

Interactive comment on “A coupled atmospheric-hydrologic modeling system with variable grid sizes for rainfall-runoff simulation in semi-humid and semi-arid watersheds: How does the coupling scale affects the results?” by Jiyang Tian et al.

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We appreciate very much the referee’s insightful comments and helpful suggestions for our manuscript. Efforts have been made to address every point of the referee’s concerns. During the revision being carried on, we are also encouraged by the positive comments from the referee “. . .the paper is interesting and well written with some minor spelling mistakes. . .” Gramma mistakes and spelling errors will carefully be checked

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before the revision is finally submitted. With the help of the referee, we hope the revised manuscript can be found rigorously and sufficiently improved.

Main comments: Point 1: The authors are actually presenting a new distributed hydrological model (why?) rather than using a distributed model widely accepted by the hydrological community. why do we need a new model? Reply: Thanks for the reviewer’s question. The Hebei rainfall-runoff model is specially developed to describe the runoff generation mechanisms in the semi-humid and semi-dry area of Northern China, which has been successfully applied in Hebei Province for rainfall-runoff modeling and real-time flood forecasting. Due to the perennial water shortage and groundwater over-exploitation, both storage-excess and infiltration-excess is found with great seepage along the river channel during the storm season. The obvious advantage of the Hebei model is the consideration of both storage-excess and infiltration-excess mechanisms for rainfall-runoff generation. It is a well-known conceptual model in China, as popular as the Xin’anjinag model. The model is easily used, and can widely be applied to other semi-humid and semi-arid watersheds with complicated (both storage-excess and infiltration-excess) mechanisms for rainfall-runoff generation. The description for the storage-excess part in the Hebei model is the same as the Xin’anjiang model. On the other hand, to reflect the heterogeneity of the infiltration capacity across the catchment, a distribution curve is adopted and expressed as Eqn. (2). The Horton infiltration model is also applied to obtain the infiltration volume for the river channel seepage. When the calculation interval is one hour, the infiltration volume can be calculated by the Eqn. (16). In order to clarify this issue, the following sentences are added in Line 22-28, Page 6: “Due to the perennial water shortage and groundwater overexploitation, both storage-excess and infiltration-excess is found in the study area with great seepage along the river channel during the storm season. The obvious advantage of the Hebei model is the consideration of both storage-excess and infiltration-excess mechanisms for rainfall-runoff generation. The model is easily applied and can be used in other semi-humid and semi-arid watersheds. In the Hebei model, the description for the storage-excess part is the same as that in the Xin’anjiang model. On the other

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hand, the infiltration capacity across the watershed is described by a distribution curve described below, and the Horton model is applied to calculate the seepage along the river channel during the river routing.”

Point 2: In order to answer your research question, you need to demonstrate that your hydrological model works well for a number of storms and that the model has been properly calibrated. I think this area is a bit weak in the paper and needs further results. For instance, the authors calibrate the model with 7 storms, but they do not give any indication on how the calibration was achieved, what type of storms are used and so on. Usually to calibrate hydrological models, continuous rainfall-runoff time series are needed rather than individual events in order to account for initial conditions in the model such as soil moisture, catchment wetness, etc. In addition, the authors quoted a model calibration efficiency (NSE) of 0.686, but three out of four events used in the validation showed an efficiency higher than 0.75. Normally the performance of the model in the validation phase is worse than in the calibration, but this is not the case in your analysis. why? For the storms used for validation, how do you account for the initial catchment conditions? It is obvious that if the model starts completely dry, the results will be affected by this even if the model is calibrated properly. please expand on this.

Reply: We agree with the referee that the model calibration part needs to be strengthened in the manuscript. The SCE-UA (Shuffle Complex Evolution) method is used to calibrate the parameters of the Hebei model (Duan et al., 1994). Actually, we have very limited choices when selecting the calibration data. Considering the semi-humid and semi-dry conditions of the study area, the soil is relatively dry before the storm season, and there is not many storm events leading to significant peak discharges. In this case, 7 storm events in Fuping and 6 storm events in Zijinggguan are chosen to calibrate the model, and 2 from each sub-watersheds are used to validate the model. Detailed information (i.e., the cumulative rainfall amounts and the peak discharges) of the events are summarized in the tables below. Considering there are already many table in the manuscript, these tables are not shown. When calibrating the model, the calibration events are bounded together to calculate one NSE value as the objective

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function. In order to guarantee reasonable values for the initial model conditions, the 24-h storm event is not independently used, but with a continuous antecedent period of data with the length of 15-days before the start of the event. In this sense, the events used for calibrating the model is some kind of “continuous” time series data. As for the calibration and validation values, it should be clarify firstly that the NSE value of 0.686 we quoted in Section 3.3 is for model verification (validation), not for calibration. This is an average value resulted from the 2 validation events from each sub-watershed, indicating that the calibrated model is reliable for further applications in the study area. It is also noted by the referee that there are three out of the four events with a NSE value higher than 0.75 in Table 9. The four storm events used for testing the grid sizes are different from those used for calibration and validation, but as “further applications”. Moreover, the validation results are from the lumped Hebei model, whereas the results of the four storm events are grid-based averaged from different grid sizes (each grid establishing a lumped model using the calibrate parameters). Therefore, the NSE values of the four events (some are higher than 0.75) are not comparable to those of the validation events (an average of 0.686). The following sentences are added in Line 29, Page 10 and Line 1-7, Page 11 to supplement more details about the model calibration and validation: “The SCE-UA (Shuffle Complex Evolution) method (Duan et al., 1994) is used to calibrate the parameters and the calibrated values are shown in Table 6. Due to the limited observational data, 7 storm events in Fuping and 6 storm events in Zijinggguan are selected and used to calibrate the Hebei model, and another 2 from each sub-watersheds are used for model validation. In order to guarantee reasonable values for the initial model conditions, the storm events are not independently used, but with an antecedent period of data with the length of 15-days before the start of the event. The validation results show an average NSE value of up to 0.686, indicating the calibrated models are reliable for further applications. It should be noted that the four storm events in Section 2.2 are different from those used for calibration and validation.”

Reference: Duan, Q., Sorooshian, S., Gupta, V. K. Optimal use of the SCE-UA global optimization method for calibrating watershed models, *J. Hydrol.*, 158(3-4), 265-284,

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Point 3: In addition, it is unclear from all the equations used to describe the model, which are model parameters. maybe you can summarise all the model parameters in a table and include their range of values. what about model parameter calibration uncertainty? What ranges of model parameters did you use and why? It is well known that different parameter sets can produce a similar model performance (equifinality). You need to look at parameter uncertainty and maybe produce an ensemble of hydrographs rather than a deterministic one. what about the uncertainty in the observations (e.g. rain. gauges, flow stations, etc)? You did not mention any of this in the paper. Reply: Thanks for the reviewer's suggestion. In the revised manuscript, efforts are made to explain all parameters and the uncertainty. Firstly, a new table below (Table 6) is added to show the calibrated parameter values on Page 6. The ranges of the parameter values are determined based on the application experiences of the Hebei model, which has been used in Northern China for more than two decades. The importance of hydrological uncertainty analysis has been emphasized in recent years and there is a necessity to incorporate parameter uncertainty estimation wherever a hydrological model is used. However, the parameter uncertainty estimation needs considerable observational data (Hughes et al., 2010). As mentioned above, there are not enough historical storm-flood processes for us to use in the study area, which makes the estimation work rather difficult. In this study, the research focus is how the coupling scale affects the flow results from the atmospheric-hydrologic coupling system, and we believe the conclusions from the comparative analyses would be quite similar even if the ensemble simulations with perturbed model parameters were carried out. Nevertheless, the parameter uncertainty estimation and ensemble simulations are suggested in the future study when sufficient observational data are available. The following paragraphs are added to address this issue in Line 6-9, Page 14: "It should be mentioned that there is a necessity to incorporate parameter uncertainty analysis in this study. However, this will need a considerable set of the observational data (Hughes et al., 2010). Due to the lack of sufficient historical storm-flood processes, it is impossible

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to carry out such analyses. Nevertheless, parameter uncertainty estimations and ensemble simulations with perturbed parameters are suggested in the future study when sufficient observational data are available." Reference: Hughes, D. A., Kapangaziwiri, E., Sawunyama, T. Hydrological model uncertainty assessment in southern Africa, *J. Hydrol.*, 387, 221–232, doi: 10.1016/j.jhydrol.2010.04.010, 2010.

As for the uncertainty of the observations, the rainfall and flow data are estimated before the establishment of the coupled atmospheric-hydrologic system. Hourly rainfall data are obtained from the rain gauges in the two sub-watersheds. Although the density of the rain gauges is a little sparse, the accuracy of the observations has been verified by the radar data, which can be found in our previous study (Liu et al., 2018). The hydrologic station observes the flow data at discrete time intervals (at least once an hour), which becomes more intense during the flood season to guarantee more realistic flow responses. The following sentence is added in Line 8-10, Page 4: "Before the establishment of the coupled atmospheric-hydrologic system, quality controls of the observational rainfall-runoff data are carried out. Rainfall observations from the rain gauges are verified by the weather radar and interpolations are done to guarantee the continuity of the flow observations (Liu et al., 2018)." Reference: Liu, J., Tian, J., Yan, D., et al. Evaluation of Doppler radar and GTS data assimilation for NWP rainfall prediction of an extreme summer storm in northern China: from the hydrological perspective, *Hydrol. Earth Syst. Sci.*, 22, 4329–4348, doi: 10.5194/hess-22-4329-2018, 2018.

Point 4: The second point is about the use of different WRF microphysics parameterisations (MP). Table 4 shows that different MPs are used to simulate each storm event. How do you isolate the impact of the WRF MPs in your results? How do you justify that the difference in the results is due to the different grid resolutions rather than the fact that different MPs are used to simulate each event? What is the performance of WRF simulating those storms? Reply: Actually in this study, the impacts of different coupling scales are compared for each certain storm event with the same MPs. The

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comparison is not carried out among the four events with different MPs. In order to eliminate the modeling errors caused by choosing inappropriate WRF parameterisations, the most suitable physical parameterisations resulting the most realistic rainfall simulations are used for each of the four storm event (as shown in Table 5). The reason why these physical parameterisations are the best choices has been discussed in detail in our previous study (Tian et al., 2017a). The following sentence is added in Line 27-31 Page 5 to address this issue: "In order to eliminate the modeling errors caused by choosing inappropriate WRF parameterisations, the most suitable physical parameterisations resulting the best rainfall simulations (Tian et al., 2017a) are used for each of the four storm events, as shown in Table 5. It should be clarified that the comparison of different coupling scales is carried out for each of storm event under the same MPs, thus using different MPs for different events will not cause difficulties in analyzing the final results." Reference: Tian, J., Liu, J., Wang, J., et al. A spatio-temporal evaluation of the WRF physical parameterisations for numerical rainfall simulation in semi-humid and semi-arid catchments of Northern China, *Atmos. Res.*, 191, 141–155, doi: 10.1016/j.atmosres.2017.03.012, 2017a.

Other comments: Point 5: The abstract talks about the "Hebei model", but this model is not known in the literature and you have not introduced this model yet. Reply: The introduction of the Hebei model is added in the abstract Line 14-16 Page 1: "The Hebei model is a conceptual rainfall-runoff model designed to describe a mixed runoff generation mechanism, including both storage-excess and infiltration-excess, in the semi-humid and semi-dry area of Northern China."

Point 6: In page 3: "a higher error rate". Do you mean "larger errors"? the use of "error rate" might be confusing. "variation pattern" - again unclear what you mean here. Reply: Revised accordingly. The sentences in page 3 are revised as below: "...but the accuracy of the rainfall information decreases such that the final simulation results may be subject to larger errors. Therefore, finding the underlying law of how the performance of the coupled atmospheric-hydrologic system is impacted by the coupling

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scale is of great importance in enhancing the accuracy of rainfall-runoff simulation."

Point 7: "Hebei model" Is this model published? If so, you need to include a proper reference. If not, then you should describe the model in the methodology and do not use "Hebei model" until this has been described. Reply: The "Hebei model" is a published model. There have already been many literatures in Chinese, but quite rare in English. The authors recently have published some English work, which gives a detail description of the Hebei model. The following reference is added in Line 13, Page 3 and the sentence in Line 13-14 is also revised: "...the Hebei model (Tian et al., 2019), a conceptual model with mixed runoff generation mechanisms of both saturation-excess and infiltration-excess, is used to construct the gridded hydrologic model." Reference: Tian, J., Liu, J., Yan, D., et al. Ensemble flood forecasting based on a coupled atmospheric-hydrological modeling system with data assimilation, *Atmospheric Research*, 224, 127-137, doi: 10.1016/j.atmosres.2019.03.029, 2019.

Point 8: "1x1km...9x9km" Are these spatial scales within the WRF model domains or are you talking about the spatial scales of the hydrological model? Reply: "1x1 km, 3x3 km, and 9x9 km" are the coupling scales for the coupled atmospheric-hydrologic system. Therefore, they are not only the spatial scales of the WRF model outputs (three domains with the grid cell size being 1x1 km, 3x3 km, and 9x9 km), but also the spatial scales of the gridded Hebei model.

Point 9: Section 2. The description of the catchments should be placed before the description of the events otherwise how do you know which catchment outlet are you talking about here? Reply: Agreed and the order of section 2.1 and 2.2 is changed.

Point 10: Is Cv a standard metric to characterise the spatial and temporal distribution of precipitation processes? If so, you need to include a reference. I believe to characterise the spatial distribution of precipitation is better to use semi variograms, correlograms or by looking at the spatial correlation of the precipitation field. Likewise, with the temporal correlation. Could you please expand and justify why a simple metric

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like this was used? Cv here is highly dependent on the number of rain gauge stations available. Is the WRF rainfall field used to compute Cv? Reply: Cv is a standard metric to describe the dispersion of measures, thus is used in this study to describe the evenness of rainfall distribution. The main advantage of the statistic is that the evenness of rainfall distribution can easily be quantified in both time and space by following the proposed rules in Section 2.2. We have a series of publications regarding the WRF model and the simulation of storm events with different spatio-temporal evenness, where the same Cv statistic and calculation rules are adopted (Tian et al., 2017a, b). We have also found other studies using the same statistic to describe the rainfall distributions (Yue et al., 2014; Jha et al., 2015). It is true that the statistics may depend on the number of rain gauges (actually any statistic may involve this uncertainty), but this is the best we can do. Considering the WRF simulations are not the “ground truth”, and in this study observations from the rain gauges are used to evaluate the WRF simulations, hence the rainfall evenness are obtained from the rain gauges rather than the WRF simulations. The following references are added in Section 2.2 when introducing the use of Cv as an evaluation of the rainfall distribution evenness. References: Tian, J., Liu, J., Wang, J., et al. A spatio-temporal evaluation of the WRF physical parameterisations for numerical rainfall simulation in semi-humid and semi-arid catchments of Northern China, *Atmos. Res.*, 191, 141–155, doi: 10.1016/j.atmosres.2017.03.012, 2017a. Tian, J., Liu, J., Yan, D., et al. Numerical rainfall simulation with different spatial and temporal evenness by using a WRF multiphysics ensemble, *Nat. Hazards Earth Syst. Sci.*, 17, 563-579, doi: 10.5194/nhess-17-563-2017, 2017b. Yue, B. J., Shi, Z. H., Fang, N. F. Evaluation of rainfall erosivity and its temporal variation in the Yanhe River catchment of the Chinese Loess Plateau, *Nat. Hazards*, 74, 585-602, doi: 10.1007/s11069-014-1199-z, 2014. Jha, S. K., Zhao, H., Woldemeskel, F. M., et al. Network theory and spatial rainfall connections: An interpretation, *J Hydrol.*, 527, 13-19, doi: 10.1016/j.jhydrol.2015.04.035, 2015.

Point 11: table 2 - I think you need to explain what values of Cv correspond to a highly variable event in space and time. how did you come up with the critical values of Cv

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(0.4 and 1 for spatial/temporal distribution respectively)? you need to justify these values. Reply: In this study, the spatial and temporal Cv of the historical storms from 1985 to 2018 is calculated to analyse the characteristics of the rainfall evenness. A threshold of 5% is used to separate even and uneven storms. It is found that the storm events with a spatial Cv < 0.4 or with a temporal Cv < 1.0 account for 5% of the total storm events from 1985 to 2018. The methodology is also adopted in our previous publication (Tian et al., 2017a). However, the critical values of 0.4 and 1.0 are based on statistical analyses of historical storm events, thus are not transferable to other areas with different meteorological conditions. In order to avoid misunderstanding, the description of the critical values are removed in the revised manuscript. Instead, the spatial and temporal evenness of rainfall distribution is ranked among different storm events. The following sentences can be found in Line 22-24 Page 4: “The smaller is the value of Cv, the more even is the rainfall distribution in space or time. According to Table 3, the ranking of the distribution evenness of rainfall in space is event 2 > event 1 > event 4 > event 3 and that in time is event 1 > event 2 > event 4 > event 3.” Reference: Tian, J., Liu, J., Wang, J., et al. A spatio-temporal evaluation of the WRF physical parameterisations for numerical rainfall simulation in semi-humid and semi-arid catchments of Northern China, *Atmos. Res.*, 191, 141–155, doi: 10.1016/j.atmosres.2017.03.012, 2017a.

Point 12: Section 2.2 can you provide more information of these catchments (e.g. catchment descriptors)? e.g. apart from catchment area, mean annual rainfall, catchment slope, mean flow, predominant land use and soil types, type of geology, percentage of urban area, etc. you can summarise all this info in a table. this is important to understand the catchment response to precipitation. Reply: Thanks for the reviewer's suggestion. Section 2.2, which changed as section 2.1 in the revised manuscript, has been revised as follows and the information table is also added to provide more details about the study area: “In this study, two mountainous sub-watersheds of the Daqing River basin (Fuping of the south branch and Zijingguan of the north branch) were chosen as the study area (Fig. 2 and Fig. 3). The Fuping sub-watershed has

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a total area of 2,219 km² and is located in the upper reaches of the Zhishahe River, a south branch of the Daqing River. The Zijinguan sub-watershed has a total area of 1,760 km² and is located in the upper reaches of the Juma River, a north branch of the Daqing River. The two sub-watersheds are the most concentrated and typical cinnamon regions. The land use mainly includes grassland, farmland and forestland. The soil erosion is severe. Due to the dry soil conditions and the groundwater over-exploitation, the river has great seepage during the storm season. More information about the two sub-watersheds is shown in Table 1. The study area embodies the representative rainfall-runoff characteristics of the sub-humid and sub-arid area in Northern China. Rainfall in Northern China is characterized by summer storms with short durations and large intensities, which are likely to result in severe flood disasters especially in mountainous areas like Fuping and Zijinguan.”

Point 13: Section 3.1.1. “in consistence” consistent? The grid resolution of the coarser domain is 9km, but the NCEP analysis is about 100km. How does the WRF model handles this discrepancy with the initial and boundary conditions between the outer domain and the analysis? What’s the temporal resolution of the analysis? Reply: “consistent” is right and the sentence is revised accordingly. The WRF model is a next-generation mesoscale numerical weather prediction system. It can run on a variety of computing platforms and handle a broad range of applications across scales ranging from tens of meters to thousands of kilometers by dynamical downscaling. Actually when we run the WRF model, four nested domains are initially adopted, with the outermost domain being 27 km in order to deal with the discrepancy with the NCEP data. In order to avoid misunderstanding, the following sentences are added in Line 10-13, Page 5. In addition, the temporal resolution of the NCEP analysis data is 1 h. “The NCEP/NCAR final operational global analysis (FNL) data with spatial resolution of 1°×1° and temporal resolution of 1 h were used to provide the lateral and boundary conditions of the WRF model (Wang et al., 2013a,b). In order to eliminate the discrepancy of the initial and boundary conditions with the driven data, another outermost domain was set beyond the WRF three nested domains to downscale the FNL data to a spatial resolution of 27

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km.”

Point 14: Section 3.1.2. Need to define all acronyms for the physical parameterisations. Reply: Revised accordingly as below: “. . .which include two microphysics parameterisations, i.e, Purdue-Lin (Lin) (Lin et al., 1983) and WRF Single-Moment 6 (WSM6) (Hong et al., 2006), two cumulus parameterisations, i.e., Kain-Fritsch (KF) (Kain, 2004) and Grell-Devenyi (GD) (Grell and Freitas, 2014), and two PBL (planetary boundary layer) parameterisations, i.e., Mellor-Yamada-Janjic (MYJ) (Hong et al., 2006) and Yonsei University (YSU) (Janjić, 1994). Besides, Rapid Radiative Transfer Model (RRTM) and Dudhia (Evans et al., 2012) usually cooperated well as the long/short wave radiation parameterisations and Noah was chosen to be the land surface model (Chen et al., 2014).”

Point 15: Section 3.2.2 “. . . has widely been applied in Hebei Province . . .” add references. Reply: The sentence is revised as below with the following reference added: “. . .has widely been applied in Hebei Province by considering both infiltration-excess and saturation-excess mechanisms of the runoff generation (Tian et al., 2019).” Reference: Tian, J., Liu, J., Yan, D., et al. Ensemble flood forecasting based on a coupled atmospheric-hydrological modeling system with data assimilation, *Atmospheric Research*, 224, 127-137, doi: 10.1016/j.atmosres.2019.03.029, 2019.

Point 16: How do you account for different soil types, land use, etc? These will have important implications in terms of runoff production. Do you use ET (evapotranspiration) or only E (evaporation)? do you have forest in any of the basins? Expand on the calibration and validation of the hydrological model, including both, lumped and distributed models. Reply: Thanks for the referee’s remind. The type of soil in the study area is typical cinnamon soil with considerable soil erosions. The land use mainly includes grassland, farmland and forestland. Due to the perennial water shortage and groundwater overexploitation, the river has great seepage during the storm season. The Hebei model is specially developed as a very simple conceptual model for rainfall-runoff modeling in this region. As mentioned in section 3.2, the calculation of the lumped Hebei

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model has no relation with the soil type or the land use, and the grid-based model only considers the spatial distribution of the rainfall, the soil water storage capacity and the soil infiltration capacity. As mentioned by Eqn. (4), E (evaporation) is used in the models instead of ET. The details of the study area are supplemented in a new table (Table 1). Please refer to our reply to Point 12.

Point 17: fig 11. Unclear if these results are for catchment A or B. Reply: As shown in Table 1, Event 1, 2 and 3 is from Fuping and Event 4 from Zijingguan. In order to make this clear, the caption of Fig. 11 is revised as follows: "Figure 11. Flood process simulations of the coupled atmospheric-hydrologic systems for the four storm events: (a) Event 1 in Fuping; (b) Event 2 in Fuping; (c) Event 3 in Fuping; (d) Event 4 in Zijingguan."

Point 18: Table 5. how do you isolate the impact of the different WRF microphysics parameterisations in your results? Reply: In this study, the impacts of different coupling scales are compared for each single storm event with the same parameterisations. The comparison is actually not carried out among the four events with different parameterisations. Please see our reply to Point 4 for more details.

Interactive comment on Hydrol. Earth Syst. Sci. Discuss., <https://doi.org/10.5194/hess-2019-587>, 2020.

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Two tables for the reply of Point 2

The storm events used to calibrate the model			
Date	Sub-watershed	24-h rainfall accumulations (mm)	Peak discharges (m ³ /s)
02/07/1997	Zijingguan	51.31	163
05/07/1998	Zijingguan	68.37	129
29/06/2006	Zijingguan	48.55	100
03/07/2007	Zijingguan	36.96	25
01/09/2012	Zijingguan	38.58	20
11/08/2013	Zijingguan	40.74	33
06/07/2000	Fuping	60.86	330
24/07/2001	Fuping	56.03	105
13/08/2004	Fuping	32.85	102
14/08/2006	Fuping	24.64	41
30/08/2010	Fuping	20.43	33
29/06/2012	Fuping	25.42	48
29/06/2013	Fuping	19.68	38

The storm events used to verify the model			
Date	Sub-watershed	24-h rainfall accumulations (mm)	Peak discharges (m ³ /s)
06/06/1997	Zijingguan	57.44	128
30/06/2005	Zijingguan	46.22	98
14/08/2005	Fuping	34.92	103
24/07/2007	Fuping	42.16	106

Fig. 1.

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Table 6 Calibrated parameters in the Hebei model

Parameters	Units	Suggested values	Descriptions	parameter values for Fuping	parameter values for Zijingsuan
u	none	0-0.1	Decreasing speed of the infiltration rate with the increase of the soil moisture	0.02	0.02
f_c	mm/h	1-2	stable infiltration rate	1.5	1.5
n	none	0.3-0.8	exponent of the distribution curve for the infiltration capacity	0.53	0.50
b	none	0.3-0.5	exponent of the distribution curve for the moisture storage capacity	0.49	0.50
WAM	mm	80-300	maximal moisture storage capacity of a certain grid cell	240	238
f_m	mm/h	20-200	maximum infiltration capacity of a certain grid cell	120	120
A	(m ³ /s) ⁻¹ =s	0-1	confluence parameter	0.85	0.85

Fig. 2.

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Table 1. The characteristics of the two sub-watersheds

Descriptors	Fuping catchment	Zijingsuan catchment
catchment area	2,219 km ²	1,760 km ²
mean annual rainfall	490 mm	650 mm
longitudinal river slope	5.7%	5.5%
Annual average runoff	2.85 × 10 ⁹ m ³	2.81 × 10 ⁹ m ³
predominant land use	grassland, farmland and forestland	grassland, farmland and forestland
soil type	cinnamon soil	cinnamon soil
type of geology	granitic gneiss	granitic gneiss
percentage of residential area	0.63%	0.52%

Fig. 3.

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