changing climate <u>– changes in moderate extremes and their</u> <u>seasonality</u>

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Abstract. Future changes in <u>river_runoff_will</u> impact many sectors such as agriculture, energy production, or ecosystems. Therefore, assessments of runoff characteristics under climate change are crucial for decision-makers and water management planners. We<u>Here, we</u> study changes in moderate runoff extremes, i.e. low and high flows that occur once every year or season in today's climate. Daily runoff is simulated for 93 Swiss catchments for the period 1981 2099 under the Representative Concentration Pathway 8.5 using 20 downscaled regional climate models from the newest transient Swiss climate change scenarios.

The the seasonality, frequency, and magnitude of moderate annual low flows is projected to decrease in lower lying catchments and to increase in Alpine catchments. Seasonal low flows in summer are projected to decrease and seasonal low flows in winter to increase. Moderate annual high flows are projected to slightly increase in most catchments but to decrease in high Alpine catchments. However, the climate model agreement on the sign of change in moderate high flows is not robust. The projected decrease in Alpine catchments contradicts results for extreme high flows from previous studies. This difference may be due to different indicators used (moderate extremes vs. extremes).low and high flows and their time of emergence. The time of emergence indicates the timing of significant changes in the flow magnitudes. For low flows-Daily runoff is simulated for 93 Swiss catchments for the period 1981–2099 under Representative Concentration Pathway 8.5 with 20 climate model chains from the most recent transient Swiss climate change scenarios.

In the present climate, annual low flows typically occur in the summer half-year in lower-lying catchments (<1500 m.a.s.l.) and in the time of emergence is early in 21st century in high Alpine catchments due to early changes in winter low flows. In lower lying catchments, significant changes in low flows emerge later in the century. For moderate high flows, only few catchments indicate a significant change.

Shifts in the seasonality of moderate low flows due to climate change are found in many catchments-winter half-year in Alpine catchments (>1500 m.a.s.l.). By the end of the 21st century, annual low flows are projected to occur in late summer and early autumn in most catchments-indicating. This indicates that the lack of decreasing precipitation and increasing evapotranspiration in summer and autumn exceedsexceed the water contributions from other processes such as snow and glacier melt contributions. For moderate high flows, changes in seasonality are found in Alpine catchments with a shift towards earlier occurrence in summer due to a reduced contribution of snow and glacier melt in summer. In the projections, low flows occur more frequently in lower-lying catchments and less frequently in Alpine catchments. For high flows, the frequency of annual low flows increases slightly, but their magnitude decreases and becomes more severe. In Alpine catchments, annual low flows occur less often and their magnitude increases. The magnitude of seasonal low flows is projected to decrease in the summer half-year in most catchments, but models often disagree on the sign of change. Changes in the annual co occurrence of moderate low and highto increase in the winter half-year in Alpine catchments. Early time of emergence is found for annual low flows are mainly due to changes in the frequency of low flows that increases in lower lying catchments and decreases in Alpine catchments. In the 21st century due to early changes in low flows in the winter half-year. In lower-lying catchments, significant changes in low flows emerge later in the century.

Annual high flows occur today in lower-lying catchments in the winter half-year and in Alpine catchments in the summer halfyear. Climate change will change this seasonality mainly in Alpine catchments with a shift towards earlier seasonality in summer due to the reduced contribution of snow and glacier melt in summer. Annual high flows tend to occur more frequent, and their magnitude increases in most catchments except some Alpine catchments. The magnitude of seasonal high flows in most catchments is projected to increase in the winter half-year and to decrease in the summer half-year. However, the climate model agreement on the sign of change in moderate high flows is weak.

1 Introduction

Assessments of climate change impacts on hydrology are crucial for future water management and adaptation planning. This is especially true for extreme events, which potentially have severe ecological and societal impacts. In this study, we focus on moderate runoff extremes in both tails of the runoff distribution: moderateparticular, annual and seasonal moderate low and high flows. Focusing on moderate extremes is motivated are relevant for several reasons. First, even moderate extremes are important for water management planning. Second, very extreme floodshigh flows and very extreme streamflow droughtslow flows are difficult to simulate, because many processes are not fully understood or not yet resolved in hydrological models. Third, hydrological models are calibrated on observed flow conditions and may miss plausible but unexperienced extreme events; that have not been experienced. Fourth, climate change projections incorporate large uncertainties regarding small-scale extremes, particularly for extremes in precipitation, which are potential flood triggers. Therefore, we focus on moderate extremes, i.e., events that occur on average once every year or season in today's climate. The larger sample size (number of events) increases the robustness of the changes estimated-changes.

Low flows have a strong impact on water quality, freshwater ecosystems, and human water use such as power production, drinking water production, irrigation for agriculture, fisheries, and recreation (IPCC, 2014). Today, long-term water management planning for Switzerland must rely on low-flow assessments from past observations. Since Because climate change is projected to alter low-flow characteristics, low-flow projections for the 21st century need to be integrated into water management planning. Changes in low--flow indicators in the pastrecent decades have been already identified in Europe (Stahl et al., 2010) and in Switzerland (Weingartner and Schwanbeck, 2020). For Switzerland, increasing low-flow magnitudes in winter low flows and decreasing low-flow magnitudes in summer low flows have been observed in nival (snow-driven) and pluvial (rain-driven) catchments. Low Magnitudes of low flows in glaciated catchments have increased in all seasons (Weingartner and Schwanbeck, 2020). Previous studies have assessed climate change impacts on low flows mainly for macroscalemacroscale catchments or and regions. Van Vliet et al. (2013) investigated low--flow changes on a global scale while, and other studies focused on European scales (e.g. Feyen and Dankers, 2009; Forzieri et al., 2014; Alderlieste et al, 2014; Papadimitriou et al., 2016; Vidal et al., 2016; Marx et al., 2018). For Switzerland, previous climate impact studies on low flows exist for lower-lying catchments in the Swiss Plateau (Meyer et al., 2011), for large-scale catchments (Bernhard and Zappa, 2012), and for very extreme low-flow regimes (100-year return periods) low flow regimes in aggregated regions (Brunner et al., 20192019a). The studies found decreasing low-flows-flow magnitudes in the lower-lying parts of Central Europe but increasing low-flows-flow magnitudes in Alpine areas, where runoff generation is mainly dominated by snow and glacier melt.

High flows <u>may also can</u> cause <u>severe damagesdamage</u> and <u>significant costs</u>. <u>Hence, potential changes in high flows have to</u> <u>be integrated in_and are also important for</u> water management and <u>infrastructure planning</u>, as <u>well.ecology</u>. Assessing future changes in <u>floodthe</u> magnitude, <u>flood</u>-frequency, and <u>flood</u>-timing <u>of high flows</u> is thus crucial for <u>planning and</u> decision <u>makers</u>. <u>Past events can help to put potential future changes into perspective.making</u>. Previous studies <u>have</u> investigated past trends in floods in Europe (e.g. Stahl et al., 2012; Hall et al, 2014; Mangini et al., 2018; Blöschl et al., 2019; Bertola et al., 2020) and in Switzerland (e.g. Birsan et al., 2005; Allamano et al., 2009; Schmocker-Fackl and Naef, 2010a,b; Castellarin and Pistocchi, 2012). No clear andor significant trend in flood magnitude was found inby these studies since the studies sometimesbecause they disagree on the direction of trends. Various factors make it difficult to compare trends in flood magnitude between catchments and between different studies. The assessments depend heavilystrongly on the quality and homogeneity of the observations, the underlying methods such as the selection of indicators orand statistical tests, and the time periods investigated time periods. Flood frequencies have increased in northern Switzerland and decreased in southern Switzerland in the recent past (Schmocker-Fackl and Naef, 2010a, Blöschl et al., 2019). Periods with many floods were found in the end of the 19th century and after 1968 in northern Switzerland (Schmocker-Fackl and Naef, 2010a). Several assessments of studies have also assessed future changes in floods in Switzerland have also been made (Allamano et al., 2009; Köplin et al., 2014; Beniston et al., 2016; Ragettli et al., 2020). Even though thoseAlthough these studies differ substantially in methodological aspects and catchment selection, they found in generalgenerally increasing but not necessarily significant changes in annual runoff maxima under climate change. Seasonal patterns of change were detected, with increasing flood magnitudes in summer floods-(Allamano et al., 2009). Also, futureFuture shifts in the seasonality of floods also depend on the regime type of the catchments (Köplin et al. 2014).

Here, we complement these assessments with a focus on moderate low and moderate high flows, i.e. annual or seasonal: 7-day runoff minima and daily runoff maxima. The new Hydro-CH2018-Runoff dataset (Muelchi et al., 2020a2020; Muelchi et al., 2020b in review2021a) is used. It consists of 119 years (1981-2099) long daily runoff simulations for 119 years (1981-2099) driven by the most up-to-date climate change scenarios for Switzerland CH2018 (CH2018, 2018). For the Representative Concentration Pathway RCP8.5 emission pathway (Moss et al., 2010; van Vuuren et al., 2011), we analyzeanalyse (1) changes of moderate low and high flows under climate change, (2) the point in time when significant changes emerge, (3) changes in the seasonality of moderate extremes, and (4) changes in thetheir frequency of and their (co-)-cocurrence. In a companion paper, Muelchi et al. (2020c, in review(2021b) assessed changes in runoff regimes and their time of emergence. Here, we extend this analysis with assessments of moderate low and high flows. SinceBecause both studies are based on the same simulations (Hydro-CH2018-Runoff ensemble), they complement each other and giveprovide a comprehensive overview on of hydrological changes in Switzerland. They also complement the above mentioned existing studies on future changes in extreme hydrological events.

2 Data

We analyse daily runoff simulations for 93 medium sized (14–1700 km²) catchments distributed in Switzerland (catchment areas range between 14 km² and 1700 km²) and covering a wide range of different runoff regime types, including glaciated catchments (22 catchments, glaciation between 0.2-% and 22%), mainly snow--driven catchments in the Alpine area, and lower

-lying catchments mainly driven by precipitation and evapotranspiration. Due to strong elevation dependence in the runoff response to climate change (Koeplin et al., 2014, Muelchi et al., 2021b), the catchments are divided into two groups: 39 catchments in the Alpine area with mean altitudes greater than 1500 m.a.s.l. (including 22 glaciated catchments) and 54 catchments in the lower-lying areas in Switzerland with mean altitudes lower than 1500 m.a.s.l. The locations of the catchments are depicted in Fig. 1_x with six representative catchments highlighted in greenblue. These representative catchments cover the most important regime types in Switzerland (Weingartner & Aschwanden, 1992): Rosegbach—x highly glaciated (22%); Signe Si

The data used for the analysis is the Hydro-CH2018-Runoff ensemble, consisting of daily mean runoff simulations for each of these 93 catchments (Muelchi et al., 2020a2020; Muelchi et al., 2020b in review 2021a). These simulations were run with the semi-distributed hydrological modelling system "semidistributed PREecipitation-Runoff-EVApotranspiration HRU Model" (PREVAH →) hydrological modelling system (Viviroli et al., 2009), PREVAH accounts for important hydrological processes such as evapotranspiration, soil moisture dynamics, snow accumulation, and snow melt. A glacier module washas been incorporated to account for glacier melt in glaciated catchments. PREVAH was calibrated (with even years between 1985- and 2014) and validated (uneven with odd years between 1985- and 2014) for each of the 93 catchments individually. Using observed discharge for calibration may put too much emphasis on overemphasize high-flow conditions and potentially overestimates overestimate low-flow conditions. Therefore, the calibration was simultaneously performed on four observational-groups: observed of observations: daily discharge measurements, inverted transformed daily discharge, monthly mean runoff, and the annual volume. This ensures good performance for the general catchment response to meteorological forcing as well as and for the discharge volume. Also low flows are represented in a satisfactory performance. The hydrological model is driven withby daily temperature and precipitation data for each catchment separately from the new high-resolution (2 by 2 km) climate change scenarios for Switzerland CH2018 (CH2018, 2018) for each catchment separately.). In nonglaciated nonglaciated catchments, the, land use was assumed to be constant over the simulation period. In glaciated catchments, the glaciated area was updated every 5 years in line with glacier projections by Zekollari et al. (2019) that), which were driven by the same climate model chains. Land use in areas where glaciers disappear during the simulation period were replaced by bare soil for areas below 3000 maslm.a.s.l. and by rock for areas above 3000 maslm.a.s.l. The Hydro-CH2018-Runoff ensemble includes simulations for three different emission pathways: Representative Concentration Pathways (RCPs): RCP2.6, RCP4.5, and RCP8.5. Because the number of available simulations per emission scenario differs, we We constrained our analysis to the RCP8.5 pathway (Moss et al., 2010; van Vuuren et al., 2011) wherefor three reasons. Firstly, the RCP8.5 pathway is considered the worst-case scenario, and simulations based on this RCP pathway are expected to cover the full range of changes likely to occur. Secondly, the changes in low and high flows from these simulations are more pronounced and thus easier to interpret than simulations based on the RCP2.6 and RCP4.5 pathways, although the direction of change in most catchments is the same. Finally, the RCP8.5 pathway includes the largest number of simulations is available, model chains. The larger the number of simulations within an emissions pathway, the more robust are the results, and this is particularly relevant for analysing the

time of emergence. In total, 20 daily simulations under the RCP8.5 emission pathway for the period 1981–2099 are available for each of the 93 catchments. Table 1 shows the climate model combinations used in this study.

3 Methods

The analysis focuses on moderate low flows and moderate high flows. Several indicators for low-_flow analysis exist focusingfocus on differentvarious properties of low flows (Tallaksen and Van Lanen, 2004). For low flowflows, we use the minimum 7-day moving average runoff (MAM7) within an extended seasona half-year or a year. This indicator is proposed by the Swiss Federal Office for the Environment (FOEN) for low-_flow statistics. The 30-year average of MAM7 is then considered as moderate low flow and used to assess changes in moderate low flows under climate change. For moderate high flows, we use the 30-year average of the annual maxima per extended season-half-year or year as a moderate high-_flow indicator. The seasons are defined as extended-summer half-year (May to October) and extended-winter half-year (November to April) season-). The seasonal distinction is motivated byused because low and high flows in the fact that winter and summer low flowshalf-years are governed by different processes and that they have different impacts. The indicators for the extended winter season and extended summer season will be referred to as the lowest and high flows while the indicators for the extended winter season and extended summer season will be referred to as the lowest and highest seasonal flows, respectively.

PercentPercentage changes are calculated <u>for each simulation</u> as the relative difference between the 30-year mean <u>offor</u> the future period (2070–2099) and the 30-year mean <u>infor</u> the reference period (1981–2010) for each simulation.). The <u>multi-modelmultimodel</u> median <u>for 20 simulations</u> of the relative changes by <u>the</u> end of the century is regarded as the best estimate. To <u>get an indication of indicate</u> the robustness of the projected changes, catchments are highlighted in the <u>Figuresfigures</u> when at least 90% of the simulations show the same direction of change-(<u>. This corresponds to "very likely" in the terminology of the Intergovernmental Panel on Climate Change (IPCC) that changes in moderate low and high flows are either positive or negative (Mastrandrea et al., 2010).</u>

To evaluate potential changes in the seasonality the day of the year for each event (low-flow and high-flow) event is extracted. <u>SinceBecause</u> moderate low flows are calculated from 7-day averages, the last day of the 7-day period is considered as the day of the low-flow event. Median seasonality is then derived by transforming the day of the year into angular values and by applyingcalculate the circular statisticsmedian. Finally, the angular values are transformed back to the day of the year. To assessThe time of emergence is defined as the time when significant changes in the distribution of moderate low-flow and high flows occur, the time of emergence is used (Mahlsteinflow magnitudes emerge from natural variability (Giorgi and Bi, 2009; Leng et al., 20112016). For each simulation and catchment, moderate low-flow and high-flow magnitude distributions of moving 30-year windows are tested against the 30-year reference period using thea Kolmogorov-Smirnov test. This test has been found to result in a more robust and earlier estimation of time of emergence than other methods (Gaetani et al., 2020). The time of emergence is then defined following Mahlstein et al. (2011) as the last year of the first 30-year moving window where the Kolmogorov-Smirnov test is rejected with a *p*-value lower than 0.05 (95% significance). We highlight<u>consider</u> the time of emergence<u>robust</u> when at least 66% of the models detect a significant change in the same 30-year window for the first time. Note that the time of emergence may not necessarily be stable over time. This threshold was also used by Mahlstein et al. (2011) and is referred to as "likely" in the IPCC terminology (Mastrandrea et al., 2010). The testing procedure described above also has some disadvantages. Because runoff responses to climate change are subject to natural climate variability, runoff responses may not always show significant changes after the first detection of a time of emergence, so changes in moderate low and high flows may not be significant in all subsequent periods. Therefore, we also analyse the behaviour of the *p*-values over time (see supplement of this paper) to gauge the persistence of the significance of changes.

Changes in the frequency of moderate low and high flows are quantified by counting years when a pre-definedpredefined runoff threshold is exceeded or undercut. We use the median magnitudemagnitudes of moderate low and high flows in the reference period as thresholdthresholds. For moderate high flows_a we count years with high flows exceeding this threshold. For low flows_a we consider years with low flows below the threshold. This was done is calculated separately for each seasonal and annual time window and each simulation-separately. Finally, the percentual change in occurrence frequency is calculated. We also investigate the co-occurrence of moderate low and high flows. Co-occurrence is considered whendefined as high flows exceeding the reference threshold and low flow flows undercutting the reference threshold occur in the same time window (year, extended half-year, or summer)- half-year.

4 Results

4.1 Future changes in moderate low flows

Median seasonal occurrenceseasonality of moderate annual low flows is shown in Fig. <u>22 (left panels)</u> for the reference period (Fig. 2a) and by <u>the</u> end of the century (Fig. 2e). In <u>most</u> Alpine catchments, annual low flows occur in late winter or early spring in the reference period. By <u>the</u> end of the century, low flows occur in autumn. However, low flows in very high Alpine catchments do not change their seasonality. Median seasonal occurrenceseasonality of low flows in pre-Alpine catchments shifts from late autumn to early autumn. In southern Alpine catchments, <u>the median seasonal occurrencechanges</u> from winter and spring to early autumn. No clear change in seasonality is found for lower <u>-lying catchments with low flows occurring</u>, <u>which occur</u> in late summer and early autumn. <u>DespiteExcept</u> in very high Alpine catchments low flows occur between August and October by <u>the</u> end of the century.

The moderate annual low flows_flow magnitudes show distinctly different patterns of change in magnitude for Alpine and non-Alpine catchments (Fig. 3 left panels). Please note that the The scale bar is limited to -60% and +60% for readability. While Whereas the magnitude of annual low flows (Fig. 3a) decrease decreases by up to -66% in most of the lower-_lying catchments (68 out of 93 catchments in total), the Alpine catchments (25 catchments with mean altitude above 1500 masl) show strong increases in magnitude (up to +200%). Lowest-Low flows in the winter flowshalf-year in Alpine catchments

coincide with the typical low-_flow season in the reference period while lowest, but low flows in the summer flows inhalfyear coincide with the typical low-_flow season in lower-_lying catchments. Lowest-Low-flow magnitudes in the winter flowshalf-year increase on average by +22%. An increase isIncreases are found in two thirds of the catchments, again with stronger increases in very high Alpine catchments-(Fig. 3e), In the summer half-year, the lowest flowslow-flow magnitudes decrease on average by -40% (maximum decrease -74%) (Fig. 3e). However, three high Alpine catchments still show an increase in lowestlow-flow magnitudes in the summer flowshalf-year due to an increase in lowestlow flows in late spring (May). The. More catchments show good model agreement (>90%) is stronger-in_the summer half-year (87 catchments) than in the annual (63) low flows and in the winter half-year (30)-low flows.).

Transient changes of moderate low-flow intensitymagnitudes and seasonality throughout the 21st century are shown for three representative Alpine catchments and three representative lower-lying catchments are shown-in Figs. 4 and 5, respectively. The relative changes for each of the catchments and each of the time windows are summarized in table 2. The high Alpine catchment, Rosegbach is highly glaciated in the reference period with a glacier coverage of 22% but loses most of the glacier coverage by end of the century (glacier coverage: 1%; not shown). The catchment, shows strong increases in both annual low-flows-flow magnitudes and lowest seasonal flows-low-flow magnitudes in the two half-year periods (Fig. 4 top row). While Although the seasonality of the annual low flows and lowest-low flows in the winter flowshalf-year does not change, the occurrenceseasonality of lowestlow flows in the summer flowshalf-year shifts from early summer to autumn. This indicates a change in the underlying processes, leading to lowest-low flows in the summer flows; half-year: The catchment is highly glaciated in the reference period, with a glacier coverage of 22%, but loses most of the glacier coverage by the end of the century (glacier coverage: 1%). In the reference period, the retention of water in snow inand ice still takes place in May. Under climate change, enhanced snowmelt increases runoff in early summer. Towards the end of the century Consequently, the contribution of snow and glacier melt in summer decreases. At the same time, precipitation in summer will decrease also decreases. This leads to a strong decrease in summer runoff. The combination of increasing runoff in early summer and decreasing runoff in late summer and early autumn results in a shift of lowest-low-flow seasonality in the summer flowshalfyear from early summer to early autumn. In the Kander catchment Kander with only, which has little glacier influence (Fig. 4 middle row), magnitudes of the annual low flows and low flows and lowestin the winter flowshalf-year increase mainly due to enhancedincreased winter precipitation falling as rain instead of snow. LowestLow-flow magnitudes in the summer flowshalf-year decrease by the end of the century. The annual low flows and lowest seasonallow flows in the half-year periods occur earlier by the end of the century. The nivalsnow influenced catchment, Plessur, shows a strong shift in seasonality in the annual low flows from winter to autumn but no change in the magnitude of annual low flows (Fig. 4 bottom row). Lowest Low-flow magnitudes in the winter flowshalf-year increase and low flows occur earlier in the season-while. The magnitudes of low flows in the lowest summer flowshalf-year decrease but do not change seasonality- does not change. The magnitude of annual low flows and lowest seasonallow flows in the half-years decrease by the end of the century in the pre-Alpine snow and rain influenced dominated catchment, Emme. The annual low flows show a clear shift in seasonality from late autumn/ and early winter to early autumn (Fig. 5 top row). A shift towards earlier occurrences easonality is also found for lowest-low flows

in the winter flowshalf-year but not for low flows in the lowest-summer flowshalf-year. In the mainly rain_ and evaporationdriven catchment, Venoge, the annual low flows and lowestlow flows in the summer flowshalf-year do not change their seasonality, but lowestlow flows in the winter flowshalf-year tend to occur earlier in the season (Fig. 5 middle row). The magnitude of annual low flows and lowest seasonallow flows in the two half-years decrease. The southern Alpine catchment, Verzasca, shows a decrease in the-magnitude and a strong shift in occurrenceseasonality from late winter to early autumn for annual low flows (Fig. 5 bottom row). LowestLow-flow magnitude in the winter flowshalf-year increase while lowestand lowflow magnitude in the summer flows decreasehalf-year decreases, both without change in the seasonality.

<u>Fig.</u>Figure 6 (left panels) shows the time of emergence for moderate low flows and <u>lowest seasonal flowslow flows in the two</u> <u>half-year periods; this is</u> when at least 66% of the models show significant changes in the distribution. <u>In total</u>, 43 catchments show a time of emergence for annual low flows, with particularly early significant changes in glaciated and/or high Alpine catchments (earliest 2018–2047). <u>TheA total of</u> 20 catchments <u>showingshows</u> a time of emergence in <u>lowestlow flows in the</u> winter <u>flows-half-year</u>, and these have a mean altitude higher than 1600 <u>maslm.a.s.l</u>. In <u>the</u> summer <u>half-year</u>, 80 catchments show significant changes in <u>lowestlow</u> flows, with an early time of emergence again found in high Alpine catchments, <u>alsoa later time of emergence in</u> lower–lying catchments. <u>Most of the catchments</u> show <u>persistent significant changes after the</u> <u>first detection of</u> a time of emergence <u>later(Fig. S1</u> in the <u>century.-supplement)</u>.

4.2 Future changes in moderate high flows

The median seasonal occurrenceseasonality of annual high flows is shown in Fig. 2 for the reference period (Fig. 2b) and by the end of the century (<u>Fig. 2d</u>). In Alpine catchments, the median seasonal occurrenceseasonality shifts from summer to late spring and early summer. However, highly glaciated catchments do not change their high-_flow seasonality. Moderate high flows in pre-Alpine catchments occur in spring in the reference period and in winter in future.by the end of the century. A change in seasonality is also found in southern Alpine catchments, where high-flows shift-low seasonality shifts from late summer and early autumn to late autumn in future. In lower-_lying catchments, no change is found in high-_flow seasonality-is found.

Relative changes of magnitude for moderate-high flows by the end of the century are depicted in Fig. 3 (right panels). The 30-year meansmagnitudes of annual high flows (Figure 3b) increase in 71 catchments (up to +28%) and decrease in 22 catchments (up to -22%). Compared to the changes in moderate-low flows, the magnitude of changechanges in high-flows is -flow magnitudes are smaller. There are no clear spatial patterns or elevation dependences, and good model agreement (>90%) is only found in 12 catchments. Highest winter-High flows in the winter half-year in lower--lying catchments coincide with the typical high--flow season in the reference period, while highestwhereas high flows in higher elevationAlpine catchments are mainly found in the summer half--year. Highest-High-flow magnitudes in the winter flowshalf-year increase in all catchments, and model agreement is higher, with 45 (out-of 93) catchments showing a-good agreement-(Fig. 3d). Strongest increase. The strongest increases in magnitude and good model agreement are found in high Alpine catchments. However, highest-high flows

in the winter flowshalf-year in high Alpine catchments are still small in magnitude. HighestHigh flow magnitudes in the winter flowshalf-year in the-lower--lying catchments increase only moderately, and model agreement is generally weak. Highest High-flow magnitudes in the summer flows (Fig. 3f)half-year decrease in 74 catchments (up to -26%) and increase in 19 catchments (up to +15%). StrongestThe strongest reductions in highesthigh-flow magnitudes in the summer flowshalf-year are found in high Alpine catchments, including the only-six catchments showing good model agreement. A spatial cluster of increasing highesthigh-flow magnitudes in the summer flowshalf-year is found in the Jura mountains (catchments in north-westwestern Switzerland)-.

Annual<u>The magnitude of annual</u> high flows and highestof high flows in the summer flowshalf-year in the Rosegbach catchment decrease towards the end of the century and tend to occur earlier in summer-while highest-, but high-flow magnitudes in the winter flowshalf-year increase and tend to occur more often later in the season (FigureFig. 7 top row). A similar pattern is also found for the Plessur (FigureFig. 7 bottom row). In the Kander, the annual high flows-flow magnitudes increase slightly-and, shift to earlier in the year, and can also occur in the winter half-year by the end of the century (FigureFig. 7 middle row). Also, highest-High-flow magnitudes in the winter flowshalf-year in the Kander also increase, and highest-high-flow magnitudes in the summer flowshalf-year show a small decrease without a significant shift in the occurrence-seasonality. The high flows in the Emme and the Verzasca do not change their seasonality, but highesthigh-flow magnitudes in the winter flowshalf-year (Fig. 8 top and bottom rows). The pluvialrain driven catchment, Venoge, shows increasing moderate annual high flows and seasonal highest flowshigh-flow magnitudes with no change in the seasonality (Fig. 8 middle rows).

The time of emergence of moderate high flows is depicted in Fig. 6 (right panels). Compared to moderate low flows, there are fewer catchments exhibitingexhibit significant changes, and these catchments are mostly high Alpine catchments. For annual high flows, three high Alpine (>2000 masl)m.a.s.l.) catchments show a time of emergence, with of the earliest time of emergence of in 2078 (2049–2078). The 27 catchments showing significant changes in highest-high flows in the winter flowshalf-year (earliest 2044, 2013–2044) are also located in the Alpine ridge (>1500 maslm.a.s.l. mean altitude). In the summer half-year, only six catchments (>1800 masl)m.a.s.l.) show a time of emergence, and the earliest time of emergence is 2071 (2042–2071). Most of the few catchments showing a time of emergence also show persistent significant changes for the rest of the century a few years after the first detection (Fig. S2 in the supplement). If the time of emergence analysis revealed a significant deviation from current high- and low-flow magnitudes in a catchment, then this change was most often persistent over time. However, the number of catchments showing a time of emergence was found to be small.

4.3 (Co-)Changes in frequency and co-occurrence of low- and high flows-flow events

So far, we have assessed changes in the magnitude and seasonality of low and high flows, in. In this section, we address changes in frequency and the frequency and co-occurrence of low-flow and high and flow events. For To do this, we need to set a threshold to identify events. The threshold dischargerunoff value is defined as a value occurring every second year in the reference period (i.e., median in the reference period).

Fig.Figure 9 illustrates relative changes in the occurrence frequency of low flows (Fig. 9 a - eupper panels) and high flows (Fig. 9 d-fmiddle panels) by the end of the century (2070-2099) that fall below resp.or exceed median values. For annualCooccurrence of low and high flows, are defined as the occurrence of low-a high-flow event and a low-flow event in the same time window (lower panels). The frequency of annual low-flow events increases in 70 catchments. These catchments are mainly found in rain dominated areasdriven catchments and to a lesser extent also in snow dominated areasdriven catchments. Catchments showing fewer occurrences less frequent low flows are only found in high Alpine areas. Good model agreement is found in 56 catchments. LowestLow flows also occur also-more oftenfrequent in summer-(, sometimes occurring every year), in almost all regions except few very high--elevation catchments. Most of the catchments (82) show a good model agreement on the increase of lowestlow-flow frequency in the summer flow occurrence. Lowesthalf-year. Low flows in the winter flowshalf-year tend to occur less often in mountainous areas-while, but lower--lying catchments still show an increase in occurrence/frequency. However, only mountainous catchments show good model agreement.

Changes in the <u>occurrencefrequency</u> of high flows are less clear than for low flows. For annual high flows, 58 catchments show increasing <u>occurrencesfrequency</u>. However, the changes are often small. <u>Also, noNo</u> clear spatial or elevation pattern emerges, and model agreement is weak. For the highest winter flows, all<u>All</u> catchments will face more years with more frequent high_flow events in the winter half-year than they do today, particularly in the high Alpine regions. In contrast, the <u>occurrencefrequency</u> of <u>highest summer high</u>-flow events in the summer <u>half-year</u> will decrease by the end of the century in most catchments. Model agreement is weaker in the summer <u>half-year</u> than in the winter half-year.

Figure 9 (g i) shows changes in the co-occurrence of moderate high and low flows defined as the occurrence of a high flow event and a low flow event in the same time window. Annual co-occurrence increases in most catchments, particularly in the lower_lying catchments. In high Alpine catchments, this co-occurrence decreases mainly due to the strong increase in winter runoff. Winter coCo-occurrence in the winter half-year decreases mainly in high altitude catchments but also in a few of the lower_lying catchments. In the summer half-year, most catchments (85 catchments) show increasing co-occurrence by end of the century. Only <u>seight</u> high Alpine catchments show decreasing co-occurrence. In contrast to high—flow occurrence frequency, the model agreement for co-occurrence is stronger in the summer half-year (48 catchments) than in the winter half-year (14) co-occurrence. catchments).

5 Discussion

5.1 Changes in moderate low flows

Low flows in Alpine regions typically occur in winter and early spring when precipitation falls as snow and accumulates. The storageStorage as snow limits the direct runoff, and only little runoff (baseflow) runoff occurs in winter. SinceBecause higher temperatures result in both a change of the precipitation type (more precipitation fallsfalling as rain instead of snow) and an

enhancement of earlier snow melt, the lowest-low-flow magnitudes in the winter flowshalf-year are projected to increase-in the future. Also, the seasonal occurrence. Furthermore, seasonality shifts from late winter to late autumn. This shift indicates that snow storage no longer dominates low flows, but. Instead summer droughtand autumn droughts in combination with lack of snow and glacier melt becomesgenerally become the main driver of low flows. However, this is not the case in highly glaciated catchments with very high mean altitudes, (>2300 m.a.s.l.), where the seasonality of low flows does not change. ConsideringIn the summer half-_year, the lowest summermagnitudes of low flows in Alpine catchments decrease due to the combination of decreasing summer precipitation, enhanced evapotranspiration_a and the reduced contribution of snow and glacier melt to the runoff. An exceptionExceptions are the catchments inat very high altitudes, which show increasing lowestlow-flow magnitudes in the summer flowshalf-year by the end of the century. Increasing lowestlow-flow magnitudes in the winter flowshalf-year in Alpine areas have been identified in observations (Weingartner and Schwanbeck, 2020), and our results show that this trend eontinueswill continue and intensify with climate change. The findings are also in agreement with results for projections of very extreme low-flow regimes (100-year return period; Brunner et al., 20192019a).

In the present climate, low flows occur mostly in late summer and autumn in lower--lying catchments. In these catchments, runoff volumes during low-_flow conditions are projected to decrease in all time periods, with the reduction in the summer half -year being much stronger than in the winter half--year. The reasons for the reduction in the summer half-year are the decreasing summer precipitation and the higher temperatures enhancing, which enhance evapotranspiration. The projected lowest summer low-flow reduction in the summer half-year is in line with observed trends (Weingartner and Schwanbeck, 2020)), but the changes getare amplified under climate change. Even though the climate change scenarios project increasing winter precipitation, the lowest-magnitudes of low flows in the winter flows-half-year are projected to decrease, mainly due to a shift in the occurrence-seasonality from winter to late autumn. The seasonality of annual low flows shifts from late autumn to early autumn.

In the <u>The</u> snow- and rain-<u>driven influenced</u> southern Alpine regions, there are typically <u>undergo</u> two periods of low flows: one in late summer and one in winter, with the winter minimum often being lower in the reference period. Under climate change, the <u>seasonal occurrenceseasonality</u> of low flows shifts from winter to late summer and early autumn. At the same time, runoff in low-<u>flow situations decreases by the</u> end of the century.

Increasing lowest winter flows in Alpine regions may be beneficial for energy production, but the decreasing lowest summer flows may have severe impacts in agricultural regions where water is needed for irrigation. Also, the decreasing water availability during low flows may have implications on the cooling of infrastructures and in combination with increasing water temperatures may foster water stress for ecosystems.

The projected changes in low flows imply potential impacts for various sectors. Changes in low flows in the winter half-year have different impacts than those in the summer half-year. The increase in mean runoff (Muelchi et al., 2021b) and in low flows in the winter half-year in Alpine regions may be beneficial for hydropower production, which is particularly sensitive to climate change (Schaefli et al., 2007; Zierl and Bugmann, 2005). Electricity demand in Switzerland is highest in winter (Hakala

et al., 2020). Even though heating degree days are projected to decrease in Switzerland (Tschurr et al., 2020), heating demand will still occur under climate change. The projected increase in winter mean runoff and low flows may allow enhanced hydropower production and possibly extend the season of hydropower production. Therefore, changes in mean runoff and low flows due to climate change should be considered in negotiations of future water use concessions (Hakala et al., 2020). However, energy demand may increase in summer due to a projected increase in cooling degree days due to higher summer temperatures (Tschurr et al., 2020). At the same time, water demand for irrigation may increase due to projected prolonged and intensified summer droughts and increased evapotranspiration (CH2018, 2018). In addition, river runoff is needed for cooling infrastructure. Because mean flows (Muelchi et al., 2021b) and low flows in the summer half-year are projected to decrease, this combination may lead to conflicts in water use. Policy makers and water management systems need to adapt to these changes in water availability, and a coordinated water use policy should be introduced to cover potential water shortages in the summer half-year. In recent years, the regional authorities in Switzerland that control public water have reviewed their laws and recommendations regarding minimum residual flows. Under climate change, these minimum residual flows are projected to decrease in summer in lower-lying catchments, indicating that residual flows defined under current conditions may not suffice to serve important water-related services. One of the measures proposed by the authorities is to increase the residual flow for environmental purposes such as protecting flora, fauna, and sediment transports (Hakala et al., 2020). In Alpine regions, an increase in the residual flow regulations may create potential deficits for hydropower providers. This may reinforce potential conflicts in water use during the summer half-year. The combination of a decrease in low flows in the summer half-year and increasing river water temperatures (Michel et al., 2021) increases water stress for river ecosystems, and some fish species such as trout may lose their habitats (FOEN, 2021).

5.2 Changes in moderate high flows

In Alpine areas-moderate, magnitudes of annual high flow and highest-high flows in the summer flows will-half-year are likely to_decrease in the projections._ This can be explained by the decreasing contribution of melt water together with decreasing summer precipitation and enhanced_increased evapotranspiration. In futureBy the end of the century, Alpine areas will face about half of the present mean runoff in summer (Muelchi et al., 2020c in review). This2021b), a decrease that is also reflected in moderate-high flows. This is in contrast to the highest-magnitudes of high flows in the winter flowshalf-year, which are projected to increase with climate change. However, highest-high-flow magnitudes in the winter flowshalf-year are still smaller in magnitude than highest-high-flow magnitudes in the summer flowshalf-year. Decreases in the runoff volume during highest high flows in the summer flowshalf-year and increases in highest-high flows in the winter flowshalf-year were also found for mountainous regions by Allamano et al. (2009) for mountainous regions... However, the decreasing annual high-flows-flow magnitudes in the Alpine areas. The reason for the different difference between results is not econclusively entirely clear, but in Köplin et al. (2014) consider very extreme floods-are considered, while, whereas this study considers moderate high flows.

Annual high flows in the Alpine region usually occur in <u>the</u> summer <u>half-year</u>, when the snow line is high, melting <u>processes</u> areis in progress, and precipitation intensities are <u>largesthighest</u>. In glaciated catchments, high flows <u>currently</u> occur at the end of summer when glacier melt reaches its peak-<u>in the reference period</u>. In snow <u>influenceddriven</u> catchments today, the high flows tend to occur in early summer during the snowmelt. In both regime types, <u>seasonal occurrence is shiftedseasonality shifts</u> to earlier months <u>suchby the end of the century so</u> that the high flows occur earlier in summer<u>-in-future</u>. An exception, <u>Exceptions</u> are <u>the</u> highly glaciated catchments with high mean elevation, which will also have snow and glacier influence in summer in the future. In these catchments, <u>the seasonal occurrenceseasonality</u> hardly changes. Köplin et al. (2014) also found shifts in the <u>occurrenceseasonality</u> of extreme floods in Alpine areas. Their results show a shift in <u>nivalsnow driven</u> catchments from summer to autumn, whereas our results show a shift to earlier spring and early summer.

In lower-lying areas-<u>magnitudes of annual high flowflows</u> and <u>highest-high flows in the</u> winter flowshalf-year tend to increase, although the increase is often not robust across models. In <u>the</u> summer <u>half-year</u>, the <u>highest flowshigh-flow magnitudes</u> tend to decrease again with no robust signals across models. Moderate high flows occur in <u>the</u> winter <u>half-year</u> in <u>pluvialrain driven</u> catchments, and this will not change <u>in future</u>. In catchments partly influenced by snow, where high flows occur in spring, the <u>sesasonal occurrence is shiftedseasonality shifts</u> from spring to late winter. This <u>behaviorshift</u> is in agreement with <u>the results</u> of Köplin et al.'s (2014)-) results. In the southern Alpine areas, the annual high flows also tend to increase and <u>will</u>-shift from late summer and early autumn to late autumn, <u>which wasa change</u> also found by Köplin et al. (2014) for extreme floods.

The increased water availability in winter in Alpine regions may be beneficial for energy production. But increasing high flows in mainly lower lying catchments may increase the potential of flood damages. However, this increase is not robust among the climate models and moderate high flows only partially reflect severe floods. The increase in moderate high flows in the winter half-year may also be beneficial for energy production because today's hydropower production capacity is limited in winter by the snow storage capacity within Alpine catchments. Under climate change, the increasing moderate high flows in Alpine catchments may create opportunities to help meet energy demand in the winter half-year. Reservoirs may be emptied more slowly due to the increase in mean flows (Muelchi et al., 2021b) and moderate high flows. In summer, the moderate high flows in Alpine catchments are projected to decrease. In lower-lying catchments, where the high-flow season is in the winter half-year, increasing moderate high flows may also indicate an increased risk of winter floods. In contrast to the results for moderate low flows, the climate model agreement on the sign of change is weaker, and thus interpretation of our results becomes more difficult.

5.3 Time of emergence of significant changes

Significant and robust changes in the magnitude of moderate low flows emerge mainly for annual low flows and <u>lowest-low</u> <u>flows in the summer flowshalf-year</u>. The majority of the catchments show a significant change in <u>the magnitude for summerof</u> low flows<u>- in the summer half-year</u>. High Alpine catchments show earlier significant changes in <u>lowestlow flows in the</u> summer <u>flowshalf-year</u> than <u>do</u> lower_lying catchments. Early times of emergence in high Alpine catchments were also found for summer mean flow in Muelchi et al. (2020c in review2021b). In the winter half-year, only Alpine catchments show a significant change in lowest winterlow flows. The main reasonreasons for this are snowpack-related processes likesuch as the change in precipitation type (from snow vs.to rain) together with, smaller snow accumulations, and associated enhancedconsequent increased direct runoff.

The magnitude of high flows significantly changes <u>for</u> only <u>for</u> few catchments. This is due to the large variability across the climate models. To detect a time of emergence, we require <u>that</u> at least 66% of the models <u>to</u> agree on significant changes in the distribution of high flows.

5.4 Changes in the (frequency and co-)-occurrence of low- and high-flow events

The frequency of annual moderate low-_flow events increases in lower-_lying catchments, whilebut fewer low flows_ flow events are detected in Alpine catchments. However, the frequency of <u>low flows in the lowest-summer flowshalf-year</u> will increase in almost all catchments. In some catchments, thethis frequency almost doubles. This may have implications in agricultural areas where irrigation plays an important role. High-_flow events in the_winter half-year will occur more often, while_but high-flow events in the_summer high_flow_eventshalf-year will occur less often. A clear pattern in occurrencefrequency of annual high-_flow events is not detectable_cannot be shown because model agreement is weak. However, most catchments show a tendency towards more occurrences.higher frequency. Co-occurrence of low-flow and high _flow events in the same year increases in most lower-_lying catchments. In contrast, high-_elevation catchments show **a** decreasing co-occurrence, mainly due to the increase in low flows. The changes in co-occurrence are dominated by changes in low-_flow occurrence. Since low_Low flows in lower-_lying (high Alpine) catchments tend to occur much less (more) often, so co-occurrence also decreases(-increases)-, but the opposite is true for low flows in high Alpine catchments. Co-occurrence of low-flow and high-and low_flow events in the same extended season are-half-year is important for ecosystems since the because this may shorten their recovery time may be shortened.times. Information about the co-occurrence is also important for insurance <u>companies for their companies</u> risk assessments.

5.5 Uncertainties and Limitations

Uncertainties in our results are larger for moderate high flows than for moderate low flows. The larger uncertainties in high flows are due to several reasons. First, high flows are difficult to model since Several sources of uncertainties affect the results of this study. A detailed discussion of sources of uncertainty in the hydrological simulations is provided in Muelchi et al. (2021a), and they are summarized only briefly here. Uncertainties arise during each modelling step: with the emission scenario (RCP8.5), the selection of climate models and their boundary conditions, the postprocessing method (Gutiérrez et al., 2018), the hydrological model (Addor et al., 2014) and its calibration, and the underlying glacier projections. All these need to be considered in adaptation planning (e.g., Wilby and Dessai, 2010). A comparison with three versions of hydrological models

for three catchments showed good agreement on the direction of change of the hydrological response to climate change (Muelchi, 2021). Therefore, we regard the simulations underlying this study as robust for climate change assessments. Our results show that projections of moderate high flows are less robust among climate models than are those of moderate low flows. Differences in the projections of moderate high flows arise for several reasons. First, high flows are difficult to model because many different processes interact with each other. In particular, small-scale precipitation patterns have a strong influence on high flows, and the input data from the climate models does not reflect small-scale precipitation processes well (Ban et al., 2015). Second, the uncertainty arising from internal variability of extreme precipitation is large, and this is thus also reflected in our results. Third, our results represent 30-year averages as well as averages across models. Therefore, a lotgreat deal of information is averaged out. Other sources of uncertainty also affectDespite these limitations, our results suchshow that it is crucial to take projected changes in moderate runoff extremes into account due to their manifold implications for various sectors, as the climate models discussed in sections 5.1 for low flows and their boundary and initial conditions, the post-processing method, the hydrological model and its calibration, and the underlying glacier projections 5.2 for high flows.

6 Conclusions

We assessed 6 Summary and conclusions

<u>This study assesses</u> changes in moderate low and high flows (annual and seasonal flow maxima and minima)under climate change for 93 catchments in Switzerland-under climate change. Runoff simulations were driven by the newestmost recent transient climate change scenarios (CH2018) for Switzerland for 1981–2099 based on the RCP8.5 scenario. This study analyzes analyzes changes not only changes in the magnitude of these moderate low and high flows but also on changes in their seasonal occurrences asonality and their frequency. Thanks to the transient property of the simulations, also the timing of significant changes (the time of emergence) could also be assessed.

The projections indicate the following results. For that changes in low flows, a strong-over time depend strongly on elevation dependence of the changes over time was found. While, Whereas low-flow magnitudes decrease in lower--lying catchments, they increase in Alpine catchments-extending, thus amplifying observed trends in the past (Weingartner and Schwanbeck, 2020). Low flows decrease by -40% in summer, and increase by +22% in winter. The results for low--flow magnitudes are in line with the projections of previous studies (e.g. Meyer et al., 2011; Bernhard and Zappa, 2012; Brunner et al., 2019). A 2019a). Moreover, a shift in seasonality was found for most of the catchments. By end of the century, with low flows will occurroccurring predominantly in late summer and autumn in most of the catchments. This indicates that the lack of precipitation by the end of the 21st century. Increasing low flows in Alpine regions in the winter half-year may be beneficial for hydropower production. However, decreasing low flows in the summer half-year may induce water use conflicts, especially

<u>because water demand</u> in summer exceeds the contribution of other processes may increase in various sectors such as snow and <u>glacier melt contributions.irrigation for agriculture and cooling infrastructures (Brunner et al., 2019b)</u>. The pronounced projected decrease in <u>summer</u>-low flows in <u>most of the summer half-year in almost all</u> the catchments (except some high Alpine eatchments)<u>ones</u> may become one of the most important challenges <u>in terms of for</u> water management. In <u>contrast, increasing</u> winter low flows in Alpine catchments may be beneficial for hydropower production.

For Relative changes in magnitude are smaller for moderate high flows, relative changes are smaller_than for low flows. Most of the catchments show an increase in moderate high-flows-flow magnitudes, but the model agreement on the changes is not robust with the exception of except in a few catchments in northern and north-western Switzerland-and the Jura mountains,. High Alpine catchments show a decrease in high-flow magnitudes in the highest summer flowshalf-year, mainly due to reduced melt water in future, and an increase in high flows in the highest-winter flowshalf-year. The magnitude of winter high flows in the winter half-year in Alpine catchments is much smaller than for those in the summer high flowshalf-year. Thus, the increasing winter high flows high-flow magnitudes in Alpine catchments in the winter half-year are not that important from a hydrological point of viewhydrologically but may become relevant for ecosystems, and energy production. Projected changes in magnitude and shifts in seasonality of moderate high flows in lower_lying catchments are in line with previous studies (e.g. Koeplin et al., 2014; Brunner et al., 20192019a). For Alpine catchments, our results do not agree with other projections in terms of magnitude andor in some cases in terms of seasonality. This contradictionlack of agreement may arise due tofrom the different yarious indicators considered. While ourOur study focuses on moderate high flows, the otherbut comparative studies focused on extreme high flows, which can be governed by different processes than moderate high flows.

Significant changes in the magnitude of low flows emerge early in the 21st century for high Alpine catchments because of an increase in winter flows. For many lower–lying catchments, a significant decrease in summer–low–flow magnitude in the summer half-year is detected but only later in the 21st century. Changes in the magnitude of high flows are mostly not robust across climate models and thus not significant.

Low-_flow events will occur more often in lower-_lying catchments and less often in high Alpine catchments. Like the weak signal in the magnitude of high flows, also changes in the occurrence frequency of high-_flow events are also small. However, most of the catchments will experience an increasing frequency in the occurrence of increasingly frequent high-_flow events. An elevational pattern was found for the co-occurrence of moderate low and high-_flow events, with increasing co-occurrence in lower-_lying catchments and decreasing co-occurrence in high Alpine catchments. This pattern is dominated by changes in the frequency and magnitude of moderate low flows.

Data availability

The data used in this study is available under <u>https://doi.org/10.5281/zenodo.3937485</u>https://doi.org/10.5281/zenodo.3937485 (Muelchi et al., <u>2020a2020</u>).

Author contributions

RM performed the analysis of the results and drafted the manuscript. JS, OR, RW_a and OM helped in <u>the</u> interpretation of the results. All authors reviewed the resulting data and assisted with paper writing.

Competing interests

The authors declare that they have no conflict of interest.

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References

Addor, N., Rössler, O., Köplin, N., Huss, M., Weingartner, R., and Seibert, J.: Robust changes and sources of uncertainty in the projected hydrological regimes of Swiss catchments, Water Resour. Res., 50, 7541-7562, https://doi.org/10.1002/2014WR015549, 2014.

Alderlieste, M. A., Van Lanen, H. A., and Wanders, N.: Future low flows and hydrological drought: how certain are these for Europe?, in: Hydrology in a Changing World, vol. 363, pp.60–65, 2014.

Allamano, P., Claps, P., and Laio, F.: Global warming increases flood risk in mountainous areas, Geophysical Research Letters, 36, https://doi.org/10.1029/2009GL041395, 2009.

Ban, N., Schmidli, J., and Schär, C.: Heavy precipitation in a changing climate: Does short-term summer precipitation increase faster?, Geophysical Research Letters, 42, 1165–1172, https://doi.org/10.1002/2014GL062588, 2015.

Beniston, M. and Stoffel, M.: Rain-on-snow events, floods and climate change in the Alps: Events may increase with warming 4 С and decrease thereafter. Science of the Total Environment. 571. 228-236. up to https://doi.org/10.1016/j.scitotenv.2016.07.146, 2016.

Bernhard, L. and Zappa, M.: Natürlicher Wasserhaushalt der Schweiz und ihrer bedeutendsten Grosseinzugsgebiete, Comissioned by the Swiss Federal Office for the Environment, 2012.

Bertola, M., Viglione, A., Lun, D., Hall, J., and Blöschl, G.: Flood trends in Europe: are changes in small and big floods different?, Hydrology and Earth System Sciences, 24, 1805–1822, https://doi.org/10.5194/hess-24-1805-2020, 2020.

Birsan, M.-V., Molnar, P., Burlando, P., and Pfaundler, M.: Streamflow trends in Switzerland, Journal of Hydrology, 314, 312–329, https://doi.org/10.1016/j.jhydrol.2005.06.008, 2005.

Blöschl, G., Hall, J., Viglione, A., Perdigao, R. A., Parajka, J., Merz, B., Lun, D., Arheimer, B., Aronica, G. T., Bilibashi, A., et al.: Changing climate both increases and decreases European river floods, Nature, 573, 108–111, https://doi.org/10.1038/s41586-019-1495-6, 2019.

Brunner, M. I., Farinotti, D., Zekollari, H., Huss, M., and Zappa, M.: Future shifts in extreme flow regimes in Alpine regions, Hydrology and Earth System Sciences, 23, 4471–4489, https://doi.org/10.5194/hess-23-4471-2019, 20192019a.

Brunner, M. I., Gurung, A. B., Zappa, M., Zekollari, H., Farinotti, D., and Stähli, M.: Present and future water scarcity in Switzerland: Potential for alleviation through reservoirs and lakes, Science of the Total Environment, 666, 1033-1047, https://doi.org/10.1016/j.scitotenv.2019.02.169, 2019b.

Castellarin, A. and Pistocchi, A.: An analysis of change in alpine annual maximum discharges: implications for the selection of design discharges, Hydrological Processes, 26, 1517–1526, https://doi.org/10.1002/hyp.8249, 2012.

CH2018: CH2018 – Climate Scenarios for Switzerland, Technical report, National Centre for Climate Services, Zurich, Switzerland, 271 pp., ISBN: 978-3-9525031-4-0, 2018.

Feyen, L. and Dankers, R.: Impact of global warming on streamflow drought in Europe, Journal of Geophysical Research: Atmospheres, 114, https://doi.org/10.1029/2008JD011438, 2009.

FOEN (ed.): Effects of climate change on Swiss water bodies. Hydrology, water ecology and water management, Federal Office for the Environment FOEN, Bern, Environmental Studies No.2101: 125p, 2021.

Forzieri, G., Feyen, L., Rojas, R., Flörke, M., Wimmer, F., and Bianchi, A.: Ensemble projections of future streamflow droughts in Europe, Hydrology and Earth System Sciences, 18,85–108, https://doi.org/10.5194/hess-18-85-2014, 2014.

Gaetani, M., Janicot, S., Vrac, M., Famien, A. M., and Sultan, B.: Robust assessment of the time of emergence of precipitation change in West Africa, Sci. Rep., 10, 1-10, https://doi.org/10.1038/s41598-020-63782-2, 2020.

Giorgi, F., and Bi, X.: Time of emergence (TOE) of GHG-forced precipitation change hot-spots, Geophys. Res, Lett., 36, https://doi.org/10.1029/2009GL037593, 2009.

Gutiérrez, J. M., Maraun, D., Widmann, M., Huth, R., Hertig, E., Benestad, R., Rössler, O., Wibig, J., Wilcke, R., Kotlarski, S., San Martin, D., Herrera, S., Bedia, J., Casanueva, A., Manzanas, R., Iturbide, M., Vrac, M., Dubrovsky, M., Ribalaygua, J., Pórtoles, J., Räty, O.E., Räisänen, J.A., Hingray, B., Raynaud, D., Casado, M., Ramos, P., Zerenner, T., Turco, M.,

Bosshard, T., Stepanek, P., Bartholy, J., Pongracz, R., Keller, D., Fischer, A., Cardoso, R., Soares, P., Czernecki, B., and Pagé, C.: An intercomparison of a large ensemble of statistical downscaling methods over Europe: Results from the VALUE perfect predictor cross-validation experiment, Int. J. Climatol., 39, 3750–3785, https://doi.org/10.1002/joc.5462, 2019.

Hall, J., Arheimer, B., Borga, M., Brázdil, R., Claps, P., Kiss, A., Kjeldsen, T., Kriauciuniene, J., Kundzewicz, Z., Lang, M., et al.: Understanding flood regime changes in Europe: A state of the art assessment, Hydrology and Earth System Sciences, 18, 2735–2772, https://doi.org/10.5194/hess-18-2735-2014, 2014.

Hakala, K., Addor, N., Gobbe, T., Ruffieux, J., and Seibert, J.: Risks and opportunities for a Swiss hydroelectricity company in a changing climate, Hydrology and Earth System Sciences, 24, 3815-3833, https://doi.org/10.5194/hess-24-3815-2020, 2020.

IPCC: Climate Change 2014: Impacts, Adaptation, and Vulnerability. Part A: Global and Sec-toral Aspects. Contribution of Working Group II to the Fifth Assessment Report of the Inter-governmental Panel on Climate Change, 1132 pp., Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA, 2014.

Köplin, N., Sch ädler, B., Viviroli, D., and Weingartner, R.: Seasonality and magnitude of floods in Switzerland under future climate change, Hydrological Processes, 28, 2567–2578, https://doi.org/10.1002/hyp.9757, 2014.

Mahlstein, I., Knutti, R.,Leng, G., Huang, M., Voisin, N., Zhang, X., Asrar, G. R., and Leung, L. R.: Emergence of new hydrologic regimes of surface water resources in the conterminous United States under future warming, Environ. Res. Lett., 11, 11, 114003, https://doi.org/10.1088/1748-9326/11/11/114003, 2016.Mahlstein, I., Knutti, R., Solomon, S., and Portmann, R. W.: Early onset of significant local warming in low latitude countries. Environmental Research Letters, 6, 034,009

R. W.: Early onset of significant local warming in low latitude countries, Environmental Research Letters, 6, 034 009, https://doi.org/10.1088/1748-9326/6/3/034009, 2011.

Mangini, W., Viglione, A., Hall, J., Hundecha, Y., Ceola, S., Montanari, A., Rogger, M., Salinas, J. L., Borzì, I., and Parajka, J.: Detection of trends in magnitude and frequency of flood peaks across Europe, Hydrological Sciences Journal, 63, 493–512, https://doi.org/10.1080/02626667.2018.1444766, 2018.

Marx, A., Kumar, R., Thober, S., Rakovec, O., Wanders, N., Zink, M., Wood, E. F., Pan, M., Sheffield, J., and Samaniego, L.: Climate change alters low flows in Europe under global warming of 1.5, 2, and 3 C, Hydrology and Earth System Sciences, 22, 1017–1032, https://doi.org/10.5194/hess-22-1017-2018, 2018.

Mastrandrea, M.D., Field, C.B., Stocker, T.F., Edenhofer, O., Ebi, K.L., Frame, D.J., Held, H., Kriegler, E., Mach, K.J., Matschoss, P.R., Plattner, G.-K., Yohe, G.W., and Zwiers, F.W.: Guidance Note for Lead Authors of the IPCC Fifth Assessment Report on Consistent Treatment of Uncertainties, Intergovernmental Panel on Climate Change (IPCC), 2010.

Meyer, R., Schädler, B., Viviroli, D., and Weingartner, R.: Klimaänderung und Niedrigwasser -Auswirkungen der Klimaänderung auf die Niedrigwasserverhältnisse im Schweizer Mittelland für 2021-2050 und 2070-2099, Comissioned by the Swiss Federal Office for the Environment, 2011.

Michel, A., Schaefli, B., Wever, N., Zekollari, H., Lehning, M., and Huwald, H.: Future water temperature of rivers in Switzerland under climate change investigated with physics-based models, Hydrology and Earth System Sciences Discussions, 1-45, https://doi.org/10.5194/hess-2021-194, 2021. Moss, R. H., Edmonds, J. A., Hibbard, K. A., Manning, M. R., Rose, S. K., Van Vuuren, D. P., Carter, T. R., Emori, S., Kainuma, M., Kram, T., et al.: The next generation of scenarios for climate change research and assessment, Nature, 463, 747–756, https://doi.org/10.1038/nature08823, 2010.

Muelchi, R., Rössler, O., Schwanbeck, J., Weingartner, R., and Martius, O.: Hydro-CH2018-Runoff ensemble (version v1), Zenodo, https://doi.org/10.5281/zenodo.3937485, 2020a2020.

Muelchi, R.: Future changes of Swiss river runoff and extreme vertically integrated moisture transport, Ph.D thesis, University of Bern, Switzerland, 198 pp., 2021.

Muelchi, R., Rössler, O., Schwanbeck, J., Weingartner, R., and Martius, O.: An ensemble of daily simulated runoff data (1981–2099) under climate change conditions for 93 catchments in Switzerland (Hydro-CH2018-Runoff ensemble), Geoscience Data Journal, 2020b, in reviewhttps://doi.org/10.1002/gdj3.117, 2021a.

Muelchi, R., Rössler, O., Schwanbeck, J., Weingartner, R., and Martius, O.: <u>FutureRiver runoff in Switzerland in a changing</u> <u>climate -</u> runoff regime changes and their time of emergence-for <u>93 catchments in Switzerland</u>, Hydrology and Earth System Sciences-<u>Discussions</u>, pp. 1–25, <u>https://doi.org/10.5194/hess 2020 516, 2020c</u>, in review, <u>2021b</u>.

Papadimitriou, L. V., Koutroulis, A. G., Grillakis, M. G., and Tsanis, I. K.: High-end climate change impact on European runoff and low flows–exploring the effects of forcing biases, Hydrology and Earth System Sciences, 20, 1785–1808, https://doi.org/10.5194/hess-20-1785-2016, 2016.

Ragettli, S., Tong, X., Zhang, G., Wang, H., Zhang, P., and Stähli, M.: Climate change impacts on summer flood frequencies in two mountainous catchments in China and Switzerland, Hydrology Research, nh2019118, https://doi.org/10.2166/nh.2019.118, 2019.

Schaefli, B., Hingray, B., and Musy, A.: Climate change and hydropower production in the Swiss Alps: quantification of potential impacts and related modelling uncertainties, Hydrol. Earth Syst. Sci., 11, 1191–1205, https://doi.org/10.5194/hess-11-1191- 2007, 2007.

Schmocker-Fackel, P. and Naef, F.: More frequent flooding? Changes in flood frequency in Switzerland since 1850, Journal of Hydrology, 381, 1–8, https://doi.org/10.1016/j.jhydrol.2009.09.022, 2010a.

Schmocker-Fackel, P. and Naef, F.: Changes in flood frequencies in Switzerland since 1500, Hydrology and Earth System Sciences, 14, 1581–1594, https://doi.org/10.5194/hess-14-1581-2010, 2010b.

Stahl, K., Hisdal, H., Hannaford, J., Tallaksen, L., Van Lanen, H., Sauquet, E., Demuth, S., Fendekova, M., and Jordar, J.: Streamflow trends in Europe: evidence from a dataset of near-natural catchments, Hydrology and Earth System Sciences, 14, 2367–2382, https://doi.org/10.5194/hess-14-2367-2010, 2010.

Stahl, K., Tallaksen, L. M., Hannaford, J., and Van Lanen, H.: Filling the white space on maps of European runoff trends: estimates from a multi-model ensemble, Hydrology and Earth System Sciences, 16, 2035–2047, https://doi.org/10.5194/hess-16-2035-2012, 2012.

Tallaksen, L. M. and Van Lanen, H. A.: Hydrological drought: processes and estimation methods for streamflow and groundwater, vol. 48, Elsevier, 2004.

Tschurr, F., Feigenwinter, I., Fischer, A. M., and Kotlarski, S.: Climate Scenarios and Agricultural Indices: A Case Study for Switzerland, Atmosphere, 11, 535, https://doi.org/10.3390/atmos11050535, 2020.

van Vliet, M. T., Franssen, W. H., Yearsley, J. R., Ludwig, F., Haddeland, I., Lettenmaier, D. P., and Kabat, P.: Global river discharge and water temperature under climate change, Global Environmental Change, 23, 450–464, https://doi.org/10.1016/j.gloenvcha.2012.11.002, 2013.

Van Vuuren, D. P., Edmonds, J., Kainuma, M., Riahi, K., Thomson, A., Hibbard, K., Hurtt, G. C., Kram, T., Krey, V., Lamarque, J.-F., et al.: The representative concentration pathways: an overview, Climatic Change, 109, 5, https://doi.org/10.1007/s10584-011-0148-z, 2011

Vidal, J.-P., Hingray, B., Magand, C., Sauquet, E., and Ducharne, A.: Hierarchy of climate and hydrological uncertainties in transient low-flow projections, Hydrology and Earth System Sciences, 20, 3651–3672, https://doi.org/10.5194/hess-20-3651-2016, 2016.

Viviroli, D., Zappa, M., Gurtz, J., and Weingartner, R.: An introduction to the hydrological modelling system PREVAH and its pre-and post-processing-tools, Environmental Modelling& Software, 24, 1209–1222, https://doi.org/10.1016/j.envsoft.2009.04.001, 2009.

Weingartner, R. and Aschwanden, H.: Discharge regime-the basis for the estimation of average flows, Tech. Rep. Plate 5.2, Hydrological Atlas of Switzerland, Bern, Switzerland, 1992.

Weingartner, R., and Schwanbeck, J.: Veränderung der Niedrigwasserablüsse und der kleinsten saisonalen Abflüsse in der Schweiz im Zeitraum 1961-2018, Comissioned by the Swiss Federal Office fort he Environment, 41pp., 2020.<u>Wilby, R. L.,</u> and Dessai, S.: Robust adaptation to climate change, Weather, 65, 180-185, https://doi.org/10.1002/wea.543, 2010.

Zekollari, H., Huss, M., and Farinotti, D.: Modelling the future evolution of glaciers in the European Alps under the EUROCORDEX RCM ensemble, The Cryosphere, 13, 1125–1146, https://doi.org/10.5194/tc-13-1125-2019, 2019.

Zierl, B. and Bugmann, H.: Global change impacts on hydrological processes in Alpine catchments, Water Resour. Res., 41, 1–13, https://doi.org/10.1029/2004WR003447, 2005.

Table 1:1: Overview of the available-climate model chains available and their initial grid spacings of 12 km (EUR-11) and 50 km (EUR-44).

Global Climate Model	Regional Climate Model	EUR-11	EUR-44
	KNMI-RACMO22E		Х
	DMI-HIRMAM5	Х	
ICHEC-EC-EARTH	CLMcom-CCLM4-8-17	Х	
	CLMcom-CCLM5-0-6		Х
	SMHI-RCA4	Х	
MOHC-HadGEM2-ES	CLMcom-CCLM4-8-17	Х	
	CLMcom-CCLM5-0-6		Х
	KNMI-RACMO22E		Х
	SMHI-RCA4	Х	
	CLMcom-CCLM4-8-17	Х	
	CLMcom-CCLM5-0-6		Х
MP1-MP1-ESM-LK	SMHI-RCA4	Х	
	MPI-CSC-REMO2009-2	Х	
	CLMcom-CCLM5-0-6		Х
MIROC-MIROCS	SMHI-RCA4		Х
CCCma-CanESM2	SMHI-RCA4		Х
CSIRO-QCCCE-CSIRO-Mk3-6-0	SMHI-RCA4		Х
IPSL-IPSL-CM5A-MR	SMHI-RCA4	Х	
NCC-NorESM1-M	SMHI-RCA4		Х
NOAA-GFDL-GFDL-ESM2M	SMHI-RCA4		Х

Table 2:2: Relative changes (in %) by <u>the</u> end of the century and the <u>seasonal-occurrence (OCCseasonality (SEAS</u>) of moderate low and high flows for the six representative catchments. The <u>seasonal-occurrenceseasonality</u> indicates the season <u>wherein which</u> the moderate <u>flow extremes</u><u>low and high flows generally</u> occur<u>most frequently.</u> Abbreviations: <u>Seasonal occurrenceSeasonality</u> in <u>the</u> winter <u>half-year</u> in the reference period and predominantly in <u>the</u> summer <u>half-year</u> in the future period (WS), <u>seasonal</u> <u>occurrenceSeasonality</u> in <u>the</u> summer <u>half-year</u> in the reference period and in <u>the</u> winter <u>half-year</u> in the future period (SW), and <u>seasonal occurrenceSeasonality</u> in <u>the</u> summer <u>half-year</u> in the in both periods (SS).

Catchment	Moderate low flows				Moderate high flows			
	YEARANN	WINTER	SUMMER	OCC<u>SE</u>	YEARANN	WINTER	SUMMER	OCC<u>SE</u>
	UAL			AS	UAL			AS
Rosegbach	+191%	+199%	+89%	WW	-20%	+97%	-25%	SS
Kander	+20%	+41%	-37%	WW	+8%	+50%	-5%	SS
Plessur	-1%	+32%	-43%	WS	-10%	+43%	-14%	SS
Emme	-53%	-17%	-66%	WS	+2%	+13%	-7%	SW
Venoge	-45%	-16%	-47%	SS	+22%	+25%	+4%	WW
Verzasca	-22%	+42%	-53%	WS	+6%	+24%	-3%	SS





Figure 1:1: Overview of the study region<u>catchments</u> and the location of the <u>corresponding</u> gauging stations (orange dots). Shadings indicate-Grey shadings show the mean altitudeelevation of the respective<u>a</u> catchment. Green<u>Catchments are divided into two</u> groups: lower-lying catchments (green) and Alpine catchments (purple). Blue contours indicate the six example catchments: Rosegbach — Pontresina (1), Kander — Hondrich (2), Plessur — Chur (3), Emme — Emmenmatt (4), Venoge — Ecublens (5), Verzasca — Lavertezzo (6). (from Muelchi et al., 2020e in reivew)





Figure 2: Median monthly occurrence of moderate low flows (left panels) and high flows (right panels) for the reference period (1981–2010, a, bupper panels) and for the end of the century (2070–2099, c, dlower panels).









(e) SUMMER



SUMMER







Figure 3: Relative changes of magnitude by <u>the</u> end of the century for moderate low flows (left panels) and moderate high flows (right panels) for the year (a,b), upper panels), the winter (c,dhalf-year (middle panels), and <u>the summer (c,f)-half-year (lower panels)</u>. Triangles indicate catchments with annual moderate low and high flows occurring in the <u>respective</u>-time window in the

reference period. Circles indicate seasonal <u>lowestlow</u> and <u>highesthigh</u> flows outside the typical low_ and high-_flow season. Black contours indicate changes with at least 90% of the models agreeing on the direction of change.

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(C) WINTER













(f)





Figure 6: Time of emergence of moderate low flows (left panels) and moderate high flows (right panels) when at least 66% of the models agree on significant changes in the distribution of moderate low and high flows, respectively.









Figure 7: <u>Multi-modelMultimodel</u> median of <u>intensitymagnitude</u> and <u>seasonal-occurrenceseasonality</u> of high flows and seasonal <u>highesthigh</u> flows in Alpine catchments: Rosegbach (top row), Kander (middle row), and Plessur (bottom row).









Figure 8: <u>Multi-modelMultimodel</u> median of <u>intensitymagnitude</u> and <u>seasonal-occurrenceseasonality</u> of high flows and seasonal <u>highesthigh</u> flows in <u>low-lower-lying</u> catchments: Emme (top row), Venoge (middle row), and Verzasca (bottom row).



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Figure 9: Relative changes in the frequency of occurrences of moderate low-_flow events (lower than the median of the reference period, top panels), of moderate high_flow events (higher than the median of the reference period, middle panels), and co-occurrence of low_ and high-_flow events (bottom panels) by <u>the</u>end of the century. Black circles indicate changes with at least 90% of the models agreeing on the direction of change.