# Influences of land use changes on the dynamics of water quantity and quality in the German lowland catchment of the Stör

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Abstract. Understanding the impacts of land use changes (LUCC) on the dynamics of water quantity and quality is necessary to identify suitable mitigation measures <u>favorable that are needed</u> for sustainable watershed management. Lowland catchments are characterized by a strong interaction of streamflow and near-surface groundwater that intensifies the risk of nutrient pollution. <u>This studyIn this study, we aims to reveal the investigated</u> relationship between the effects of long-term land use changes in individual land use classes and the on the water and nutrient balances in a typical lowland in the lowland

- catchment of the upper Stör catchment\_in northern lowland areas, Northern Germany. To this end, A using the hydrological model SWAT (Soil and Water Assessment Tool, SWAT) and partial least squares regression (PLSR) are used. SWAT model runs for three different land use maps (1987, 2010, 2019) were conducted and the outputs were compared to
   deriveRespective changes in -water quantity (i.e., evapotranspiration (ET), surface runoff (SO), base flow (BF) and water
- 15 deriveRespective changes in -water quantity (i.e., evapotranspiration (ET), surface runoff (SQ), base flow (BF), and water yield (WYLD)) and quality variables (i.e., sediment yield (SED), total phosphorus (TP) and total nitrogen (TN) loads). These changes were between any two of the three simulations conducted by SWAT using land use maps in 1987, 2010, and 2019 were found to be relatedd to to land use changes at the subbasin scale usingaecording to PLSR-results. The major land use changes that significantly affected water quantity and quality variables were related to a decrease of arable land and a
- 20 respective increase of pasture and urban land during the period of 1987-2019. Changes of landscape indictors such as area size, shape, dominance, and aggregation of each land use class, could accounted for as much as 61%-88% (75% on average) of the respective variations in water quantity and quality variables. The aggregation, contiguity degrees, and area extent of arable land were found to be most important to control the variations in most water quantity variables. Increases of arable (PLAND*a*) and urban land percent (PLAND*u*) led to morewould markedly accelerate TP and TN pollution, sediment export,
- 25 and surface runoff.PLSR. The cause-effect results of this study can provide a quantitative basis for targeting the most influential land use-change in landscape composition and configurations to mitigate adverse impacts on water quality in the future.

The change in the areal percentage of arable land positively affected the dynamics of SED, TP, TN and negatively affected BF, indicated by a Variable Influence on Projection (VIP) > 1.16 and large absolute regression coefficients (RCs: 0.6-0.88)

- 30 for SED, TP, TN; 1.65 for BF). The change in pasture area was negatively affecting SED, TP, and TN (RCs: 0.69 0.12, VIPs >1) while positively affecting ET (RC: 0.09, VIP: 0.92). The change in settlement percentage had a VIP of up to 1.17 for SQ and positively and significantly influenced it (RC: 1.16, p-value < 0.001), were used to quantify the impacts of different land use types on the variations in actual evapotranspiration (ET), surface runoff (SQ), base flow (BF), and water yield (WYLD) as well as on sediment yield (SED), total phosphorus (TP) and total nitrogen (TN) loads. To this end, the</p>
- 35 model was calibrated and validated with daily streamflow data (30 years) as well as sediment and nutrient data from two water quality measurement campaigns (3 years in total). Three model runs over thirty years were performed using land use maps of 1987, 2010, and 2019, respectively. Land use changes between those years were used to explain the modelled changes in water quantity and quality on the subbasin scale applying PLSR. SWAT achieved a very good performance for daily streamflow values (calibration: NSE=0.79, KGE=0.88, PBIAS=0.3%; validation: NSE=0.79, KGE=0.87,
- 40 PBIAS=7.2%), a satisfactory to very good performance for daily TN (calibration: NSE=0.64, KGE=0.71, PBIAS=-11.5%; validation: NSE=0.86, KGE=0.91, PBIAS=5%), a satisfactory performance for daily sediment load (NSE=0.54.0.65,

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KGE=0.58 0.59, PBIAS= 22.2% 12%), and an acceptable performance for daily TP (calibration: NSE=0.56, KGE=0.65, PBIAS= 4.7%; validation: NSE=0.29, KGE=0.22, PBIAS= 46.2%) in the Stör Catchment. The variations in ET, SQ, BF, WYLD, SED, TP, and TN could be explained to an extent of 61% 88% by changes in the area, shape, dominance, and aggregation of individual land use types. They were largely correlated with the major LUCC in the study area i.e. a decrease of arable land, and a respective increase of pasture and settlement. The change in the areal percentage of arable land positively affected the dynamics of SED, TP, TN and negatively affected BF, indicated by a Variable Influence on Projection (VIP) > 1.16 and large absolute regression coefficients (RCs: 0.6-0.88 for SED, TP, TN; -1.65 for BF). The change in pasture area was negatively affecting SED, TP, and TN (RCs: 0.69 - 0.12, VIPs >1) while positively affecting ET (RC: 0.09, VIP: 0.92). The change in settlement percentage had a VIP of up to 1.17 for SQ and positively and significantly influenced it (RC: 1.16, p-value <0.001).

# 1 Introduction

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Good water quality and quantity are essential for enhancing ecological stability and diversity, and both of which play important roles in maintaining sustainable agricultural or economic development and human health (Antolini et al., 2020;
Gleick, 2000; Lu et al., 2015; Singh et al., 2017; Srinivasan and Reddy, 2009). The <u>water resources</u> dynamics of water quality and quantity at thewithin a catchment scale are mainly governed by a combination of climate and land use, as other catchment characteristics (e.g., topography, soil, and lithology) usually do not change on a short term (Farjad et al., 2017; Shuster et al., 2005; Wagner et al., 2018). Vice versa, hydrology affects land use as well (Wagner and Fohrer, 2019; Wagner and Waske, 2016). So far, many In the past three decades, Many efforts have been made to study studies have found the

- 60 influences of land use changes with respective to urbanization, deforestation, urbanization-and agriculture intensification have exerted significant effects of the change of land use area on water quality or water balance components (Kändler et al., 2017; Shrestha et al., 2018; Wagner et al., 2016). The effects on water quality have been a concern since the 1970s (Johnson et al., 1997). Land use patterns changes They can alter surface roughness, evapotranspiration, soil infiltration, and the interaction between surface and subsurface water (Fiener et al., 2011; Lei et al., 2021; Wei et al., 2007), and promote or
- 65 hinder generation and transportation of \_ConsequentlyCorrespondingly, the amount of water and the level of carried or transported\_soil particles, chemicals, or metals (Ding et al., 2016; Nafi'Shehab et al., 2021; Taka et al., 2022)transported can be promoted or hindered\_(refs), altering water quantity and quality. The effects of land use changes on catchment water resources are manifold, e.g., urbanization results in a significant increase in surface runoff and water yield (Ayivi and Jha, 2018), expansion of farmland area poses increased risks to non-point source pollution of nitrogen (N) and phosphorus (P) as
- 70 well as soil erosion (Hacisalihoglu, 2007; Jia et al., 2013; Rajaei et al., 2017; Roberts and Prince, 2010), whereas more seminatural vegetation (e.g., forest, bushland, or grassland) increases the ability of filtering pollutants and intercepting rainfall thus reducing water pollution and streamflow (Moreno - Matcos et al., 2008; Yan et al., 2013). Given the diverse direct effects of land use changes on hydrological processes and water contaminant generation and transportation inputs, fit is of great practical significance importance to identify the key predictor variablespattern of land use change s impacting
- 75 water resources, in order to achieve an effective catchment management of water and land use \_\_\_\_and water resourcesmanagement in a particular catchment. Changes of both the composition and spatial structure of landscape can exert diverse influences on catchment hydrology and ecological systems (Allan, 2004; Ding et al., 2016; Haidary et al., 2013; Shawul et al., 2019). It is imperative to discriminate the effects of different aspects of a certain land use class to target sustainable and comprehensive land and water management (Liu et al., 2012; Shi et al., 2013).
- 80 Earlier studies <u>have often aimed at analyzing land use change effects have generally measured relationships between land use transition and water quantity and quality, using the lumped indicators of landscape composition, e.g., <u>land use areal proportion percentage of a land use class of in</u> the catchment (Kumar et al., 2022; Lei et al., 2021). However, composition indicators <u>are rather coarse to depict the relationships, because they</u> do not convey any details with respect to spatial settings</u>

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of landscape patterns. <u>The Spatial</u> configuration of<u>s</u> in the spatial land use distribution scapes can be measured is another fundamental element measured by using landscape metrics (i.e., algorithmsmetrics to quantify the spatial structure of land use patterns within a defined geographic area). Compared to the composition indicators that refers to the abundance (e.g., areal percent) of land patches (i.e., homogenous areas of the landscape) (Hesselbarth et al., 2019) belonging to one certain class

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- without considering their spatial characteristics, landscape configuration metrics describe spatial fragmentation or distribution of patches, e.g., the shape complexity.-such as, including the metrics of the dominance, diversity, shape, aggregation, and interconnection of land use patches (i.e., homogenous areas of the landscape) (Hesselbarth et al., 2019)., Landscape configuration metrics of the dominance, diversity, shape, aggregation, and interconnection of land patches They play a critical part in determining the energy and matter fluxes of e.g., solar radiation, temperature, evapotranspiration, surface runoff, nutrients, and sediments from the landscape ecology perspective (Amiri and Nakane, 2009; Forman, 1995;
- Lei et al., 2019; Wu and Lu, 2021). They therefore affect hydrological and ecological processes. With the availability of advanced spatial analysis (e.g., GIS) and remote sensing techniques (RS), various landscape metrics can be acquired efficiently for an overall assessment of landscape structure, based on classified land use maps from satellite data. Landscape configuration metricsingConfigurationThey are arewere found to be sometimes sometimes more important as descriptors of
- water quality than composition metricsindicators in some case studies: m(Ding et al., 2016; Gémesi et al., 2011). E.g., etrics;
   e.g., ; Ding et al. (2016) observed that -found that water quality is more significantly affected by the configuration i.e. poorer
   water quality was not as much more associated with areal percentage as with the higher patch densities; (PD) of cropland, orchards and grassland or-\_ and thea higher value of largest patch index (LPI) of urban land, -compared-to-the-areal percentages of than area percent of them\_than by composition of the land use type-in a low-order streams dominated catchment (drainage area: 35,340 km<sup>2</sup>) in southeastern China.\_-, <u>Gémesi et al. (2011) indicated that contagion, cohesion, and aggregation indices are more important than composition variables with regard to the variability in TN and TP in the
  </u>
- Mississippi Atchafalaya River watershed in USA. Recent studies on land use effects on water quantity mainly focus on land use percent, rarely on landscape metricsDespite little consideration of landscape configuration in the studiesy of water quantity –(Anand et al., 2018; Shrestha et al., 2018), –; However, metrics like landscape the shape, dominance or, or connectivity degree of land patches of one certain land use type-is closely linked to the intensificationmodification of eatchmentthe hydrological cycle. seape may play critical roles in altering the hydrological cycle,; For example,e.g., more
- fragmented forest patches <u>may closely relate to the capacity of favor funneling of precipitationinfiltration and interception of rainfall (Ghimire et al., 2017); hardness and straightness of land patches of farmland, urban, and natural land uses influence streamflow rates at different magnitudes and directions (Riitters, 2019; Shi et al., 2013); more concentrated grassland patches result in greater evapotranspiration (Yu et al., 2020). Therefore, it is necessary to assess influences of changes in different aspects of a land use class to better understand their impacts on water resources dynamics.</u>
- In orderT-to quantify effects of land use changes on water resources-within-catchments, hydrological models have beenare widely implementedused (Gabriels et al., 2021; Idrissou et al., 2022; Wijesekara et al., 2012), e.g., SWAT (Soil and Water Assessment Tool) (Arnold et al., 1998), HSPF (Hydrological Simulation Program-Fortran) (Bicknell et al., 2001), or DHSVM (Distributed Hydrology-Soil Vegetation Model) (Wigmosta et al., 1994). Models are particularly useful to detect historic as well as future land use change impacts usingapplying a land use scenario analysis (Anand et al., 2018; Aredo et al
- 120 2021). As a physically-based and semi-distributed hydrological model, SWAT has proven its suitability for an integrated modeling of thewater, sediment, and nutrient dynamics in hydrological processes and in nutrients cycling in different-sized rural catchments (Aghsaei et al., 2020; Tan et al., 2021). even under circumstance where observation data are limited. Furthermore, it considers the spatial heterogeneity. SWAT has been applied in many catchments worldwide to investigate the hydrological and hydro-chemical effects due to spatio-temporal changes of land use (Amin et al., 2020; Anand et al., 2020).
- 125 2018; Boongaling et al., 2018). In lowland areas, the transport of water and nutrients is strongly influenced by flat topography and shallow groundwater tables in addition to the spatially heterogonous land use. The spatially distributedhydrological model SWAT is quitehas proven its suitability useful to incorporate model the eco-hydrological

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consequences of spatio-temporal land use changes invariations in spatial patterns of these lowland featurescatchments (Guse et al., 2014; Pott and Fohrer, 2017b). Particularly in theseveral lowland catchments in northern Germany, SWAT was 130 extensively proven to be a powerful modelling tooltested forin impact studies. E.g., Lam et al. (2012) using SWAT satisfactorily modeled the long-term observations of daily streamflow and nitrate load in the Kielstau catchment and identified found that diffuse source pollution (dominated by agriculture) contributed dominantly (95%) to nitrate load; In the Upper Stör catchment, Song et al. (2015) using SWAT-coupled SWAT with HEC-RAS to revealed analyze temporal dynamics of sediment as well as different sediment-loads in subbasins covered by heterogonous land use conditions. Guse et 135 al. (2015) incorporated different spatial distributions of agricultural crops into SWAT and successfully assessed the impacts of land use changes on nitrate load in the Treene catchment. Despite a high feasibility of SWAT modelling water quantity and quality, previous studies illustrated that the original SWAT version sometimes performed relatively poorly for recession limbs and low flow periods of streamflow (Guse et al., 2014; Pfannerstill et al., 2014), In lowland catchments, groundwater contributes significantly to low flows and thus becomes a dominant component of streamflow (Pott and Fohrer, 2017b). 140 contribution from shallow groundwater affects the of streasmflow .however the abundant groundwater recharge is it is not adequatelycompletely considereddescribed in original SWAT version. To more accurately model low flows, processs, an enhanced version of SWAT, SWAT<sub>SSS</sub>, was recently-developed in the Kielstau catchment (a northern lowland catchment in Germany), by conceptually separating the shallow groundwater aquifer of the original SWAT into a fast and slow shallow aquifers (Pfannerstill et al., 2014). SWAT285 was previously evidenced to perform betterwas successfully used for 145 modelling daily streamflow and nitrogen loads in a few German lowland catchments (e.g., Kielstau and Treene) by optimizingimproving the representation of lowland- flow periods (Haas et al., 2017; Pfannerstill et al., 2014)-. Given the aforementioned strength-of SWAT3s application, SWAT3s is assumed to be moremore-suitable for assessing the impacts of land use changes on water resources in lowland areas dominated by groundwater recharge. While land use changes and the associated the changes in landscape metrics composition and configuration have a great 150 potential of influencing hydrology, soil erosion or water quality dynamics at different spatial and seasonal scales (Haidary et al., 2013; Jones et al., 2001; Kändler et al., 2017), some landscape metrics may have a high probability for collinearity. The collinear landscape metrics carry redundant information and are not independent predictor variables (Hargis et al., 1998). They can therefore result in biased or even misleading results when using conventional multivariate regression techniques like ordinary least-square regression, particularly in the case of a small number of observations (Shawul et al., 2019; Shi et

- 155 al., 2013). Compared to ordinary multivariate statistical methods which present relatively low robustness dealing with multicollinear variables, partial least squares regression analysis (PLSR) can overcome the limitation of multi-collinearity and achieve a robust performance by using techniques of multivariate statistical projection (Shi et al., 2013). The PLSR-Based on has widely been used to measure the "cause effect" relationships between land use changes and water resource, based on the powerful technique of projecting predicted and observed variables onto a new space and estimating the underlying structure /
- between projected spaces, PLSR facilitates an unbiased analysis of "cause-effect" relationships between land use changes and water resources components (Ferreira et al., 2017; Shi et al., 2013; Yan et al., 2013). Using an integrated approach of PLSR and hydrological modelling involving usingwith SWAT and PLSR, impacts of the multifacted land use changes in land use on various water resources components can be effectively identified. E.g., in the Upper Du catchment, China, Yan et al., (2013) observed that the farmland positively influenced streamflow and sediment yield, whereas forest area showed negative
- 165 correlation with them. as well as for sediment, whilebesides, Uurban expansion would cause streamflow to increase as well. Shi et., al (2013) indicated that the landscape metrics e.g., -Shannon's diversity index (SHDI), aggregation index (AI), largest patch index (LPI), contagion (CONTAG), and patch cohesion index (COHESION),-were the-important to impacts controleontrolling the watershed- soil erosion and sediment yield, and they altogether could explain contributing 65% and 74% ofto their variations inat soil-erosion and sediment yieldthem at in-subbasin levels, respectively.- Gashaw et al. (2018)
- 170 <u>identified that more shrubland would cause water yield and surface runoff to decrease while evapotranspiration and</u> groundwater flow to rise; However, increased cultivated land would result in decreases of groundwater flow and

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evapotranspiration in Blue Nile Basin, Ethiopia. The sameL-land use change contributes to the variations in annualaffects Ww differently, (e.g., evapotranspiration, base flow, and surface runoff) in different ways are related to land use changes in different ways (refs), .F, an increase of settlement areas would resulted in a higher proportion of surface runoff to streamflow

- 175 whileand lower groundwater flow (Marhaento et al., 2017). Further investigation is therefore needed to achieve a more effective and accurate means of quantifying comprehensive water quality and quantity consequences of land use changes to optimize catchment management. In summary, it has been demonstrated that PLSR is a suitable method forefficient -to quantifying distinguish complex impacts on water quantity and quality.
- The Stör River is the longest tributary of the Elbe River in the northernmost federal state of Germany, Schleswig-Holstein. Intensive agricultural activities (e.g., grazing, tillage, fertilizer, and pesticide application) are common in the catchment and increase the risk of water quality pollution (Monaghan et al., 2007). A variety of amelioration measures, e.g., tile drainage and straightening or canalizing of tributaries have been implemented in the past century to sustain agriculture productivity in lowlands dominated by shallow groundwater tables and abundant groundwater recharge. These activities brought about changes in the input and transport of nutrients and in hydrological fluxes. Meanwhile, the heterogeneity of the landscape
- pattern has been intensified due to artificial disturbances (Goldewijk and Ramankutty, 2004; Gu et al., 2007). We previously found significant relationships between land use patterns and water quality parameters at the landscape level in the upper Stör Catchment based on measurements (Lei et al., 2021). However, a A modeling approach allows for to investigating the analyze the dynamically and quantitativerepresent effects model the quantitative contribution of land use changes the land use, changes (composition and structure) measured by separate land use types classes on water quality and quantity, and it is necessary for facilitates developing effective informed and practicable strategies of improving water quality and controlling soil erosionachieving sustainable land and water management (Pott, 2014; Ripl et al., 1996).

To identify the key land use changes controlling the spatial and temporal variations in water quantity and quality, relationships between landscape characteristics of each land use type-class and water quality (represented by sediment, TP and TN) and quantity (represented by evapotranspiration, surface runoff, base flow, and water yield) are were explored at the subbasin scale in the upper Stör Catchment. To this end, the hydrological model SWAT and partial least squares regression (PLSR) are were employed. The study aims at (1) calibrating and validating a catchment model for streamflow, sediment, TP, and TN loads; (2) quantifying the changes of landscape characteristics and water quality and quantity variables at the subbasin scale; (3) investigating the relationships (depicted by the contribution and influence) between LUCC and water quality and quantity dynamics at the subbasin scale.

# 200 2 Materials and methods

# 2.1 Study area

The rural lowland catchment of the upper Stör is the focus of this study (Figure 1). It extends from the origin of the Stör River in Willingrade to the gauge in Willenscharen (Figure 1Figure 1) and is free of tidal influence. The catchment has a drainage area of approximately 462 km<sup>2</sup>, with a total length of the river network of about 221 km. Its temperate climate is characterized by an average annual precipitation of 850 mm and a mean temperature of 9.4 °C between 1990 and 2019, according to the records by weather stations Neumünster and Padenstedt (DWD, 2020a). The average daily streamflow measured at the catchment outlet in Willenscharen is 5.8 m<sup>3</sup> s<sup>-1</sup> between 1990 and 2019, with -low flows (mean value: 3.8 m<sup>3</sup> s<sup>-1</sup>) in summer (May-October) and high flows (mean value: 7.9 m<sup>3</sup> s<sup>-1</sup>) in winter (November-April) (LKN, 2020). Discharge occurring in the highest flow period (December-March) contributes most (around 50%) to the total annual amount of stream flow. The catchment is characterized by a flat topography, descending from nearly 60 m a.s.l. in the northeast and 85 m in the western part towards 20 m in the center and to 5-10 m in the southern part. Sandy soil (Cambisol, Gley-Podsol, Podsol) dominates the catchment, particularly in the central lowland part, while some Gley soils are mainly distributed in the east Formatted: Font color: Text 1

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and peat soils can be found in proximity to streams and near two major wetlands (Pott and Fohrer, 2017a). The catchment is

dominated by rural land use composed of arable land (36.1%) and pasture (31.3%), followed by forest (18.7%), urban areas (12.8%), and a minor fraction of water and wetland as indicated by a land use map for 2019 (Lei et al., 2021). The main 215 cultivated crops include winter cereals (wheat, barley, and rye), corn, and rapeseed.



Figure 1. Characteristics of the study area: Location of the upper Stör Catchment (a), spatial distributions of topography (b) (LvermA, 2008) and soil types (c) (Finnern, 1997), of subbasins, weather and gauging stations, and waste water treatment plants (WWTPs) (d) (Pott, 2014), as well as land use maps (e) (Lei et al., 2021; Rathjens et al., 2014; Ripl et al., 1996).

# 2.2 Land use data and landscape metrics

Land use maps for 1987, 2010, and 2019 have been were used to characterize changes in land use and landscape patterns. The earlier two maps (1987, 2010) have beenwere adapted from Ripl et al. (1996) and Rathjens et al. (2014), respectively, and are-were based on Landsat TM-5 image data at 30 m resolution. The land use map for 2019 has been-was derived from 10 m resolution Sentinel-2 satellite images (Lei et al. 2021). The land use types classes are were categorized uniformly as: 1) 225 arable land (winter cereals, corn, and winter rape, and other crops), 2) pasture (meadow, field grass, and rangeland); 3) forest (deciduous and coniferous forest); 4) urban (residential, commercial and industrial areas); 5) water (rivers, ponds, and lakes) and 6) wetland (Figure 1 Figure 1). Water and wetland are not considered for further analysis, as they comprise only minor and mostly constant percentages.

230 The area percentage of land use type class (PLAND) has been used as a measure of land use composition. Configuration metrics include the largest patch index (LPI), area-weighted mean shape index (AWMSI), area-weighted mean contiguity index (CONTIGAW), aggregation index (AI), and interspersion juxtaposition index (IJI), considering the dominance, shape, and interconnection of landscape (Ding et al., 2016; Gémesi et al., 2011). Composition and configuration indices of pasture, arable land, forest and urban have been were selected for subsequent analysis (Table 1 Table 1). They have been were derived 235 with the help of the software FRAGSTATS 4.2. All indices and their changes are were analyzed at subbasin scale.

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Attributes	Metrics	Unit	Description	Abbreviation at class level	Note		
Composition	Percentage of land use (PLAND)	%	Areal percentage of land use typesclasses	PLANDa, PLANDp, PLANDf, PLANDy			
Configuration	Largest patch index (LPI)	%	Percentage of the landscape composed of the	LPIa, LPIp, LPIf, LPIu	4-		Rormatted Table
Ū	Area-weighted mean shape index (AWMSI)	-	The sum of the mean shape index multiplied by the area weight of each patch type involving the	AWMSIa, AWMSIp, AWMSIf, AWMSIu	Metrics for land use type class	æ	
	Aggregation index (AI)	%	corresponding class Number of the same patch type being adjacent divided by the maximum number of adjacencies for the corresponding lund use class.	AIa, AIp, AIf, AIu	a (refers to arable land) p (refers to pasture), f (refers to forest),	,	Formatted Table
	Area-weighted mean contiguity index (CONTIGAW)	-	Measure of the patch shape based on the sum of spatial connectedness multiplied by the area weight of the patch for a certain class	CONTIGAWa, CONTIGAWp, CONTIGAWf, CONTIGAWu	<i>u</i> (refers to urban)		
	Interspersion juxtaposition index (IJI)	%	Measure of patch adjacency and interspersion or intermixing of patch types for a class	Ша, Шр, Шf, Шu			
2.3 Hydrol	ogic <u>al</u> and water quality	mode	eling				
2.3.1 SWA	T model						
The Soil a	nd Water Assessment Te	ool (S	SWAT) is a process-based and semi-	distributed eco-hydrolog	cal model with a		
continuous	time step (Arnold et a	1., 19	98). It is suitable for the simulation	n of streamflow, sedime	ent, nutrients, and		Field Code Changed
groundwate	er dynamics in catchments	s of di	fferent sizes (Aghsaei et al., 2020; Bieg	ger et al., 2014; Haas et a	l., 2016; Tigabu et		Field Code Changed
al., 2020).	The computation of water	routi	ng, nutrient cycles and soil erosion is	based on hydrologic resp	onse units (HRUs)		
characteriz	ed by the same land use,	soil t	ype, and slope in the same subbasin r	epresenting the spatial he	eterogeneity of the		
catchment	(Arnold et al., 2013). The	HRU	-based calculations for the subbasins a	re routed through the rive	rs that connect the		Field Code Changed
subbasins (	Neitsch et al., 2011).						Field Code Changed
To accurat	elv represent groundwate	r dvn	amics in this lowland catchment, we	applied SWAT <sub>35</sub> , the an	enhanced SWAT		
model <del>SW</del>	AT <sub>20</sub> s that is based on S	WAT	2012 Rev 582 (Pfannerstill et al 20	(14) In comparison to th	e standard SWAT	_	Formatted: Subscript
	lighting that uses the		EWAT was smaller three				Field Code Changed
model application that uses two aquifers, $-$ SWAT <sub>3*S</sub> uses employs three groundwater aquifers by subdividing and							Formatted: Subscript
subdivides the original shallow aquifer from SWAT into a fast and a slow aquifer. SWAT <sub>255</sub> was developed in the German							Formatted: Subscript
lowland ca	tchment of the Kielstau, 4	vhere	-toit better represented low flow period	<u>ls of streamflow s</u> and gr	oundwater storage		
and flow d	ynamics when compared t	the the	original SWAT version (Pfannerstill et	t al., 2014). It was <del>alread</del>	<u>-also</u> successfully		
applied to	the lowland catchment of	the T	Treene, -proving its usefulness for mod	delling nutrients as well	Haas et al., 2017;		Field Code Changed
Haas et al.,	2016).						
2.3.2 Mode	el databases and setup						
SWAT req	uires topography, soil, la	nd use	e, hydro-meteorological input data. To	pography data was obtain	ned from a Digital		Formatted: Font color: Text 1
Elevation N	Aodel (DEM) in 5 m reso	lution	(LyermA 2008) and used to delineate	the watershed into 21 s	bbasins Soil data	- 0	
and attribut	es for SWAT have been	vere d	erived by Pott and Fohrer (2017b) from	n a soil type man (Finner	n 1997) The land		
und attribu	2010 is use used to	build	the model. Three year eren rotation	n a son type map (i mier	wheat/acreated		
	or 2019 is was used to	bund	the model. Thee-year crop rotations	s (whiter wheat/whiter v			
ape/winter	wneau/corn; corn/corn/co	orn) <del>ar</del>	e-were adapted from Oppelt et al. (201	(2) and implemented for	he respective land		
use classes	. Agriculture managemer	nt sche	edules and fertilization (e.g., applicati	on rates of N, P fertilize	ers and manure at		
different cr	op growth stages) have b	een <u>we</u>	ere determined according to the actual	guidelines of agriculture	practices (KTBL,		
1995 and 2	008; Kühling, 2011; LW	K, 19	91 and 2011). From the DEM a four	slope classes (<1%, 1-2%	6, 2-5% and >5%)		
are were o	lefined. Slope, soil, and la	and us	e classes were combined to obtain 361	8 HRUs in the catchmen	t. The HRUs were		Formatted: Font: (Default) Times New Roman (Asian) Times New Roman
generated without excluding any HRUs by thresholds for land use, soil, or slope class percentages, to allow for a better							Formatted: Font color: Text 1
spatial representation. To accurately represent lowland hydrology, drainage tiles were considered based on the estimated						ſ	Formatted: Font: 9 pt
distribution	of drained areas in the c	atchn	nent (Venohr, 2000). We adapted drain	nage parameter values fo	r DEP_IMP (1200	15	Formatted: Tab stops: 9.79 cm, Left
mm), DDR	AIN (875 mm), TDRAIN	(24 h	a), and GDRAIN (61 h) from a previous	s modeling study in the ca	atchment (Pott and		Formatted: Font: 9 pt
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- Fohrer, 2017b). Waste water treatment plants (WWTP) were implemented as point sources using data from monthly measurement campaigns in 2009 and 2010 and WWTP data vary with space and seasons (Pott, 2014). Daily values of temperature (max. and min), solar radiation, humidity, and wind speed are available from 1990 to 2019 for the climate station Padenstedt (DWD, 2020b). Precipitation data are available for four stations (DWD, 2020b) (Figure 1). Daily streamflow is-was measured at the gauges in Padenstedt (PAD), Sarlhusen (SAR) and Willenscharen (WIL) from 1990 to 2019 (LKN, 2020). Daily sediment and nutrient data have beenwere both obtained during two measurement campaigns, i.e., between August 2009 and August 2011 and between October 2018 and November 2019 in Willenscharen. Daily mixed samples have beenwere taken by an automatic and cooled sampler from a depth of 0.30 m above the river bed at the central section of the stream. They have beenwere analyzed according to German standard procedure for water analysis (DEV) (Einheitsverfahren, 1997) in the laboratory of Department of Hydrology and Water Resources Management at Kiel University. Total suspended sediment concentration has beenwas measured by filtering 1 l of water sample through 0.45 µm celluloseacetate filter paper and drying at 105°C. The concentration of total phosphorus (TP) has beenwas determined by
- spectrophotometry, according to DEV H36 and DEV D11, while total nitrogen (TN) has beenwas measured by chemiluminescence detection according to DEV H3. Each measurement of TP or TN concentration from unfiltered samples has beenwas performed based on a blank comparison analysis of distilled water and triplicate analysis of subsamples. Their concentrations have beenwere determined by the arithmetic mean values of any two subsamples with smallest measurement differences (less than <10%). Based on the measurements of daily -concentration and streamflow, the respective daily loads of sediment, TP, and TN were calculated-calculted, respectively.</li>

# 2.3.3 Model calibration and validation

A step-wise calibration approach has been applied for dThe variables daily streamflow (1), sediment (2), TP (3), and TN (4)

- 290 data were calibrated separately and stepwise. T, and the number in the parentheses right after them denotes their respective calibration order, i.e., This means that streamflow was firstly calibrated first, followed by sediment, and then TP, and TN. RelevantAn overview of details of calibration and validation details for each variable is provided -are summarized in Table 2. Streamflow was calibrated using a fifteen year time period from 1990 to 1991 and from 2007 to 2019. The other available fifteen years (1992-2006) have been used for validation. This
- 295 Preliminary parameter ranges (Table S1)-were selected based on experiences with the SWAT model in the Stör Catchment (Pott and Fohrer, 2017b) and other German lowland catchments (i.e., Kielstau and Treene catchments) (Haas et al., 2016; Lam et al., 2012; Pfannerstill et al., 2014) as well as in relevant studies from other countries (Aghsaei et al., 2020; Boongaling et al., 2018). The final ranges of selectedcalibrationed parameters (Table S1) were determined based on the sensitivity of parameters to model outputs as derived from 2000 trial runs following the method used by Guse et al. (2020), in which model simulations are iteratively repeated with successively constrained parameter ranges.
- in which model simulations are iteratively repeated with successively constrained parameter ranges.
   Parameter sets were generated from the derived parameter ranges using Latin Hypercube Sampling in the R-package FME (Soetaert and Petzoldt, 2010). For each of these 8000 (streamflow) and 5000 (sediment, TP, and TN loads) independent parameter sets, -model runs were conducted each involving a warm-up period (four years)-respectively, and evaluated using multiple performance criteria to select the best parameter set. To this end, -the objective functions Nash-Sutcliffe efficiency
- 305 (NSE), Kling-Gupta Efficiency (KGE), and Percent Bias (PBIAS), which were proposed in Guse et al. (2014) and Moriasi et al. (2007), were applied. For an accurate representation of all segments of the hydrograph (very high, high, middle, low, and very low periods), the additional signature measure RSR (Ratio of Root Mean Square Error to the Standard Deviation of the Observations) was used (Haas et al., 2016; Zambrano-Bigiarini, 2020). The definition of each objective function is provided in Text S1 in the supplementary materials-information.
- First, streamflow was calibrated at three gauges. -The two upstream gauges Padenstedt (PAD) and Sarlhusen (SAR) were used to select the best parameter sets for the respective sub-catchments first (Figure 1). Then, the best parameter set for the area downstream of PAD and SAR and upstream of the outlet gauge Willenscharen (WIL) was selected. For each of the

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three streamflow gauges, we pre-selected the parameter sets that yielded a KGE >0.75 for the streamflow calibration period. To accurately represent streamflow dynamics during the periods of water quality measurements (Aug. 2009 - Aug. 2011 and Oct. 2018 - Nov. 2019), the mean RSR for the five flow duration curve (FDC) segments -during these periods was assessed

- Oct. 2018 Nov. 2019), the mean RSR for the five flow duration curve (FDC) segments -during these periods was assessed and the best 300 streamflow parameter sets indicated by a low RSR were selected. From these 300 sets, the final parameter set yielding the highest KGE in these periods was selected. Calibration and validation periods (Table 2) were defined based on an equal representation of dry, normal, and wet years according to the annual precipitation.
- Second, with the derived set of best hydrological parameters, -model runs for 5000 different sediment calibration parameter
   sets were carried out and the best model run was selected based on the highest NSE. Third, this model was run for 5000 different sets of TP calibration parameters and the best model run was similarly selected using the NSE. Finally, based on the so far derived best parameters, another 5000 model runs for TN calibration were carried out and the best model run indicated by the highest NSE was selected. To accurately represent peak loads and their dynamics, the NSE was selected as single criterion for the water quality variables. The calibrated parameters are provided in Table S1 in the supplementary information.—Evaluation and processing of the model data were carried out in R using the packages hydroGOF (Zambrano-

Bigiarini, 2020) and zoo (Zeileis and Grothendieck, 2005). <u>The split of the streamflow observations were split to data ensures an equal representation of dry, normal, and wet years in</u> the calibration and validation period, according to the annual precipitation. <u>As for the multi-gauge calibration, Ffirst, data</u> from the two upstream gauges Padenstedt (PAD) and Sarlhusen (SAR) have been<u>were used to calibrate parameters in the</u>

- 330 respective subcatchments (Figure 1). Then, the parameters for the area downstream of PAD and SAR and upstream of the outlet gauge Willenscharen (WIL) have been<u>were</u> calibrated. Sediment, TP and TN loads have been calibrated for two hydrologic years (sediment: 30/10/2009 07/08/2011; TP, TN: 08/08/2009 10/08/2011) using a model with the land use map in 2010 and validated for one hydrologic year (19/10/2018 05/11/2019) using a model with the land use map in 2010 and validated for one hydrologic year (19/10/2018 05/11/2019) using a model with the land use map in 2010 and validated for one hydrologic year (19/10/2018 05/11/2019) using a model with the land use map in 2010 and validated for one hydrologic year (19/10/2018 05/11/2019) using a model with the land use map in 2010 on the daily data from Willenscharen. The calibration has been performed based on 8000 (stream
- flow) and 5000 (sediment, TP, and TN loads) parameter sets generated using Latin Hypereube Sampling method (Soctaert and Petzoldt, 2010), For each parameter set a model run has been performed, allowing for a warm up period of 4 years. From experiences with the SWAT model in the Stör Catchment (Pott and Fohrer, 2017b) and other German lowland eatchments (i.e., Kielstau and Treene catchments) (Haas et al., 2016; Lam et al., 2012; Pfannerstill et al., 2014) as well as in relevant studies from other countries (Aghsaei et al., 2020; Boongaling et al., 2018), the parameters most likely to affect
- 340 hydrological and water quality processes have beenwere selected and their preliminary ranges have beenwere defined (Table 2). The final ranges of selected parameters have beenwere determined based on the sensitivity of parameters to model outputs as derived from 2000 trial runs following the method used by Guse et al. (2020), in which model simulations are iteratively repeated with successively constrained parameter ranges to obtain more precise parameter identifiability and improve model performance. The calibrated parameters and ranges were provided in (Table S1 in the supplementary informationTable 2). Calibration and validation have beenwere carried out in R using the packages FME (Soctaert and
- Petzoldt, 2010), hydroGOF (Zambrano Bigiarini, 2020) and zoo (Zeileis and Grothendieck, 2005). The performances for modeling streamflow and sediment, TP and TN loads have beenwere assessed using <u>a multi-metric</u> approach based on the objective functions Nash Sutcliffe efficiency (NSE), Kling-Gupta Efficiency (KGE), and Percent Bias (PBIAS), which were as proposed in Guse et al. (2014) and Moriasi et al. (2007). The definition and algorithm of each
- objective function are provided in Text S1 in the supplementary information. For an accurate representation of all phases of flow hydrograph for water quality simulation periods, the additional signature measure RSR (Ratio of Root Mean Square Error to the Standard Deviation of the Observations) was used to calibrate the very high, high, middle, low, and very low periods (Haas et al., 2016; Zambrano-Bigiarini, 2020). For each of the three streamflow gauges, we pre-selected the parameter sets that yielded a KGE >0.75 for the streamflow calibration period. To particularly represent runoff dynamics during the periods of water quality measurements (Aug. 2009 – Aug. 2011 and Oct. 2018 – Nov. 2019) well, the mean of
- RSR for the five flow duration segments during these periods was assessed and the best 300 streamflow parameter sets

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indicated by a low RSR were selected. From these 300 sets, the final parameter set yielding the highest KGE in these periods was selected. For sediment, TP, and TN calibration the parameter set that yielded the highest NSE during the calibration period was selected to represent peak loads and their dynamics well, as peak loads of them are often underestimated and NSE relates more to the peak values.

	Calibration				Validation			
	Streamflow	Sediment load	TP load	TN load	Streamflow	Sediment load	TP load	TN load
Simulation Evaluation period period	<u>1990-1991;</u> 2007-2019	<u>30/10/2009-</u> <u>7/8/2011</u>	<u>8/8/2009-</u> 10/8/2011	<u>8/8/2009-</u> 10/8/2011	<u>1992-2006</u>	<u>19/10/2018-</u> <u>5/11/2019</u>	<u>19/10/2018-</u> <u>5/11/2019</u>	<u>19/10/2018</u> <u>5/11/2019</u>
<u>Simulation</u> period	<u>1986-2019</u>	2005-2011	2005-2011	2005-2011	<u>1986-2019</u>	<u>2014-2019</u>	<u>2014-2019</u>	<u>2014-2019</u>
<u>Land use</u> map	<u>2019</u>	2010	<u>2010</u>	2010	<u>2019</u>	2019	2019	<u>2019</u>
Gauges	PAD/SAL/WIL	WIL	WIL	WIL	PAD/SAL/WIL	WIL	WIL	WIL
Calibration runs	8000	<u>5000</u>	<u>5000</u>	<u>5000</u>				
<u>Performan</u> ce criteria	KGE>0.75 in 1990- 1991; 2007-2019 & best KGE among 300 best mean RSR of FDC in 8/8/2009- 10/8/2011; 19/10/2018- 5/11/2019	<u>NSE</u>	<u>NSE</u>	<u>NSE</u>				

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arameters	Definition	Calibrated ra	inge		Calibrated	value	
arameters used to-	calibrate streamflow						
		WILL	SAR	PAD	WILL	SAR	PAD
SURLAG	Surface runoff lag coefficient	0.1-0.6	0.1-0.6	0.1-0.5	0.13	0.13	0.13
_GWDELAY <sub>fsh</sub>	Groundwater delay time fast shallow aquifer (days)	4 <del>8-85</del>	<del>40-80</del>	<del>65-100</del>	83	<del>42</del>	71
_ALPHABFfsh	Baseflow alpha factor - fast shallow aquifer (day+)	0.17-0.38	0.18-0.38	0.05-0.22	0.18	0.26	0.08
RCHRGssh	Aquifer percolation fraction – slow shallow aquifer	0.8-0.94	0.08-0.58	<del>0.38-0.7</del>	0.91	<del>0.34</del>	0.44
GWDELAY ss	Groundwater delay time - slow shallow aquifer (days)	<del>68-105</del>	<del>58-100</del>	80-120	80	<u>92</u>	87
ALPHABFssh	Baseflow alpha factor slow shallow aquifer (day-1)	0.0009-0.002	0.001-0.007	0.003-0.009	0.0019	<del>0.0036</del>	0.006
RCHRGdp	Aquifer percolation fraction inactive deep aquifer	0.02-0.15	0.015-0.14	<del>0.1-0.45</del>	0.14	0.03	0.15
ESCO	Soil evaporation compensation factor	0.85-0.98	0.93-1	0.7-0.95	0.86	<del>0.94</del>	0.77
EPCO	Plant uptake compensation factor	0.01-0.025	0.05-0.22	0.1-0.35	0.02	<del>0.06</del>	0.23
s_CN2	Initial SCS runoff curve number for moisture condition II	-13 1	-121	-122	-5.64	-3.27	<del>-4.89</del>
SOL_AWC	Available water capacity of the soil layer (mm)	-0.06 - 0.02	-0.060.01	<del>-0.04 - 0.03</del>	-0.006	-0.020	0.001
-SOL_K	Saturated hydraulic conductivity (mm h+)	0.7-1.3	0.8-1.2	0.8-1.2	1.052	<del>0.811</del>	1.079
arameters used to	calibrate sediment						
ADJ_PKR	Peak rate adjustment factor for sediment routing in the main channel	<del>0.55-2</del>			<del>0.61</del>		
CH_COV_1	Channel crodibility factor	0.1-0.5			0.41		
CH COV 2	Channel cover factor	0.4-0.7			0.57		
USLE_P	USLE support practice factor	0.5-1			0.93		
-SLSUBBSN	Average slope length (m)	0.8-1.08			0.88		
HRUSLP	Average slope stepness (m m <sup>+</sup> )	0.95-1.28			1.1		
LAT SED	Sediment concentration in lateral and groundwater flow (mg 1-4)	55-140			110		
USLE K	Soil crodibility (K) factor	0.06-0.2			0.09		
SOL Z	Depth from soil surface to bottom of layer (mm)	-70-20			-65		
USLE_C	Minimum value of USLE C factor for land cover/plant	0.08-0.43 (cro	pland); 0.002-0.0	17 (pasture)	0.192 (crop	land), 0.015	(pasture
arameters used to-	calibrate total phosphorus						
P_UPDIS	Phosphorus uptake distribution parameter	30-100			73.61		
PPERCO	Phosphorus percolation coefficient	<del>10-16</del>			10.3		
PHOSKD	Phosphorus soil partitioning coefficient	115-190			181.14		
PSP	Phosphorus sorption coefficient	0.01-0.5			0.21		
ERORGP	Organic P enrichment ratio	0.8-4.8			2.38		
_GWSOLP	Concentration of soluble phosphorus in groundwater contribution t stream flow from the subbasin	ө 0.04-0.4			<del>0.19</del>		

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r_SOL_SOLP	Soluble phosphorus concentration in the soil layer (mg kg+)	<del>30-90</del>	32.1
Parameters used to	o calibrate total nitrogen		
r_CMN	Rate factor for humus mineralization of active organic nitrogen	0.001-0.003	0.002
r_RCN	Concentration of nitrogen in rainfall (mg 1-1)	<del>1.3-6</del>	5
r_CDN	Denitrification exponential rate coefficient	0.09-0.18	0.16
r_N_UPDIS	Nitrogen uptake distribution parameter	20-90	<del>69.05</del>
r_NPERCO	Nitrogen percolation coefficient	0.03-0.5	0.06
r_SDNCO	Denitrification threshold water content	<del>0.3-0.95</del>	0.95
r_HLIFENGWfsh	Half-life of nitrate in fast shallow aquifer (days)	30-125	52
r_HLIFENGWssh	Half-life of nitrate in slow aquifer (days)	250-480	454
r SHALLSTNeeh	Initial concentration of nitrate in slow aquifer (mg. 1-4)	30-85	37.41

# 2.3.4 Model application

365 Applying the best respective parameter sets, Ithe model has been was run for three land use scenarios. each Each scenario simulation iswas run from 1990 to 2019 using of the three one of the three land use maps (in 1987, 2010, and 2019)-from 1990 to 2019. As agriculture in 1987 was generally classified, it has been was split as corn (12%), rapeseed (29%), and wheat (59%) randomly distributed in the catchment in SWAT model, according to the statistical data from Schleswig-Holstein Statistical Office (1992-2012). For the three scenario simulations, All-all other inputs i.e. DEM, soil data, weather data, waste water quality data, management practices, and fertilization have been were kept constant, and the calibrated parameters 370 have beenwere adapted. Hence, each model run is performed under a different land use scenario defined by one of the three land use maps. The respective differences in the mean annual value of results each response water resources variable (i.e., actual evapotranspiration (ET), surface runoff (SQ), base flow (BF), water yield (WYLD), sediment (SED), TP, or TN load) arewere obtained by comparing the results from two scenario model runs (see Text S2 and -S3) between from any thetwo se model runs using any two of the three land use maps. They can be referred to as the respective variationschanges driven by 375 land use changes during the corresponding periods such as of 1987-2010, 2010-2019, and 1987-2019. The modeleded results have been were used to explore the influences of land use changes (LUCC) on the variationschanges in eachthe water resources response, variables, actual evapotranspiration (ET), surface runoff (SQ), base flow (BF), and water yield (WYLD) as well as on sediment (SED), TP, and TN loads. Based on the model results, Furthermore, the contributions of LUCC on 380 changes in ET, SQ, BF, and WYLD as well as in SED, TP, TN at the subbasin scale are-were evaluated, and key impacts

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# 2.4 Partial least squares regression

from LUCC are-were identified.

Combining the features of principal component and multiple linear regression analyses, partial least squares regression (PLSR) is a robust multivariate analysis method of determining the relationship between two sets of variables. It is powerful to even when dealing with multi-collinear predictor variables. The principle of PLSR is to extract a few latent components from original predictor variables that carry as much variation as possible, and which are meanwhile most likely to predict the variation in the response variable. Detailed information on the underlying theory and algorithms of PLSR is available in Abdi (2010).

- In this study, PLSR is-was used to reveal the contribution of changes in land use types-classes on the variation in ET, SQ, 390 BF, WYLD, SED, TP, and TN across three time steps (1987, 2010, and 2019). The predictor variables are-were the absolute changes in area percent (PLAND) and landscape metrics (LPI, AWMSI, AI, CONTIGAW, UI) of four main land use types mean annual values of ET, SQ, BF, WYLD and SED, TP, and TN loads at the subbasin scale modelled under different land use betweenconditions conditions in 1987, 2010, and 2019. PLSR models for all of these response variables were 395 constructed. The absolute change in each land use indicator iwas calculated using equation (5)-(7) while that in each of water
- resourcesresponse variable iwas calculated using equations (8)-(10) as shown in Text S3 in the supplementary informationmaterials. A cross-validation is-was performed with 50 random repetitions on 10 equal segments of the data set.

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It is way used to determine the number of optimal components of the PLSR model to obtain a desirable balance between the explained variation in the response ( $\mathbb{R}^2$ ) and predictive power of the model (measured as cross-validated goodness of the prediction:  $\mathbb{Q}^2$ ). The cumulative predictive ability (cumulative goodness of prediction:  $\mathbb{Q}^2_{cum}$ ) and the cross-validated root mean squared error (RMSECV) as the difference between actual and predicted values, are-were determined for each model (Yan et al., 2013). The regression coefficients (RCs) signify the direction and extent of the effect of LUCC predictor variables. The variable importance for the projection (VIP) quantifies the importance of the predictors. By-According to the Wold's assessment criterion, a predictor with VIP<0.8 is assessed as less important (Boongaling et al., 2018; Wold et al., 2001). To achieve model parsimony, the following PLSR modelling procedures has beenwere conducted: First, an initial

405 2001). To achieve model parsimony, the following PLSR modelling procedures has beenwere conducted: First, an initial simulation of PLSR is run using all predictors. Next, new PLSR models are run by iteratively excluding the predictor with small variable importance (VIP) until the modelling procedure resulted in acceptable variable importance or only two predictors remained. The number of components of candidate PLSR model was determined so that the Q<sup>2</sup><sub>cum</sub> is maximized (Shi et al., 2013).

410 All the PLSR analyses were performed with the R packages pls (Mevik et al., 2020) and mdatools (Kucheryavskiy, 2020).

# 3 Results and discussion

# 3.1 Model performances for Ecalibration and validation periods

	The simulated and measured daily values of streamflow (Figure 2) and water quality (Figure 3) data are visually compared
	for the calibration and validation periods, and the statistical performance of the models is assessed (Table 3). As shown in
415	Table 3, for streamflow, the The model obtains a NSE and KGE values both above of 0.76 0.810.75 and, a KGE and
	absolute PBIAS values below or nearlyslightly above 10%, of 0.82 0.85 for streamflow at the two upstream gauges
	Padenstedt and Sarlhusen, and a slightly higher NSE (calibration: 0.79, validation: 0.79) and KGE (calibration: 0.88,
	validation: 0.87) for streamflow at the outlet in Willenscharen. The PBIAS values are within the range of -2.2% - 10.6%.
	These values indicates a good to very good model performance for depicting daily streamflow in the catchment according to
420	the assessment criteria for model evaluation <del>SWAT model performance criteria</del> (Moriasi et al. 2007). For daily TN

- LoadLikewise, the model shows a nearly goodsatisfactory to very good performance <u>for daily TN load</u>, indicated by an NSE withinbetween 0.64 and 0.86 of 0.64 for calibration and of 0.86 for validation and by a KGE ≥ 0.71 (for calibration: 0.71; for validation: 0.91), while absolute values of PBIAS values are below 15%. For sediment and TP the model shows a lower performance. For Sediment load, the model achieves a satisfactory to good performance as indicated by NSE (0.54 0.65) and PBIAS (-22.2 12%) values. during calibration (NSE = 0.54, KGE = 0.58, PBIAS = 12%) and a good performance
- during the validation period (NSE = 0.65, KGE = 0.59, PBIAS = -22.2%). For TP-The model simulates TP load with an unsatisfactory (validation) to satisfactory (calibration) performance, which is assessed by NSE below and above 0.5, respectively. the model obtains a satisfactory performance for calibration (NSE =0.56) but an unsatisfactory performance (NSE =0.29) for validation. The worse TP model performance may be due to the short and possibly different conditions
- during calibration and validation periods. Nevertheless, PBIAS for TP model is still within the acceptable performance range (±40 ≤ PBIAS < ±70) (Moriasi et al., 2007). It should be noted that the performance ranges from Moriasi et al. (2007) refer to a monthly time step, whereas we used a daily time step, a finer temporal scale (daily step), on which it is usually harder to achieve a good model representation (Pfannerstill et al., 2014; Tan et al., 2021). We therefore conclude that even for daily TP the model performance is acceptable, particularly with regard to the study purpose of analyzing long-term changes in the water and matter balance.</li>

According to a detailed model performance assessment using multiple metrics (shown in Text S2 in supplementary information), the model depicts long term dynamics of daily streamflow observations with a very good performance and TN load with nearly good performance, sediment with satisfactory performance. However, the model simulates TP load with a relatively lower performance which is still considered acceptable.



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Overall, modeled and measured daily values show clear consistency in their dynamics (Figure 2 Figure 2 and 3). Differences 440 mainly appear for low flow periods in summer and particularly for a few peak flows in winter, or low flow periods in summer. As already indicated by the goodness of fit measures (Table 3), the modeled streamflow matches the measured values most of the time from 1990 to 2019. HoweverSpecifically, a few single-flood peaks are underestimated in winter, e.g. on 27-28/Feb/2002, 5-6/Jan/2012, and 24-25/Dec/2014. This might be related to an insufficient representation of snow in the model, or and deficiencies in single-event flood routing in the model (Lam et al., 2012). The underestimation of peak 445 streamflow in winter was also observed in other rural lowland catchments of Treene (Haas et al., 2016) and Kielstau (Lam et al., 2010) in northern Germany. Sediment loads are overestimated during the calibration period and slightly underestimated during the validation period mainly for a few peak values. The incorrect estimation might be due to the fact that river sediment load is also influenced by tile drains and bank erosion in lowland catchments (Kiesel et al., 2009), while SWAT 450 primarily\_takes into account sheet erosion. A few sediment peaks in early March 2010, mid-Jan 2011 and mid-Feb 2019 are underestimated butNevertheless, -othersome peaks e.g. in Nov, Dec 2009, and Mar 2019 are very well depicted. A similar behavior ean beis observed for modelling TP load during the calibration and validation periods, with slight overestimation

of TP in summer (April - June of 2009 and 2019) and underestimation of a few peaks in winter (between-November -and March). TN is generally well represented, except for only a few underestimations of extreme peaks in winter (e.g., early 455 March or November 2010), mid March 2019). Overall, the underestimation of some peak loads of sediment, TP and TN might be attributed to the underestimation of corresponding peak flows.



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Figure 3. Comparisons between measured and modeled daily loads of sediment, total phosphorus (TP), and total nitrogen (TN), respectively for calibration (left) and validation (right) periods

Table 3. Performance metrics for the model calibration and validation periods.

	Calibration				Validation	Validation			
Index	Streamflow (PAD/SAR/WIL)	Sediment load	TP load	TN load	Streamflow (PAD/SAR/WIL)	Sediment yield	TP load	TN load	
Period	1990-1991; 2007-2019	30/10/2009- 7/8/2011	8/8/2009- 10/8/2011	8/8/2009- 10/8/2011	1992-2006	19/10/2018- 5/11/2019	19/10/2018- 5/11/2019	19/10/2018- 5/11/2019	
KGE	0.85/0.82/0.88	0.58	0.65	0.71	0.84/0.85/0.87	0.59	0.22	0.91	
NSE	0.76/0.78/0.79	0.54	0.56	0.64	0.81/0.81/0.79	0.65	0.29	0.86	
PBIAS (%)	5.6/-2.2/0.3	12	-4.7	-11.5	0.7/10.6/7.2	-22.2	-46.2	5	

# 3.2 Characteristics of land use change

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Land use changes between 1987 and 2019 vary across the catchment (Figure 4Figure 4). This is indicated by the individual / dynamics in the four main land use types-classes of arable land, pasture, forest, and settlement area. Arable land has been decreasing and primarily replaced by pasture (by 16.2% of the catchment, dark cyan in Figure 4Figure 4). The decrease of / arable land is more pronounced-in the northeast (e.g., subbasins 3 and 9-11) is more pronounced than in the northwestern part (e.g., subbasins 2, 4, 6, 8) where pasture was sometimes converted to arable land (dark pink, Figure 4Figure 4). / Conversely, pasture shows an increasing trend over the period of observation. The increase is stronger in the east is stronger as compared to the west of the catchment (Figure 4Figure 4 and 5). The change of pasture is in part associated with the / stream restoration including stabilizing river shore and increasing riparian vegetation (Dickhaut, 2005; Gessner et al., 2010). Besides, agricultural grasses may have been included in the pasture class due to the classification approach. Forest also

shows an increasing trend as indicated by green colors in Figure 4, with a strong increase in the lowlands of the middle (subbasins 6 and 13) and southern parts (subbasin 17, Figure 5Figure 5). Urban area has expanded mainly around the city of Neumünster (subbasin 15 and 17) (Figure 5Figure 5).

480 In addition, the subbasin-scale land use metrics varied substantially between 1987, 2010, and 2019 (Figure 6Figure 6). The mean area percent (PLAND) per subbasin declined for arable land (APLAND) by 16% and 3% during the periods of 1987-2010 and 2010-2019, respectively. In contrast, subbasin-averaged pasture (PPLAND) increased for the period of 1987-2010

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by 12% but decreased slightly from 2010 to 2019 by 0.8%. Both forest (FPLAND) and urban (UPLAND) areas have steadily increased from 1987 over 2010 to 2019. Similar trends are found in the metrics of the percentage of largest patch index (LPI) and the interspersion juxtaposition index (IJI). The subbasin average of LPI for arable land has decreased by 20% from 1987 to 2019, whereas the LPIs of other land use types-classes shows a slight and stable increase. The IJI of arable land displays an overall slight increase from 1987 to 2019, while the IJI values of other land uses have steadily and notably increased (with a net increase up to over 20%). Both the area-weighted mean contiguity (CONTIGAW) and aggregation (AI) of each land use type-class have decreased over time, whereas the area-weighted mean shape index (AWMSI) has continuously and slightly increased. Despite similar changing directions of the land use patterns in the periods of 1987-2010 and 2010-2019, land use has been subject to more alterations in the former period than in the latter. Additionally, CONTIGAW, AI, and IJI of arable land exhibited opposite trends in the two periods, with a decrease from 1987 to 2010, and a slight increase from 2010 to 2019.



495 Figure 4. Spatial distribution of land use changes between 1987 and 2019 in the Stör Catchment. Individual land use change types are marked by different colors. The percentage of each change type calculated as percentage of the catchment area is given in the parentheses. The strongest change is marked in bold.



Figure 5. Spatial distribution patterns of the changes of each land use type between 1987, 2010, and 2019.

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# 3.3 Differences of changes in water quantity and quality

505 Using the results from the three different <u>scenario</u> model runs based on three land use maps of 1987, 2010, and 2019, we calculated changes in water quantity and quality. The spatial distribution of the variations in modeled subbasin-scale actual

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evapotranspiration (ET), surface runoff (SQ), base flow (BF), water yield (WYLD), and loads of sediment (SED), total phosphorus (TP), and total nitrogen (TN) between 1987, 2010, and 2019 are shown in <u>Figure 7Figure 7</u>. ET and SQ are mostly characterized by increases of up to 10.8 mm and 11.4 mm, respectively from 1987 to 2019, with slight decreases by up to 3.8 mm in several subbasins between 2010 and 2019. The most significant increase in ET occurs in subbasins which show a larger increase in forest from 1987 to 2019, such as subbasins 8, 12 and 17 (Figure 5Figure 5). SQ shows a stronger increase in the middle-western subbasins whichsubbasins that <u>, which</u> experienced larger expansion of urban areas (Figure 5Figure 5), with the strongest increase of SQ occurring in subbasins 15 and 17 that experienced the largest increase of urban area between 1987 and 2019. This might be attributed to the increased surface sealingimpervious surface which facilitates the generation of surface runoff and reduces confluence time (Anand et al., 2018; Sood et al., 2021). Contrarily, BF and WYLD have decreased by up to 20 mm and 13 mm, respectively in most subbasins in the periods 1987-2010 and 1987-2019. However, the slight increase in a few subbasins in central part of the catchment exhibit a slight increase in base flow, which is probably attributed to a greater contribution of shallow groundwater in the central lowland areas to low flow periods than in the steeper eastern and western steeper areasexplains the higher possibility of groundwater recharge to low flow

- 520 proportions of streamflow there as compared to the the surrounding steeper areas. The loads of SED, TP, and TN show notable decreasing trends from 1987 to 2019. Pronounced reductions of SED (7.8-18.2 t km<sup>-2</sup>) occur in the relatively steeper northeastern corner (e.g., subbasins 3, 9-10) and the southwestern corner (e.g., subbasins 5 and 12) and subbasin 17, while the decrease is weaker in the mid-west. Overall, the changes in TP and TN loads show a weak decrease in the (mid) west and more pronounced decreases in the east and steeper southwest of the catchment (Figure 7Figure 7). The spatial differences
- 525 illustrate that the decrease of SED, TP, and TN is stronger in steeper subbasins decrease to a greater extent than in lowland subbasins due to land use changes, may be related to the more intense exchange between groundwater and surface water and a higher contribution of nutrients from groundwater to stream in lowlands. which is partly explained by a higher fraction of contribution from groundwater to soil particles and nutrients of streamflow in middle lowland areas. The most pronounced net decrease of TP and TN loads are observed in subbasins 12 and 17, corresponding to the largest decrease of arable land
- 530 percentage (50% in subbasin 17, 30% in subbasin 12) between 1987 and 2019. The single subbasin that has experienced a slight increase of sediment or TP load is subbasin 1, which is characterized by the least reduction of arable land and minor decrease of forest. The most significant decrease in nutrients and sediment has occurred in subbasins which have experienced notable increases of pasture or forest and a decrease of arable land, e.g., subbasins 12 and 17 (Figure 5Figure 5). Overall, variations in surface runoff, sediment, TP, and TN are depicted by spatially explicit patterns on the subbasin scale. It
- 535 is necessary to consider this spatial heterogeneity, when establishing management measures in order to improve water quality.

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# 540 3.4 Influences of changes in land use metrics on water quantity and quality

# 3.4.1 Contributions of LUCC to variations changes in water quantity and quality

A summary of the PLSR models separately constructed for ET, SQ, BF, WYLD, SED, TP and TN, is provided in <u>Table 4</u>Table 4. The prediction plots for the seven variables by applying the PLSR models are shown in Figure 8. The changes in water quantity and quality could be reasonably explained by the constructed PLSR models (0.61<R<sup>2</sup><0.88, 0.57<Q<sup>2</sup><0.85, Table 4). The comparison of the actual and predicted values (in Figure 8) illustrates the accuracy of the model calibration and cross-validation. For the ET and WYLD models, the percentage of unexplained variation decreases with increasing number of components, whereas the prediction error of cross-validated observations (indicated by cross-validated root mean squared error, RMSECV) is minimal with one or two components, respectively. This indicates that adding more components does not improve the correlation with the residuals of the response variables (Onderka et al., 2012). Overall, 60.5% and 68.3% of the variations in the changes in ET and WYLD can be explained by the first component and the first two components, respectively. Adding other components does not strongly increase the cumulative explained variations (only by +4.2-5.4%) in ET and WYLD changes from 1987 to 2019 (<u>Table 4Table 4</u>). For SQ, two components are extracted for the PLSR model, with 58.9% of variation is explained on the first component and cumulative explained variations increase to

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81.3% when adding the second component. For all other variables, the minimum RMSECV is achieved with models using

- five components. For base flow, 37.4% of the variation in the dynamics is explained by the first component, cumulatively 64.2% adding the second component, and ultimately 87.7% with a consecutive addition of third, fourth, and fifth component. For the changes in loads of sediment, TP, and TN, the first component of the models always explains the majority of the variation (43.7-63%, <u>Table 4Table 4</u>). With all water quality variables together, approximately 75% of the variation changese is accurately explained on average.
- 560 Approximately 70-80% of the variations in water quantity and quality dynamics were explained by LUCC, underlining the importance of LUCC on catchment water resources. Better explanations (over 81%) of SQ and BF by LUCC confirmed the significant influences of landscape heterogeneity on surface runoff and groundwater dynamics (Kändler et al., 2017; Xu et al., 2020; Zhang and Schilling, 2006). Only a quarter of the variations in sediment, TP, or TN cannot be interpreted by LUCC, which demonstrates that changes of rural landscape patterns are essentially important in controlling nutrients.
- 565 pollution. The proportion and spatial arrangement of agriculture land play an important role in the generation and transportation of nutrient pollutants as previously reported in different catchments worldwide; e.g., Zhang et al. (2020b) found that agricultural cultivation on steeper hillsides intensified N and P entries in ponds in the hilly Tianmu Lake catchment of Eastern China. Gémesi et al. (2011) identified the cohesion and contagion of cropland were more important than other land use indicators to account for the variability in TN and TP in the relatively plain Lowa Lake catchment of the
- 570 <u>central USA.</u> –The minor unexplained fraction may be attributed to potential changes in waste water treatment which sometimes remained constant in our modeling approach. Lower explanation of TP may be additionally due to the lower SWAT model performance for TP, the susceptibility of P to soil or geomorphology properties (Maranguit et al., 2017; Noe et al., 2013). More than 60% of the variations in ET and WYLD are explained by LUCC. The unexplained fraction may be attributed to the different contributions from specific crops (included in SWAT) and the lumped agriculture land use class as well as the compensating effect of subbasins (Wagner et al., 2013).

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Table 4. Summary of the PLSR models of evapotranspiration (ET), surface runoff (SQ), base flow (BF), water yield (WYLD), sediment yield (SED), total phosphorus load (TP) and total nitrogen load (TN) at subbasin scale.

Response Variable Y	$\mathbb{R}^2$	$\mathbf{Q}^2$	Component	Explained variability in Y (%)	Cumulative explained variability in Y (%)	RMSECV	$Q^2_{cum}$
E.E.	0.61	0.57	1	60.5	60.5	2.32 (mm)	0.568
EI			2	2.4	62.9	2.35 (mm)	0.558
			3	1.2	64.1	2.44 (mm)	0.524
			4	0.2	64.3	2.41 (mm)	0.535
			5	0.4	64.7	2.41 (mm)	0.534
SQ	0.81	0.78	1	58.9	58.9	1.70 (mm)	0.561
			2	22.4	81.3	1.20 (mm)	0.783
BF	0.88	0.85	1	37.4	37.4	4.61 (mm)	0.230
			2	26.8	64.2	3.92 (mm)	0.442
			3	9.7	73.9	3.15 (mm)	0.640
			4	8.8	82.7	2.59 (mm)	0.757
			5	5.0	87.7	2.05 (mm)	0.847
WYLD	0.68	0.61	1	64.6	64.6	2.43 (mm)	0.611
			2	3.7	68.3	2.43 (mm)	0.614
			3	0.9	69.2	2.46 (mm)	0.602
			4	0.4	69.6	2.47 (mm)	0.598
			5	0.4	70.0	2.49 (mm)	0.592
SED	0.77	0.67	1	43.7	43.7	2.76 (t km <sup>-2</sup> )	0.382
			2	19.2	62.9	2.50 (t km <sup>-2</sup> )	0.493
			3	11.1	74.0	2.13 (t km <sup>-2</sup> )	0.630
			4	1.6	75.6	2.08 (t km <sup>-2</sup> )	0.650
			5	1.0	76.6	2.03 (t km <sup>-2</sup> )	0.667
TP	0.76	0.65	1	51.5	51.5	12.03 (kg km <sup>-2</sup> )	0.468
			2	10.7	62.2	11.14 (kg km <sup>-2</sup> )	0.544
			3	10.4	72.6	10.32 (kg km <sup>-2</sup> )	0.608
			4	3.0	75.6	9.80 (kg km <sup>-2</sup> )	0.647
			5	0.7	76.3	9.71 (kg km <sup>-2</sup> )	0.653
TN	0.73	0.68	1	63.0	63.0	43.04 (kg km-2)	0.597
			2	5.8	68.8	40.56 (kg km <sup>-2</sup> )	0.643
			3	3.1	72.1	39.20 (kg km <sup>-2</sup> )	0.666
			4	0.5	72.6	38.90 (kg km <sup>-2</sup> )	0.671
			5	0.7	73.3	$38.51 (kg km^2)$	0.678

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# 3.4.2 Effects of LUCC predictors on water quantity and quality

According to the PLSR results, each category of the landscape indices including percentage (PLAND), largest patch (LPI), 585 shape (AWMSI), contiguity (CONTIGAW), aggregation (AI), or interspersion (IJI), plays an essential role in influencing as least one water quantity or quality variable (Table 5Table 5). The effects on the changes in ET, SQ, BF, WYLD, SED, TP, and TN are measured using weights, regression coefficients (RCs), and VIP values in the PLSR models. VIPs for predictors included into the models are greater than 0.8. For the ET model, the highest VIPs are obtained in predictors aggregation index for arable land (AIa) and contiguity index for arable land (CONTIGAWa) (VIP = 1.25, RCs = -0.122), followed by 590 PLANDa (VIP = 1.037, RC = -0.101) and AI $\mu$  (VIP = 1.03, RC = -0.1). ET tends to decrease with larger aggregation (AIa) and contiguity (CONTIGAWa) indices, and arable land percent (PLANDa) (negative RCs), whereas it increases with more pasture (PLANDp) (positive RC). In the case of surface runoff, the first and second components of the model are dominated by PLANDu on the positive side, with minor positive effect from PLANDa on the second component (Table 5). The urban area percent (PLANDu) obtains largest VIP of 1.173, and are identified as most important influencing the change in surface runoff. Surface runoff increases with an increase in arable (PLANDa) and urban areas (PLANDu) (RCs=0.403, 1.161, 595 respectively). For base flow, in addition to arable land, pasture plays a key role in explaining its variation. Arable land (PLANDa), pasture (PLANDp) percent and area-weighted shape index of pasture (AWMSIp) obtain the largest VIPs of 1.259, 1.03, and 1.063, respectively. All show negative correlations with base flow. Ala and CONTIGAMa are important predictors for water yield with large VIPs of 1.226 and 1.218, respectively. Their higher values result in an increase of water 600 yield. For sediment, TP or TN models, the selected components are dominated by areal percentages of arable land and pasture, in addition to the landscape metrics of arable land. The models obtain the largest regression coefficients or VIPs for PLANDa, LPIa, or PLANDp. They have VIPs of 1.0113-1.173 for sediment, 1.089-1.305 for TP, 1.005-1.232 for TN, respectively. Inferred by the RCs, an increase in sediment, TP, or TN occurs with increasing arable land (RCs: 0.602-0.884), while a decrease may occur with higher percentage of arable land in largest patches (LPIa) (RCs: -0.74 - -0.225), or with

605 more pasture area (RCs: -0.693 - -0.122). rLPIa, AIa and CONTIGAWa are the mo

<sup>T</sup>LPI*a*, AI*a* and CONTIGAW*a* are the most important landscape structure indicators affecting water quantity or quality (VIP  $\geq 1$  most of the time, Table 5). AI*a* and CONTIGAW*a* have positive impacts on WYLD while negative impacts on ET. By

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Formatted: Font color: Text 1 Formatted: Font: 9 pt Formatted: Tab stops: 9.79 cm, Left Formatted: Font: 9 pt definition, AI*a* and CONTIGAW*a* would increase, respectively, when arable landscape patches are more clumped and contiguous (Shi et al., 2013; Uuemaa et al., 2009). Clumped and connected aAgriculture in more clumped and connected land patches with fewer edges has been proven to show a higher capability of reducing thehave reduced more runoff

- 610 <u>land</u> patches with fewer edges <u>has been proven to show a higher capability of reducing thehave reduced more runoff</u> infiltration of runoff, compared to <u>that in</u> small scattered patches (Boongaling et al., 2018), <u>thus resultingwhich may result</u> in the increase of <u>the</u> water yield <del>amount in the catchment</del>. Our results also corroborate with Ayivi and Jha (2018) who reported that increased water yield and base flow occur with increasing cohesive and aggregated agriculture <u>in a moderate-altitude</u> <u>catchment (i.e., -Reedy Fork-Buffalo Creek catchment, USA)</u>. –Negative impacts on ET may be explained by the interactive
- 615 changes between arable and pasture, i.e., arable land has been increased at the cost of losing pasture, and vice versa. Likewise, Shawul et al. (2019) observed that reduction in pasture would result in a decrease of ET in an agriculturedominated and moderate-altitude catchment (the Upper Awash Catchment, Ethiopia). The negative effect of AWMSIp on base flow implies that the coarse grass landscape has a higher capacity of absorbing and intercepting rainfall thereby resulting in lower base flow. Though landscape metrics are more often used to explain water quantity than quality variables (Table 5), the negative influences of LPIa on sediment and nutrients, and positive influences of AWMSIa on sediment and
- TP cannot be overlooked. The sSimilar finding waswere observed also is is in agreement with previous\_findings in the hilly catchments, where that scattered and complicated agriculture patches are susceptible to soil erosion and thus water quality deterioration\_(Nafi'Shehab et al., 2021; Yan et al., 2013).
- The change in the percentage of arable land is most responsible for water quantity and quality dynamics, with VIP values greater than 1 for all response variables but WYLD. This may be explained by the fact that the decrease in arable land is the strongest. The negative correlations between PLAND*a* and evapotranspiration (ET) and base flow (BF) imply that conversion of arable land to e.g., pasture or forest would result in increased ET and BF, due to higher capability of plant evapotranspiration and slower water transmission, which is in agreement with previous findings that perennial vegetation is more likely to increase ET (Li et al., 2017; Peel et al., 2010) and the decrease in agriculture leads to increased annual base
- 630 flow (Basuki et al., 2019). <u>Changes of the percentage of arable land positively influence SQ, WYLD, SED TP, and TN loads</u>. Less <u>runoff</u>-interception by crops and additional <u>surface</u> runoff <u>routes</u>-resulting from implementation of tillage practices (e.g., tractor road) can result in increased surface runoff (SQ). The lower ET amount of crops compared to pasture and forest is in part responsible for the increase in WYLD. <u>Soil erosion might be accelerated due to uncovered and fragile</u> soil by tillage practices implemented in cultivated areas as well as the increased surface runoff. N and P pollution is prone to occur in arable areas, which have a high risk of generating nutrient pollutants from excessive fertilizer or manure and eroded

soil particles. The positive relationships between arable land percent and SQ, WYLD, SED TP, and TN loads are found in

- other studies <u>around the world</u> as well (Mirghaed et al., 2018; Sood et al., 2021; Wagner et al., 2013; Wang et al., 2019; Zhang et al., 2020a). Pasture shows a positive influence on ET and negative influences on sediment, TP, and TN. This also illustrates that more grassland (or rangeland) would increase plant evapotranspiration process, <u>Pasture can improve water</u> quality due to reduced soil erosion and nutrient transportation rate, as well as the high uptake and infiltration of nutrients by vegetation cover, <u>Relevant studies</u> (Ding et al., 2016; Hatano et al., 2005; Li et al., 2008; Zhang et al., 2020a) <u>have often</u> observed that semi-natural vegetation (e.g., forest, bushland or grassland) is beneficial for good water quality in river- or lake-dominated catchments, <del>bydue to</del> higher capability of filtering contaminants and reducing their inputs as well as decreasing surface runoff.
- By applying the quantitative results that the increases in arable or pasture areas most significantly intensify or reduce the risk of soil erosion and nutrient pollution, respectively, individual subbasins can be identified as nutrient pollution "source" or "sink". Based on these results, it is possible to develop a set of more targeted strategies to effectively control diffuse pollution at a spatial scale. At the same time, best management practices such as proper fertilization, abate of traditional tillage, crop rotation, vegetation buffer, are important to improve water quality in rural catchments (Haas et al., 2017; Pott
- and Fohrer, 2017a). Urban expansion is most important influencing surface runoff, the increase in urban area percent results in an increase of it (regression coefficient value > 1.16, <u>Table 5</u>. Similar results have been found, e.g., by Shi et al.

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(2007) who discovered that increased urbanized land led to increased surface runoff, by increasing flood peaks-flood runoff and decreasing surface runoff confluence time, in a typical urbanized region (Shenzhen) in China. It is therefore necessary to increase the frequency of measuring runoff, sediment and nutrient, particularly during the course of storm flood events in settlement area. Unlike previous findings (Wang et al., 2018; Yan et al., 2013), forest properties have not exerted significant influences, probably due to only minor temporal changes in some landscape metrics, e.g., area percent (PLAND), dominance (LPI), and shape (AWMSI) of forest (Figure 6).

Table 5. Regression coefficients (RCs), VIP and weight values of each PLSR model.

	Table	5. Reg	ression	a coeffi	cients (	RCs), V	/IP and	l weigh	t value	es of ea	ch PLS	R mod	el.								Fie
	ET			SQ				BF							WYLD						
Predictors	RC	VIP	W*[1]	RC	VIP	$W^{\circ}[1]$	W*[2]	RC	VIP	$W^{*}[1]$	W*[2]	W*[3]	W* [4]	W*[5]	RC	VIP	W*[1]	W*[2]			
PLANDa	-0.101	1.037	-0.017	0.403	0.790	-0.048	0.189	-1.654	1.259	-0.001	-0.128	-0.135	-0.208	-0.201	0.043	0.882	0.017	-0.042			
PLANDp	0.089	0.918	0.015					-1.474	1.030	-0.034	0.024	-0.117	-0.304	-0.256	0.011	0.866	-0.015	0.072			
PLANDf								-0.575	0.915	-0.035	-0.074	-0.072	-0.045	0.092							
PLANDu	0.080	0.818	0.013	1.161	1.173	0.090	0.173														
LPIa	-0.088	0.906	-0.015																		
AWMSIp								-0.143	1.063	-0.052	-0.058	0.059	0.093	-0.013							
AWMSIf	0.085	0.870	0.014												-0.039	0.837	-0.016	0.041			
AIa	-0.122	1.254	-0.020												0.187	1.226	0.024	0.025			
AIP	-0.094	0.961	-0.016												0.100	0.924	0.018	-0.009			
AIu	-0.100	1.030	-0.017												0.212	1.068	0.020	0.058			
CONTIGAWa	-0.122	1.251	-0.020												0.184	1.218	0.024	0.024			
CONTIGAWP	-0.087	0.891	-0.015												0.112	0.880	0.018	0.004			
CONTIGAWu	-0.094	0.959	-0.016					0.281	0.805	0.040	0.029	-0.078	0.064	0.011	0.198	1.007	0.019	0.054			
IJla								0.038	0.859	0.040	0.024	0.098	-0.142	-0.091							
Des listers	SED							TP							TN						
Fiediciois	RC	VIP	$W^*[1]$	W*[2]	W*[3]	W*[4]	W*[5]	RC	VIP	W*[1]	W*[2]	W*[3]	W*[4]	W*[5]	RC	VIP	$W^{*}[1]$	W*[2]	W*[3]	W*[4]	W*[5]
PLANDa	0.602	1.165	0.027	0.038	0.106	0.037	0.040	0.755	1.305	0.029	0.031	0.117	0.142	0.059	0.884	1.232	0.033	0.103	0.133	0.166	0.333
PLANDp	-0.693	1.173	-0.026	-0.022	-0.124	-0.096	-0.099	-0.499	1.089	-0.025	-0.007	-0.099	-0.074	0.002	-0.122	1.005	-0.030	-0.054	-0.049	0.031	0.324
PLANDu	0.013	0.908	-0.022	-0.033	0.020	0.097	0.116	-0.045	1.038	-0.025	-0.033	0.005	0.057	0.137	0.028	1.013	-0.024	-0.032	0.052	0.197	0.093
PLANDf								-0.009	0.821	-0.016	-0.053	0.061	0.047	0.004							
LPIa	-0.632	1.113	0.015	-0.095	-0.117	-0.037	-0.070	-0.740	1.205	0.017	-0.064	-0.208	-0.091	-0.057	-0.225	0.945	0.023	-0.054	-0.209	0.028	0.019
LPIp	0.397	0.819	-0.009	0.075	0.086	-0.043	0.020														
AWMSIa	0.472	0.902	0.007	0.103	-0.017	0.073	0.080	0.492	0.817	0.008	0.087	0.020	0.093	0.085							
AWMSIp	-0.445	1.087	-0.023	-0.077	-0.050	-0.022	0.107	-0.152	0.872	-0.019	-0.031	-0.057	0.127	-0.001							
CONTIGAWa	0.039	0.877	0.023	-0.001	-0.042	-0.024	0.075	0.079	0.864	0.021	-0.027	-0.013	0.015	0.069	0.114	0.840	0.022	-0.072	0.037	0.019	0.077
AIa	-0.053	0.876	0.022	-0.006	-0.055	-0.039	0.041	0.008	0.856	0.021	-0.030	-0.025	0.000	0.052	-0.034	0.833	0.022	-0.081	0.015	-0.024	-0.038

Note: VIP values greater than 1 were marked in bold; the absolute weights greater than 0.1 were marked in Italic.

### 660 4 Conclusion

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In this study\_--the separate contributions of changes in land use on the dynamics of seven water quantity and quality variables, i.e., actual evapotranspiration (ET), surface runoff (SQ), base flow (BF), water yield (WYLD), sediment (SED), total phosphorus (TP), and total nitrogen (TN) loads have beenwere quantified by applying an integrated approach of hydrological modeling (SWAT) and partial least squares regression (PLSR). The influences of the changes in individual land use indicators dseape metrics on variations changes in water quantity and quality have beenwere measured and identified using a scenario analysis for three different land use maps of the past.

With an exceptional data set that covers land use changes and three water quality campaigns over a period of three decades, a hydrologic model was set up and showed reasonable performance on the daily time scale. The modelling analysis of the effects of past land use changes showed that Driven by land use changes, water quality and quantity variables are modelled to variedy in different ways on the subbasin scale. SED, TP, and TN decreased more strongly in the eastern and western Formatted: Font color: Red

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parts than in the middle lowlands, implying that a higher contribution of nutrients by groundwater can mediate the influences of land use change. The results of the scenario analysis indicate that the dynamics of all water quantity and quality variables are largely explained Based on a PLSRn analysis-using PLSR, about 75% (on average) of the modelled variations in water quality and quantity variables can be accurately explained by land use indicators, indicating land use changes contribute to

- 675 the majority of water resources dynamics. The change of arable land is inferred to be most responsible important for water quality and quantity dynamics, as arable land indicators mostly holdshowed a greater importance (measured by VIP>1) for more response variables compared to other indicators. Looking at the most significant impacts, expansion of arable land (PLANDa) may cause a BF to decrease and. Uurbanization expansion will resultresulted in increased SQ. WithMore aggregated and connected arable land patches become more aggregatedled to a decrease of WYLD
- 680 are expected to decrease and increase, respectively. Arable land expansion tend to exacerbated soil erosion and P and N pollution, whereas an increase of pasture ean helphelped to relieve nutrient pollution problems. These results pointunderline to the fact that water quality and quantity variables are affected by land use changes in different ways. Overall,T-to achieve reasonable water balance and good water quality, more attention should be attached to the dynamics in the extent and the spatial configuration of arable land require special attention.extent as well as its spatial configuration. The spatial assessment of changes in water quantity and quality variables in se results this study camprovides a basis -be referred to provide for an

<u>Solution in the results of the variations in other water quantity and quality variables are most influenced by arable land change.</u> (61–68% of the variations in ET and WYLD; 75–88% of the variations in other water quantity and quality variables) by land use changes (LUCC) between 1987 and 2019. Landscape metrics show a stronger effect on water quantity than on water quality. Moreover, water quantity

- 690 and quality variables are most influenced by arable land change. The percentage (PLANDa), contiguity (CONTIGAWa), and aggregation (AIa) of arable land are identified as primary landscape metrics controlling the variations in BF, ET and WYLD. Greater percentages of settlement area and arable land may significantly accelerate runoff processes. Land planners and decision makers probably need to control land use patterns in runoff-sensitive areas to minimize negative impacts. Sediment, TP, and TN loads are closely associated with pasture and arable land. The expansion of arable land (PLANDa)
- 695 may exacerbate soil erosion and P and N pollution. The arable land in large and aggregated (LPIa) or simpler shape (AWMSIa) patches can help to mitigate soil erosion and water quality deterioration. The results indicate that the smaller changes in forest did not exert significant influence on water quantity and quality.

The approach applied in this study identifies the important influences of land use changes on water quantity and quality, which are helpful for formulating an informed and targeted plan with regard to land and water resource management. This approach is applicable to other catchments to predict both the water quality and hydrological responses to land use changes with the help of time sequenced land use data.

# Data availability

The datasets used in this study may be available upon request to the corresponding author.

### Author contribution

705 Chaogui Lei, Paul D. Wagner, and Nicola Fohrer designed the experiments and Chaogui Lei carried them out. Chaogui Lei and Paul D. Wagner developed the model codes. Chaogui Lei performed the simulations with the supervision by co-authors. Chaogui Lei prepared the manuscript with many contributions from all co-authors.

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# **Competing interests**

The authors declare that they have no conflict of interest.

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