

Response to Anonymous Referee #1

This is a review of “The evaluation of the potential of global data products for snow hydrological modelling in ungauged high alpine catchments” by Weber et al. The paper investigates the impacts of using a series of climate data products to force a hydrological model and its advanced snow modules and compare the behavior in terms of snow process representation and runoff. The authors also compare the impact of modifying the DEM resolution to investigate the impact of using coarser (but free) DEMs compared to more refined (but expensive) ones such as LiDAR. The study takes place on a small, 12km² catchment in Bavaria near the summit of Zugspitze. The authors find that the choice of DEM is not as critical as first thought, and that the use of global climate products can yield reasonable results in hydrological modelling but that there is still improvements to be made. I have read the paper and found it very interesting and complete. The text flows generally well, although some expressions and sentences don't “sound” right and should be corrected by a native speaker. Scientifically, I have some issues with a few aspects of the work and I also have some suggestions to improve the work and make it more useful to the community. I will start by mentioning the more general points and end with smaller, more technical points.

Author's answer: We would like to thank the reviewer for carefully reading the manuscript and providing us with a very constructive and overall positive feedback. We have thoroughly considered all of his/her comments and address them point by point in the following.

General comments:

1 - The authors use a variety of climate data products to drive the hydrological model to simulate the snow accumulation and melt processes. There are two station datasets (local and from a somewhat distant but similar catchment), satellite products and the ERA20C product. I have a problem with the latter. It can be argued that the 0.05, and to some extent the 0.2 / 0.25 products can be “reasonable” in terms of spatial resolution to represent the 12km² catchment. However, the 125km resolution ERA-20C has a resolution of 125x125 = 15625km², or more than three orders of magnitude difference. The catchment represents less than 0.08% of the tile size. It seems unreasonable to me to include it in the analysis. I think no researcher would use this product for such a small catchment in real world applications. The authors talk about using ERA-20C because of the correction of Gao et al. 2012, but I am positive that using a product such as ERA5-Land (With a 0.1 resolution) would be a better proposition.

Author's answer: This is a very important comment and we are also convinced that showing results of the quite new ERA5-Land product will make our work more up to date. Thereafter, we absolutely agree with your suggestion and performed model runs with ERA5-Land (approx. 9 km) as well as ERA5 (approx. 31 km) data for the updated manuscript version. In our first response, we stated that we could obtain quite good simulation results with these data. Unfortunately, when we checked the ERA5 and ERA5-Land input data again, we found that they were wrong due to a rounding error in the latitude. Subsequently we ran the model again with the right input data. However, the results, we obtained are not satisfactory. Both data sets are too cold and too wet resulting in way too much snow with piling up to distinctive snow towers. We added statements for ERA5 and ERA5-Land in the following lines:

II. 91-94: “To answer the first question, we investigated globally and publicly available meteorological driver data from the Climate Forecast System Reanalysis (CFSR), different versions of the Global Land Data Assimilation System (GLDAS), ERA-20C as well as ERA5 and ERA5-Land datasets and precipitation information from the Climate Hazards group Infrared Precipitation with Stations (CHIRPS).”

II. 215-220: “Regarding the hybrid data, we forced CRHM with the well-known Climate Forecast System Reanalysis (CFSR) dataset provided by the National Centers for Environmental Prediction (NCEP) (Saha et al., 2010). Moreover, we used NASA's Global Land Data Assimilation Systems (GLDAS) (Rodell et al., 2004), ERA-20C (Poli et al., 2016) and the new ERA5 (Hersbach et al., 2020) and ERA5-Land (Muñoz Sabater, 2019). CFSR, GLDAS, ERA-20C, ERA5 and ERA5-Land products combine in situ measurements, remote sensing and atmospheric modelling. ERA5 and ERA5-Land are the latest members of the ERA family and ERA5-Land has a relatively high spatial resolution with 9 km.”

II. 286-287: “The newer and higher resolved ERA5 and ERA5-Land setups are in a similar range with -1.8 °C and -2.5 °C.”

I. 300: “Only precipitation data from ERA5 and CFSR are close to the reference data over the last 30 years.”

II. 308-309: “However, some curves, e.g. the ones of ERA5, ERA5-Land and DWD_Wendelstein almost show a plateau from spring to summer in contrast to the reference.”

II. 487-488: “The same can be observed for the CFSR, ERA5 and ERA5-Land setups, but for those even at both measurement stations.”

II. 532-539: “Regarding the quite new and higher resolved ERA5 and ERA5-Land data, we were quite surprised that both data sets are not able to better capture the meteorological situation in the RCZ. In particular from the 9 km resolution ERA5-Land data, we expected better results. One reason for the bad performance is the direct vicinity of the RCZ to the surrounding much lower forelands. Therefore, even ERA5-Land might be too coarse to capture this situation. The comparably low pixel altitude of this product also supports this assumption. We assume that ERA5 and ERA5-Land could be applied with much more success in a central alpine catchment where the investigated catchment is surrounded by high-alpine regions, which might be more representative for the landscape domain.”

II. 628-631: “Another step in this direction is snow cover data from reanalysis products like ERA5 or ERA5-Land (Hersbach et al., 2020) that might provide information without modelling as well as for model evaluation or initialization purposes. However, its spatial resolution of 31 km and 9 km, respectively, might still be too coarse for heterogeneous high alpine terrain.”

II. 598-600: “To answer our first research question on the potential applicability of globally available meteorological data, we examined data from CFSR, different versions of GLDAS, CHIRPS, ERA5, ERA5-Land and ERA-20C including a specific downscaling approach, as well as a data transfer from another alpine station.”

II. 615-617: “The CFSR as well as ERA5 and ERA5-Land setups performed worst and produced so called ‘snow towers’ in all parts of the catchment due to too low temperatures, leading to the fact that the snowpack could not melt completely, even in summer.”

2 - The GTOPO30 product seems to give reasonable results and the authors state this in several places in the paper. However, it seems that it performs well because it is biased and it is “counteracting” the bias of the meteorological products. Therefore it is better, but for the wrong reasons. I think it would be warranted to add a section (or sentence) in the discussion to clarify this to prevent readers from getting the wrong impression of the quality of GTOPO30. Again, I think users working on very small catchments with high gradients would never use such a coarse product, it was not designed for this.

Author’s answer: We agree with you that the GTOPO30 product was not designed for applications in small and highly heterogeneous catchments. Nonetheless, such a product might be useful in large mountain domains and if computational power is limited. Moreover, we clarified why the GTOPO30 product performed so well at a first glance. The reasonable results derived with the GTOPO30 DEM are, however, just specious due to the coarse resolution, where small scale topographic effects are averaged out. This leads to the fact that the meteorological products are interpolated in a too coarse topographic description of the catchment, which might level out the meteorological data. Therefore, for some parts of the catchment the values are too high, while they are too low in other parts. Nevertheless, it might well capture the mean topographic/meteorological situation of the catchment as it is the case in our study. We added this to our discussion:

II. 558-562: “However, the almost 100 m lower DEM-dependent average catchment altitude of GTOPO30 plays a minor role, since it is still high enough and temperatures are low enough that the snow cover can develop in the model as it was measured. Regarding the catchment scale, the differences among HRUs are averaged out. The GTOPO30 DEM was not created for applications in small basins with highly heterogeneous terrain and we generally would refrain from using it for such applications.”

3 - I think the authors should have compared the snow modelling results they obtain with those from a reanalysis directly, such as ERA5. This could be a much simpler way than using reanalysis meteorological data to drive a hydrological model. It seems to me that this would be a much simpler alternative than using these convoluted methods? I think it could be appropriate to at least mention the possibility here as it fits the bill perfectly: using publicly available global datasets to model snow hydrological modelling in alpine catchments. Furthermore, it could be used to force the initial states of the hydrological model to simulate runoff.

Author’s answer: Indeed, the quite new ERA5-Land reanalysis product includes many very interesting variables such as snow height and SWE as an average over the entire pixel size of the product. We also agree with your statement that integrating ERA5-Land snow would be an interesting addition. In this regard, we promised to make a comparison of the ERA5-Land SWE with specific snow hydrologic variables MSWE, DMSWE, snow cover duration, ablation shown in Table 6 for the entire RZC. However, we suggested this when we had the old, false

data as explained in our response to comment 1. As the right ERA5-Land data represents values of a pixel that is approx. 800 m lower such a comparison does not make sense in our opinion. This is supported by the temperature and precipitation data that are out of the climatological range. Therefore, we refrained from showing such a comparison. Nonetheless, we expect further developments regarding such data and stated that in Section 4.3 and specifically mentioned the ERA5 and ERA5-Land products.

ll. 586-589: "Another step in this direction is snow cover data from reanalysis products like ERA5 or ERA5-Land (Hersbach et al., 2020) that might provide information without modelling as well as for model evaluation or initialization purposes. However, its spatial resolution of 31 km and 9 km, respectively, might still be too coarse for heterogeneous high alpine terrain.

4 - I notice that there is no section on model calibration, as this model does not require calibration but is instead "parameterized" to the environment. I suggest adding this information as it is atypical for a model to not require calibration.

Author's answer: We followed your recommendation and added the following:

ll. 191-193: "Unlike most hydrological models, CRHM does not require calibration. The only parameters that we adjusted for modelling are the previously named catchment/HRU specific physiographic characteristics."

Specific comments:

Lines 155-160: Do I understand correctly that all precipitations were multiplied by 1.5? The SWE technically also includes the effects of ablation/sublimation/transport, so I think it is dangerous to correct precipitation in this manner as the actual real factor is probably different. Perhaps add some limitations in the text here.

Author's answer: We only corrected snow precipitation with this factor and agree with you that the real factor might be different for the named effects. However, the factor was determined in a way that minimizes these effects. To minimize the influence of sublimation, data with a 10 minute temporal resolution were used for the comparison of SWE and precipitation. Before the comparison, the effect of wind on the measured snow water equivalent was investigated. Therefore, time periods without snow fall, temperatures below freezing and strong wind were checked. This revealed no wind induced snow redistribution that could be captured by the snow gauge. Nevertheless, there are effects which could not be taken into account such as sublimation from the snowflakes during the snow fall event due to strong winds. As you suggested, we added a respective paragraph.

ll. 142-147: "Therefore, we used the factor of 1.5 to correct snow precipitation measured at the DWD station which is used for model forcing. The factor was determined using 10-minute values to minimize the effect of sublimation from the snow cover. However, the potential effect of snow redistribution by wind on this factor could not be detected by analysing the SWE in periods with temperatures below the freezing point, strong winds and without snowfall. Besides, it has to be noted that strong winds rarely occur without precipitation in RCZ, pointing to some limitations of this under catch factor determination."

Lines 296-301: This section is a bit confusing. Also, there are 2 peaks of runoff caused by snow accumulation periods? There are two in the year?

Author's answer: We agree with you that this section was a bit confusing and thus we rewrote it. We just wanted to hint at the fact, that the two in situ measured DWD station datasets (the reference and the Mt. Wendelstein setups) both have their precipitation maximum during the snow accumulation period in March. Furthermore, the reference has a second small peak in November. On the other hand, all other data sets have their maximum in summer.

ll. 302-305: "In general, not all precipitation regimes follow the same pattern. The reference and DWD_Wendelstein data sets follow a regime with predominant precipitation during the snow accumulation period. All other meteorological datasets reach their maximum precipitation in summer but also show a local maximum in March. The CHIRPS precipitation is similar to the reference, however, with a more pronounced maximum in August."

Line 322: missing year for reference "Danielson and Gesch".

Author's answer: We added the year.

Line 336: “Adjusted it to the ALOS...” : Should this be altitude-corrected? Please clarify

Author’s answer: We clarified the sentence. ll. 325-326: “...adjusted it to the ALOS, SRTM and GTOPO30 DEM specific HRU altitude, slope and aspect (Table 2) as explained in Section 2.3”

Line 378: “shorter” should be “less”

Author’s answer: We corrected that.

Line 394: “snow towers” needs to be defined better.

Author’s answer: We added the following. ll. 383-386: “Snow towers are an effect in snow hydrological modelling that occurs mainly at higher altitudes and describes the unrealistically high accumulation of snow over several years. Reasons can be the insufficient description of redistribution processes in the model or unrealistic meteorological forcing data (Freudiger et al., 2017) as the latter will be true in our case.”

Line 413-422: This section is not clear upon reading. I needed to read more of the paper before coming back and understanding this section. Please simplify and/or clarify.

Author’s answer: We made some changes to the text to clarify this section. ll. 394-403: “Regarding the snow depth results simulated by using the model setups with different topographic characteristics on basis of the 2.5 m DEM for the reference run and the three different globally and publicly available DEMs ALOS, SRTM and GTOPO30, the snow depth development was simulated realistically at both measurement sites (Fig. 7g, h, Table 5). This results in R2 values above 0.8, NSE values above 0.77 and MAE values below 0.33 for all three global data setups at the LWD site being close to the statistical values of the reference. For GTOPO30, the quality measures for the DWD site are even slightly better, whereas the ALOS and SRTM setups show slightly weaker NSE, R2 and MAE values at the DWD site (Table 5). The two 30 m DEMs SRTM and ALOS show slight differences in the NSE values at the measurement stations, which can be attributed to differences in the sensing angles and post processing algorithms leading to topographic differences and as a consequence also to differences especially in Qsi as mentioned in Section 3.2.”

Response to Anonymous Referee #2

Review of “The evaluation of the potential of global data products for snow hydrological modelling in ungauged high alpine catchments” Weber et al.

General overview

The authors present an evaluation of snowpack and hydrological modeling outcomes in a well-characterized and gauged catchment using parameterizations derived from regional and global datasets. They also evaluate the influence of three different global digital elevation models on the derivation of model inputs such as slope, aspect, and solar insolation as well as the impact on model results. The goal was to illustrate and quantify the impacts of using these products to estimate snowpack and runoff in ungauged catchments in snow-dominated mountainous regions.

The purpose and need for the study are reasonably clear and such work is important for the advancement of snow and mountain hydrology, in general. While it appears that the work and results are technically sound, I found the manuscript very difficult to follow and as such am unable to fully evaluate their results. I recommend a major revision. Specifically, I recommend: ...

Author’s answer: We would like to thank the reviewer for taking his/her time to read the manuscript and providing us with a very constructive feedback. Furthermore, we thank him/her for mentioning the importance of this work regarding advancements in snow/mountain hydrology. We have thoroughly considered all of his/her comments and address them point by point in the following.

General comments:

1 – I recommend reorganizing the methods and results to more concisely lay out the study (methods then results), including the use of additional tables or figures as recommended in the “Specific comments” section

Author’s answer: We followed your recommendation and restructured the manuscript. Please see also our answers to the respective points to the specific comments.

2 – I recommend rewriting the manuscript with an eye toward brevity and clarity, including working with an English speaking editor. The manuscript is overly long, repetitive, and full of awkward and difficult to follow sentences.

Author’s answer: We followed your recommendation and restructured the paper. We also took care to avoid repetitions and gave the paper to a native speaker for proof reading.

Specific comments:

Introduction: At some point in the introduction, there needs to be a clear statement or bulleted list of the research questions. These questions should then guide the organization of the rest of the paper.

Author’s answer: In the last paragraph of the introduction (ll. 102-114), we now present the structure of the paper, as we agree that this improves readability. Moreover, we formulated our two main goals as research questions with bullet points in the introduction, regarding the overall question on how far it is possible to use globally available input data for snow-hydrologic modelling in ungauged high alpine catchments. ll. 87-90:

- “Is it possible to use globally available meteorological data products to reliably simulate snow depth, specific snow hydrological parameters and runoff in high alpine catchments?”
- “What impact have DEM products with different spatial resolutions on snow hydrological simulations in complex topography?”

Section 2, 3, 4: These appear to be all methods, yet there are partial results mixed in. It would be much clearer, if it were a single methods section with subsections on the catchment, the model method, and the input

datasets. Then the first subsection under results could compare the inputs of the various datasets, followed by the rest of Section 5. Otherwise, it is very difficult to follow.

Author's answer: As noted in the general comments, we restructured the manuscript according to your suggestions. The method section now contains the catchment description, the model description and simulation methods, as well as the input datasets. The former Sections 4.2 and 4.4 were moved as subsections to the results section. The structure now is

- 1 Introduction
- 2 Methods
 - 2.1 The Research Catchment Zugspitze
 - 2.2 Structure of the Cold Regions Hydrological Model
 - 2.3 Meteorological input data and their preparation for model input
 - 2.4 Applied DEMs to describe the land surface characteristics
- 3 Results
 - 3.1 Comparison of the different meteorological input data sets
 - 3.2 Influence of the DEMs and associated land-surface characteristics on meteorological conditions
 - 3.3 Evaluation of simulated snow depth with measured data
 - 3.4 Comparison of snow hydrologically relevant indices and the simulated runoff regime
- 4 Discussion
 - 4.1 Application of global meteorological data
 - 4.2 Application of global DEM products
 - 4.3 General remarks and potential future developments
- 5 Conclusions

Page 5, line 138: The term “gradient” should be replaced with “relief”. “Relief” refers to the absolute difference in elevation of a region, whereas “gradient” refers to the slope.

Author's answer: Thank you for pointing out this linguistic subtlety to us. We changed “gradient” to “relief”.

Page 5, line 142: The term “knee wood” should be replaced by “krumholz”, which is the more internationally recognized term for dwarf woody alpine vegetation.

Author's answer: We replaced “knee wood” by “krumholz”.

Introduction, lines 112-124: Much of this should be moved to the methods section as it is repetitive.

Author's answer: We shortened this paragraph. Some aspects are still mentioned in ll. 102-114. The definition and explanation of HRUs is now in Section 2.2 ll. 153 ff.

Page 5, line 160-199: A figure or table would greatly help the reader understand the model structure and underlying component calculations.

Author's answer: CHRM is a widely used model for investigating snow hydrological questions. The modules and underlying component calculations are described in detail in Pomeroy et al. (2007) and the other papers cited in the paragraph. We included a reference and pointed out more clearly the structure and description of the model in the updated version. In addition, a very general structure description for the application of the model in the RCZ is presented in Weber et al. (2016) – we included the reference to this paper in this context as well. The paragraph can be found in ll. 175 ff.

Page 8, line 201-202: Why is the reference referred to as representing the “gauged basin mode”? It would help if the methods section started out with statement about how the study is organized, such as “The approach we use is to model snow and runoff in the well-characterized RCZ using the CRHM and in-situ measurements. Then we repeat the modeling process using alternative datasets. In order to do this, we first describe the watershed, the model, the derivation of the datasets...” Moreover, the derivation of the reference model needs to be more clearly explained. Are you using a single meteorological station's data (Mt. Zugspitze DWD station) to drive the catchment model and compare the results at two other stations in the catchment (LWD, DWD)?

Use of standard three-letter station abbreviations for all stations in the study would help clarify this, too.

Author's answer: We followed your suggestion to shortly outline the method steps undertaken at the beginning of the methods section, which should now increase readability (ll. 117-118). The reference setup is referred to as 'gauged basin mode' since it uses in situ measured driver data directly measured within the gauged RCZ.

Indeed, we are using data from a single meteorological station. We pointed this out more clearly in the updated version along with a better description the reference data (ll. 195-196). The reason why we 'just' used the DWD dataset as meteorological input is now clarified in more detail in the manuscript: For the reference setup, we used meteorological data measured at the DWD station on Mt. Zugspitze since it is the best dataset in terms of continuity and data quality measured directly within the catchment. Due to some longer data gaps, especially at the end of the winter season and the summer time, we did not use the meteorological LWD data as driver data. We will add in that besides the gaps in summer, there are also some gaps during winter, mainly at the end of season (ll. 196-198).

In general, we agree that three-letter abbreviations for all stations would be good. For the two in situ stations LWD and DWD in the RCZ we used three-letter abbreviations. To avoid confusion, we named the DWD station situated at Mt. Wendelstein 'DWD_Wendelstein'. We think that for this study it is ok to extend the latter station name, as we also do not use three letters to describe the results of the global meteorological and DEM datasets, which have in general longer abbreviations and whereof it would not make sense to shorten them.

Page 10, line 257: What does PUB stand for?

Author's answer: The acronym PUB stands for predictions in ungauged basins and is defined in l. 67.

Figures 3 and 4: These figures are introduced in the text and appear before section 4.3 which describes the DEM parameterizations. It was unclear, at first, why the DEM datasets were part of the graphs. It would be clearer to present methods in one full section and results in the next. As it is, some results are mixed in with each method section.

Author's answer: We followed your recommendation and restructured the manuscript to clearly distinguish between methods and results. Regarding the two graphs you mentioned (Fig. 3 & 4) we think that in order to exceed a not too high number of graphs in total, and not to be repetitive that adding the information related to the DEMs is definitely ok. However, we mentioned more clearly in the caption that the graphs contain information of the meteorological and DEM setups.

Page 18, line 432: What does 'thunderstruck' mean? Do you mean that there are data gaps to power outages caused by lightning strikes?

Author's answer: We changed it to lightning strikes.

References

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- Weber, M., Bernhardt, M., Pomeroy, J. W., Fang, X., Härer, S., and Schulz, K.: Description of current and future snow processes in a small basin in the Bavarian Alps, *Environ Earth Sci*, 75, 962, doi:10.1007/s12665-016-6027-1, 2016.

Response to Anonymous Referee #3

Comments on the manuscript: “The evaluation of the potential of global data products for snow hydrological modeling in ungauged high alpine catchments” by Weber et al.

This manuscript examines the sensitivity of the simulated snow characteristics in a small alpine basin related to meteorological input data and DEMs using a hydrological model with a sophisticated snow module. Studies like this are of a great importance as high elevation snowpack critically affects alpine ecosystems and water resources in a number of regions in the world. The experiment is comprehensively designed and well executed. The paper is informative and interesting.

Author’s answer: We would like to thank the reviewer for thoroughly reading the manuscript and for providing us with valuable suggestions for improvement. Moreover, we thank him/her for mentioning the importance of this work regarding improvements in the investigation of high alpine water resources which is relevant for numerous (ungauged) catchments worldwide. We have carefully considered all of his/her comments and address them point by point in the following.

General comments:

(1) The analysis of the experimental results and the organization as well as the presentation of the analysis output need improvements,

Author’s answer: In general, we reorganized the paper structure, which was also mentioned as an issue by reviewer 2 and we better streamlined and improved the presentation and analysis of the results as well as the discussion. We also had an eye on avoiding repetitions. As you make some clear suggestions on these points in the specific comments below, please also consider our answers to those points in the following. The general structure of the paper now is:

- 1 Introduction
- 2 Methods
 - 2.1 The Research Catchment Zugspitze
 - 2.2 Structure of the Cold Regions Hydrological Model
 - 2.3 Meteorological input data and their preparation for model input
 - 2.4 Applied DEMs to describe the land surface parameters
- 3 Results
 - 3.1 Comparison of the different meteorological input data sets
 - 3.2 Influence of the DEMs and associated land surface parameters on meteorological conditions
 - 3.3 Evaluation of simulated snow depth with measured data
 - 3.4 Comparison of snow hydrologically relevant indices and the simulated runoff regime
- 4 Discussion
 - 4.1 Application of global meteorological data
 - 4.2 Application of global DEM products
 - 4.3 General remarks and potential future developments
- 5 Conclusions

(2) Insufficient evaluation of the reference simulation across entire HRUs is also problematic. Because of the large elevation range within the test basin (RCZ), the snow field within the basin is expected to vary widely. Without the evaluation of the elevation dependent model performance with the reference data, the accuracy of the reference run cannot be well established.

Author’s answer: We agree with you that the snow cover in the basin can vary widely as it is the case in any high alpine region with complex topography and that it would be valuable if it was possible to evaluate the accuracy of the elevation dependent snow cover variation in more detail.

We discussed this point in more detail in the updated version and added the paragraph in ll. 486-491: “It could be argued that the evaluation of the simulation results on an HRU basis is insufficient due to the large heterogeneity of the catchment. However, the chosen HRU delineation partly accounts for this heterogeneity since it directly relies on the LIDAR measured dominant snow depth distribution (Weber et al., 2020). Moreover, dominant snow depth patterns in high alpine regions are largely persistent over the years as e.g.

Grünewald et al. (2013) show for various high alpine catchments. We are thereafter confident that the spatial distribution of the snow cover is quite well simulated.”

Moreover, we think that the availability of two snow depth gauges in such a small catchment is already a lot and we are quite lucky with the situation in this well gauged high alpine catchment. Of course, it would be nice if more or even each of the ten HRUs would be represented with a snow depth gauge; however, this is not realistic and will most probably not be found in such a potential station density anywhere in the world. Thereafter, the only possible accuracy analysis for this study period is the investigation of snow depth with data measured at the DWD and LWD station which are situated at least at different altitudes. Moreover the locations of these stations were chosen from the two operating institutions DWD and LWD to be as representative as possible for the respective altitude, e.g. no shadowing against wind but also no particular exposure, relatively flat terrain, etc.

(3) The lack of the analysis and evaluation on monthly time scales (i.e., annual-cycle resolving analysis). This is important because the forcing and snow fields undergo strong seasonal cycle, hence, seasonal cycle-resolving analysis & evaluation can be useful in assessing key sources of the simulation errors. This will also help to link the seasonal cycle of the runoff to monthly snow ablation.

Author’s answer: In addition to the already presented mean monthly precipitation (Fig. 4) and the mean monthly runoff (Fig. 8), we present the mean monthly SWE in Fig. 8 and the mean monthly Qsi in Fig. 6 in the revised version of the manuscript.

Regarding Qsi, we wrote in ll. 307-310: “The 10-year monthly mean of incoming shortwave radiation (Qsi) for the RCZ, relevant for snow cover depletion, is presented in Fig. 6. All setups show very similar data. However, some curves, e.g. the ones of ERA5, ERA5-Land and DWD_Wendelstein almost show a plateau from spring to summer in contrast to the reference. A possible reason might be that convective clouds form during this period while Mt. Zugspitze is frequently above the condensation level.”

Among the setups that showed plausible results, the accumulated amount of snow, which is a result of the precipitation regime, is the most important factor for the magnitude of runoff. Moreover, it also determines the period in which snow is available for melt and thus the occurrence of the peak runoff. Regarding the topographically different setups, the individual radiation budget of the HRU influences the timing and magnitude of runoff. We discussed this in ll. 450-469.

The precipitation regimes presented in Fig. 4 show strong differences. The reference’s regime has its peak precipitation during the snow accumulation period whereas others, like ERA5, have their peak precipitation during summer. We presented this in ll. 300-306.

(4) The terminology “topography parameterization” is confusing. The true meaning of “topography parameterization” in the manuscript is “characteristics of topographic parameters” such as slope and azimuth that vary according to the DEM resolutions. “Parameterization” typically means representing unresolved properties using resolved values or representing a property using a related variable(s).

Author’s answer: You are right, this was misleading. We changed the wording in the text where necessary.

(5) The writing is OK, but I found occasional awkward/unfamiliar sentences. As a speaker of American English as the second language, I don’t like to suggest any specific changes in grammar and writing. I strongly recommend the authors to consult native English speakers to go through the entire writing.

Author’s answer: We gave the manuscript to a native speaker for prove reading.

Specific comments:

(1) The authors present results from elaborate model runs and analyses, but the key messages are not clearly presented and are sometimes confusing.

Author’s answer: We revised the entire paper and streamlined the text so that the key messages become clear. In the introduction, we added our two main research questions as bullet points, which are both dealing with the overall question of how far it is possible to use globally available input data for snow-hydrologic modelling in ungauged high alpine catchments.

II. 87-90:

- “Is it possible to use globally available meteorological data products to reliably simulate snow depth, specific snow hydrological parameters and runoff in high alpine catchments?”
- What impact have DEM products with different spatial resolutions on snow hydrological simulations in complex topography?”

The answers to our questions and thus our key messages are now clearly presented in the abstract (see comment below) and the conclusions part.

II. 625-636: “Based on our results, the answer to our first research question is that it is not possible to exclusively use the tested global meteorological data products to reliably simulate snow depth and further snow hydrological parameters as well as runoff in the high alpine RCZ. One reason is that up to now, these global products are neither able to describe the meteorological heterogeneity of such complex catchments with steep terrain, nor its average conditions. This is reflected in a range of 3.5°C in the catchment mean decadal temperature and 1510 mm in the catchment mean decadal annual precipitation sum over all input data for the RCZ. However, we assume that results could be different if the investigated catchment was not a topographic outlier in the landscape region as it is the case with RCZ with its adjacent lower forelands. The answer to our second question is that compared to the influence of the different meteorological forcings on simulated snow hydrological parameters, the influence due to different characteristics of topographic parameters like slope, aspect and altitude due to different DEM products is smaller, even in complex terrain. Nonetheless, there are considerable differences mainly due to DEM product dependent variations in the radiation balance and due to mean HRU altitude induced variations in temperature and precipitation.”

Moreover, we summarized in a few sentences the more specific key findings of the results part in each of the four sub-sections regarding

- the comparison of the different meteorological input data sets (see II. 311-317 (Section 3.1)),
- the influence of the DEMs and associated land surface parameters on meteorological conditions (see II. 345-349 (Section 3.2)),
- the evaluation of simulated snow depth with measured data (II. 404-412 (Section 3.3)),
- and the comparison of snow hydrologically relevant indices and the simulated runoff regime (II. 492-496 (Section 3.4)).

At the end of the conclusions section, we gave the following outlook (see II. 636-641): “Despite the weak performance of the global meteorological products, we assume that they might produce better results if the analyzed catchment is more representative for the surrounding larger scale landscape region. Furthermore, we expect a growing importance of such data in future snow hydrological modelling, also in ungauged basins, due to their constant and rapid evolution including temporal and spatial refinements. In order to generalize the findings from our study and to intensively test newly developed meteorological and snow hydrological products, we suggest to conduct further investigations in the well monitored catchments from INARCH.”

(2) Please improve Abstract so that the key findings in the experiment are presented more concisely clearly. Separating the sensitivity to DEMs and associated orographic parameters from the sensitivity to the forcing data may help organizing with more clarity.

Author’s answer: We improved the abstract according to your suggestions. Key questions and the main findings are now clearly pointed out:

Abstract: “For many ungauged mountain regions, global datasets of different meteorological and land surface parameters are the only data sources available. However, their applicability in modelling high alpine regions has been insufficiently investigated so far. Therefore, we tested a suite of globally available datasets by applying the physically-based Cold Regions Hydrological Model (CRHM) for a 10-year period in the gauged high alpine Research Catchment Zugspitze (RCZ), which is 12 km² and located in the European Alps. Besides meteorological data, snow depth is measured at two stations. We ran CRHM with a reference run with in situ measured meteorological data and a 2.5 m high-resolution digital elevation model (DEM) for the parameterization of the surface characteristics. Regarding different meteorological setups, we used ten different globally available datasets (including versions of ERA, GLDAS, CFSR, CHIRPS) and additionally one transferred dataset from a similar station in the vicinity. Regarding the different DEMs, we used ALOS and SRTM (both 30 m) as well as GTOPO30 (1 km). The following two main goals were investigated: a) the reliability of simulations of snow depth, specific snow hydrological parameters and runoff with global meteorological products and b) the influence of different global DEMs on snow hydrological simulations in such a topographically complex terrain. The range between all setups in mean decadal temperature is high at 3.5°C

and for the mean decadal precipitation sum at 1510 mm, which subsequently leads to large offsets in the snow hydrological results. Only three meteorological setups, the reference, the transferred in situ dataset and the CHIRPS dataset, substituting precipitation only, showed agreeable results when comparing modelled to measured snow depth. Nevertheless, those setups showed obvious differences in the catchment's runoff regime and in snow depth, snow cover, ablation period, the date and quantity of maximum snow water equivalent in the entire catchment and in specific parts. All other globally available meteorological datasets performed worse. In contrast, all globally available DEM setups reproduced snow depth, the snow hydrological parameters and runoff quite well. Differences occurred mainly due to differences in radiation model input due to different spatial realizations. Even though SRTM and ALOS have the same spatial resolution, they showed considerable differences due to their different product origin. Despite the fact that the very coarse GTOPO30 DEM performed relatively well on the catchment mean, we advise against using this product in such heterogeneous high alpine terrain since small-scale topographic characteristics cannot be captured. While global meteorological data is not suitable for sound snow hydrological modelling in the RCZ, the choice of the DEM with resolutions in decameter-level is less critical. Nevertheless, global meteorological data can be a valuable source to substitute single missing variables. For the future, however, we expect an increasing role of global data in modelling ungauged high alpine basins due to further product improvements, spatial refinements and further steps regarding assimilation with remote sensing data."

(3) The statement in Abstract, L23:24, contradicts the statement in Conclusion L625-634.

Author's answer: We rewrote the entire abstract including this statement to avoid contradiction.

(4) L149:151: How the mean snow field (e.g., SWE, SCA) over the entire RCZ is calculated? This is among the key evaluation variables.

Author's answer: CRHM calculates for each HRU and each time step values of SWE and snow depth. For these calculations, CRHM requires a meteorological input for each HRU. Since such data is not available for each HRU, we used the method developed by Liston and Elder (2006) to generate it for each HRU. We use the modelled values of SWE and snow depth to calculate the area weighted mean value for the entire RCZ. We are confident, that the chosen HRU delineation is suitable for realistic snow cover simulations since it is the same as in Weber et al. (2020) who established these HRUs and evaluated them with spatially distributed LIDAR snow depth measurements. We described this better in the revised manuscript (ll. 155-167) and hope that this is the answer to your question since L149:151 do not deal with the calculation of the mean snow field over the entire RCZ. The snow indices as they are analyzed in Section 3.4 are HRU-area weighted catchment means.

(5) L156:157: The met data at LWD are not used as the forcing. Why the snow precipitation at LWD is corrected for undercatch?

Author's answer: The meteorological data at the LWD station, as already stated in the manuscript, has some longer data gaps (especially at the end of the winter seasons and during summer time in the study period of 2000-2010), which is why we refrained from using it as meteorological forcing data. Since 2014, a snow scale was installed at the LWD station, which enabled us to use measured SWE for more recent years as a good possibility to investigate the precipitation undercatch. Similar to the literature reported undercatch of snow precipitation of up to 50% (WMO, 2011; Grossi et al., 2017), we found out in our former study (Weber et al., 2020) that we also can expect an undercatch of 50% for the RCZ. Assuming that undercatch variations in precipitation over the entire catchment are rather negligible, the DWD snow precipitation has been corrected with the factor derived at the LWD station.

We clarified this in the revised manuscript. ll. 142-147: "Therefore, we used the factor of 1.5 to correct snow precipitation measured at the DWD station which is used for model forcing. The factor was determined using 10-minute values to minimize the effect of sublimation from the snow cover. However, the potential effect of snow redistribution by wind on this factor could not be detected by analysing the SWE in periods with temperatures below the freezing point, strong winds and without snowfall. Besides, it has to be noted that strong winds rarely occur without precipitation in RCZ, pointing to some limitations of this under catch factor determination."

(6) Spell out HRU at its first appearance.

Author's answer: We spelled it out at its first appearance

(7) L227: Check the spatial resolution of the CFSv2 data. The finest resolution I could find is T382 that correspond to approximately 0.313degree.

Author's answer: We checked the spatial resolution again and thank you for the hint! In course of this we realized that we misunderstood the data description of the database where we downloaded the data. As you wrote, the spatial resolution of CFSv2 is approx. 0.313 degree. However, CFSv2 only exists from 2011 onwards. All data before and thus the data we used is data from the Climate Forecast System Reanalysis (CFSR) which has a 0.2 degree resolution. We changed this throughout the manuscript and renamed this setup CFSR.

(8) L265: Section 3 → Section 4

Author's answer: We now refer to the right section.

(9) L288-291: I don't understand what "contrary development" indicates. This sentence is ambiguous. Please provide more explanations.

Author's answer: We referred to the wrong figure and corrected that. It is Figure 3b instead of 3a. This shows that the precipitation situation in the reference and at the DWD_Wendelstein setup is contrary over the years 2005 – 2010.

(10) L293-295: This sentence is ambiguous. Please rewrite.

Author's answer: We rewrote the paragraph. ll. 302-306: "In general, not all precipitation regimes follow the same pattern. The reference and DWD_Wendelstein data sets follow a regime with predominant precipitation during the snow accumulation period. All other meteorological datasets reach their maximum precipitation in summer but also show a local maximum in March. The CHIRPS precipitation is similar to the reference, however, with a more pronounced maximum in August. As we focus on the seasonal snow cover, we assume that precipitation amount and regime is of major importance for snow cover simulations."

(11) L288:289: This sentence cannot explain the peak in March.

Author's answer: We guess you mean L.298-298 and thank you for pointing this out. Indeed, this sentence made no sense and was deleted.

(12) May change the title of Section 4.3. I suggest "Landsurface parameterization on basis of DEMs" → "Sensitivity of the land-surface parameters to DEMs (or DEM resolutions)"

Author's answer: We changed the title of the respective section to "Applied DEMs to describe the land surface parameters" since we think this fits even better to the content.

(13) Also suggest a new title for Section 4.4: "Influence of the DEMs and associated land-surface parameters on meteorological conditions"

Author's answer: We changed the title of the respective section according to your suggestion.

(14) L339: 100 m → 200 m (195 m). (200m elevation difference also corresponds to a lapse rate of 5K/km, slightly more stable than the standard atmosphere which is understandable over a cold surface like snow/ice.)

Author's answer: The difference in catchment mean altitude is 105 m which is roughly 100 m. We wrote the exact value (l. 327) instead of roughly 100 m to avoid confusion and also added the catchment mean values in Table 2.

(15) Section 5: The evaluation based only on NSE and R2 is insufficient. Need more metrics, at least the 'mean bias' and RMSE. Also provide additional evaluations for each month (i.e., annual-cycle resolving model evaluations)

Author's answer: We additionally provided the mean average error (MAE) in Table 5. For the analysis of the results in Section 3.3, we fully took this statistical measure into account (see ll. 357-404).

(16) Section 5.1: If there are problems in evaluating the snow simulations at DWD and DLW as stated in Section 5.2, how can you justify evaluating the daily snowdepth against the observations at these sites?

Author's answer: In Section 3.4 (formerly 5.2) we stated that we do not validate the snow cover duration and the number of ablation days with measured values. This is mainly due to data gaps at the LWD station in spring, which makes it impossible to determine the melt out day as well as the exact length of snow cover duration for some years. In addition, at the DWD station, the melt out day may be delayed due to the fact that the snow stake is installed directly on the small glacier 'Nördlicher Schneeferner'. This might bias the definition of snow cover duration and ablation days.

Nonetheless, we can very well quantitatively validate the daily modelled snow depth with the measurements for the entire season when data is available (this is true for most days except the days, where we had data gaps at the LWD station at the end of the season). Of course, there might be years regarding the DWD station, in which the model is not able to capture the exact melt out date of the snow cover due to the mentioned effect (ice is conserving the snow cover). However this can almost be neglected in a day based comparison since this snow cover is usually very thin and hardly affects the total seasonally accumulated snow volume. We clarified the above mentioned points in more detail in the updated version.

ll. 362-366: "The ice has a cooling effect on the snow cover which leads to a delayed melt out of a thin snow cover compared to a faster melt of snow directly on top of rocks. This effect is particularly strong in spring and autumn, and is not considered in the model. However, this effect is negligible in case larger amounts of snow are piled up on top of the glacier, as it is the case for most of the time during the winter season. Thus, modelled snow cover at the DWD station lasts on average 20 days less than the measured one."

(17) L395: "snow towers pile up" → Is this due to the forcing errors or model errors or combined forcing-model errors?

Author's answer: This is a well-known effect, which might occur if the meteorological input is not realistic, which is the case for some input setups we showed. The model structure, however, is always the same in each of our setups and such snow towers could never be observed when forcing the model with in situ measured meteorological data. The reason for this snow towers is mentioned in the following sentences in ll. 386-387: "These simulated snow towers are caused by much lower temperatures and much higher precipitation compared to the reference (Section 3.1)."

To clarify this in more detail we also added the following definition in ll 383-386: "Snow towers are an effect in snow hydrological modelling that occurs mainly at higher altitudes and describes the unrealistically high accumulation of snow over several years. Reasons can be the insufficient description of redistribution processes in the model or unrealistic meteorological forcing data (Freudiger et al., 2017) as the latter will be true in our case."

(18) L395: 'downscaled temperature' → 'temperature downscaling'

Author's answer: We guess the reviewer means L399. We changed it.

(19) L406:421: This paragraph repeats the statements in the previous paragraph. May be removed and present a summary of this paragraph in Conclusion.

Author's answer: We removed this paragraph and we revised the manuscript with an eye on deleting repetitions in general.

(20) Section 5.2

A) Statements in this section is too qualitative without solid supports from acceptable level of evaluations.

Author's answer: We rewrote parts of this section (now Section 3.4) including more 'hard' numbers.

See ll. 430-437: "For the catchment mean and varying meteorological products, Table 6 shows an overestimation of MSWE in the CHRIPS and DWD_Wendelstein setup of 15% and 20%, respectively, compared to the reference setup. This overestimation is related to the 378 mm and 534 mm higher mean annual precipitation sum in the CHIRPS and the DWD_Wendelstein setups. For the DWD_Wendelstein setup the combination with the lower mean temperature leads to a 12 days later occurrence of the DMSWE, a 22 days longer snow cover duration and a two weeks longer ablation period on average for the entire catchment. In contrast, the later DMSWE, the longer snow cover duration and the longer melting period of the CHIRPS setup is solely due to higher winter precipitation and the resulting higher MSWE. As a consequence, the higher MSWE

values in these two meteorologically different setups result in a higher runoff of 29 and 30%, respectively (Table 6)."

See also II. 441-448: "Regarding the simulated snow hydrological results using the setups with different DEMs, the mean catchment MSWE corresponding to the GTOPO30 DEM is closest to the reference with a negligible deviation of 1% (Table 6). This good performance goes in line with the comparison of the measured and modelled snow depth data in Section 3.3 (Table 5). Both 30 m resolution DEMs result in a considerably higher catchment mean MSWE than the reference setup at +18% in the ALOS setup and +17% in the SRTM setup. The GTOPO30 setup shows for the DMSWE and the length of the snow cover as well as the ablation periods also the smallest deviations compared to the reference setup. The mean catchment DMSWE of the ALOS and SRTM setups occur eleven and 12 days later, respectively, when compared with the reference setup. Moreover, the snow cover duration in these setups is almost one month longer, which is reflected in their higher number of ablation days, too."

B) If missing data and site characteristics at LWD and DWD, respectively, prevent using these data for evaluation, how can we trust the reference data are accurate enough for evaluating model data?

Author's answer: This concern is reasonable, probably for all measurement stations worldwide. In the following, we provide you some specific information to the stations we used.

DWD station: this data is quality controlled by the DWD, the German Weather Service. Measurement errors are filtered and there are virtually no gaps.

LWD station: The instruments of this station are regularly maintained by the LWD, the Bavarian Avalanche Warning Service, since they rely on good quality data. All the large gaps at this station used to occur in late spring when the avalanche season was over and the LWD no longer produces an avalanche warning bulletin. So they simply had no need for data in this time of the year. Consequently, the data recorded at that location is very accurate and reliable.

Regarding the locations of the two measurement stations, both institutions chose locations, which are representative for the respective altitude and general site conditions. Further restrictions particular concerning the use of DWD data for snow cover duration evaluations are explained in the answer to your comment 16.

C) If DMSWE cannot be validated, how can MSWE can be validated? MSWE is supposed to occur on DMSWE.

Author's answer: We never validated the MSWE and never stated that in our manuscript. We do not have measured SWE in the investigated time period. We only have measured snow depth at the LWD and DWD stations to validate our model runs. In our analysis in Section 3.4 (former 5.2) we compared all SWE related indices resulting from the three best meteorological setups and all DEM setups to the reference run that was forced with in situ measured driver data and parameterized with the high resolution 2.5 m DEM.

(21) L457:470: How can the data of largest warm bias, precipitation underestimation and insolation overestimation perform best in simulating the amount and timing of runoff? This may indicate major flaw in the model physics and/or forcing combinations. Need further discussions. This also indicates the need for a more rigorous evaluation of the reference data and other model data.

Author's answer: We rewrote the respective passage in the manuscript to clarify:

II. 452-465: "The fact that the snow cover indices of the GTOPO30 setup are so close to the reference entails a very similar runoff behavior despite it has the largest warm bias, precipitation underestimation and insolation overestimation (Table 4). To explain this, the differences in the process of discharge formation have to be examined. Although the GTOPO30 DEM is lower in altitude than the reference DEM, it is still high enough that monthly mean temperature is below freezing until March and only at 0.1 °C for April while it is at -0.7 °C April in the reference. Both, the reference and GTOPO30 setup are clearly above freezing with 4.2 °C and 5.1 °C, respectively, in May. The main melt, which is temperature induced in spring starts almost at the same time in both setups illustrated by the SWE decline in Fig. 8. Melt induced runoff starts slightly earlier in the reference setup (Fig. 8), because there are HRUs which receive considerably more Q_{si} than in the GTOPO30 setup (Fig. 4). Nonetheless, the major part of the RCZ receives more Q_{si} in the GTOPO30 setup. The longer the days last in spring, the stronger becomes the radiation induced melt effect which is why in the results of the GTOPO30 setup, melt induced runoff becomes stronger in June than in the reference setup. In summer the lower elevation has a greater effect on the temperature than in winter which leads to higher melt induced runoff in the GTOPO30 setup in June and July. This is illustrated in the faster decline of SWE in June and July in Fig. 8. This effect decreases in August since less snow is left to melt."

(22) L475: Monthly Qsi needs to be evaluated as it is directly involved in snowmelt.

Author's answer: As we explained in our answer to comment 3, we added Fig. 6 to present the mean monthly Qsi for the RCZ. Moreover, we added in ll. 307-310 the following: "The 10-year monthly mean of incoming shortwave radiation (Qsi) for the RCZ, relevant for snow cover depletion, is presented in Fig. 6. All setups show very similar data. However, some curves, e.g. the ones of ERA5, ERA5-Land and DWD_Wendelstein almost show a plateau from spring to summer in contrast to the reference. A possible reason are convective clouds that form during this period while Mt. Zugspitze frequently is above the condensation level."

Among the setups that show plausible results, the accumulated amount of snow, which is a result of the precipitation amount and regime, is the most important factor for the magnitude of runoff. Moreover, it also determines the period in which snow is available for melt and thus the occurrence of the peak runoff.

(23) L506:507: This may be an overstatement for the CHRPS data. CHRPS yields good results for snow cover and runoff, but not in NSE and R2 of the daily snow depth.

Author's answer: The NSE and R² of the CHIRPS setup for daily snow depth is clearly the best among all presented globally available meteorologically different setups. Despite the quality measures are lower than for the reference, they are still considerably high. Studies from Moriasi et al. (2007) or Rahman et al. (2013) come to the conclusions that NSE values greater 0.5 can be considered as good in hydrological modelling.

(24) 511-530: Can the inter-model difference also be related to their snow models? Do all of these models use the same snow model?

Author's answer: The inter-model difference can also be related to the used snow models. As explained in the paragraph, the studies cannot be compared one by one since they are conducted on different scales, in different (micro-)climatic regions and also with different models. However, these studies show that it is generally possible to obtain reasonable results with global data, if conditions are appropriate, e.g., data is available for downscaling or the investigation of larger scales. We tried to improve the respective passage in the manuscript.

(25) Please discuss the poor NSEs and R2s with the transferred and CHRP data.

Author's answer: As stated in our answer for comment 23, we consider the quality measures of the CHIRPS setup to be good. According to Moriasi et al. (2007) an NSE > 0.5 can be considered as satisfactory in hydrological modelling. This threshold is also used in alpine hydrological modelling (e.g. Rahman et al., 2013). We added the following in ll. 370-379 : "The NSE values of the CHIRPS setup range between 0.54 and 0.55; the R² values range between 0.66 and 0.79 (Table 5). The reason for the good performance of the CHIRPS setup is that the forcing data is the same as in the reference setup except from precipitation. The overestimation of snow depth in some years, especially at the DWD site, is due to the higher precipitation values as described in Section 3.1. For the snow depth simulations based on the transferred data from the alpine DWD AWS at Mt. Wendelstein, we also obtained reasonable but weaker results, e.g., with an R² of 0.59 and 0.66 at LWD and DWD, respectively but higher MAE values. The NSE values at the stations are 0.33 and 0.11, respectively. The statistical values are rather moderate, but still hint at a plausible representation of the temporal development compared to the results of the other global data. The low NSE values can be traced back to the years 2002, 2005 and 2010 in which the peaks are strongly overestimated in the DWD_Wendelstein run at both stations. Both, the CHIRPS and DWD_Wendelstein show temperature and precipitation data which are within the total range of climatology."

(26) L550-551: This statement ignores the substantial differences in runoff between GTOPO30 and ALOS/SRTM.

Author's answer: The statement you mention here is about the modelled snow cover development and not the runoff. Regarding the snow cover, the difference between the different DEM setups is indeed rather negligible.

(27) L571: "the choice of the DEM has far less impact" Incorrect. DEMs have large impacts on the simulated streamflow annual cycle

Author's answer: We rewrote the paragraph (see ll. 367-369): "The presented results show that the right choice of the meteorological forcing data still is the biggest challenge for snow hydrological modelling in

ungauged alpine headwater catchments. The choice of the DEM has far less impact on the snow cover modelling but can still result in considerable differences in snow cover as well as runoff generation.”

(28) L600:604: Misleading. The ‘borrowed’ data performed poorly in terms of NSE and R2.

Author’s answer: We rewrote the paragraph (see ll. 608-610): “Those two plausible setups are the CHIRPS setup in which precipitation only is substituted whilst taking all other meteorological data from the reference in situ data set, and the meteorological dataset which was ‘borrowed’ from another alpine AWS in the catchment’s greater vicinity at Mt. Wendelstein.”

Regarding the quality measures, please refer to our answer to comment 25, too.

(29) L510-512: There is a mystery. How can the forcing data sets of such a wide variation can produce such similar simulations? This needs answers from the authors.

Author’s answer: We are not sure if you really mean L510-512 since we did not find a statement there that fits to your comment. We assume you perhaps could have meant L610-612, which is about differences due to DEMs. Our answer why the different DEM setups performed quite similar is that on the one hand, the ALOS and SRTM DEMs have the same spatial resolution, which is high enough that the HRU surface characteristics are relatively well represented as with the reference DEM. On the other hand, if results at the catchment scale are considered, in the coarser GTOPO30 setup, differences are leveled out. Moreover, as written in our answer to comment 21, the effect of altitude and enhanced radiation input is not as strong in winter, as in summer. Nonetheless, regarding the individual HRUs there are considerable differences as Figure 9 illustrates.

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