



Seasonal watershed-scale influences on nitrogen concentrations across the Upper Mississippi River Basin

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14 Abstract. Humanity's footprint on Earth systems has engendered water quality impoverishment in streams, lakes, 15 and coastal waters globally. In agricultural areas, stream nitrogen concentrations are often high where excess 16 nitrogen fertilization and wetland loss via artificial drainage degrade water quality. While the watershed-scale 17 influence of fertilization and wetland loss on annual nitrogen loads has been studied, little is known about the 18 watershed-scale effects of these wetland losses at seasonal time scales. Here we apply machine learning and linear 19 statistical analyses in a big data framework to improve understanding of the role wetlands play in influencing the 20 seasonality of down-gradient water quality. We confirm the seasonal role of wetlands in improving water quality at 21 the watershed scale and uncover evidence demonstrating the importance of contemporary watershed nitrogen inputs 22 to in-stream total nitrogen concentrations [TN]. We observe that in the Upper Mississippi River Basin, United 23 States, after the application of spring fertilizers, [TN] drops by 70% from June to September suggesting the 24 importance of seasonal nutrient loading. Our data mining approach affords exploration of the potential influence of 25 numerous landscape and wetland hydrologic processes on [TN], some of which are shown to exert seasonal 26 influence. Our counterfactual analysis-in which wetlands are restored to their historic extent-points to the 27 substantial water quality benefits of wetland restoration, including particular water quality improvements in the 28 spring when [TN] are highest. Water quality benefits due to wetland restoration would make water safer for human 29 consumption and improve the security of aquatic ecosystems. 30

31 1 Introduction

32 Numerous rivers (Meybeck and Helmer, 1989), lentic inland waters (Brooks et al., 2016), and receiving coastal 33 waters (Diaz and Rosenberg, 2008) are in the throes of a water quality crisis largely driven by increasing rates of 34 anthropogenic nutrient loading since the 1850s (Vitousek et al., 1997) to sustain Earth's human population of 7.8 35 billion. With projections of continued population growth (Gerland et al., 2014) requiring expansion or further 36 intensification of agriculture, in the absence of dramatic measures, already realized environmental degradation will 37 persist or increase in extent or severity (Liu et al., 2012). Water quality degradation due to excess nutrient loads can 38 lead to deterioration of ecosystems, "impoverishment of aquatic biodiversity" (Bogardi et al., 2012), harmful algal 39 blooms (Brooks et al., 2016), and hazards to human health (Falkenmark, 1990;Bouwer, 2000). In coastal waters 40 impacted by eutrophication, hypoxia-driven habitat compression may hinder life cycle functions of pelagic species 41 (Diaz and Rosenberg, 2008). Integrating the economic consequences of those excess nitrogen (N) impacts yields a 42 staggering economic burden—on the order of hundreds of billions USD per year (Houlton et al., 2019;Compton et 43 al., 2011;Sutton et al., 2011;Sobota et al., 2015), a cost that might be partially offset through wetland restoration 44 (Rankinen et al., 2014;Hey, 2002;Houlton et al., 2019).

45 At the global scale, 64% of inorganic nitrogen export originates from anthropogenic sources, and 54% 46 occurs as a consequence of diffuse agricultural inputs (Seitzinger et al., 2005), though urban contributions are 47 important as well (Chen et al., 2016). At regional scales as much as 78% of N loading can be a consequence of 48 agricultural fertilization (Compton et al., 2019). Riverine export rates have been observed ranging from 5% (David 49 et al., 2010) to 38% (Compton et al., 2019) of N inputs. While some have attributed this observed nitrogen loading 50 in part to legacy effects (Van Meter et al., 2018; Basu et al., 2010), others have attempted to distinguish between 51 legacy versus contemporary loading hypotheses and have found stronger evidence for the role of ongoing large-scale 52 fertilization (Ballard et al., 2019;Stackpoole et al., 2019). However, in the absence of large-scale critical 53 experimentation (Platt, 1964) the relative importance of possible legacy effects across different systems remains 54 uncertain.





55 Nitrogen, along with phosphorus, is a key limiting nutrient that, in excess, contributes to observed 56 widespread eutrophication in coastal (Ryther and Dunstan, 1971), riverine (Dodds and Smith, 2016), and lacustrine 57 (Conley et al., 2009) environments. Nitrogen-laden leachate and surface runoff that flows downgradient toward 58 rivers, lakes, and oceans may first intersect connecting floodplain (Noe and Hupp, 2005;Sanchez-Perez et al., 2003) 59 and/or non-floodplain wetlands (Lane et al., 2018; Mushet et al., 2015). In the anaerobic sediments underlying these 60 wetlands, biologically available nitrogen in the form of nitrate (NO3⁻) may-in the presence of labile carbon and 61 anoxic conditions for microbial activity (Soares, 2000)-be converted to nitrogen gas (N2) by the denitrification 62 process (Sanchez-Perez et al., 2003). Evidence has therefore emerged that wetlands may substantially reduce 63 nitrogen loading of surface waters at the watershed scale (Hansen et al., 2018; Quin et al., 2015; Fisher and Acreman, 64 2004;Golden et al., 2019). The efficacy of wetlands in reducing nitrogen loading has been observed across a range 65 66 of flow conditions and seasons (Uuemaa et al., 2018).

Investigations of fertilizer inputs and wetland impacts on nitrogen loading are commonly conducted at an annual timescale (Basu et al., 2010;Thompson et al., 2011;Van Meter et al., 2016;Van Meter et al., 2018, 2019;Golden et al., 2019). Yet temporal variability is critically important because it influences the timing and frequency of events in which nutrient concentrations exceed levels safe for human consumption (e.g., 10 mg l⁻¹ in the US; 4.4 mg l⁻¹ in Germany) and ecological integrity. Despite recent advances in understanding the role of wetlands in influencing water quality across spatial scales, improved understanding of the seasonal variability remains a priority (Bloschl et al., 2019).
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Golden et al. (2019) suggest the possible role of big data in improving understanding of the influence of
 non-floodplain wetlands on water quality. To deal with the challenges associated with big data, the hydrologic
 science community has tentatively and successfully tested machine learning (Shen, 2018;Tyralis et al., 2019), a data
 driven approach that contrasts with and supplements the normative physically based methods. Application of
 machine learning to seasonal water quality data and process-based drivers of nutrient loads, presents an opportunity
 for discerning the seasonal dynamics of wetlands as they relate to watershed-scale nitrogen conditions.

79 In this paper, we explore the seasonal role nitrogen inputs and wetlands play in influencing downgradient 80 water quality at the watershed scale. To advance this goal we ask: 1) What is the role of watershed characteristics in 81 mediating intra-annual stream TN concentrations ([TN]) at the catchment to large watershed scales? 2) To what 82 extent is restoration of historic wetland distributions a viable means to improve intra-annual stream [TN] and how 83 might water quality improvements be seasonally dependent? To answer these questions, we capitalize on recently 84 developed novel databases of wetland metrics that reflect the structural and functional characteristics of hydrological 85 flowpaths into and out of wetlands (Mengistu et al, In Revision; Leibowitz et al., In Review), in addition to 86 distributed measurements of streamflow and [TN]. We combine machine learning (random forest) with a linear 87 mixed effects model to assess seasonal variability of [TN] in watersheds across the Upper Mississippi River Basin 88 (UMRB) in the Midwestern US. The results of our combined machine learning and linear statistical approach points 89 to the key drivers of variations in [TN] across seasons in the UMRB and provide insights on how machine learning 90 may be used in future watershed-scale water quality analyses. 91

92 2 Methods

93 2.1 Study area

94 The Mississippi is the longest river in the United States, originating at Lake Itasca, Minnesota. The study area 95 consists of the Upper Mississippi River basin (UMRB; 492,000 km²; Figure 1), the largest contributor of residual 96 nitrogen to the Mississippi River basin (Burkart and James, 1999;Qi et al., 2020). The UMRB consists principally of 97 the Great Plains, northern forests, and eastern temperate forests ecoregions (Omernik and Griffith, 2014). The 98 surficial geology of the Upper Mississippi River basin is dominated by thick silty glacial till sediments interspersed 99 with thinner units (Soller and Reheis, 2004). Additionally, thick coarse-grained proglacial sediments are present, 100 typically toward the basin's northern extent. Precipitation averages 920 mm yr⁻¹, two-thirds of which falls in spring 101 and summer. Precipitation varies spatially from a low of 600 mm yr⁻¹ in the northwest to a high of 1200 mm yr⁻¹ in 102 the southeast (Daly et al., 2008). Potential evaporation increases from a low in January to a peak in July. Elevation, 103 which serves as a hydraulic driver in topographically driven flow regimes, ranges from 520 m in the northeast to 140 104 m in the southeast.

105 Overlaid on climate dynamics and natural physiography, human alterations in the form of conversion of
 106 perennial vegetation to seasonal crops (Zhang and Schilling, 2006) and artificial drainage have pervasively
 107 amplified the hydrologic cycle by increasing stream discharge (Blann et al., 2009;Belmont et al., 2011;Schottler et
 108 al., 2014). The Upper Mississippi River is also influenced by the presence of numerous locks (Gramann et al.,
 109 1984). Integrating the natural environmental conditions with human alterations yields a hydrologic regime in which





110 specific discharge (the quotient of streamflow volume and contributing area) increases throughout the winter to a 111 peak in the spring (May), decreases through the summer, and remains low in the fall.

112 The UMRB drains some of the continent's most fertile arable land, which is predominantly in either corn or 113 soybean production. Because of these and other land uses a variety of nutrients and contaminants have been 114 observed in the waters of the Mississippi River. The majority of fertilizer is applied in the spring prior to or after 115 planting, though dry fall fertilizer applications range from 0-25% of the total annual amount (Cao et al., 2018). 116 Nutrient loading has been exacerbated by land management practices, such as artificial drainage, that enhance 117 discharge in the Upper Mississippi River (Schilling et al., 2010). River sediment cores have revealed an order of 118 magnitude increase of sediment deposition, since the 1830s (Engstrom et al., 2009). In the UMRB mean N fertilizer use rate in 2015 was 49 (maximum = 173) kg N ha⁻¹ yr⁻¹, corresponding to $2.4 \cdot 10^6$ t N yr⁻¹ (Cao et al., 2018). The 119 120 use rate within the UMRB exceeds NO₃-N loads discharged from the Mississippi River into the Gulf of Mexico as 121 reported by Van Meter et al. (2018). Within the Upper Mississippi River, water quality degradation is known to 122 impact aquatic communities (Houser and Richardson, 2010). Excess nutrients-including a tripling of NO3 N input 123 to the Gulf of Mexico (Goolsby et al., 2001;McIsaac et al., 2002)—are credited with causing the dead (hypoxic) 124 zone in the Gulf of Mexico, which has been measured at 20,700 km² in extent (Rabotyagov et al., 2010) and has 125 raised questions about the current approaches to improving water quality in the Mississippi River (McLellan et al., 126 2015).

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128 2.2 Water quality data

129 We obtained quality controlled stream [TN] (unfiltered total nitrogen as N) data (1995-2007) from the SPARROW 130 (Spatially Referenced Regression on Watershed Attributes) Major River Basin 3 water quality modeling group, 131 which compiled data from federal, state, and local government monitoring (Saad et al., 2011). Sites included at least 132 25 stream [TN] measurements per site (over the course of the 13-year sampling window) distributed throughout the 133 year to ensure representation of all seasons. This resulted in a total of 6895 [TN] measurements divided amongst 82 134 sites (Table S1), all of which included corresponding discharge measurements. [TN] values in this dataset range 135 from 0.1 to 25.1 mg l⁻¹ (median=3.6). These measurements are taken from streams draining watersheds of median 136 area 2,580 km² (mean=7,229), ranging from small catchments (45 km²) to large watersheds (52,048 km²).

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138 2.3 Derivation of variables

To describe the variation in our response variable [TN], we used static (i.e., time-invariant) watershed-scale
predictor variables developed in recent work (Mengistu et al., in revision) that aimed at quantifying watershed and
wetland characteristics as well as the structural and functional attributes of flowpaths between wetlands and rivers.
The variables (Table S2) reflect potentially important elements of watershed-scale nutrient cycling, including source
(e.g., nutrient loading via cultivated areas), sink (e.g., denitrification via wetland and open water areas), and
transport processes (e.g., soil types of overland flowpaths).

145 Specifically, for the 82 watersheds, the variables describe watershed characteristics (e.g., land cover, 146 watershed area) and average watershed-scale wetland-to-stream flowpath characteristics (e.g., wetland-to-stream 147 flowpath Mannings values or maximum soil porosity along that flowpath) that are intended as proxies for hydraulic 148 and hydrologic processes. The variables represent source contributions to wetlands and the structural characteristics 149 of hydrological flowpaths from wetlands to the nearest streams. For example, maximum porosity along the flowpath 150 (a structural characteristic) is a proxy for infiltration and conversion of overland flow to shallow and/or deep 151 subsurface flows, which attenuate TN from reaching the stream (functional characteristics). The derivation and 152 detailed lists of these variables are described in detail in Mengistu et al. (in revision).

153 The static variables are supplemented by time varying values developed in this study-monthly soil NO₃-N 154 (Wu and Liu, 2012), daily discharge (from USGS gages; see Table S1), monthly wetness index (the quotient of 155 spatially averaged precipitation and potential evaporation) derived from PRISM climate data (Daly et al., 2008), 156 year of [TN] sampling, and day of year of [TN] sampling. Potential evaporation (PET) was estimated following 157 Hargreaves (1994), where daily minimum and maximum temperatures were extracted from PRISM for each 158 watershed. Soil NO₃-N consists of one value for each month, estimated from a Soil and Water Assessment Tool 159 (SWAT) simulation for the Iowa River basin, which is located within UMRB (Wu and Liu, 2012). In the absence of 160 watershed-specific information on temporal evolution of soil NO₃-N (kg ha⁻¹), the secular values from Wu and Liu 161 (2012) were allowed to remain uniform across watersheds. Areal extent calculations were performed in the Albers 162 Equal Area projection for the conterminous US. All geospatial analyses were conducted in ArcGIS 10.7 and R 163 (Mengistu et al. (in revision)).

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165 2.4 Modeling approach





166 We used a three-phased model approach to investigate the potential watershed-scale wetland and landscape drivers 167 of seasonal [TN] variability: (1) a nonlinear machine learning approach followed by (2) a linear statistical model 168 with seasonal harmonics (i.e., periodic functions) and (3) counterfactual simulations using the statistical model to 169 ascertain the potential influence of wetland restoration on [TN]. The three-phased approach helped us ascertain 170 which hydrologic processes potentially serve as key drivers of [TN] variability across the UMRB, because the 171 relative importance of the hydrologic processes influencing the seasonal evolution of water quality-including their 172 interactive influence and scale dependence on drivers of [TN]-are unknown a priori. Hence, we first interrogated 173 watershed metrics (see Table S2) as predictor variables for [TN] using random forest (see Machine learning). Based 174 on the scores of explanatory variables in random forest, we then selected the top candidate variables to develop

linear mixed effect (LME) models that capture and reproduce the cyclical seasonal nature of [TN] (see *Mixed effects modeling*). We subsequently linked these [TN] LME models to the seasonal influence of wetlands using an additive

variable approach (see *Mixed effects modeling*). Using the linked [TN] LME model, we performed the

- 178 counterfactual simulations (see *Wetland restoration scenarios*).
- 179

180 2.5 Machine learning

181 We chose the nonlinear random forest machine learning algorithm to analyze watershed proxies explaining [TN] 182 variability. Random forest quantifies individual variable importance and minimizes overfitting and bias, while 183 remaining "competitive" with other methods (Breiman, 2001) and providing the flexibility of fewer assumptions 184 than traditional linear statistical methods. Despite limitations associated with interpretability (Shen, 2018), random 185 forest has numerous perceived advantages for water resource applications-including accounting for interactions 186 among variables (Cutler et al., 2007), appropriateness for big data applications, and computational efficiency 187 (Tyralis et al., 2019). Random forest is increasingly applied to extract information relevant to water resources within 188 a big data context (Cho et al., 2019), including its use in determining the role of competing drivers in mediating 189 observed water quality across spatial scales (Read et al., 2015).

190 We used the randomForest package (v. 4.6-14) in R (v. 3.6.1) and tuned the m_{try} parameter to minimize out 191 of bag error, a procedure described in detail by Tyralis et al. (2019) and citations therein. Specifically, out of bag 192 refers to those samples that were withheld from model training for verification purposes. We principally considered 193 candidate variables for the LME models amongst the top five highest ranking in the random forest model results. 194 However, noting the large number of variables considered (n=53; Table S2) and substantial cross-correlation among 195 variables, we supplemented the random forest approach with expert knowledge by removing variables deemed 196 important (by random forest) but not known in the scientific literature as important sources or sinks of nitrogen. 197 (Removed variables are reported in the Results section, below.) In the interest of incisively determining the limits of 198 random forest for the application at hand we ran the algorithm on the dataset as a whole, a random subset of 70% of 199 observations at each measurement location, and on all measurements at a random subset using 70% of sampling 200 locations.

201 Our final selected variables based on random forest and expert judgement were input into the LME model.
 202 Specifically, to build an LME, we relied on a subset of the most important predictors that emerged from random forest, our system understanding, and common metrics of model performance—the Akaike and Bayesian
 204 Information Criteria (AIC and BIC, respectively; Helsel et al., 2020).

206 2.6 Mixed effects modeling

207 Mixed effects modeling is an extension of simple linear modeling and is widely used in investigating complex water 208 resource problems (Bart, 2016;Wine et al., 2018a;Wine et al., 2018b;Bywater-Reves et al., 2018;Hurley and 209 Mazumder, 2013: Araujo et al., 2012: Ahearn et al., 2005). This is particularly true for those situations in which part 210 of the natural variability is associated with measured phenomena (i.e., fixed effects) and part of the natural 211 variability results from complex phenomena (i.e., random effects) such as site-specific characteristics. LME also 212 offers tools to overcome heteroscedasticity. In this way, LME relaxes certain assumptions commonly associated 213 with application of simpler methods (Zuur, 2009). Assumptions of LME models include homogeneity of variance 214 215 and correct model specification-that all relevant terms and interactions are included.

In developing an LME model, we sequentially added variables starting from those assigned highest importance by random forest, ensuring that variables representing key concepts from the advection diffusion reaction equation (ADRE)—including nitrogen sources, fluid advection, reactions (i.e., denitrification), and scale are represented. While we did not intend to solve the transient ADRE (Clairambault, 2013;Oldham et al., 2013) here, we nonetheless considered it briefly as a lens into the physical processes underlying the temporal dynamics of nonconservative solute concentrations (*c*):





222 $\frac{\partial c}{\partial t} = -\nabla \cdot (cu) + \nabla \cdot (D\nabla c) - \frac{\partial q}{\partial t} + S,$ Eq. 1 223

224 which are controlled by velocity (u), the diffusion coefficient (D), concentration change due to reactions (a), and 225 sources and sinks (S). These dynamics are transient in time (t) and distributed across space. This transient equation allows for seasonal variations in [TN] as a consequence of advective $(\nabla \cdot (cu))$, diffusive $(\nabla \cdot (D\nabla c))$, reactive $(\frac{\partial q}{\partial t})$, 226 227 or additive (S) processes. The quotient of the respective rates of advective transport and diffusion, the Peclet 228 number, is directly related to velocity and length scale and inversely related to diffusivity. Empirically, increasingly 229 chemostatic behavior has indeed been observed as catchment scale increases (Creed et al., 2015), consistent with 230 ADRE. With respect to the source term, the timing of fertilization is expected to yield a key seasonal source whereas 231 crop growth and denitrification serve sink functions. In UMRB it is reasonable to expect that [TN] is influenced by 232 coupled surface and subsurface flow and transport dynamics. Though a physically based approach to understanding 233 seasonal nitrogen dynamics is beyond the scope of this work, this approach nonetheless remains as a reference.

To build our LME models, we considered first order interactions for those variables perceived as having an interactive influence on [TN], i.e., in cases where a predictor's influence on [TN] was expected to depend on the value of another predictor. Model development was ceased—following successive AIC improvements—when the aforementioned key concepts and interactions had been represented.

238 We anticipated seasonally cyclic behavior in [TN] across the UMRB, particularly because seasonal nutrient 239 loading from agricultural fertilization, as well as from other diffuse nutrient sources and point discharges, is 240 prominent across the UMRB. Helsel et al. (2020) suggest representing this cyclic behavior with periodic functions 241 or harmonics (i.e., sine and cosine) and identifying and addressing cases where cycles shorter than one year may 242 occur. We used this approach for our intra-annual [TN] LME models. However, we did not anticipate that all 82 243 watersheds would exhibit identical periodic behavior (i.e., amplitude and timing of peak). Hence, we included these 244 periodic functions as random effects, which allows the parameters defining each harmonic to vary by watershed. To 245 ensure that modeling assumptions were met we examined the model residuals of each watershed for 246 heteroscedasticity at the completion of the modeling analyses.

Once the LME models were developed and to further explore factors that may be driving the seasonality in
 [TN], we calculated the amplitudes of the first harmonic in the final selected LME models and applied Spearman
 rank correlations between these amplitudes and the watershed variables derived from random forest. This
 nonparametric approach minimizes distributional assumptions and sensitivity to extreme values and provides further
 insights into how [TN] varies seasonally with our random forest-based watershed variables.

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255 2.7 Wetland restoration scenarios

256 To evaluate the extent to which wetland restoration might serve as an effective means of enhancing water quality by 257 decreasing [TN], we engaged in counterfactual modeling. Our counterfactual modeling used the final selected LME 258 model and we altered two variables to create our various wetland restoration scenarios: the proportion of the 259 watershed covered by wetland and the proportion of the watershed covered by cultivated land-cover types. (Note 260 that since wetland area is not strictly independent of cultivated area-due to issues of complementarity-the results 261 must be interpreted in the context of this limitation.) For our counterfactuals, we inferred historic wetland areas from 262 Horvath et al. (2017), which provides fine resolution (30 m) estimates of potential wetland restoration on 263 agricultural lands using soils and topography. In contrast to the widely referenced GLWD, derived from small-scale 264 maps at 1:1,000,000 to 1:3,000,000 spatial resolution, Horvath et al. (2017) relies in part on SSURGO (USDA 265 NRCS Soil Survey Geographic database), which is based principally on 1:24,000 to 1:12,000 spatial resolution. We 266 developed two counterfactual scenarios: (1) 50% wetland restoration to historic conditions and (2) 100% wetland 267 restoration to historic conditions. We assumed that increasing wetland area proportionally decreased cultivated area. 268 Further, potential [TN] reductions were subject to the natural limit that concentration reductions cannot exceed 269 observed concentrations. For each scenario, we predicted [TN] across the 82 UMRB watersheds.

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271 3. Results

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273 3.1 Seasonal variability of [TN] across UMRB

[TN] varies strongly in the UMRB—by a factor of 250— among all sites and across the 13-year study period.

275 Concentrations range from 0.1 mg l⁻¹ to 25.1 mg l⁻¹. In 12% of the measurements, [TN] exceeds 10 mg l⁻¹, the

276 maximum contaminant level goal for nitrate as N in the US, and in 44% of the measurements [TN] exceeds 4.4 mg l





277 ¹, the drinking water standard in Germany (as an example of more stringent global water quality standards).
278 Considering all watersheds, [TN] is lowest in September, increases during the fall when fertilizer is sometimes
279 applied, remains steady through the winter, and increases further in the spring when fertilizer is commonly applied
280 (Figure 2). Between June and September, median [TN] drops by 70% (Figure 2). Though this pattern is generally

exhibited by the dataset as a whole, the relatively small watersheds (<350 km²) have a wide range of [TN] seasonal
 patterns (Figure 3).

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284 3.2 Random forest

Random forest predicted, with low bias, all training datasets as well as the verification dataset that consisted of a subset of observations from each site (Figure 4). Prior to running random forest, the optimization procedure incrementally reduced out of bag error for [TN] predictions. However, when predictions were attempted by a random forest trained on a random subset of sites, substantial positive biases (i.e., over-prediction) were observed on sites that had been excluded from the training dataset. This is a consistent and known challenge in predictive modeling of complex systems. Ultimately, random forest assigned candidate predictor variables a range of importance values, calculated as the percent increase in mean square error when a variable is withheld.

Discharge was identified as the single best correlate of [TN] (Table 1), though this correlation cannot be
 interpreted independently of other terms in ADRE. Other important variables that emerged included *year of sampling*, a proxy for interannual climate or land-management variability; *day of year of sampling*, a proxy for intra annual variability in nitrogen loading and the hydrologic cycle; and *monthly wetness index*, a proxy for drivers of
 hydrologic dynamics. While wetland metrics tended to achieve low importance ranks, this does not necessarily
 indicate an inability on their part to influence water quality. Rather, this may instead point to the limited remaining
 distribution, or historic loss, of wetlands across large areas of the UMRB (Figure 5).

299 Forest emerged as the most important watershed metric in predicting [TN]. (As anticipated, the nature of 300 the correlation was inverse.) Forests are not expected to serve as a major source of nitrogen, and forests are typically 301 nitrogen sinks only in locations where atmospheric deposition is the most important nitrogen source (e.g., in forest 302 of the Northeastern United States (Goodale et al., 2002)). Therefore, based on expert opinion that forests in the 303 UMRB do not match either criteria, we removed the forest predictor. Its ranking in the random forest model was 304 subsequently replaced by total watershed N inputs, i.e., total annual agricultural inputs of TN plus annual 305 atmospheric deposition of TN (Mengistu et al, in revision), along with average watershed Manning's roughness 306 coefficients for wetland-to-stream flowpaths. Manning's coefficients were estimated based on land cover values, 307 with the highest values occurring in forested areas (Table 2).

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309 3.3 Linear mixed effects: sequential results

Our linear mixed effects modeling aimed initially to reproduce the cyclical variability in [TN] and then to link [TN]
 to the seasonal influence of wetlands. The resulting idealized model (Table 3, equation 14) represents cyclical trends
 in [TN]. To determine the final model, we fit 14 sequential, increasingly complex models.

313 In the first four models (Table 3), our goal was to reproduce the seasonally cyclic behavior in [TN] seen 314 across the study area (e.g., see Figure 2 and Figure 3, and also large watersheds, as seen in Figure 6) with the first 315 and second harmonics (Figure 7, Table 3 equations 1-4). We next included discharge as a random effect (Table 3, 316 Eq. 6), noting the importance of discharge in random forest (Tables 1 and 2) and that concentration-discharge 317 relationships may be direct, inverse, or weak (Figure 8). (Fitting discharge as a random effect allows LME to assign 318 positive or negative coefficients of appropriate magnitudes.) Most commonly, higher [TN] was observed at higher 319 stream stage (particularly in midsized, untiled watersheds), though the strength of this relationship was variable. It is 320 also important to note that spring fertilization occurs coincident with spring rains, which thereby reflects a high 321 source availability during a high flow period-a combination anticipated to enhance solute transport. Attempts at 322 inclusion of additional random effects increased time to model convergence and the possibility of instability.

Several fixed effect terms further improved the model, including year, monthly soil nitrate concentration, cultivated area, wetland area, watershed area, and monthly wetness index. In our final step, interactions between monthly wetness index and monthly soil nitrate concentrations, as well as wetness index and cultivated area, were added and improved model fit. The final model form (Table 3, Equation 14) explained the greatest amount of the observed variability in [TN] (Figure 9). We accounted for observed heteroscedasticity in model residuals by quantifying the variance of model residuals by watershed in an initial model run (of Eq. 14 in Table 3) and assigning these variances in a fixed variance structure for the final model.

Spearman rank correlations of the first harmonic, representing seasonal [TN] dynamics, with watershed
 variables revealed small amplitudes in association with low watershed nutrient loading (e.g., the presence of
 grassland (r=-0.39)) and watershed area (r=-0.21), where the convolution of many flowpaths across large watersheds





333 is expected to dampen peaks and troughs in [TN] temporal variability (Table 4). The total number and number of 334 wetlands per watershed area (wetland density) were inversely related with the [TN] first harmonic amplitude (r=-335 0.26). A wetland-to-stream structural characteristic that attenuates flow (Manning's coefficient along the flowpath) 336 was also negatively related to the [TN] harmonic (r=-0.25). Because Manning's values are based on land cover and 337 forested areas are assigned the highest Manning's values, this result might simply reflect smaller amplitudes in 338 predominantly forested watersheds. Impervious areas in wetland drainage areas were positively correlated with the 339 [TN] amplitude (r=0.24). Collectivity, the seasonal results from our LME model and correlations with the first 340 harmonic correspond to our process understanding as derived from first principles, i.e., those embedded in the 341 ADRE. Specifically, we see that the LME model and the associated amplitude of the harmonics captures seasonality 342 in the sources and sinks represented in dynamic water quality models.

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345 3.4 Counterfactual models

346 The results of our counterfactual model suggest that the potential water quality gains (i.e., [TN] reductions) 347 associated with 50% and 100% wetland restoration are substantial, even in large watersheds of the UMRB (Figure 348 10). Reductions reach 25 mg l⁻¹ with Horvath et al. (2017) wetland estimates. For example, median measured [TN] 349 was 3.6 mg l^{-1} , which decreases to 1.4 mg l^{-1} following simulated (100%) wetland restoration. The decrease in the 350 third quartile of [TN] is more dramatic, dropping from 7 mg l⁻¹ to 2.3 mg l⁻¹. If complementarity of cultivated areas 351 with wetland loss is neglected, and full wetland restoration alone is considered, the third quartile decreases only to 352 5.4 mg l⁻¹, implying that denitrification by wetlands is secondary to the reduction in fertilization, which is of primary 353 importance. The [TN] reductions are largest during the spring when water quality degradation is most pronounced 354 (Figure 10).

It is clear that [TN] reductions are most prevalent when the seasonal [TN] model developed herein is forced by historic wetland distribution from Horvath et al. (2017), which accounts for wetland areas not detected by GLWD and therefore captures a more complete spatial coverage of potential wetlands across the UMRB. Further, while we assume wetland and cultivated areas are complementary, this is typically reasonable given that most wetland loss in the UMRB has been a consequence of drainage to facilitate expansion of cultivated lands.

361 4 Discussion

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363 4.1 Seasonality of TN

In the presence of recurrent anthropogenic nitrogen loading from fertilizer across the UMRB, which coincides in time with peak discharges (Figure 2), flowpaths were likely activated proximal to the nitrogen source at the land surface, resulting in relatively high [TN] during high discharge periods (Domagalski et al., 2008). Lower flows tended to occur from July through February in the UMRB, out of phase with spring fertilizer application.

Artificial tile drainage decreases the residence time of water in the vadose zone (Danesh-Yazdi et al., 2016) in many of the watersheds throughout the UMRB, thereby facilitating rapid transport of recurrently applied N and engendering high [TN] under baseflow conditions. This is clear in inverse concentration-discharge relationships, particularly in smaller watersheds of the UMRB (see example in Figure 8b). As flows increase in watersheds that are heavily artificially drained, concentrations rapidly decrease, suggesting [TN] is source-limited across a wide range of flow conditions in artificial drained watersheds.

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375 4.2 Wetland restoration and decreased [TN]

376 While we are unable to separate the relative effects of reduced loading from wetland restoration, our counterfactual 377 modeling of potential effects of wetland restoration (Figure 10) further supports this axiom: the presence of wetlands 378 occurs in association with enhanced water quality (Mitsch et al., 2001;Carpenter et al., 1998;Creed et al., 379 2017;Cohen et al., 2016;Vymazal, 2007;Jordan et al., 2011;Lane et al., 2018;Lane et al., 2015;Golden et al., 380 2019;Marton et al., 2015). Our work comports with McLellan et al. (2015) that wetlands, together with other 381 measures, can dramatically improve [TN] in the Mississippi River basin. Whereas much past nutrient removal work 382 has focused on the annual time step (Basu et al., 2010; Cheng and Basu, 2017) or snapshots in time (Hansen et al., 383 2018), here we show that those water quality improvements associated with the absence of loading and nitrogen 384 removal by wetlands correspond to large reductions in stream [TN] during the spring when concentrations are 385 highest (Figure 10).

Our results regarding the potential of wetland restoration to improve water quality contrast with past
 assertions that Gulf of Mexico water quality goals may not be achieved because of legacy nitrogen (Van Meter et al.,
 2018). Our observations of a 70% drop in [TN] between June and September and [TN] increases that tend to occur





coincident with fertilization—during the spring and fall—instead appear to be indicative of the importance of
 contemporary basin-scale nitrogen loading (Van Meter et al., 2020). Based on empirical observations (Figure 2), if
 excess fertilizer applications ceased, we speculate that a large portion of the observed water quality impairment
 would resolve at the time scale of months.

Our findings agree with retrospective analyses observing rapid recovery of groundwater nitrate
 concentrations from initial values exceeding 10 mg 1⁻¹ to values below this threshold at the timescale of months with
 concentrations from initial values exceeding 10 mg 1⁻¹ to values below this threshold at the timescale of months with
 concentrations from initial values exceeding 10 mg 1⁻¹ to values below this threshold at the timescale of months with
 group and to grassland conversion—and if environmental protection was assigned a high priority (Van Meter and
 Basu, 2015). Specifically, our data analyses suggest measures to reduce [TN] in streams in the UMRB, such as
 wetland restoration and associated decreases in fertilization, may not be substantially confounded by legacy effects.
 For these reasons, we suggest that additional critical experiments (Platt, 1964) interrogating the legacy effects of
 nitrogen are needed.

400 401

402 4.3 Big data, machine learning, and uncertainty

403 When we engage in statistical modeling or data mining, we do so with two implicit assumptions-that variables are 404 independent of one another and the amount of measurement error is similar across variables. Our analysis questions 405 these widespread assumptions. Instead we observe the risk in data mining of incorrectly selecting a feature (e.g., 406 forest) whose proportional cover in the watershed is correlated-for reasons of complementarity-with the true 407 cause of nitrogen loading (i.e., surplus nutrient inputs in cultivated areas). Why is a correlated variable assigned—by 408 random forest-higher importance than the causal driver? Perhaps, forested areas are simply more readily quantified 409 (i.e., by satellite remote sensing) with lower error (Wickham et al., 2017) than nitrogen loading-whose estimation 410 is subject to greater uncertainties. To be sure (inverse) dependence of nitrogen export on forest cover is well known 411 (Wickham and Wade, 2002). However, the complexity inherent in coupled human-natural systems means that the 412 question of how to best infer cause-effect relations will persist as a formidable challenge (Ferraro et al., 2019;Muller 413 and Levy, 2019).

414 Following our foray into big data, we agree with McCabe et al. (2017) that big data presents "our 415 community with a unique opportunity to develop new insights that advance fundamental aspects of the hydrological 416 sciences", though big data and machine learning present a double-edged sword (Karpatne et al., 2019). On the one 417 hand an unprecedented quantity of information is available, with the clear potential to improve hydrologic process 418 understanding. On the other hand, this treasure trove of information raises the risk that any number of spurious 419 variables will be assigned high importance by non-physical machine learning algorithms. For this reason, shallow 420 learning algorithms such as random forest are best applied to hypothesis generation and gap filling (of incomplete 421 time series).

422 Among the most informative aspects of this work was the importance assigned to each of scores of 423 candidate predictor variables. In contrast to the typical approach of testing a single hypothesis, we show here that (if 424 we think of each predictor variable as an alternative hypothesis) scores of candidate hypotheses find some level of 425 support. As an example of the implications of the challenge this presents, our framework does not allow for 426 distinguishing between the effects of denitrification versus the simple absence of fertilization. Whereas structural 427 uncertainty is acknowledged as a "major scientific and engineering challenge" (Renard et al., 2010), the level of 428 support assigned to scores of different processes relevant to hydrology suggests the possibility that conventional 429 modeling approaches may underestimate structural uncertainty.

430 When the challenges of model structure are combined with those of the parameterization of those models 431 with complex structures, Beven (1993) concludes that "application of distributed hydrological models is more an exercise in prophecy than prediction". However, we agree with Sivapalan (2009) that by changing the question that 432 433 we pose we can circumvent the uncertainty that arises when uncertain model predictions are used to interrogate the 434 possible effects of secondary, tertiary, or quaternary drivers-often a convenient focus of modern hydrology. For 435 example, in the case of this study, we shifted our research question from a focus on observed wetland influence to 436 wetland restoration scenarios. By asking questions related to the main drivers of unprecedented degradation of 437 natural water resources (e.g., the environmental impacts of growth of humanity's nutrient footprint), we can 438 minimize uncertainty and provide the hydrologic science basis required to remedy the degradation of natural water 439 resources observed to threaten human wellbeing and ecological integrity today.

440

441 4.4 Study Implications

We find ourselves at a point in the Anthropocene in which the measures humans take to secure our well-being
simultaneously threaten our health and the health of the environment. For example, as we observed here, widespread
nitrogen loading and the increase in agricultural land at the expense of wetlands is expected to improve agricultural





productivity, though it simultaneously has impaired water quality. For this reason, humanity's environmental
footprint remains unsustainable (Hoekstra and Wiedmann, 2014;Ehrlich and Holdren, 1971) and diverse approaches
are needed to address ever-expanding nutrient-driven water quality issues.

448 Recently, greater emphasis has been placed on approaching complex water security challenges from an 449 integrative approach that considers perspectives from all relevant disciplines (Zeitoun et al., 2016; Melsen et al., 450 2018). In this case the challenge involves securing high quality freshwater for humans and aquatic ecosystems in the 451 face of powerful agricultural interests together with an unprecedented human population. Indeed, there is growing 452 acknowledgement within the water resource community of the role of water resource securitization-the 453 characterization of an issue as an existential threat requiring implementation of extraordinary measures—in attaining 454 hydrologic process understanding (Wine, 2020, 2019; Wine and Laronne, 2020; Goulden et al., 2009; Brooks and 455 Trottier, 2010;Yu et al., 2015;Schmeier and Shubber, 2018;Farnum, 2018;Grech-Madin et al., 2018). With respect to 456 water quality in the UMRB, there is a need to examine the role played by water resource securitization in hydrologic 457 process understanding, including how it influences the relative importance of such foci as legacy effects, climate 458 change, or the uncertainty associated with non-point source pollution origin or best management practice siting. 459

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462 5 Conclusion

463 Principally as a consequence of ongoing seasonal nitrogen loading in agriculture, [TN] in the UMRB is elevated to 464 the extent that it regularly exceeds the US federal maximum contaminant level for nitrate-nitrogen, with exceedance 465 most likely in June. Expansion of agriculture into former wetland areas together with widespread contemporary 466 nitrogen fertilization (in excess of crop uptake) are primary drivers of the observed degraded conditions whose 467 hazards to human health and biotic security are well known. Here we show that spring rains incident on fertilized 468 agricultural fields increase [TN]. Our model results suggest that restoring historic wetland extent would substantially 469 reduce [TN]-particularly in the spring and early summer when water quality is most severely degraded. This 470 improved water quality, with anticipated benefits for human and biotic health, is caused by the denitrification that 471 takes place in wetland sediments as well as the decrease in fertilization. 472

473 Data availability

474 Data presented in figures will be posted to SciHub following article acceptance. Precipitation and temperature data
 475 were obtained from the PRISM Climate group (<u>https://prism.oregonstate.edu/</u>). Solute concentration data are

476 available from the SPARROW (Spatially Referenced Regression on Watershed Attributes) Major River Basin 3
 477 water quality modeling group.

478

479 Competing interests: The authors declare that they have no known competing financial interests or personal480 relationships that could have appeared to influence the work reported in this paper.

481

482 Author contribution: Conceptualization, MLW, HEG, JRC, CRL, and OM; Data curation, MLW; Formal analysis,
483 MLW; Investigation, MLW; Methodology, MLW, HEG, JRC, CRL, and OM; Software, MLW and OM;
484 Supervision, HEG, JRC, CRL, and OM; Writing—original draft, MLW; Writing—review & editing, MLW, HEG,
485 JRC, CRL, and OM.

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- 838





- 839 Table 1. Importance of top-ranked predictors of TN concentrations as measured by % increase in mean square error
- (MSE) if a particular predictor is omitted. Wetness index (WI) of the month of [TN] measurement (WI_t) and WI of the four preceding months are considered.

Predictor	% Increase in
	MSE
Discharge	112.5
Year	56.1
WI _{t-1}	53.0
Day of Year	50.8
WI _{t-2}	49.7
Forest	45.6
WI _{t-3}	42.8
WIt	40.6
WI _{t-4}	40.2
Soil nitrate (Wu and Liu, 2012)	36.1
Wetland count	22.9
Shallow subsurface flowpath from wetland to stream	22.9
Watershed area	21.5

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- 844 Table 2. Omitting forest area as a predictor reveals the importance of nitrogen loading. See Table 1 caption for
- 845 <u>abbreviations</u>

	% Increase in
Variable	MSE
Discharge	105.8
WI _{t-1}	54.7
Total N input	54.0
Year	52.8
Day of year	48.4
WI _{t-2}	48.2
WIt	42.8
WI _{t-4}	42.7
WI _{t-3}	41.4
Soil Nitrate (Wu and Liu,	
2012)	40.2
Average Manning's	
the flowmath from wetland to	
stroom	33.5
Sucall	55.5
Average Impervious	26.4
Wetland Count	23.4

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Table 3. Total Nitrogen models of increasing complexity, where y_{ijt} refers to \log_{10} [TN] from the *t*th day of year of the *i*th year measured at the *j*th watershed. Predictors include discharge (x_0), seasonal soil nitrogen (x_1) as simulated by Wu and Liu (2012), cultivated area (x_2), wetland area (x_3), watershed area (x_4), wetness index during the preceding month (x_5), and year (α). The mixed effects model involves both fixed effects, which involve fitting scalar coefficients (β) and random effects, which fit vectors (u).

Model	Model Form	New term	AIC	BIC
1	$y_{ijt} = \beta_0 + u_{0j} + \beta_1 \cdot \sin\left(\frac{2\pi t}{365}\right) + \beta_2 \cdot \cos\left(\frac{2\pi t}{365}\right) + e_{ijt}$		11151.78	11185.97
2	$y_{ijt} = \beta_0 + u_{0j} + \beta_1 \cdot \sin\left(\frac{2\pi t}{365}\right) + \beta_2 \cdot \cos\left(\frac{2\pi t}{365}\right) + \beta_3 \cdot \sin\left(\frac{4\pi t}{365}\right) + \beta_4 \cdot \cos\left(\frac{4\pi t}{365}\right) + e_{ijt}$	2 nd harmonic	10886.61	10934.48
3	$y_{ijt} = \beta_0 + u_{0j} + u_{1j} \cdot \sin\left(\frac{2\pi t}{265}\right) + u_{2j} \cdot \cos\left(\frac{2\pi t}{265}\right) + \beta_1 \cdot \sin\left(\frac{4\pi t}{265}\right) + \beta_2 \cdot \cos\left(\frac{4\pi t}{265}\right) + e_{ijt}$	1 st harmonic, random effect	9479.995	9562.049
4	$y_{ijt} = \beta_0 + u_{0j} + u_{1j} \cdot \sin\left(\frac{2\pi t}{365}\right) + u_{2j} \cdot \cos\left(\frac{2\pi t}{365}\right) + u_{3j} \cdot \sin\left(\frac{4\pi t}{365}\right) + u_{4j} \cdot \cos\left(\frac{4\pi t}{365}\right) + u_{4j} \cdot \cos\left(\frac{4\pi t}{365}\right)$	2 nd harmonic, random effect	9053.358	9196.952
5	$y_{ijt} = \beta_0 + u_{0j} + u_{1j} \cdot \sin\left(\frac{2\pi t}{365}\right) + u_{2j} \cdot \cos\left(\frac{2\pi t}{365}\right) + u_{3j} \cdot \sin\left(\frac{4\pi t}{365}\right) + u_{4j} \cdot \cos\left(\frac{4\pi t}{365}\right)$	Discharge, fixed effect	7303.277	7453.706
6	$y_{ijt} = \beta_0 + u_{0j} + u_{1j} \cdot \sin\left(\frac{2\pi t}{365}\right) + u_{2j} \cdot \cos\left(\frac{2\pi t}{365}\right) + u_{3j} \cdot \sin\left(\frac{4\pi t}{365}\right) + u_{4j} \cdot \cos\left(\frac{4\pi t}{365}\right)$	Discharge, random effect	6193.4	6384.855
7	$ + u_{5j} \bullet x_{0ijt} + e_{ijt} $ $ y_{ijt} = \beta_0 + u_{0j} + u_{1j} \bullet \sin\left(\frac{2\pi t}{365}\right) + u_{2j} \bullet \cos\left(\frac{2\pi t}{365}\right) + u_{3j} \bullet \sin\left(\frac{4\pi t}{365}\right) + u_{4j} \bullet \cos\left(\frac{4\pi t}{365}\right) $	Year, fixed effect	6131.496	6404.933
8	$ + u_{5j} \bullet x_{0ijt} + a_{(i)} + e_{ijt} $ $ y_{ijt} = \beta_0 + u_{0j} + u_{1j} \bullet \sin\left(\frac{2\pi t}{365}\right) + u_{2j} \bullet \cos\left(\frac{2\pi t}{365}\right) + u_{3j} \bullet \sin\left(\frac{4\pi t}{365}\right) + u_{4j} \bullet \cos\left(\frac{4\pi t}{365}\right) $	Soil nitrogen, fixed effect	6113.457	6393.724
9	$ + u_{5j} \bullet x_{0ijt} + \alpha_{(i)} + \beta_1 \bullet x_{1t} + e_{ijt} $ $ y_{ijt} = \beta_0 + u_{0j} + u_{1j} \bullet \sin\left(\frac{2\pi t}{365}\right) + u_{2j} \bullet \cos\left(\frac{2\pi t}{365}\right) + u_{3j} \bullet \sin\left(\frac{4\pi t}{365}\right) + u_{4j} \bullet \cos\left(\frac{4\pi t}{365}\right) $	Cultivated area, fixed effect	6059.893	6346.99
10	$ + u_{5j} \bullet x_{0ijt} + \alpha_{(i)} + \beta_1 \bullet x_{1t} + \beta_2 \bullet x_{2j} + e_{ijt} $ $ y_{ijt} = \beta_0 + u_{0j} + u_{1j} \bullet \sin\left(\frac{2\pi t}{365}\right) + u_{2j} \bullet \cos\left(\frac{2\pi t}{365}\right) + u_{3j} \bullet \sin\left(\frac{4\pi t}{365}\right) + u_{4j} \bullet \cos\left(\frac{4\pi t}{365}\right) $	Wetland area, fixed effect	6057.12	6351.047
11	$ + u_{5j} \bullet x_{0ijt} + a_{(i)} + \beta_1 \bullet x_{1t} + \beta_2 \bullet x_{2j} + \beta_3 \bullet x_{3j} + e_{ijt} $ $ y_{ijt} = \beta_0 + u_{0j} + u_{1j} \bullet \sin\left(\frac{2\pi t}{365}\right) + u_{2j} \bullet \cos\left(\frac{2\pi t}{365}\right) + u_{3j} \bullet \sin\left(\frac{4\pi t}{365}\right) + u_{4j} \bullet \cos\left(\frac{4\pi t}{365}\right) $ $ + u_{5i} \bullet x_{0iit} + a_{(i)} + \beta_1 \bullet x_{1t} + \beta_2 \bullet x_{2i} + \beta_3 \bullet x_{3i} + \beta_4 \bullet x_{4i} + e_{iit} $	Watershed area, fixed effect	6044.125	6344.88





12	$y_{ijt} = \beta_0 + u_{0j} + u_{1j} \cdot \sin\left(\frac{2\pi t}{365}\right) + u_{2j} \cdot \cos\left(\frac{2\pi t}{365}\right) + u_{3j} \cdot \sin\left(\frac{4\pi t}{365}\right) + u_{4j} \cdot \cos\left(\frac{4\pi t}{365}\right) + u_{5j} \cdot x_{0ijt} + \alpha_{(i)} + \beta_1 \cdot x_{1_t} + \beta_2 \cdot x_{2_j} + \beta_3 \cdot x_{3_j} + \beta_4 \cdot x_{4_j} + \beta_5 \cdot x_{5ijt}$	Wetness index, fixed effect	6030.475	6338.06
13	$ \begin{aligned} &+ e_{ijt} \\ y_{ijt} &= \beta_0 + u_{0j} + u_{1j} \cdot \sin\left(\frac{2\pi t}{365}\right) + u_{2j} \cdot \cos\left(\frac{2\pi t}{365}\right) + u_{3j} \cdot \sin\left(\frac{4\pi t}{365}\right) + u_{4j} \cdot \cos\left(\frac{4\pi t}{365}\right) \\ &+ u_{5j} \cdot x_{0ijt} + \alpha_{(i)} + \beta_1 \cdot x_{1t} + \beta_2 \cdot x_{2j} + \beta_3 \cdot x_{3j} + \beta_4 \cdot x_{4j} + \beta_5 \cdot x_{5ijt} \end{aligned} $	Interaction of wetness index, soil NO ₃ -N	6022.122	6336.535
14	$\begin{aligned} &+\beta_{6} \bullet x_{5_{ijt}} \bullet x_{1_{t}} + e_{ijt} \\ y_{ijt} &= \beta_{0} + u_{0_{j}} + u_{1_{j}} \bullet \sin\left(\frac{2\pi t}{365}\right) + u_{2_{j}} \bullet \cos\left(\frac{2\pi t}{365}\right) + u_{3_{j}} \bullet \sin\left(\frac{4\pi T}{365}\right) + u_{4_{j}} \bullet \cos\left(\frac{4\pi t}{365}\right) \\ &+ u_{5_{j}} \bullet x_{0_{ijt}} + \alpha_{(i)} + \beta_{1} \bullet x_{1_{t}} + \beta_{2} \bullet x_{2_{j}} + \beta_{3} \bullet x_{3_{j}} + \beta_{4} \bullet x_{4_{j}} + \beta_{5} \bullet x_{5_{ijt}} \end{aligned}$	Interaction of wetness index, cultivated area	6018.495	6339.736
	$+ \beta_6 \bullet x_{5_{ijt}} \bullet x_{1_t} + \beta_7 \bullet x_{5_{ijt}} \bullet x_{2_i} + e_{ijt}$			





Table 4. S	pearman correlation	between the	first harmonic	: [TN] amp	olitude	associated	with each	watershed and	watershed metri	cs.

Variable	Correlation
Watershed area covered by grassland	-0.393
Watershed area covered by barren land	-0.278
Number of wetlands in a watershed	-0.259
Average Manning's roughness coefficient along the flowpath from wetland to stream	-0.246
Watershed area covered by pasture	-0.237
Average proportion of wetland drainage areas with impervious surfaces	0.236
Number of wetlands per unit watershed area	-0.209
Area of a watershed	-0.207
Total wetland area in a watershed	-0.202







Figure 1: Land cover, study basins, and stream gauges in the Upper Mississippi River basin, USA.







Month

Figure 2: Stream total nitrogen (TN) concentrations peak in June following spring fertilizer loading. By September concentrations have decreased by 70%, consistent with contemporary nutrient loading as the primary cause of observed water quality impoverishment. Data are from 82 watersheds in the Upper Mississippi River basin (1995-2007). Recommended timing for fertilizer application (i.e., spring and fall) are shaded gray. (The gray dashed lines respectively refer to the highest and lowest monthly median [TN].)







Month Month Figure 3: A range of conditions are observed in smaller UMRB watersheds (<350 km²), e.g., here TN concentrations peak in September, March, January, and June, respectively. (The red dashed line references the current maximum contaminant level for nitrate.)







Figure 4: (a)The m_{try} parameter in random forest is optimized by minimizing out of bag (OOB) error. (b) Random forest predictions in which all observations are made available for training. (c) Random forest predictions on the training (black) and verification (gray) datasets, where 70% of samples from each site were selected. (d) Random forest predictions when training was performed on all observations from each of 70% of all sites.







Figure 5: Proportion of wetlands within each of the 82 study watersheds in the Upper Mississippi River. Historic distribution of wetlands from the (Top) Global Lakes and Wetlands Database (Lehner and Doll, 2004) and (Middle) Horvath et al. (2017). (Bottom) Contemporary distribution of wetlands from the National Wetland Inventory.







Month Month Figure 6: Four examples of larger watersheds in the UMRB (>27,000 km²) showing Total Nitrogen [TN] concentrations tend to peak in June or July and reach a trough in September or October. (The grey dashed line references 1 mg l⁻¹.)







Day of Year Day of Year Figure 7: Linear Mixed Effect Modeling [TN] with a random intercept and first harmonic (a), adding a second harmonic (b), assigning the first harmonic a random effect (c), and assigning both harmonics as random effects (d). These model iterations correspond to equations 1-4 of Table 3. The red line represents the fixed effects and the blue lines indicate the random effects (by watershed).







Figure 8: Concentration-discharge relationships can be a) direct, b) inverse, or c) weak. Inverse relationships are observed in watersheds in which 50% or more of the area is drained artificially by tiles. d) Typically, the Spearman correlation between concentration and discharge is direct, whereas fewer watersheds exhibit inverse relationships. The red stippled line in a) and b) corresponds to the maximum contaminant level for nitrate in the US.







Figure 9: Mixed effects modeling described TN variability (top), including correspondence between the seasonal distributions of measurements (black) and model predictions (blue). Assessment of the Spearman correlation between measured and modeled values indicates strong (>0.6) to very strong (>0.8) agreement in most watersheds.







Figure 10: Observed (black) and estimated [TN] for 50% (blue) and 100% (green) restoration of the historic wetland extent using our final model forced by Horvath et al. (2017). Red lines, provided as a reference, indicate the maximum contaminant levels for nitrate in the US 10 mg l^{-1} and Germany 4.4 mg l^{-1} .