



# 1 Quantifying the uncertainty of precipitation forecasting

# <sup>2</sup> using probabilistic deep learning

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9 Abstract. Precipitation forecasting is an important mission in weather science. In recent years, data-10 driven precipitation forecasting techniques could complement numerical prediction, such as precipitation 11 nowcasting, monthly precipitation projection and extreme precipitation event identification. In data-12 driven precipitation forecasting, the predictive uncertainty arises mainly from data and model 13 uncertainties. Current deep learning forecasting methods could model the parametric uncertainty by 14 random sampling from the parameters. However, the data uncertainty is usually ignored in the forecasting 15 process and the derivation of predictive uncertainty is incomplete. In this study, the input data uncertainty, 16 target data uncertainty and model uncertainty are jointly modeled in a deep learning precipitation 17 forecasting framework to estimate the predictive uncertainty. Specifically, the data uncertainty is 18 estimated a priori and the input uncertainty is propagated forward through model weights according to 19 the law of error propagation. The model uncertainty is considered by sampling from the parameters and 20 is coupled with input and target data uncertainties in the objective function during the training process. 21 Finally, the predictive uncertainty is produced by propagating the input uncertainty and sampling the 22 weights in the testing process. The experimental results indicate that the proposed joint uncertainty 23 modeling and precipitation forecasting framework exhibits comparable forecasting accuracy with 24 existing methods, while could reduce the predictive uncertainty to a large extent relative to two existing 25 joint uncertainty modeling approaches. The developed joint uncertainty modeling method is a general 26 uncertainty estimation approach for data-driven forecasting applications.

# 27 1 Introduction

28 Precipitation is a key hydrometeorological variable in earth system science, and is the main driving





29 factor of floods and droughts (Xu et al., 2019). In the year of 2019, the flood disaster driven by extreme 30 precipitation caused a direct economic loss of 29.6 billion dollars in China, and the drought disaster led 31 to a crop production loss of 23.6 billion kilograms (http://www.mwr.gov.cn/sj/#tjgb). Accurate 32 precipitation forecasting is vital for the early warning of flood and drought, smart city management and 33 agricultural water resources allocation (Van Den Hurk et al., 2012; Pozzi et al., 2013). However, the 34 precipitation forecasting problem suffers from uncertainties from data, algorithms and random factors 35 (Reeves et al., 2014; Kobold and Sušelj, 2005; Xu et al., 2020b). The predictive uncertainty is a 36 measurement of the spread of precipitation forecasting and could indicate how much the forecasted 37 precipitation values fluctuate around the mean (Papacharalampous et al., 2020). Therefore, the 38 uncertainty range should be given when generating precipitation forecasting results.

39 The precipitation forecasting methods can be divided into two categories: numerical weather 40 forecasting and statistical machine learning. Numerical models consider the physical process of earth 41 system and could simulate the interactions between atmospheres, oceans and lands (Sikder and Hossain, 42 2016; Molinari and Dudek, 1992). Numerical models have strong physical meaning and are the dominant 43 ways of operational precipitation forecasting. However, the forecasting ability of numerical models is 44 limited due to the uncertainty in initial and boundary conditions, the imperfection of parameterization 45 schemes and the uncertainty in parameters (Reeves et al., 2014). With the development of computer 46 technology and machine learning algorithms, using random data-driven techniques for precipitation 47 forecasting is becoming popular in recent years (Shi et al., 2015; Trebing et al., 2021; Sønderby et al., 48 2020). The accuracy of data-driven methods is comparable to currently advanced numerical models in 49 short-term (e.g. from hours to weeks) precipitation forecasting. For example, the convolutional long-50 short term memory network is shown to outperform the physical optical flow method in precipitation 51 nowcasting based on radar images (Shi et al., 2015). Another deep learning model called MetNet showed 52 advantages over traditional numerical models in terms of the forecasting accuracy and running time for 53 hourly precipitation prediction (Sønderby et al., 2020). The data-driven methods also exhibit appealing 54 results in subseasonal to seasonal precipitation forecasting relative to numerical models (Boukabara et 55 al., 2019; Chantry et al., 2021; Hwang et al., 2019). A key drawback of data-driven precipitation 56 forecasting method is the lack of physical meaning, also known as black-box model. Despite this feature, 57 data-driven statistical machine learning methods have been widely used for parameter calibration, data





58 processing, submodel replacement and process understanding among physical simulations (Ardabili et 59 al., 2019; Sahoo et al., 2017; Reichstein et al., 2019). The data-driven learning techniques are strong 60 complements to numerical models for the improvement of precipitation forecasting accuracy. 61 The predictive uncertainty in precipitation forecasting arises mainly from data and models (Gal, 62 2016). The data uncertainty comes from external observation conditions, instruments and processing 63 algorithms. The data uncertainty is usually examined by perturbing initial conditions in numerical models 64 and producing a perturbed multi-model ensemble, which is widely seen in hydrometeorological ensemble forecasting (Xu et al., 2019; Gneiting and Raftery, 2005; Duan et al., 2019; Vitart et al., 2017). The data 65 66 uncertainty is rarely investigated in data-driven precipitation forecasting and is often assumed to be 67 accurate without error. The model uncertainty is often represented by an ensemble of perturbed model 68 physics and parameters in numerical weather forecasting (Vitart et al., 2017; Kirtman et al., 2014; Taylor 69 et al., 2012). In data-driven models, the model uncertainty is generally modeled by random regularization 70 of parameters (Gal, 2016; Kendall and Gal, 2017). For linear regression, the parametric uncertainty is 71 indicated by the standard deviation of trained parameters. In deep learning, the network layers could be 72 randomly abandoned to prevent overfitting and generate a forecasted ensemble by Monte Carlo sampling 73 (Kendall and Gal, 2017; Srivastava et al., 2014; Loquercio et al., 2020; Ghahramani, 2015).

74 The data and model uncertainties should be considered jointly in an integrated modeling framework 75 to get the predictive uncertainty, as the data and model uncertainties could both inflate the predictive 76 spread considerably (Gal, 2016; Kendall and Gal, 2017). It is expected that, the forecasting result would 77 be more or less different if the used data and parameters are randomly sampled from the population. Data 78 uncertainty is usually assumed as a constant or Gaussian distribution and could be propagated into final 79 forecasting through error forward propagation (Loquercio et al., 2020; Xu et al., 2020a). If the data 80 uncertainty is unknown, it can be learned from the training process by considering the data uncertainty 81 as a trainable parameter (Kendall and Gal, 2017). However, the joint learning of data errors and model 82 weights will increase the number of training parameters and may mix the error flow from data and 83 parameters. A prior estimation of data uncertainty could help unravel the data error and facilitate the 84 training process. On the other hand, previous forecasting studies usually model the input data uncertainty 85 and ignore the uncertainty in the target (predictand) data (Kendall and Gal, 2017; Loquercio et al., 2020). 86 The uncertainty in the target dataset also plays an important role in the parameter training process and





87 could influence the forecasting accuracy.

88	There are two ways to estimate the data uncertainty. One is to use in-situ ground stations to calculate
89	the systematic and random errors within the data and the other is to use multisource datasets to compute
90	random error by intercomparison (Xu et al., 2021; Gruber et al., 2016; Sun et al., 2018). The in-situ
91	validation method is limited to the number and density of ground stations and are suitable for small areas
92	with enough station coverage. The second method is independent of the in-situ stations and requires
93	multisource datasets with independent error distribution (Gruber et al., 2016). There are numerous
94	precipitation datasets from various sensors and models and could be used to calculate precipitation data
95	error at a large spatial scale (Xu et al., 2020b; Sun et al., 2018). Three-cornered hat (TCH) and triple
96	collocation (TC) are two commonly used methods to evaluate the random error among multisource
97	datasets, which do not require ground measurements as references (Premoli and Tavella, 1993; Mccoll
98	et al., 2014; Stoffelen, 1998). The basic assumption of TCH and TC methods is the stationarity of both
99	the raw dataset and its error, which may not be always satisfied for real-world data. Most of the existing
100	studies assume that the used multisource datasets obey the stationarity condition when using TCH or TC
101	methods (Xu et al., 2020b; Gruber et al., 2016; Gruber et al., 2017), which is useful for the determination
102	of relative prior random error.

103 In this study, we aim to quantify the predictive uncertainty of data-driven precipitation forecasting 104 by fully considering the uncertainty from data and models. The data uncertainty is estimated by the TCH 105 method a priori and is assumed as Gaussian distribution. The data uncertainty is propagated within model 106 training by the law of error propagation. The parametric uncertainty is modeled by randomly abandoning 107 some network layers during the training process. The data and model uncertainties are jointly considered 108 in the objective function within a deep learning encoder-decoder framework. The forecasting 109 experiments are conducted to see whether the accuracy of precipitation forecasting can be improved by 110 joint data-model uncertainty modeling relative to several uncertainty processing strategies from the 111 existing studies.

#### 112 2 Study area and data

The study area is located at southern and northern China, East Asia (Figure 1). The annual rainfall 113 decreases from the southeast to the northwest, with an approximately average rainfall of 1500 mm in the 114





115 southeast regions and 300 mm in the northwest areas. Most of the southern areas feature a subtropical 116 monsoon climate and the rainfall is relatively larger in summer and smaller in winter. From June to July in 2020, extreme precipitation hit the southern China (Wei et al., 2020) and caused a direct economic 117 118 loss of 13.2 billion dollars. The precipitation forecasting in southern area of China is very challenging 119 and meaningful. Previous studies use numerical models for precipitation forecasting in this area and show 120 some values (Yuan et al., 2012; Luo et al., 2017). The northern area of China features the temperate moon 121 and continental climates, with an annual rainfall of 400 to 800 mm and the main rainy season of July and 122 August. Here we would like to explore the possibility of weekly precipitation forecasting by a data-driven 123 deep learning method.





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Multisource precipitation datasets are used here to obtain the data error and to measure the precipitation forecasting ability of different uncertainty processing strategies, including the Modern-Era Retrospective Analysis for Research and Applications, version 2 (MERRA-2) (Gelaro et al., 2017), the National Centers for Environmental Prediction Reanalysis version 2 (NCEP R2) (Saha et al., 2014) and the European Centre for Medium-Range Weather Forecasts Reanalysis version 5 (ERA-5) (Hersbach et al., 2020) datasets from 1980 to 2020. The surface 2-meter temperature and the geopotential height at 500 hPa datasets are collected from the three datasets accordingly as predictors. All these daily datasets





- 133 are converted to weekly data and are bilinearly interpolated into 0.25° resolution. In the forecasting
- 134 process, the temperature and geopotential height predictors in the historical three consecutive weeks are
- 135 used to forecast the precipitation in the target week.

#### 136 3 Methods

137 **3.1. Estimation of data uncertainty** 

138 The TCH method (Xu et al., 2020b; Premoli and Tavella, 1993) is used to estimate the uncertainty in temperature, geopotential height and precipitation datasets. The collected three datasets are all 139 140 reanalysis data, which is generated from different physical models and data assimilation algorithms. The 141 different reanalysis datasets and their errors are generally not closely correlated and are regarded as 142 collocated datasets for the uncertainty estimation, similar with existing studies (Xu et al., 2021; Mccoll et al., 2014; Gruber et al., 2017). In the TCH algorithm, one arbitrary dataset is chosen as the reference 143 144 among the three datasets, and then the differencing operation is conducted between the reference and the 145 other two datasets to get the differencing series. The covariance of the differencing series is connected 146 to the variance-covariance matrix of precipitation datasets through matrix transformation. The 147 parameters of the variance-covariance matrix are iteratively resolved by minimizing the global 148 correlation of the covariance of the differencing series. A detailed introduction of TCH method could 149 refer to Premoli and Tavella (1993) and Xu et al. (2020b).

150 The uncertainties of the predictors and predictands are estimated seasonally by the TCH method. 151 The weekly datasets are grouped according to the weekly climatology and then used to estimate the 152 uncertainty. For example, all the precipitation datasets which belongs to the first week of each year are concatenated to apply the TCH method in order to get the uncertainty of the datasets on the first week of 153 154 each year. Similarly, the data uncertainty on the second week, third week and until the fifty-two week is 155 evaluated sequentially. This strategy enables a time-variant uncertainty estimation, which is more 156 reasonable as the precipitation climatology is different for different seasons. The NECP R2 and ERA-5 157 data are used to assist the uncertainty estimation of MERRA-2 data by the TCH method, and the 158 precipitation forecasting experiments are conducted based on MERRA-2 data to evaluate the proposed 159 forecasting framework.





# 160 **3.2. Variational Bayesian inference**

161	Here we introduce the variational inference theory (Hoffman et al., 2013), which is a standard
162	Bayesian modeling technique for the estimation of model uncertainty. Given the input data $X = \{x_1, \dots, x_N\}$
163	and the output data $Y = \{y_1, \dots, y_N\}$ , the Bayesian regression is to find suitable parameters within the
164	function $y=f^{w}(x)$ which could generate the output Y according to the input X. The parameters w is assumed
165	to obey a prior distribution $p(w)$ before the observations are known. When the observed data is obtained,
166	it is possible to determine which parameters are more suitable for the function according to the data. A
167	likelihood distribution $p(y x,w)$ is defined to describe the probability of y generated by x and w. For
168	example, a Gaussian likelihood function is defined as
1.00	
169	$p(y x,w) = \mathcal{N}(y; f^w(x), \tau^{-1}I)$
170	
171	
172	
173	(1)
174	where $\tau^{-1}$ is the observation noise.
175	Given the input data $X$ and the output data $Y$ , the Bayesian theorem is to find the posterior
176	distribution of parameters in the parameter space.
177	$p(w X,Y) = \frac{p(Y x,w)p(w)}{p(Y X)}$
	p(Y X)
178	
179	
180	
181	(2)
182	where the numerator $p(Y X)$ is the normalization factor, also named as model evidence.
183	$p(Y X) = \int p(Y X, w) p(w) dw$
184	

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(=
The solution of Equation (3) needs to marginalize the likelihood over w, which is tractable
analytically for some simple models such as Bayesian regression, while is intractable for complex mode
such as deep learning methods (Gal, 2016).
Given the new input data $x'$ , the forecasted value is generated by the integral of probability over the
parameter space, which is called the inference process.
$p(y' x',X,Y) = \int p(y' x',w)p(w X,Y)dw$
(4
Since the posterior distribution of parameters $p(w X,Y)$ cannot be obtained analytically, a
approximate analytical distribution $q_{\theta}(w)$ could be defined, with $\theta$ as the parameter to be estimated, t
be as close as the posterior distribution. The Kullback-Leibler (K-L) divergence (Kullback and Leible
1951) is an indicator to measure the similarity of two distributions, also known as relative entropy. The
objective function is to minimize the K-L divergence between the two distributions.
$KL(q_{\theta}(w)  p(w X,Y)) = \int q_{\theta}(w) \log \frac{q_{\theta}(w)}{p(w X,Y)} dw$
(5
The optimal variational distribution $q'_{\theta}(w)$ is obtained when the K-L divergence is minimized
The estiamted variational distribution could be regarded as the posterior distribution of parameters an
then the predictive distribution could be generated.





211	$p(y' x',X,Y) \approx \int p(y' x',w)q'_{\theta}(w)dw =: q'_{\theta}(y' x')$	
212		
213	(6)	
214	The above inference process is the variational Bayesian inference. Variational inference replaces	
215	the integral of the likelihood with optimization, which simplifies the estimation of posterior distribution.	
216	3.3. Monte Carlo sampling	
217	Monte Carlo method is a kind of stochastic simulation technology, proposed by Stanislaw Ulam	
218	and John von Neumann during the second world war (Von Neumann and Ulam, 1951). Monte Carlo	
219	methods are used to estimate unknown parameters by random sampling and are widely applied in	
220	mathematics, physics, game theory and finance (Brooks, 1998; Jacoboni and Lugli, 2012; Metropolis	
221	and Ulam, 1949; Rubinstein and Kroese, 2016).	
222	In Equation (4), the posterior distribution $p(w X,Y)$ cannot be solved analytically. Assume U <sub>i</sub> as the	
223	weight matrix $K_i \times K_{i-1}$ from <i>i</i> -1 layer to <i>i</i> layer, i.e. $w = \{U_i\}_{i=1,,L}$ , a variational weight distribution $q(w)$	
224	is defined to randomly replace the columns with zero (dropout process).	
225	$U_{i} = H_{i} \cdot diag\left(\left[z_{i,j}\right]_{j=1}^{K_{i}}\right)$	
226		
227	(7)	
228	$[z_{i,j}] \sim Bernoulli(p_i), i = 1,, L, j = 1,, K_{i-1}$	
229		
230	(8)	
231	where $p_i$ and $H_i$ are variational parameters; $z_{i,j}$ is a binary variable, with a value of zero representing the	
232	abandoning of <i>j</i> th unit in <i>i</i> -1 layer and a value of one the keeping, based on the <i>Bernoulli</i> distribution at	
233	the probability $p_i$ .	
234	The predictive distribution is estimated after minimizing the K-L divergence.	
235	$q(y' x') = \int p(y' x', w)q(w)dw$	





235
 (9)

 236
 The predictive mean and variance can be obtained after repeating the dropout process multiple times.

 239
 
$$\mathbb{F}_{q(y'|x')}(y') = \int y' q(y'|x') dy' =$$

 240
  $\int y' N(y', y', y', x', u_1, ..., u_1), \pi^{-1}) Bern(U_1) \cdots Bern(U_1) dU_1 \cdots dU_1 dy' =$ 

 241
  $\int y' (x', u_1, ..., u_1) Bern(U_1) \cdots Bern(U_1) dU_1 \cdots dU_L = \frac{1}{T} \sum_{i=1}^{T} y'(x', 0_{1,i}, ..., 0_{i,i})$ 

 242
 (10)

 243
 (10)

 244
  $Var_q(y'|x')(y') \approx \tau^{-1}l + \frac{1}{T} \sum_{i=1}^{T} y'(x', 0_{1,i}, ..., 0_{i,i})^T y'(x', 0_{1,i}, ..., 0_{i,i}) - (1)

 245
  $\mathbb{E}_{q(y'|x')}(y') \approx \tau^{-1}l + \frac{1}{T} \sum_{i=1}^{T} y'(x', 0_{1,i}, ..., 0_{i,i})^T y'(x', 0_{1,i}, ..., 0_{i,i}) - (1)

 246
  $Var_q(y'|x')(y') \approx \tau^{-1}l + \frac{1}{T} \sum_{i=1}^{T} y'(x', 0_{1,i}, ..., 0_{i,i})^T y'(x', 0_{1,i}, ..., 0_{i,i}) - (1)

 247
 where  $u_{i,i}$  is the forecasted value for *i*th pixel and *t*th ensemble. The calculation of predictive variance

 248
 is based on the standard deviation of the ensemble, which represents the spread of the forecasted values.

 249
 Dropout is a Bayesian method to model the model uncertainty in forecasting (Srivastava et al.,

 250
 Dropout is a Bayesian method to model the ensemble. Kendall and Gal (2017) regarded the

 251
 Dropout is a Bayesian method to model the learning of data uncertainty increases the number of

 25$$$ 





261	(12)
262	$\sigma_x \sim \mathcal{N}(0, \sigma)$
263	
264	(13)
265	$\sigma_{\gamma} \sim \mathcal{N}(0, \sigma)$
266	
267	(14)
268	The data uncertainty is randomly sampled $T$ times to generate an ensemble of predictors and
269	predictands. In the meantime, the parameters are randomly dropped out for $T$ times to construct a
270	parametric ensemble. The perturbed data and parameter values are jointly used to calculate the training
271	loss. The objective function is expressed as follows (Kendall and Gal, 2017), which is obtained from the
272	likelihood of a Gaussian process (Srivastava et al., 2014).
273	$\mathcal{L}(\theta, p) = -\frac{1}{N} \sum_{i=1}^{N} \log p\left(y_{i,\sigma} \middle  f^{\mathcal{Q}_i}(x_{i,\sigma})\right) + \frac{1-p}{2N} \ \theta\ ^2$
274	
275	(15)
276	$\sigma = \sqrt{(\sigma_x^{(l)})^2 + \sigma_y^2}$
277	
278	(16)
279	where N is the sample size; p is the dropout probability; $\hat{U}_{l} \sim q'_{\theta}(U)$ ; $\theta$ is the parameter to be estimated;
280	$\sigma_x$ and $\sigma_y$ are the data uncertainty for predictor and predictand, respectively. The negative log-likelihood
281	function can be deduced according to the objective function.

282 
$$-\log p\left(y_{i,\sigma} \left| f^{\vartheta_i}(x_{i,\sigma}) \right) \propto \frac{1}{2\sigma^2} \left\| y_i - f^{\vartheta_i}(x_i) \right\|^2 + \frac{1}{2} \log \left( \sigma^2 \right)$$
283
284 (17)





285	where $\sigma$ is the regression noise, with the mean of zero in a Gaussian distribution.
286	The objective function consists of a mean square error (MSE) term adjusted by data uncertainty and
287	a regularization term, which is the negative logarithm of the Gaussian likelihood function. The objective
288	function includes an uncertainty parameter $\sigma^2$ , which is determined by the sum of propagated uncertainty
289	and target data uncertainty. The minimization of the negative log-likelihood function could be reached
290	by differentiating the optimization function and setting to zero.
291	$\frac{\partial}{\partial \sigma^2} \left[ \frac{1}{2\sigma^2} \left\  y_i - f^{\theta_i}(x_i) \right\ ^2 + \frac{1}{2} \log \left( \sigma^2 \right) \right] = 0$
292	
293	(18)
294	$\Rightarrow -\frac{1}{2\sigma^4} \left\  y_i - f^{\mathcal{O}_i}(x_i) \right\ ^2 + \frac{1}{2\sigma^2} = 0$
295	
296	(19)
	- H - G - H <sup>2</sup>
297	$\Rightarrow \sigma^2 = \left\  y_i - f^{\hat{U}_i}(x_i) \right\ ^2$
298	
299	(20)
300	where the minimum value of the negative log-likelihood function could be reached when the data
301	variance equals to the square of the difference between the forecasted value and the observation.
302	Once the network weights are determined according to the objective function, the new input data
303	uncertainty is propagated and the weights are randomly sampled to produce the forecasted ensemble.
304	The predictive mean and variance are calculated from the predictive ensembles.
305	$\mu_i = rac{1}{T} \sum_{t=1}^T \mathbf{y}_{t,i}$
306	
307	(21)

308  $Var_i \approx \frac{1}{T} \sum_{t=1}^T y_{t,i}^2 - \left(\frac{1}{T} \sum_{t=1}^T y_{t,i}\right)^2 + \frac{1}{T} \sum_{t=1}^T \sigma_{t,i}^2$ 





(22)

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310



311

Figure 2: The proposed integrated data-model uncertainty modeling framework in precipitationforecasting.

#### 314 **3.5. The deep learning forecasting framework**

315 In deep learning, encoder-decoder is a commonly used forecasting model (Badrinarayanan et al., 316 2017; Cho et al., 2014). In the encoder process, an input signal is converted into a one-dimension vector 317 with fixed length. In the decoder process, the one-dimension vector is transformed into the target data 318 with variable length. The available networks used for encoder and decoder processes are arbitrary and 319 depend on the specific problem, such as convolutional neural network (CNN), recurrent neural network 320 (RNN) and long-short term memory (LSTM) network (Hochreiter and Schmidhuber, 1997; Goodfellow 321 et al., 2016). Here we designed a deep learning encoder-decoder model for weekly precipitation 322 forecasting (Figure 3). The temperature and geopotential height data for previous three weeks are 323 regarded as inputs, with an image size of 64×64×6. In the encoder process, the input image is down-324 sampled by a series of convolution, pooling and dropout operations, resulting in a one-dimension vector 325  $(1 \times 1 \times 2048)$ . In the decoder process, the one-dimension vector is up-sampled by deconvolution, dropout 326 and convolution operations, resulting in a forecasted precipitation image (64×64×1). The down-sampling 327 and up-sampling procedures are used to learn the nonlinear mapping relationships between predictors 328 and predictands.

337





329 In the training process, the optimization algorithm is set to Adam (Kingma and Ba, 2014), which is 330 a stochastic learning algorithm based on adaptive moment. The network learning rate is set to 0.001 and 331 the stopping rule of iteration is that the validating error does not decrease for at least 100 times. The data 332 uncertainty is propagated forward according to the law of uncertainty propagation and the dropout 333 process is repeated 10 times with a dropout rate of 0.5. The random seed is set to 1 to enable the 334 reproducibility of the experiment. The experimental data spans from 1980 to 2020 (2139 weeks), of 335 which 60%, 20% and 20% of the data are used for training, validating and testing, respectively. The 336 optimal model parameters are determined based on the minimal validating loss.



338 Figure 3: The developed deep learning model for precipitation forecasting. 339 The proposed deep learning framework for precipitation forecasting is demonstrated in Figure 4. 340 The input data and its uncertainty are prepared and are considered as inputs for the forecasting model. 341 The model weights are initialized and the input uncertainty is propagated forward according to the 342 weights. The loss function value is calculated according to the forecasted value, the propagated uncertainty, target data and its uncertainty. The forecasting model is trained according to the optimization 343 344 algorithm and then the trained model is obtained. Next the test data is used to produce the forecasted 345 value and variance based on the model weights and uncertainty propagation. Finally, the forecasted value 346 and variance are evaluated and compared with several precipitation forecasting methods.







347 348

Figure 4: The proposed deep learning framework for precipitation forecasting.

349 We designed a series of comparison experiments to investigate the effect of different uncertainty 350 processing strategies on the forecasting performance. The precipitation forecasting experiment without 351 considering uncertainty is used as the baseline (Experiment 1). The mean square error is used as the loss 352 function and the data and model uncertainties are not considered in Experiment 1. The uncertainty 353 sources are incorporated differently into the experiments, including predictor uncertainty (Experiment 354 2), predictor and predictand uncertainties (Experiment 3), model uncertainty based on Srivastava et al. (2014)'s method (Experiment 4), data and model uncertainties based on Kendall and Gal (2017)'s method 355 356 (Experiment 5), data and model uncertainties based on Loquercio et al. (2020)'s method (Experiment 6), 357 and data and model uncertainties (Experiment 7) based on the proposed framework here. The data 358 uncertainty only includes the propagated uncertainty from the input data in Equation (16) in Experiment 359 2, while the propagated uncertainty from the input data and the target data uncertainty are both included 360 in Experiment 3. In Experiment 4, the data uncertainty is ignored and the model parameters are randomly 361 sampled for 10 times to get the model spread. In Experiment 5, the input uncertainty is regarded as the 362 trainable parameter and the trained uncertainty value is used as the data uncertainty in Equation (16), and 363 the model uncertainty is considered the same way as Experiment 4. In Experiment 6, the input data 364 uncertainty is propagated and the model uncertainty is modeled by sampling the parameters. In

365





Experiment 7, the input uncertainty is propagated and the target data uncertainty is included in Equation

# 373 4 Results and discussion

# 374 **4.1. The uncertainty of input and output datasets**

375 The data uncertainty of predictors and predictand is calculated based on the TCH method and is

376 shown in Figure 5. The precipitation data uncertainty is much higher than







- 378 Figure 5: The data uncertainty calculated by the TCH method. The uncertainty distribution is plotted
- 379 according to the uncertainty over all the pixels of the study area. The red cross indicates the outliers.
- 380 the temperature and geopotential height variables, with a median of  $\sim$ 43% relative uncertainty fraction 381 for precipitation,  $\sim 2\%$  for temperature and  $\sim 1\%$  for geopotential height. Therefore, the precipitation data 382 suffer from greater uncertainty relative to the input data and the predictand uncertainty should not be 383 ignored in the training process. The combination of the propagated input uncertainty and the predictand 384 uncertainty is used as the adjusted parameter to regularize the loss function, which is relatively reasonable 385 as the data with larger uncertainty should contribute less to the total training loss. It should be noted that 386 the predictor and predictand data are normalized to [0,1] before uncertainty estimation to ensure a fair 387 comparison of uncertainty value. The high uncertainty for precipitation data is related to strong 388 spatiotemporal heterogeneity of precipitation and the high inconsistency among the reanalysis data (Xu 389 et al., 2020b), while temperature and geopotential height data are much more homogeneous in space and 390 time.

# 391 **4.2. Overall precipitation forecasting performance**

392 As for the predictive uncertainty, the forecasting method that only considers model uncertainty 393 (Srivastava et al., 2014) obtains the minimum predictive uncertainty (Table 1). However, the data 394 uncertainty is not considered when only sampling from the parameters and thus the impact of data error 395 on forecasting is not evaluated. In Kendall and Gal (2017)'s method, the data uncertainty is regarded as 396 a trainable parameter and the model uncertainty is modeled by random parameter sampling. Whether the 397 learned parameter value for data uncertainty parameter could represent the real data error needs further 398 investigation. Loquercio et al. (2020) used the law of uncertainty propagation to propagate the data 399 uncertainty and sampled the parameters randomly during training. In our proposed method, the input 400 data uncertainty, target data uncertainty and model uncertainty are jointly coupled by uncertainty 401 propagation and random parameter sampling. The average predictive uncertainty (3.460) based on the 402 proposed method is smaller than the Loquercio et al. (2020)'s and Srivastava et al. (2014)'s methods. In 403 this regard, the proposed method could reduce the predictive uncertainty of precipitation forecasting to 404 some extent, when jointly modeling data and model uncertainties. The proposed method could slightly 405 improve the precipitation forecasting performance and could improve the reliability of precipitation 406 forecasting by reducing the uncertainty.

407 When only the input uncertainty is modeled in the forecasting model, the predictive uncertainty is 408 12.950. If the target data uncertainty is coupled with input uncertainty, the predictive uncertainty is 409 substantially reduced (3.022). In Equation (20), when the predictive error on the right side of the equation 410 reaches local minimum and remains unchanged basically, the left side of the equation includes the input





- 411 uncertainty propagation and the target data uncertainty. When new data is used to make prediction, the 412 predictive uncertainty is generated by the input uncertainty and the law of error propagation. Thus, when 413 only the input uncertainty is modeled in Equation (20), the left side of this Equation equals to the 414 propagated uncertainty from the input data. If the left side of the Equation (20) is replaced from the 415 propagated input uncertainty with the combination of propagated input uncertainty and target uncertainty, 416 the propagated input uncertainty after replacement will be smaller than that of no replacement, i.e. 417  $(\sigma_x^{(l)})^2 = \sigma^2 - \sigma_y^2 < \sigma^2 = ||y_i - f^{\theta_i}(x_i)||^2$ .
- 418 Table 1. The accuracy of precipitation forecasting based on different uncertainty processing strategies.

Uncertainty processing	RMSE	Uncertainty
No uncertainty	25.357	-
Predictor uncertainty	25.500	12.950
Predictor and predictand uncertainties	25.932	3.022
Model uncertainty (Srivastava et al., 2014)	25.368	1.778
Data and model uncertainties (Kendall and Gal, 2017)	25.225	4.914
Data and model uncertainties (Loquercio et al., 2020)	25.324	13.169
Data and model uncertainties (This study)	25.199	3.460

419 The best RMSE is shown in bold for each column.

# 420 **4.3. Spatial patterns of precipitation forecasting**

421	In Figure 6, the spatial patterns of the RMSE for precipitation forecasting demonstrate some
422	similarities and differences between different uncertainty processing strategies. Overall, the spatial
423	distribution of RMSE is similar with each other and is smaller in the northwest region but larger in the
424	southeast region. In the places where the annual rainfall is abundant, the water cycle process is
425	accelerated and the precipitation observations may suffer from large uncertainty. The difficulty of
426	forecasting extreme high precipitation volume also increases the average RMSE in the southeast region
427	relative to the northwest region (Yuan et al., 2012; Huang et al., 2013). There are some differences of the
428	forecasting error among different forecasting methods in local areas. For example, the forecasting
429	performance based on our proposed method could outperform the methods in Experiments 1, 2 and 4
430	and is comparable with the methods in Experiments 3, 5 and 6 for the local areas covered by black circles
431	in Figure 6.







432

Figure 6: The spatial patterns of RMSE for precipitation forecasting. In this figure, x means the modeling
of input uncertainty; x+y represents the modeling of input and output uncertainty; x+y+0 indicates the
modeling of input uncertainty, output uncertainty and model uncertainty. The black circles represent the
highlighted areas.

437

438 Figure 7 demonstrates the impact of different uncertainty processing methods on the predictive 439 uncertainty of precipitation. If only the input uncertainty is considered in the forecasting model, the 440 predictive uncertainty is large in the central and southwest regions. The predictive uncertainty could be 441 substantially reduced when incorporating the target data uncertainty besides the input uncertainty. The 442 modeling of model uncertainty only could produce the minimum predictive uncertainty spatially in the 443 experiment. The predictive uncertainty is relatively small based on Kendall and Gal (2017)'s method. In 444 Loquercio et al. (2020)'s method, the predictive uncertainty is close to the result of input uncertainty 445 modeling in space, suggesting that the coupled modeling of input and model uncertainties fails to help 446 reduce the forecasting spread. Our proposed method could include the input, target and model 447 uncertainties jointly and could help reduce the predictive uncertainty to a large extent, relative to the 448 methods in Experiment 2 and 6.

19







450 Figure 7: The spatial patterns of uncertainty for precipitation forecasting.

# 451 4.4. Uncertainty analysis and discussion

449

452 In precipitation forecasting, data and model uncertainties both bring uncertainty to the forecasting 453 result. The higher the data and model uncertainties, the more divergent the forecasting, suggesting the 454 forecasting less reliable. Therefore, the data and model uncertainties should be jointly considered in the 455 forecasting process (Gal, 2016; Kendall and Gal, 2017; Loquercio et al., 2020; Parrish et al., 2012). 456 Although the predictive error is close to each other among different forecasting methods in Figure 6 and 457 Table 1, the predictive uncertainty has some discrepancies. The modeling of input uncertainty only in the 458 forecasting model would bring high predictive uncertainty and ignore the target data uncertainty. The 459 joint modeling of input and target uncertainties could reduce the predictive uncertainty substantially, 460 which is related to the change of the variance in Equations (16) and (20) corresponding to the minimum 461 value of the forecasting error term. The propagation of input uncertainty is constrained by refining the 462 uncertainty representation in Equation (16) after incorporating the target uncertainty term and thus 463 changing the weight training process.

464 The proposed method in this study could model the input uncertainty, target uncertainty and model





465 uncertainty jointly and could reduce the predictive uncertainty relative to Kendall and Gal (2017)'s and 466 Loquercio et al. (2020)'s methods. The developed method does not increase the training parameter and 467 is a general forecasting uncertainty method for geophysical applications such as temperature forecasting, 468 runoff forecasting and wind speed forecasting, especially for data-driven forecasting models (Ham et al., 469 2019; Zheng et al., 2020; Hossain et al., 2015). 470 In numerical precipitation forecasting systems, ensemble forecasting is commonly used to quantify the predictive uncertainty (Duan et al., 2019). In ensemble forecasting, the model parameters and data 471 472 are perturbed to produce a forecasted ensemble and thus the data and model uncertainties are both 473 considered. However, it would be time-consuming and cost-expensive to conduct large-sample sampling 474 for complex physical models. In our developed method, the law of error propagation is used to propagate 475 the data uncertainty. The uncertainty propagation of convolution, max-pooling and deconvolution in the 476 deep learning forecasting model is tractable in an analytical form. However, the uncertainty propagation 477 process is generally intractable analytically for complex statistical or physical models. Therefore, the 478 theory and implementation technology for uncertainty modeling require further development, such as 479 surrogate modeling, Monte Carlo methods, polynomial chaos expansions and Bayesian approaches 480 (Linde et al., 2017; Sudret et al., 2017; Zhu and Zabaras, 2018; Schiavazzi et al., 2017; Nitzler et al., 481 2020).

# 482 5 Conclusion

483 In this study, we proposed a data-model uncertainty coupling framework to estimate the predictive uncertainty of precipitation forecasting. In this framework, the predictor and predictand uncertainties are 484 485 estimated a prior by the TCH method and are assumed as Gaussian distribution. The predictor uncertainty 486 is propagated forward during training and testing processes by the law of error propagation. The model 487 uncertainty is represented by randomly abandoning model weights from deep learning layers. The data and model uncertainties are jointly modeled in the objective function during training and are also used 488 489 during the testing process. The loss function is constructed by the MSE statistic adjusted by data 490 uncertainty and a regularization term based on logarithmic data uncertainty. In the loss function, the 491 adjusting parameter is determined by the combination of the square of predictor and predictand 492 uncertainties. The forecasted ensembles are used to calculate the predictive mean and variance to estimate





493 the predictive uncertainty of precipitation.

494 The weekly precipitation forecasting in southern and northern China is used as an example to 495 examine the effectiveness of the proposed joint uncertainty modeling framework. Temperature and 496 geopotential height data in previous three weeks are used to forecast the precipitation in the target week. 497 The forecasting model is developed based on an encoder-decoder deep neural network, with multivariate 498 spatiotemporal predictor data as inputs and spatiotemporal precipitation data as output. The results 499 exhibit comparable precipitation forecasting accuracy for the proposed method with several existing 500 uncertainty processing strategies, while the predictive uncertainty is reduced relative to two data-model 501 uncertainty modeling methods. The reduction of predictive uncertainty is significative for quantitative 502 precipitation forecasting from a data-driven view.

503 The data-driven precipitation forecasting method has limitations in the interpretation part relative 504 to numerical weather prediction. The precipitation forecasting accuracy for numerical models could still 505 be improved by improving the parameterization schemes and resolving the uncertainties in observations, 506 parameters and models. The proposed uncertainty modeling framework may also provide some insights 507 for the uncertainty quantification in numerical prediction models. For example, the uncertainty 508 propagation for the input data and the coupling with target data uncertainty could be used in a data 509 assimilation scheme to estimate the propagated uncertainty in weather forecasting.

510 Data-driven precipitation forecasting could be used as a tool to assist regional prediction and 511 warning of extreme weather events together with numerical models. The proposed joint data-model 512 uncertainty modeling framework could help estimate the forecasting spread and is a general approach to 513 derive predictive uncertainty for geophysical forecasting applications. Further research should focus on 514 the non-Gaussian uncertainty modeling for complex integrated statistical-physical models.

#### 515 Data availability

The meteorological data are publicly available and can be obtained via the website <u>https://gma</u>
 <u>o.gsfc.nasa.gov/reanalysis/MERRA-2/</u> for MERRA-2, <u>https://psl.noaa.gov/data/gridded/data.ncep.rea</u>
 <u>nalysis2.html</u> for NCEP R2 and <u>https://climate.copernicus.eu/climate-reanalysis</u> for ERA-5.

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#### 519 Author contributions

- 520 LX conceptualized and wrote the paper. CY provided supervision of this study. NC gave suppo
- 521 rt in developing the manuscript.

# 522 Competing interests

523 The authors declare that they have no conflict of interest.

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