



Estimation of evapotranspiration through an improved daily global solar radiation in SEBAL model: a case study of the middle Heihe River Basin

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Abstract: The agricultural activities, hydrologic cycle, and ecological environment are seriously influenced by evapotranspiration (ET), especially in arid and semi-arid areas. A new method for estimating daily global solar radiation (GSR) over rugged terrains in the middle Heihe River Basin is developed on the basis of Iqbal model C. And with the land surface parameters retrieved from multisource remote sensing data, a daily surface ET on June 21–24, 2009, is simulated by using surface energy balance algorithm for land (SEBAL) model. The results show:

1. An improved daily GSR with a resolution of 100 m×100 m is implemented. The mean absolute bias error (MABE) is 9 W/m², and the mean absolute relative bias error (MARBE) is 2.5%. The MABE of the daily GSR using the SEBAL model is 122.2 W/m², and the MARBE is 33.9%.

2. The spatial distribution of the daily GSR is more reasonable using the improved model than the original model. The GSR is larger on a sunny slope (an open place) than on a shady slope (a rugged place).

3. Bringing the new model into SEBAL significantly improves the accuracy of the ET. The MABE of ET decreases from 2.1 mm (original scheme) to 0.6 mm (improved scheme),



26 and the MARBE declines from 44% to 13% accordingly. Moreover, the spatiotemporal
27 resolution of the ET simulation is effectively improved by the combined moderate-resolution
28 imaging spectroradiometer and thematic mapper surface parameters.

29 4. All highest ET value appeared in all types of water bodies, followed by farmland,
30 forest, wetland, and residential areas, the lowest values appeared over bare rock land. The
31 water consumption in these areas is dominated by agriculture. The new results provide better
32 theoretical basis and scientific guidance for ecosystem protection and sustainable utilization
33 of water resources.

34

35 1. Introduction

36 About 60% of global precipitation is consumed by evapotranspiration (ET),
37 while 99% of water in farmland system is consumed by ET (Kit, 2000). As an
38 important part of water cycle, ET is of great significance for understanding the water
39 cycle process, regulating the hydrological process by vegetation, and rational
40 utilization of limited water resources (Sellers et al., 1996), especially in arid and
41 semi-arid areas. Therefore, the accurate determination and estimation of the daily ET
42 can satisfy the requirements of not only agricultural production but also other
43 ecological water requirements and provide scientific guidance for the rational
44 utilization and distribution of water resources (Nie et al., 2004; Rahman et al., 2016),
45 especially during the growing season when numerous water resources are exhausted
46 in the middle reaches of the Heihe River Basin in China.



47 Surface energy balance algorithm for land (SEBAL) model (Bastiaanssen et al.,
48 1998a and 1998b) is currently a crucial method for estimating water and heat fluxes.
49 Using remote sensing data and meteorology knowledge, calculating water and heat
50 fluxes on a daily scale is feasible and widely applied (Bastiaanssen et al., 2000;
51 Teixeira et al., 2009; Liu, 2008; Yin, 2014; Li et al., 2010; Ahmad et al., 2014;
52 Usman et al., 2015). Liu (2008) verified that the SEBAL model is sensitive to daily
53 global solar radiation (GSR). Therefore, the daily GSR is an important input
54 parameter for ET estimation using the SEBAL model.

55 The solar radiation that reaches the ground is affected by astronomical,
56 atmospheric (i.e., air molecules, water vapor, and cloud), and surface factors (i.e.,
57 slope, aspect, terrain shading, and surface coverage). Thus, the simulation of GSR in a
58 rugged terrain is significantly complex (Zeng et al., 2005; Yeom et al., 2016; Shi et al.,
59 2018). Geographical information system (GIS) and remote sensing techniques have
60 provided new methods for estimating solar radiation, including the above mentioned
61 factors. Digital elevation model (DEM) data have been extensively used to simulate
62 solar radiation in mountainous areas (Williams et al., 1972; Bocquet, 1984; Dozier
63 and Frew, 1990; Qiu et al., 2003; Chen et al., 2013; Shi et al., 2018). However,
64 research on the daily GSR is limited to accurately estimating climate factors on a
65 daily scale. Chen et al., (2013) developed a DEM-based radiation model to estimate
66 instantaneous clear-sky solar radiation for surface energy balance system to obtain
67 accurate energy absorbed by the mountain surface. However, terrain shading is
68 insufficiently considered. Shi et al., (2018) used Iqbal model C (1983) and DEM to



69 estimate the daily GSR to investigate the effects of topography. The daily GSR was
70 acquired in a relatively coarse grid of 1 km×1 km. However, their method was not
71 applied to a high spatiotemporal resolution estimation of the GSR directly, thereby
72 making the performance of agriculture activities infeasible.

73 In the present study, a distributed model for daily GSR is proposed on the basis
74 of possible solar radiation, which includes astronomical, atmospheric, and land
75 surface factors in a relatively high spatiotemporal resolution. Based on the DEM data
76 with 100 m resolution, weather station information (radiation, wind, and pressure),
77 and remote sensing images of thematic mapper (TM) and moderate-resolution
78 imaging spectroradiometer (MODIS), this study aims to 1) realize the estimation of
79 daily GSR in a 100 m×100 m resolution, 2) compare and validate the estimated results
80 with measured data, and 3) improve the application of simulated daily GSR by
81 calculating ET using the SEBAL model.

82 **2. Methodology**

83 **2.1. Study Area**

84 The study area is located in the middle reaches of the Heihe River Basin in
85 Northwestern China, including Linze County and Ganzhou District, with an elevation
86 range of 1,360–3,600 m and an area of approximately 60,000 km² (Fig. 1). An oasis
87 is situated along the river, surrounded by a desert (Fig. 2). This area is typical warm
88 temperate desert arid climate.



89 **2.2. Data**

90 **2.2.1. Remote sensing data**

91 Landsat TM data can provide high-resolution information about land surface
92 temperature, land cover classification, albedo, normalized difference vegetation
93 index(NDVI), and emissivity. This study adopts the Landsat TM5 image that was
94 captured on June 24, 2009, in which the cloud cover in the study area was below 5%
95 due to a favorable weather condition (Fig. 5a).

96 MODIS Levels 2 and 3 images on June 21–24, 2009, without clouds are
97 collected. NDVI, emissivity, and surface albedo are retrieved by using Level 2
98 products (500 m×500 m). Daily land surface temperature (LST) with a spatial
99 resolution of 1 km comes from Level 3 MOD11A1.

100 **2.2.2. Meteorological data**

101 To supplement the TM and MODIS images, the heights of vegetation and the
102 following meteorological data are required: air pressure, relative humidity, air
103 temperature, sunshine duration, LST, wind speed, net radiation, and daily GSR. In
104 addition, two ET measurements are placed in a farmland and a wetland station (Fig. 1).
105 All the meteorological data must be measured hourly throughout the day.

106 **2.2.3. GIS data**

107 Considering the spatiotemporal resolution of remote sensing data, this study uses
108 the DEM with a spatial resolution of 100 m×100 m and the vector map of the
109 administrative boundary with the map scale of 1:250,000. All data are obtained from



the website of Digital Heihe (<http://heihe.westgis.ac.cn>). The land cover map is created by a computer-assisted visual interpretation at the scale of 1:100,000 on the basis of the TM image (Fig. 2).

2.3. Methodology

2.3.1. Improved daily GSR model

The daily GSR in the SEBAL model is calculated as follows (Bastiaanssen et al., 1998a and 1998b; Chen et al., 2000):

$$R_{a24} = G_{sc} \times d_r \times \int_{\omega_1}^{\omega_2} \cos \theta d\omega, \quad (1)$$

where R_{a24} is the daily GSR, G_{sc} is the solar constant, that is, 1367 W/m^2 , d_r is the earth-sun distance factor (dimensionless), ω_1 and ω_2 represent the solar hour angles (radians) at 5 min after sunrise and at 5 min before sunset, respectively, and θ is the solar zenith angle. $\cos \theta$ is calculated by Chen et al., method(2013).

In Eq. (1), only sun–earth spatial relations on a specific date, slope, and aspect relations are considered. The effects of wet–clean air conditions and the terrain shading are disregarded. The improved daily GSR model includes the above mentioned factors, which can be categorized into two main models.

(1) Determination of GSR over a horizontal surface

The effects of air molecules, O_3 , CO_2 , oxygen, and other mixed gas and water vapor on short-wave solar radiation are considered in this part by using Iqbal Model C. A detailed description is presented in Shi et al., method (2018).

We obtain the daily global radiation under wet–clean air conditions using

$$Q_{0w} = \frac{T}{2\pi} \int_{-\omega_b}^{\omega_b} I_t d\omega, \quad (2)$$



where $-\omega_0$ and ω_0 represent the solar hour angles at sunrise and sunset, T is the length of a day, and I_t is global irradiance.

(2) Determination of GSR over rugged terrains

The effects of aerosol, land surface factors (slope, aspect, terrain shading, and surface cover), and cloud are considered in this part.

The daily GSR received on land surface consists of three parts (Fu, 1983).

$$Q_{\alpha\beta} = Q_{b\alpha\beta} + Q_{d\alpha\beta} + Q_{r\alpha\beta}, \quad (3)$$

where $Q_{\alpha\beta}$ is the daily GSR over rugged terrains, $Q_{b\alpha\beta}$ is the daily direct solar radiation over rugged terrains, $Q_{d\alpha\beta}$ is the daily diffuse solar radiation over rugged terrains, and $Q_{r\alpha\beta}$ is the daily reflected solar radiation over rugged terrains.

To determine $Q_{b\alpha\beta}$, we assume that

$$Q_{b\alpha\beta} = \frac{Q_{0\alpha\beta}}{Q_{0w}} Q_b = R_b Q_b, \quad (4)$$

$$Q_{b\alpha\beta} = Q_{0\alpha\beta} \frac{Q_b}{Q_{0w}} = Q_{0\alpha\beta} \frac{Q_b}{Q} \frac{Q}{Q_{0w}} = Q_{0\alpha\beta} f_b k_t, \quad (5)$$

where $Q_{0\alpha\beta}$ and Q_{0w} are the daily astronomical solar radiation over a terrain and a horizontal surface, correspondingly, and Q_b is the direct solar radiation over a horizontal surface. R_b represents the effects of slope, aspect, and topographic shadow, Q is the daily GSR over a horizontal surface, f_b is the direct component coefficient, and $k_t = a_G + b_G s$ (where s is the percentage of sunshine duration, and a_G and b_G are empirical coefficients) is the clear sky coefficient. The effects of aerosol and cloud are considered in the term $a_G + b_G s$.

Similar to Eq. (5), we derive an expression for the diffuse component.

$$Q_{d\alpha\beta} = Q_d [f_b k_t R_b + V(1 - f_b k_t)], \quad (6)$$



154 where Q_d is the diffuse solar radiation in the horizontal surface, and V is the terrain
 155 openness (terrain openness + terrain shading = 1). The method for determining V is
 156 described by Qiu (2003).

157 $Q_{0\alpha\beta}$ is calculated using Qiu's (2003) model, and Q_b and Q_d are calculated on the
 158 basis of the empirical model exhibited by daily direct solar radiation and diffuse solar
 159 radiation in the radiation station considering the effects of aerosol and cloud.

160 Finally, the reflected radiation from the sloped surface can be computed by the
 161 following expressions:

$$162 \quad \begin{cases} Q_{r\alpha\beta} = Q\rho_g(1-V) & V \leq 1 \\ Q_{r\alpha\beta} = 0 & V > 1 \end{cases} \quad (7)$$

163 where ρ_g is the surface albedo, which can be retrieved from Landsat TM5 and
 164 MODIS09GA. ρ_g is described in detail in Appendix.

165 2.3.2. SEBAL model principle

166 The SEBAL procedure consists of a series of algorithms. In this study, this
 167 procedure is implemented using the ModelMaker module of ERDAS software. The
 168 algorithms solve the complete energy balance equation

$$169 \quad \lambda ET = R_n - G - H, \quad (8)$$

170 where λET is the latent heat flux, R_n is the surface net radiation flux, G is the soil
 171 heat flux, and H is the sensible heat flux.

172 The parameterization of R_n and G is mature. Thus, the core issue of the model
 173 is calculating H and ET . SEBAL model introduces a Monin-Obukhov loop iteration
 174 to estimate H .



175 The above mentioned instantaneous results (H , R_n , and G_0) are substituted in the
 176 energy balance Eq. (8) to calculate instantaneous latent heat flux λET . The daily
 177 time scaling extension of the model is implemented on the basis of the evaporative
 178 fraction (EF) method.

$$179 \quad \frac{\lambda ET}{R_n - G} = EF, \quad (9)$$

180 Following Shuttleworth et al., (1989), the instantaneous EF is assumed to be
 181 similar to its 24 h counterpart (Brutsaert et al., 1992; Crago, 1996) and the assumption
 182 supported by numerous field studies (Bastiaanssen et al., 1998a, b; Morse et al.,
 183 2000).

184 Soil heat flux for 24 h periods is assumed to be nearly 0 for vegetated surfaces
 185 given the canceling effect of positive G during daylight and negative G during
 186 nighttime. Therefore, an actual λET_{24} for the 24-h evaporation can be computed as
 187 $\lambda ET_{24} = EF \times R_{n24}$, (10)

188 where R_{n24} is the daily net radiation. The accuracy of the daily GSR estimation
 189 largely determines the accuracy of R_{n24} estimation.

190 2.3.3. Combination of multi-source remote sensing data for ET simulation

191 Liu (2008) conducted a sensitivity analysis of the SEBAL model and suggested
 192 that the LST and emissivity are highly sensitive to the simulated ET, whereas the
 193 surface albedo, wind speed, NDVI, and aerodynamic roughness of a surface are
 194 slightly sensitive to ET.

195 Only MODIS data can be used to calculate the daily ET because no TM data



196 from June 21 to 23, 2009, are available. Therefore, a co-simulation experiment on the
197 ET simulation, which uses three strategies of land surface parameters, is performed in
198 the present study. TM strategy refers to all the land surface parameters retrieved by
199 the TM image. MODIS strategy refers to all the land surface parameters retrieved by
200 the MODIS image. The TM/MODIS hybrid strategy, which combines the advantages
201 of the first two methods, refers to the LST from the MODIS11A1 image and the
202 surface albedo and NDVI from the TM image. The addition of the TM surface albedo
203 and NDVI can improve spatial resolution, whereas adding the MODIS LST can
204 improve temporal resolution. We can approximate that no change will occur in the
205 surface albedo and NDVI within a few days.

206 The TM image cycle is 16 days with a high spatial resolution, whereas the
207 MODIS image cycle is 1 day. However, the spatial resolution is less than the TM
208 image. Many local features retrieved by the TM image are generalized in the image
209 retrieved by MODIS. The TM/MODIS hybrid combines the advantages of the
210 spatiotemporal resolution of both strategies.

211 **3. Results and discussion**

212 **3.1. Daily GSR**

213 **3.1.1. Accuracy of simulated result**

214 The validation results of the two models are summarized in Table 1. The mean
215 absolute bias error (MABE) of the improved daily GSR model is 9 W/m^2 , and the
216 mean absolute relative bias error (MARBE) is 2.5%. The MABE of the daily GSR in



217 the SEBAL model is 122.2 W/m^2 , and the MARBE is 33.9%.

218 The results show that the GSR on June 24, 2009, calculated in the SEBAL model
 219 ranged from 438 W/m^2 to 465.7 W/m^2 with a mean of 446 W/m^2 , which was much
 220 higher than the improved daily GSR model with a range of $350\text{--}394 \text{ W/m}^2$ and a mean
 221 of 370 W/m^2 .

222 In Eq. (7), the surface albedo is calculated by two remote sensing data. The result
 223 indicates that a slight difference of 0.4 W/m^2 in the daily GSR emerges by using the
 224 surface albedo from the TM and MODIS images on June 24, 2009.

225 3.1.2. Spatial pattern of daily GSR

226 Theoretically, the GSR is larger on a sunny slope (an open place) than on a shady
 227 slope (a rugged place), thereby indicating that the daily GSR is large where the
 228 sunshine duration is long.

229 The spatial distribution of the daily GSR calculated using the SEBAL model
 230 presents discontinuous and improper stripes (Fig. 3a). However, the calculation of the
 231 improved model is more reasonable, because the effects of the spatial position
 232 relations, percentage of sunshine, slope, aspect, terrain shading, and atmospheric
 233 influence are comprehensively reflected.

234 An obvious difference between the SEBAL and the improved models is observed
 235 in the southern area and the Longshou Mountain (Fig. 3b). The daily GSR in the
 236 SEBAL model has reached the maximum value in the two areas. However, the daily
 237 GSR in the improved model is near the minimal value. This result can be attributed to
 238 the terrain shading that included a diffuse and reflected solar radiation in the new



239 model. In addition, the percentage of sunshine is higher in the northwest, thus
 240 implying that sunshine duration is long, and the corresponding daily GSR is high (Fig.
 241 3b). However, these distribution features are unclear in the daily GSR in the SEBAL
 242 model.

243 The daily GSR of 0 °, 45 °, 90 °, and 360 ° in the slope direction of the study area
 244 are statistically analyzed, and then the mean value is calculated (Fig. 4). In terms of
 245 local surface distribution pattern, the daily GSR must reach the maximum value near
 246 the south slope at 180 °. However, the simulated results using the SEBAL model show
 247 that the daily GSR is only the minimum value (Fig. 4a), and the error is corrected by
 248 the improved model (Fig. 4b).

249 Therefore, the improved daily GSR model improves the calculation accuracy and
 250 makes the spatial distribution more reasonable than the original model.

251 3.2. Daily ET

252 3.2.1. Accuracy of simulated Daily ET

253 The comparative analysis of oasis stations (farmland [plain area]) is listed in
 254 Tables 2 and 3. The accuracy of the daily ET is higher in the improved scheme than in
 255 the original scheme, and ET is near the measured ET. The results show that the
 256 improved scheme using the TM/MODIS hybrid strategy is the optimum among the
 257 three models.

258 The simulated ET analysis of the TM/MODIS hybrid strategy of oasis stations
 259 from June 21 to 24, 2009, is summarized in Table 4. The mean measured ET for 4
 260 days was 4.8 mm. The MABE of the improved scheme-simulated ET is 0.6 mm, and



the MARBE is 12%. For the original scheme ET, the MABE is 1.85 mm, and the MARBE is 39%. The mean ET of the wetland station (saline land) in 4 days is 2.2 mm. The simulated ET changes from 3.5 mm (original scheme) to 2.2 mm (improved scheme).

According to the research purpose and China's land use classification system, the study area is divided into 22 underlying surface coverage types (Fig. 2). Therefore, the TM/MODIS hybrid strategy of the original and improved schemes are used to calculate the ET of the 23rd and 24th days, respectively. The mean ETs of the two days in each surface coverage type are calculated and listed in Table 5. The mean measured ET for the two days is 4.6 mm in the farmland station. It changes from 6.3 mm (original scheme) to 4.3 mm (improved scheme). The mean measured ET for the two days is 2.2 mm in the wetland station. It changes from 3.7 mm (original scheme) to 2.4 mm (improved scheme).

In Table 5, the highest ET value has appeared in all types of water bodies, followed by farmland, forest, wetland, and residential areas. The ET values in the Gobi Desert and bare rock land are low. The water consumption in these areas is dominated by agriculture. The mean daily water consumption in the study area is $1.46 \times 10^7 \text{ m}^3$. The farmland, which accounts for 44.8%, has consumed $6.5 \times 10^6 \text{ m}^3$ in the improved scheme, and $9.5 \times 10^6 \text{ m}^3$ in the original scheme.

3.2.2. Spatial pattern of daily ET

Figs. 5 and 6 illustrate that, for all land cover types, except for water bodies, the highest ET value is found over the farmland given the presence of irrigation water. In



283 addition, the ET is low in the area around the oasis. The differences in ET over the
284 various land types mainly depends on the NDVI. A large NDVI value indicates a high
285 ET.

286 In Fig. 5 and 6, the simulation ET based on the improved scheme declines more
287 than the original scheme. In the improved scheme, the desert ET is mainly distributed
288 in the vicinity of 0 mm, and the oasis ET is in the vicinity of 4.8 mm. In the original
289 scheme, the desert ET is mainly distributed in the vicinity of 0 mm, and the oasis ET
290 is in the vicinity of 6.5 mm.

291 The TM strategy has the highest resolution (Figs. 5a and 6a), whereas the
292 MODIS strategy has the lowest (Fig. 5b and 6b); many features retrieved by the TM
293 strategy are generalized in the image. The TM/MODIS hybrid strategy (Fig. 5c and 6c)
294 combines the advantages of the spatiotemporal resolution of both strategies to provide
295 a feasible scheme for the estimation of the daily ET on improving temporal resolution
296 and maintaining a relatively detailed spatial resolution.

297 Figs. 5a and 6a are images with cloud and cloud shadows at the northwest
298 corner, correspondingly. Several parts of the LST image in Figs. 5b and 6b are
299 missing; this phenomenon mostly occurs in desert areas. These areas are smaller than
300 the entire study area and cannot affect the statistical characteristics of various types of
301 land ET.

302 4. Conclusions

303 The new model fully considers the atmospheric and surface factors, especially



304 the influence of terrain shading on diffuse solar radiation and reflected solar radiation
305 over rugged terrains.

306 In order to estimate daily ET accurately, a new method for estimating the daily
307 GSR over the rugged terrains in the middle reaches of the Heihe River Basin is
308 proposed in this study. This method is based on Iqbal model C, fully considering the
309 atmospheric and surface factors, especially the influence of terrain shading on diffuse
310 solar radiation and reflected solar radiation over rugged terrains. The daily surface ET
311 on June 21–24, 2009, is simulated by using the SEBAL model, and a portion of the
312 daily GSR in the SEBAL model is improved. The results can be summarized as
313 follows:

314 The improved daily GSR with a resolution of 100 m×100 m is implemented; the
315 MABE of the simulated results is 9 W/m², and the MARBE is 2.5%. The MABE of
316 the daily GSR in the SEBAL model is 122.2 W/m², and the MARBE is 33.9%.

317 Theoretically, the spatial distribution of the daily GSR is more reasonable using
318 the improved model than the original model. The GSR is larger on a sunny slope (an
319 open place) than on a shady slope (a rugged place). In addition, if the percentage of
320 sunshine in the northwest is high, then the daily GSR is also high. The estimated
321 result of the improved daily GSR model is consistent with the distribution law, but the
322 estimated result of the SEBAL daily GSR model cannot fully reflect.

323 The co-simulation experiment on the ET simulation is designed for three
324 strategies, namely, TM, MODIS, and TM/MODIS hybrid strategies. They are used in
325 the original SEBAL model (original scheme) and the improved daily GSR SEBAL



326 model (improved scheme). The results show that the improved scheme ET is more
327 accurate than the original scheme. Moreover, the TM/MODIS hybrid strategy in the
328 improved scheme is the most reasonable in terms of accuracy and spatial distribution.

329 The simulated ET declines more based on the improved scheme than based on
330 the original scheme. The mean measured ET of the oasis station in 4 days is 4.8 mm.
331 The original scheme ET is 6.9 mm, and the improved scheme ET is 4.2 mm using the
332 TM/MODIS hybrid strategy. The MABE of ET decreases from 2.1 mm (original
333 scheme) to 0.6 mm (improved scheme), and the MARBE declines from 44% to 13%
334 accordingly. And in all schemes desert ET is mainly distributed in the vicinity of 0
335 mm.

336 All simulated ETs show that the highest ET value is found in farmland, except
337 for water bodies considering the presence of irrigation water, followed by farmland,
338 forest, wetland, and residential areas, the lowest values appeared over bare rock land.

339 Different combination strategies show that the TM strategy has the highest
340 resolution, and the MODIS strategy has the lowest; however, the TM image cycle is
341 16 days, whereas the MODIS image cycle is 1 day. The TM/MODIS hybrid strategy
342 combines the advantages of the spatiotemporal resolution of both strategies.

343 During the growing season, the surface albedo and the height of the vegetation
344 change daily. The calculation of ET is influenced by the satellite image quality, and
345 the clouding images cannot be used. Thus, simulating continuously is difficult,
346 especially using the TM/MODIS hybrid scheme. An approach that combines the ET
347 estimates obtained from the TM with the drainage lysimeters has been verified useful



348 in computing cumulative ET (Morse et al., 2000).

349 In the growing season, the water of Heihe River is mainly consumed in the
 350 middle reaches. Therefore, accurate calculation of water consumption of various land
 351 types in the middle reaches, especially agricultural production, can effectively use
 352 limited water resources, leave more water to the downstream, and make the
 353 downstream ecological environment better.

354 Appendix

355 MODIS09GA

356 Broadband shortwave surface albedo is calculated from the normalized reflection
 357 values of Channels 1, 2, 3, 4, 5, and 7 using the following equation (Liang et al.,
 358 2003):

$$359 \rho_g = 0.160 \alpha_1 + 0.291 \alpha_2 + 0.243 \alpha_3 + 0.116 \alpha_4 + 0.112 \alpha_5 + 0.081 \alpha_7 - 0.0015. \quad (1)$$

360 TM5

$$361 \rho_g = \frac{\alpha_{toa} - \alpha_{path_radiance}}{\tau_{sw}^2}, \quad (2)$$

362 where α_{toa} is the atmospheric top reflectivity and can be obtained from the weighted
 363 average of the reflectivity of the shortwave bands (1, 2, 3, 4, 5, and 7 bands of TM) in
 364 remote sensing images. $\alpha_{path_radiance}$ is the albedo path radiance, and τ_{sw}^2 is the
 365 two-way transmittance (Morse et al., 2000).



366 **Data availability.** Meteorological data are available from the authors by request.

367 DEM and vector map of the boundary were provided by Digital Heihe
 368 (<http://heihe.westgis.ac.cn>). Remote sensing data are available from
 369 <https://ladsweb.modaps.eosdis.nasa.gov/> and
 370 <http://www.gscloud.cn/sources/?cdataid=263&pdataid=10>

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 372 technical guidance and data support, Doc. Shi gives help in daily GSR calculation.

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Figures

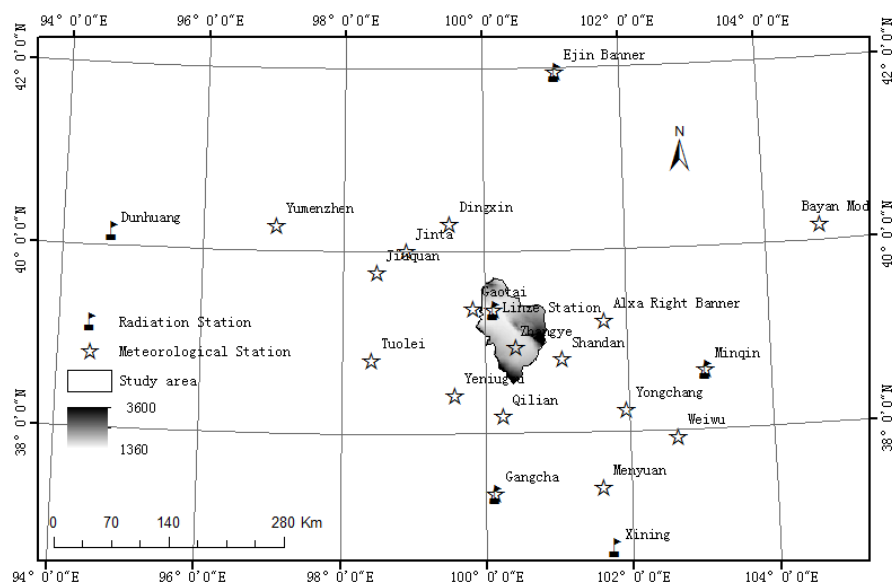


Fig.1. Distribution of meteorological stations and DEM of the middle reach of the Heihe River (unit: m)

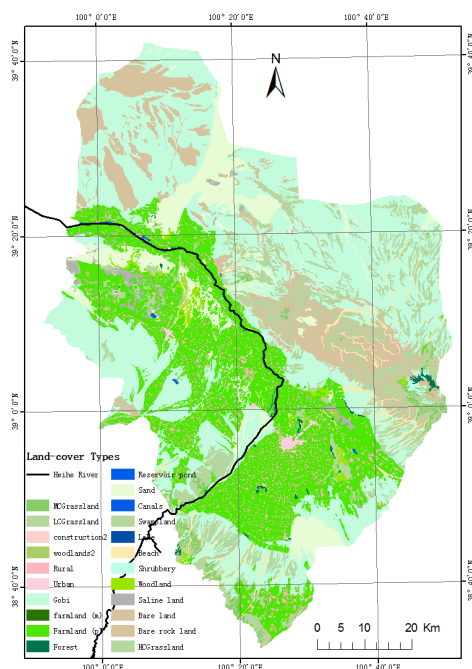


Fig.2. Land-cover map of the study area

(High-coverage grassland: HCGrassland, Moderate-coverage grassland: MCGrassland,
 Low-coverage grassland: LCGrassland, Sand desert: Sand, Gobi desert: Gobi, Other
 woodlands:woodlands2, Other construction:construction2, Rural residential: Rural, farmland
 (mountain):Farmland (m), Farmland (plain):Farmland(P))

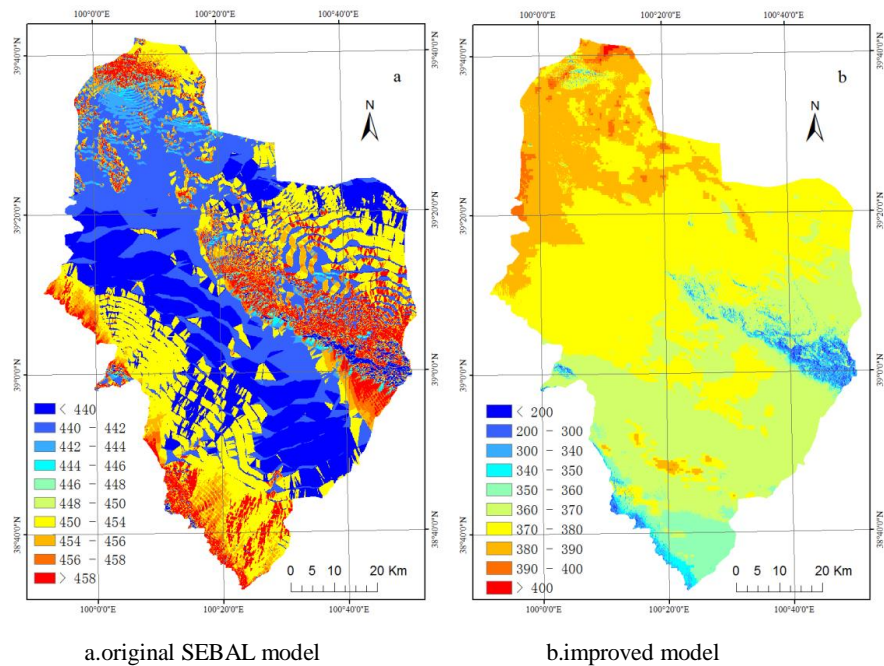
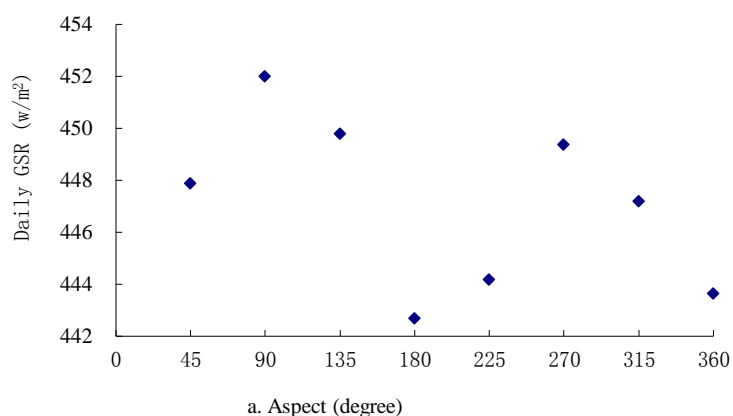
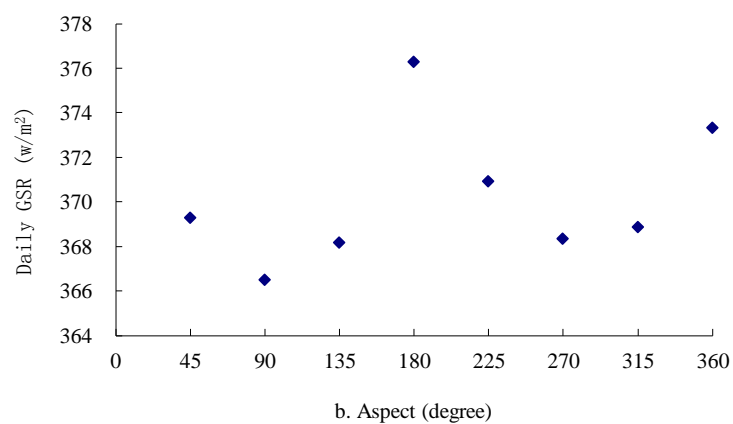


Fig.3. Spatial distribution of daily GSR in the middle Heihe River Basin on 24 June, 2009 (Unit: W/m²)



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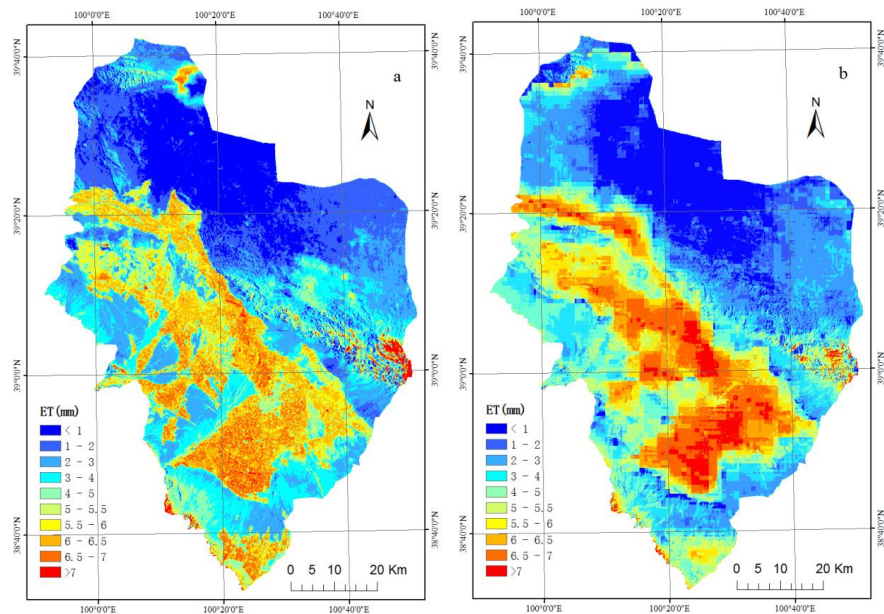
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Fig.4. Relation between daily GSR and aspect in the middle Heihe River Basin on 24 June, 2009 (Unit: mm)
 (a. original SEBAL model, b. improved model)

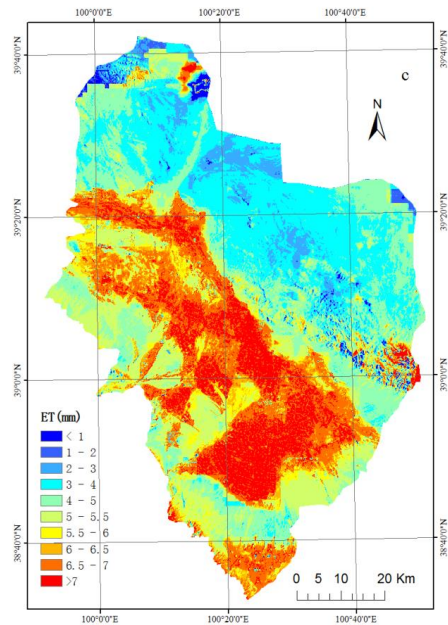


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a. TM strategy

b. MODIS strategy



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c. TM/MODIS hybrid strategy

Fig.5. Spatial distribution of ET retrieved using original scheme on 24 June, 2009 (Unit: mm)

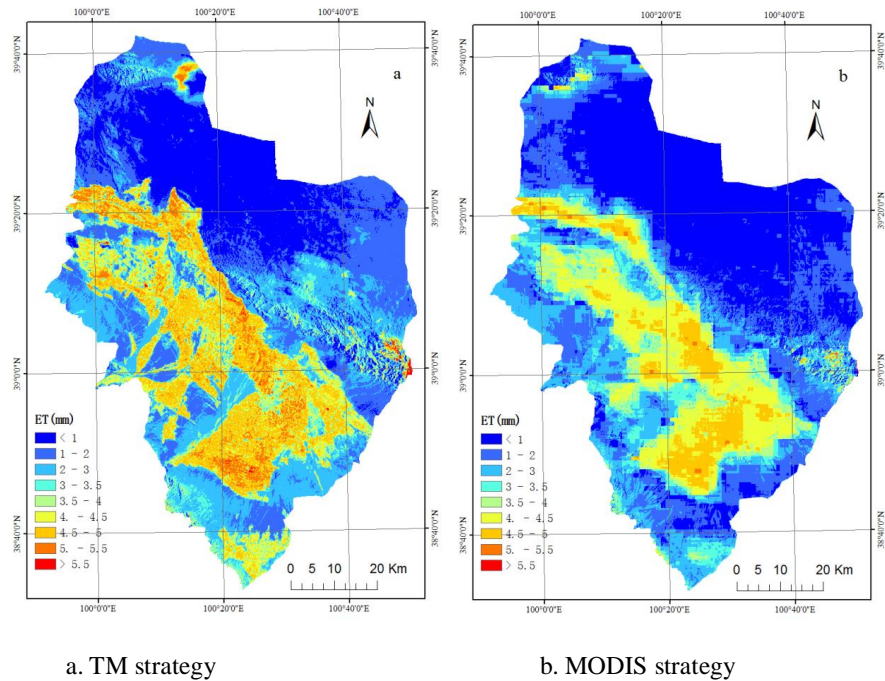


Fig.6. Spatial distribution of ET retrieved using improved scheme on 24 June, 2009 (Unit: mm)



502 **Tables**

503 Table 1

504 Daily GSR of middle Heihe River Basin on June 21–24, 2009 (Unit: W/m²)

Date	measure daily GSR	simulated daily GSR		MABE		MARBE (%)	
		improved	original	improved	original	improved	original
6.21	354.2	346.1	484	8.1	129.8	2.3	36.6
6.22	360	355.3	484.1	4.7	124.1	1.3	34.5
6.23	377.3	375	484.3	2.3	107	0.6	28.4
6.24	356.5	377.3	484.4	20.8	127.9	5.8	35.9
Mean	362	363.4	484.2	9	122.2	2.5	33.9

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537 Table 2

538 ET of three combination schemes on 24 June, 2009(Unit: mm)

Classification	measured ET	original scheme	improved scheme
TM strategy	4.8	6.3	5.0
MODIS strategy	4.8	6.3	4.2
TM/MODIS Hybrid strategy	4.8	7.0	4.8

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575 Table 3
 576 Errors of ETs of three combination strategies on 24 June, 2009(Unit: mm)

simulation strategy	simulation scheme	MABE	MABRE (%)
TM strategy	original scheme	1.5	31
	improved scheme	0.2	4
MODIS strategy	original scheme	1.5	31
	improved scheme	0.6	13
TM/MODIS strategy	original scheme	2.2	46
	improved scheme	0	0

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608 Table 4
 609 Daily ET of middle Heihe River Basin on June 21–24, 2009 (Unit: mm)

Date	Measure ET	simulated ET		MABE		MARBE (%)	
		improved	original	improved	original	improved	original
6.21	4.9	3.8	6.6	1.1	1.7	22	35
6.22	5.1	3.9	6.5	1.2	1.4	24	27
6.23	4.5	4.4	6.6	0.1	2.1	2	47
6.24	4.8	4.8	7	0	2.2	0	46
Mean	4.8	4.2	6.7	0.6	1.85	12	39

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640 Table 5

641 ET of each land-cover types

Underlying Surface	Type	Area (km ²)	improved scheme Mean ET (mm)	original scheme Mean ET (mm)
Forest land		16.43	3.5	5.6
Shrubbery		15.73	2.1	3.6
Woodland		56.99	3.4	5
Other woodlands		7.89	3.5	5.3
High-coverage grassland		36.43	3	4.9
Moderate-coverage grassland		76.43	2.5	4.2
Low-coverage grassland		731.89	1.9	3
Canals		20.24	4.5	6.6
Lake		0.16	3.6	5.4
Reservoir pond		4.2	4	5.6
Beach		106.18	2.1	3.2
Urban Land		15.33	3.8	5.5
Rural residential land		164.1	4	5.9
Other construction lands		4.51	3.8	5.6
Sand desert		472.64	1.4	2.3
Gobi desert		2321.37	1.4	2.1
Saline land		49.2	2.4	3.7
Swampland		4.48	4.2	6
Bare land		78.74	1.1	1.7
Bare rock land		558.52	1.5	2.3
farmland (mountain area)		0.13	3.7	5.9
Farmland (plain area)		1515.16	4.3	6.3

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