

Detecting hydrological connectivity using causal inference from time-series: synthetic and real karstic study cases

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Abstract. We investigate the potential of causal inference methods (CIMs) to reveal hydrological connections from time-series. Four CIMs are selected from two criteria, linear or nonlinear ~~and~~ and bivariate or multivariate. A priori, multivariate and nonlinear CIMs are best suited for revealing hydrological connections because they suit-fit nonlinear processes and deal with confounding factors such as rainfall, evapotranspiration, or seasonality. The four methods are applied to a synthetic case and a real karstic study case. The synthetic experiment indicates that, confirms our expectation: unlike the other methods, the multivariate nonlinear framework has a low false-positive rate and allows for ruling out a connection between two disconnected reservoirs forced with similar effective precipitation. However, the-for the real study case, the multivariate nonlinear method appears-unstable when it comes to real cases, making the overall meaning of the causal links uncertain was unstable because of the uneven distribution of missing values affecting the final sample size for the multivariate analyses, forcing us to cope with the results' robustness. Nevertheless, if we recommend a nonlinear multivariate framework to reveal actual hydrological connections, all CIMs bring valuable insights into the system's dynamics, making them a cost-effective and recommendable comparative tool for exploring data. Still, causal inference remains attached to subjective choices and-operational constraints while building the dataset or constraining the analysis, operational constraints, and hypotheses challenging to test. As a result, the robustness of the conclusions that the CIMs can draw deserves-to-be questioned always deserves caution, especially with real and-imperfect, imperfect, and limited data. Therefore, alongside research perspectives, we encourage a flexible, informed, and limit-aware use of CIMs, without omitting any other approach that aims at the causal understanding of a system.

1 Introduction

Causal inference methods (CIMs) aim at identifying causal interactions between variables from variables-data only (Spirtes et al., 2000; Pearl, 2009). When applied to time-series, these empirical methods are built upon the principle of priority of the cause, which goes back to Hume (Hume, 1748). They-: CIMs infer causation from the expected time-dependencies between causes and effects, i.e., the causes must occur before preceding the effects. They have evolved throughout the 20th century to go beyond the well-known correlation, or cross-correlation, between two time-series (see Runge et al., 2019a, for a broad review). Although widely used, the correlation or cross-correlation method is criticized for its inability to cannot identify non-

linear causal relationships. ~~Besides, the correlation cannot discriminate nor discriminate actual~~ causal links from ~~associations~~
25 ~~dependencies~~ resulting from confounding factors. Indeed, ~~dependencies between variables can be explained either by a direct~~
~~causal link or through~~ the common cause principle (Reichenbach, 1956; Runge et al., 2019a). The common cause principle
~~'s principle~~ (Reichenbach, 1956; Runge et al., 2019a) tells us that dependencies between ~~the two~~ variables could result from
a third ~~cause acting on the variables~~ confounding cause influencing both. Nowadays, ~~a plethora of CIMs has many CIMs~~
have been developed, differing in hypotheses or application fields. Some CIMs explicitly deal with, either or both, nonlin-
30 ear dependencies or confounding factors through multivariate ~~analysis~~ analyses. These new CIMs are of growing interest in
Earth, land, and hydrological sciences (Meyfroidt, 2016; Runge et al., 2019a; Goodwell et al., 2020). In hydrology, appli-
cations of CIMs remain rare and cover, for example, the potential causal feedbacks between soil moisture and precipitation
~~(Salvucci et al., 2002; Tuttle and Salvucci, 2017), cross-scales rainfall interactions~~ (Salvucci et al., 2002; Tuttle and Salvucci, 2017; Wang
~~, the differential effects of environmental drivers of evapotranspiration~~ (Ombadi et al., 2020), interactions across rainfall time
35 scales (Molini et al., 2010), the ecohydrological feedback processes (Ruddell and Kumar, 2009), or the study of hydrological
connectivity (Sendrowski and Passalacqua, 2017; Rinderer et al., 2018) ~~as in this article~~.

~~The study of hydrological connectivity~~ Our study compares four CIMs on hydrological study cases in the spirit of other recent
comparative studies (Rinderer et al., 2018; Ombadi et al., 2020). Our application differs in the choice of CIMs and the research
questions. We selected four CIMs, all operating on the time-domain following the principle of priority up to a maximum causal
40 delay d_{max} . We use two bivariate CIMs, the linear Cross-Correlation Function (CCF), the nonlinear Convergent Cross Mapping
(CCM) method (Sugihara et al., 2012; Ye et al., 2015), together with two multivariate methods, one linear testing for Partial
Correlation (ParCorr) and one nonlinear testing for Conditional Mutual Information (CMI). The multivariate CIMs are parts of
the same causal inference framework, PCMCI, a sequential procedure based on the PC algorithm (Spirtes and Glymour, 1991)
, followed by a test for Momentary Conditional Independence (MCI) (Runge et al., 2019b).

45 Like Rinderer et al. (2018), we discuss if CIMs are suitable for studying hydrological connectivity, which aims at identifying
the paths taken or that could be taken by water. ~~There is a priori no causal interaction, i.e., flow, possible between two points~~
~~in space that do not benefit from a hydrological connection. This link between the concepts of hydrological connectivity and~~
~~causality motivates the study of one given the other and vice versa. There are different types of connections and various~~
~~ways to refer to them~~ (Bracken et al., 2013). We ~~refer to the terminology of Rinderer et al. (2018), which is inspired by~~
50 and borrowed from align with Rinderer et al. (2018) 's terminology inspired by the field of neurological neurology and brain
connectivity (Friston, 2011). ~~There~~ Accordingly, there are three types of connectivity: (i) structural, (ii) functional, and (iii)
effective connectivity. The structural connectivity is derived from the medium and highlights the potential, static, and time-
invariant water flow paths from the geological environment's topography, spatial adjacency, or contiguity. The functional ~~one~~
connectivity is dynamic and ~~is~~ retrieved from statistical time-dependencies between local hydrological variables. ~~A statistical~~
55 ~~association may result from confounding factors, e.g., rainfall acting on two disconnected reservoirs or a shared seasonal~~
~~pattern. Therefore, dependencies do not necessarily imply factual causation, such as process-based water flows. Then, the~~
~~functional~~ Functional connectivity is a matter of cross-predictability and ~~still reflects potential rather than actual flow paths for~~
~~water. CIMs with a multivariate framework address confounding factors. They offer the promises of discriminating functional~~

connectivity from the effective one, which reveals actual flow paths and processes within the system. From the structural to the effective connectivity through the functional one, the search for hydrological connections can be seen as a progressive constraint from the potential paths to the actual paths taken by water. reflects dynamic links between the variables. These dynamic links are potential connections subject to confounding factors, i.e., they may or may not be related to a flow process between variables. Effective connectivity precisely refers to actual connections linked through hydrological processes and flows.

Because of their hidden and heterogeneous structures, karst systems are regularly studied through time-series analysis or other empirical approaches to derive a causal or functional representation of the system (Bakalowicz, 2005). The application of CIMs to karst is therefore very relevant. Rinderer et al. (2018) reviewed a broad ensemble of CIMs from the literature and tested them to assess groundwater connectivity. Like Ombadi et al. (2020), they report spurious effective connections (or False Positives), highlighting the imperfect, yet, variable detection performances of CIMs. To deal with False Positives, Rinderer et al. (2018) proposed to constrain or limit the results by an assessment of structural connectivity. However, karst systems present some challenges for causal inference (Bakalowicz, 2005). In particular, the heterogeneity of the fractured geological environment is difficult or impossible to observe, characterize, or hypothesize. This hiddenness jeopardizes the derivation of a reliable map of structural connectivity to guide causal inference, as recommended in Rinderer et al. (2018).

To date, the most commonly used method remains the linear cross-correlation function (CCF), including for karsts (e.g., Angelini, 1997; 1998). Some study cases are now involving linear multivariate methods (e.g., Salvucci et al., 2002; Tuttle and Salvucci, 2017), including for karst (Kadić et al., 2018). Hydrological systems were first theoretically characterized by linear methods (Dooge, 1973). Indeed, we expect some linearity in mass transfers and positive time-dependent correlations between causes and effects, e.g., precipitation and discharge. This delayed linear transfer allows hydrological signals to be related by the unit hydrograph theory as a simple convolution window (Dooge, 1973). Besides, linear methods also benefit from the computational efficiency of linear algebra, which makes linear CIMs less time-consuming. Despite their monotonic behavior, hydrological systems, especially karsts, are nevertheless sensitive to initial conditions, i.e., nonlinear. Nonlinearity is imputed to inherently our concern is the hydrological connectivity in the vadose zone of karst systems. Assessing structural connectivity in karst systems is a challenging task because of their hidden and heterogeneous structure (Bakalowicz, 2005). Thus, without structural connectivity assessment, we investigate hydrological connectivity from CIMs alone. Besides, karst systems are known for their nonlinear behavior, which could be imputed to nonlinear hydrological processes such as power laws, e.g., taking the form of power laws, or threshold effects triggering flows (Bakalowicz, 2005; Blöschl and Zehe, 2005). For this reason, nonlinear CIMs may be more suited, although not specifically designed to deal with threshold effects.

To investigate how the CIMs can help to understand the dynamics of the karst system and how the method hypotheses impact this understanding, we compare the connectivity inferred from the four CIMs (Table 1), all operating on the time domain: the linear and bivariate CCF, the bivariate and nonlinear Convergent Cross Mapping (CCM) method (Sugihara et al., 2012); and two multivariate methods, one linear (ParCorr) and one nonlinear (CMI), both part of the same causal inference framework called PCMCI implemented in the Tigramite Python package (Runge et al., 2019b). First, the four methods are used in a synthetic case study. As a result, by design, multivariate nonlinear CIMs (e.g., PCMCI-CMI) would seem best suited to retrieve

effective hydrological connections: they account for nonlinear dependencies and deal with confounding effects, e.g., seasonality or the forcing of precipitation and evapotranspiration, through a multivariate framework. Following a theoretical introduction of the different CIMs, where two hydrological discharge variables with similar meteorological forcing are either connected or disconnected. Then, the same CIMs are applied to a real-time-series dataset from the Rochefort cave in Southern Belgium. The real dataset include rainfall, we test this assertion and get hands-on the CIMs using a toy model to conduct a virtual experiment reproducing a simple case of two parallel hydrological reservoirs forced by the same effective precipitation. In a real study case, we apply the CIMs on data acquired at the Lorette cave (Rochefort, Southern Belgium). This dataset includes rainfall and potential evapotranspiration data, electrical resistivity patterns in time-series patterns from the subsurface obtained from a geophysical monitoring experiment using time-lapse Electrical Resistivity Tomography (ERT) (Watlet et al., 2018) (Watlet et al., 2018; Delforge et al., 2020b), and drip discharge time-series with distinct dynamical patterns monitoring percolation at three spots within the cave. Time-series also have different numbers of missing values, unevenly distributed over time, allowing a discussion of the impact of missing values on the analysis. To appreciate the results, In particular, previous dye tracing tests have revealed fast connected preferential flow between the surface and a particular spot in the cave (Poulain et al., 2018). This prior knowledge can be seen as a reality check on the blind CIMs (Poulain et al., 2018). We expect CIMs to reveal this specific connection. Time-series also have different numbers of missing values, unevenly distributed over time, allowing a discussion of the impact of temporal gaps or short sample size on the analysis.

2 Materials and Methods

2.1 Causal Inference Methods (CIMs)

Four CIMs are selected based on two binary criteria: linear or nonlinear and bivariate or multivariate (Table 1). They operate in the time-domain and investigate time-dependencies up to a maximum time delay d_{max} . The methods are presented here in a nutshell; a more detailed description of the methods and their practical implementation is available in the supplementary materials (SM1). The Cross-Correlation Function (CCF) and the Convergent Cross-Mapping (CCM) (Sugihara et al., 2012; Ye et al., 2015) methods respectively check for linear and nonlinear bivariate dependencies.

The two other methods are part of a multivariate causal inference framework implemented in the Tigramite Python package (v.4.1) and based on the PCMCI conditional independence algorithm (see Runge et al., 2019b, and SM1.3). Being multivariate, those methods can cope with confounding variables. The general principle of causal inference using conditional independence is the following: identifying a causal lag between two variables is checking whether or not time-dependencies pertain while removing the effect of the other variables and potential causal lags. Therefore, the analysis framework is stochastic because a background noise must persist to assess the variables' independence after removing deterministic effects. Both methods test the independence between the series either with a linear test, the partial correlation method (ParCorr, section 2.1.3), or a nonlinear one, the Conditional Mutual Information (CMI, section 2.1.4).

In practice, however, one cannot remove the effect of all potentially causal delays without falling into a curse of dimensionality affecting the causal detection power of the approach (Runge et al., 2019b). For this reason, ParCorr (and CMI) are coupled with

the PC and MCI algorithms (or PCMCI by concatenation). The first step is the PC algorithm, named after its authors Peter Spirtes and Clark Glymour (1991). Using the independence test, PC iteratively selects a subset of potential causal variables and delays for each variable to avoid the curse of dimensionality (Runge et al., 2019b, SM1.3). The second step, Momentary Conditional Independence (MCI), applies the independence test based on the subsets identified during the PC step. The PC procedure has a tuning hyperparameter named α_{PC} , which controls the number of potential causes. α_{PC} varies from 0 to 1, where 1 is the less restrictive case which implies not pre-selection.

While the CMI method is the most promising in that it does not assume linearity and accounts for confounding effects, as for the other CIMs, the reliability of the reported causal relationships nevertheless depends on underlying hypotheses (discussed in Runge, 2018a). Perhaps, the most important but the most challenging to verify and conceptualize in practice is the hypothesis of causal sufficiency. Causal sufficiency implies that the analysis should include all potential common causes. This is indeed difficult to verify because (i) potential causes, as variables are generally dictated by data availability, (ii) one is not supposed to know potential causes before the analysis, or (iii) the concept of variable is mathematically abstract and building a finite and parsimonious time-series dataset require a prior or tacit exercise of conceptualization of the continuous system. In this way, one can say that the dataset is a hypothesis in itself, which, if improperly framed, can induce spurious causal links even with the best CIM available.

Selected causal inference methods (CIMs) **Linear Nonlinear Bivariate** Cross-Correlation Function (CCF) **Convergent Cross-Mapping (CCM)** **Multivariate** Partial Correlation (ParCorr) **Conditional Mutual Information (CMI)**

2.1.1 Cross-Correlation Function (CCF)

The CCF method is the most common to analyze time-dependencies and address causality (e.g., Angelini, 1997; Larocque et al., 1998; Labat et al., 2015). A variable X_t is said to be a cause of variable Y_t if the Pearson's correlation coefficient ρ between Y_t and Y_t and X_{t-d} is significant on their overlapping domain for at least one realistic value of d up to d_{max} . The significance is assessed with a Student's t-test (see supplementary materials SM1 for more details on the CIMs' implementation).

Usually, the CCF method is not explicitly presented as a CIM. Nevertheless, we consider it as such because the method is simple, intelligible as linearly interpretable, and, in practice, widely used for the implicit purpose of making causal inference, despite its limits as a bivariate linear method, in most scientific domains like hydrology (e.g., Angelini, 1997; Larocque et al., 1998; Labat et al., 2015). In general, linear methods have been popular in hydrology and attractive for their computational efficiency (Dooge, 1973). Besides, as a result of CCF popularity, the outcomes of any other CIMs are best appreciated when benchmarked against the well-known CCF method.

2.1.2 Convergent Cross-Mapping (CCM)

The CCM method CCM is primarily designed to reveal weak nonlinear interactions between time-series (Sugihara et al., 2012; Ye et al., 2015). CCM tests if dynamic trajectories behave consistently is a nearest-neighbor forecasting approach. It tests whether dynamic trajectories between two variables behave consistently, i.e., with some cross-predictive skills, while the

160 system revisits the same states ~~, i.e., dynamic recurrence~~(dynamic recurrence). The system states are usually unknown; they are approximated by trajectory segments found in a trajectory matrix M_Y given by the Takens' embedding theorem (Takens, 1981): ~~$M_Y = \{Y_t, Y_{t-1}, \dots, Y_{t-(m-1)}\}$,~~

$$\underline{M_Y = \{Y_t, Y_{t-1}, \dots, Y_{t-(m-1)}\}} \quad (1)$$

where m is the embedding dimension. In this case, with unit lags between time-series, m corresponds to the length of the segments and can be optimized using self-forecasting performance while predicting points in Y_t from their nearest neighbors in M_Y (Sugihara and May, 1990; Delforge et al., 2020a). This length m is set to two days in this study ~~,~~ due to the overall good performance of this value during our preliminary testing, i.e., $M_Y = \{Y_t, Y_{t-1}\}$.

For a causal analysis and to check if X_t is a cause of Y_t , CCM makes forecasts of the points in X_t from points other values in X_t , whose time indices are identified from Y_t . ~~For a single forecast at a~~ as similar states in M_Y for a given forecast time of reference t^* , ~~CCM selects the time indices of the $m+1$ nearest neighbors of the state Y_{t^*} . The points in X_t mapping to the nearest neighbors' time indices are averaged to make a forecast of X_{t^*} . The nearest neighbor time indices are selected such that they are at least.~~ The algorithm is detailed and illustrated in the supplementary materials (SM1.2). Our CCM implementation is found in Delforge et al. (2020a). CCM is an ensemble method, i.e., producing many (N_{SAM}) forecasts of the time-series at a given lag from N_{SAM} bootstrapped samples of size L from the trajectory matrix M_Y to identify similar states. In our case, we measure the predictive skills as suggested in Sugihara et al. (2012) with the average Pearson's correlation $\bar{\rho}$ between time-series and their forecasts. Like CCF, the significance of the mean Pearson's correlation is appreciated with a Student's t-test.

There are few recent applications of CCM in hydrology (Wang et al., 2018; Medina et al., 2019; Ombadi et al., 2020). We chose the approach of Ye et al. (2015) over the original one of Sugihara et al. (2012). The latter infers causality without predictive delays ($d_{max} = 0$) and investigates convergence instead of time-dependencies. The convergence criterion for causality suggests that predictive skills $\bar{\rho}$ must increase by considering longer sample sizes L in M_Y to identify similar states until a plateau is reached. Yet, Sugihara et al. (2012) recognized that convergence could be met when variables are subject to strong forcing or seasonality, a case referred to as synchrony, which could yield spurious bidirectional causal relationships, in particular regarding hydrological variables (e.g., Sugihara et al., 2017). Although the parallel is not explicit in the literature, we considered that synchrony relates to confounding and adopted the approach of Ye et al. (2015), which attempt to overcome synchrony by investigating the asymmetric patterns of time-dependencies and solve causality from the principle of priority up to d_{max} , i.e., as CCF but considering nonlinear dependencies (see also Wang et al., 2018). We do not check for convergence and apply CCM with $L = 100$ based on our preliminary testing (similarly to Ombadi et al., 2020), assuming that convergence is occurring if $\bar{\rho}$ is significantly high for this given L . Secondly, we use a Theiler window (Theiler, 1986), ~~tw~~ days away from the reference state t^* , where tw is the Theiler window (Theiler, 1986). This window ensures that nearest neighbors are not direct temporal neighbors but recurrent ones remote in time, following the CCM philosophy, which is not considered in most applications of CCM (e.g., Wang et al., 2018; Ombadi et al., 2020), but well in applications related to the chaos theory from which CCM is derived (see Huffaker et al., 2018). The Theiler window is a time window that forces similar states to be sampled outside its

range, for a time of reference. It prevents undesired predictive skills from being imputed to auto-correlation rather than based on the consistency of dynamic recurrent trajectories, i.e., what CCM is supposed to test. Given the auto-correlated nature of hydrological series, we consider its use recommendable, especially for short time-series (Theiler, 1986). Although, the effect of tw is set to 10 days in this study such that nearest-neighbor states are on predictive skills is expected to be marginal if the time-series are long enough (Delforge et al., 2020a). In our case, we systematically consider $tw = 10$ days such that similar states are most likely to be separated from their reference point by at least by one rainfall event.

In practice, the time to prediction tp varies between $[-d_{max}, 0]$ to evaluate the time-dependencies between X_{t+tp} . Finally, since CCM is rooted in the chaos theory, it could be argued that CCM applies under the strict assumptions of its parent theory built on deterministic mathematical models, i.e., on low-dimensional systems with infinite length, noiseless, and non-cyclic and Y_t and infer causality from the principle of priority (Ye et al., 2015). In this case, the time-indices of the nearest neighbors of Y_t are simply shifted by tp to identify the states in X_t that are averaged into a single prediction of state X_{t+tp} . For each tp , we forecast the full vector X_{t+tp} of prediction 100 times based on bootstrapped samples of M_T (See SM1.2). CCM forecast skills and time-dependencies are appreciated with the mean Pearson's correlation between the 100 estimates and the original non-intermittent series or relations (Kantz and Schreiber, 2003; Sugihara et al., 2012; Sivakumar, 2017; Huffaker et al., 2018; Runge et al., 2020). While we acknowledge that these hypotheses are not met with real hydrological time-series, our implementation of CCM is the one developed by Delforge et al. (2020a). (Koutsoyiannis, 2006), CCM and its underlying framework were introduced as capable of operating on short and noisy series as a nearest-neighbor regressor combined with its bootstrapping strategy (Sugihara and May, 1990; Sugihara et al., 1994, 2012), and the interest of any CIM is ultimately to be applied to real data while raising awareness on potential problems related to the real context. CCM, as well as other CIMs (see Runge et al., 2019a; Ombadi et al., 2020) or any statistical methods, meet issues with sample length, noise, cycles, dynamic intermittency, as well as the stationarity or significance testing (Huffaker et al., 2018; Medina et al., 2019), which should be considered in the result interpretation.

2.1.3 PCMCI: the Partial Correlation Test (ParCorr) and Conditional Mutual Information (CMI) tests

The ParCorr conditional independence test is like Granger causality and relies on linear vector autoregressive models (Granger, 1969). The principle is the following: PCMCI is based on a stochastic framework testing conditional independence (Spirtes et al., 2000; Pearl, 2000), adapted to account for highly time-dependent time-series, and implemented in the Tigramite Python package (Runge et al., 2019b). It follows a two-step procedure: PC and Momentary Conditional Independence (MCI). The PC acronym refers to its authors Peter Spirtes and Clark Glymour (1991). PCMCI's grounding lies in a multivariate definition of the Wiener-Granger causality (Wiener, 1956; Granger, 1969), here referred to as MCI (Runge et al., 2019b, a): a lagged variable X_{t-d} causes is said to be a cause of Y_t if a significant correlation exists between X_{t-d} and has a significant dependence or predictive power over Y_t after removing the linear, while removing the effect of all the potentially causal other potential variables influencing X_{t-d} or Y_t , except X_{t-d} . These potential variables are called Parents and symbolized $\mathcal{P}(X_{t-d})$ and $\mathcal{P}(Y_t) \setminus \{X_{t-d}\}$. Conditioning the analysis on the Parents allows dealing with confounding variables.

Parents are subset variables found in the dataset that includes delayed variables up to $2d_{max}$, as $\mathcal{P}(X_{t-d})$ has to be estimated beyond the maximum causal lag d_{max} . Parents are estimated during the conditions selection step performed with PC. The

resulting estimates are noted $\hat{\mathcal{P}}(X_{t-d})$ and $\hat{\mathcal{P}}(Y_t) \setminus \{X_{t-d}\}$. PC removes irrelevant conditions from the full set of variables and delays, including their own past. There are recent applications of Granger causality or partial correlations in hydrology (e.g., Salvucci et al., 2002; Tuttle and Salvucci, 2017) or in the case of karst (Kadić et al., 2018). In the case of ParCorr, α_{PC} is automatically optimized by iterative independence testing until there is no more condition to drop out (see Runge et al., 2019b). The resulting size of Parents' set are controlled by α_{PC} , a liberal parameter varying between 0 and 1, with the latter being the less restrictive case that includes all possible variables. Then, MCI is defined as:

$$MCI : X_{t-d} \perp\!\!\!\perp Y_t | \hat{\mathcal{P}}(Y_t) \setminus \{X_{t-d}\}, \hat{\mathcal{P}}(X_{t-d}) \quad (2)$$

To assert that X_{t-d} causes Y_t , MCI must be rejected by a given independence test. PCMCI is paired with the linear Partial Correlation (ParCorr) and the nonlinear Conditional Mutual Information (CMI) tests. ParCorr is estimated by ordinary least squares by regressing the variables against their covariates considering a multivariate linear model. Then, the dependency between the residuals is tested using Pearson's correlation and a Student's t-test for the significance. For the PC step with ParCorr, Tigramite optimizes α_{PC} between 0.05 and 0.5 for ParCorr-based on Akaike's Information criterion of model performance (Akaike, 1974). In the information theory, CMI, or $I_{X,Y|Z}$ for possibly multivariate and continuous random variables X, Y, Z is defined as the mutual information I between X and Y conditioned to Z (Runge, 2018b):

$$I_{X,Y|Z} = \iiint dx dy dz p(x, y, z) \log \frac{p(x, y|z)}{p(x|z)p(y|z)} \quad (3)$$

2.1.4 Conditional Mutual Information Test (CMI)

CMI is a nonlinear alternative to the ParCorr independence test running with the PCMCI algorithm (SM1.3). CMI is a multivariate extension of the concept of entropy transfer (Schreiber, 2000; Sendrowski and Passalacqua, 2017), i.e., another bivariate nonlinear CIM such as CCM. CMI is evaluated with a nearest-neighbor estimator (Frenzel and Pompe, 2007; Vejmelka and Paluš, 2008) coupled with a shuffling significance test (Runge, 2018b). Several methods exist to estimate mutual information within the Tigramite package. However, the nearest-neighbor estimator is recommended for time-series below. If $I_{X,Y|Z} = 0$, X and Y are conditionally independent to Z , and, therefore, not directly causally related, assuming that the probability densities are correctly estimated, among other assumptions (Runge, 2018a). This is not a trivial task, especially in the case of high dimensionality or small sample sizes. Regarding PCMCI, dimensionality varies between variables according to the size of the sets of Parents (Eq.2), hence, depending on d_{max} , α_{PC} , the independence test and its parameters. Based on numerical experiments covering sample sizes from 50 to 2,000 and dimensions up to 10, Runge (2018b) recommends using nearest-neighbors estimators of CMI (Frenzel and Pompe, 2007; Vejmelka and Paluš, 2008) for small sample sizes (< 1000 samples, which is the case in this study. Unlike ParCorr, the). As no analytical significance test is available, like the Student's t-test used in the other cases, Runge (2018b) provides a shuffling significance test that is, in return, computationally demanding. Because of this computational requirement, PC with CMI has no implemented procedure for the selection of α_{PC} value is not optimized and has to be set manually. We considered two values, a less restrictive one of 0.2 and we limited ourselves to two values among suggested ones: the default $\alpha_{PC} = 0.2$ and a more restrictive one of 0.05 $\alpha_{PC} = 0.05$ given the strong interdependence of hydrological series.

CMI, or nonlinear independence tests in general, are very computationally expensive and quickly require high-performance computers or significant computational time depending on the size of PCMCI is based on a strict framework of assumptions: faithfulness, causal sufficiency, the dataset and the hyperparameters of the analysis (see Runge, 2018b). In comparison, the synthetic experiment at section 3 required two weeks of computation with CMI, while the result with ParCorr was almost instantaneous. absence of contemporaneous dependencies, the Causal Markov Condition, stationarity, and the assumptions behind the selected independence tests such as linearity or nonlinear constraints, or hyperparameters related to the estimators (see Runge, 2018a). In this study, we are mainly concerned with the first three. Faithfulness relates the absence of causality to conditional independence and imposes that time-series contains noise signals to test it. Causal sufficiency implies that the monitored variables include all common causes, following the principle (Reichenbach, 1956). Finally, contemporaneous links are related to the principle of priority, as some time lags are desired to infer the direction of causal relationships.

Given its flexibility, PCMCI is comparable to several other CIMs found in the literature. First, the PC algorithm itself is a CIM (Spirtes and Glymour, 1991), evaluated in hydrology as well (Ombadi et al., 2020). Yet, in comparison, the sequential PCMCI procedure is expected to reduce the number of False Positives (Runge et al., 2019b). PCMCI-ParCorr is also related to Granger Causality (GC) (Granger, 1969), which has several applications in hydrology (e.g., Salvucci et al., 2002; Tuttle and Salvucci, 2017), including karst (Kadić et al., 2018). Still, there are substantial differences. GC relies on multivariate vector autoregressive models, which do not account for contemporaneous dependencies. PCMCI reports significant contemporaneous dependencies, although without any causal direction since it cannot be inferred from the principle of priority. Also, GC relies on an F-test checking for significant reductions of the variance in the models' residuals. Besides, GC is not a sequential procedure like PCMCI. The GC conditioning is performed on all possible past variables, as if $\alpha_{PC} = 1$. As results, GC suffers from the curse of dimensionality, lowering its detection performance (Runge et al., 2019b, a). For this reason, GC is often applied in a bivariate way, considering only X_t and Y_t 's past. The Transfer Entropy (TE) method (Schreiber, 2000) is the nonlinear and information-theoretic equivalent of bivariate GC (Barnett et al., 2009), making PCMCI-CMI a multivariate and sequential extension of TE. In hydrology, there are some applications of TE (e.g., Ruddell and Kumar, 2009; Sendrowski and Passalacqua, 2017; Rind). Regarding PCMCI-CMI, Goodwell et al. (2020) reviews information-theoretic approaches and their potential applications in hydrology. Jiang and Kumar (2019) provides a framework built upon PCMCI-CMI to characterize short and long-term dependencies of observed stream chemistry data.

2.2 Study Site and Data

The karstic study site is located in the Calestienne, a band of outcropping Devonian limestone crossing southern Belgium. The time-series dataset results from the monitoring of the Lorette cave (Fig. 1.a), one of the largest caves sited. The karstic study site is the Lorette cave, next to the city of Rochefort in southern Belgium (Fig. 1.a) (Watlet et al., 2018; Poulain et al., 2018, and references therein). The dataset is completed by At 3 km of the study site, a PAMESEB agrometeorological station provides estimates of daily potential evapotranspiration data (E_{PET}, Fig. 1.c) estimated with the Penman-Monteith FAO-56 method (Allen et al., 1998) and data from a PAMESEB agrometeorological station located 3 km from the study site in Jemelle. Other time-series are obtained from sensors on site (Fig. 1.a). Rainfall data (RF, On the Lorette site, at the Rochefort

295 cave laboratory, daily rainfall data (RF, Fig. 1.a) are ~~daily-average-aggregated~~ from a Lufft tipping bucket rain gauge with a 1 min sample rate located at the surface (elevation ~ 225 m AOD). Inside the cave (elevation ~ 190 m AOD), two drip discharge monitoring devices (~~P1, P2~~P1, P2, Fig. 1.a) are installed within the main chamber accessible from a sinkhole, which constitutes the cave entrance. In particular, ~~P1~~P1 monitors an active dripping point due to a visible fracture on the chamber's ceiling (Triantafyllou et al., 2019, for a 3D model). Based on dye injection at the surface and in-cave tracing, a connection and
300 preferential flow path between the dye injection point (DT, Fig. 1.a) and ~~P1~~P1 was identified (Poulain et al., 2018). The breakthrough curve showed an initial arrival time of 3.75 hours, a sustained peak for 80 hours, and a tail lasting up to 120 days. However, sporadic peaks in concentration were observed after every rainfall event, reacting after 1.48 hours, peaking after 7.2 hours, and lasting up to 30 hours on average. ~~P2~~P2 monitors a dripping spot draining a porous limestone area. The last one, ~~P3~~P3, located in the North gallery, monitors the slow discharge from drops falling from one single stalactite below a massive
305 limestone layer. ~~P1, P2, and P3~~P1, P2, and P3 (Fig. 1.c) are daily means of the percolation rate.

An Electrical Resistivity Tomography (ERT) profile was installed to investigate the hydrology of the subsurface and potential connections above the cave (Watlet et al., 2018). ERT is a geophysical monitoring tool to study various types of hydrological processes (see Slater and Binley, 2021, and references therein). ~~At the study site, the~~The ERT profile is not flat ~~as it and~~ starts from the depression of a sinkhole where the entrance to the cave is located (ERT, Fig. 1.a). The ERT experiment allowed
310 collecting ERT ~~datasets data~~ daily between 2014 and 2017, which ~~still represents, to the best of the author's knowledge, the longest, represents the longest~~ high-resolution ERT monitoring experiment conducted in a karst environment to our knowledge (see Watlet et al., 2018, for details). This dataset consists of processed ERT data, defined over 1558 spatial cells. ~~Each of them is assigned to a resistivity time-series defined on and~~ 465 daily time-steps defining the temporal dimension of the dataset. As a necessary prior step to causal inference, the 1558 time-series were dimensionally reduced to six time-series clusters (~~R0 to~~
315 ~~R5~~R0 to R5, Fig. 1, b, and c). We used hierarchical agglomerative clustering with the Ward linkage method to minimize the squared distance between time-series within clusters. This algorithm is similar to k-means. As clustering was applied on the standardized ERT dataset, clusters represent groups of linearly correlated resistivity dynamics. The optimal number of clusters of 6 was selected using the optimal Silhouette Index as a clustering evaluation metric (Rousseeuw, 1987). The methodological aspects associated with clustering are extensively covered in another issue (Delforge et al., 2020b).

320 ~~R0~~R0 is associated with a dense limestone area in the model's center (Fig. 1.b, X=22 m, Z=10 m). ~~R1 shows~~R1 covers responsive resistivity dynamics at the plateau's surface (Fig. 1.b, X=32 m, Z=13 m), ~~as well as in and~~ a fractured area around coordinate X=15 m below the surface ~~patterns of R3. R2~~pattern R3. R2 is rather representative of the resistivity patterns of the limestone matrix. ~~R4~~R4 is associated with a clayey limestone layer located below the ~~patterns~~pattern of slope surface ~~R5~~
325 ~~R5~~R5 (Watlet et al., 2018; Delforge et al., 2020b). We expect causal links to appear primarily between P1 and the near-surface resistivity patterns ~~R1 or R3~~R1 or R3.

Time-series of Fig. 1.c are the 11 inputs for the four selected CIMs. Table ~~2-1~~2-1 shows their statistics. Bivariate CIMs (CCF and CCM) are applied between each pair of time-series on their overlapping time-domain with a maximum causal delay $d_{max} = 5$ days. ~~This maximum delay allows covering the~~We considered this delay for our final results as it covers the span of bivariate dependencies (Fig. SM2.1, SM2.2), and the full time-span of preferential flow peaks lasting up to 80 h (Poulain et al., 2018).

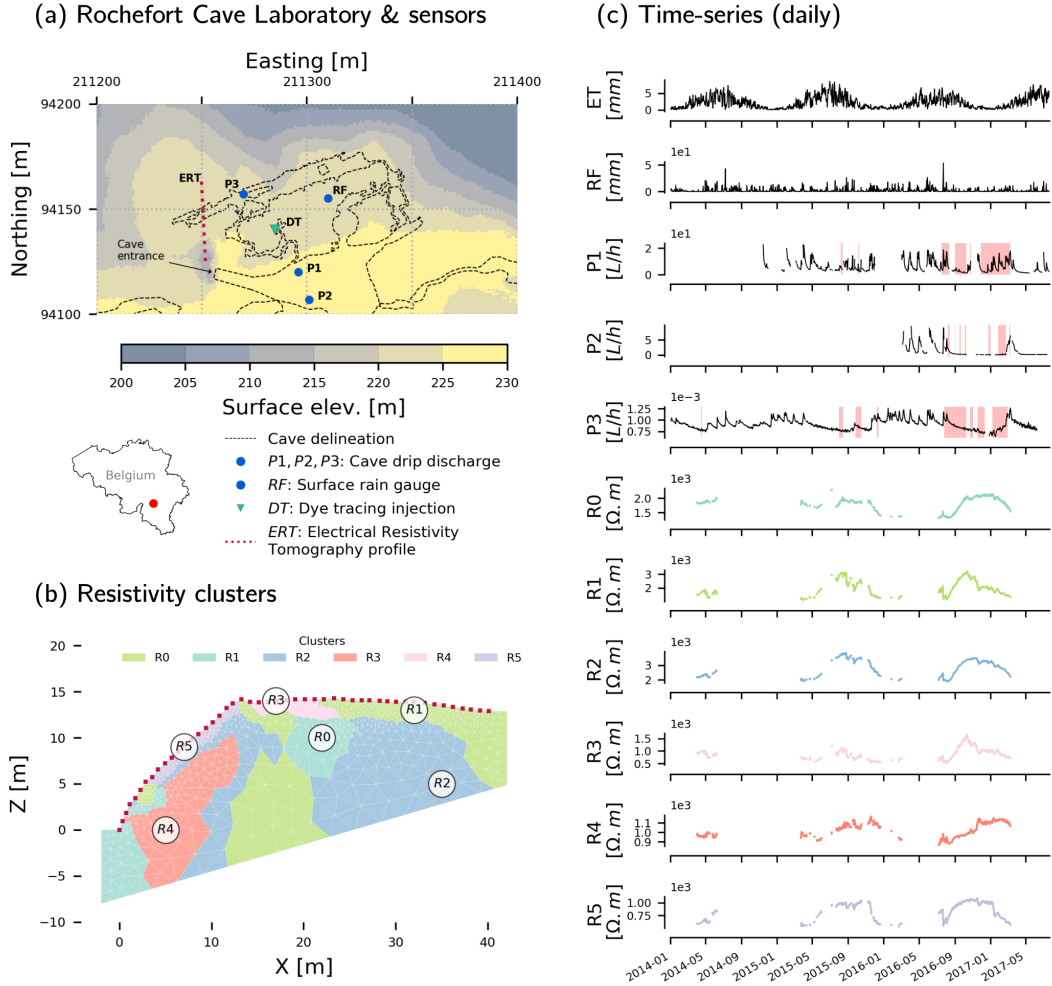


Figure 1. Study site and data: (a) Rochefort cave laboratory and sensors (EPSG: 31370), (b) resistivity clusters obtained from hierarchical agglomerative clustering of standardized resistivity data (Watlet et al., 2018; Delforge et al., 2020b), (c) daily time-series dataset. Resistivity time-series (e , $R0$ – $R0$ to $R5$ – $R5$) are the mean resistivity variations per cluster (b). Potential evapotranspiration data (ET – ET) are obtained from an agrometeorological station (PAMESEB) located 3 km from the site. The red background areas on (in Fig. 1.c) show the time-domain resulting from the conditioning on past delays with $d_{max} = 5$ days while considering respectively $P1$ – $P1$, $P2$ – $P2$, and $P3$ – $P3$ only in the causal dataset. A hydrological connection was identified by dye injection and tracing from the surface (DT) to $P1$ – $P1$ (Poulain et al., 2018). Source: Digital Elevation Model from Service Public de Wallonie, Cave delineation from Watlet et al. (2018).

Table 1. Summary statistics of the time-series variable.

Statistic	ET [mm]	RF [mm]	P1 [L/h]	P2 [L/h]	P3 [L/h]	R0 [Ω .m]	R1 [Ω .m]	R2 [Ω .m]	R3 [Ω .m]	R4 [Ω .m]	R5 [Ω .m]
Count	1297	1297	718	366	1223	465	465	465	465	465	465
Mean	2.2	2.0	6.13	1.30	9.06E-04	1.80E+03	1.95E+03	2.80E+03	8.78E+02	1.02E+03	8.28E+02
Std dev.	1.8	4.2	4.33	1.92	1.11E-04	2.38E+02	5.68E+02	5.64E+02	2.35E+02	7.75E+01	1.73E+02
Min	0.0	0.0	1.05	0.00	6.35E-04	1.30E+03	1.11E+03	1.88E+03	5.07E+02	8.67E+02	5.33E+02
10%	0.3	0.0	1.91	0.05	7.76E-04	1.37E+03	1.28E+03	2.03E+03	5.92E+02	9.23E+02	5.89E+02
25%	0.7	0.0	3.00	0.09	8.21E-04	1.69E+03	1.48E+03	2.27E+03	7.06E+02	9.59E+02	6.45E+02
50%	1.6	0.1	4.57	0.27	8.99E-04	1.85E+03	1.85E+03	2.83E+03	8.51E+02	1.03E+03	8.52E+02
75%	3.3	2.1	8.46	1.73	9.75E-04	1.96E+03	2.36E+03	3.29E+03	1.01E+03	1.10E+03	9.92E+02
90%	5.0	6.2	12.74	4.20	1.06E-03	2.08E+03	2.86E+03	3.48E+03	1.21E+03	1.13E+03	1.03E+03
Max	8.6	53.8	22.66	9.45	1.26E-03	2.30E+03	3.21E+03	3.83E+03	1.65E+03	1.17E+03	1.09E+03

330 For the same reason, we use the same d_{max} for the multivariate methods. In general, ~~it is preferable to use large d_{max} values ;~~
~~because any delay that can be considered as a potential cause must be included in the analysis are recommended~~ to satisfy the
hypothesis of causal sufficiency. ~~On the other hand while PCMCI sequential procedure deal with the potentially high dimension~~
~~resulting from d_{max} (section 2.1.3). Still, d_{max} cannot take a larger value because of the~~ ~~be higher than 5 days because of our~~
~~uneven time distribution of~~ missing data. Indeed, ~~the PCMCI algorithm PCMCI~~ dismisses all time slices of samples where
335 missing values occur in any variable and their lags up to $2d_{max}$, which limits the overall overlapping time-domain to 48 days.
~~This low value is mainly imputed to,~~ ~~mostly because of~~ the short time-domain of ~~P2-P2~~ (Fig. 1.c, Table 21). Concerned that
this number would impact the robustness of the analysis, we also applied PCMCI ~~with $d_{max} = 5$ days~~ while considering one
percolation data at a time. In this way, ~~P1, P2, P3-P1, P2, P3~~ are considered separate and independent drainage systems,
and the overlap domains become larger ~~while keeping $d_{max} = 5$ days~~: 184, 62, and 218 days respectively. These ~~conditioning~~
340 domains are shown in red for ~~the respective time-series considered (P1, P2, P3 in Fig. 1.c.)~~.

3 Synthetic Study Case

3.1 Conceptual Model

~~The four CIMs are first applied to test and explore the following theoretical assertion : nonlinear and multivariate CIMs~~
~~We aim at assessing CIMs performances for the detection of effective hydrological connectivity, in particular, the assertion~~
345 ~~that multivariate nonlinear methods are best suited to detect effective hydrological connectivity. For this purpose, for that~~
~~purpose. To this end, we built~~ a simple hydrological reservoir model~~is~~, inspired by the ~~problematic case of the common cause~~
~~common cause problem~~ (Fig. 2). ~~The common cause problem is easily conceived through two~~ ~~Two~~ separate and independent
reservoirs~~(, A and B)~~ ~~benefiting from the same meteorological forcing. Without an effective connection, they will nevertheless~~
~~show a strong temporal dependence. In this case, two reservoirs, A and B, are subject to the same forcing, and their discharge~~

350 Q_A, Q_B , are forced with the same effective precipitation P_{eff} (i.e., precipitation minus evapotranspiration). Then, without any effective connection, they will nevertheless show a temporal dependence. If A and B are disconnected, the ideal CIM would then reject the effective connection between ~~the reservoirs and their discharges~~ Q_A and Q_B if P_{eff} is included in the multivariate analysis. However, B responds systematically one day later than A to P_{eff} . Hence, with bivariate methods and the priority principle only, Q_A would seemingly cause Q_B . For comparison, we consider a case where Q_A ~~and is effectively~~ connected to another series Q'_B are effectively connected as if, as they were contributing to the same drainage network, with Q_A upstream of Q'_B . Notworthily, this experiment does not cover the case of nonlinearities arising from threshold effects and intermittent processes. Instead, we assume ~~a continuous causal relationship~~ sustained causal interactions and no intermittency over time, as the CIMs do.

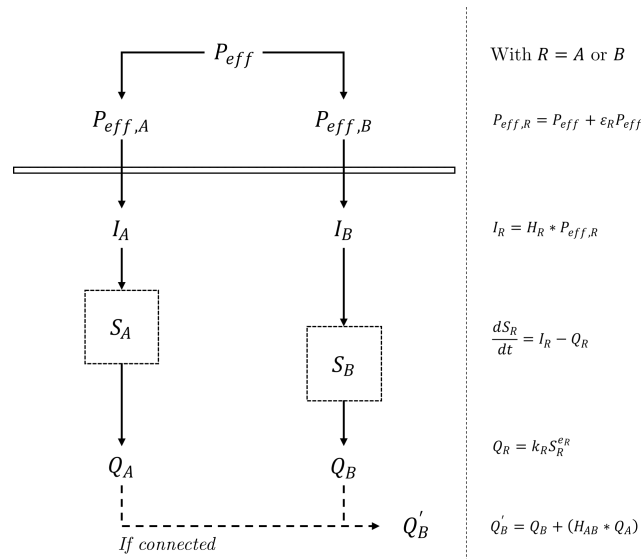


Figure 2. The conceptual and mathematical model for the synthetic study case. Two reservoirs A and B are forced by inflows I_A or I_B resulting from the unit hydrograph H_A and H_B forward convolution on noisy variant $P_{eff,A}$ and $P_{eff,B}$ of the same effective precipitation P_{eff} computed from real data (section 2.2). The storage S_A and S_B dynamics follows a typical continuity equation $dS/dt = I - Q$, where the discharge Q_A and Q_B follow a nonlinear power law $Q = kS^e$ with two parameters k and e . The reservoir B responds mainly 1 day after A , which introduced a systematic time-dependency between the reservoir suggesting causation. Q'_B is a flow downstream of Q_B draining Q_A and transferred considering the unit hydrograph H_{AB} . The causal analysis involves either the disconnected case Q_A and Q_B or the connected case Q_A and Q'_B . Multivariate methods includes P_{eff} within the analysis.

360 The model is forced by real effective precipitation data P_{eff} monitored at the study site (section 2.2). Both reservoirs take as input a net inflow term I_A and I_B resulting from the application of the unit hydrographs ~~H_A and H_B~~ H_R , with R either A or B , as linear transfer functions convolved forwardly on a noisy precipitation input ($H_R * P_{eff,R}$ with $*$ the convolution operator). Adding some noise is mandatory to check for conditional independence (faithfulness, section 2.1.3). A multiplicative noise term is preferred as hydrological variables are often characterized by multiplicative noise (e.g., Rodriguez-Iturbe et al., 1991).

365 ~~With R either A or B~~ Accordingly, $P_{eff,R} = P_{eff} + \varepsilon_R P_{eff}$, with ε_R being randomly generated from a normal distribution with zero mean and standard deviation equal to ε_{lvl} times the standard deviation of P_{eff} . The parameter ε_{lvl} is always identical for A and B , such that ε_A and ε_B have the same distribution. The continuity equation gives the reservoirs storage dynamics: $dS_R/dt = I_R - Q_R$. The outflow Q_R introduces some nonlinearities through a typical nonlinear storage-discharge relationship $Q_R = k_R S_R^{e_R}$, with k_R and e_R the discharge coefficient and the nonlinear exponent. Such power-law formulations are typical in hydrology (Dooge, 1973) and common while modeling karsts as well (Hartmann et al., 2014; Jourde et al., 2015).

Table 2. Model parameters for the synthetic cases

Model	H_A	k_A	e_A	H_B	k_B	e_B
1	[0.7, 0.2, 0.1]	0.1	1	[0.1, 0.8, 0.1]	0.1	1
2	[0.7, 0.2, 0.1]	0.1	1	[0.1, 0.8, 0.1]	0.01	1.5
3	[0.7, 0.2, 0.1]	0.01	1.5	[0.1, 0.8, 0.1]	0.1	1
4	[0.7, 0.2, 0.1]	0.01	1.5	[0.1, 0.8, 0.1]	0.01	1.5

370 For the synthetic cases, we derived four models based on ~~the parameters presented in Table 3.~~ Table 2. The unit hydrographs H_A and H_B are constant, with their maxima differing by one daily time-step. ~~The~~ This lag introduces the desired constant time-dependencies between the two reservoirs despite the absence of connection. The recession parameters allow generating distinct dynamic patterns with various degrees of nonlinearity thanks to $\varepsilon_R e_R$. In addition, we also considered 14 stochastic noise level $\varepsilon_{lvl} \in \{0.05, 0.1, \dots, 0.65, 0.70\}$. With ~~the such noise levels, the correlation between two generated Q_A with model configuration 1 (Table 2) should vary on average between 0.96 ($\varepsilon_{lvl} = 0.05$) and 0.24 ($\varepsilon_{lvl} = 0.7$).~~ With the four combinations of ~~Table 3-2~~ and the 14 noise levels, 56 datasets were generated from four years of effective precipitation data P_{eff} (2014-2018) and initial storages S_A and S_B equal to 30 mm. Only the last year of the three variables P_{eff} , Q_A , Q_B were considered for the causal inference experiment, ~~the first three being considered as a warming-up period~~. The data generation is repeated to produce 56 additional datasets with an effective connection. Q_A and Q_B are causally related by overwriting Q_B such that $Q'_B = Q_B + (H_{AB} * Q_A)$ where ~~$H_{AB} = [0.1, 0.8, 0.1]$~~ $H_{AB} = [0.1, 0.8, 0.1]$ is a linear transfer function convolved forwardly on Q_A . Finally, the whole synthesis process is repeated to generate first-order differenced datasets, ~~making (i.e., $Y_t - Y_{t-1}$), that is~~ a total of 224 datasets with 4 model combinations, 14 noise levels, connected or not, and differenced or not. The primary purpose of the ~~difference is to create a case where the shared seasonality that affects the bivariate dependencies between the data is eliminated~~ differenced data is to simply illustrate the effect and value of removing past dependencies (auto-correlation, seasonality) on the CIMs results.

385

3.2 Result

Figure 3 ~~shows the results of the experiment. It~~ depicts the average and interquartile range of $Q_A - Q_B - Q_B$ time-dependencies, or $Q_A - Q'_B$ if connected, obtained with the four CIMs ~~on the datasets synthesized from the numerous model combinations (Table 3) and noise levels with~~ (a to d) on the synthetic datasets for a maximum delay $d_{max} = 5$ days. The multivariate analysis

390 includes P_{eff} . We distinguish between cases where the reservoirs are connected or not and where the data are differenced or not. Regarding the bivariate methods (a and b), CCF and CCM both exhibit sustained time-dependencies ~~when the data is not differenced for not-differenced data~~ due to the auto-correlation ~~and seasonal patterns left in the series~~ and seasonality. ~~From~~. ~~For~~ differenced data, the results better screen the expected peak at lag one. ~~However, it is also the case while the reservoirs are disconnected because of~~, including disconnected reservoirs. This illustrates that bivariate CIMs cannot deal
 395 ~~with~~ the confounding effect related to the common forcing of ~~the two~~ reservoirs. This is not an effective connection but a functional and apparent one resulting from ~~the delayed responses of the two reservoirs~~ their delayed responses. The sustained time-dependency of CCM over the lag of 2 days is an artifact of the embedding dimension, $m = 2$ (Eq.1), defining the length of trajectory segments, which is two days in this case (see SM1.2).

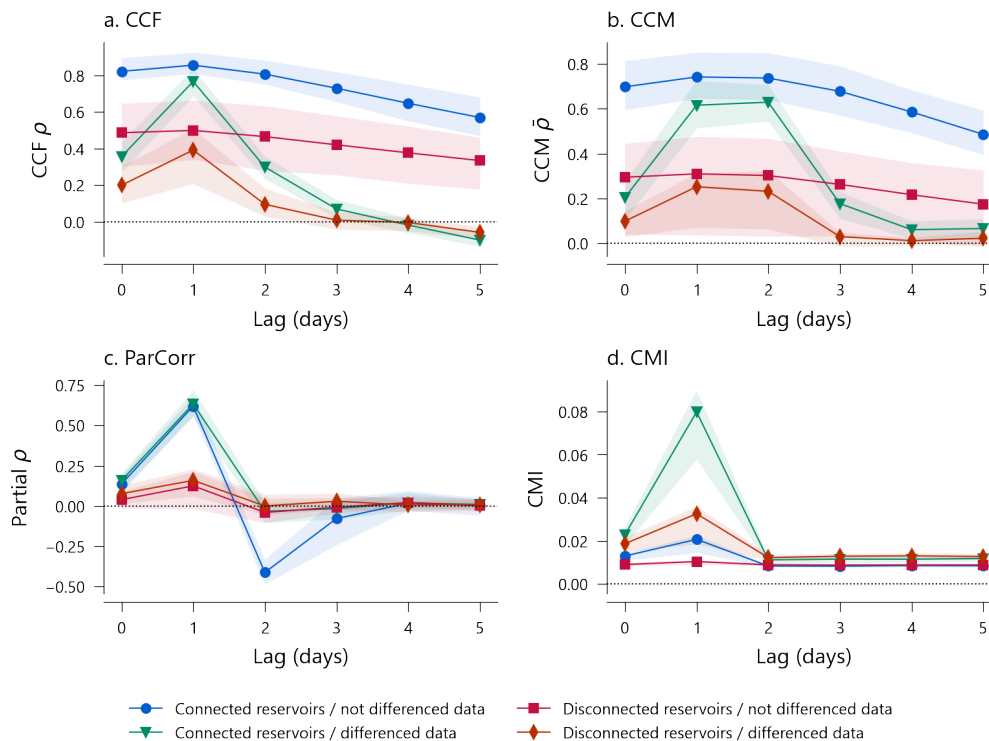


Figure 3. Patterns of statistical time-dependencies between Q_A and Q_B (Q'_B if connected) for the four CIMs (a. to d.). Lines are the average statistics for 56 synthetic datasets obtained from four different model structures (Table 3) and 14 distinct noise levels. The envelope represents the interquartile range of the statistics. In general, connected reservoirs show a $Q_A \rightarrow Q'_B$ causal dependencies at a lag of one day. However, except for CMI on the not differenced data, disconnected reservoirs show a non-causal, yet, significant statistical dependencies $Q_A \rightarrow Q_B$ at lag one day, since reservoir A mostly reacts one day before B to the effective precipitation.

400 Regarding the multivariate CIMs ~~patterns, the~~ (c and d), differencing time-series is theoretically useless as PCMCI already ~~conditions on the past variables~~ (Eq. 2) and can deal with auto-correlation and seasonality. The linear ParCorr method seems

to discriminate the connected case from the disconnected case. Still, it always shows a peak at lag 1, whatever the cases, which could be misinterpreted as an effective connection. Only the nonlinear CMI method applied ~~on the not-differenced to the not-differenced~~ data seems to reject the idea of connection when it is effectively absent. This finding supports our theoretical assertion: the multivariate nonlinear method is ~~the~~-best suited to address effective hydrological connectivity. Furthermore, the method appears to perform better if seasonality is left present in the time-series. Still, Fig. 3 shows the pattern of the statistics, not the result of a causality test and its p-value.

Since we know for each simulation whether or not there is ~~an-actual-a~~ causal link between A and B , Table 4 ~~assesses the performance of the statistical test from true positives-3 reports True Positives~~ (TP), ~~false positives-False Positives~~ (FP), ~~true negatives-True Negatives~~ (TN), and ~~false negatives-False Negatives~~ (FN) for the problematic lag of 1 day. We consider the multivariate PCMCI methods ParCorr and CMI, with the latter having two different α_{PC} ~~values for pre-selection of potential causes for preselection of Parents~~ (PC stage). Two levels of significance are considered at 99% and 95% based on the p-values obtained by the tests. Table 4-3 shows that, for a similar level of accuracy, CMI for ~~not-differenced-not-differenced~~ data has higher precision and a lower false-positive rate, meaning that positive tests are likely to detect actual causal relations. ~~This Differencing increased the PCMCI-CMI false-positive rate. The low false-positive rate for not-differenced data~~ is particularly contrasting with other methods and provides a valuable piece of information. However, the high precision comes at the cost of a low recall (~~as reported in Runge, 2020~~): CMI misses about half of the actual causal links. On the contrary, ParCorr misses none but has a bad precision, i.e., many ~~false positives-False Positives~~. This analysis thus provides an overview of the contrasts between methods. Of course, this virtual three-variable configuration is far from representative of the great variety of natural hydrological systems and their spatiotemporal organizations.

4 Real Study Case

4.1 Bivariate Methods

Figure 4 ~~shows-reports the causal graphs for significant~~ pairwise dependencies between first-order differenced time-series, for a better screening of time-dependencies using bivariate methods (Fig. 3). Detailed time-dependencies are also reported in the Supplementary Materials (SM2.1, SM2.2). ~~The links displayed are those for which the correlation value is significantly different from zero following a Student's t-test, allowing us to consider that $d_{max} = 5$ days is sufficient.~~ The CCF method (Fig. 4.a) ~~is reporting-reports~~ many potential linear causal associations. ~~If causality is hard to infer from such a diagram, the results make sense in general. A typical pattern is that the sign of time-dependencies tends to flip after a few delays due to RF's forcing and the fact that dry periods come after the rain. Considering low delays, ET is positively related to the resistivity patterns, mostly at the surface (R1(arrows), R3). ET is negatively correlated with P1 only, which is known to drain fast flow from the surface through the epikarst. All variables are dependent on RF, the main confounding factor, but R0, associated with a dense limestone area, depends to a lesser extent. R0 in red for positive correlations and R4 put apart, the quartet R1, R2, R3, and R5 exhibit strong positive and contemporaneous correlations together. P1, P2, and P3 also are instantaneously related. P1 and P2 have strong dependencies with all resistivity patterns, but inconsistent and positive correlations are reported between~~

Table 3. Causal test statistics for the synthetic cases at lag 1 day

Method	ParCorr		CMI		CMI	
	Not diff.	Diff.	Not diff.	Diff.	Not diff.	Diff.
Datasets						
Data Statistics						
Total count	112	112	112	112	112	112
Actual +	56	56	56	56	56	56
Actual -	56	56	56	56	56	56
Test results						
TP at 99%	56	56	28	56	15	56
(95%)	(56)	(56)	(34)	(56)	(27)	(56)
FP at 99%	29	29	3	28	1	28
(95%)	(35)	(32)	(4)	(42)	(2)	(35)
TN at 99%	27	27	53	28	55	28
(95%)	(21)	(24)	(52)	(14)	(54)	(21)
FN at 99%	0	0	28	0	39	0
(95%)	(0)	(0)	(22)	(0)	(29)	(0)
Test metrics						
Accuracy ¹ 99%	0.74	0.74	0.72	0.75	0.64	0.75
(95%)	(0.69)	(0.71)	(0.77)	(0.62)	(0.72)	(0.69)
Precision ² 99%	0.66	0.66	0.90	0.67	0.94	0.67
(95%)	(0.62)	(0.64)	(0.89)	(0.57)	(0.93)	(0.62)
Recall ³ 99%	1.00	1.00	0.5	1.00	0.3	1.00
(95%)	(1.00)	(1.00)	(0.61)	(1.00)	(0.48)	(1.00)
FP rate ⁴ 99%	0.52	0.52	0.05	0.5	0.02	0.5
(95%)	(0.63)	(0.57)	(0.07)	(0.75)	(0.04)	(0.63)

¹(TP+TN)/(TP+FP+FN+TN); ²TP/(TP+FP); ³TP/(TP+FN); ⁴FP/(FP+TN)

TP: True Positive; FP: False Positive; TN: True Negative; FN: False Negative

the anomalous resistivity series R4 representative of the clayey limestone. P3 seems rather dependent on R5 (slope) and R2 (mostly limestone matrix), making sense since P3 most likely drains the matrix's delayed flow.

Graph of pairwise cross-dependencies: (a) with the linear Cross-Correlation Function (CCF), (b) with the nonlinear Convergent Cross-Mapping (CCM) method. An undirected line represents contemporaneous dependencies. Delayed dependencies are shown using directed curved arrows. All corresponding delays d are displayed in the middle of its corresponding arrow. The color of arrows maps to the strength of dependencies. Solid and dash-dotted arrows represent respectively significant dependencies with p -value < 0.001 and < 0.01 .

Compared to CCF, the CCM results in blue for negative ones. Delayed dependencies are represented with curved arrows with the delay d (in days) printed in the middle. Contemporaneous dependencies ($d=0$) are represented by straight links, with no direction since it cannot be inferred from the principle of priority. For CCM (Fig. 4.b) are more intelligible since fewer significant links are reported. However, CCM as a nonlinear method gives no clue about the nature of the underlying dynamics or the dominant sign of the dependencies. P2 is now exclusively related to R5, and P3 has no dependencies on resistivity patterns. P1 is CCM-related to the surface resistivity patterns R1, R3, significant dependencies are fewer, unsigned as nonlinear, and R5, but also R2, which is representative of the limestone matrix resistivity. Compared to CCF, CCM supports the particular conclusion of connected and preferential flows occurring between the surface and P1, therefore, only represented in red. The results and their meaningfulness are appreciated in their dedicated discussion sections.

4.2 Multivariate methods

For multivariate methods, we chose to report causal graphs for the raw (not differenced not differenced) data since differencing did not impact ParCorr but is theoretically unnecessary and reduced the precision of CMI in the virtual experiment (Table 3). Hence, Fig. 5 shows linear conditional dependencies (ParCorr) obtained from the raw time-series for the full dataset (All data, Fig. 5.a) and considering the discharge series one by one (P1, P2, and P3, Fig. 5, b to d). The P1, P2, and P3 datasets allow the analysis to be performed over larger time domains (Fig. 1.c). Except for R4, the dominant relationships between resistivity and meteorological variables are maintained between the graphs, demonstrating stability in the ParCorr results despite differences in the considered time-domain. As for CCM (Fig. 4.b), P1 is associated with R1, R2, R3, and R5 (Fig. 5, a and b). The rainfall RF remains significantly related to P1, suggesting that resistivity patterns are not sufficient causes of P1. Regarding P2, similarly to CCM (Fig. 4.b), ParCorr on all data (Fig. 5.a) reveals a significant link between R5 and P2. However, the link is absent in Fig. 5.c, and two direct links from RF appear and bypass the resistivity patterns. Yet, R3 and R4 seem to influence P2 at lag 2, but the relationship is positive, which cannot be interpreted as a linear mass transfer. We also denote two upward links to R4 and ET. These links seem physically unrealistic and potentially problematic since the effect of P2 is to be removed from these variables, which may alter the whole causal graph. P3, i.e., the low rate stalactite drip discharge, remains unrelated (Fig. 5, a and d).

Regarding the nonlinear analysis (CMI) PCMCI-CMI, we found unstable results on all datasets: All data, P1, P2, and P3 (see SM2.3). The causal graphs varied substantially when we repeated the analysis with the same parameters due to the stochastic nature of the independence test (Runge, 2018b). Consequently, we developed a sensitivity analysis by varying the hyperparam-

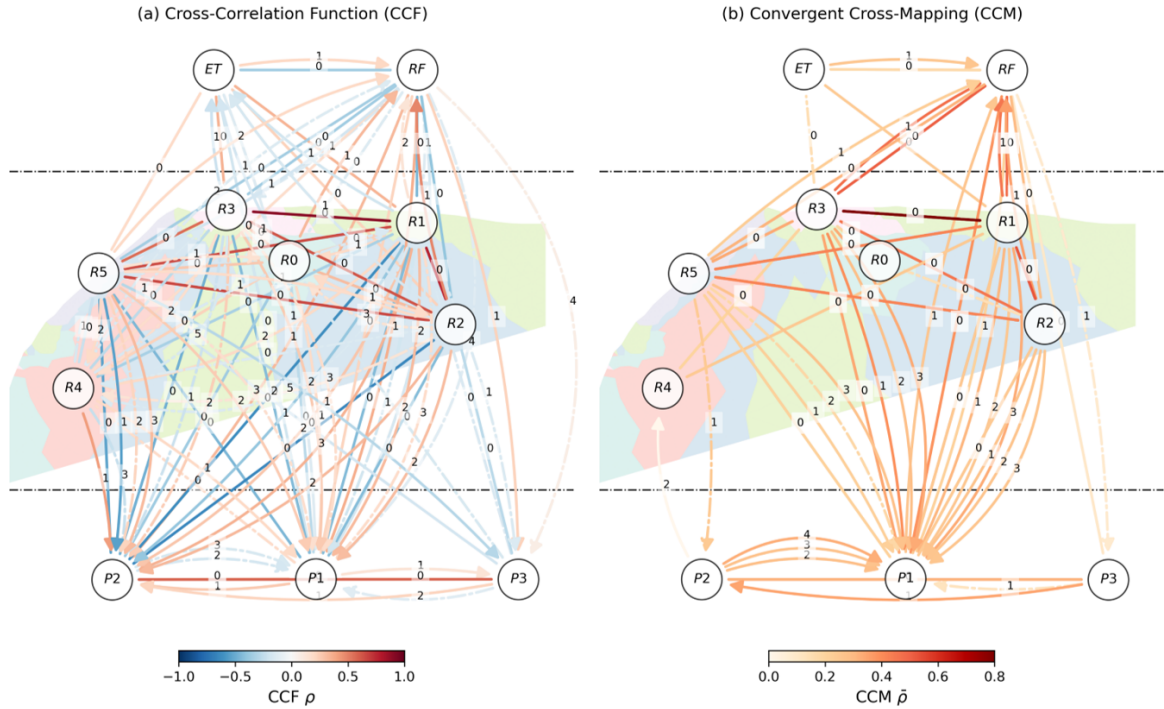


Figure 4. Graph of pairwise cross-dependencies: (a) with the linear Cross-Correlation Function (CCF), (b) with the nonlinear Convergent Cross-Mapping (CCM) method. An undirected line represents contemporaneous dependencies. Delayed dependencies are shown using directed curved arrows. All corresponding delays d are displayed in the middle of its corresponding arrow. The color of arrows maps to the strength of dependencies. Solid and dash-dotted arrows represent respectively significant dependencies with p -value < 0.001 and < 0.01 .

eters of the method, hoping to isolate more stable configurations (SM2.3). Whatever the configurations configuration, the results remained unstable. The instability is attributed to the lack of data and the fact that CMI, the most sensitive method, is applied to highly correlated data from a smooth inverted electrical resistivity model. This creates anomalies in the representation of causality between resistivity variables potentially due to the overly deterministic relationships that link these series. This problem is illustrated and further detailed in the supplementary materials (SM2.3). Consequently, to achieve a causal representation of the system with the nonlinear multivariate method, we had to adopt another logic, that of the consensus brought by the set of models from the sensitivity analysis. We considered all simulations done for the sensitive analysis as an ensemble of models rendering each a causal graph. Figure 6 reports the links achieving majority (50% considering a p -value of 0.05) among the ensemble of causal graphs from the sensitivity analysis. Following the consensus logic, the results again suggest a preferential connection with P1, while the two other drip discharge series remain unrelated.

5 Discussion

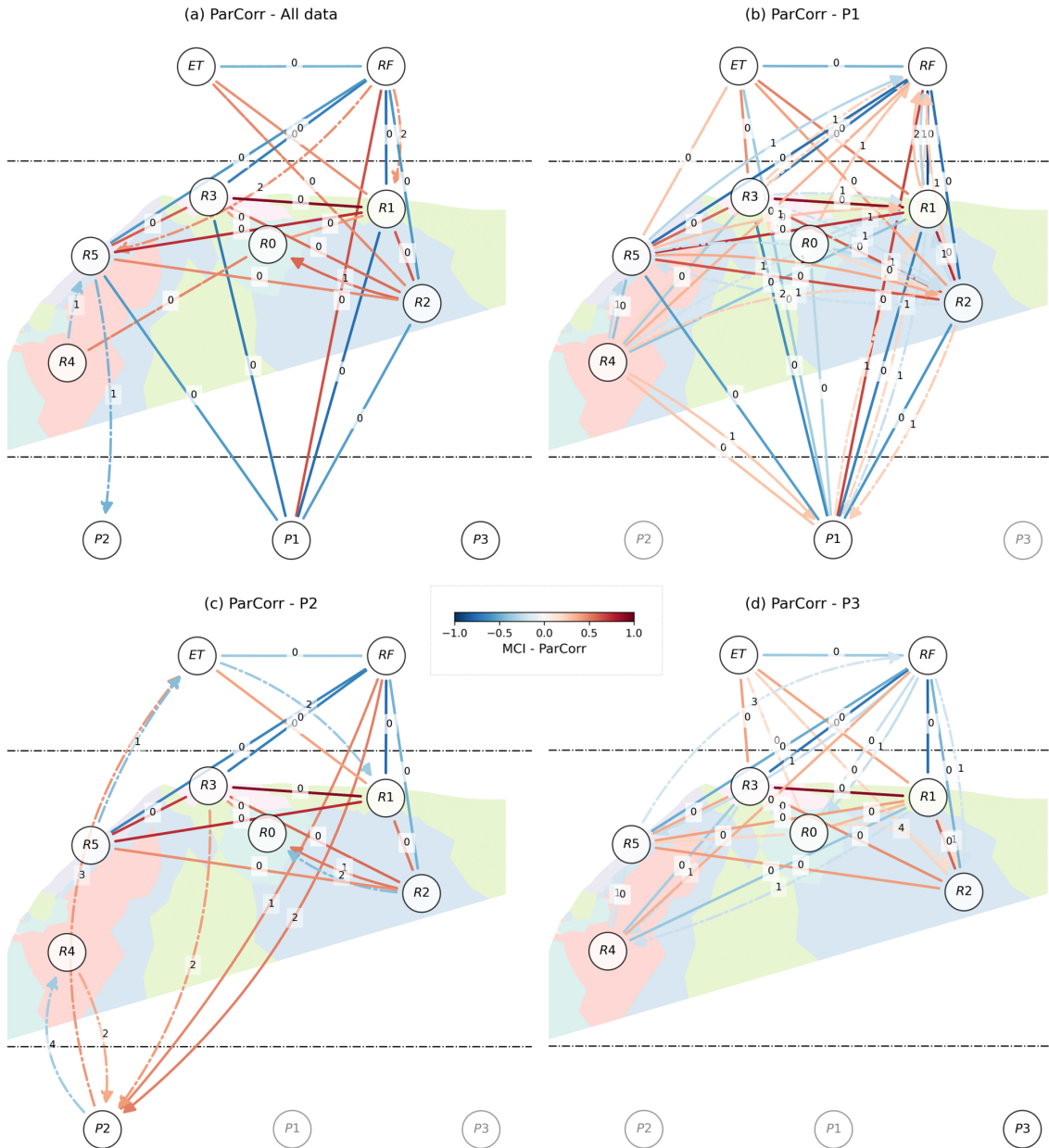


Figure 5. Graph of ParCorr cross-dependencies: considering (a) all data or one unique discharge series (b) P1, (c) P2, or (d) P3. An undirected line represents contemporaneous dependencies. Delayed dependencies are shown using directed curved arrows. All corresponding delays d are displayed in the middle of its corresponding arrow. The color of arrows maps to the strength of dependencies. Solid and dash-dotted arrows represent respectively significant dependencies with p -value < 0.001 and < 0.01 . For each graph, the size of the overlapping time-domain between the variables changes as follows: 48 days (a), 184 days (b), 62 days (c), and 218 days (d).

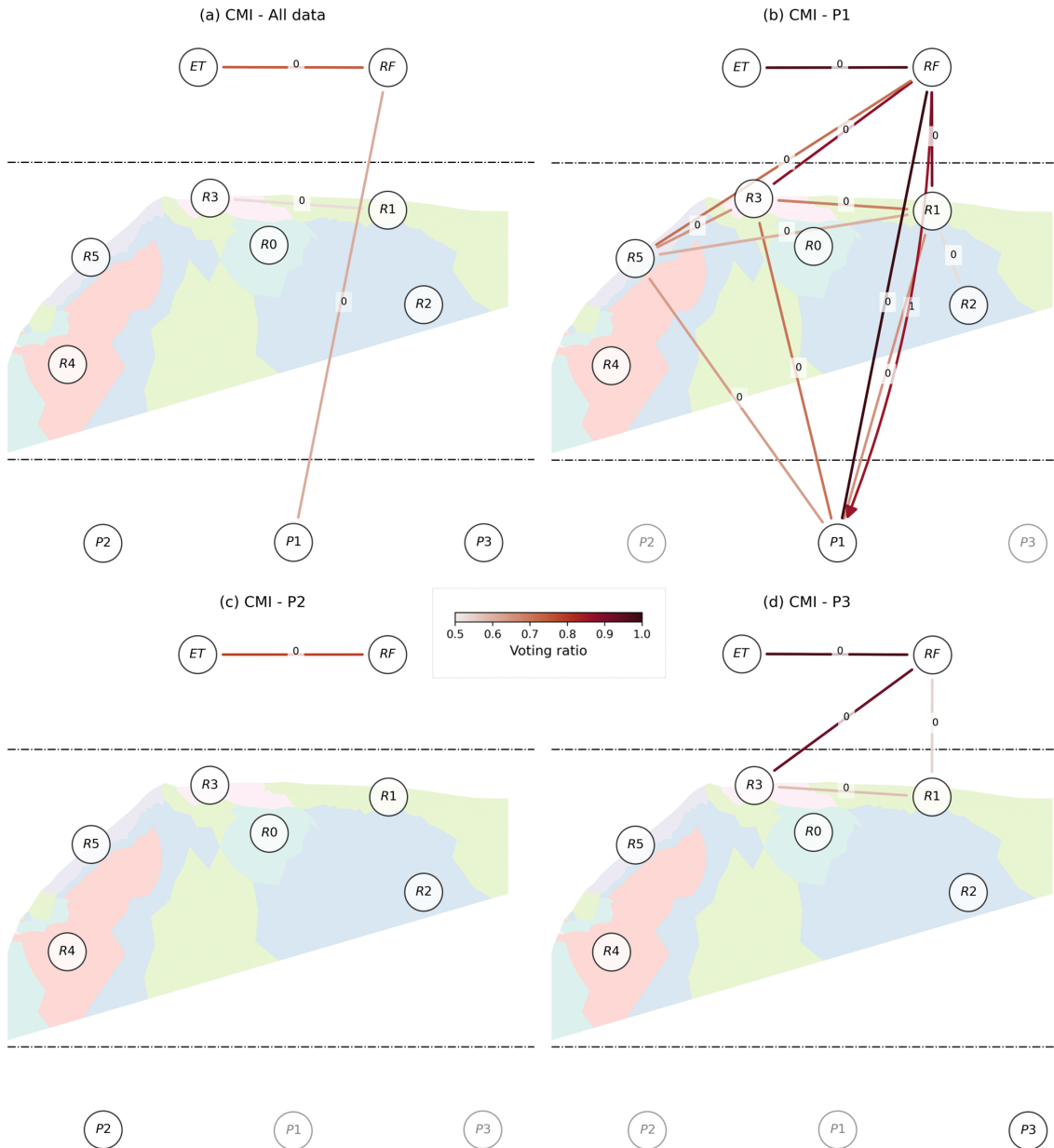


Figure 6. Consensual graph of CMI cross-dependencies obtained from the ensemble of simulations performed in the sensitivity analysis: (considering (a) all data or one unique discharge series (b) P1, (c) P2, or (d) P3. An undirected line represents the Contemporaneous dependencies. Delayed dependencies are shown using directed curved arrows. All corresponding delays d are displayed in the middle of its corresponding arrow. The links displayed are those reaching the majority (50%) for all causal graphs where the connections are established with a p-value of significance at 0.05. The color of arrows maps to the strength of the voting ratio. For each graph, the size of the overlapping time-domain between the variables changes as follows: 48 days (a), 184 days (b), 62 days (c), and 218 days (d).

~~We introduced CMI as the most relevant for effective connection detection. Our results showed that CMI differs from the other
480 CIMs in its low false-positive rate for the virtual case. However, CMI remains imperfect and missed a significant amount of
effective connections. Moreover, the method proved to be unstable in our real case study, forcing us to adopt a logic initially
absent from causal inference methods, that of a consensus graph.~~

5.1 CIMs' specific discussion

5.1.1 Cross-Correlation Function (CCF)

~~The CCF results for the real case, on the differenced data, show a dubiously high number of significant connections; we
485 interpret those as an indication that the method does not only reveal effective hydrological connections and reflects the
complexity of inferring connectivity from CIMs with strongly correlated, synchronous series forced by common drivers.
As a result, statistical dependencies, or functional connections, are ubiquitous. A corollary would be: one could commonly
build a point-to-point model based on two time-series and their statistical dependencies (functional connections) in hydrology.
490 Somehow, it explains why hydrological systems may be modelled with simple functional lumped model or linear transfer
functions (Dooge, 1973). However, when effective connections matter, e.g., to predict the spread of a pollutant, it might
become problematic. Similarly, Rinderer et al. (2018) reported many connections with the CCF method and bivariate Granger
Causality. Our primary explanation is that CCF on differenced data, i.e., with seasonality roughly removed, does not account
for the common forcing by rainfall, evapotranspiration, or residual harmonic signals. This failure leads to dependencies despite
495 the absence of causal links or effective connections (as in Fig. 3a), and the system's variables remain highly correlated and
synchronized.~~

~~Besides, we see that our results contain many False Positives reasonably identified from the sign of the dependencies
(Fig. 64a). The instability is possibly due to a large amount of missing data. Another hypothesis is that the ERT time-series
are too smooth and deterministically related, which could contribute to the instability of the method (Runge, 2018a). The
500 supplementary materials indeed show a strange behavior in the links with the CMI method, especially on the reduced time
domains due to missing data (see SM2.3, Fig. SM.5 P2, and Fig. SM.6, All data)fact that linear methods report signed
dependencies makes it easier to interpret the results and potential False Positives. For instance, positive relationships between
resistivity and drip discharge are unexpected if water transfers from low resistivity areas to the drip outlets. Often, positive
relationships, e.g., $R5 \rightarrow P2$ with $d = \{2, 3\}$, follow negative ones that are more interpretable as a transfer (e.g., $R5 \rightarrow P2$
505 with $d = \{0, 1\}$). This pattern is rather a phasing artifact captured by CCF interpretable as: "after the rain, the good weather",
and vice-versa. In general, we observe strong linear dependencies between resistivity series (except $R4$, the clayey limestone),
which may be a problem for PCMCI-CMI (see 5.1.4). Some links are more intriguing and are hardly interpreted as a causal
relationship, e.g., $R5$ causing RF .~~

5.2 **On the Practical Use of CIMs**

510 From this latter observation, we recall that dependencies can appear by chance, in particular, if the series are short, hence the importance of significance tests. We could have further controlled the number of links by considering lower p-values, but these were already as low as 0.001, or by comparing the dependencies obtained with those obtained with surrogates (Rinderer et al., 2018). Surrogates are random substitutes for the original data that share statistical or dynamic properties with the original data, for example, auto-correlation patterns (Schreiber and Schmitz, 2000). In our opinion, these approaches are convenient to rule out many connections that prevent us from focusing on the most significant ones. They will, however, not overcome the method drawback that is imputed to its design. As a bivariate method, CCF may fail to retrieve effective connections due to confounding factors (Runge et al., 2019a, and section 3.2).

Optimistically, we note that CIMs tend to reveal the preferential connection between $P1$ and the surface, as expected from the dye tracing test (Poulain et al., 2018, and Fig. 1.a). Yet, no statistical test provides the actual probability of a causal relationship, but rather that of an association under the assumptions of the CIM's underlying model and. Though clearly not perfect, CCF is simple, linearly interpretable, and widespread or popular. Considering that other CIMs carry their own imperfection, we do not discourage using CCF while knowing its limitations, testing for linearity, removing harmonics (or the use of suitable surrogates' comparisons), or confounding factors using either a statistical or physical model. Manual handling of confounding factors is close to assessing connectivity/causality in a multivariate framework, although not automated but supervised by our knowledge and expertise. Here, we only removed the seasonal signal by differencing, which is not sufficient and results in many dependencies.

5.1.1 Convergent Cross Mapping (CCM)

CCM results present fewer connections (Fig. 4.b), with many connections pointing from RF or resistivity series (except for $R4$, the clayey limestone) to $P1$. In particular, this result is encouraging as we expect a fast preferential flow and effective connection between the adequacy of the dataset, notably, surface and $P1$ (Poulain et al., 2018, and Fig. 1). Still, even when applied to deterministic time-series (Runge et al., 2019a), for a system not subject to synchrony (Sugihara et al., 2012, 2017), or, in such case, by using the time-delayed approach of Ye et al. (2015), CCM can be deceiving because it cannot deal with confounding factors, by design, as a bivariate method. Our result in the synthetic experiment tends to confirm it (Fig. 3.b). CCM identifies sustained significant dependencies with the not-differenced data imputed to slow seasonal variations. As for CCF, some of these dependencies may have been ruled out if we had considered surrogate significance testing with a surrogate model that accounts for seasonality (as seen in Nes et al., 2015; Huffaker et al., 2018; Medina et al., 2019). Yet, on the differenced data removing most of the seasonal signal, CCM captures significant time-dependencies at delays reflecting the response time of the reservoirs and not effective connections since they are disconnected. Besides, for the real experiment (Fig. 4.b), we found dependencies between drip discharge data (e.g., $P2$ and $P1$), which cannot be interpreted as effective connections, since they represent different outlets, similarly to the virtual experiment. In parallel, Ombadi et al. (2020) reported a high false-positive rate while using the original CCM approach (Sugihara et al., 2012), whatsoever the sample length, explained by CCM's inability to deal with confounding and synchrony.

In addition, noise is expected to reduce both the CCM's true and the false-positive rates, i.e., through the causal sufficiency hypothesis (Runge, 2018a). In other words, claiming that a link inferred from CIMs is a causal or effective connection is a delicate step. On this front, the CIMs are all equivalent, and they only differ on the types of connections that they can or cannot detect with a fair rate of false alarm. We demonstrate that the CMI method has a low general mapping skills (see Ombadi et al., 2020). Regarding our experiment as well, although hidden in the interquartile range shown in Fig. 3b, we found that higher noise levels led to lower $\bar{\rho}$. Theoretically, however, this decrease in performance might not be systematic. It may, in some cases, depend on the auto-correlated nature of noise and deterministic signals, the sample length, and the size of a Theiler window (Theiler, 1986). In general, we consider the idea of CCM being restricted to strictly deterministic systems (Runge et al., 2019a) to be overly conservative, potentially misleading, and impractical to real study cases. The theory does not impose anything on the practitioners, but it raises awareness of potential problems if the assumptions are not tested or violated, which is why it deserves careful considerations (further developed in Kantz and Schreiber, 2003; Sivakumar, 2017; Huffaker et al., 2018)

Currently, we consider CCM to be best suited to test whether or not there is a functional connection between two points, similarly to CCF but considering nonlinear dependencies. Therefore, our recommendations align with those of CCF. However, if CCM has better predictive capabilities than CCF, it can be concluded that the dependency is nonlinear.

5.1.2 PCMCI: Partial Correlation (ParCorr)

ParCorr is linearly interpretable and computationally efficient. As CCM, the ParCorr results for the real experiment (Fig. 5) seem promising because they generally favor the expected preferential flow and effective connection between the surface and $P1$ (Poulain et al., 2018, and Figure 1). ParCorr did not experience any stability problem, first, because ParCorr always provides the same results for the same dataset. Secondly, we deemed that the most significant connections remain between the results from the dataset variations that we tested (Fig. 5.a-d). Although there are few differences between the links that are evaluated in the four variations, these are often related to relationships of lesser significance or magnitude. Some of them are alarming, such as $R5 \rightarrow RF$ at lag $d = 1$ in Fig. 5.b or $d = 3$ in Fig. 5.d. They are even more alarming because the whole causal graph can be affected by conditioning on an irrelevant variable (Eq. 2).

Based on the results of the synthetic experiment (Fig. 3 and Table 3), PCMCI-ParCorr, or its variant multivariate GC, may suffer from a relatively high false-positive rate (> 50% in Table 3), most likely due to the inability to deal with direct nonlinear dependencies. Similarly, Ombadi et al. (2020) reported a high false-positive rate for GC in their synthetic experiment, although below CCM's one. A high rate was also observed in Rinderer et al. (2018), although using bivariate GC. Accordingly, claiming an effective hydrological connection based on PCMCI-ParCorr requires caution because of the high false-positive rate reported here and in the literature. We consider that its applications would require deeper testing to ensure that a multivariate linear model fits the data.

5.1.3 PCMCI: Conditional Mutual Information (CMI)

575 The result of the virtual experiment (Fig. 3 and Table 3) confirmed our general expectation that a multivariate nonlinear method is best suited to assess effective connectivity. PCMCI-CMI had the lowest false-positive rate (Table 3), which is particularly desired given the confounding problem in hydrological systems. Yet, it had a low recall relative to other approaches, meaning that the results contain False Negatives. Ombadi et al. (2020) also reported a relatively low false-positive rate compared to the others using different methods, PC (Spirtes and Glymour, 1991) and the bivariate conditional Transfer Entropy (TE) method (Schreiber, 2000). Still, ~~bivariate methods are helpful if only to select potential links and check which ones are dismissed while using multivariate CMI. The value of CCM compare to CCF is the opportunity to reveal weak nonlinear interactions for cases where such relations are suspected. Besides, linear methods are faster, more interpretable, and explainable to an audience~~ according to Runge et al. (2019b), PCMCI-CMI is supposed to perform better than PC thanks to its sequential procedure and TE by being multivariate and accounting for confounding effects.

585 In the real case study, PCMCI-CMI was found unstable, providing different results across consecutive runs (see, Fig SM.5 and SM.6). Two main reasons might explain this instability. First, PCMCI conditions dependencies on the Parents (Eq. 2) and, therefore, build a dataset containing the initial variables and their delayed versions up to $2d_{max}$. The overlapping time-domain over which all variables and delays are defined is small (see Fig. 1). It covers only 48 up to 218 time-stamps (P3). Ruddell and Kumar (2009) reported that 500 to 1000 samples are generally sufficient to obtain a qualitatively robust estimate of the TE, estimated using a binning approach. The nearest-neighbor approach that we used (Runge, 2018b) is more suited for short datasets. Yet, we found that our final sample sizes, accounting for missing data, are lower and the dimensionality is potentially higher, though variable and depending on the size of the Parents' sets (Eq.2), than those tested in the evaluation of the CMI estimator (Runge, 2018b). This could be the main reason for the instability. In addition, we applied our analysis to highly correlated data from a smooth inverted electrical resistivity model (Fig. 4.a). PCMCI-CMI encountered highly interdependent anomalies regarding resistivity (reported in Fig. SM.5 and 6). We suspect that these anomalies may be related to a violation of the faithfulness hypothesis (Runge, 2018a). PCMCI-CMI may fail because the resistivity series ($R1$ to $R6$) are overly deterministically related while accounting for nonlinear dependencies, i.e., ~~many reasons that could favor their practical use.~~ without a sufficient stochastic variability. This complementary reason for the instability is further illustrated and detailed in the supplementary materials (SM2.3).

600 In general, ~~the CMI allow us to focus on a limited number of relationships: the strongest, the most cross-predictable, the most robust, or consensual if the approaches or parameters are varied.~~ For this reason, the high number of connections found by the CCF method (Fig. 1) is not surprising. For short datasets, or with unevenly distributed missing values, we consider that building a consensual graph from multiple runs is a good strategy as our results suggested a fast preferential flow in the case of P1 only (Fig. 6.a and b), as expected (Poulain et al., 2018, and Fig. 1). This is, however, computationally expensive. Reducing the number of variables involved in the analysis is another option. It reduces the dimensionality and limits the reduction of the overlapping time-domain due to missing values. Considering a bivariate and nonlinear conditional case (i.e., TE), the low false-positive rate obtained by Ombadi et al. (2020) is encouraging in that regard. Besides, regarding multivariate approaches, Rinderer et al. (2018) performed the conditioning on the main confounding variables, i.e., effective precipitation, only. This

strategy would deserve further consideration, since it makes the study of effective connectivity from multivariate CIMs between two points in the system a systematic and potentially efficient "three-body problem" representation.

We generally consider that multivariate nonlinear CIMs are preferred to assess effective connectivity based on the theoretical background and the results obtained from our synthetic and real experiments. Still, PCMCI-CMI may miss effective connections because of its low recall evidenced by the virtual case (Table 3) and the low number of arrows reported in our consensual graph (Fig. 6). ~~4.a) should act as a warning light. Nonetheless, and this is true for all methods, we can adjust the number of causal connections retrieved and limit the results to~~ In addition, other assumptions may be violated (Runge, 2018a). In particular, for a real study case, one cannot test the hypothesis of causal sufficiency; that is, all common drivers should be included in the analysis. In short, we consider causal sufficiency to be challenging to test and conceptualize since hydrological variables can also be represented in a spatially explicit manner in a high-dimensional continuum. CIMs go hand in hand with a dimension reduction task like ours (Delforge et al., 2020b). In the end, there remain conceptual uncertainties and a risk that these CIMs may provide relationships or connections that are spurious and not effective.

5.2 Further limitations, recommendations, and perspectives

For our comparative study, we limited ourselves to specific methods and set aside some of their hypotheses to focus on fewer relevant elements for potential CIMs' users in hydrology. More details can be found regarding the underlying frameworks of the newly introduced CIMs, for CCM and/or the chaos theory (Kantz and Schreiber, 2003; Sugihara et al., 2012; Sivakumar, 2017; Huffaker et al., 2018), the PCMCI framework (Runge, 2018a; Goodwell et al., 2020), or both (Runge et al., 2019a). In particular, we assumed but did not discuss stationarity and its various definition depending on the framework. We did not test the significance of our results using surrogate data test, which stands as a common practice (e.g., Schreiber and Schmitz, 2000; Huffaker et al., 2018). We did not address the impact of noise quantitatively (see Ombadi et al., 2020, for that matter). Also, we hypothesized sustained interactions between variables and did not discuss dynamic intermittency, which could be explicitly visible, such as zeros values in rainfall time-series (Sivakumar, 2017), or hidden and imputed to threshold effects (Blöschl and Zehe, 2005). Arguably, hydrological connectivity is highly dynamic and potentially intermittent with some portion of the hydrological system getting connected and disconnected depending on its state (Bracken et al., 2013). To explore dynamic intermittency, applying the CIMs on hydrograph segments (e.g., high or low flows; see also Delforge et al., 2020a), or for different seasons (e.g., Ombadi et al., 2020) are two potential options provided that the sample size requirements are met.

Besides, the selected CIMs operate on the time-domain, limiting us to studying close temporal connections. Yet, hydrological connections and processes can be spread over much longer time-scales. This further questions our tacit assumption of causal sufficiency. Accordingly, studying methods that operate on the frequency domain or that couple the frequency domain with the time-domain may deserve a particular interest (e.g., Granger, 1969; Molini et al., 2010; Rinderer et al., 2018).

The PCMCI algorithm is still actively developed. A more recent implementation of PCMCI is now available in the new version of the Tigramite Python package (v4.2), with refined default parameters and improved computational performance. In particular, a new algorithm, PCMCI+, deals with contemporaneous links and strong auto-correlation in series, with the promises of stronger recall and well-controlled false-positive rate (Runge, 2020). Besides auto-correlation, we found that

contemporaneous links are numerous and compromise the recovery of causal direction based on the principle of priority. As contemporaneous links may concern hydrological systems at all spatiotemporal scales, we recommend exploring PCMCI+ for future studies. In particular, while using PCMCI-CMI or any CIM rooted in the information theory, specific attention should be paid to the issue of small sample size (or temporal gaps from missing values). We did not provide any heuristic values linking robustness and the sample size. Such an indicator is contextually linked to the study case and depends on many elements. From the scope of statistical inference with PCMCI, it depends on the Parents' preselection procedure (PC and its parameter α_{PC}), the strongest dependencies by selecting a proper method-parameterization: lowering the p-values, using more restrictive tests, or confronting the statistical dependencies with those obtained from surrogate data sets (Schreiber and Schmitz, 2000). The predictive potential of the most significant links gives them value, and their robustness should prevent futile discussions on an evanescent singularity. Altogether, it appears difficult to rely on any particular link obtained with a single method and in a unique configuration. Indeed, this study showed how volatile links CIMs could be. Therefore, CIMs' output, like many others in statistics, should be taken with caution. Their robustness should be assessed by varying the methods, CMI estimator and its parameters, the number of variables, the maximum causal delay d_{max} , the nature of the dependencies through their joint probability distributions, or signal/noise characteristics and ratios. From the scope of the hydrological system, it relates to its complexity reflected in the spatiotemporal scale of the hydrological problem, the heterogeneity of its geomorphology (or the patterns of structural connectivity), the datasets, or the time domains to address the robustness of the links. On the other hand, we cannot exclude that an elusive but actual causal connection would be only detectable by one or the other method, using the right parameterization. nonlinearity of hydrological interactions, their number and spatial organization. Hence, we instead recommend testing the robustness of the results for each particular study case. Potentially, combining any CIM with virtual models mimicking the dynamic properties of the data while knowing the actual causal graph may help select the right CIM, its related estimator, and adequate parametrization. Yet, our toy model (Fig. 2) was not meant for that purpose but to evaluate different CIMs with a parsimonious representation of the confounding problem in hydrology.

By being labeled as causal, we may have more expectations of CIMs than other types of experiments or investigation methods, hoping to get simple answers to complex problems from them. Nevertheless, A final consideration is more epistemological: should hydrological connectivity (or causality) be studied from a purely empirical and single automated perspective, as with CIMs? We remind that all types of methods can contribute to our causal understanding of environmental systems (e.g., dye tracing tests or spatially detailed inverse resistivity models) (Bakalowicz, 2005; Watlet et al., 2018; Poulain et al., 2018). However, the potential sensitivity of CIMs to their assumptions and parameterization, parameters, or conceptual uncertainties makes them hazardous to use alone. Despite their limitations, CIMs, despite their recent noticeable improvements. For this reason, even if PCMCI-CMI appeared to stand out, we recommend comparative studies using several CIMs. Linear CIMs may remain attractive for their intelligibility and computational efficiency. Bivariate methods remain helpful to realize what functional connections multivariate CIMs exclude. We further consider that CIMs complement other methods, beyond time-series analysis, e.g., field experiments or physically-based approaches, and they could be combined to narrow the range of possible causal representations of the system under study.

5.3 Research Perspectives

Today, CIMs benefit from a growing interest; they evolve and progress. CIMs pursue the ideal of causality solely inferred by data and machines. They are very generic methods and should be studied and appropriated for the specificity of hydrological data. In general, Regarding CIMs only, Klemeš (1982) was particularly critical of letting the data speak for itself. ~~For the multivariate methods, we have chosen to let the causal graphs be formed.~~ Two avenues are possible and should be kept in mind. As mentioned for CCF in section 5.1.1, we see no issues in applying a bivariate method, with awareness of its limits and appropriate hypothesis testing, in a sequential way, e.g., while removing confounding factors using either a statistical or a physically-based model. That is a supervised causal inference. Another option is to integrate physical constraints in the causal inference algorithm. Here, we adopted the CIM's philosophy of letting the causal graphs be inferred from the data ~~alone.~~ We alone. For PCMCI, we have not prescribed any constraint on the conditioning of variables. ~~This means that variables can be conditioned on potentially aberrant links, negatively impacting the whole.~~ Yet, physically irrelevant Parents (Eq. 2) may negatively impact the causal graph. ~~Studying how to constraint causal inference is an exciting research perspective. It should be done with caution.~~ Constraining must be framed to prevent us from ~~constraining the analysis forcing CIMs~~ on perceptually biased ~~hypotheses on the system's assumptions on the systems'~~ functioning. In particular, this could be done by reintroducing some physical concerns or spatial dimensions into the analysis. Rinderer et al. (2018) already proposed to constrain CIMs with structural connectivity, ~~which is the first level of potential connectivity allowed by the geophysical environment. Currently, distances nor the length of flow path or energy potentials do not.~~ However, structural connectivity is unknown, hidden, or too complex to be hypothesized for most karst systems (Bakalowicz, 2005; Hartmann et al., 2014). Still, mass balances, energy potentials, or spatial attributes such as distances or flow path lengths, none currently matter in CIMs. ~~This~~ The spatial dimension is initially present in Hume's contiguity principle in time and space (Hume, 1748). To Schrodinger, the spatial continuum is also a causal paradigm in physics (Schrödinger, 1954). Then, we recommend research avenues on reconciling CIMs with space and physics in geosciences.

~~Other research perspectives are more fundamentally related to the nature of the hydrological data.~~

1. ~~Long-time-scale. The methods proposed here work on the temporal domain. However, the assumption of causal sufficiency is questionable since hydrological connections can be spread over much longer time-scales. Thus, methods that work in the frequency domain or that couple the frequency domain with the time-domain deserve a particular interest (e.g., Molini et al., 2010).~~
2. ~~Intermittent connectivity. We did not address the nonlinearity associated with threshold effects or intermittent hydrological connections (Blöschl and Zehe, 2005; Bracken et al., 2013). The CIMs assume constant interactions, even if modulable in the case of nonlinear methods. Since a hydrograph at the output of a hydrological system is a good representation of the hydrological state of the system, applying the CIMs on a segmentation of the hydrograph (e.g., high or low flows) is another interesting application. Segmentation would, however, reduce the available temporal domain for conditioning the variables, similarly to missing values in this study.~~

- 710 3. ~~Contemporaneous dependencies. Most hydrological dependencies are contemporaneous, preventing the inference of causality from the principle of priority. Some CIMs are not built upon the principle of priority (e.g., Spirtes et al., 2000; Pearl, 2009; F~~
~~and could be investigated. In particular, the Tigramite package has been recently updated to include an improved and~~
~~faster PCMCI+ algorithm that deals with contemporaneous links and strong auto-correlation in series, with the promises~~
~~of stronger recall and well-controlled false positive (Runge, 2020).~~
- 715 4. ~~Statistical models and their parametrization. The CMI method deserves to be studied in more detail for its potentiality~~
~~revealed in the virtual experiment. There are other estimators of the CMI that could be more adapted to longer series or~~
~~other time-scales (Runge et al., 2019b). Furthermore, the parameterization of estimators could be studied in more detail,~~
~~supported by the current interest in hydrology about the information theory (see Goodwell et al., 2020, and other debaters)~~
~~. Virtual experiments also provide an opportunity to study in more detail how to parameterize CIMs.~~
- 720 5. ~~Inversion or processing artifacts. Many hydrological datasets are not directly measured but inverted, modeled, or processed~~
~~so that artificial dependencies between the data may exist. These dependencies should not be confused with causal~~
~~relationships, and further study could be conducted to assess the sensitivity of CIMs to these artifacts and propose~~
~~strategies to mitigate the problem.~~

6 Conclusions

725 We applied four causal inference methods for detecting hydrological connections. These CIMs study the temporal dependencies
between variables. They are either linear or nonlinear, and bivariate (pairwise dependencies) or multivariate (conditional
pairwise dependencies). The four CIMs are applied to both a synthetic and a real case in a karstic study site. The synthetic
data are effective precipitation and modeled discharge from two reservoirs, either parallel and disconnected or contributing
in series to the same drainage channel. The real karstic dataset involves precipitation and evapotranspiration data, subsurface
730 resistivity series clustered from an inverted electrical resistivity model associated with a time-lapse ERT dataset, and three drip
percolation discharge series in the cave below the surface. In this case, we know that a preferential flow exists between the
surface and the one spot of drip discharge in the cave.

For the synthetic case, bivariate methods cannot discriminate causal dependencies from those arising from confounding by
the meteorological forcing of the model. Bivariate methods generally only account for potential connections, while multivariate
735 methods attempt to extricate actual ones. In accordance with our theoretical expectations, we found that the nonlinear and
multivariate CIM based on conditional mutual information (CMI) is indeed the most precise: the identified connections are
likely to be actual connections. Yet, this method has a bad recall: it misses many connections. Moreover, this same method
proved to be unstable for

The results highlighted that the real-easenonlinear multivariate method, PCMCI coupled with the Conditional Mutual
740 Information test (PCMCI-CMI), shines by its low false-positive rate relative to the other three methods. Hence, statistical
dependencies revealed by PCMCI-CMI are more likely to be effective hydrological connections. We believe that the temporal

domain, further reduced by missing data and the conditioning process, was too small to converge. Possibly, the instability is also due to the electrical resistivity model and the clustered series being too smooth, correlated, and therefore too deterministically related to establishing causality based on the concept of conditional independence. To overcome this issue, we built a consensual causal graph from multiple iterations of the method and its parameters, with the outcome in phase with our perceptual understanding of the system. Yet, consensus introduces a new logic in the causal inference process that is not part of the initial theory. The causal graph of the linear bivariate method, This advantage is particularly valuable since hydrological systems present highly interdependent time series (or functional connections), favoring a high false-positive rate. This finding confirms our introduced expectation that multivariate nonlinear CIMs are best suited to infer effective connectivity while dealing with nonlinear dependencies and confounding factors resulting from seasonality or meteorological forcing. However, PCMCI-CMI has a low recall, i.e., it misses effective connections, and particular attention should be paid to the robustness of the outcome for small sample sizes or temporal gaps in the data, as evidenced in our real study case. Furthermore, PCMCI-CMI relies on challenging hypotheses to test (Runge, 2018a). Given these uncertainties, PCMCI-CMI, like any other CIM, would always present a risk of spurious results. For this reason, we do not discourage the use of other CIMs as well, for a comparison purpose, with awareness of their limits. Alongside other limitations, recommendations, and perspectives, we remind that framing the causality of hydrological systems is not restricted to the cross-correlation function (CCF), shows ubiquitous dependencies between variables. Not much can be learned about connectivity and preferential flow paths without considering more restrictive tests. The three other methods support the idea of a preferential connection where it is suspected. use of the strictly empirical CIMs, but benefits from the vast panel of scientific investigation techniques (e.g., Bakalowicz, 2005; Bracken et al., 2013). In line with Rinderer et al. (2018), hybrid approaches could be developed by reintroducing the physical aspects of the problem to exclude or control the risk of CIMs physically irrelevant outcomes.

~~Nonetheless, any multivariate CIM will assess causality under the hypothesis of causal sufficiency, meaning that all the potential causes are monitored alongside other hypotheses. The sufficiency of the data is hardly verified in practice, as well as the complete adequacy of one particular CIM and its parameters. As a result, different causal links may appear with different configurations. Accordingly, it is delicate to interpret a causality test as the probability of an actual hydrological connection. While caution and flexibility are always warranted, CIMs are interesting data mining methods, undoubtedly capable of making causal discoveries and constantly improving. We propose suggestions for research perspectives considering the specificity of hydrological dynamics. In the meantime, we promote non-exclusive approaches while using CIMs by varying the methods, their parameters, the time domain, and the data. They participate in a causal understanding of hydrological systems, but causal understanding is not limited to them. They synergize with other leading approaches such as physically-based modeling, other empirical data analysis approaches, or field investigations.~~

Code and data availability. CCF and the Student's t-test are computed using the Scipy Python package (Virtanen et al., 2020). The CCM python implementation is available from Delforge et al. (2020a). The official R implementation is available from the CRAN repository: <https://CRAN.R-project.org/package=rEDM>. PCMCI and independence tests are implemented within the Tigramite (v.4.1 in this case) Python

775 package: <https://jakobrunge.github.io/tigramite/>. Evapotranspiration data were obtained from the agrometeorological PAMESEB network for the station of Jemelle: <https://agromet.be>. All other environmental time-series can be obtained from Watlet et al. (2018) and the related repository: <https://zenodo.org/record/1158631>. Resistivity clustered time-series can be reconstructed following Delforge et al. (2020b) and the example available from the repository: <http://dx.doi.org/10.17632/zh5b88vn78.2>

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