response: We apologize for these oversights, and thanks very much for your reminder. We have added more detailed information about the characteristics of paired catchments, data and the history of vegetation change treatment in section 2 (Line 84-109). The analysis of changes in potential evapotranspiration and rainfall have been mentioned in section 5.1 (Fig. 7; Line 394-402). More specific analysis of the application of traditional methods have been added in section 5.1 (Line 376-383; Line 389-409).

Relevant text reads (Line 84-109): “The Red Hill catchment (1.95 km²) and the Kileys Run catchment (1.35 km²) were paired catchments located 23 km northeast of Tumut and 100 km west of Canberra in New South Wales, Australia (35.322°S, 149.137°E) (Fig. 1). The catchments are adjacent, and the soil texture, topographic characteristics, and climatic conditions are similar. The altitude of the two catchments ranges from 590 m to 835 m above sea level. The slope in the lower part of catchments is mostly gentle, and gradually increases towards the ridge in a convex form. Geology of the Red Hill catchment is predominately Young granodiorite, while the Kileys Run catchment is dominated by Alkali diorite. Valley floor, midslope yellow duplex, shallow red soils and upslope red duplex are four main soil types in these two catchments. Upslope red duplex soils has the highest saturated hydraulic conductivities and valley floor soils has the lowest saturated hydraulic conductivities (Major et al., 1998). The climate of these two catchments is temperate with highly variable and winter-dominated rainfall. In 1988, *P. radiata* was planted in the Red Hill catchment (0.5 km²), and the remainder (1.45 km²) was planted in April 1989. By 1997, pine occupied 78% of the Red Hill catchment. During multiyear drought period, no trees died in the treated catchment (Bren et al., 2006). The neighboring catchment (Kileys Run) was the control catchment, which has been maintained as a grazed pasture control over the entire observation period (Webb and Kathuria, 2012). Daily rainfall and runoff from these two catchments were collected during the period of 1990–2015. The daily rainfall was measured by tipping bucket rain gauges had been located at catchment outlet and the daily runoff was measured by a flat-v style crump weir at a gauging station at the outlet of each catchment (Major et al., 1998). Mean annual rainfall and mean annual runoff of the Red Hill catchment were 817 mm and 75 mm, respectively, during the study period. Mean annual rainfall and runoff were 817 mm and 161 mm, respectively, in the Kileys Run catchment over the same period. Monthly potential evapotranspiration (PET) records were obtained from the SILO Data (www.longpaddock.qld.gov.au/silo/point-data/). The daily data were only used for the analysis of flow duration curves (FDCs). The monthly data were used for the paired-catchment method (PCM), time-trend method (TTM), the new framework and the analysis of double mass curves (DMCs). The annual data are used in the sensitivity-based method (SBM). Figure 2 shows the change of rainfall anomaly (%) in the Kileys Run and Red Hill catchments. Rainfall anomaly (%) is defined as the percentage deviation of annual rainfall to mean annual rainfall. It can be seen that three-year moving average of the rainfall anomaly (the black line) is lower than zero from 2000 to 2009. According to the method of determining multiyear drought period proposed by Saft et al. (2015), two catchments experienced prolonged drought lasted 10 years from 2000 to 2009 and this coincided with the period of the Millennium Drought of Australia (van Dijk et al., 2013). The minimum measured annual rainfall from 1990 to 2015 were 388.6 mm.”

(Line 394-402): “Figure 7 shows the change of annual *P* and PET. Over the entire study period from 1990 to 2015, *P* showed an insignificant decreasing trend of 3.4 mm year⁻¹ ($p > 0.1$) and PET showed an insignificant increasing trend of 3.5 mm year⁻¹ ($p > 0.1$). Both *P* and PET decreased before 1996 and then

increased after 1996. The rates of increase for annual P and PET during 1997-2015 were 12.0 mm year⁻¹ and 2.6 mm year⁻¹, respectively, and the contributions of P and PET to runoff changes caused by climate variability were -22.0 mm and -11.0 mm, respectively. The mean annual PET during the period of 1990-1996, 1997-2009, 2010-2015 and 1997-2015 were 1168 mm, 1262 mm, 1186 mm and 1238 mm, respectively. Compared with the period of 1990-1996 and 2010-2015, the mean annual PET during the period of 1997-2009 (the multiyear drought occurred) increased by 94 mm and 76 mm, respectively. Compared with the period of 1990-1996, the mean annual PET during the period of 1997-2015 increased by 70 mm. It was consistent with the cognition that afforestation and drought can make PET increase.”

(Line 376-383): “At Red Hill experiment site, non-stationary changes of the treated catchment are caused by both vegetation change and multiyear drought, and stationary changes are caused by climate variability. Non-stationary changes of the control catchment are only caused by multiyear drought, and stationary changes are only caused by climate variability. According to the paradigm of PCM, shift in the rainfall-runoff relationship separated from the runoff correlation between the treated and control catchments should be caused only by the treatment of the treated catchment and effects of any other drivers can induce either stationary or non-stationary changes should be eliminated by making use of the control catchment. Therefore, the PCM is still the most reliable method compared with other methods and separated effect by the PCM is only caused by vegetation change (i.e., afforestation).”

(Line 389-409): “The SBM is sourced from the Budyko framework (Budyko, 1974). It assumes that steady state of catchment water balance is fundamentally determined by water input (represented by precipitation) and energy demand (represented by potential evapotranspiration) and transition from one steady state to another without any change in catchment properties should moving on the Budyko curve (Roderick and Farquhar, 2011; Sun et al., 2014; Wang et al., 2021). Therefore, stationary changes driven by climate variability during post-treatment period can be separated by sensitivity of runoff to P and PET established during the pre-treatment period. Figure 7 shows the change of annual P and PET. Over the entire study period from 1990 to 2015, P showed an insignificant decreasing trend of 3.4 mm year⁻¹ ($p>0.1$) and PET showed an insignificant increasing trend of 3.5 mm year⁻¹ ($p>0.1$). Both P and PET decreased before 1996 and then increased after 1996. The rates of increase for annual P and PET during 1997-2015 were 12.0 mm year⁻¹ and 2.6 mm year⁻¹, respectively, and the contributions of P and PET to runoff changes caused by climate variability were -22.0 mm and -11.0 mm, respectively. The mean annual PET during the period of 1990-1996, 1997-2009, 2010-2015 and 1997-2015 were 1168 mm, 1262 mm, 1186 mm and 1238 mm, respectively. Compared with the period of 1990-1996 and 2010-2015, the mean annual PET during the period of 1997-2009 (the multiyear drought occurred) increased by 94 mm and 76 mm, respectively. Compared with the period of 1990-1996, the mean annual PET during the period of 1997-2015 increased by 70 mm. It was consistent with the cognition that afforestation and drought can make PET increase. The result estimated by the SBM is the impact of climate variability (without changing the catchment characters/non-stationary changes in the rainfall-runoff relationship) on runoff, that is, $(\overline{\Delta Q_t^{clim}})$. It ignored the impact of multiyear drought on runoff, which have proved to cause non-stationary changes (Kinal and Stoneman, 2012; Peterson et al., 2021; Saft et al., 2016; van Dijk et al., 2013), which may violate the assumptions of the SBM. Estimated change of the SBM, which is close to the result of the TTM (see Table 2) and is about 2.3 times greater than the result of the PCM, includes the non-stationary changes not only caused by vegetation change but also caused by multiyear drought.”

R5. For some statements I suggest to add some additional more recent papers.

e. at line 33.

e. at lines 52-53 “in many catchments around the world” two papers are cited one from 1980 and one from 2005, may be implying that these methods are not used that often anymore?

Response: Thanks for your suggestions. Changes have been made as suggested. (Line 36-38; Line 56-58, see references below).

Relevant text reads (Line 36-38): “However, separating the effects of vegetation change and climate variability on runoff remains a great challenge due to the complex interactions between climate variability and vegetation change (Cavalcante et al., 2019; Jones et al., 2006; Zhang et al., 2021).”

(Line 56-58): “This method has been applied in many paired catchments around the world to provide fundamental understanding and knowledge for water resource management under vegetation change (Brown et al., 2005; Stoof et al., 2012; Van Loon et al., 2019).”

Cavalcante, R. B. L., Pontes, P. R. M., Souza Filho, P. W. M., and Souza, E. B.: Opposite Effects of Climate and Land Use Changes on the Annual Water Balance in the Amazon Arc of Deforestation, *Water Resour. Res.*, 55, 3092-3106, <https://doi.org/10.1029/2019WR025083>, 2019.

Zhang, J., Zhang, Y., Sun, G., Song, C., Dannenberg, M. P., Li, J., Liu, N., Zhang, K., Zhang, Q., and Hao, L.: Vegetation greening weakened the capacity of water supply to China's South-to-North Water Diversion Project, *Hydrol. Earth Syst. Sci.*, 25, 5623-5640, <https://doi.org/10.5194/hess-25-5623-2021>, 2021.

Stoof, C. R., Vervoort, R. W., Iwema, J., van den Elsen, E., Ferreira, A. J. D., and Ritsema, C. J.: Hydrological response of a small catchment burned by experimental fire, *Hydrol. Earth Syst. Sci.*, 16, 267-285, <https://doi.org/10.5194/hess-16-267-2012>, 2012.

Van Loon, A. F., Rangelcroft, S., Coxon, G., Breña Naranjo, J. A., Van Ogtrop, F., and Van Lanen, H. A. J.: Using paired catchments to quantify the human influence on hydrological droughts, *Hydrol. Earth Syst. Sci.*, 23, 1725-1739, <https://doi.org/10.5194/hess-23-1725-2019>, 2019.

R6. Lines 79-93: The data and location description are very brief. I suggest to give some more extended information. i.e. It can be seen that Kileys Run experienced a multiyear drought that lasted ... with the period of the Millennium Drought”, explain why? What is lowest amount of rainfall measured?

Response: Thanks for your advice. We have added more extended information about paired catchments and the determination of multiyear drought in the revised manuscript (Line 84-95; Line 96-109).

Relevant text reads (Line 84-95): “The Red Hill catchment (1.95 km²) and the Kileys Run catchment (1.35 km²) were paired catchments located 23 km northeast of Tumut and 100 km west of Canberra in New South Wales, Australia (35.322°S, 149.137°E) (Fig. 1). The catchments are adjacent, and the soil texture, topographic characteristics, and climatic conditions are similar. The altitude of the two catchments ranges from 590 m to 835 m above sea level. The slope in the lower part of catchments is mostly gentle, and gradually increases towards the ridge in a convex form. Geology of the Red Hill catchment is predominately Young granodiorite, while the Kileys Run catchment is dominated by Alkali diorite. Valley floor, midslope yellow duplex, shallow red soils and upslope red duplex are four main soil types in these two catchments. Upslope red duplex soils has the highest saturated hydraulic conductivities and valley

floor soils has the lowest saturated hydraulic conductivities (Major et al., 1998). The climate of these two catchments is temperate with highly variable and winter-dominated rainfall. In 1988, *P. radiata* was planted in the Red Hill catchment (0.5 km²), and the remainder (1.45 km²) was planted in April 1989. By 1997, pine occupied 78% of the Red Hill catchment. During multiyear drought period, no trees died in the treated catchment (Bren et al., 2006). The neighboring catchment (Kileys Run) was the control catchment, which has been maintained as a grazed pasture control over the entire observation period (Webb and Kathuria, 2012).”

(Line 96-109): “Daily rainfall and runoff from these two catchments were collected during the period of 1990–2015. The daily rainfall was measured by tipping bucket rain gauges had been located at catchment outlet and the daily runoff was measured by a flat-v style crump weir at a gauging station at the outlet of each catchment (Major et al., 1998). Mean annual rainfall and mean annual runoff of the Red Hill catchment were 817 mm and 75 mm, respectively, during the study period. Mean annual rainfall and runoff were 817 mm and 161 mm, respectively, in the Kileys Run catchment over the same period. Monthly potential evapotranspiration (PET) records were obtained from the SILO Data (www.longpaddock.qld.gov.au/silo/point-data/). The daily data were only used for the analysis of flow duration curves (FDCs). The monthly data were used for the paired-catchment method (PCM), time-trend method (TTM), the new framework and the analysis of double mass curves (DMCs). The annual data are used in the sensitivity-based method (SBM). Figure 2 shows the change of rainfall anomaly (%) in the Kileys Run and Red Hill catchments. Rainfall anomaly (%) is defined as the percentage deviation of annual rainfall to mean manual rainfall. It can be seen that three-year moving average of the rainfall anomaly (the black line) is lower than zero from 2000 to 2009. According to the method of determining multiyear drought period proposed by Saft et al. (2015), two catchments experienced prolonged drought lasted 10 years from 2000 to 2009 and this coincided with the period of the Millennium Drought of Australia (van Dijk et al., 2013). The minimum measured annual rainfall from 1990 to 2015 were 388.6 mm.”

R7. Lines 346-357, the paragraph about “Multiyear drought induced changes in the rainfall-runoff relationship” is short and doesn’t discuss the subject. I suggest to compare your results with other locations where multiyear drought led to changes in the rainfall-runoff relationship.

Response: Thanks for your suggestions. We have added a few sentences to discussion the subject that multiyear drought induced changes in the rainfall-runoff relationship in section 5.2 (Line 414-427).

It can be seen that multiyear drought has led to significant changes in the rainfall-runoff relationship in Fig. 4 and 5, which is similar to the significant downward shift of rainfall-runoff regression lines in basins in southeast Australia, the United States and China (Avanzi et al., 2020; Saft et al., 2015; Tian et al., 2018). In this study, the runoff coefficient decreased by 87.8% and 63.3% during drought period in the Red Hill and Kileys Run catchment, respectively. The latter was close to the decrease of runoff coefficient of 65.8% in Texas caused by extreme drought (Allen et al., 2011). T Runoff coefficient decrease of the Red Hill catchment was higher than that of the Kileys Run catchment because runoff of the Red Hill catchment was also affected by afforestation, which can increased annual evaporation and decreased streamflow (Cheng et al., 2017; Hoek Van Dijke et al., 2022; Wang-Erlandsson et al., 2018). This will lead to more runoff reduction than predicted based on the rainfall-runoff relationship established in pre-drought period, and ignoring the impact of non-stationary change may cause large errors in the results (Zhao et al., 2010). Compared with the line during drought period, the rainfall-runoff regression line moved up after the drought due to heavy rainfall in 2010, but it did not completely return to the state before the

drought. Peterson et al. (2021) suggested that these changes may be due to water loss from increased transpiration during drought period.

Relevant text reads (Line 414-427): “According to the results in session 4.2, multiyear drought has led to shift in the rainfall-runoff relationship of paired catchments, which is similar to the significant downward shift of rainfall-runoff regression lines in basins in southeast Australia, the United States and China (Avanzi et al., 2020; Saft et al., 2015; Tian et al., 2018), and the increases in zero flow days with low flows being more affected than high flows of FDCs in 10 catchments from southeastern Australia, New Zealand and South Africa (Lane et al., 2005). In this study, the runoff coefficient decreased by 87.8% and 63.3% during the drought period in the Red Hill and Kileys Run catchments, respectively. The latter was close to the decrease of runoff coefficient of 65.8% in Texas caused by extreme drought (Allen et al., 2011). Runoff coefficient decrease of the Red Hill catchment was higher than that of the Kileys Run catchment because runoff of the Red Hill catchment was also affected by afforestation, which can increased annual evaporation and decreased streamflow (Cheng et al., 2017; Hoek Van Dijke et al., 2022; Wang-Erlandsson et al., 2018). Multiyear drought can lead to more runoff reduction than predicted based on the rainfall-runoff relationship established during pre-drought period as ignoring the impact of non-stationary changes may cause large errors in the results (Zhao et al., 2010). Compared with the line during drought period, the rainfall-runoff regression line moved up after multiyear drought due to heavy rainfall of 2010, but it did not return to the state before multiyear drought. Peterson et al. (2021) suggested that these changes may be due to severe water loss from transpiration during drought period.”

R8. In this paragraph you do mention a change in physical processes (runoff, soil moisture and evapotranspiration?), but not really compared with studies elsewhere. Do you have any specific evidence available about changes processes?

Response: Thanks for your constructive comments. Changes have been made as suggested. More comparisons with studies elsewhere and descriptions of changes in annual lowest seven-day flow (it reflects the storage state of groundwater to a certain extent) have been added in section 5.2 (Line 428-441). Because only rainfall, runoff and potential evapotranspiration data can be obtained, the conjecture about the reasons is based on the analogy of similar catchments also affected by multiyear drought in Australia.

For precipitation input, inter-annual rainfall variability decreased and high rainfall years were missing during the drought period in Fig. 2. Similar changes also occurred in 124 watersheds in Australia during drought period (Saft et al., 2015). The reduction of rainfall input reduces runoff at the source. For runoff process, in the Kileys Run and Red Hill catchments, rainfall primarily occurs in autumn and winter, less rainfall in autumn may result in lower antecedent soil moisture, which means more precipitation were used to replenish the soil water deficit in winter (Fig. 8). The changes of the monthly averages of rainfall and runoff were very similar to that of Murray Darling Basin, where rainfall-runoff relationship has changed caused by multiyear drought (Potter et al., 2010). As a result, runoff in winter during drought period was less than that during pre-drought period and the decrease of rainfall in next spring further aggravated runoff reduction. The decrease of GRACE satellite-observed average monthly terrestrial water storage and estimated groundwater storage in Murray–Darling Basin may support the above speculation about runoff reduction (van Dijk et al., 2013). The decline in groundwater levels may also be the reason for runoff reduction. Decline in precipitation usually resulted in a decline in groundwater levels (Peters et al., 2003), and may cause the connection between groundwater and surface water to be disrupted (Kinal and Stoneman, 2012). Brutsaert (2008) demonstrated that annual lowest seven-day

flow can be used to indicate the change of ground water storage in the absence of observations of groundwater level. The annual lowest seven-day flow in the Kileys Run catchment generally declined from 1990 to 1999, and was reduced to zero from 2001 to 2010.

Relevant text reads (Line 428-441): “Inter-annual rainfall variability decreased and high rainfall years were missing during the drought period (see Fig. 2). Similar changes were also reported in 124 watersheds in Australia during the drought period (Saft et al., 2015). The reduction of rainfall reduced runoff. In the Kileys Run and Red Hill catchments, rainfall primarily occurred in autumn and winter, less rainfall in autumn may resulted in lower antecedent soil moisture, which means more precipitation were used to replenish the soil water deficit in winter (Fig. 8). As a result, runoff in winter during drought period was less than that during pre-drought period and the decrease of rainfall in next spring further aggravated runoff reduction. It was consistent with less runoff during the second period under the same rainfall in Fig. 4. The decrease of GRACE satellite-observed average monthly terrestrial water storage and estimated groundwater storage in Murray–Darling Basin may support the above speculation about runoff reduction (van Dijk et al., 2013). The decline in groundwater levels may also be the reason for runoff reduction. Decline in precipitation usually resulted in a decline in groundwater levels (Peters et al., 2003), and may cause disconnection between groundwater and surface water (Kinal and Stoneman, 2012). Brutsaert (2008) demonstrated that annual lowest seven-day flow can be used indirectly to indicate the change of ground water storage. The annual lowest seven-day flow in the Kileys Run catchment generally declined from 1990 to 1999, and was reduced further to zero from 2001 to 2010, suggesting ground water storage have dried up for a long time during multiyear drought.”

R9. In addition to that, add some references about global and local knowledge about land use changes and their effects, such as the infiltration trade-off hypothesis (i.e. Bruijnzeel, 1989) and regional water availability caused by tree restoration such as (i.e. Hoek van Dijke, et al. 2022).

Response: Thanks for your advice. Changes have been made as suggested. (Line 420-422, see references below).

Relevant text reads (Line 420-422): “Runoff coefficient decrease of the Red Hill catchment was higher than that of the Kileys Run catchment because runoff of the Red Hill catchment was also affected by afforestation, which can increase annual evaporation and decrease streamflow (Bruijnzeel, 1989; Cheng et al., 2017; Hoek Van Dijke et al., 2022).”

Bruijnzeel, L. A.: Forestaion and dry season flow in the tropics: a closer look, *J. Trop. For. Sci.*, 1, 229-243, 1989.

Hoek Van Dijke, A. J., Herold, M., Mallick, K., Benedict, I., Machwitz, M., Schlerf, M., Pranindita, A., Theeuwens, J. J. E., Bastin, J., and Teuling, A. J.: Shifts in regional water availability due to global tree restoration, *Nat. Geosci.*, 15, 363-368, <https://doi.org/10.1038/s41561-022-00935-0>, 2022.

Cheng, L., Zhang, L., Chiew, F. H. S., Canadell, J. G., Zhao, F., Wang, Y., Hu, X., and Lin, K.: Quantifying the impacts of vegetation changes on catchment storage-discharge dynamics using paired-catchment data, *Water Resour. Res.*, 53, 5963-5979, <https://doi.org/10.1002/2017WR020600>, 2017.

R10. Lines 412-426, I suggest to re-introduce abbreviations.

Response: Thanks for your advice. The words have been modified to “paired-catchment method (PCM)”,

“time-trend method (TTM)” and “sensitivity-based method (SBM)” (Line 548; Line 549).

R11. Lines 417-418, conclusion. Add numbers so the reader is able to “agree” with your conclusion.

Response: Thanks for your suggestions. We have added a few sentences to support our conclusion in the new manuscript (Line 547-548; Line 553; Line 555-556; Line 556-557).

Relevant text reads (Line 547-548): “The runoff coefficient decreased by 87.8% and 63.3% during the drought period in the Red Hill and Kileys Run catchments, respectively.”

(Line 553): “Estimated afforestation impacts were 32.8%, 93.5%, and 76.1% of total runoff changes by the PCM, TTM and SBM, respectively.”

(Line 555-556): “Impacts of afforestation, multiyear drought, and climate variability on runoff of the treated catchment (Red Hill) were 32.8%, 54.7% and 23.9%, respectively.”

(Line 556-557): “The contribution of vegetation change to runoff reduction using the three methods under the new framework become consistent (32.8%, 38.8% and 21.4%).”

R12. Idem line 420.

Response: Thanks for your suggestions. The lines have been modified to “The contribution of vegetation change to runoff reduction using the three methods under the new framework become consistent (32.8%, 38.8% and 21.4%).” (Line 556-557).

Technical corrections

R13. Line 82, what about Red Hill?

Response: Thanks for bringing this to our attention and more information about the Red Hill catchment have been mentioned in the new manuscript (Line 88-91).

Geology of the Red Hill catchment is predominately Young granodiorite, while the Kileys Run catchment also has Young granodiorite but is dominated by Alkali diorite. The ridges and upper slopes of each catchment have predominantly given rise to red duplex soils that comprise dark brown organic loam and silty loam overlying light to medium reddish-brown clay. In Red Hill catchment lower slopes comprise sandy soils consisting of dark grey-brown organic loam/silty loam overlying grey-brown bleached and heavily cemented loamy silt grading to yellow-brown or grey-brown mottled medium-heavy clay at depths > 45 cm. This soil type is absent from the Kileys Run catchment and mean saturated hydraulic conductivities were 42.84 cm/day, 11.12 cm/day and 6.62 cm/day at depths of 20–50 cm, 50–80 cm and 80–100 cm, respectively.

Relevant text reads (Line 88-91): “Geology of the Red Hill catchment is predominately Young granodiorite, while the Kileys Run catchment is dominated by Alkali diorite. Valley floor, midslope yellow duplex, shallow red soils and upslope red duplex are four main soil types in these two catchments. Upslope red duplex soils has the highest saturated hydraulic conductivities and valley floor soils has the lowest saturated hydraulic conductivities (Major et al., 1998).”

R14. Line 87, it would be nice to apply your analysis with data up to 2020?

Response: Thanks for bringing this to our attention. Unfortunately, we do not have measured runoff and rainfall data up to 2020. We can obtain the annual grid rainfall from 1990 to 2020, the average value is 940.0 mm. Based on the analysis of the grid rainfall anomaly from 1980 to 2020 and from 1990 to 2020, the two series both experienced multiyear drought from 2001 to 2009 (Fig. R1), which is consistent with the period of multiyear drought in the manuscript (the drought period is from 2000 to 2009). There is an alternate change of wet and dry after 2015, especially, a large negative rainfall anomaly (%) is existed after 2017. Therefore, the extended data may not greatly improve the results.

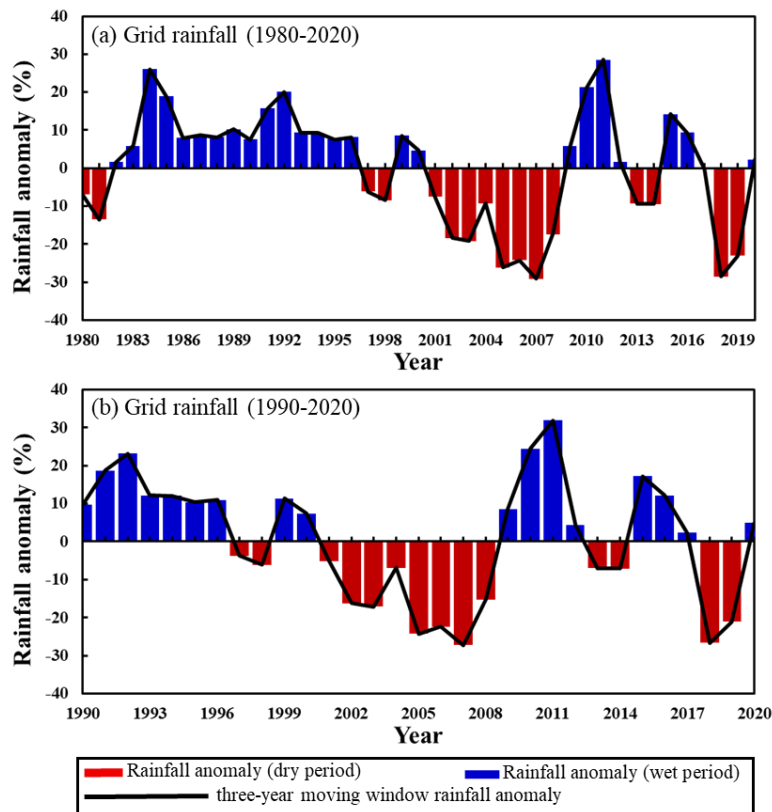


Figure R1 Rainfall anomaly (%) as a percentage of the mean annual rainfall of Kileys Run and Red Hill catchment. (a) Grid rainfall data (1980-2020); (b) Grid rainfall data (1990-2020). Red bars represent dry years and blue bars represent wet years. The black line represents the three-year moving average of the rainfall anomaly.

R15. Line 87, which method did you used to measure the runoff? Of where did you collected the data?

Response: Thanks for bringing this to our attention and the source of the data have been mentioned in the revised manuscript (Line 96-98).

Runoff data were measured on site by government agency and were collected by Dr. Lu Zhang from CSIRO Land and Water, Black Mountain, Canberra, Australia. Observations were made at the outlet of each catchment, a gauging station was installed and began operation on 26 May, 1989. Daily runoff data are therefore available from June 1989 onwards. The gauging stations comprise a flat-v style crump weir at the Red Hill and Kileys Run catchments. The rating curves (stage-discharge relationships) for the Red

Hill and Kileys Run catchments had been developed and tested by a series of velocity-area gauging and, as a result, the runoff data from those stations was considered reliable (Webb and Kathuria, 2012).

Relevant text reads (Line 96-98): “The daily rainfall was measured by tipping bucket rain gauges had been located at catchment outlet and the daily runoff was measured by a flat-v style crump weir at a gauging station at the outlet of each catchment (Major et al., 1998).”

R16. Lines 88-89, add more information/background. You use daily rainfall and runoff, but monthly PET?

Response: Thanks for your advice. More information about data have been added in section 2 (Line 101-104).

The daily rainfall and runoff data were measured on site by government agency and were collected by Dr. Lu Zhang from CSIRO Land and Water. The potential evapotranspiration (PET) data were monthly data from the SILO Data (www.longpaddock.qld.gov.au/silo/point-data/). The daily data were only used for the analysis of flow duration curve. The monthly data were used for the PCM, TTM, the new framework and double mass curve. The annual data were used in the SBM.

Relevant text reads (Line 101-104): “The daily data were only used for the analysis of flow duration curves (FDCs). The monthly data were used for the PCM, TTM, the new framework and double mass curves (DMCs). The annual data are used in the SBM.”

R17. Line 90, “figure 2 shows the rainfall anomaly that was calculated by the method proposed by”, I suggest to describe the method (define anomaly). Which in this case is the annual percentage of rainfall being larger or less than the average rainfall.

Response: Thanks for your advice. Changes have been made as suggested in section 2 (Line 104-108).

Rainfall anomaly (%) is defined as the percentage deviation of annual rainfall to mean annual rainfall. The period of drought is determined by rainfall anomaly (%) smoothed with a three-year moving window. The method for determining the multiyear drought period is as follows (Saft et al., 2015):

The first year of the drought remains the start of the first three-year negative anomaly period;

The end year is set as the last year of this three-year negative anomaly period (unless: if the last two years have slightly positive anomalies (but each <15% of the mean), in which case the end year is set to the first year of positive anomaly);

The length of dry periods must be not less than seven years;

The mean dry period anomaly must be less than -5%;

Relevant text reads (Line 104-108): “Rainfall anomaly (%) is defined as the percentage deviation of annual rainfall to mean annual rainfall. It can be seen that three-year moving average of the rainfall anomaly (the black line) is lower than zero from 2000 to 2009. According to the method of determining multiyear drought period proposed by Saft et al. (2015), two catchments experienced prolonged drought lasted 10 years from 2000 to 2009 and this coincided with the period of the Millennium Drought of Australia (van Dijk et al., 2013).”

R18. Line 91-92, I suggest add to add the reference again.

Response: Thanks for your advice. Changes have been made as suggested. (Line 108-109, see references below).

Relevant text reads (Line 108-109): “this coincided with the period of the Millennium Drought of Australia (Peterson et al., 2021; van Dijk et al., 2013).”

van Dijk, A. I. J. M., Beck, H. E., Crosbie, R. S., de Jeu, R. A. M., Liu, Y. Y., Podger, G. M., Timbal, B., and Viney, N. R.: The Millennium Drought in southeast Australia (2001-2009): Natural and human causes and implications for water resources, ecosystems, economy, and society, *Water Resour. Res.*, 49, 1040-1057, <https://doi.org/10.1002/wrcr.20123>, 2013.

Peterson, T. J., Saft, M., Peel, M. C., and John, A.: Watersheds may not recover from drought, *Science*, 372, 745-749, <https://doi.org/10.1126/science.abd5085>, 2021.

R19. Figures 3 – 5, and 7: used colours (combination of red and green) are not very suitable for colour blind readers (HESS - Submission (hydrology-and-earth-systemsciences.net))

Response: We apologize for these oversights. Fig. 3 – 5, and 7 have been modified in the new manuscript (Line 220; Line 275; Line 307; Line 442, refer to Fig. 3 – 5 and 8).

R20. Lines 363-365, add reference?

Response: Thanks for your advice. Changes have been made as suggested (Line 449-450, see references below).

Relevant text reads (Line 449-450): “which is the essence of the limitations of traditional application (Dey and Mishra, 2017; Li et al., 2012; Zhang et al., 2019).”

Li, H., Zhang, Y., Vaze, J., and Wang, B.: Separating effects of vegetation change and climate variability using hydrological modelling and sensitivity-based approaches, *J. Hydrol.*, 420-421, 403-418, <https://doi.org/10.1016/j.jhydrol.2011.12.033>, 2012.

Dey, P., and Mishra, A.: Separating the impacts of climate change and human activities on streamflow: A review of methodologies and critical assumptions, *J. Hydrol.*, 548, <https://doi.org/10.1016/j.jhydrol.2017.03.014>, 2017.

Zhang, L., Nan, Z., Wang, W., Ren, D., Zhao, Y., and Wu, X.: Separating climate change and human contributions to variations in streamflow and its components using eight time - trend methods, *Hydrol. Process.*, 33, 383-394, <https://doi.org/10.1002/hyp.13331>, 2019.

R21. Line 372, add reference?

Response: Thanks for your advice. Changes have been made as suggested (Line 455-457, see references below).

Relevant text reads (Line 455-457): “and multiyear drought weakened the impact of vegetation change on runoff (see Table 2), which was important for us to design ecological engineering projects for

sustainable water resources management (Brodribb et al., 2020; Newman et al., 2006; Xiao et al., 2020).”

Xiao, Y., Xiao, Q., and Sun, X.: Ecological Risks Arising from the Impact of Large-scale Afforestation on the Regional Water Supply Balance in Southwest China, *Sci. Rep.-UK*, 10, 4150, <https://doi.org/10.1038/s41598-020-61108-w>, 2020.

Newman, B. D., Wilcox, B. P., Archer, S. R., Breshears, D. D., Dahm, C. N., Duffy, C. J., McDowell, N. G., Phillips, F. M., Scanlon, B. R., and Vivoni, E. R.: Ecohydrology of water-limited environments: A scientific vision, *Water Resour. Res.*, 42, <https://doi.org/https://doi.org/10.1029/2005WR004141>, 2006.

Brodribb, T. J., Powers, J., Cochard, H., and Choat, B.: Hanging by a thread? Forests and drought, *Science*, 368, 261-266, <https://doi.org/10.1126/science.aat7631>, 2020.

R22. Figure 4, perhaps add cumulative rainfall and the separated periods.

Response: Thanks for your advice. Changes have been made as suggested in Fig. 4 and Line 261-263.

The rainfall data used in these two catchments are the same. The measured cumulative rainfall shown in Fig. 4 has two significant abrupt change points in October 1999 and March 2012. They are consistent with the beginning and end of multiyear drought (2000 – 2009).

Relevant text reads (Line 261-263): “Cumulative monthly rainfall of paired catchments shown in Fig. 4 (e) also had two significant abrupt change points in October 1999 and March 2012 that were consistent with the beginning and end of multiyear drought (2000–2009).”

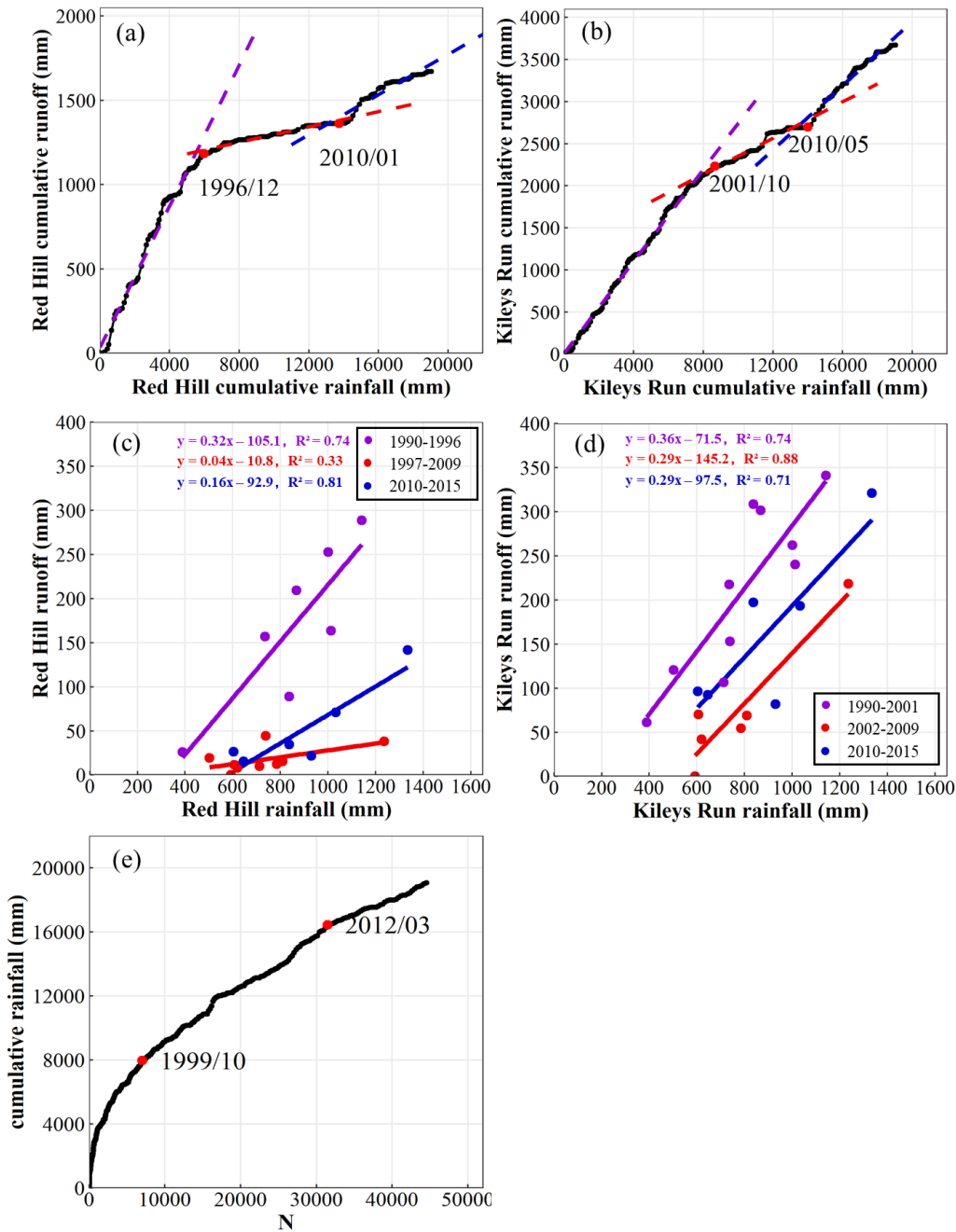


Figure 4: (a) Double mass curve of monthly rainfall and runoff of the Red Hill catchment (treated catchment), (b) Double mass curve of monthly rainfall and runoff of the Kileys Run catchment (control catchment), (c) Relationships between annual rainfall and runoff of the Red Hill catchment (treated catchment), (d) Relationships between annual rainfall and runoff of the Kileys Run catchment (control catchment), and (e) Cumulative monthly rainfall of paired catchments during the period of 1990–2015. The dashed lines in (a) and (b) represent the linear regression lines between cumulative rainfall and cumulative runoff during three different periods (January 1990 to December 1996 (purple), January 1997 to January 2010 (red), and February 2010 to December 2015 (blue) in Red Hill; January 1990 to October 2001 (purple), November 2001 to May 2010 (red), and June 2010 to December 2015 (blue) in Kileys Run). The purple, red, and blue lines in (c) and (d) represent the linear regression lines for three different periods (1990–1996, 1997–2009, and 2010–2015 in Red Hill; 1990–2001, 2002–2009, and 2010–2015 in Kileys Run).

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a preliminary water yield analysis, Forest Research and Development Division, State Forests of New South Wales, Sydney, 24 pp., 1998.

Newman, B. D., Wilcox, B. P., Archer, S. R., Breshears, D. D., Dahm, C. N., Duffy, C. J., McDowell, N. G., Phillips, F. M., Scanlon, B. R., and Vivoni, E. R.: Ecohydrology of water-limited environments: A scientific vision, *Water Resour. Res.*, 42, <https://doi.org/https://doi.org/10.1029/2005WR004141>, 2006.

Peters, E., Torfs, P. J. J. F., van Lanen, H. A. J., and Bier, G.: Propagation of drought through groundwater- A new approach using linear reservoir theory, *Hydrol. Process.*, 17, 3023-3040, <https://doi.org/10.1002/hyp.1274>, 2003.

Peterson, T. J., Saft, M., Peel, M. C., and John, A.: Watersheds may not recover from drought, *Science*, 372, 745-749, <https://doi.org/10.1126/science.abd5085>, 2021.

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Response to Reviewer #3

General Comments:

R1. This is an interesting paper on an important topic. In fact it could be argued the issues discussed in this paper are absolutely fundamental in hydrology. This is particularly true in light of the way “data” is used uncritically in many studies.

Response: Thanks very much for your great efforts to assess our manuscript. We have studied your and reviewers’ comments carefully and have made corrections/revisions as suggested. The point-to-point responses to the comments and our plans for revision are listed below. In the following, we have detailed how these comments (in black) are raised and our responses (in deep sky blue).

R2. The paper is generally well written and structured, although there is a need for a more careful read-through as there is some awkward syntax and grammar (particularly in the Discussion) where the quality of the writing seems to wander a little. My main questions relate to the data set used and the essential paradigm of the PCM. The Redhill site did not have a calibration period. Although the authors suggest that the first period of treatment may be thought of as non-treated in that the trees were very small and not high water using, I feel the implications of this may be important. This then connects to the PCM paradigm; that is, that the length of calibration or the approach developing the calibrations in theory should account for the type of non-stationarity that is discussed (ie. drought). Putting this another way, how can we decide what is non-stationarity and what is variability? This is particularly germane to Australian hydrology where we experience significant variability. I am not suggesting climates are stationary, but disentangling non-stationarity from variability with a relatively short data period (in climate terms) is a question.

Response: Thanks for your constructive comments. We have added more contents about the calibration period, non-stationarity and variability in discussion section in the revised manuscript (Line 458-474; Line 526-544).

- (1) We agree with the question of the calibration period. However, for the Red Hill catchment, it was unfortunate that rainfall and runoff data before treatment were not measured. In Zhao et al. (2010), the impacts of the calibration period determined by treatment period and period before abrupt change point of annual runoff on the results at four paired catchments were compared, it was found that runoff changes caused by vegetation change were not sensitive to the different calibration periods. Moreover, considering that the runoff may not change significantly in the first few years after plantation of seedlings of *P. radiata*, we re-estimated the impact of vegetation change on runoff based on a calibration period with the data of the previous three years. The contribution of vegetation change to total runoff changes was 34.2%, and the difference with the result in the manuscript was only 1.4%. Therefore, the calibration period set in this manuscript was reasonable.
- (2) In this study, we distinguish climate variability and non-stationarity by considering the influence on the rainfall-runoff relationship. For climate variability, we think that it will not lead to non-stationary changes in the rainfall-runoff relationship, that is, rainfall and runoff changes at the same rate. For non-stationarity, it will lead to non-stationary changes in the rainfall-runoff

relationship, that is, it can be reflected by the significant abrupt change point of double mass curve and the significant up/down movement of rainfall-runoff linear regression line (Avanzi et al., 2020; Li et al., 2018). For large-scale watersheds, it is difficult to detect and confirm non-stationarity from variability because of the complexity and regional differences of positive and negative fluctuations or feedback of climate. However, for the two small studied catchments, the impact of climate fluctuations is very intense, and persistent fluctuations below the average are easy to cause non-stationary changes in the rainfall-runoff relationship.

Relevant text reads (Line 458-474): “Climate variability and multiyear drought are supposed to have essentially different influences on the rainfall-runoff relationship in this study. Climate variability is not supposed to result in non-stationary changes in the rainfall-runoff relationship, that is, rainfall and runoff change at the same rate. While multiyear drought is assumed to result in non-stationary changes in the rainfall-runoff relationship, that is, it can be demonstrated by the significant abrupt change point on the double mass curves (DMCs) and the significant downward shift of rainfall-runoff linear regression line (Avanzi et al., 2020; Li et al., 2018). The multiyear drought in this study refers to drought with long duration and severe intensity, which can cause non-stationary changes in the rainfall-runoff relationship of catchments as shown in Fig. 4 and discussed in section 5.2. It is quite different from the wet/dry periods fluctuating near the average line (i.e., climate variability) (Han et al., 2019). For the two small studied catchments, the impact of climate fluctuations is very intense, and persistent fluctuations below the average are easy to cause non-stationary changes in the rainfall-runoff relationship because the long-term rainfall reduction may lead to changes of catchment characteristics, that is, lose connection between surface and groundwater. However, for large-scale watersheds, it is difficult to detect and to separate non-stationarity from variability because of the complexity and regional differences of positive and negative fluctuations or feedback of climate (Clark et al., 2016; Murakami et al., 2020). Negligence of non-stationarity induced by multiyear drought can result in significant differences in estimated effects of vegetation change as shown in Fig. 6, which has also been reported by Zhao et al. (2010) using about 16 years data at the same site. In the new framework, effect of multiyear drought is estimated between pre- and post-change periods as that of vegetation change, although two years of rainfall is above the average after 2000. Because the slope of DMCs is still very close to that during the period after 2009 (post-drought) (see Fig. 4).”

(Line 526-544): “For Red Hill experiment site, the calibration period was from one year after treatment to the abrupt change point of annual runoff (1990–1996, seven years), because rainfall and runoff data before treatment were not measured. Zhao et al. (2010) compared the influences of two different methods for determining the calibration period on the estimated vegetation impacts at four paired catchment sites. One is determined by the time of treatment. The other is determined by the abrupt change point of annual runoff. It was found that runoff changes caused by vegetation change were not sensitive to different calibration periods. Considering that runoff may not change significantly during the first few years after plantation of seedlings of *P. radiata*, we re-estimated the impact of vegetation change on runoff based on a calibration period with the data of the previous three years (1990–1992). Impacts of vegetation change were 34.2%, 74.2% and 61.0% of total runoff changes by the PCM, TTM and SBM, respectively. The contributions of vegetation change, multiyear drought and climate variability to total runoff changes using the new framework were 34.2%, 37.4% and 39.0%, respectively. Comparing to those in Table 2, the difference of the contribution of vegetation change to total runoff changes was only 1.4%. It indicates that selection of the length of the calibration period may have little impact on the estimation of runoff changes caused by vegetation change before the treated catchment reaches a new equilibrium state. This issue has also been discussed in Bren and Lane (2014) and they found that runoff

of paired catchments had good calibrations (Nash-Sutcliffe efficiency (N-S) = 0.8) with 100 days of data and very little improvement after three years. For Red Hill experiment site, the change of N-S is close to that reported by Bren and Lane (2014). Good calibration (N-S > 0.85) is achieved with about 150 days. Similar results are obtained with monthly flows, good calibration (N-S > 0.35) is achieved with about 24 months. It suggests that runoff of the Red Hill and Kileys Run catchments will be well calibrated with calibration period exceeds 150 days (daily data) or 24 months (monthly data). Considering longer calibration period has lower mean error, calibration period is set from beginning of available data to the time of the abrupt change of annual runoff in this study.”

R3. The calibration period issue is a vexed one as there is no longer an appetite by funding bodies to set up a paired-catchment experiment and then wait for a lengthy period before anything happens. Bren and Lane (2014, JH 519) explored this issue and proposed a method using daily flows that rather obviously increases the number of data points. Somewhat surprisingly the analysis showed that good calibrations (Nash-Sutcliffe E = 0.8) with 100 days of data, and very little improvement after 3 years. Apologies for the treatise, but I wonder if this approach might be useful in thinking about the PCM. That is, if such an analysis was performed and compared with the other analyses it might be very useful. Data could be randomly pulled out of the Kileys Run data. At the very least it should be discussed.

Response: Thanks for your suggestion. We have added discussions regarding to your concerns in the revised manuscript (Line 536-544).

We used the data of seven years before the abrupt change point of annual runoff and three years after the beginning of the data series to explore the change of Nash-Sutcliffe coefficient (N-S) with the increasing length of the observation set. For seven years (1990-1996), the calibration set was from 1990 to 1992 and the verification set was from 1993 to 1996 (Fig. R1 (b) and (d)). For three years (1990-1992), the calibration set was from 1990 to 1992 and the verification set was from 1993 to 1996 (Fig. R1 (a) and (c)). According to Fig. R1, it can be found that the change of N-S was close to Bren and Lane (2014). It showed that good calibration (N-S > 0.85) was achieved in about 150 days, and then maintained at the high-level value. Similar results were obtained with monthly of flows, calibration (N-S > 0.35) was achieved in about 24 months.

We used the data of the previous three years (1990-1992) as pre-treatment period, and re-estimated the effects of vegetation change on runoff by the PCM, TTM, SBM and the new framework. Impacts of vegetation change were 34.2%, 74.2% and 61.0% of total runoff changes by the PCM, TTM and SBM, respectively. The contributions of vegetation change, multiyear drought and climate variability to total runoff changes were 34.2%, 37.4% and 39.0%, respectively, using the new framework. The sum of the three terms was 110.6%, with a difference of 10.6% from 100%. When the pre-treatment period was as pre-change period of runoff (1990-1996) used in the manuscript, impacts of vegetation change were 32.8%, 93.5% and 76.1% of total runoff changes by the PCM, TTM and SBM, respectively. The contributions of vegetation change, multiyear drought and climate variability to total runoff changes were 32.8%, 54.7% and 23.9%, respectively, using the new framework. The sum of the three terms was 111.4%, with a difference of 11.4% from 100% and a difference of 0.8% from 110.6%. It also can be found that the difference in the contribution of vegetation change to total runoff changes estimated by the two different ways to determine the calibration period was only 1.4%. It showed that taking period before the abrupt change point of annual runoff as the calibration period was also reasonable considering a lower mean error.

Relevant text reads (Line 536-544): “It indicates that selection of the length of the calibration period may have little impact on the estimation of runoff changes caused by vegetation change before the treated catchment reaches a new equilibrium state. This issue has also been discussed in Bren and Lane (2014) and they found that runoff of paired catchments had good calibrations (Nash-Sutcliffe efficiency (N-S) = 0.8) with 100 days of data and very little improvement after three years. For Red Hill experiment site, the change of N-S is close to that reported by Bren and Lane (2014). Good calibration (N-S > 0.85) is achieved with about 150 days. Similar results are obtained with monthly flows, good calibration (N-S > 0.35) is achieved with about 24 months. It suggests that runoff of the Red Hill and Kileys Run catchments will be well calibrated with calibration period exceeds 150 days (daily data) or 24 months (monthly data). Considering longer calibration period has lower mean error, calibration period is set from beginning of available data to the time of the abrupt change of annual runoff in this study.”

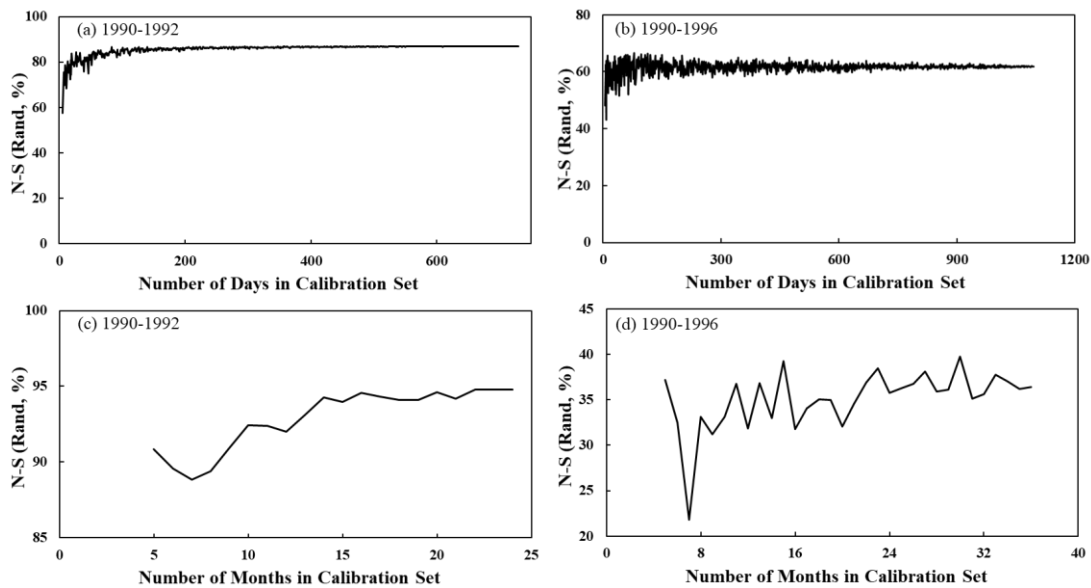


Figure R1. Mean of 10 values of the N-S coefficient of the Coefficient of Determination as a function of the number of days and months in the calibration data set for the Kileys Run and Red Hill data using the 10-fold cross-validation approach to monitor the development of the calibration. “Rand” refers to the data being randomized.

Specific Comments:

R4. Line 40 - there are more updated references for the research in Australian catchment behaviour that are relevant (eg Petersen et al, 2021, Science 372)

Response: Thanks for your advice. Changes will be made as suggested (Line 46-47, see references below).

Relevant text reads (Line 46-47): “It is widely known that Australia experienced multiyear drought (known as the Millennium Drought) between 1997 and 2009 (King et al., 2020; Peterson et al., 2021).”

King, A. D., Pitman, A. J., Henley, B. J., Ukkola, A. M., and Brown, J. R.: The role of climate variability in Australian drought, *Nat. Clim. Change*, 10, 177-179, <https://doi.org/10.1038/s41558-020-0718-z>, 2020.

Peterson, T. J., Saft, M., Peel, M. C., and John, A.: Watersheds may not recover from drought, *Science*,

R5. L46. I think this statement about PCM and non-stationarity requires more justification. How does it not deal with the issue given that is the paradigm of PCM?

Response: Thanks for your constructive comments. We have added more justification about the paired-catchment method (PCM) and non-stationarity in the revised manuscript (Line 376-383).

At Red Hill site, the treated catchment suffered from non-stationary changes caused by vegetation change and multiyear drought, and stationary changes caused by climate variability. The control catchment suffered from non-stationary changes only caused by multiyear drought, and stationary changes caused by climate variability. The PCM can offset the effects on both paired catchments induced by multiyear drought and climate variability based on data of the control catchment, and the effects of vegetation change on the treated catchment can then be separated. That is, for the paradigm of PCM, shift in the rainfall-runoff relationship separated from the runoff correlation between the treated and control catchments should be caused only by the treatment of the treated catchment and effects of any other drivers can induce either stationary or non-stationary changes should be eliminated by making use of the control catchment. Therefore, the PCM is still the most reliable method compared with other methods.

Relevant text reads (Line 376-383): “At Red Hill experiment site, non-stationary changes of the treated catchment are caused by both vegetation change and multiyear drought, and stationary changes are caused by climate variability. Non-stationary changes of the control catchment are only caused by multiyear drought, and stationary changes are only caused by climate variability. According to the paradigm of PCM, shift in the rainfall-runoff relationship separated from the runoff correlation between the treated and control catchments should be caused only by the treatment of the treated catchment and effects of any other drivers can induce either stationary or non-stationary changes should be eliminated by making use of the control catchment. Therefore, the PCM is still the most reliable method compared with other methods and separated effect by the PCM is only caused by vegetation change (i.e., afforestation).”

R6. L 65+ Both TTM and SBM require lengthy records; is this an issue with these analyses at Redhill? There may be an argument to see the drought as a plus in terms of a record with wet/dry periods.

Response: Thanks for your constructive comments. The discussion about this question have been mentioned in the revised manuscript (Line 462-472; Line 516-525).

(1) The length of data is essentially a problem of reaching an equilibrium state for catchments. Han et al. (2020) provided a global assessment of the steady-state assumption in catchment water balance calculations for 1,057 global unimpaired catchments. Results showed that ~70% of the catchments attain steady state within 10 years. The time needed for a catchment to reach steady state showed a close relationship with climatic aridity and vegetation coverage, with arid/semiarid and sparsely vegetated catchments generally having a longer time. For small catchment, it may need a shorter time to reach steady state. It can be seen in Fig. 4 (a), the double mass curve becomes a straight line in a very short time. It demonstrates that the length of data used in this study (26 years) is acceptable.

(2) Multiyear drought mentioned in the manuscript is a drought with longer duration and greater intensity, which will cause non-stationary changes in the rainfall-runoff relationship of catchments. It is quite different from the wet/dry periods floating near the average line. The effects of multiyear drought on runoff are very significant and cannot be ignored. It is necessary to consider impacts of multiyear drought on runoff in the new framework.

Relevant text reads (Line 462-472): “The multiyear drought in this study refers to drought with long duration and severe intensity, which can cause non-stationary changes in the rainfall-runoff relationship of catchments as shown in Fig. 4 and discussed in section 5.2. It is quite different from the wet/dry periods fluctuating near the average line (i.e., climate variability) (Han et al., 2019). For the two small studied catchments, the impact of climate fluctuations is very intense, and persistent fluctuations below the average are easy to cause non-stationary changes in the rainfall-runoff relationship because the long-term rainfall reduction may lead to changes of catchment characteristics, that is, lose connection between surface and groundwater. However, for large-scale watersheds, it is difficult to detect and to separate non-stationarity from variability because of the complexity and regional differences of positive and negative fluctuations or feedback of climate (Clark et al., 2016; Murakami et al., 2020). Negligence of non-stationarity induced by multiyear drought can result in significant differences in estimated effects of vegetation change as shown in Fig. 6, which has also been reported by Zhao et al. (2010) using about 16 years data at the same site.”

(Line 516-525): “It shows that the increase of data length has little effect on the estimation of runoff change caused by vegetation change after runoff of catchment experiencing vegetation change has reached a new stable equilibrium state. The time required for runoff in different catchments to reach a new equilibrium state is different. For example, the Red Hill catchment takes seven years, Australia and New Zealand have suggested that three to 10 years, or even more (18 years for an afforested catchment in Biesievlei, South Africa (Brown et al., 2005)), majority of the time is between five and 10 years (Lane et al., 2005), are required for the treated catchment to reach a reasonably stable rainfall-runoff relationship after vegetation change (Zhao et al., 2010). Han et al. (2020) provided a global assessment of the steady-state assumption in catchment water balance calculations for 1,057 global unimpaired catchments and shown that ~70% of the catchments attained steady state within 10 years. For small catchment, it may need a shorter time to reach steady state. Thus the length of data used in this study (26 years) is enough to reach a steady state.”

R7. Lane et al. 2005 (JH) used FDCs that included Redhill – might be worth including these results as a comparison from a different method. This paper also has some estimates of time to equilibrium.

Response: Thanks for your advice. Changes have been made as suggested in section 5.2 and 5.4 in the new manuscript (Line 416-418; Line 520-522).

Relevant text reads (Line 416-418): “and the increases in zero flow days with low flows being more affected than high flows of daily flow duration curves (FDCs) in 10 catchments from southeastern Australia, New Zealand and South Africa (Lane et al., 2005).”

(Line 520-522): “majority of the time is between five and 10 years (Lane et al., 2005), are required for the treated catchment to reach a reasonably stable rainfall-runoff relationship after vegetation change.”

R8. L 69 – syntax not great “issues about this hypothesis” could be improved

Response: We apologize for these oversights. We have modified these sentences in the new manuscript (Line 75-77).

Relevant text reads (Line 75-77): “However, this question has not been explored and verified, and clarifying whether multiyear drought will have an important impact on the application ability of the three widely used methods will provide a meaningful reference for ecological engineering under changing climate with frequent extremes in future.”

R9. L 107 – should be double mass curves, FDCs etc. There are quite a few examples of this, need a careful read.

Response: We apologize for these oversights. We have modified them in the revised manuscript (Line 101-103; Line 127-128; Line 254; Line 461).

Relevant text reads (Line 101-103): “The daily data were only used for the analysis of flow duration curves (FDCs). The monthly data were used for the PCM, TTM, the new framework and double mass curves (DMCs).”

(Line 127-128): “Double mass curves (DMCs), flow duration curves (FDCs), and rainfall-runoff linear regression curves were employed to detect changes in the rainfall-runoff relationship caused by vegetation change and multiyear drought.”

(Line 254): “The double mass curves (DMCs) of monthly rainfall and runoff of the two paired catchments are shown in Fig. 4 (a) and Fig. 4 (b).”

(Line 461): “it can be demonstrated by the significant abrupt change point on the double mass curves (DMCs)”

R10. L 168- See earlier general discussion. It does trouble me that a site with no calibration is used for this study. In addition, the vegetation effect is dynamic; growing from seedlings to (presumably, given there are no growth data) a closed canopy. I do wonder if this really is the best data set for such a study, or what might be gained from using more data sets.

Response: Thanks for your constructive comments. Except for Red Hill site, there are no paired catchments that have experienced vegetation change and multiyear drought at the same time. The use of data and the calibration period were also further discussed in section 5.4 in the new manuscript (Line 526-544). It was found that the difference in the contribution of vegetation change to total runoff changes estimated by the two different ways to determine the calibration period (the previous three years after treatment and seven years before the abrupt change point of annual runoff in the manuscript) was only 1.4%. In this study, we not only found the reasons why results of the three traditional methods were inconsistent, but also thought more deeply about the paradigm of PCM. For PCM, shift in the rainfall-runoff relationship separated from the runoff correlation between the treated and control catchments should be caused only by the treatment of the treated catchment and effects of any other drivers can induce either stationary or non-stationary changes should be eliminated by making use of the control catchment. Therefore, the PCM is still the most reliable method compared with other methods.

Relevant text reads (Line 526-544): “For Red Hill experiment site, the calibration period was from one year after treatment to the abrupt change point of annual runoff (1990–1996, seven years), because rainfall and runoff data before treatment were not measured. Zhao et al. (2010) compared the influences of two different methods for determining the calibration period on the estimated vegetation impacts at four paired catchment sites. One is determined by the time of treatment. The other is determined by the abrupt change point of annual runoff. It was found that runoff changes caused by vegetation change were not sensitive to different calibration periods. Considering that runoff may not change significantly during the first few years after plantation of seedlings of *P. radiata*, we re-estimated the impact of vegetation change on runoff based on a calibration period with the data of the previous three years (1990–1992). Impacts of vegetation change were 34.2%, 74.2% and 61.0% of total runoff changes by the PCM, TTM and SBM, respectively. The contributions of vegetation change, multiyear drought and climate variability to total runoff changes using the new framework were 34.2%, 37.4% and 39.0%, respectively. Comparing to those in Table 2, the difference of the contribution of vegetation change to total runoff changes was only 1.4%. It indicates that selection of the length of the calibration period may have little impact on the estimation of runoff changes caused by vegetation change before the treated catchment reaches a new equilibrium state. This issue has also been discussed in Bren and Lane (2014) and they found that runoff of paired catchments had good calibrations (Nash-Sutcliffe efficiency (N-S) = 0.8) with 100 days of data and very little improvement after three years. For Red Hill experiment site, the change of N-S is close to that reported by Bren and Lane (2014). Good calibration (N-S > 0.85) is achieved with about 150 days. Similar results are obtained with monthly flows, good calibration (N-S > 0.35) is achieved with about 24 months. It suggests that runoff of the Red Hill and Kileys Run catchments will be well calibrated with calibration period exceeds 150 days (daily data) or 24 months (monthly data). Considering longer calibration period has lower mean error, calibration period is set from beginning of available data to the time of the abrupt change of annual runoff in this study.”

R11. Figure 3 is a great figure!

Response: Thanks for your comments.

R12. Table 2 – total flow changes would be useful. They appear later in the text but I think having totals in the table make it easier to evaluate the methods.

Response: Thanks for your advice. The total flow changes have been mentioned in Table 2 (Line 330).

R13. This also brings up another point that I don't think has been discussed properly. Q_{clim} is conceptualised as the climate effect, encompassing wet and dry and mean climate inputs. I am not sure there is adequate discussion of how this does not deal with the climate issue as formulated.

Response: Thanks for your constructive comments. The discussion about this question have been mentioned in the revised manuscript (Line 389-394; Line 402-407).

Q_{clim} in this study refers to the impact of climate variability (wet and dry spells) on runoff that does not lead to non-stationary changes in rainfall-runoff relationship. It represents the runoff changes caused by climate stability changes and it is estimated by the SBM. The SBM is derived from the Budyko framework. In the formula, the impact of climate change on runoff is estimated by rainfall and potential

evapotranspiration changes. In the Budyko curve, it is reflected in moving from one point to another point on the same curve. There is a necessary assumption in the SBM, that is, a transition from one steady state to another with no change in catchment properties (the same curve) (Roderick and Farquhar, 2011; Sun et al., 2014). When the SBM is applied to the Red Hill catchment that has experienced multiyear drought, the result is actual the impact of climate variability (without changing the catchment characters/non-stationary changes in rainfall-runoff relationship) on runoff, that is, Q_{clim} . It may ignore the impact of multiyear drought on runoff and indirectly overestimate the impact of vegetation change on runoff. Because catchment properties of the Red Hill catchment may have changed due to multiyear drought (Kinal and Stoneman, 2012; Peterson et al., 2021; Saft et al., 2016; van Dijk et al., 2013), which violates the assumptions of the SBM.

Relevant text reads (Line 389-394): “The SBM is sourced from the Budyko framework (Budyko, 1974). It assumes that steady state of catchment water balance is fundamentally determined by water input (represented by precipitation) and energy demand (represented by potential evapotranspiration) and transition from one steady state to another without any change in catchment properties should moving on the Budyko curve (Roderick and Farquhar, 2011; Sun et al., 2014; Wang et al., 2021). Therefore, stationary changes driven by climate variability during post-treatment period can be separated by sensitivity of runoff to P and PET established during the pre-treatment period.”

(Line 402-407): “The result estimated by the SBM is the impact of climate variability (without changing the catchment characters/non-stationary changes in the rainfall-runoff relationship) on runoff, that is, ΔQ_t^{clim} . It ignored the impact of multiyear drought on runoff, which have proved to cause non-stationary changes. Recent studies have also reported that multiyear drought can cause catchment properties changes and hydrological functionings (Kinal and Stoneman, 2012; Peterson et al., 2021; Saft et al., 2016; van Dijk et al., 2013), which may violate the assumptions of the SBM.”

R14. 5.2.1 – this paragraph brings up the interesting point (that is the subject of the Saft/Peterson/Fowler etc studies); is it the climate that is non-stationary or is the processes (obviously driven by the climate).

Response: Thanks for your constructive comments. We think that climate change is a combination of non-stationary changes and process. The non-stationary change may be due to the sudden and drastic changes of human activities (explosive increase of industrial activities (Bauska et al., 2015)) and/or the amount of solar energy that gets to earth (Karl and Trenberth, 2003) and other factors. Most of these non-stationary changes occur suddenly and change from one equilibrium state to another through a relatively short period, such as non-stationary change in rainfall-runoff relationship caused by multiyear drought (Fowler et al., 2018; Peterson et al., 2021; Saft et al., 2015). The process is long-term, it may keep the equilibrium state and change continuously, or show a trend change (increase or decrease) over time. For example, the air temperature shows a gradual upward trend in a longer period (years and decades), and in a shorter period (days, months and seasons), the temperature constantly fluctuates around the average temperature, which means that there will still be those days which are cool and those days which are warm (Hoegh-Guldberg et al., 2019).

R15. The Paragraph around Line 375 needs some rewriting, the syntax is jarring. For example “the” control..

Response: We apologize for these oversights. Changes have been made as suggested in the revised manuscript (Line 484-488).

Relevant text reads (Line 484-488): “In the new framework, the control catchment plays an irreplaceable role in estimating the impact of vegetation change and multiyear drought on runoff. Because the control catchment can eliminate the impact of climate variability and multiyear drought on runoff when the PCM is used to quantify runoff change caused by vegetation change, and the impact of multiyear drought on the treated catchment is transferred from the control catchment. The former must use the runoff data of the control catchment, and the latter needs both the rainfall and runoff data of the control catchment.”

R16. L 388 “Because Saft..” this is a poor sentence

Response: We apologize for these oversights. We have modified these sentences in the manuscript (Line 499-501).

Relevant text reads (Line 499-501): “Saft et al. (2016) re-evaluated a wide range of factors that may be responsible for the additional runoff reductions and suggested that the shifts were mostly influenced by catchment characteristics related to pre-drought climate and soil and groundwater storage dynamics but less affected by the percentage of woody cover.”

R17. L 399 “pines” should not be italicized. *P.radiata* would be

Response: We apologize for these oversights. The word “*P. radiata*” has been modified to “*P. radiata*” (Line 92; Line 510; Line 531).

Relevant text reads (Line 92): “In 1988, *P. radiata* was planted in the Red Hill catchment (0.5 km²)”

(Line 510): “the treated catchment was covered by *P. radiata*”

(Line 531): “the first few years after plantation of seedlings of *P. radiata*”

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