

# ***Interactive comment on “Turbulence in the stratified boundary layer under ice: observations from Lake Baikal and a new similarity model” by Georgiy Kirillin et al.***

**Anonymous Referee #2**

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## General Comments

The authors present compelling evidence – both theoretical and observational – for a new boundary layer model of flow beneath the ice cover of a freshwater lake. The approach taken is relevant for many large lakes with significant boundary currents, and in some cases may be important for sea ice as well: as such this will be of interest for readers of HESS. Nevertheless the manuscript would benefit from some clarification as outlined below.

## Scaling Arguments

The theoretical development (§2) is relatively straightforward but could be more clear.

C1

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It seems a bit unfortunate that both  $Q_{cw}$  and  $Q_{iw}$  are used for the water-ice heat flux. I expect  $Q_{cw}$  is used to emphasize that the flux is purely conductive (eqn. 2). Later it is shown that the flux is more than conductive, as exemplified by the turbulent diffusivity  $K$  in the equation between eqn. 19 and eqn. 20. It would be more clear to use  $Q_{iw}$  throughout, and argue that the new scaling shows that the diffusivity in eqn. 2 should be turbulent and not molecular. Thus the water-ice heat flux will be larger, and the ice at station S1 is thinner. Some additional minor points for this section: It would be nice to see a reference for eqn. 8. If it is Mironov et al. then state this before you write the equation. Also,  $\nu$  is the kinematic viscosity – not the viscosity though it is clear what you meant (pg. 7, line 7).

## Results

Fig. 3 has no time/date indicated. Are these data averaged over some period or a single snapshot? Also, there are important differences between Fig. 3a (station S1- strong flows) and Fig. 3b (S2- weaker flows) that are not discussed. It is pointed out the S1 shows 2 stratified layers above the mixed layer below 10 m depth – but S2 clearly shows 3 layers, evidently related to some process near the ice-water interface. Both are different from the canonical no-flow situation, exemplified in Fig. 13 (gray curve) which has 1 stratified layer. So as flows are increased from zero to strong the number of stratified layers seems to increase from 1 to 3 and then fall to 2. Can you explain this? Also, one assumes the total heat content in Fig. 3b (thicker ice) is larger than in Fig. 3a (thinner ice) – it might be of interest to compute this, comparing to the latent heat of the ice cover difference.

Other minor points:

Pg. 10, L.2-6: This paragraph appears to describe S1, though some points are relevant to both stations. You should clarify. Pg. 15, L. 18: Fig. 9b should be Fig. 9a. Likewise, on L.23 Fig. 9a should be Fig. 9b. Pg. 21, L.33: Eq. 4.5 should be Eq. 20.

Interactive  
comment

