

1 **Performance of the Global Forecast System’s Medium-Range**
2 **Precipitation Forecasts in the Niger River Basin Using Multiple**
3 **Satellite-Based Products**

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14 **Abstract.** Accurate weather forecast information has the potential to improve water resources management, energy,
15 and agriculture. This study evaluates the accuracy of medium-range (1 – 15 day) precipitation forecasts from the
16 Global Forecast System (GFS) over watersheds of eight major dams (Selingue Dam, Markala Dam, Goronyo Dam,
17 Bakolori Dam, Kainji Dam, Jebba Dam, Dadin Kowa Dam, and Lagdo Dam) in the Niger river basin using NASA’s
18 Integrated Multi-satellitE Retrievals (IMERG) “Final Run” satellite-gauge merged rainfall observations. The results
19 indicate that the accuracy of GFS forecast varies depending on climatic regime, lead time, accumulation timescale,
20 and spatial scale. The GFS forecast has large overestimation bias in the Guinea region of the basin (wet climatic
21 regime), moderate overestimation bias in the Savannah region (moderately wet climatic regime), but has no bias in
22 the Sahel region (dry climate). Averaging the forecasts at coarser spatial scales leads to increased forecast accuracy.
23 For daily rainfall forecasts, the performance of GFS is very low for almost all watersheds except for Markala and
24 Kainji dams, both of which have much larger watershed areas compared to the other watersheds. Averaging the
25 forecasts at longer time scales also leads to increased forecast accuracy. The GFS forecasts, at 15-day accumulation
26 timescale, have better performance, but tend to overestimate high rain rates. Additionally, the performance assessment
27 of two other satellite products was conducted using IMERG Final estimates as reference. The Climate Hazards Group
28 InfraRed Precipitation with Station data (CHIRPS) satellite-gauge merged product has similar rainfall characteristics
29 with IMERG Final, indicating the robustness of IMERG Final. The IMERG “Early Run” satellite-only rainfall product
30 is biased in the dry Sahel region, however in the wet Guinea and Savannah regions, IMERG “Early Run” outperforms
31 GFS in terms of bias.

32

33

34 1. Introduction

35 Global climate forecasts, with lead times ranging from hours to several months, are becoming increasingly available
36 (Saha et al. 2014; Abdalla et al. 2013; NCEP 2015; JMA 2019). Significant societal benefit could be realized from
37 research to reduce common barriers in climate forecast utilization blocking the path to improving water resources
38 management, energy, and agriculture. One such a barrier is the lack of understanding of climate forecast accuracy in
39 different regions of the world. This focus is timely given the recent advances in numerical atmospheric models, and
40 in the wealth of new observing capabilities including satellite remote sensing. These combined models and
41 observational datasets provide opportunity for researchers to quantify the accuracy of climate forecasts.

42
43 The Niger River is the principal river of West Africa, and is shared among nine riparian countries (Fig. 1): Benin,
44 Burkina Faso, Cameroon, Chad, Guinea, Ivory Coast, Mali, Niger and Nigeria. The basin is facing multiple pressures
45 from increasing population, water abstraction for irrigation, and risk of extreme hydrological events due to climate
46 change (Sylla et al. 2018). A number of hydropower dams exist in the region, and additional dam projects are
47 envisaged in order to alleviate chronic power shortages in the countries of the Niger basin. Optimal management of
48 water resources is key to maximizing benefits, such as hydropower generation, and minimize disasters, such as
49 flooding. Climate forecast information has the potential to improve water resources management, energy, and
50 agriculture (e.g., Patt et al. 2007; Breuer et al. 2010; Mase and Prokopy 2014; Pandya et al. 2015; Koppa et al. 2019;
51 Alexander et al. 2020). For example, in a recent study, Koppa et al (2019) showed that the use of seasonal precipitation
52 forecasts in reservoir planning of Omo Gibe dam in Ethiopia can increase annual hydropower generation by around
53 40%.

54
55 Several studies have investigated the accuracy of seasonal forecasts in West Africa (e.g., Bliefert et al. 2019;
56 Pirret et al. 2020). Seasonal forecasts are important for water resource planning, while medium-range (1-day to 15-
57 day) forecasts are important for operational decisions, such as reservoir operations. The availability of medium-range
58 global climate forecasts has grown in recent years. Examples of such forecast products include Global Forecast System
59 (GFS; NCEP 2015), NCEP climate forecast system (NSF CFS, Saha et al. 2014), European Centre for Medium-Range
60 Weather Forecasts (ECMWF; Abdalla et al. 2013), and Global Spectral Model (GSM; JMA 2019). For these
61 precipitation forecasts to be effectively used in applications, their accuracy must be known, which is usually performed

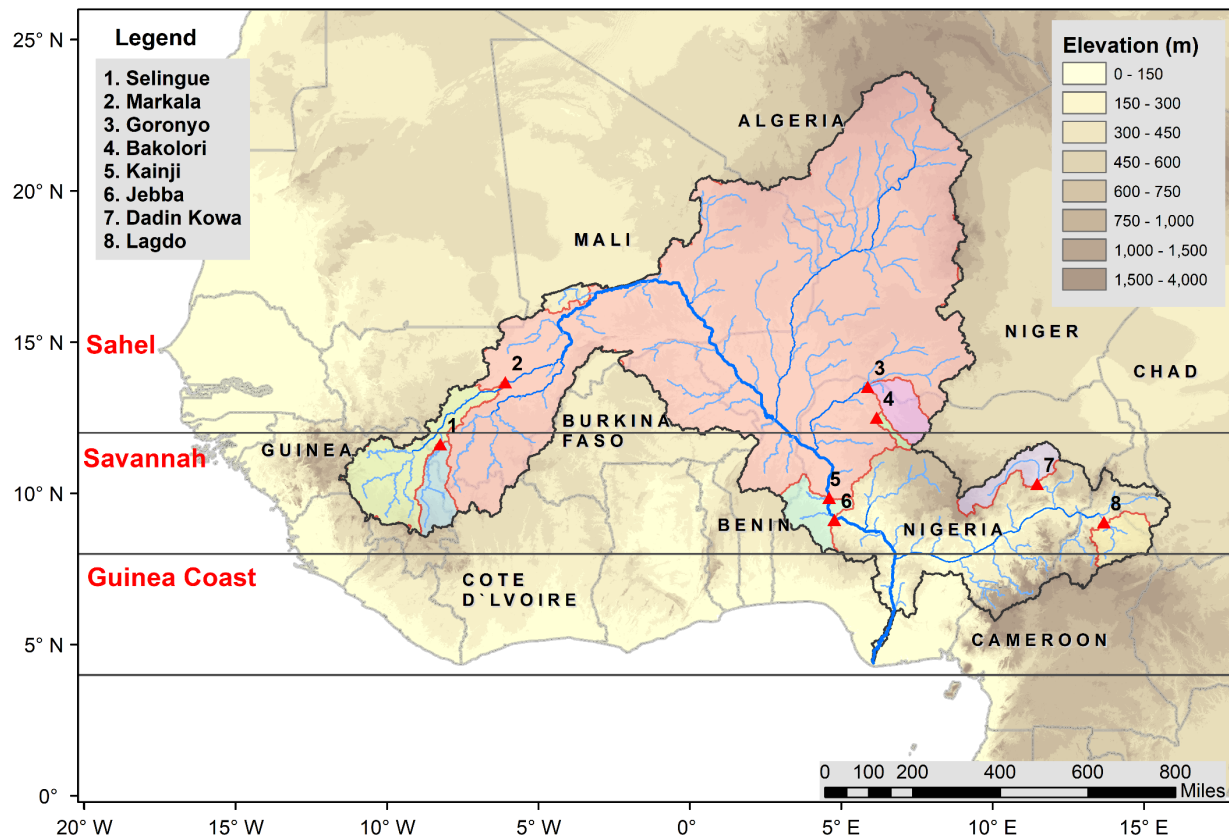


Figure 1. The Niger River Basin, and locations of major reservoir dams in the basin: (1) Selingue, (2) Markala, (3) Goronyo, (4) Bakolori, (5) Kainji, (6) Jebba, (7) Dadin Kowa, and (8) Lagdo.

62 through comparison of precipitation forecasts to observations (e.g., Tian et al. 2017; Yuan et al. 2014). Wang et al.
 63 (2019) performed numerical experiment to examine the sensitivity of GFS to inclusion or exclusion of additional
 64 observations collected over the eastern Pacific during the El Niño Rapid Response (ENRR) field campaign, type of
 65 data assimilation method to prepare the initial conditions, and inclusion or exclusion of stochastic parameterizations
 66 in the forecast model. They reported that the GFS forecast errors are only slightly sensitive to the additional ENRR
 67 observations, more sensitive to the DA methods, and most sensitive to the inclusion of stochastic parameterizations in
 68 the model. In addition, they reported that GFS forecasts have difficulty to capture the location and magnitude of heavy
 69 rain rates. Sridevi et al. (2018) evaluated the performance of GFS in India by using rain gauge and satellite rainfall
 70 product, and reported that the GFS forecast shows some skills in 1-day and 2-day lead times, but low skills from 3-
 71 day onwards. Lien et al. (2016) compared the statistical properties of GFS forecasts and Tropical Rainfall Measuring
 72 Mission (TRMM) Multisatellite Precipitation Analysis (TMPA; Huffman et al. 2007, 2010) observations. They
 73 reported that the GFS model has positive bias in precipitation amount compared to TMPA observations, and that the

74 GFS forecasts have large random errors at higher resolutions, especially for convective precipitation. According to
75 Jiang et al. (2015) the lack of consideration of the Aerosol-Cloud Interactions (ACIs) in the GFS model leads to
76 significant bias in the GFS precipitation forecasts.

77
78 In our study region of the Niger River basin, there has not been any performance evaluation of GFS precipitation
79 forecasts to date. The Niger basin lies in three different climate regimes (wet regime, moderately wet regime, and dry
80 regime), and is home to nine major irrigation and hydropower dams (Selingue, Markala, Goronyo, Bakolori, Kainji,
81 Jebba, Dadin Kowa, and Lagdo). Recent advances in satellite rainfall products, particularly following the Global
82 Satellite Measurement satellite mission (GPM; Hou et al. 2014), and extensive evaluation of GPM rainfall products
83 in West Africa, provides us with opportunity to use GPM rainfall products as reference for evaluation. Many studies
84 have conducted to evaluate the accuracy of the satellite rainfall estimates in West Africa. Dezfuli et al. (2017a)
85 evaluated the performance of NASA's Integrated Multi-satellitE Retrievals (IMERG) "Final Run" (IMERG Final)
86 (version 4; Huffman et al. 2019a, b) in comparison with two, high-resolution, experimental rain gauge station data
87 provided by the Trans-African Hydro-Meteorological Observatory (TAHMO; van de Giesen et al. 2014), and reported
88 the capability of IMERG Final to represent well the diurnal cycle of rainfall. Using the same dataset, Dezfuli et al.
89 (2017b) showed that IMERG Final is able to capture the propagation of large Mesoscale Convective Systems (MCSs),
90 a significant advantage over its predecessor's (TMPA) 3-hourly temporal resolution, which misses the time evolution
91 of most of these systems. Gossett et al. (2018) evaluated the performance of a number of satellite rainfall products
92 (focusing only on versions that do not include rain gauge data) by comparison with rain gauge station networks in
93 Benin and Niger, and reported that the satellite products (especially IMERG Early) exhibit high performance in Niger
94 but relatively lower performance in Benin. Satgé et al. (2020) evaluated the accuracy of a number of gridded
95 precipitation datasets over West Africa through comparison against rain gauge station data, and reported that CHIRPS
96 and TMPA (the predecessor to IMERG) provided reliable estimates at both daily and monthly timescales, while the
97 remaining satellite products considered (CMORPH, PERSIANN, GSMaP, ARC, and TAMSAT) and all atmospheric
98 reanalysis products considered (MERRA and JRA) were deemed unreliable. Furthermore, they found out that satellite
99 products that incorporated rain gauge information outperformed satellite-only products. Maranan et al. (2020)
100 compared IMERG Final products against experimental rain gauge station data in the moist forest region of Ghana,

101 West Africa, and showed that IMERG Final datasets are able to capture monthly rainfall with a correlation coefficient
102 close to unity.

103
104 The objective of this study is to evaluate the accuracy of medium-range precipitation forecasts derived from the Global
105 Forecast System (GFS) for the major reservoir dams of the Niger basin through comparison against IMERG Final.
106 We chose GFS model due to its relatively high spatial ($0.25^\circ \times 0.25^\circ$) and temporal resolution (3-hourly to 6-hourly)
107 as well as free-of-charge data availability to users. The main questions addressed in this study are as follows. First,
108 how does the accuracy of GFS forecast vary across different reservoir dams in the same basin? Second, how does the
109 accuracy vary with lead time in the range 1- to 15-day? Third, what is the effect of spatial averaging (from 0.25° all
110 the way to basin-scale) and temporal aggregation (from 1-day to 15-day) on the forecast accuracy? Fourth, how does
111 the accuracy of GFS forecast compare with the accuracy of satellite-only rainfall products that are available in near-
112 real-time, as the latter may have the potential to calibrate and improve the accuracy of GFS?

113

114 **2. Data and Methodology**

115 **2.1 Global Forecast System (GFS) Medium-Range Precipitation Forecasts**

116

117 The Global Forecast System (GFS) is a global numerical weather prediction system run by the U.S. National Weather
118 Service (NWS). The GFS forecast products with a resolution of 0.25° by 0.25° are obtained from National Center for
119 Atmospheric Research (NCAR) Research Data Archive (RDA) GFS Historical Archive (NCEP 2015). The GFS is
120 run four times a day at UTC 00, UTC 06, UTC 12, and UTC 18 hours. One of the GFS model output variables is
121 accumulated precipitation, where the precipitation forecasts are accumulations starting from the model run time. We
122 obtained the 1-day lead daily rainfall forecast by subtracting the 24-hour rainfall accumulation forecast from the 48-
123 hour rainfall accumulation forecast. Similarly, in order to obtain the 5-day lead daily rainfall forecast, we subtracted
124 the 120-hour rainfall accumulation forecast from the 144-hour rainfall forecast. We only considered the model runs at
125 UTC 00 hour.

126

127 The GFS model went through a major upgrade, and its version-15 forecasts are available since June 12, 2019. In
128 version 15, the Finite Volume Cubed Sphere dynamical model (FV3) replaced the Global Spectral Model (GSM) as

129 the core model. In the GSM model, the horizontal resolutions were T1543 (12.5km) from 0 to 240 hours (0-10 days)
130 and T574 (~34km) from 240 to 384 hours (10-16 days) (NCEP 2021a). However, in the FV3 model, the horizontal
131 resolution of the model is about 13 km for days 0-16 (NCEP 2021b). The model runs are re-gridded to produce
132 precipitation forecasts at 0.25° resolution (NCEP 2015). The Key FV3 model schemes include (Putman and Lin 2007):
133 (1) the Rapid Radiative Transfer Method for GCMs (RRTMG) scheme for shortwave/longwave radiation (Mlawer et
134 al. 1997; Iacono et al. 2000; Clough et al. 2005), (2) the Hybrid eddy-diffusivity mass-flux (EDMF) scheme for
135 Planetary Boundary Layer (PBL) (NCEP, 2021a), (3) the Noah Land Surface Model (LSM) scheme for land surface
136 option (Chen et al. 1997), (4) the Simplified Arakawa-Schubert (SAS) deep convection for cumulus parameterization
137 (Arakawa et al. 1974; Grell 1993), and (5) an advanced GFDL microphysics scheme for microphysics (NCEP, 2021b).

138

139 **2.2 IMERG Final Satellite Precipitation Products**

140 IMERG Final rainfall products are used in this study as reference to evaluate the performance of GFS precipitation
141 forecasts. IMERG Final combines all available microwave precipitation estimates, microwave-calibrated infrared
142 estimates, and rain gauge data to provide rainfall estimates at very high resolution (30-minute, 0.10°) (Hou et al. 2014;
143 Huffman et al. 2015). The IMERG products are categorized into three types, namely early run, late run, and final run.
144 It is only the final run or “final” version that incorporates rain gauge data. The data latency of IMERG Final is about
145 3.5 months. Details of IMERG algorithm developed by NASA are available at Huffman et al (2019a, b). The latest
146 version (V6B) of IMERG datasets have been accessed from the NASA’s Earth Data Goddard Earth Sciences Data
147 and Information Services Center (GES DISC) web portal.

148

149 **2.3 Other Satellite Precipitation Products**

150 In order to put the GFS forecast performance into perspective, we also evaluated two other state-of-the-art satellite
151 rainfall products:

- 152 • IMERG Early provides un-calibrated IMERG rainfall fields, which do not include correction from rain gauges.
153 The data latency of IMERG Early is near-real-time, about 4 hours. We have used the latest version (V6B) of
154 IMERG Early datasets. Post-processing calibration of GFS forecasts (in order to improve the accuracy of GFS
155 forecasts) requires the use of “relatively better performing” and “available in near-real-time” independent
156 rainfall observations to correct real-time dynamical GFS model forecasts. Comparison of the performance of

157 IMERG Early with the performance of GFS would indicate to what extent the IMERG Early products could be
158 used for calibration of GFS forecasts.

159 • The Climate Hazards Group InfraRed Precipitation with Station data (CHIRPS) is derived primarily from
160 thermal infrared data using the cold cloud duration (CCD) approach, calibrated using TRMM Multi-satellite
161 Precipitation analysis (TMPA 3B42 v7; Huffman et al. 2007) precipitation datasets by local regression, and
162 include rain gauge station data from multiple sources (regional and national meteorological services). CHIRPS
163 data are available at a spatial resolution of 0.05° and a temporal resolution of 1-day, with a data latency period
164 of about 3 weeks. Details of CHIRPS algorithm are available at Funk et al. (2015). Agreement between the
165 reference (IMERG Final) and CHIRPS would indicate that the IMERG Final estimates are robust.

166

167 **2.4 Study Region**

168 The Niger river, with a drainage basin of 2,117,700 Km², is the third longest river in Africa. The source of the main
169 river is in the Guinea Highlands, and runs through Mali, Niger, on the border with Benin and then through Nigeria,
170 discharging through a massive delta, known as the Niger Delta (the world's third largest wetland), into the Atlantic
171 Ocean. The rainfall regimes in the region follow the seasonal migration of the Inter-Tropical Convergence Zone
172 (ITCZ), which brings rainfall primarily in the summer season (Animashaun et al. 2020; Sorí et al. 2017).
173 Climatologically, the Niger basin lies in three latitudinal sub-regions (Akinsanola et al. 2015, 2017): (1) the Guinea
174 coast (latitude 4°–8°N), which borders the tropical Atlantic Ocean in the south; (2) the Savannah (latitude 8°–12°N),
175 an intermediate sub-region; and (3) the Sahel (latitude > 12°N) to the north. The Guinea coast experiences a
176 bimodal rainfall regime that is centered in the summer monsoon period of June–September, with August being the
177 period of a short dry season, while the Savannah and Sahel sub-regions experience a unimodal rainfall regime, with
178 maximum rainfall occurring in August (Akinsanola and Zhou 2018). The ranges of annual rainfall amounts are: 400–
179 600 mm in the Sahel, 900–1200 mm in the Savannah; and 1500–2000 mm in the Guinea coast (Akinsanola et al.
180 2017).

181

182 The Niger basin is home to eight major reservoir dams (see Table 1 and Fig. 1): (1) Selingue Dam in Mali: a primarily
183 hydropower dam, (2) Markala Dam in Mali: a primarily irrigation dam, serving about 75,000 ha of farmland, (3)
184 Goronyo Dam in Nigeria: a multi-purpose dam for flood control, provision of downstream water supply and the release

185 of water for irrigation in the dry season, (4) Bakolori Dam in Nigeria: a primarily irrigation dam with a command area
 186 of about 23,000 ha, (5) Kainji Dam in Nigeria: the largest Dam on the Niger supplying power for most towns in
 187 Nigeria, (6) Jebba Dam in Nigeria: a primarily hydropower dam, (7) Dadin Kowa Dam: a multi-purpose dam for
 188 water supply, electricity and irrigation, (8) Lagdo Dam in Cameroon: a multi-purpose dam providing electricity to the
 189 northern part of the country and supplying irrigation water for 15,000 hectares of cropland. The watersheds of the
 190 dams are primarily either in the Savanna (Selingue, Markala, Jebba, Dadin Kowa, an Lagdo), or in the Sahel (Goronyo,
 191 Kainji), or partly in both (Bakolori). The watershed sizes vary over a large range, from 4,887 Km² (Bakolori Dam) to
 192 1,464,092 Km² (Kainji Dam). The average elevations of the watersheds are close to each other at 500 ± 50 m.a.s.l.

193
 194 In order to make the results of this study meaningful to reservoir managers, the Niger basin was divided into
 195 watersheds according to the locations of the dam reservoirs (see Fig. 1). Then the sub-basin of each dam was defined
 196 as the drainage between the dam itself and the upstream dam. For example, the drainage basin of the Markala Dam
 197 does not include the drainage basin of the Selingue Dam.

198
 199 Table 1. Selected dams and their watershed characteristics

Dam	Country	Operational since*	Capacity (million m ³)*	Power (MW)*	Primary Purpose*			Area of Drainage Basin (km ²)**	Elevation of Drainage Basin (m)**
					Irrigation and Water Supply	Flood Control	Hydroelectricity		
Selingue	Mali	1982	2170	44			x	32685	473
Markala	Mali	1947	175		x			102882	442
Goronyo	Nigeria	1983	942		x	x		31547	446
Bakolori	Nigeria	1978	450		x			4887	519
Kainji	Nigeria	1968	15000	960			x	1464092	406
Jebba	Nigeria	1984	3600	540			x	40268	308
Dadin Kowa	Nigeria	1988	2855	35	x		x	32936	535
Lagdo	Cameroon	1983	7800	72		x	x	31352	452

200 * information obtained from the Global Reservoir and Dam Database (Lehner et al. 2011) and Food and Agriculture Organization
 201 of the United Nations (FAO)'s Global Information System on Water and Agriculture (AQUASTAT).

202 ** Calculated from HydroSEHDS (Lehner et al. 2008).
 203
 204

205 2.5 Evaluation

206 IMERG Final rainfall products are used in this study as reference to evaluate the performance of GFS precipitation
 207 forecasts. The comparison period is 15 June 2019 to 15 June 2020 to match the period for which the version-15 of

208 GFS model forecasts is available. The spatial resolutions of the forecast and satellite products are different: 0.25°(GFS),
 209 0.10° (IMERG Final and IMERG Early), and 0.05° (CHIRPS). The temporal resolutions of the satellite products are:
 210 30-minute (IMERG Final and IMERG Early) and daily (CHIRPS). Our comparison is mostly based on sub-basin (i.e.
 211 watershed for each dam) average values, in which case we average all the datasets to the sub-basin spatial scale. In
 212 some cases, where we compare the spatial patterns of rainfall, we resample both IMERG products and CHIRPS to
 213 0.25° using the bilinear interpolation technique to match the spatial resolution of GFS.

214
 215 For evaluation metrics, we used the modified Kling-Gupta Efficiency (KGE; Gupta et al. 2009; Kling et al. 2012) and
 216 its components: Bias Ratio (BR), correlation (R), and variability ratio (γ). *KGE* measures the goodness-of-fit between
 217 estimates of precipitation forecasts and reference observations as:

$$KGE = 1 - \sqrt{(R - 1)^2 + (BR - 1)^2 + (\gamma - 1)^2},$$

$$BR = \frac{\mu_f}{\mu_o},$$

$$\gamma = \frac{CV_f}{CV_o},$$

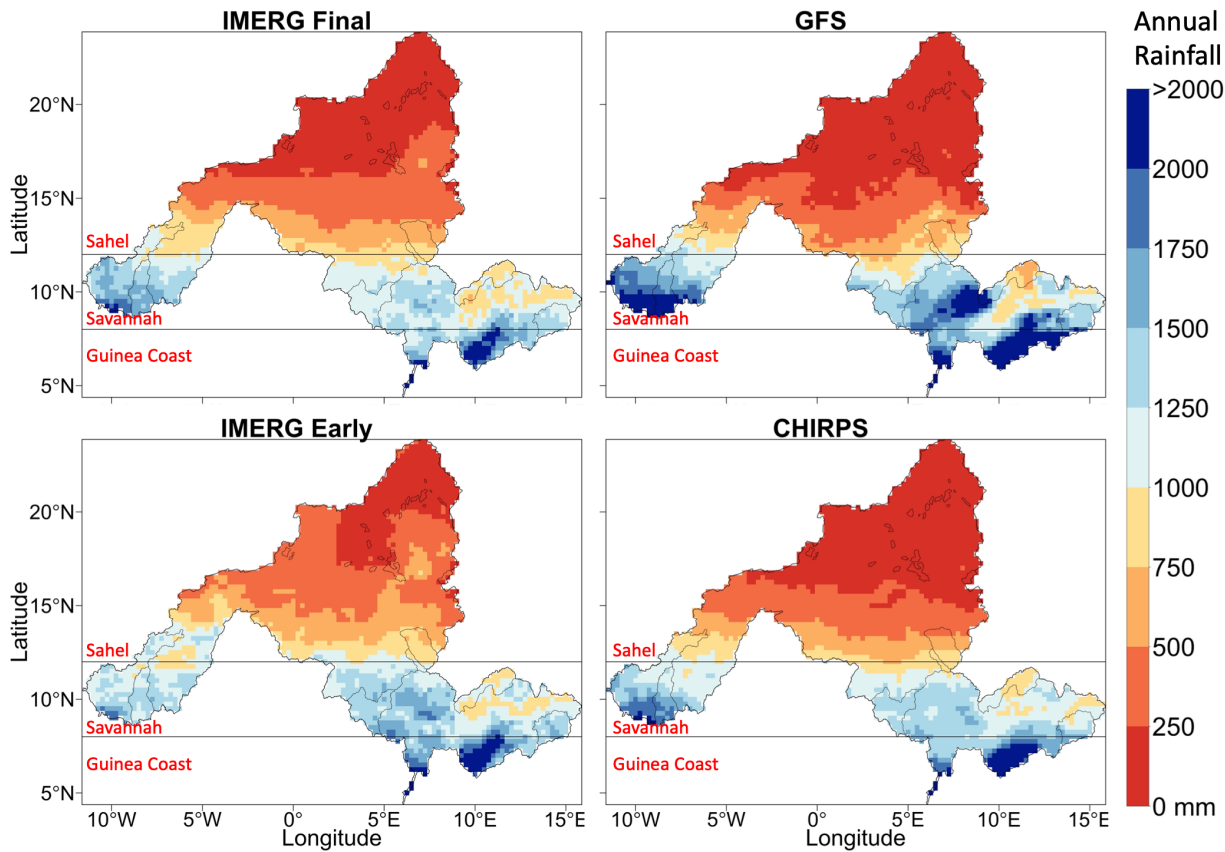
221 where R is the linear correlation coefficient between forecasted and observed precipitation, BR is the bias ratio, γ is
 222 the variability ratio, μ is the mean precipitation, CV is the coefficient of variation, and the indices *f* and *o* represent
 223 forecasted and observed precipitation values, respectively. KGE values range from $-\infty$ to 1, with values closer to 1
 224 indicating better model performance. Towner et al. (2019) suggested the following classifications: “Good” ($KGE \geq$
 225 0.75), “Intermediate” ($0.75 \geq KGE \geq 0.5$), “Poor” ($0.5 \geq KGE > 0$), and “Very poor” ($KGE \leq 0$). The *BR* values
 226 greater than 1 indicate a positive bias whereby forecasts overestimate precipitation relative to the observed data,
 227 while values less than 1 represent an underestimation. The γ values greater than 1 indicate that the variability in the
 228 forecast time series is higher than that observed, and values less than 1 show the opposite effect. The *R* measures the
 229 strength and direction of the linear relationship between the forecast and observed values, and to what extent the
 230 temporal dynamics of observed rainfall is captured in the forecasts. The correlation values of 0.6 or more are
 231 considered to be skillful (e.g., Alfieri et al. 2013). In addition, the root mean-square-error normalized by reference
 232 precipitation mean (NRMSE) was also used.

233

234 **3. Results and Discussion**

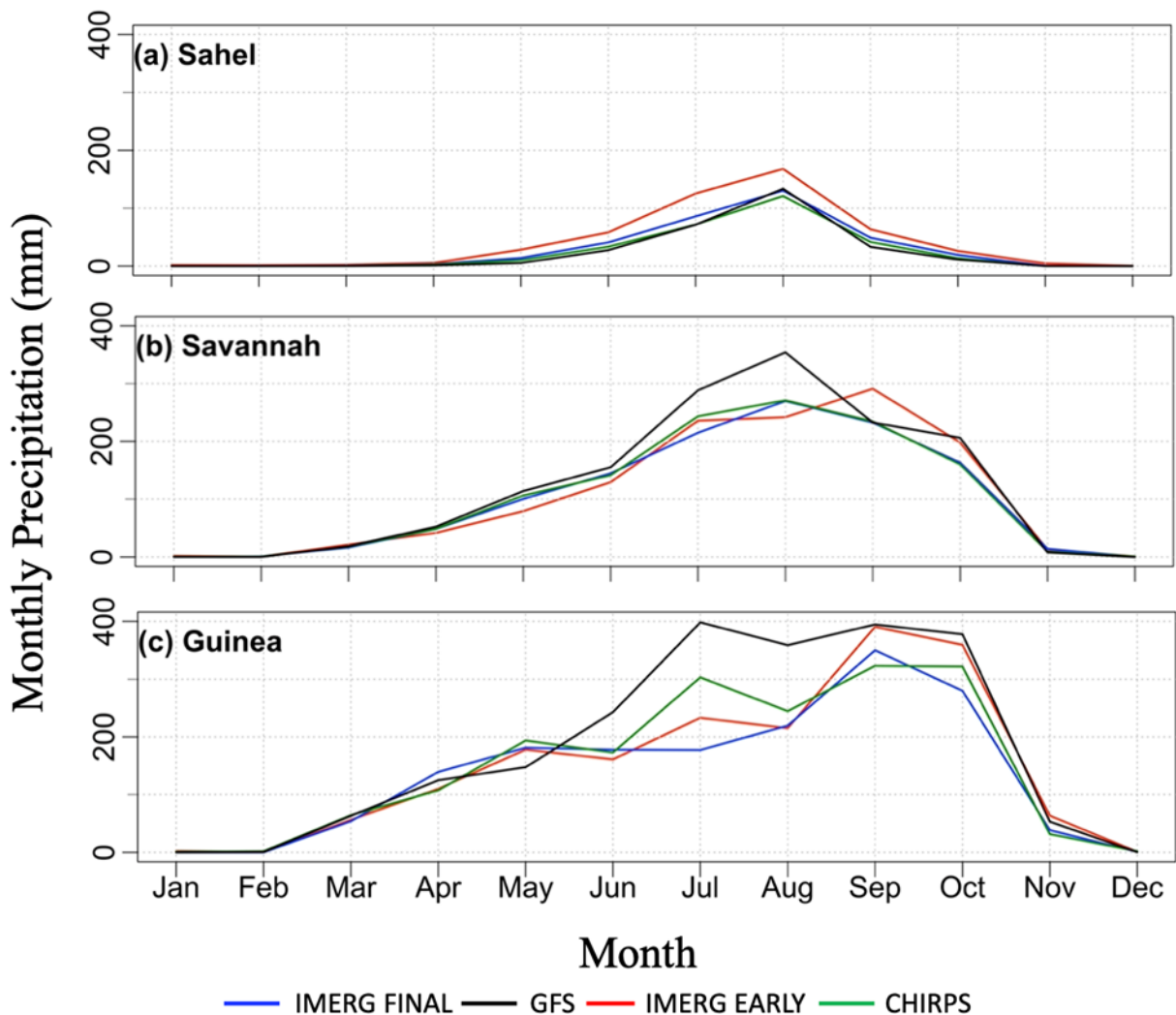
235 **3.1 Annual Spatial Variability and Seasonal Characteristics**

236 The spatial map of annual (15 June 2019 – 15 June 2020) rainfall from the various rainfall products is given in Figure
237 2. According to the reference rainfall product (i.e. IMERG Final), the Niger basin experiences average annual rainfall
238 of 700 mm. The spatial rainfall distribution shows north-to-south increasing gradient, with the Sahel region (> 12°N)
239 receiving on average 346 mm per year, the Savanna region (8°N – 12°N) receiving on average 1,206 mm per year,
240 and the Guinea region (4°N – 8°N) receiving on average 1,620 mm per year. The spatial structures (climatology and
241 north-south gradient in rainfall) of GFS, IMERG and CHIRPS rainfall fields are quite similar to those of IMERG
242 Final. However, the 1-day GFS tends to overestimate in the wet Guinea region of the basin, whereas both IMERG
243 Early and CHIRPS give values that are very close to IMERG Final.



244
245 Figure 2. Spatial map of annual rainfall (in mm), for the period 15 June 2019 to 15 June 2020, derived from (a)
246 IMERG Final, (b) GFS (1-day lead time), (c) IMERG Early, and (d) CHIRPS.
247

248 Figure 3 shows the seasonal rainfall pattern for each climatological region. According to the reference IMERG Final,
 249 as one goes from north to south, the rainy season expands from 3 months (June – September) in the Sahel to 6 months
 250 (March – November) in the Savanna and Guinea regions. The peak rainfall also shows north-south gradient, with peak
 251 rainfall of 130 mm in the Sahel, to 269 mm in the Savanna, and 350 mm in the Guinea. The rainfall pattern is unimodal
 252 with a peak rainfall value in August for both Sahel and Savanna, but becomes bimodal with one peak in May and the
 253 other in September for Guinea.
 254



255 Figure 3. Monthly precipitation regime for the three climatological zones of the Niger river Basin: (a)
 256 Sahel, (b) Savanna, and (c) Guinea. Analyses are based on rainfall fields derived from IMERG Final, 1-
 257 day-lead GFS, IMERG Early, and CHIRPS. The time period covers from 15 June 2019 to 15 June 2020.

258 When validated against IMERG Final, the performance of GFS in capturing the seasonal rainfall characteristics
259 deteriorates as one goes from north to south. GFS captures both the seasonal rainfall pattern and rainfall peak in the
260 Sahel, and captures the seasonal rainfall pattern but tends to moderately overestimate the peak in the Savannah, while
261 it has large overestimation (almost twice as much as the reference) in the Guinea particularly during summer. As far
262 as the other satellite products are concerned, IMERG Early tends to slightly overestimate in the Sahel across all rainy
263 months, but performs relatively well in the Savannah and Guinea regions. CHIRPS is very close to IMERG Final in
264 all regions and months, with the exception of modest overestimation of the July rainfall in Guinea.

265

266 **3.2 How well do GFS forecasts capture annual rainfall?**

267 Here, we aggregate the 1-day lead GFS forecasts to annual time scale and compare the results against corresponding
268 annual precipitation estimates from IMERG Final. Figure 4 presents the watershed-averaged annual rainfall for each
269 dam watershed. According to IMERG Final, the annual rainfall varies from 434 mm (in Kainji) to 1,481 mm (in
270 Selingue). Watersheds 1 (Selingue) and 2 (Markala), located in the western part of the Savannah, receive the largest
271 amount of rainfall, i.e., 1481 mm and 1406 mm, respectively. Watershed 3 (Markala), located in the eastern part of
272 the Sahel, receives 741 mm of annual rainfall. Watershed 4 (Bakolori), characterized by the smallest watershed area
273 compared to the rest of the watersheds, lies partly in the Sahel and partly in the Savannah region and receives 921 mm
274 of annual rainfall. Watershed 5 (Kainji), characterized by the largest watershed area of all, lies mostly in the Sahel
275 region and receives the lowest amount of annual rainfall (434 mm). Watersheds 6 (Jebba), 7 (Dadin Kowa), and 8
276 (Lagdo), located in the Savannah, receive annual rainfall amounts of 1190 mm, 941 mm, and 1295 mm, respectively.

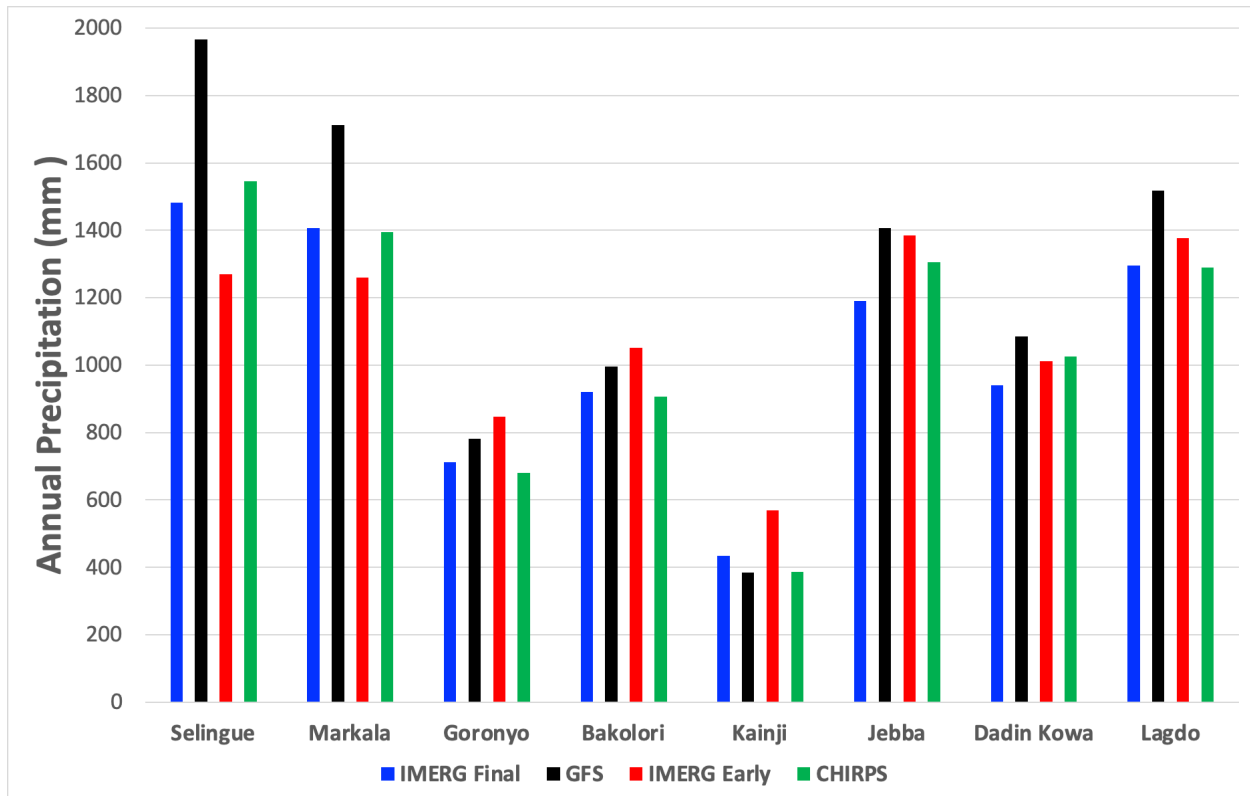
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278 Validated against IMERG Final, the GFS tends to overestimate rainfall in all watersheds located in the Savannah (or
279 watersheds that receive relatively large rainfall amounts), with an overestimation varying in the range 8% to 33%,
280 with larger bias for watersheds receiving higher rainfall amount. For watersheds in the Sahel (watersheds receiving
281 low rainfall amount), GFS gives less bias (-11% for the driest Kainji watershed and +10% for Bakolori).

282

283 In contrast, IMERG Early tends to underestimate rainfall in all watersheds located in the Savannah (with larger
284 negative bias in watersheds with large rainfall amount) but tends to overestimate in all watersheds located in the Sahel
285 (with very large overestimation bias for the driest watershed) Therefore, GFS and IMERG Early have different bias

286 characteristics: whereas GFS outperforms IMERG Early in the Sahelian climate where well-organized convective
 287 systems dominate the monsoon, IMERG Early outperforms GFS in the Savannah and Guinea climate which are
 288 characterized by short-lasting and localized systems and wet land surface conditions. CHIRPS estimates are
 289 reasonably close to IMERG Final, indicating that the choice of reference product between CHIRPS and IMERG Final
 290 would not substantially affect the findings on the accuracy of GFS forecasts.
 291

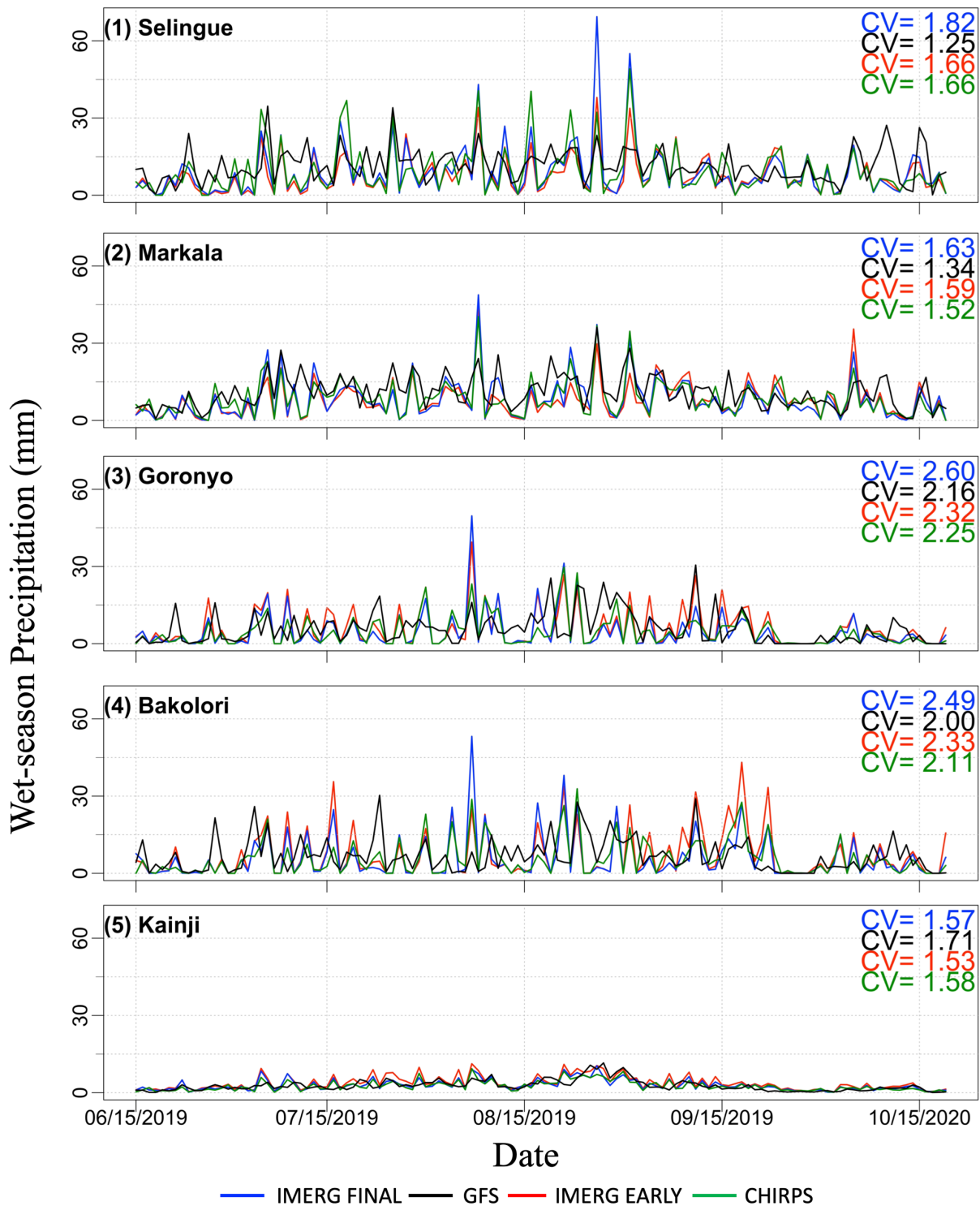


292
 293 Figure 4. Sub-basin averaged annual precipitation (mm) for the period, 15 June 2019 to 15 June 2020, for each
 294 of the Niger’s sub-basin, derived from the 1-day lead GFS forecast and different satellite precipitation
 295 products.

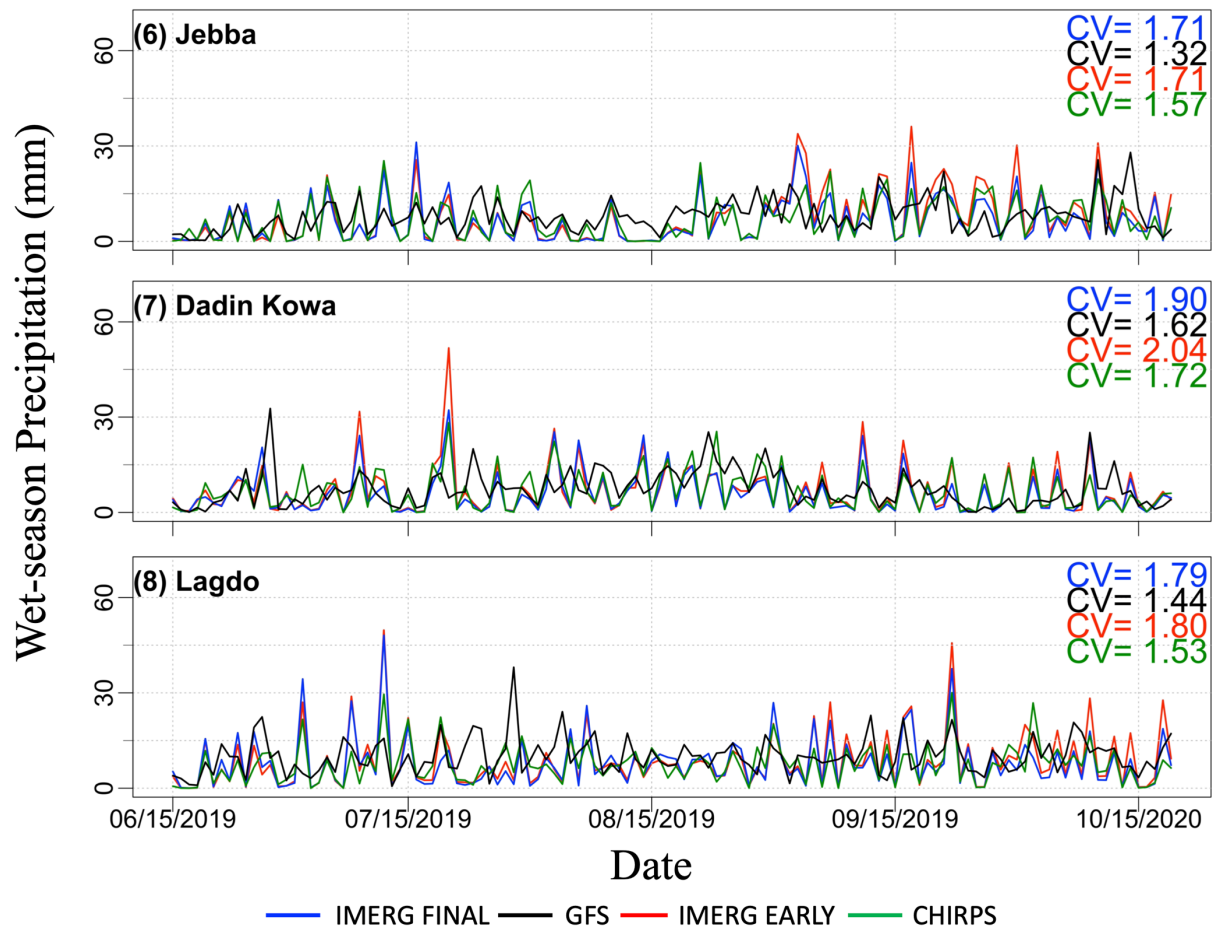
296 **3.3 How well is the time series of daily precipitation forecasted?**

297 Figures 5 and 6 present the time series of watershed-averaged daily rainfall, for the wet period June – October.
 298 According to IMERG Final, the temporal variability (as measured through coefficient of variation or CV) varies from
 299 1.22 to 2.60. Validated against IMERG Final, the GFS tends to underestimate the temporal variability and particularly
 300 underestimate large spikes in rainfall, at almost all sites except at Kainji. The GFS’s relatively better performance for

301 Kainji could be attributed to the watershed's large area that results in relatively smooth temporal variability. Both
 302 IMERG Early and CHIRPS provide CV values that are very close to IMERG Final.



303 Figure 5. Time series of sub-basin averaged precipitation total (mm) for the wet period (June – September 2019 for
 304 all sub-basins, derived from various precipitation products, for five sub-basins. The Figure also shows the
 305 coefficient of variation (CV) as a measure of temporal variation.



306 Figure 6. Same as in Figure 5 but for the remaining three watersheds.

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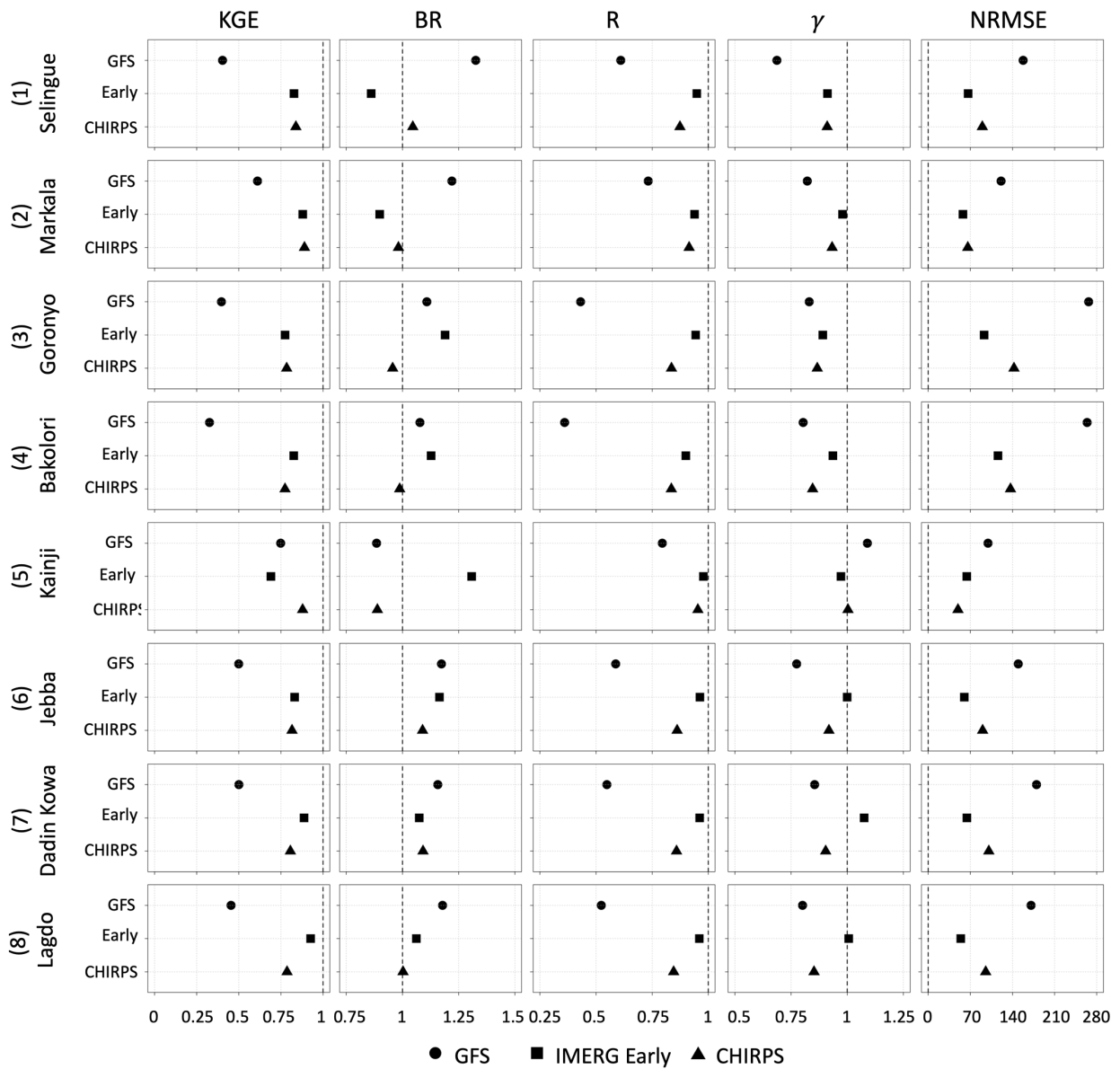
308 Figure 7 displays the performance statistics of watershed-averaged daily rainfall (validated against IMERG Final) in
 309 terms of Kling-Gupta Efficiency (KGE), Bias Ratio (BR), correlation (R), variability ratio (γ), and root mean square
 310 error normalized by reference precipitation mean (NRMSE). First, the performance results for the 1-day lead GFS are
 311 considered. The KGE scores are poor ($0.3 < KGE < 0.5$) for half of the watersheds considered (Selingue, Goronyo,
 312 Bakolori, and Lagdo) and intermediate ($0.5 < KGE < 0.75$) for the remaining half watersheds (Markala, Kainji, Jebba,
 313 and Dadin Kowa). The breakdown of the KGE scores (BR, R, and γ) reveals the key factors contributing to the KGE
 314 estimates. The GFS tends to overestimate daily precipitation for most sub-basins, as BR is higher than one, except for
 315 Kainji. The overestimation is particularly high for Selingue and Markala, where BR is 1.33 and 1.22, respectively.
 316 The correlation coefficient between GFS and IMERG Final is mostly low ($R < 0.60$), and is particularly lower for
 317 Bakolori ($R=0.36$) and Goronyo ($R=0.43$). The variability ratio of GFS is mostly between 0.69 to 0.83 (except for

318 Kainji, where γ is 1.09), indicating that the GFS tends to give lower temporal variability of rainfall. The NRMSE is
319 very high, ranging from 100% to 266%, and is particularly high for Goronyo (266%) and Bakolori (264%), which are
320 relatively small-sized watersheds.

321
322 Next, the performance of IMERG Early was examined with respect to IMERG Final, mainly to assess if it is possible
323 to use the near-real-time IMERG Early product to calibrate and improve the accuracy of GFS forecasts. The IMERG
324 Early performs much better with KGE values higher than 0.75 (except for Kainji where KGE is 0.69), correlation
325 higher than 0.90, and variability ratio close to the optimum value. The high performance of IMERG Early is due to its
326 similarity with the IMERG Final product, as the main difference between the two products is that IMERG Early,
327 unlike IMERG Final, does not use monthly rain gauge observations for bias correction. Such monthly bias correction
328 techniques would not alter the pattern and variability of IMERG Early compared to IMERG Final. Therefore, the
329 performance of IMERG Early should be evaluated using bias statistics, the other statistics (correlation and variability
330 ratio) are presented for completeness. IMERG Early overestimates rainfall in most watersheds in the range 11%
331 (Lagdo) to 28% (Kainji) except for two watersheds, where it slightly underestimates by 14% (Selingue) and 11%
332 (Markala). Comparison of the performance of GFS and IMERG Early indicates that both products have different bias
333 characteristics. In some watersheds (e.g., Kainji), GFS outperforms IMERG Early in terms of bias, whereas in other
334 watersheds (e.g., Markala), IMERG Early outperforms GFS.

335
336 CHIRPS was also compared with IMERG Final to assess how the use of different reference products may affect the
337 finding about the performance of GFS forecasts. The KGE scores of CHIRPS are higher than 0.75 in all cases,
338 indicating that CHIRPS and IMERG Final have comparable KGE performance. Therefore, the performance of GFS
339 is expected to be about the same even if the reference product used this in this study (IMERG Final) changes to
340 CHIRPS.

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348 Figure 7. Summary of performance statistics (Kling-Gupta Efficiency KGE, Bias Ratio BR, correlation R,
 349 variability ratio γ , and root mean square error normalized by reference rainfall [%], for the 1-day lead time GFS
 350 forecasts and other satellite products. The time period considered was June 15, 2019 – June 15, 2020.
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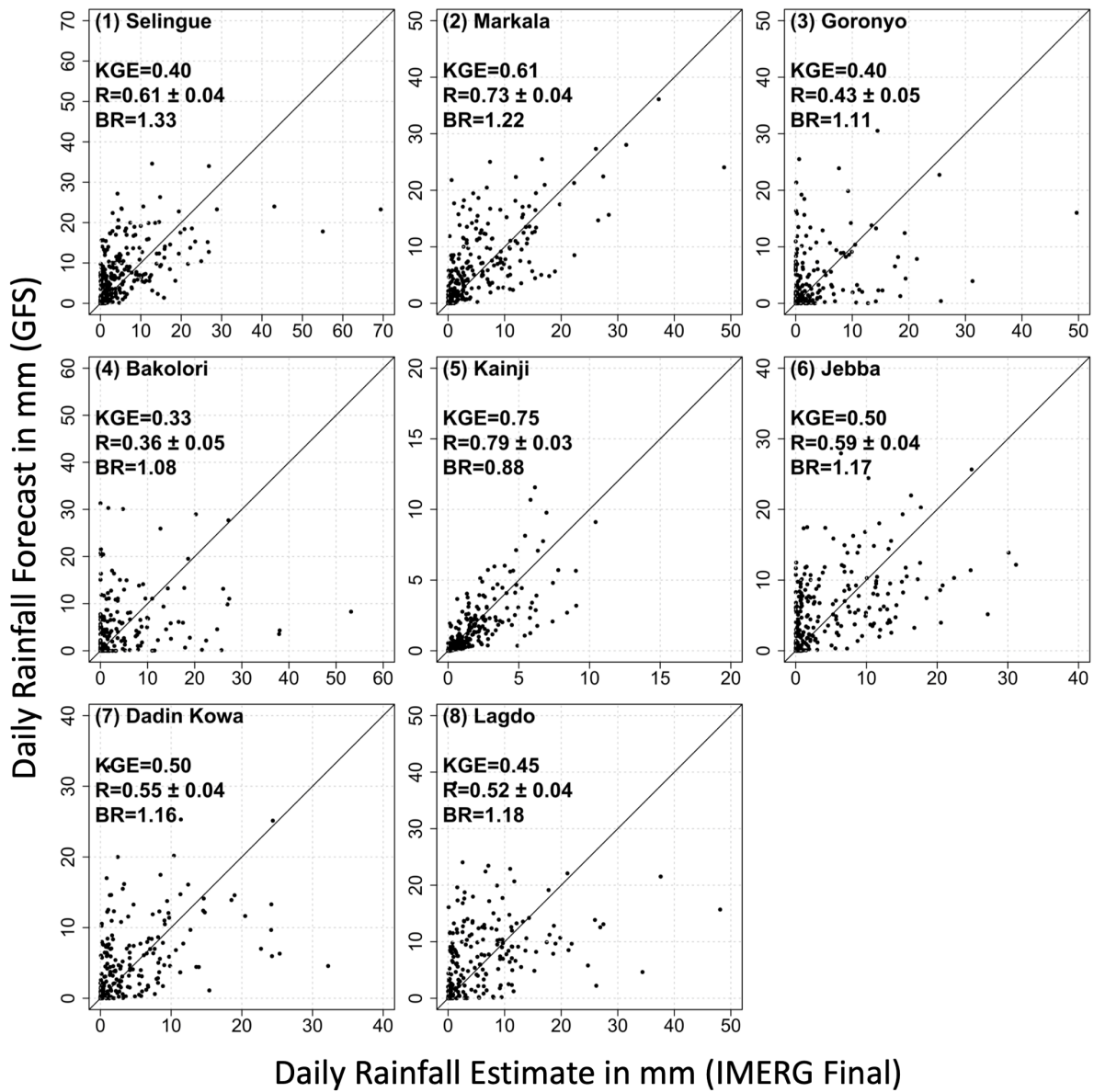


Figure 8. Scatterplot of watershed-averaged daily precipitation forecast obtained from 1-day lead GFS forecasts against corresponding values from IMERG Final.

355 3.4 Dependence of Forecast Performance on Precipitation Rate

356 Figure 8 presents the scatterplot of 1-day lead GFS forecasts and IMERG Final at daily and watershed-average scales.

357 The performance of GFS varies between watersheds. In the Markala and Kainji watersheds, GFS forecasts agree well

358 with IMERG Final at almost all rain rates. In the Selingue watershed, GFS agrees well with IMERG Final for rain

359 rates under 30 mm/day, but GFS substantially underestimates all rain rates above 30 mm/day. In the remaining five

360 watersheds, GFS has poor performance, replete with large scatter, high false alarm, and large underestimation bias of
361 heavy rain rates.

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364 **3.5 Dependence of Daily Forecast Performance on Lead Time and Spatial Scale**

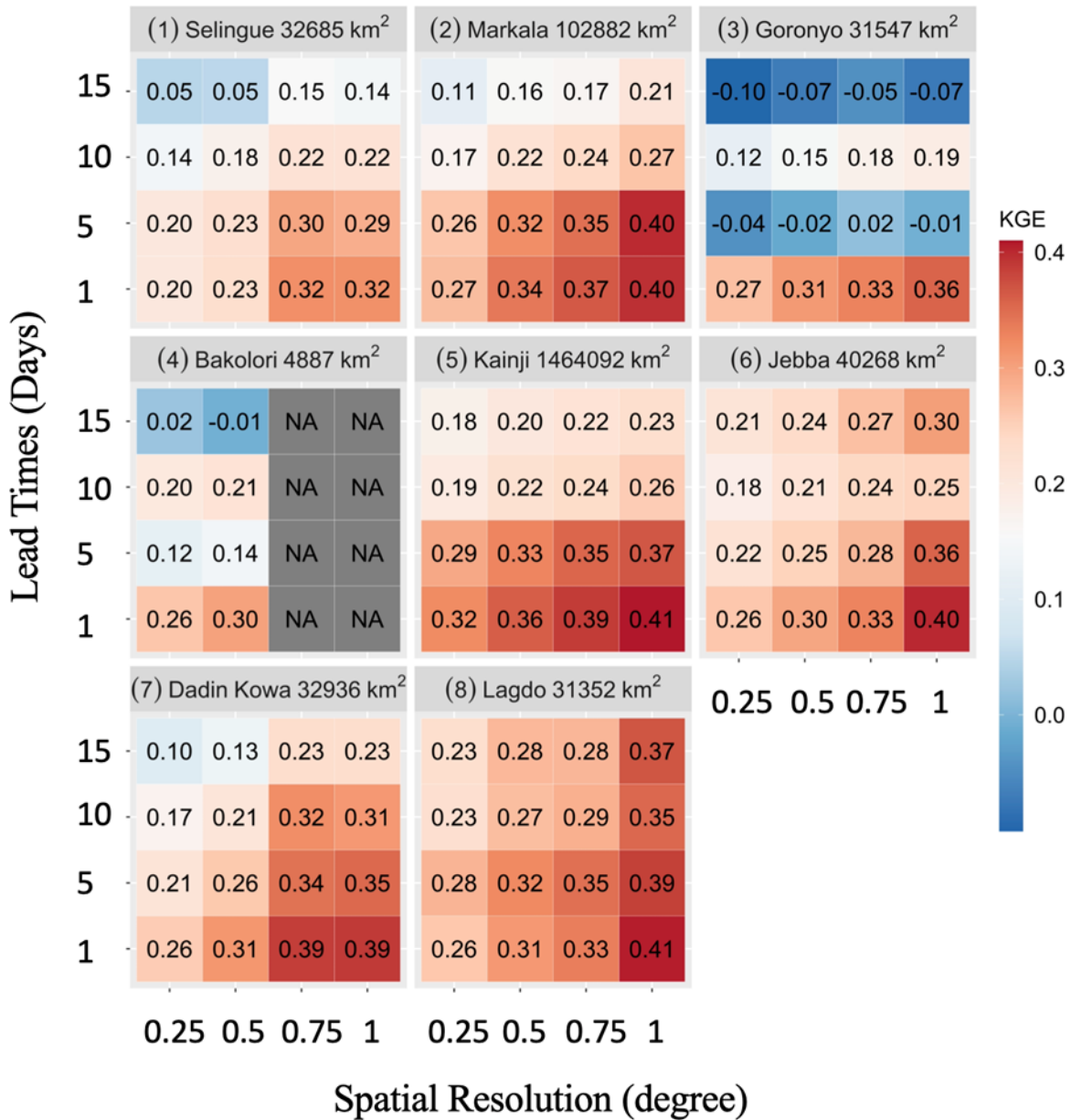
365 In order to assess the effect of various lead times and spatial scales on forecast performance, we obtained daily GFS
366 forecasts at various lead times (1-day, 5-day, 10-day, and 15-day), and aggregated the forecasts at spatial scales from
367 0.25° to coarser scales (0.5° , 0.75° , and 1°) by averaging grids. The purpose of degrading the resolution is to determine
368 at which resolution the forecasts have acceptable performance. The KGE value at each spatial resolution was
369 calculated in the following steps: (i) average the data at the required spatial resolution, (ii) extract pairs of data (one
370 from IMERG Final, and the other from GFS), (iii) concatenate the pairs to form one large series of data, and (4)
371 compute a single KGE from this data series. The resulting KGE values are shown in Fig. 9.

372

373 With regard to the effect of spatial scales, the KGE at the GFS native resolution (i.e. 0.25°) is very low. As the spatial
374 scale increases, KGE increases, as expected. For instance, for Markala watershed KGE increases from 0.27 (0.25°) to
375 0.40 (1°) for a 1-day lead. This indicates that the variation in KGE values between the watersheds could be partly
376 explained by the watershed size. For example, based on Fig. 5, the KGE for the 1-day lead daily GFS forecast was the
377 highest for the largest Kainji watershed (watershed area of $1,464,092 \text{ Km}^2$) and the lowest for the smallest Bakolori
378 watershed ($4,887 \text{ Km}^2$). With regard to the effect of lead time for daily forecasts, KGE decreases significantly as lead
379 time increases. For instance, for Markala watershed and a grid size of 1° , KGE decreases from 0.40 (1-day lead) to
380 0.21 (15-day lead).

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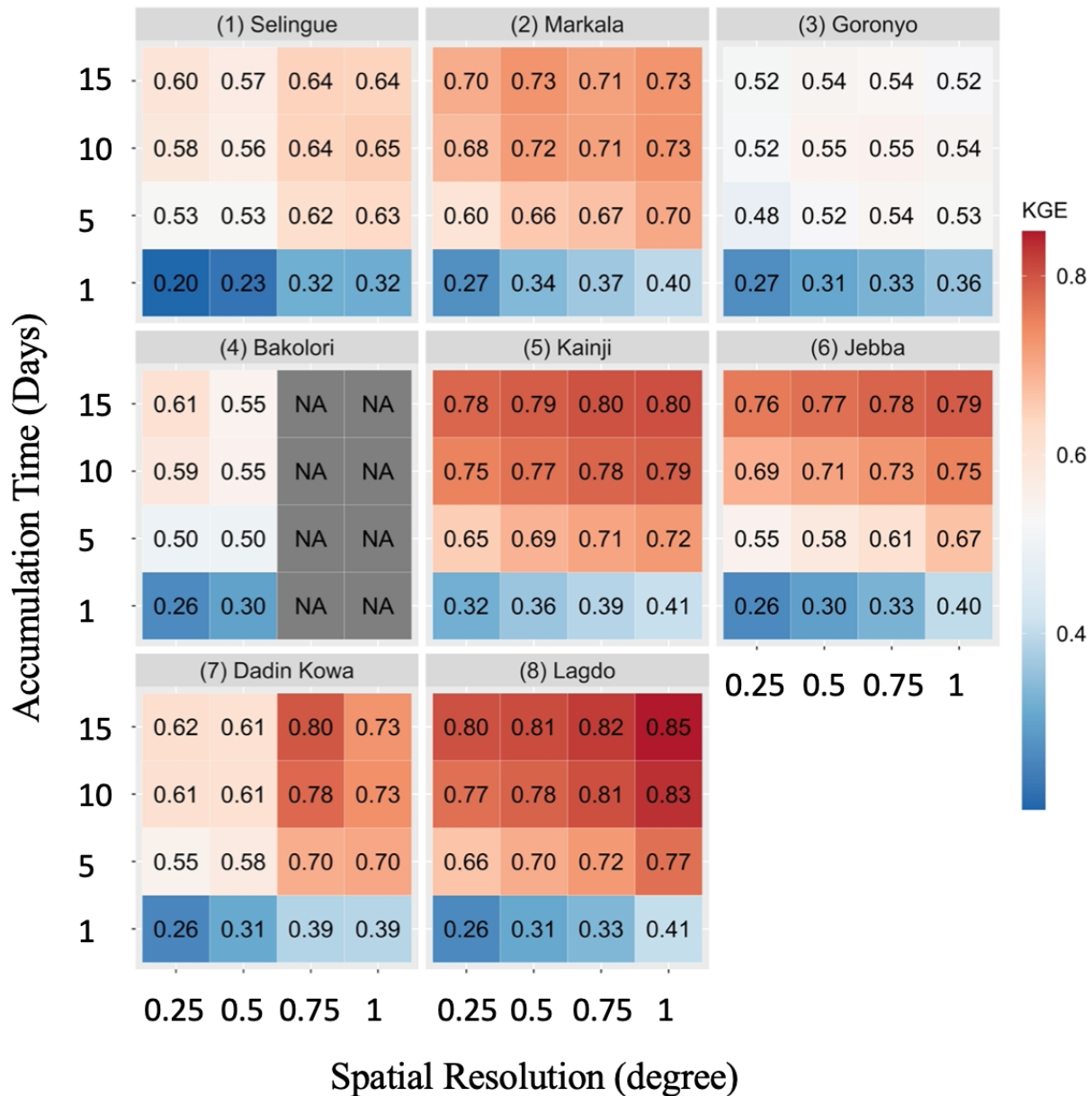
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Figure 9. Kling-Gupta Efficiency (KGE) for daily precipitation of GFS as a function of lead time (1-day, 5-day, 10-day, and 15-day) and spatial scale (0.25°, 0.50°, 0.75°, 1.0°). The dam names and corresponding watershed areas are given in the titles.

389 3.6 Effect of Temporal Aggregation Scale on Forecast Performance

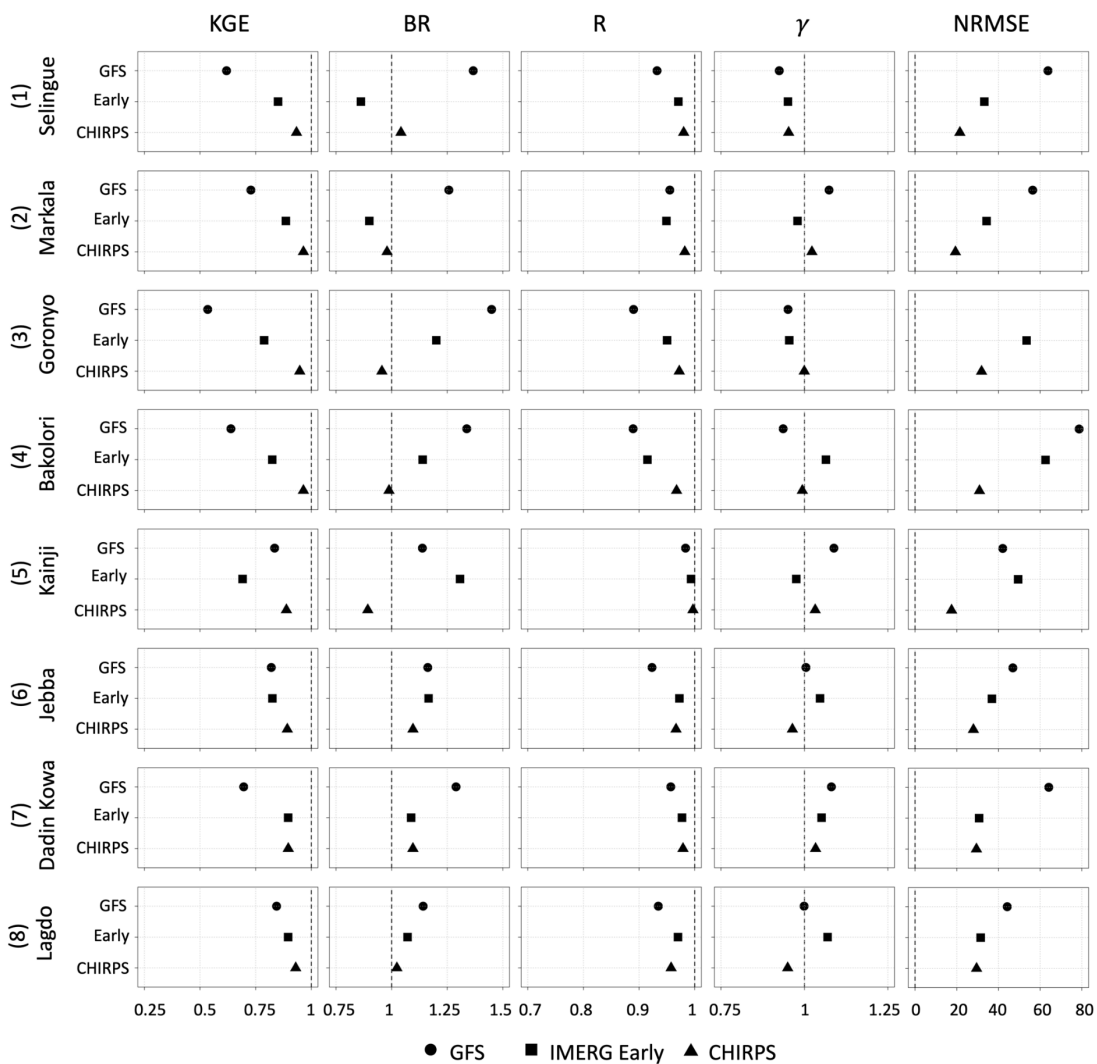
390 To assess the effect of temporal aggregation scale, we obtained the 1-day total, 5-day total, 10-day total, and 15-day
391 total GFS precipitation forecasts. These multi-day forecasts are constructed by combining multiple lead-time forecasts.
392 For instance, the 5-day total forecast is obtained by adding the 1-day lead, 2-day lead, 3-day lead, 4-day lead, and 5-

393 day lead daily forecasts. Figure 10 presents the KGE values for GFS forecasts over different temporal aggregation
 394 scales, and different grid sizes. Temporal aggregation substantially increases KGE at all spatial scales. For example,
 395 at the grid size of 1° over Markala watershed, the KGE values jump from 0.40 at daily timescale to 0.73 at 15-day
 396 total timescale.

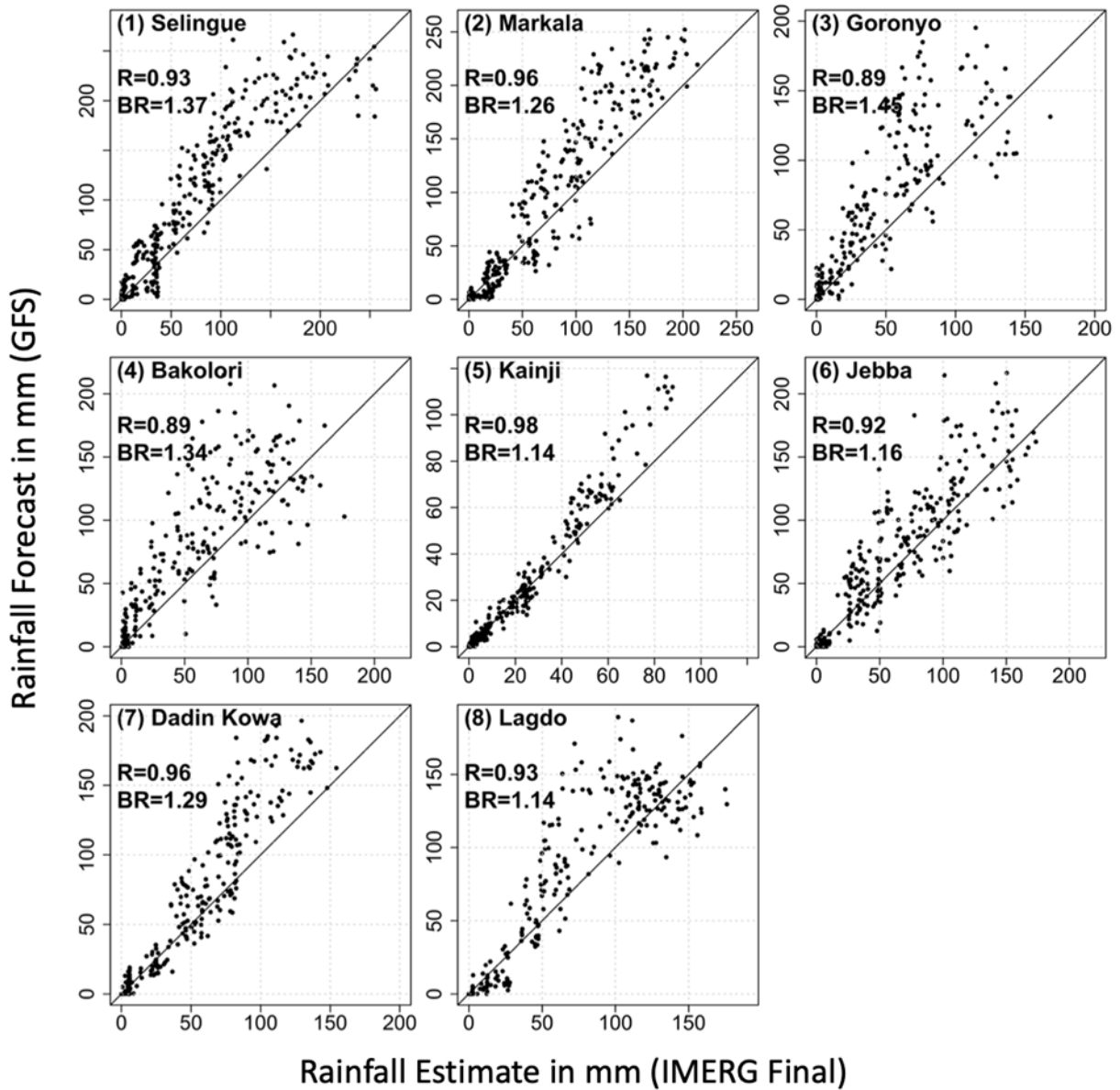


397 Figure 10. Kling-Gupta Efficiency (KGE) of GFS as a function of accumulation time scale (1-day, 5-day, 10-day,
 398 and 15-day) and spatial scale (0.25°, 0.50°, 0.75°, 1.0°).
 399

400 In Figure 11, we show the performance statistics of GFS for 15-day accumulated watershed-averaged rainfall forecast.
 401 The KGE values are intermediate ($0.5 < KGE < 0.75$) for four watersheds and good ($KGE > 0.75$) for the remaining
 402 four watersheds. Analysis of the components of KGE reveals that the improvement of KGE at longer timescales comes
 403 as a result of improved correlation and variability ratio. At the 15-day accumulation timescale, IMERG Early estimates
 404 have less bias than GFS at all watersheds, except at Kainji watershed. Figure 12 presents the scatterplot of 15-day
 405 accumulated GFS forecast vs IMERG Final. In general, the GFS estimates perform well for low to moderate rain rates,
 406 but tend to overestimate higher rain rates. This is consistent with Wang et al. (2019) who reported the difficulty of
 407 capturing the magnitude of high rain rates in GFS model.



408 Figure 11. Summary of performance statistics (Kling-Gupta Efficiency KGE, Bias Ratio BR, correlation R,
 409 variability ratio γ , and root mean square error normalized by reference rainfall [%], for the 15-day accumulated GFS
 410 forecast and other satellite products.
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412
 413 Figure 12. Scatterplot of watershed-averaged 15-day accumulated precipitation forecast obtained from GFS forecast
 414 against corresponding values from IMERG Final.
 415

416 **4. Conclusions**

417 This study has evaluated the accuracy of medium-range (1-day to 15-day lead time) forecasts available from the Global
 418 Forecast System (GFS), for the watersheds of large dams in the Niger river basin. Despite the limited temporal
 419 coverage, some consistent features emerged from this evaluation. The accuracy of GFS forecast depends on climatic
 420 regime, lead time, accumulation timescale, and spatial scale. With regard to the role of climatic regimes, the GFS

421 forecast has large overestimation bias in the Guinea (wet climatic regime), moderate overestimation bias in the
422 Savannah (moderately wet climatic regime), but has no bias in the Sahel (dry climate). With regard to lead time, as
423 the lead time increases, the forecast accuracy decreases. Averaging the forecasts at coarser spatial scales leads to
424 increased forecast accuracy. For daily rainfall forecasts, the performance of GFS is very low ($KGE < 0.32$) at almost
425 all watersheds except at Markala ($KGE = 0.44$) and Kainji ($KGE = 0.68$), both of which have much larger watershed
426 areas compared to the other watersheds. Averaging the forecasts at longer time scales also leads to increased forecast
427 accuracy. For 15-day rainfall accumulation timescale, the KGE values are either “intermediate” (i.e., $0.50 \leq KGE \leq$
428 0.75) for half of the watersheds (Selingue, Goronyo, Bakolori, and Daddin Kowa) or “good” (i.e., $KGE \geq 0.75$) for
429 the remaining half (Markaa, Kainji, Jebba, and Lagdo). With regard to the effect of rainfall rate, the 15-day
430 accumulated GFS forecasts tend to perform better for low to medium rain rates, but contain large overestimation bias
431 at high rain rates.

432
433 The performance statistics of GFS indicate the need for calibrating GFS forecasts in order to improve their accuracy.
434 Post-processing calibration of GFS forecasts requires the use of “relatively better performing” and “available in near-
435 real-time” independent rainfall observations to correct real-time dynamical GFS model forecasts. This study has
436 compared the performance of IMERG Early satellite rainfall products with the performance of GFS in terms of bias.
437 In the Guinea and Savannah regions, IMERG Early outperforms GFS in terms of bias, while in the dry Sahel region,
438 IMERG Early is outperformed by GFS.

439
440 We acknowledge that the reference dataset used in our evaluation (i.e., IMERG Final) has its own estimation errors.
441 We conducted additional assessment to evaluate the performance of IMERG Final with respect to another independent
442 and high-quality (i.e. satellite-gauge merged) rainfall product (i.e. CHIRPS). Our results show that IMERG Final and
443 CHIRPS have similar rainfall characteristics, indicating the robustness of IMERG Final.

444
445 Overall, we conclude that the GFS forecasts, at 15-day accumulation timescale, have acceptable performance,
446 although they tend to overestimate high rain rates. The shorter the time scale, the lower is the GFS performance. Given
447 that IMERG Early outperforms GFS particularly in the wet region of the Niger, we recommend testing the suitability
448 of IMERG Early to serve as input into post-processing of GFS in order to improve the accuracy of GFS forecasts.

449 Possible post-processing techniques that could be explored include: simple bias (multiplicative) correction
450 (Gumindoga et al. 2019), multi-resolution bias correction through wavelet analysis (Xu et al. 2019) or empirical mode
451 decomposition method (Wang et al. 2020, Prasad et al. 2019), and Artificial Intelligence-based methods such as Feed
452 Forward Neural Network (Cloud et al. 2019), Support Vector Machine ((Du et al. 2017; Yu et al. 2017), and Adaptive
453 Neural Fuzzy Inference System (Jehanzaib et al. 2021).

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479 **5. Data and Code Availability**

480 We acknowledge the National Center for Atmospheric Research (NCAR) for providing public access to the GFS
481 rainfall forecast data products (<https://rda.ucar.edu/datasets/ds084.1/>), NASA for providing public access to IMERG
482 Final and IMERG Early rainfall data products (<https://disc.gsfc.nasa.gov>), and the University of California Santa
483 Barbara's (UCSB) Climate Hazard's group for providing public access to CHIRPS rainfall data
484 (<https://www.chc.ucsb.edu/data>).

485

486

487 **6. Author Contribution**

488 H. Yue: data processing, data analysis, and manuscript preparation; M. Gebremichael: project oversight, method
489 design, contribution to manuscript text; V. Nourani: method design, contribution to manuscript text.

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492 **7. Competing Interests**

493 The authors declare that they have no conflict of interest.

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496 **8. Acknowledgement**

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