

Response to Anonymous Referee #1

Major comment: Given that gas analyzers were used that simultaneously measure both CH₄ and CO₂, I really think that the authors should show the system's performance for CO₂ as well. In L363, the authors write that have CO₂ data but focus on CH₄ for simplicity, but this choice makes life much less simple for all the researchers that want to measure both CH₄ and CO₂, and therefore need to do all the CO₂ testing themselves. Showing the CO₂ results would make this study much more useful and applicable for a much wider community, and certainly render more citations to this paper. At the very minimum, include the CO₂ performance tests in the supplementary information, but I'd rather see that the CO₂ data is fully integrated in the paper, including the title.

Response: We totally agree with the reviewer that CO₂ should be included in the manuscript. We integrated CO₂ results in the manuscript and mention CO₂ in the title as well. Please see the manuscript.

Minor comments:

Title: Include "carbon dioxide".

Response: We included CO₂ in the revised manuscript.

L13. Freshwater lakes and reservoirs are aquatic systems, so that's a repetitive formulation. Simplify.

Response: We deleted "in freshwater lakes and reservoirs" in Line 13. Please see line 14-15 in the revision.

L48-49. This sentence omits that dissolved CH₄ concentration is very strongly a function of methanogenesis, this should be added.

Response: We added "In addition to formation processes that lead to CH₄ accumulation" in the revised manuscript. Please see line 61-62 in the revision.

L69. Not only phytoplankton, but also other microbial life forms. I suggest to reword to "biological".

Response: We agree that many microorganisms might be involved and thus the word “biological” is more appropriate. We made this change in the revised manuscript (Line 91 in the marked revision).

Figure 1. The heart of the equilibrator is the gas-water mixing unit, and the gas-water separation unit. These should be illustrated much more clearly, as a technical drawing, such that people can build them themselves. The pictures in the SI don't really help very much.

Response: We have replaced Figure S1 with a technical drawing in which the gas-water mixing unit and the gas-water separation unit were both described in detail.

L150-154. This text could go to the figure caption.

Response: This text has been removed from the main text and now is part of the figure caption. Please see line 172-176 in the revision.

L155. Use full word “Laboratory”

Response: We changed “Lab” to “Laboratory”. Please see line 177.

L178. The 13 m tubing length is not mentioned in this list, but it's shown in the figures.

Response: Thanks for pointing this out. We added 13 m to the text. Please see line 204.

L193. Mention which these various methods were.

Response: We described explicitly what the methods are in the text. Please see line 218-219.

L204. From what I read, the depth from which peristaltic pumps can pump up water is physically limited to about 14 m. So how come you could pump water from 30 m depth?

Response: Pump head that the pump needs to overcome is related to vertical distance of the pump to water surface only and unrelated to the vertical position of water intake below the water surface. Thus, often there is only < 0.5 m pump head when the FaRAGE is placed in a small boat.

L210. Was the effect of boat speed on equilibration tested? Depending on the type and placement of water intake, bubbles might start to form when driving too fast.

Response: We did not try higher boat speed than 10 km h⁻¹. The driving speed should be chosen according to the spatial resolution that the users would like to have. In our case, 17 m spatial resolution (spatially averaged) was achieved at 5 km h⁻¹ driving speed corresponding to 12 s response time (see line 377-378). Bubbles were not observed at 10 km h⁻¹ speed when the water intake is mounted on the side wall of the boat, 0.5 m below the water surface. Driving too fast is not recommended as it may harm the CTD probe as well.

L229. Please give this correction equation.

Response: The correction equations for CH₄ and CO₂ are shown in Fig. 2a and b, respectively.

L255. No details on statistical methods or tests are given, yet it says “significantly” here. Which test were performed, and what test statistics did they return?

Response: Thanks. Indeed, statistical tests were not performed. We changed this word to “substantially” and gave mean ± standard deviation. Please see line 289-290.

L260. Please rephrase, “while extended response times” is unclear. Fig.2, panel d. Change the right y-axis colour to red (such as in panel c). Also, why is the red point for 13 m tubing length not connected to the other red points, and how come that its response time is so much longer for high-to-low than for low-to-high, and also so much longer than for the 8 m tubing?

Response: Thanks. This is an incomplete sentence. We rephrased the sentence to “A 91.8% equilibration ratio can be achieved by extending the tube length to 13 m while extended response times are expected.”

We have changed the right-handed y-axis color to red in panel d. Please see the new Fig. 3c-d. The red point for the 13 m tube length severely deviated from the well-fitted power function. The sharp increase in response time of high-to-low is a result of increased resistance of the gas-water mixture flow. The instability (abnormal sharp increase in response time) started from 13 m tube length and became unacceptable when the tube length is 18 m. Please see our explanations in line 294-302.

L373. I would be more careful with this statement. You can state that the equilibrator was not negatively affected by high phytoplankton density, but you haven't tested suspended sediment, so it's not sure it would work in e.g. in a turbid river. You can't exclude that for sustained operation in a turbid system, a filter in the water intake might be required.

Response: Thanks. Indeed, so far the device has never been tested particularly in turbid rivers with suspended sediment particles. We confined this statement to lakes without high sediment loads and also pointed out explicitly that a filtration unit for the water intake might be needed in turbid rivers. Please see line 427-433.

L383. Unclear what this sentence means, please rephrase.

Response: Thanks. They are indeed misleading and unnecessary, therefore, we have removed them from the manuscript.

L408. "potentially be" instead of "be potential".

Response: We corrected this in the text. Please see line 466.

Supporting information:

L32. Coupling instead of couple.

Response: We corrected this in the text (line 33).

Fig.S1. The pictures of the syringes don't show much, and don't help those who want to build their own. Use better pictures and include a technical drawing.

Response: We improved these by replacing them with technical drawings.

L120. This is not a complete sentence.

Response: Thanks. We rephrased this sentence. Please see line 127-128.

Fig.S3. Please also show the corresponding depth profiles of CH₄ and CO₂ at this sampling occasion, such that the reader can judge in how far phytoplankton density might have affected measurements.

Response: We added depth profiles of CH₄ and CO₂ to Fig. S3.

Response to Anonymous Referee #2

I have two remarks to be considered and few suggestions

1) I was very interested in the CH₄ profiles and - while impressed by the high vertical resolution they achieve – I understood the authors tested their system in different lakes, covering some “typical” methane concentration range. Indeed they show two profiles: one from mesooligotrophic Lake Stechlin and one from eutrophic Lake Arend – I supposed they choose trophic state as a proxy for GHG content (Beaulieu et al., 2019). Thus I was confused seeing that Lake Arend, that they present as eutrophic, shows especially low methane, even lower than what they show for the meso-oligotrophic (Lake Stechlin). So what is the criterion behind this choice? Why picking an eutrophic lake that has even less methane than the meso-oligotrophic? Why not a typical eutrophic lake with methane building up below the oxycline during summer stratification? (They show a large literature on this issue in the Intro). For calibration the authors limited the upper range of methane to <2 micromoles L⁻¹. While for high concentrations (microM to mM range) the authors suggest dilutions (line 357) but they don't deal with the problem in this paper. I think the range they show is fair enough, but they should clarify this “lower range test”, directly relatable to surface waters but not to littoral methane rich or eutrophic lake bottom waters, AND change the sentence in the abstract “The FaRAGE is capable of continuously measuring dissolved CH₄ concentrations in the nM-to-mM range” as it may be capable of that, but is not shown here.

Response:

Thanks for the reviewer's very detailed comments. We have such CH₄ profiles with a typical anoxic hypolimnion in summer where CH₄ was enriched to sub milli-molar high concentration. We have added these data into the revised manuscript. Please see Fig. 4c-d.

We further added additional laboratory tests for high CH₄ concentration (e.g., 33 μM). The measured concentration is about 245 ppm with the FaRAGE. For the Gas Scouter 4301 we use, linearity can only be guaranteed up to 500 ppm. Adjustment of water-gas mixing ratio is needed for CH₄ concentrations in the mM range. Therefore, we agree with the reviewer's suggestion. We rephrased the measurement range-related statement accordingly.

2) One big step forward the authors claim is the “extraordinary fast response relative to all existing gas equilibration devices” (line 274). However, when their system is used with PICARRO G2132-1 + a desiccant to measure stable isotopes of gas species, it does not get that much faster than Hartmann 2018 “High Spatio-Temporal Dynamics of Methane Production and Emission in Oxic Surface Water” (line 107 of supplementary material). If I understand correctly then, what showed in table S4 is not entirely correct since their system response time when using PICARRO G2132-1 is obtained without any dryer. As far as I know, for using a PICARRO G2132-1 with a moisty flow a dryer is absolutely necessary. How humid is the water entering the analyzer? In case the authors think that a dryer is needed, they should change the table S4 and the statement in line 274 “The comparison between up-to-date previous studies and this study (Table S4) demonstrated the extraordinary fast response relative to all existing gas equilibration devices. A 53 s response time was achieved when the FaRAGE was adapted to the Picarro G2132-i, which is significantly faster than others (171-6744 s).”

Response:

Thanks for suggesting this. We did not have appropriate dryers for testing when the tests were performed. Therefore a dryer made from silicone beads was tested and a 150 s extension in response time was observed. But we are aware that this has been well tested in Webb et al (2016), in which they tested both Drierite and magnesium perchlorate ($Mg(ClO_4)_2$) as dryers. We reproduced their results below. They show both types of dryer have no effect on CH_4 and CO_2 , except for a 1.5 min time delay on CO_2 was caused by using Drierite.

Yes, the water vapor content of the gas sample flow is above 1% and should be dried before entering the PICARRO G2132-i. We are currently using the drying materials as suggested by Webb et al (2016) and they work quite well.

So the numbers in table S4 and the statement in line 274 are all valid. We understand it is not well clarified. We clarified this in the revised version of the manuscript. Please see the revision in line 113-117 in the supporting information.

Reference

Webb, J. R., Maher, D. T., and Santos, I. R.: Automated, in situ measurements of dissolved CO_2 , CH_4 , and $\delta^{13}C$ values using cavity enhanced laser absorption spectrometry: Comparing response times of air-water equilibrators, *Limnol. Oceanogr.: Methods*, 14, 323-337, <https://doi.org/10.1002/lom3.10092>, 2016.

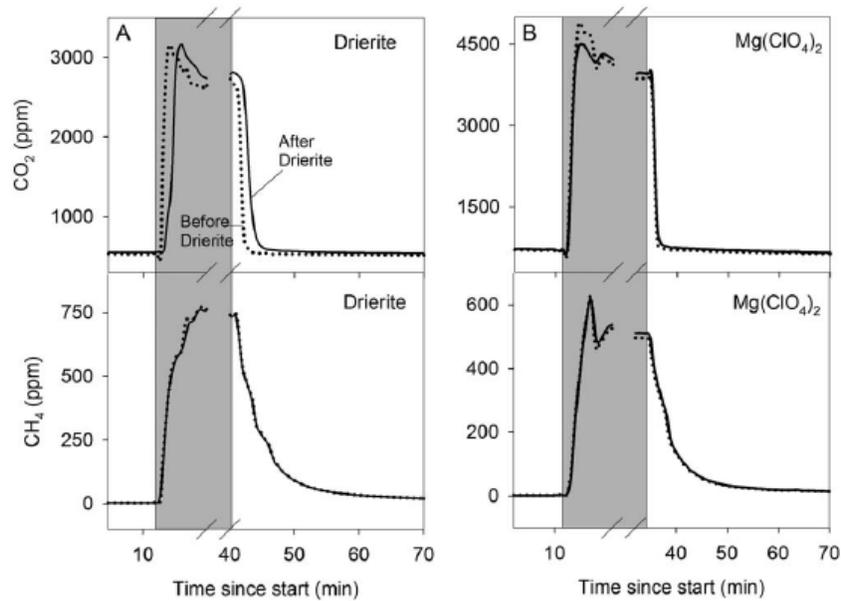


Fig. 2. Step experiments carried out to determine the effect of two desiccants, Drierite (A) and magnesium perchlorate ($\text{Mg}(\text{ClO}_4)_2$) (B), on the response times of CO_2 and CH_4 through a showerhead equilibrator. A time delay of 1.5 min was observed for CO_2 concentrations when Drierite was used as a desiccant. $\text{Mg}(\text{ClO}_4)_2$ had no effect on concentration response times and CH_4 remained unaffected by both desiccants.

Suggestions

When describing the set up the authors often refer to a “bubble unit”, which I suppose in the scheme (Fig.1) is called “gas-water mixing unit”. Consider harmonize.

Response:

Thanks for pointing this out. We checked this carefully and now the terms are used consistently throughout the manuscript.

Line 363- I think it would be better to rephrase the reason why CO_2 is not shown. “for simplicity” for them or for the reader? Maybe they can mention which non-simple problems we meet when applying the system to CO_2 .

Response:

We have added CO_2 as well and incorporated it throughout the full manuscript.

As to line 366 it is possible for the authors to use their system at sea enhancing the liquid to HS ratio to achieve low concentrations.

Response:

Thanks for suggesting this. We replaced one calibration point in Fig. 2, in which 5.5 nM dissolved CH₄ concentration can be well characterized (with the 500 mL min⁻¹ to 1000 mL min⁻¹ water-gas mixing ratio). Indeed, lower concentrations in the ocean system can be also well measured by increasing water-gas mixing ratio.

I would recommend to make sure that the scientific community that works on GHG air-water exchange in oceans gets interested too (add to abstract and line 97?). For different reasons from the ones highlighted here for inland waters (to name one the massive lack of ground data to calibrate satellite inferences) this system could be applied to voluntary observing ship programs to map CO₂ and CH₄ surface concentrations. In case the authors find a major obstacle to this it would be good to mention – making a suggestion for adapting their system for oceanographic applications.

Response:

Thanks for suggesting these. We agree that the FaRAGE can be a good method for studying GHGs from oceans. We added a sentence in abstract (line 35-36) and also in text (line 117-118).

Line 382- they mention how temperature should be corrected for the change along the hose- can give an example on how off can it get and does that mean to always measure temperature in situ at depth along with the profile?

Response:

We included an example in SI (Fig. S4) to show how much the water temperature (water flow inside the gas-water separation unit) can differ from the in situ water temperature. It's more important to monitor temperature of water flow inside the gas-water separation unit of the device. We recommend to installing a temperature logger in the device if someone would like to rebuild the device, but can only afford one thermometer.

Line 400 replace "filthy"

Response:

Thanks. We rephrased this sentence (see line 458).

Line 410- wouldn't a scrubber serve for that? Would that slow down the system RT?

Response:

Thanks. Yes, a copper scrubber can help to remove H₂S from the gas samples (Malowany et al. 2015). According to Malowany et al. (2015), no time delay was observed when a copper scrubber was used. We also reproduced the figure they included in the publication. We added this reference to the revised manuscript (line 469-471).

Reference

Malowany, K., Stix, J., Van Pelt, A., and Lucic, G.: H₂S interference on CO₂ isotopic measurements using a Picarro G1101-i cavity ring-down spectrometer, *Atmos. Meas. Tech.*, 8, 4075–4082, <https://doi.org/10.5194/amt-8-4075-2015>, 2015.

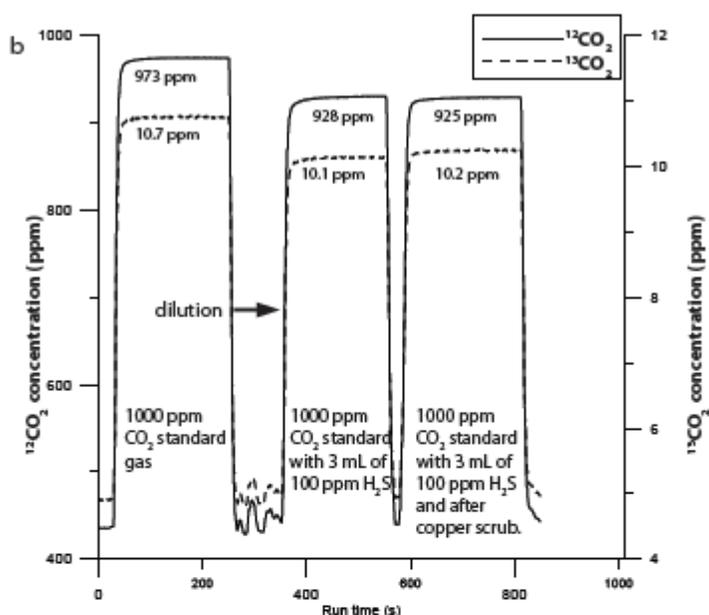


Figure 3. Change in the ¹²CO₂ and ¹³CO₂ concentrations with addition of H₂S to the standard gas. (a) Plot showing the percentage change in CO₂ concentration between gas with H₂S and gas scrubbed of H₂S. There is a visible increase in the ¹²CO₂ concentration and a decrease in the ¹³CO₂ concentration with addition of H₂S. The percentage decrease for ¹³CO₂ is significantly greater than the percentage increase for ¹²CO₂. (b) Plot showing the 1000 ppm standard CO₂ gas with the addition of 3 mL of 100 ppm H₂S and the subsequent response after the H₂S was removed with the copper scrub. There is a small, yet visible, increase in the ¹³CO₂ concentration and decrease in the ¹²CO₂ concentration when H₂S is removed.

References

Beaulieu, J.J., DelSontro, T. & Downing, J.A. Eutrophication will increase methane emissions from lakes and impoundments during the 21st century. *Nat Commun* 10, 1375 (2019). <https://doi.org/10.1038/s41467-019-09100-5>

1 **A Fast-response automated gas equilibrator (FaRAGE) for continuous *in situ***
2 **measurement of ~~methane~~ CH₄ and CO₂ dissolved in water**

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12 **Abstract**

13 Biogenic greenhouse gas methane (CH₄)-emissions, e.g. of methane (CH₄) and carbon
14 dioxide (CO₂) from inland waters contribute substantially to global warming. In aquatic systems,
15 dissolved CH₄ in freshwater lakes and reservoirs is greenhouse gases are highly heterogeneous
16 both in space and time. To better understand the biological and physical processes that affect
17 sources and sinks of both CH₄ and CO₂ in lakes and reservoirs, itstheir dissolved CH₄
18 concentrations needs to be measured with a-highest spatial and temporal resolution. To achieve
19 this goal, we developed the **Fast-Response Automated Gas Equilibrator (FaRAGE)** for real-
20 time *in situ* measurement of dissolved CH₄ and CO₂ concentration_s at the water surface and in
21 the water column. FaRAGE can achieve an exceptionally short response time ($t_{95\%} = 12$ s when
22 including the response time of the gas analyzer) while retaining an equilibration ratio of 6362.6%
23 and a measurement accuracy of 0.5% for CH₄. A similar performance was observed for

24 dissolved CO₂ (t_{95%} = 10 s, equilibration ratio 67.1%). An equilibration ratio as high as 91.8%
25 can be reached at the cost of a slightly increased response time (16 s). The FaRAGE is capable
26 of continuously measuring dissolved CO₂ and CH₄ concentrations in the nM-to-sub mM (10⁻⁹ -
27 10⁻³ mol L⁻¹) range with a detection limit of sub-nM (10⁻¹⁰ mol L⁻¹), when coupled with a
28 cavity ring-down greenhouse gas analyzer (Picarro GasScouter). ~~It enables the possibility of~~
29 ~~mapping dissolved CH₄ concentration in a “quasi” three-dimensional manner in lakes.~~
30 ~~Additional tests demonstrated a~~ Similarly good performance of FaRAGE ~~the equilibrator could~~
31 ~~be demonstrated for measuring dissolved CO₂.~~ FaRAGE enables allows for the possibility of
32 mapping dissolved concentration in a “quasi” three-dimensional manner in lakes and ~~The~~
33 ~~FaRAGE~~ provides an inexpensive alternative ~~to other commercial gas equilibrators.~~ It is
34 simple to operate and suitable for continuous monitoring with a strong tolerance to suspended
35 particles. While the FaRAGE is developed for inland waters, it can be also applied to ocean
36 waters by slightly tuning the gas-water mixing ratio. ~~The FaRAGE is e~~ easy adaptability ~~easily~~
37 ~~adapted to suit to~~ other gas analyzers ~~such as Ultra-portable Los Gatos and stable isotopic gas~~
38 ~~analyzer (Picarro G2132 i) also provides the potential for many further~~ expanding the range of
39 potential applications, including nitrous oxide and isotopic composition of the gases. ~~e.g.~~
40 ~~measuring dissolved ¹³C-CH₄ and ¹³C-CO₂ and CO₂.~~

41

42 1 Introduction

43 Despite the well-established perception of inland waters as a substantial source of
44 atmospheric methane (CH₄) and carbon dioxide (CO₂) (Bastviken et al., 2011; Cole et al., 2007;
45 Tranvik et al., 2009), the magnitude of these greenhouse gases remains large uncertainties
46 uncertain owing to the fact that poorly constrained some key processes affecting CH₄ (e.g.
47 bubbling) and CO₂ budget are still poorly constrained sources and sinks (Saunois et al., 2019).
48 Most freshwater lakes and reservoirs are often oversaturated with CH₄ and CO₂ (relative to
49 atmosphere) and their distribution are characterized by high spatio-temporal heterogeneity
50 (Hofmann, 2013). Point-based and short-term measurements can result in biases in estimating
51 diffusive CH₄ flux (Paranaíba et al., 2018). Thus, resolving the spatio-temporal dynamics of
52 both dissolved CH₄ and CO₂ concentration in lake water is a prerequisite for a better budgeting
53 understanding of sources and sinks, production and loss processes of these gases in freshwater
54 lakes.

55 The distribution of CH₄ and CO₂ in lakes is often characterized by pronounced vertical
56 and horizontal concentration gradients, which can occur either below or above often coincides
57 with the position of the thermocline. In many deep stratified lakes, a sharp vertical gradient of
58 CH₄, for instance, below the thermocline can develop in the anoxic hypolimnion (mM range)
59 (Encinas Fernández et al., 2014; Liu et al., 1996). In contrast, in some stratified lakes with a
60 fully oxygenated hypolimnion CH₄ can accumulate above the thermocline (~µM range)
61 (Grossart et al., 2011; Donis et al., 2017; Günthel et al., 2019). In addition to formation
62 processes that lead to CH₄ accumulation, the concentration of dissolved CH₄ is also regulated
63 by losses due to oxidation and emission to the atmosphere (Bastviken et al., 2004; Juutinen et
64 al., 2009). Both rates can be Particularly the e Emission rates, in particular, are highly variable,
65 particularly for the flux term as which is strongly affected it depends dependent by-on turbulence
66 generated induced by wind and-or convective mixing (Read et al., 2012; Vachon and Prairie,

2013). ~~In addition to the uneven v~~Vertical ~~CH₄~~ distributions of CH₄ and CO₂, ~~can~~ be further confounded the contribution of littoral sediments which can result in distinct horizontal gradients of CO₂ and CH₄. ~~apparent horizontal gradients have been observed in lakes where littoral sediments are identified as a CH₄ source~~ (Murase et al., 2003). ~~This~~ Accounting for horizontal CH₄ gradients ~~can also~~ is therefore critical as lateral transport may account for ~~contribute to the~~ a proportion of the epilimnetic CH₄ peak ~~observed~~ in pelagic waters via ~~lateral transport~~ (Hofmann et al., 2010; Fernández et al., 2016; Murase et al., 2005; Peeters et al., 2019).

~~These spatial~~Spatial ~~pattern~~distributions of CH₄ and CO₂ in aquatic systems, ~~however,~~ are subject to temporal variation in response to ~~vary~~ over time, particularly as factors which control their production, consumption and loss to the atmosphere fluctuate. ~~variable drivers as well as water motion.~~ Nevertheless, ~~dissolved CH₄ in lake water is not only featured with variable spatial patterns, it also changes at different time scales as most processes that contribute to the spatial heterogeneity are not always synchronized.~~The rise and fall of ~~ILake~~ CH₄ concentration ~~Concentrations of CH₄ and CO₂ in lakes demonstrate often show strong~~ profound seasonality, ~~that are is~~ driven primarily by thermal stratification (Encinas Fernández et al., 2014) and phytoplankton dynamics (Günthel et al., 2019). While the build-up of hypolimnetic CH₄ storage is a slow process that is closely related to the development of lake hypoxia, ~~the~~ epilimnetic CH₄ ~~and CO₂ maximum~~ can be highly variable even at a daily basis as ~~it is~~ they are strongly affected by phytoplankton dynamics (Günthel et al., 2019; Hartmann et al., 2020; Bižić et al., 2020). In addition, storms can act as another driver for short-term ~~CH₄~~ dissolved gas dynamics in the lake because ~~it they~~ often ~~leads~~ contribute to higher evasion rates caused by strong vertical turbulent mixing (Zimmermann et al., 2019) and enhanced horizontal transport (Fernández et al., 2016). While the seasonal patterns of dissolved CH₄ ~~and CO₂~~ concentration in lake water seem recurrent and can be simulated (Stepanenko et al., 2016), the unpredictable

91 effects of short-term ~~phytoplankton-biological~~ dynamics and storm events can present a
92 challenge in modeling ~~lake CH₄-the~~ dynamics of greenhouse gases in lakes.

93 While there is an urgent need for resolving the spatio-temporal ~~variabilities~~variability
94 of CH₄ in large water bodies (e.g. ~~deep, stratified~~ lakes), we recognize limitations in the
95 available methodology. Like most gases in dissolved phase, CH₄ and CO₂ cannot be measured
96 directly in water. Instead, a carrier gas (synthetic air or at air concentration) is added to achieve
97 (full/partial) gas-water equilibration. The headspace gas sample is then measured with a gas
98 spectrometer and the concentration of targeted gas can be calculated according to Henry's law
99 (Magen et al., 2014). To save sampling effort, continuous gas equilibration devices have been
100 developed, which generally can be classified ~~to~~in four categories: 1) Membrane type (Schlüter
101 and Gentz, 2008; Boulart et al., 2010; Gonzalez-Valencia et al., 2014; Hartmann et al., 2018) -
102 gases are extracted from water ~~by~~ using a gas-permeable membrane; 2) Marble type
103 (Frankignoulle et al., 2001; Santos et al., 2012) - gas exchange is enhanced by pumping water
104 through marbles that increases the gas-water contact area; 3) Bubble type (Schneider et al.,
105 1992; Körtzinger et al., 1996; Gülzow et al., 2011) - dissolved gases are stripped out by
106 bubbling the water sample; 4) Showerhead type (Weiss-type) (Johnson, 1999; Rhee et al., 2009;
107 Li et al., 2015) - water is pumped from top and then mixed with a circulated headspace carrier
108 gas. A full evaluation on the performance of these devices was provided in a recent review
109 (Webb et al., 2016), where, the most important parameter, response time, was found to vary
110 between 2-34 min for dissolved CH₄. While it is already encouraging, improvements are
111 expected to further shorten the response time.

112 Driven by the need to resolve temporal and spatial ~~variabilities~~variability of dissolved
113 CH₄ and CO₂ in inland waters~~lakes or /reservoirs~~ with sufficient precision, we developed a
114 novel, low-cost equilibrator to achieve fast gas-water equilibration. The **Fast-Response**
115 **Automated Gas Equilibrator (FaRAGE)** can be coupled with a portable gas analyzer, which

116 makes it perfect for field use. Here, the performance of the FaRAGE is evaluated by
117 investigating its response time, detection limit and equilibration ratio. Although FaRAGE is
118 developed for inland waters, it can be also adapted for oceanographic applications. Applications
119 are provided exemplarily to demonstrate the potential of the FaRAGE for improving our
120 understanding on the spatial distribution and temporal dynamics of dissolved CH₄ and CO₂ in
121 inland waters.

122 **2 Materials and Methods**

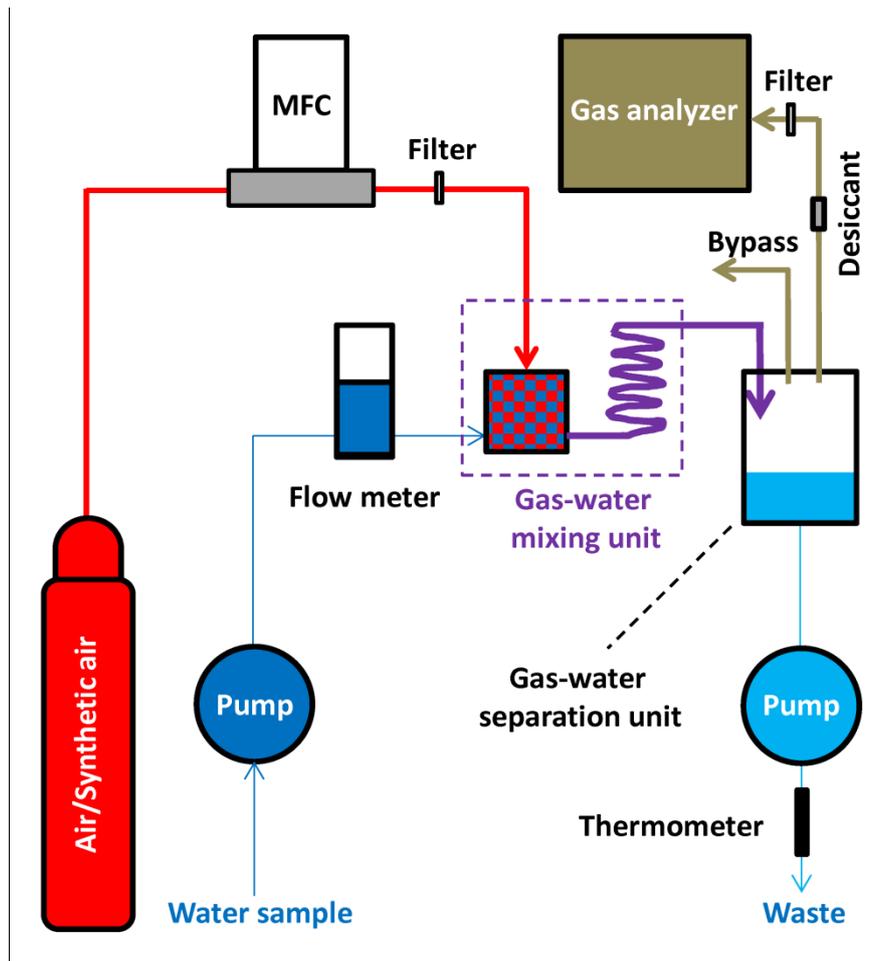
123 **2.1 Device description**

124 The design of the FaRAGE is modified from two types of equilibrators: Bubble-type
125 (Schneider et al., 1992) and Weiss-type (Johnson, 1999). In contrast to the traditional bubble-
126 type and Weiss-type equilibrators that create a large-volume headspace and circulates air back
127 to the headspace, the FaRAGE is a flow-through system that adds gas flow into a constant water
128 flow to produce a minimal headspace for continuous concentration measurement of CO₂ and
129 CH₄ dissolved in water.

130 The operation principle of the FaRAGE is depicted in Fig. 1 and ~~photos-technical~~
131 drawings of the main parts of the prototype are provided in Fig. S1. A list of information on
132 suppliers and cost of each part can be found in Table S1. A mass flow controller (SIERRA
133 C50L, Netherlands) is used to generate a constant carrier gas (normal air/synthetic air) flow (1
134 L min⁻¹) from a compressed air tank coupled with a pressure regulator. Water samples are taken
135 continuously using a peristaltic pump (500 mL min⁻¹), and the flow is monitored using a flow
136 meter (Brooks Instrument, Germany). The two flows mix in a gas-water mixing unit ~~that is~~
137 ~~composed of a gas bubble generating unit~~ and then travel through a coiled hose for further gas-
138 water turbulent mixing. In the gas-water mixing unit ~~bubble unit~~ (modified from a 10 mL plastic
139 syringe), a jet flow is created by adapting narrowed tubing (2 mm inner diameter) to the water

140 pumping hose (3.2 mm inner diameter). Degassing occurs when the jet flow enters the chamber
141 with a sudden enlarged diameter (14 mm). Degassing is further enhanced by micro-bubbles that
142 are generated by a bubble diffuser attached to the carrier gas hose (inside the ~~bubble unit~~plastic
143 syringe). The gas-water mixture flows through the 2-m long Tygon tube (3.2 mm inner diameter)
144 where additional equilibration occurs. The flow is finally introduced to a gas-water separation
145 unit (a 30 mL plastic syringe) where the headspace gas is separated from the water. In this
146 chamber, water falls down freely to the bottom while the headspace gas is taken directly to a
147 greenhouse gas analyzer (1 L min⁻¹ gas pumping rate; GasScouter G4301, Picarro, USA). A 2-
148 m long Tygon tube (3.2 mm inner diameter) is attached to the top of the chamber for venting
149 excess gas flow while stabilizing gas pressure in the headspace. The bottom water is discharged
150 back to the lake using another peristaltic pump (500 mL min⁻¹). To protect the gas analyzer
151 from damaging high water vapor content, a Teflon membrane filter (pore size 0.2 μm) is placed
152 before the gas intake (resulting in a ~210 mL min⁻¹ reduction in flow rate of gas sample, which
153 is vented from the bypass at the top of the gas separation unit). A desiccant (a 20 mL plastic
154 syringe filled with dried silicone beads) is used to reduce moisture concentration ~~to < 0.1%~~
155 when attaching to a Picarro G2132-i isotope analyzer (Picarro, USA), in which < 1% moisture
156 level is required for ¹³δC-CH₄ measurement. The temperature of the water sample at the point
157 of equilibration with the headspace gas is monitored using a fast thermometer (precision
158 0.001 °C, 1 Hz, TR-1050, RBR, Canada) attached to the end of the water discharging hose.

159 ~~As concerns might arise from the availability of gas analyzer coupled to the FaRAGE,~~
160 In addition to Gas Scouter from Picarro, two additional widely used models of greenhouse gas
161 analyzers were tested. They are the Ultraportable Los Gatos (Los Gatos Research, USA) and
162 stable isotopic CH₄ analyzer (G2132-i, Picarro, USA). The main technical details of all three
163 tested gas analyzers are listed in Table S2.



164
 165 **Fig. 1** Schematic design of the FaRAGE. The components include: Air tank containing
 166 compressed carrier gas (air or synthetic air) with a pressure regulator, a mass flow controller
 167 (MFC) for generating constant carrier gas flow, two peristaltic pumps for taking and
 168 discharging water, respectively, a flow meter for monitoring water sample flow, a gas-water
 169 mixing unit, a gas-water separation unit, a gas analyzer, and a thermometer for measuring water
 170 temperature at phase equilibration. A Teflon membrane filter is placed after the MFC and
 171 another is added before the gas analyzer to protect from being flooded. A desiccant is used to
 172 dry the gas flowing to the gas analyzer (if Picarro isotopic analyzer is used). The red color
 173 marks the flow of carrier gas, dark blue line indicates the water sample, purple line shows the
 174 flow of gas-water mixture, the light brown line shows the flow of gas sample (after partial
 175 equilibration) and the light blue line depicts the water discharged back to lake. The thickness
 176 of the lines scales with the gas/water flow rates. The arrows show the flow directions.

2.2 Laboratory validation

The FaRAGE prototype was first tested intensively in the laboratory to determine both the equilibration ratio and response time. The tests were performed for both CH₄ and CO₂ as with a GasScouter G4301 (Picarro, USA), which measures both gases simultaneously. For simplicity, only the results for one gas (CH₄) are explained in the following text. The equilibration ratio is defined as the percentage concentration of the gaseous CH₄ gas concentrations at the outlet of the gas equilibrator in comparison to the equilibrium concentration (full gas-water equilibration). The equilibration ratio was established by across measuring a range of CH₄-stock solutions (nano-to-milli molar dissolved gas concentrations). These standard solutions were prepared by adding different amounts of either CH₄ or CO₂ into a 200 mL headspace of a 2 L Schott bottle filled with Milli-Q water. The exact dissolved CH₄ concentrations in these solutions were tested with the ~~traditional~~ manual headspace method: a 400 mL headspace was created in a 500 mL plastic syringe with nitrogen gas. The CH₄-gas concentration of the headspace gas was then measured using GasScouter G4301 ~~(Picarro, USA)~~. At the same time, CH₄-dissolved CH₄ and CO₂ concentrations of these standard solutions were measured with the FaRAGE for at least 2 min and an average was calculated from more than 60 individual data points. We directly compared dissolved CH₄-gas concentrations measured using the two different methods, i.e., our equilibrator and manual headspace method.

The response time of the device was investigated by switching the water sample inlet between two water samples with different CH₄-concentrations of either CH₄ or CO₂. Triplicated measurements were performed. An exponential fit was applied to the concentration change curve and the response time was determined as time needed to reach 95% of the final concentration.

The effect of water-to-gas mixing ratio on equilibration ratio and response time of the device was investigated. By fixing the carrier gas flow rate to 1 L min⁻¹, the water-to-gas mixing

202 ratio was varied from 0.04, 0.08, 0.12, 0.15, 0.24, 0.29, 0.36, 0.43 and 0.5 by adjusting the water
203 sample flow rate. The effect of tube length on performance of the device was also examined by
204 adapting 1, 2, 4.4, ~~and 8.4~~ and 13 m Tygon tube onto the gas-water mixing unit. For all these
205 tests, triplicated measurements of the equilibration ratio and response time were performed
206 corresponding to different mixing ratios and the mean values were used for analysis.

207 Tests were performed to investigate the performance of the device when adapting to two
208 other types of gas analyzers. As the equilibration ratio is unaffected by the model of gas
209 analyzers, only response time was determined. This was done by fixing carrier gas and water
210 sample flow rates to 1 and 0.5 L min⁻¹, respectively. The surplus gas was vented to the air as
211 Ultraportable Los Gatos and Picarro G2132-i have a gas intake flow rate of only 500 and 25
212 mL min⁻¹, respectively. The effect of desiccant on response time of Picarro G2132-i was
213 checked by measuring gas samples with and without a desiccant installed.

214 2.3 Field tests

215 ~~Two~~ Four lakes in Germany were chosen for field tests. Lake Stechlin is a deep meso-
216 oligotrophic lake with a maximum depth of 68 m and Lake Arend is a eutrophic lake with a
217 maximum depth of 48 m. Pronounced CH₄ peaks in the epilimnion of Lake Stechlin have been
218 previously reported that were measured with ~~various~~ two different methods (manual headspace
219 method in Grossart et al. (2011) and Tang et al. (2014); membrane-based gas equilibrators in
220 Hartmann et al. (2018)). This makes it ideal for our testing purpose. While CH₄ profiles at Lake
221 Arend have never been reported, the metalimnetic oxygen minimum in the lake observed during
222 summer (Kreling et al., 2017) renders it interesting for CH₄ profiling throughout the entire water
223 column. ~~Another two~~ Additionally, we selected both eutrophic lakes with an anoxic
224 hypolimnion lakes (Lake Großer Pälitz and Lake Zotzen), where CH₄ and CO₂ can accumulate
225 are both eutrophic with an anoxic hypolimnion during the period of thermal stratification.

226 Measurements were conducted in these two lakes to test the capability of FaRAGE to measure
227 anoxic hypolimnetic water with high dissolved CH₄ and CO₂ accumulated concentrations.

228 Due to the high potential of the FaRAGE for real-time *in situ* measurement of dissolved
229 CH₄ and CO₂ concentrations, we explored potential field applications. These field tests included
230 depth profiling of dissolved CH₄ concentrations in ~~Lake Arend and Lake Stechlin~~ the four lakes
231 and investigations of the horizontal distribution of surface dissolved CH₄ and CO₂
232 concentrations across the entire Lake Stechlin. For the first application, a fast-response CTD
233 (conductivity, temperature and depth) profiler (XR-620 CTD+, RBR, Canada) was mounted
234 onto a winch with a 30 m long water hose (4 mm inner diameter) attached. The CTD profiler
235 with hose was lowered down continuously at a constant speed (1 m min⁻¹). The exact depth and
236 temperature of sampled water can be extracted from the CTD profiler by correcting for the
237 travel time of water sample flow in hose. For the spatial mapping, a GPS antenna (Taoglas,
238 AA.162, USA) was attached to the Picarro gas analyzer. The water intake was submerged 0.5
239 m below the water surface together with the CTD profiler and fixed to one side of the boat. The
240 boat was driven at a constant speed of 5 km h⁻¹.

241 **2.4 Theoretical background and data processing**

242 The FaRAGE shares a similar working principle to the Weiss-type gas equilibrator
243 described by Johnson (1999). The theoretical background and equations are provided in S3.

244 A simplified calculation is described by referring to the manual headspace method. In
245 principle the gas-water mixture is analogous to the static headspace method with the final CH₄
246 gas concentration in the gas phase assumed to reach a full equilibrium with that dissolved in the
247 aqueous phase. Therefore, by specifying the mixing ratio of air and water, the total mass of CH₄
248 for instance, can be calculated by summing up the CH₄ in the headspace with the dissolved CH₄
249 (at equilibrium according to Henry's law, which is temperature and pressure dependent) in the

250 aqueous phase and subtracting the mass of background CH₄ (from the carrying gas with known
251 concentration). The dissolved CH₄ gas concentration is then expressed as the volumetric
252 concentration of the total net mass of either CH₄ or CO₂ in the dissolved phase in the given
253 sample volume. A separated exemplary calculation sheet (excel file S5) is provided, which
254 allows for correction for temperature and pressure change (Goldenfum, 2010).

255 As the equilibration is only partially reached (< 92%), a correction coefficient is needed.
256 This can be obtained by measuring the water samples with known concentrations across a large
257 gradient. By referring to the results measured with the manual headspace method assuming full
258 equilibration (Magen et al., 2014), an equation for precise correction of the measured CH₄
259 dissolved gas (alsoand CO₂) concentrations can be obtained.

260 **3 Results and Discussion**

261 **3.1 Detection limit, equilibration ratio and response time**

262 The FaRAGE is capable of achieving a high gas equilibration ratio. We observed a high
263 correlation ($R^2 = \underline{0.9991.000}$, $p < 0.01$) between the concentrations obtained using the
264 ~~traditional~~ headspace method and those measured using the FaRAGE (Fig. 2a) across a wide
265 range of dissolved CH₄ and CO₂ concentrations. The measurement accuracy is 0.5% (standard
266 deviation in relation to final concentration) once a stable plateau was reached (Fig. ~~2b~~2c). For
267 CH₄, the FaRAGE reaches a high equilibration ratio (~~63~~62.6%) and ensures a rapid response.
268 The determined response time $t_{95\%}$ is only 12 ± 1 s when switching from low-to-high (nano-to-
269 sub micro molar) dissolved CH₄ concentrations while the $t_{95\%}$ is a little longer (15 ± 2 s) when
270 switching from high-to-low concentration (Fig. ~~2b~~2c). For the current design specifications that
271 allow for a high equilibration ratio, the detection is theoretically limited by the sensitivity of
272 the coupled gas analyzer. In the lab tests, a clear response was observed at least for CH₄
273 concentration at air saturation (~~16.95.5~~ nM inside the lab building). The measurable CH₄

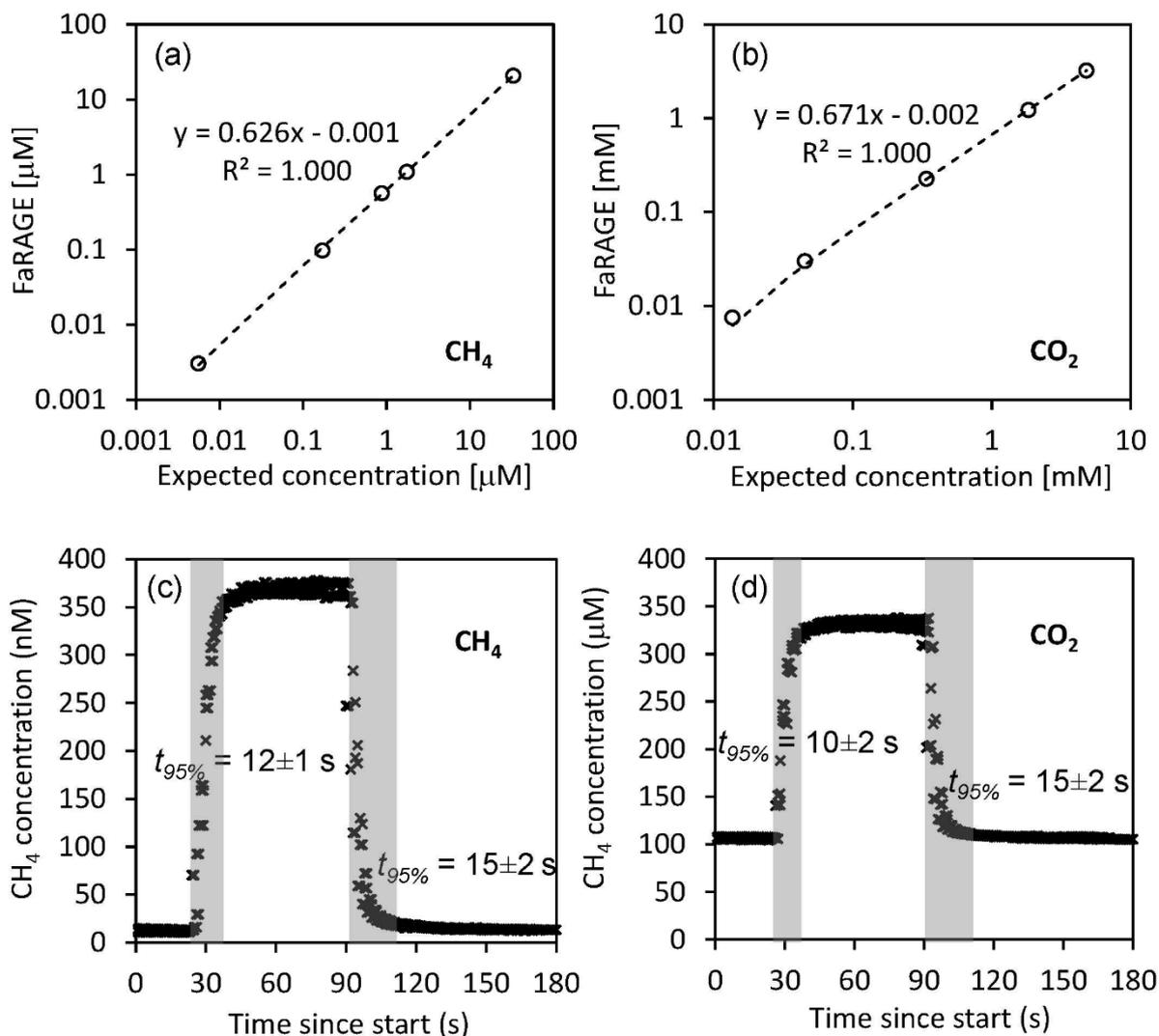
274 concentrations should be at least sub-nM (10^{-10} mol L⁻¹) given the high performance of cavity-
275 ring-down gas analyzers. This is more than sufficient for applications in inland waters where
276 dissolved CH₄ concentrations are often above air saturation. Despite CO₂ (Weiss, R. F., 1974)
277 is an order of magnitude more soluble in water than CH₄ (Wiesenburg and Guinasso, 1979),
278 similar performances of the FaRAGE were observed when measuring dissolved CO₂. An
279 equilibration ratio of 67.1% (Fig. 2b) was achieved with a fast response (Fig. 2d; $t_{95\%} = 10 \pm 2$
280 and 15 ± 2 for low-to-high and high-to-low, respectively) when a 2 m mixing tube was used.

281 The response time for the FaRAGE results from two components: 1) the response of the
282 gas analyzer to changes in gas concentration and 2) the physical gas-water exchange process.
283 The response time for the gas analyzer is 5 s when the CH₄ concentration increases (Fig. S2).
284 The FaRAGE itself needs < 10 s to reach 95% of the final steady-state concentration.

285 Equilibration ratio and response time of the FaRAGE is not sensitive to the water-to-gas
286 mixing ratio (Fig. 2e3a) but rather to the length of the tube in-installedattached after the gas-
287 water mixing unit (Fig. 2d3c). Little-A small effect, of the increased water-to-gas mixing ratio
288 was also observed on the equilibration ratio in response to the increase of water to gas mixing
289 ratio. Also, tThe increase of water-to-gas mixing ratio did not significantly-substantially
290 change the response time of the device (on average 9.5 ± 1.5 s for low-to-high and 13.9 ± 2.4 s
291 for high-to-low, respectively). This is in contrast to other types of equilibrators in which an
292 increase of water-to-gas mixing ratio was found to result in a faster response (Webb et al., 2016).
293 However, a sharp enhancement of equilibration ratio was observed due to the extended length
294 of the tube for the gas-water mixing unit. A 91.8% equilibration ratio can be achieved by
295 extending the tube length to 13 m while extended response times are expected (low-to-high 17
296 s and high-to-low 47.5 s, respectively). The instabilityIncreases in response time were notable
297 when the tube-length exceeded 13 m and were considered excessive at a tube length of 18 m
298 (Fig 3c-d).-(abnormal sharp increase in response time) started from at the 13 m tube length (Fig-

299 ~~3c d) and became unacceptable when tube length is reached 18 m.~~ Further enhancement of the
300 equilibration ratio was thus not possible when a longer tube (e.g. 18 m) was used. The gas flow
301 rate cannot be stabilized at 1 L min⁻¹ due to the increased resistance in response to the further
302 extension of tube length. Equilibration ratio and response time were affected by the length of
303 the tube after the gas-water mixing in a similar way as it was for CH₄ (Fig. 3b, d) with only one
304 exception in the response time when the dissolved CO₂ concentration changed from high to low.
305 The response time increased linearly ($R^2 = 0.910, p < 0.01$) from 11 s to 18 s in response to the
306 increase of water-to-gas ratio from 0.04 to 0.5.

307 As shown in Table S2 and Fig. S2, the fast response of the FaRAGE is partly due to the
308 extremely fast response of the Picarro Gas Scouter. ~~This makes it unfair to compare with other~~
309 ~~equilibrators in which different gas analyzers were used.~~ Tests were performed by adapting the
310 FaRAGE to two other greenhouse gas analyzers (Ultraportable Los Gatos and Picarro G2132-
311 i) and the response times are listed in Table S3. Comparisons were made in Webb et al. (2016)
312 and Hartmann et al. (2018) where both CH₄ and ¹³C-CH₄ were measured using a Picarro
313 G2201-i (Picarro, USA). Here we used a similar Picarro stable isotopic gas analyzer (Picarro
314 G2132-i) and unified all previous reported response times τ to $t_{95\%}$ by applying the equation $t_{95\%}$
315 $= 3\tau$. The comparison between up-to-date previous studies and this study (Table S4)
316 demonstrated the extraordinary fast response relative to all existing gas equilibration devices.
317 A 53 s response time was achieved when the FaRAGE was adapted to the Picarro G2132-i,
318 which is significantly substantially faster than ~~others previously reported~~ (171-6744 s).



320

321 **Fig. 2** Performance of the Fast-Response Automated Gas Equilibrator (FaRAGE with a 2-m
 322 tube in the gas-water mixing unit) for both dissolved CH₄ and CO₂. (a)-(b) Correction equations
 323 for dissolved CH₄ and CO₂, respectively by referring FaRAGE measurements to expected
 324 concentrations measured using the manual headspace method. The dashed lines show a linear
 325 fit and the equations are shown next to the lines. Note that in the two graphs both axes are log
 326 transformed. (c)-(d) Exemplary response time of FaRAGE for low-to-high and high-to-low
 327 concentration changes (water-to-gas mixing ratio 0.5). Triplicated tests were performed and the
 328 average response time was taken at the time point when 95% of the final concentration was
 329 reached.

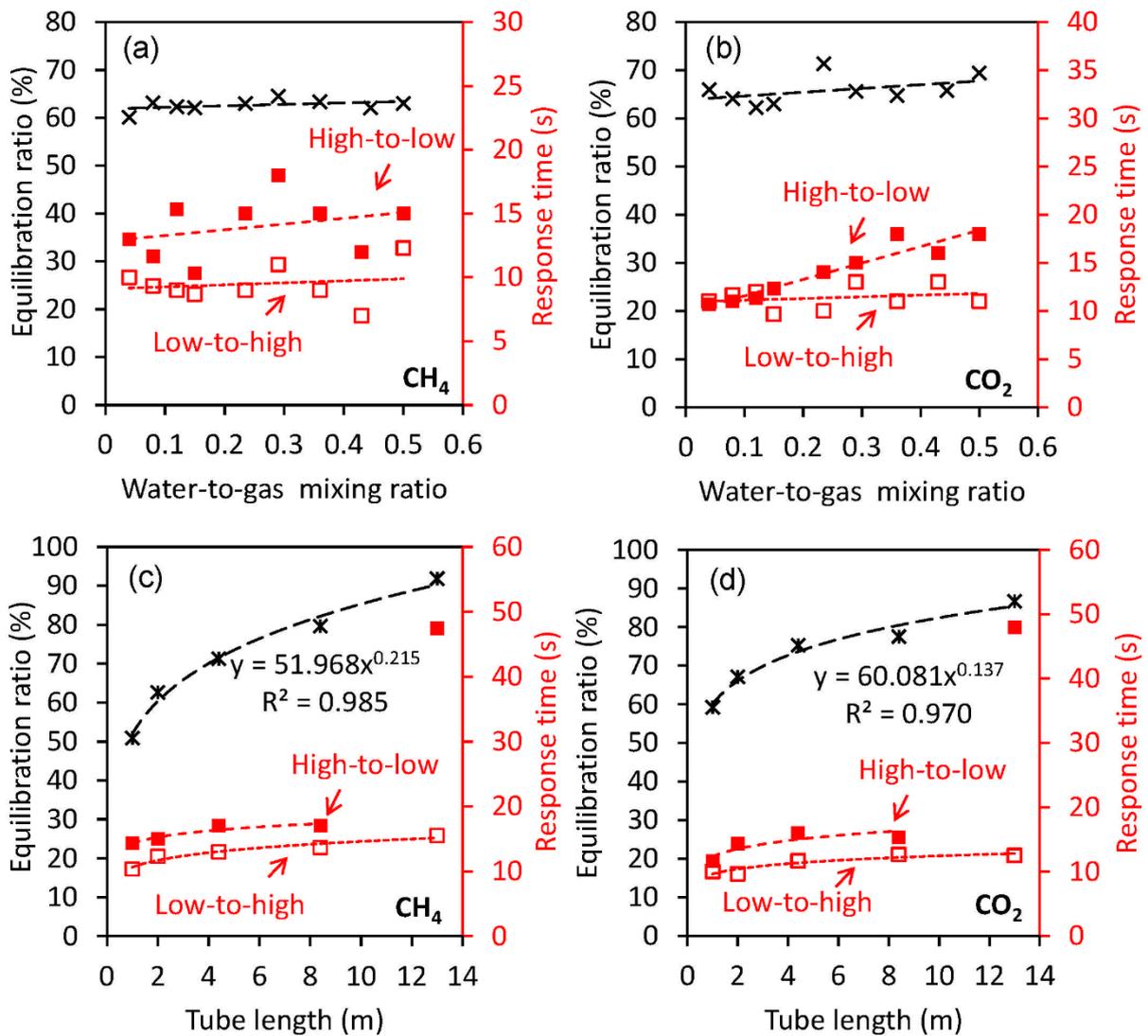


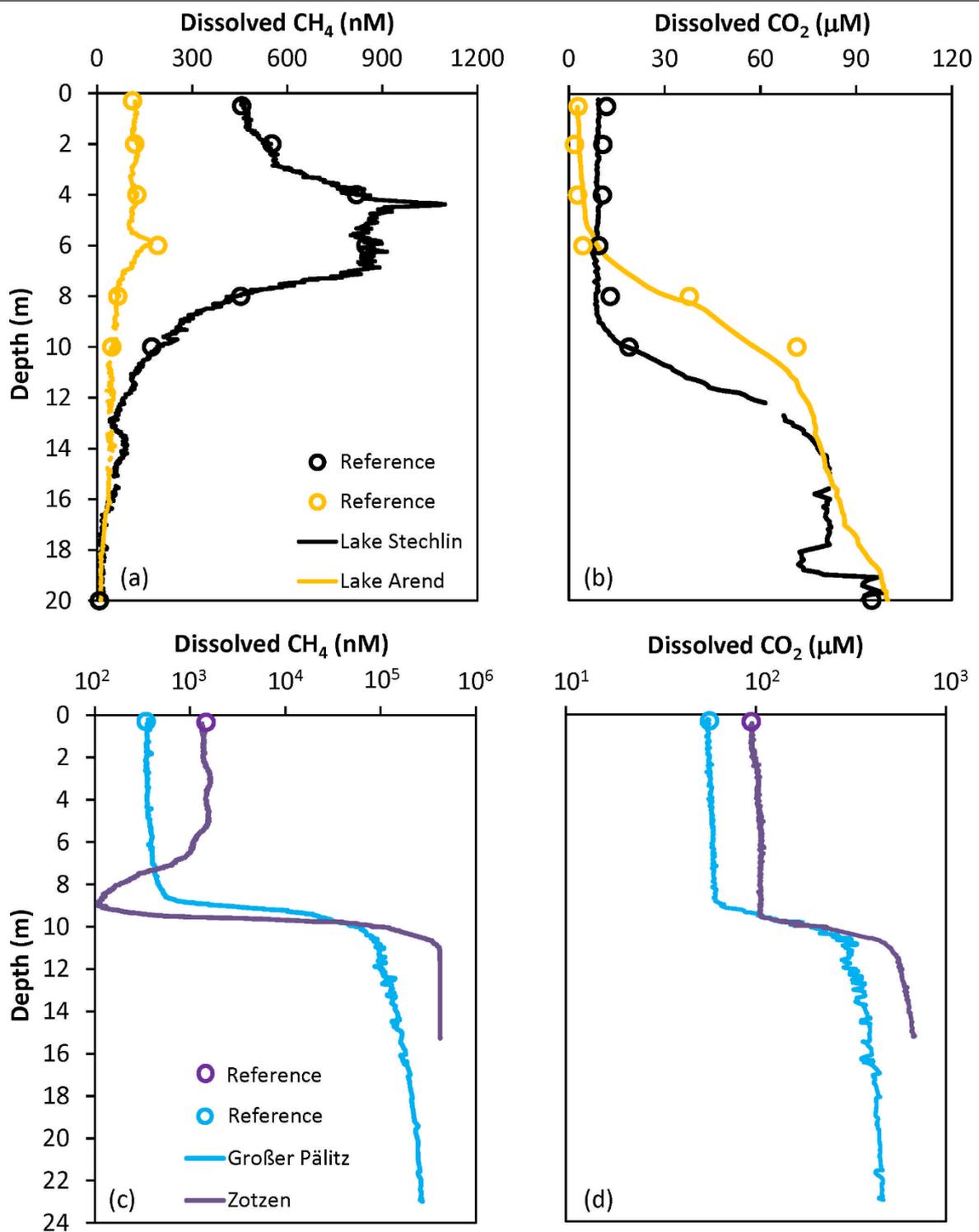
Fig. 3 Factors affecting performance of the gas equilibrators for both dissolved CH₄ and CO₂.

(a)-(b) Equilibration ratio and response time in response to changing water/gas mixing ratio (with a 2-m tube in the gas-water mixing unit). Black cross symbols are equilibration ratios, and low-to-high and high-to-low response times are represented by red open and solid squares, respectively. (c)-(d) Equilibration ratio and response time in response to changing tube length of gas-water mixing unit (with a fixed water-to-gas mixing ratio of 0.5). Black cross symbols are equilibration ratios, and low-to-high and high-to-low response times are represented by red open and solid squares, respectively.

3.2 Depth profiles of dissolved CH₄ and CO₂ from multiple lakes

340 Good agreement was observed between depths profiles of dissolved CH₄ and CO₂
341 concentration measured using ~~two different methods~~ the FaRAGE and the manual headspace
342 method (Fig. ~~34~~). ~~The observed~~ The occurrence of a maximum in the vertical profile of
343 dissolved CH₄ concentration in the upper layer of Lake Stechlin (Fig. ~~3b4a~~) is consistent with
344 previous observations (Grossart et al., 2011; Tang et al., 2014; Hartmann et al., 2018). In Lake
345 Arend we also observed a CH₄ peak (Fig. ~~3a4a~~) although the overall concentration was lower.
346 The opposite was observed at Lake Großer Pälitz and Lake Zotzen (Fig. 4c) with an anoxic
347 hypolimnion, where the dissolved CH₄ concentration was three orders of magnitude higher than
348 at in the epilimnion. Higher dissolved CO₂ (10² - 10³ μM) was also observed in the hypolimnion
349 of these two lakes (Fig. 4d) in comparison to the Lake Stechlin and Lake Arend (< 10² μM in
350 Fig. 4b).

351 In contrast, ~~with to~~ the headspace method, the FaRAGE allowed for ~~the localized profiles~~
352 of CH₄ and CO₂ concentration maximum to be described at a high vertical resolution, similar
353 to that obtained with more sophisticated membrane filter equilibrators (Hartmann et al., 2018;
354 Gonzalez-Valencia et al., 2014). The FaRAGE was capable of resolving differences in
355 dissolved CH₄ and CO₂ concentrations in lake water at decimeter ~~scales~~ resolution with ease.
356 Whilst care should be taken to ensure the sampling hose moves smoothly and slowly through
357 the water column, continuous profiling of a 20 m deep lake can be completed in 30 min. This
358 is a big advantage since *in situ* CH₄ concentrations can vary at very short time scales (hours to
359 days) subject to internal production, oxidation, weather conditions, ~~and~~ etc. (cf. Hartmann et al.
360 (2020)).



362

363 **Fig. 4** Depth profiles of dissolved CH₄ and CO₂ concentration from a set of lakes in Germany:

364 (a)-(b) Lake Stechlin and Lake Arend with an oxygenated hypolimnion in summer; (c)-(d) Lake

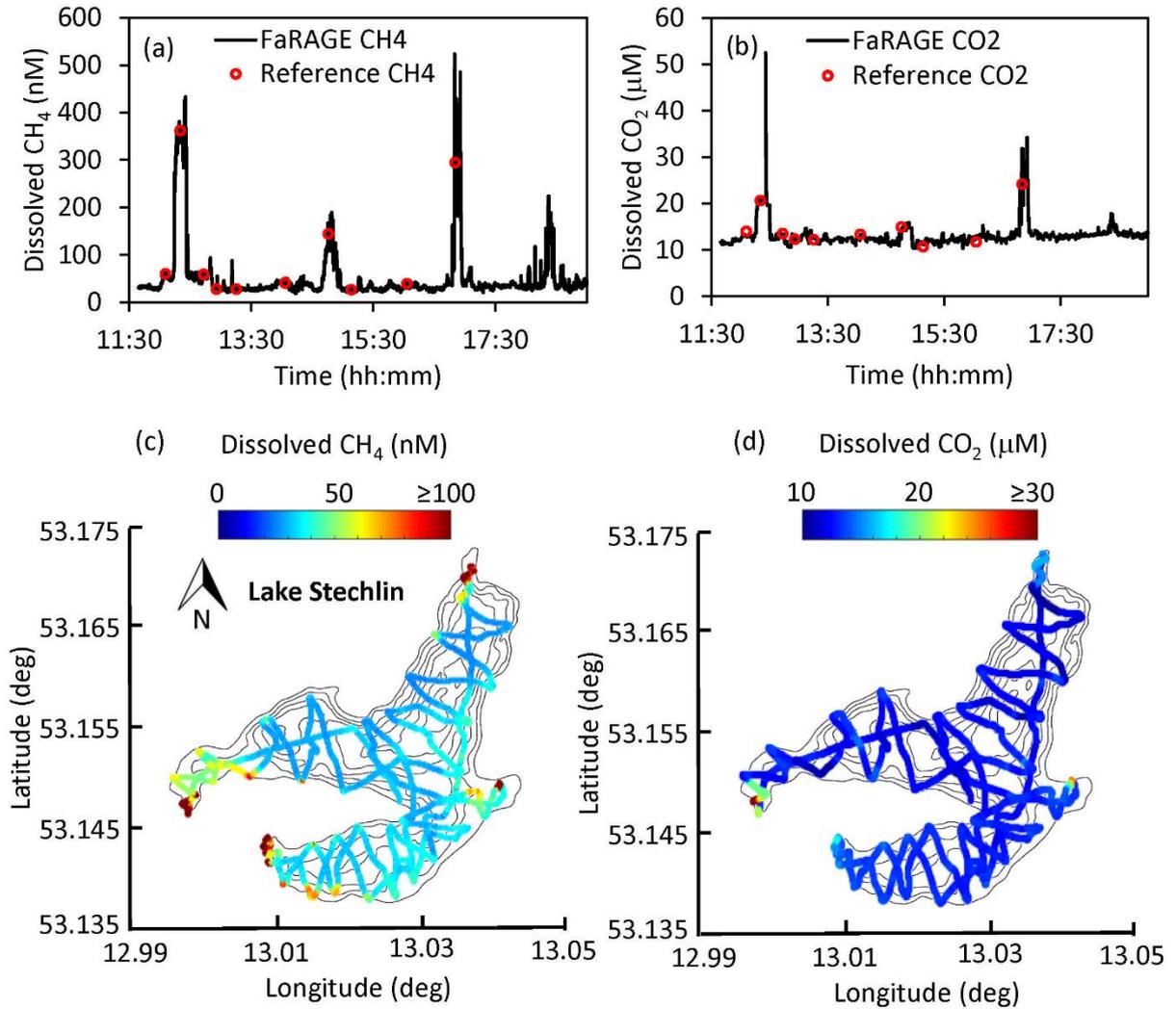
365 Großer Pälitz and Lake Zotzen, both with an anoxic hypolimnion in October. Note the log

366 transformed x-axis is used in (c)-(d). References using the headspace method are designated as

367 red open circles and measurements using the FaRAGE are shown as solid lines.

3.3 Resolving spatial ~~variabilities~~variability of dissolved CH₄ and CO₂ concentrations

We confirmed the capability of the FaRAGE to operate continuously over a 7-h period without notable decreases in performance (Fig. [4a5a-b](#)). Benefitting from its fast response rate, surface water dissolved CH₄ and CO₂ concentrations across the 4.52 km² Lake Stechlin ~~was~~ were mapped with great detail within one day ([Fig. 5c-d](#)). During the cruise, 10 reference measurements were made at different sites and times, which were consistent with nonstop online *in situ* measurements. The cruising survey demonstrated the capability of this device for resolving not just vertical dynamics of CH₄ and CO₂ in lake water, but also the potential for studying horizontal gas distributions ~~of CH₄~~ across large distances, for instance large lakes and rivers. With a driving speed of 5 km h⁻¹ and a response time of 12 s, a spatial resolution of 17 m can be achieved, which is sufficient for such a medium-sized lake. ~~The relative higher dissolved CH₄ concentrations in the shallow littoral zone of Lake Stechlin (Fig. 4b) reflect higher CH₄ release from the local sediment.~~



382

383 **Fig. 5** Map of surface dissolved CH₄ concentration at Lake Stechlin. (a)-(b) Time series of 7-h
 384 continuous surface water CH₄ and CO₂ measurement on March 28, 2019. The reference
 385 headspace measurements are shown as red circles. (c)-(d) Spatial distribution of surface water
 386 CH₄ and CO₂ concentration is given on top of the lake's bathymetry. Colored symbols show
 387 CH₄ and CO₂ concentrations according to the color bars. Black lines show the outline of the
 388 lake with depth contours.

389 4 Comments and Recommendations

390 4.1 Adaptability to different gas analyzers

391 The reasons for the ~~significantly-substantially~~ shortened response time of the FaRAGE
392 compared to other types of gas equilibrators are two-fold. While the working principle of the
393 FaRAGE is based on the bubble-type (Schneider et al., 1992) and Weiss-type equilibrators
394 (Johnson, 1999), a reduced headspace volume is adopted, which enhances the physical gas-
395 water exchange. Another reason is the use of an extremely fast-response gas analyzer (Picarro
396 Gas Scouter 4301). It is a highly recommended combination for ~~concentration~~-measurement of
397 dissolved gasses when the best time-wise performance is preferred due to its great mobility
398 (Table S2). However, coupling to other Cavity-Ring-Down gas analyzers is also possible (Table
399 S3). This feature enables a possibility to investigate stable isotopic nature of dissolved CH₄ and
400 CO₂, which is important when sources of CH₄ and CO₂ need to be identified.

401 When a portable gas analyzer (Picarro Gas Scouter or Ultraportable Los Gatos) is used
402 for measuring CH₄ and CO₂ concentrations only, the gas equilibrator can be optimized for
403 different application environments. The length of coiled tube for gas-water mixing can be
404 adjusted to change the response time (Fig. ~~2d3c-d~~). For smaller lakes a higher spatial resolution
405 can be obtained by shortening the equilibration tubing, which shortens the response time, and
406 hence increases the spatial resolution, whilst maintaining an acceptable equilibration ratio (51%
407 when tube length is 1 m). In environments with extremely low dissolved CH₄ concentrations,
408 e.g. ocean waters, a longer gas-water mixing tube should be used to ensure a high gas
409 equilibration ratio.

410 To measure stable isotopic CH₄ and CO₂ in water, the sensitivity of the FaRAGE can be
411 modified to better adapt to the choice of gas analyzer (~~e.g., when Picarro G2201-i or G2132-i~~
412 ~~is used~~). For example, high dissolved CH₄ concentrations (e.g. μM-to-mM range) can be
413 measured with greater accuracy by increasing the flow rate of the carrier gas relative to the
414 sample water flow, therefore diluting the CH₄ concentrations to the range of the gas analyzer.
415 This can be particularly useful, for instance, when an instrument has an optimal precision at a

416 low concentration range (1.8-12 ppm, ~~for e.g., Picarro G2201-i or G2132-i analyzers~~ ~~Picarro~~
417 ~~isotopic gas analyzer~~) for ^{13}C -CH₄ measurements. By using pure N₂ gas or carrier gases (e.g.
418 Helium and Argon) and corresponding gas analyzers, it would be possible to measure other
419 dissolved trace gas concentrations, e.g. ~~CO₂ can be measured simultaneously (CO₂ was tested~~
420 ~~in this study, but not shown for simplicity)~~ N₂O. ~~In addition, benefited due to from the high~~
421 ~~equilibration ratio of this device (max. 91.8%), it would be possible to measure dissolved CH₄~~
422 ~~(and other gases) close to equilibrium concentrations.~~

423 **4.2 Uncertainties due to suspended solids, temperature and pressure change**

424 The FaRAGE is proven to be resistant to suspended solids in freshwater lakes without
425 having to use additional accessories. As shown in Fig. S3, apparent phytoplankton blooms were
426 observed in the two studied lakes each with a high biomass (Chl-a > 30 $\mu\text{g L}^{-1}$) in the epilimnetic
427 water. The measurements were unaffected, without any interruptions during measurements. As
428 algal particles are a large component of suspended particle concentration in freshwater
429 clearwater lakes without a high suspended sediment load ~~concentration~~ systems, it is safe to
430 claim the resistance of this device to suspended solids in such a-systems. However, care must
431 be taken to avoid the water intake hose hitting the bottom sediment, which could cause blockage
432 of the water hose. An additional filtration unit for the water intake might be needed when the
433 device is to be applied to turbid rivers.

434 The temperature and hydrostatic pressure could both change when water is pumped out
435 through a water hose. To consider the temperature effect, a fast temperature logger is used (Fig.
436 1) which allows for corrections in calculation. Instead of using *in situ* lake temperature, the
437 temperature measured at the gas equilibrator, ~~should be used~~ where gas equilibration occurs, should be used. Our measurements found a minor effect when measuring surface waters but an
438 apparent warming for hypolimnetic water in deep lakes (Fig. S4). ~~While a calibration can be~~

440 ~~done directly by taking water samples from multiple depths of the lake (e.g., Fig. 3) to consider~~
441 ~~this effect, one could make the calculation without taking many samples by applying~~
442 ~~temperature correction.~~

443 The temperature correction can be made by referring to the manual headspace method.
444 The constant gas and water flow can be used as headspace and water volume, respectively. By
445 considering the temperature and pressure effects on gas solubility, the dissolved CH₄ and CO₂
446 concentrations can be calculated (an example calculation sheet is provided in Table S5). The
447 calibration curve can be established using the manual headspace measurements as standards.
448 The final concentrations can be corrected for partial equilibration by applying the equation from
449 the calibration curve (e.g., Fig. 2a-b). The response time should be deduced when calculating
450 CH₄ and CO₂ depth profiles and spatial distributions, in addition to the time lag caused by
451 pumping water samples by using an extended water intake hose.

452 **4.3 Calibration, maintenance and mobility**

453 The FaRAGE can be readily adopted for measuring other trace gases when coupled with
454 other portable gas analyzers. Due to differences in gas solubility (Duan and Sun, 2003;
455 Wiesenburg and Guinasso Jr, 1979), for each new gas, it would be necessary to establish the
456 relative equilibration efficiency and response time, following the approach we outlined here for
457 CH₄ and CO₂. Once set, a new calibration is only required when the tubing diameter or length
458 is changed (when the old one is filthy-no longer usable due to biofilm growth). This can be done
459 by referring to a number of known concentrations that covers a wide range (at least 5), e.g.,
460 taking water samples from different water depth of the lake or a gradient from littoral to pelagic
461 zones. Once this full calibration is made, the calibration curve can be used for calculating the
462 subsequent measurements. A one-point reference measurement should be performed between
463 depth profiles or transects to check for apparent drifting. This can usually be done by taking
464 one surface water sample from a lake for manual headspace measurement. Care should be taken

465 when measuring in lakes with an anoxic hypolimnion where hydrogen sulfide is likely to
466 accumulate. The performance of Cavity-Ring-Down gas analyzers can be potentially affected
467 by ~~organics, ammonia, ethane, ethylene, or sulfur containing compounds~~ H₂S gas (Kohl et al.,
468 2019). At these sites, ~~it is always recommended to take additional samples and measure them~~
469 ~~with traditional methods (e.g., with a Gas Chromatograph Analyzer)~~ it is recommended to use
470 a copper scrubber to remove H₂S from the gas samples (Malowany et al. 2015) and no time
471 delay will be induced.

472 The gas equilibrators should be carefully maintained. Replacement of parts is
473 recommended at a monthly basis provided the device is heavily in use. They include bubble
474 diffusor and the coiled gas-water mixing tube. In addition, to ensure the performance and
475 prevent biofilm formation the gas-water mixing and separation units should be cleaned after
476 use. Running with distilled or Milli-Q water would help to rinse the device and reduce the risk
477 of biofilm development in the inner tubes. The performance of peristaltic pumps should be also
478 regularly checked and the inner pump tubes need to be replaced to ensure a constant water flow.

479 The combination of FaRAGE with the Picarro Gas Scouter provides the most mobility.
480 The system can be easily carried by one person and work in a small aluminum or inflatable boat
481 ~~with where~~ a maximum capacity of three people is possible. The device can also work in bad
482 weather with additional measures based on protecting the gas analyzer from water damage by
483 rain or flooding.

484 **Code availability**

485 Not applicable.

486 **Data Availability**

487 An example calculation sheet (raw data of Fig. 2a) is provided as part of supporting
488 information for device calibration and for temperature and pressure correction when calculating

489 dissolved methane concentration. The full data sets associated with lab and field tests are
490 available upon request.

491 **Supplement link**

492 From Copernicus.

493 **Author contributions**

494 SBX and WW proposed the idea and built the first prototype. LL improved the prototype
495 and conducted lab and field tests. JW contributed to the field tests. AL contributed to the
496 derivation of equations; HPG [led the project and](#) advised the development of the modified
497 prototype. LL drafted the initial manuscript. All authors discussed the results and commented
498 on the manuscript.

499 **Competing interests**

500 The authors declare that they have no conflict of interest.

501 **Acknowledgements**

502 This work was financially supported by the National Natural Science Foundation of
503 China (grant No. 51979148 [& 91647207](#)). L.L., J.W. and H.P.G. were financially supported by
504 [the “Aquameth” project of](#) the German Research Foundation (DFG GR1540/21-1+2). [We thank](#)
505 [Andreas Jechow, Christine Kiel, Igor Ogashawara, Katrin Kohnert, Sabine Wollrab and Stella](#)
506 [Berger for providing support with collecting field test data under project CONNECT \(SAW-](#)
507 [K45/2017\) which is funded by the Leibniz Association, Germany.](#) The authors would like to
508 thank Hannah Geisinger and Truls Hveem Hansson for helping collecting field data.

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1 **Supporting Information for**

2 **A Fast-response automated gas equilibrator (FaRAGE) for continuous *in situ***

3 **measurement of methaneCH₄ and CO₂ dissolved in water**

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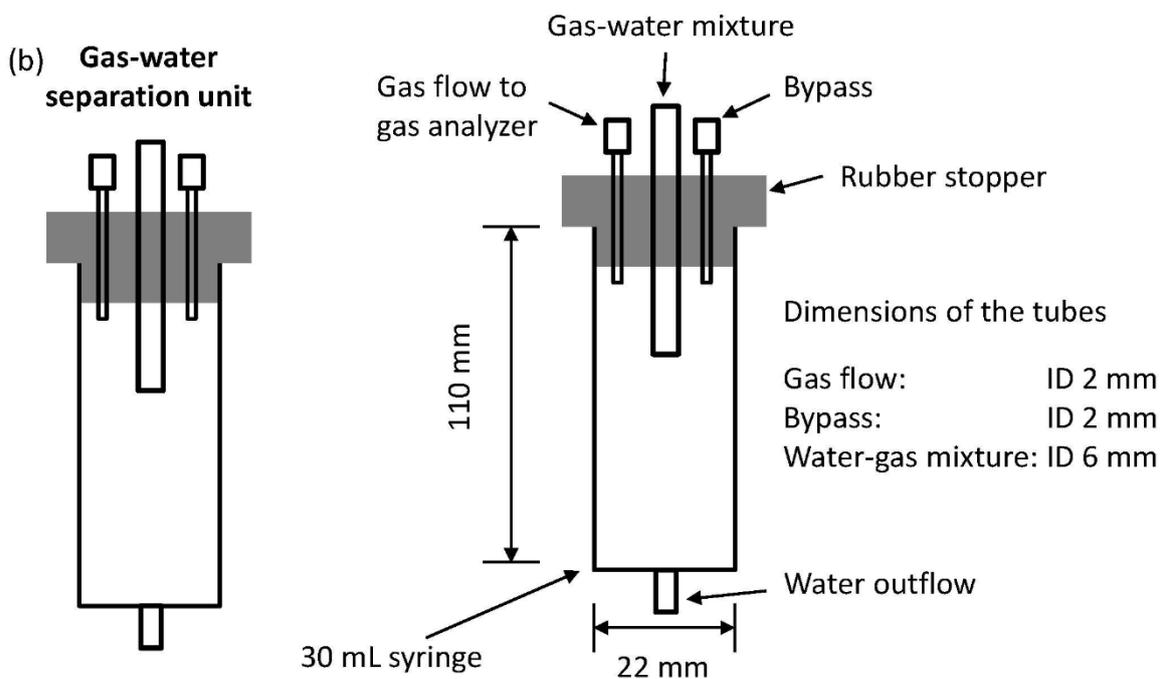
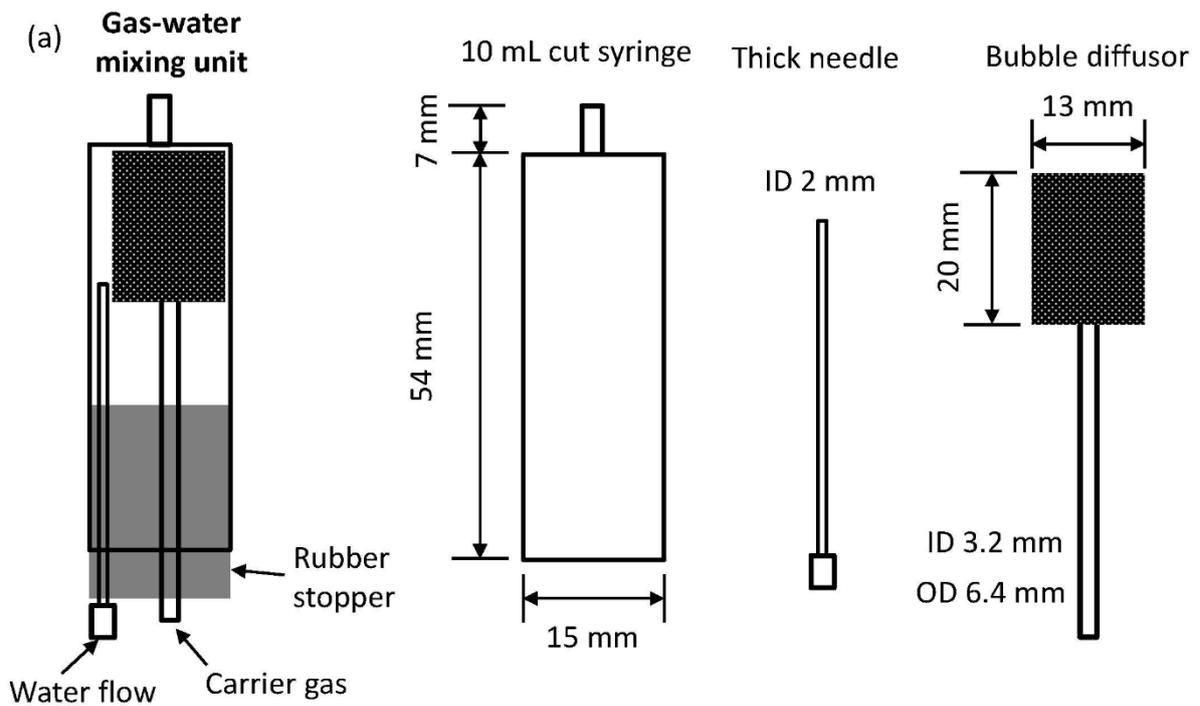
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21 **S1. Details of parts, gas analyzers and costs**

22 To make the FaRAGE field deployable, parts were tightly packed into an aluminum box
23 with a built-in power supply. The electric parts were separated from other parts containing
24 water in the box by using a plastic board. Ports were well labelled on the right-handed side so
25 that even somebody new to the system can work with it. To help interested readers rebuild the
26 device, the two key components (gas-water mixing unit and gas-water separation unit) were
27 shown in the detailed technical drawings (Fig. S1). The suppliers and costs for these parts
28 were listed in Table S1. A total of 3,560 € was calculated for building the complete device
29 excluding the costs for the power supply. As the expensive RBR temperature logger is not a
30 necessity since we happen to have it in storage, a cheaper temperature logger can always be
31 used. For example, a fast HOBO temperature logger (HOBO U12 with a Temperature probe
32 TMC1-HD) is available for < 200 €. The total cost can be cut down significantly to < 3,000 €.

33 The FaRAGE is capable of coupling with different greenhouse gas analyzers, depending
34 on the research question and instrument availability. Three most widely used field-deployable
35 gas analyzers were compared in Table S2 to provide a reference for readers when choosing a
36 gas analyzer. They are GasScouter G4301 (Picarro, USA), Ultraportable Greenhouse Gas
37 Analyzer (Model 915-0011, LosGatos Research, USA) and Picarro G2132-i isotope analyzer
38 (Picarro, USA). We noticed Picarro 2201-i has been more often used, but our Picarro G2132-i
39 is an equivalent instrument except that the module for isotopic CO₂ is not installed. The
40 former two instruments measure CH₄, CO₂ and H₂O and the last one additionally measures
41 stable isotopic CH₄. As shown in Table 2, clearly GasScouter G4301 is most suitable for field
42 measurement of dissolved CH₄ concentrations due to its extremely high mobility. The built-in
43 battery pack can support 8 h continuous measurements and the ability to amount GPS antenna
44 offers the advantage in doing spatially-resolved measurements. The Picarro G2132-i isotope
45 analyzer is most immobile because of it is heavy and relative high power consumption in

46 addition to its particularly long time to warm up (30 min). However, Picarro G2132-i
 47 measures stable isotopic CH₄, while the other two instruments cannot. Care must be taken and
 48 a proper boat with stable power supply is needed in order to use Picarro G2132-i as a coupling
 49 unit for the FaRAGE.



50

51 **Fig. S1** ~~Assembled Technical drawings of FaRAGE key components, in an aluminum box~~
52 ~~with built in power supply. Main parts and ports are marked on the figure. The power supply~~
53 ~~is made from a 100 Ah (12 V) Li-ion battery coupled with a power transformer that converts~~
54 ~~12 V DC power to 230 V AC power.~~(a) Gas-water mixing unit and (b) gas-water separation
55 unit. Note: ID and OD are the abbreviations of inner diameter and outside diameter,
56 respectively.

57 **Table S1.** List of materials for parts of the FaRAGE prototype. Details on dimensions, model, producer/supplier and cost are provided.

Items	Dimensions	Model specifications	Producer/Supplier	Quantity	Cost
Diving tank	10 L	Pressure up to 230 bar	Atlantis Berlin	1	199 €
Pressure regulator		200 bar / 0 - 10 bar, HERCULES CK1401	Gase Dopp	1	59.98 €
Mass flow controller (for air)		SIERRA Model C50L SMART-TRAK	SCHWING Verfahrenstechnik GmbH	1	995 €
Flow meter (for water)		0.082-0.82 L min ⁻¹ , 1355GAF3CBXN1AAA	Brooks Instrument GmbH	1	943.91 €
Peristaltic pump	9 x 11 x 16 cm	0-500 mL min ⁻¹ , 24V/1A DC power	Purchased from Taobao, China	2	200 €
Temperature logger		Precision 0.001 °C, maximum 6 Hz measurement frequency, TR-1050	RBR, Canada	1	1,000 €
Tygon tube	3.2/6.4 mm in./out. Ø	Saint-Gobain Schlauch Tygon S3 E-3603 2.5bar	RS Components GmbH	15 m	68.78 €
Plastic syringe for mixing unit	5 mL	Cut to 3 mL, sealed with a rubber stopper	BD plastipak	1	1 €
Plastic syringe separation unit	30 mL	Sealed with a rubber stopper	BD plastipak	1	1 €
Plastic syringe for desiccant	50 mL	Filled with silicone beads, sealed with a rubber stopper	BD plastipak	1	1 €
Rain pipe			Toom	1	10 €
Bubble diffusor	12 mm Ø, 16 mm length	Pawfly 0.6 Inch Air Stone, UL266	Ebay	1	1 €
Teflon membrane filter	25 mm Ø	PTFE 0.2 µm	Lab Logistics Group GmbH	2	2 €
Tube connector	for 3.2-4.2 mm	LL male, barbed hose connection: PP, 10 pcs/pack 2- 1882	neoLab Migge GmbH	10	12 €
Aluminium box	38.3 x 57 x 37.5 cm	65 L, Stier aluminium box	Amazon	1	64.95 €
Total					3,560 €

58 **Table S2.** Summary of technical details for the three greenhouse gas analyzers tested in this study.

Analyzer	Gases	Gas flow rate	Cavity pressure	Measurement frequency	Concentration range	Precision	Response time	Dimensions	Weight	Power consumption	GPS Kit	Mobility
GasScouter G4301	CH ₄ CO ₂ H ₂ O	1 L min ⁻¹	> 700 Torr	1 Hz	CH ₄ : 0-800 ppm CO ₂ : 0-3% H ₂ O: < 3%	CH ₄ : 3 ppb CO ₂ : 0.4 ppm	5 s	35.6 × 17.7 × 46.4 cm	10.4 kg	25 W, built-in Li-ion battery	Yes	Very high
Ultraportable Greenhouse Gas Analyzer 915-0011	CH ₄ CO ₂ H ₂ O	0.5 L min ⁻¹	140 Torr	1 Hz	CH ₄ : 0.01-100 ppm CO ₂ : 1-2% H ₂ O: < 7%	CH ₄ : 2 ppb CO ₂ : 0.6 ppm	~10 s	17.8 x 47 × 35.6 cm	17 kg	70 W, on battery/AC power	No	High
Picarro G2132-i	CH ₄ $\delta^{13}\text{C}$ -CH ₄ CO ₂ H ₂ O	25 mL min ⁻¹	148 Torr	0.5 Hz	CH ₄ : 1.8-10 ppm high-performance mode; 10-1000 high-range mode CO ₂ : 200 - 2000 ppm guaranteed range H ₂ O: <2.4 % guaranteed range	CH ₄ : 5 ppb + 0.05 % of reading (12C); 1 ppb + 0.05 % of reading (13C) CO ₂ : 1 ppm + 0.25 % of reading (12C)	~30 s	43.2 x 17.8 x 44.6 cm	27.4 kg	205 W, AC power	No	Fair

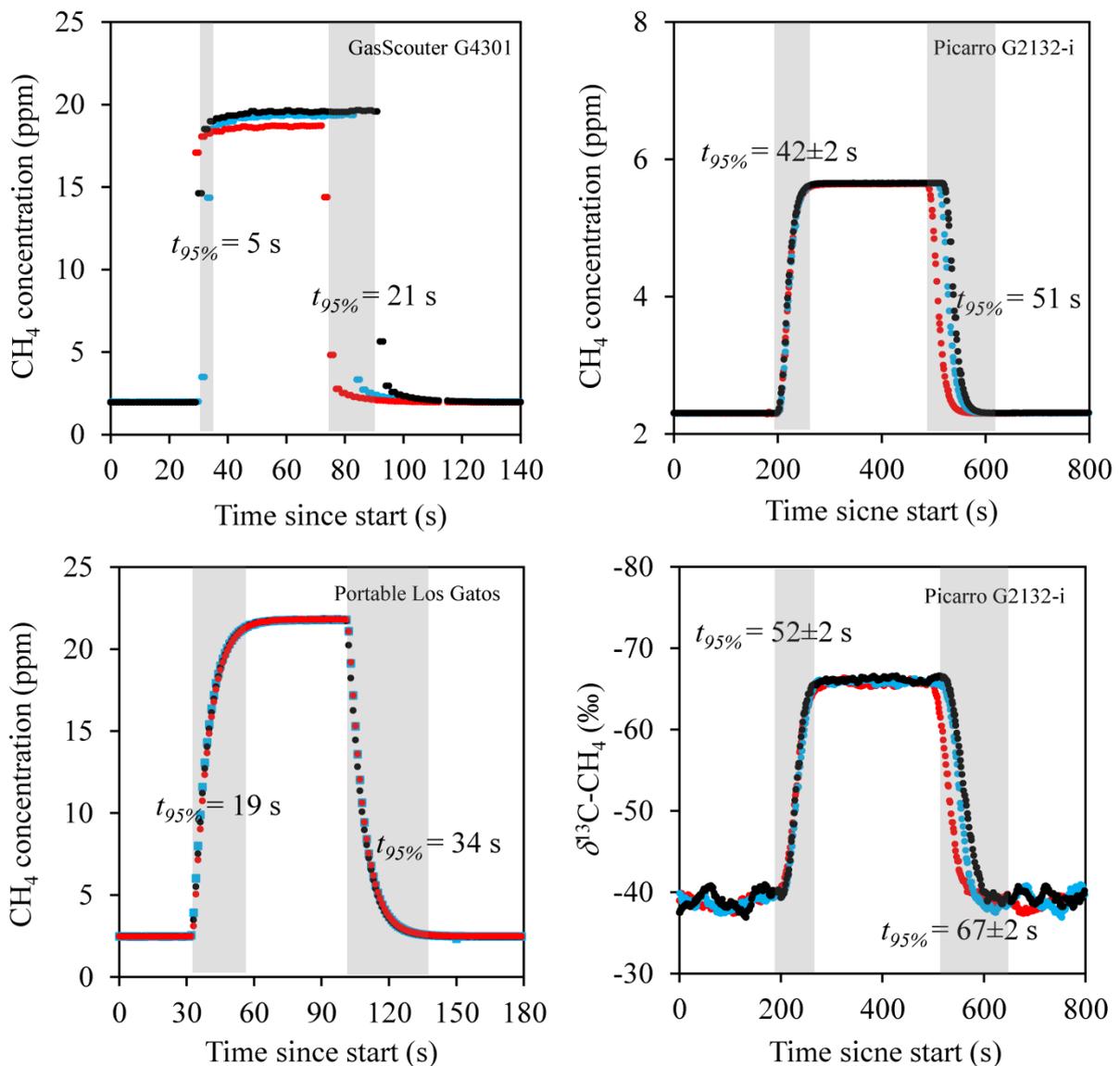
59 Note: 1) GasScouter G4301 does not use a vacuum pump to maintain a stable cavity pressure and the gas flow rate should be stable but slightly
60 above/below the recommended value.

61 2) All gas analyzers are sensitive to liquid-phase water, therefore a hydrophobic filter is normally placed before the gas intake to protect
62 instrument from being flooded.

63 3) According to Picarro, interference can occur for concentrations of H₂O and CO₂ well above normal ambient levels, as well as other
64 organics, ammonia, ethane, ethylene, or sulfur containing compounds.

65 S2. Re-evaluation of response time of gas analyzers

66 While response time for each gas analyzer has been provided by its manufacturer (Table
67 S2), a large difference was found when they were re-evaluated (Fig. S2). Picarro GasScouter
68 has the fastest response to concentration increase, in comparison to four-fold and eight-fold
69 slower response for portable Los Gatos and Picarro G2132-i, respectively. All three gas
70 analyzers were seen longer response time when concentration changed from high to low. The
71 Picarro GasScouter still has the best performance compared to the other two.



72

73 **Fig. S2** Response times of gas analyzers. Triplicated measurements were performed. Low-to-
74 high and high-to-low concentration changes were investigated. The response time was
75 determined by taking the time when 95% of final concentration was reached. For $\delta^{13}\text{C-CH}_4$,
76 30 s moving average data was used.

77 **S3. Theoretical background**

78 With the present design of the **Fast-Response Automated Gas Equilibrator (FaRAGE)**,
79 a continuous dynamic gas-water mixing occurs and the carrier gas is partially equilibrating
80 with the CH_4 dissolved in water sample. The gas composition reaching the gas analyzer
81 depends on equilibration time and flow rates. The equilibration between the carrier gas and
82 the water sample during flowing through the FaRAGE depends on the concentration
83 difference between the gas stream (C in $\mu\text{mol L}^{-1}$) and the dissolved (aqueous) concentration
84 in the sample water (C_a):

$$85 \quad \frac{dC}{dt} = k \times \left(\frac{1}{HRT} C_a - C \right) \quad (1)$$

86 Where H is the temperature-dependent Henry constant ($\text{mol L}^{-1} \text{atm}^{-1}$), R the universal
87 gas constant ($8.31 \text{ J mol}^{-1} \text{K}^{-1}$), T is temperature (K) and k (s^{-1}) is an exchange coefficient. The
88 equilibrium gaseous concentration $C_{eq} = \frac{1}{HRT} C_a$ corresponds to the headspace concentration
89 of a fully equilibrated water sample. k is expected to depend on the relative flow rates of gas
90 and water as well as on the flow regime and mixing of both phases in the FaRAGE. For an
91 initial concentration of CH_4 in the carrier gas C_{ini} , the time-dependent concentration during the
92 passage through the equilibrator is:

$$93 \quad C(t) = (C_{ini} - C_{eq})e^{-kt} + C_{eq} \quad (2)$$

94 After a device-specific partial equilibration time t_e , the CH_4 concentration in the carrier
95 gas has changed to C_{pe} , which is measured by the gas analyzer

96
$$C_{pe} = C(t_e) = K(C_{ini} - C_{eq}) + C_{eq} \quad (3)$$

97 With $K = e^{-kt_e}$ being a device-specific coefficient, which can be obtained by calibrating
 98 the FaRAGE with at least one water sample of known dissolved concentration (C_{eq}) through:

99
$$K = \frac{C_{pe} - C_{eq}}{C_{ini} - C_{eq}} \quad (4)$$

100 The equilibrium headspace concentration of CH₄ in the water sample and the
 101 corresponding dissolved concentration can be estimated from the initial and final carrier gas
 102 concentration as:

103
$$C_{eq} = \frac{1}{HRT} C_a = (KC_{ini} - C_{pe}) / (K - 1) \quad (5)$$

104 For a high flow rate of the carrier gas, the response time of the system to changing
 105 dissolved concentrations at the sample intake is predominantly determined by the gas venting
 106 rate, i.e. by the total volume of carrier gas that is in contact with the water sample, divided by
 107 the volumetric gas flow rate (cf. level two model of Johnson (1999)), as well as by the
 108 response time of the gas analyzer.

109 **Table S3.** Response times when adapting to different gas analyzers. Tests were performed
 110 with a water/gas mixing ratio of 0.5. Triplicates were made and mean values are shown here.

Gas analyzer	Treatment	t _{95%} response time (s)		
		CH ₄	CO ₂	¹³ C-CH ₄
Gas Scouter G4301	Low-to-high	13	6	-
	High-to-low	13	6	-
Ultraportable Greenhouse Gas Analyzer 915-0011	Low-to-high	34	32.3	-
	High-to-low	37	30	-
Picarro G2132-i	Low-to-high	53	53	53
	High-to-low	65.3	60.7	65.3

111 Note: Response time for Picarro G2132-i was determined without using a desiccant. ~~Repeated~~
 112 ~~tests were done to examine the effect of desiccant. The response time was extended by 150 s~~
 113 ~~when the 20 mL desiccant was used.~~ A desiccant should be used to keep the moisture content
 114 in gas samples < 1%. Drierite and magnesium perchlorate (Mg(ClO₄)₂) are recommended for
 115 such a purpose due to their high performance. It was shown by Webb et al. (2016) that both
 116 types of dryer had no effect on CH₄ and CO₂, except for a 1.5 min delay in response time for
 117 CO₂ when using Drierite.

118 **Table S4.** Comparison of response times for simultaneous measurement of dissolved CH₄ and
 119 δ¹³C-CH₄ in water from previous studies using different devices (after Webb et al., 2016,
 120 Hartman et al., 2018) compared with response times in this study. Response time was unified
 121 here to *t*_{95%} to allow for meaningful comparison. The *t*_{95%} values were taken from literature by
 122 applying *t*_{95%} = 3τ and the mean were used.

Device	<i>t</i> _{95%} response time (s)	Study
Weiss-type (small)	6744	Li et al. (2015)
General oceanics	6123	Webb et al. (2016)
Shower head	4971	Webb et al. (2016)
Weiss-type (large)	3600	Rhee et al. (2009)
Marble	2679	Webb et al. (2016)
Bubble-type	2034	Gülzow et al. (2011)
Liqui-Cel (medium)	1251	Webb et al. (2016)
Liqui-Cel (small)	531	Webb et al. (2016)
Liqui-Cel (large)	351	Webb et al. (2016)
Liqui-Cel (small) in vacuum mode	171	Hartmann et al. (2018)
Combined Weiss-type with bubble-type	53	This study

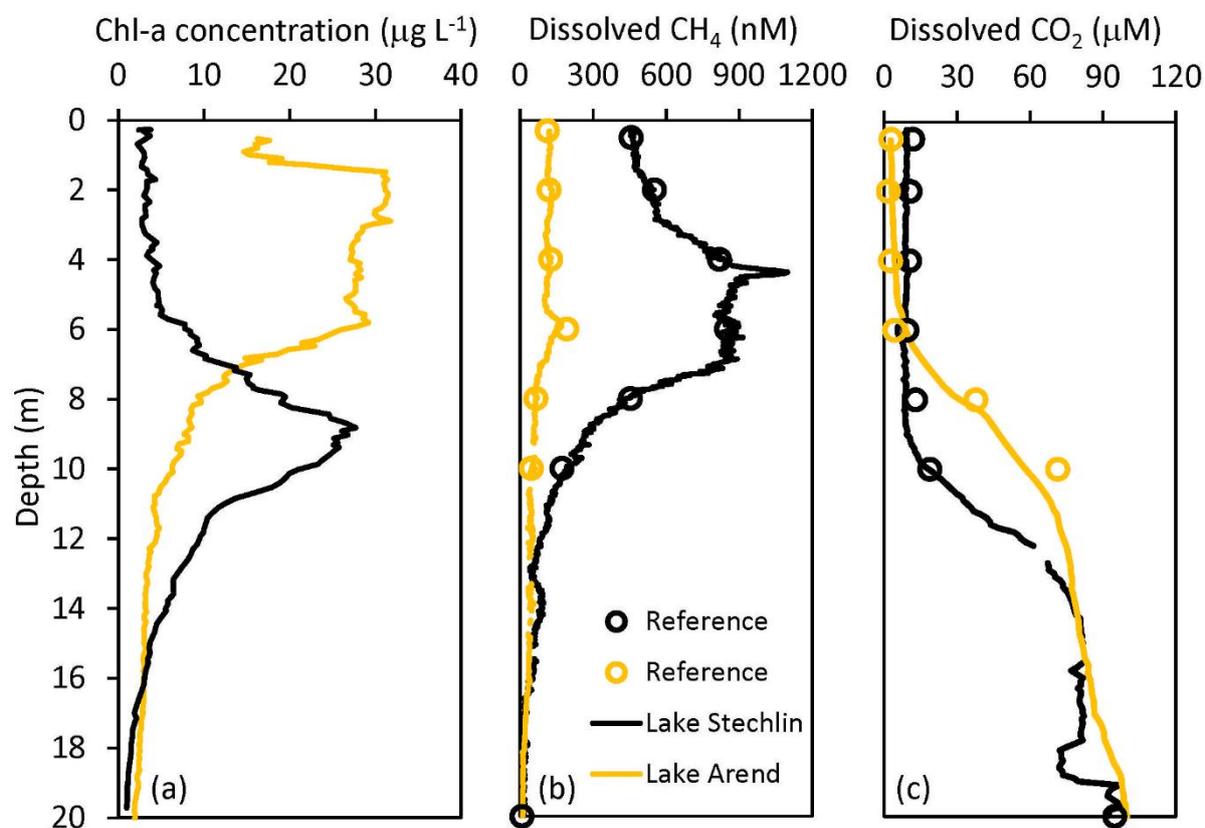
123

124 **S4. The depth profiles of phytoplankton biomass at Lake Arend and Lake Stechlin**

125 As in most freshwater lakes phytoplankton is a large component of suspended solids in
 126 water column, the effect of phytoplankton biomass on the performance of the gas equilibrators

127 was evaluated. Fig. S3 shows ~~that the presence of in both study lakes~~ high phytoplankton
128 biomass (represented by Chl-a) within the surface 20 m water depth in the both study lakes.

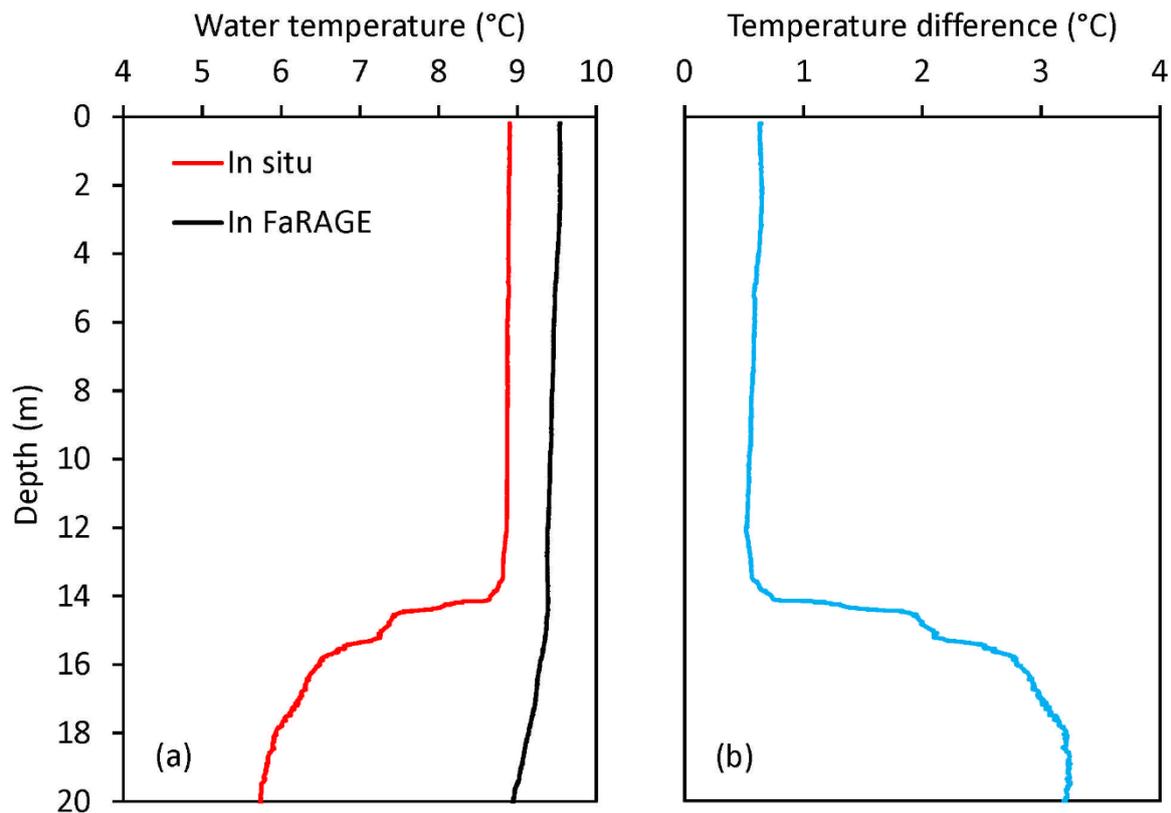
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130

131 **Fig. S3** Depth profiles of Chlorophyll-a (Chl-a) at Lake Arend and Lake Stechlin on June 17
132 and July 23, 2019 with (b)-(c) dissolved CH_4 and CO_2 profiles. The profiles were measured
133 using a BBE FluoroProbe (Moldaenke, Germany) simultaneously with dissolved gas profiles.

134



135

136 **Fig. S4** An example of altered depth profile of water temperature at Lake Stechlin in autumn
 137 2019. (a) Comparison of in situ water temperature (red line) with water temperature measured
 138 in the FaRAGE (black line). (b) The difference between the two temperature measurements
 139 (In FaRAGE - In situ).

140

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