

Interactive comment on “HESS Opinions: Improving the evaluation of groundwater representation in continental to global scale models” by Tom Gleeson et al.

John Ding

johnding_toronto@yahoo.com

Received and published: 26 August 2020

Groundwater storage as a quadratic reservoir and the value of new streamflow observations

In addition to the assumption of a linear reservoir, $Q = kS$ (Lines 144 & 155; Table 1, Row Baseflow ... recession (k)), the groundwater storage need be considered as a quadratic one.

[Printer-friendly version](#)

[Discussion paper](#)



In the context of model evaluation (Figure 1 and Section 3.1), the importance of acquiring new streamflow data at a project site cannot be overemphasized (Table 1, Row Streamflow).

Considering a catchment on a continental scale as either a linear, $Q = kS = C_1S$, or a quadratic storage, $Q = (C_2S)^2$, the baseflow can be linearized as follows (Beck et al., 2013, cited by authors; Azmi et al., 2020, SC1 therein; Ding, 2020):

$$\log Q(t) = \log Q(t_0) - C_1(t - t_0), \quad (1)$$

Or:

$$-1/\sqrt{Q(t)} = -1/\sqrt{Q(t_0)} - C_2(t - t_0), \quad (2)$$

Equations (1) and (2) are based on the logarithmic and NISR (negative inverse square root) transformation of the streamflow Q , respectively.

On an ungauged catchment, all these two equations will need to determine their discharge coefficients, C_1 and C_2 , are a minimum of three new flow measurements in the field over a period of days or weeks. These will be used to falsify the hypothesis of a linear or a quadratic storage. Direct measurements of the low to mean flow on river on a regional scale are doable, though a logistic and technical challenge (Lines 424-427). Legend or hearsay has it, ancient Egyptians measured the Nile River flow by diverting it to a side chamber where the volume of water was measured.

Figure 1 shows the data transformation diagram for the log and the negative inverse m^{th} root (NI m R) transform (Santos et al., 2018, SC5 therein; Ding, 2018). Compared to the log transform, differences are small among the first, second and third root of the NI m R transform. The NI2R or NISR transform, which is derived for the outflow from a hillslope, maybe considered a representative of these fractional power ones. For

Printer-friendly version

Discussion paper



application, a catchment can be classified either as linear or quadratic, unless dictated theoretically otherwise. Between the two, the log transformation is a low- to mid-pass filter, and the NISR a low-pass one.

References

Azmi, E., Ehret, U., Weijs, S. V., Ruddell, B. L., and Perdigão, R. A. P.: Technical note: "Bit by bit": A practical and general approach for evaluating model computational complexity vs. model performance, *Hydrol. Earth Syst. Sci. Discuss.*, <https://doi.org/10.5194/hess-2020-128>, in review, 2020.

Ding, J. Interactive comment on "Technical note: Pitfalls in using log-transformed flows within the KGE criterion" by Léonard Santos et al., *Hydrol. Earth Syst. Sci. Discuss.*, <https://doi.org/10.5194/hess-2018-298-SC5>, 2018.

Ding, J. Interactive comment on "Technical note: "Bit by bit": A practical and general approach for evaluating model computational complexity vs. model performance" by Elnaz Azmi et al., *Hydrol. Earth Syst. Sci. Discuss.*, <https://doi.org/10.5194/hess-2020-128-SC1>, 2020.

Santos, L., Thirel, G., and Perrin, C.: Technical note: Pitfalls in using log-transformed flows within the KGE criterion, *Hydrol. Earth Syst. Sci.*, 22, 4583-4591, <https://doi.org/10.5194/hess-22-4583-2018>, 2018.

Interactive comment on *Hydrol. Earth Syst. Sci. Discuss.*, <https://doi.org/10.5194/hess-2020-378>, 2020.

Figure 1. Streamflow data transformation diagram for a nonlinear storage-discharge function, $Q = (C_n S)^n$.
 If $n = 1$, $J_1(Q) = \log Q$; if $n > 1$, $J_n(Q) = -1/Q^{1-1/n}$.
 Adapted from Ding, 2018, SC5.

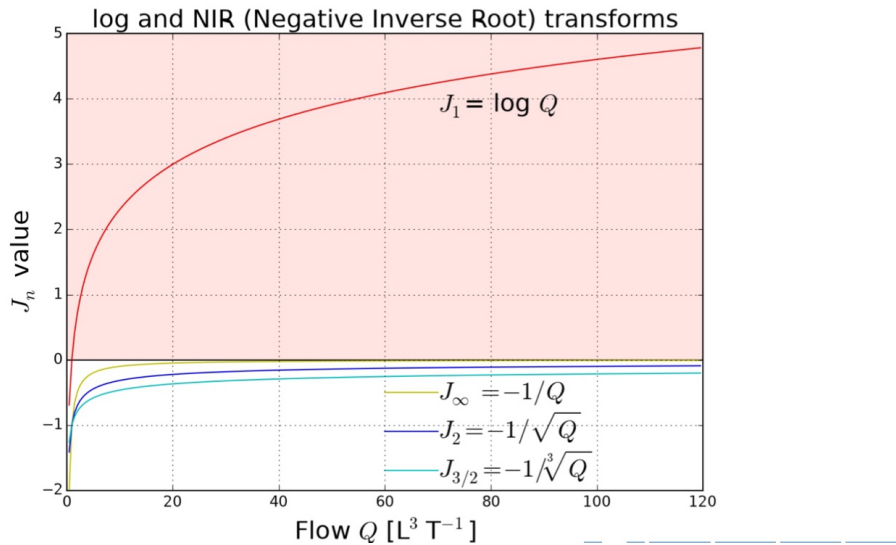


Fig. 1.

Printer-friendly version

Discussion paper

