



1 **Climate change and snow cover trends in Iceland**

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6

7 Abstract

8 We studied ~~the~~ trends in climate ~~change~~ and snow cover in Iceland. Climate was classified based on
9 the Köppen-Geiger (KG) classification system for 1950-2100 using the ensemble average of + NASA-
10 NEX downscaled CMIP5 projections for RCP 4.5. Snow Cover Frequency (SCF) was calculated based on
11 ~~in~~-MODIS10A1 snow product in days/year and SCF trends were calculated for 2001-2016. Trends in
12 climate and snow cover changes were evaluated in 4 different elevation bands: Coastline (0-100
13 m.a.s.l.), Lowland (100-500 m.a.s.l.), Highland (500-1000 m.a.s.l.) and Mountains/Glaciers (1000+
14 m.a.s.l.). The results showed that in all elevations zones warmer climate classes have been replacing
15 colder climates and polar tundra since 1950's, based on climate projections we expect these trends to
16 continue throughout the present century. We observed that in large areas of the country a significant
17 increase in SCF had occurred during the period 2001-2016. These changes were most pronounced in
18 the highlands where SCF had increased by 3.5 days/year on average. The only locations where we
19 observed decreasing SCF was around the termini of the country's outlet glaciers. The results suggest
20 that by the end of the present century polar tundra climate (ET) will have decreased from 20% to 5%
21 coverage ~~and~~ cold climate with warm summers (Dfb), will extend around the island and spread into
22 the highlands.





23 1 Introduction

24 We studied recent climate change in Iceland and compared these changes to changes in snow cover
25 in the country. Past and predicted future climate characteristics were assessed according to the
26 Köppen-Geiger (KG) classification method using Global Circulation Model (GCM) projections. The
27 climate trends were combined with MODIS satellite observations of snow cover to investigate the
28 impact of climate change on snow cover in Iceland. Understanding the past and future impacts of
29 climate change on snow resources is essential for water management in cold regions, not least in
30 Iceland. The specific research questions were: Has the Icelandic climate been undergoing recent
31 changes, and if so, have these changes coincided with changes in local snow resources? Furthermore,
32 what changes can be expected to these regimes in the future?

33 Icelandic climate is highly controlled by the sea conditions in the surrounding North Atlantic
34 ocean (e.g. Massé et al., 2008) and seasonal mass balance trends of Icelandic glaciers have been shown
35 to correlate with large-scale oceanic circulations in the North Atlantic (e.g. Eythorsson, 2018). Air
36 temperatures in Iceland have been increasing over all long-term records. However, a marked cooling
37 trend was observed between the 1940-1980, this cooling trend has been ascribed to changes in large
38 scale atmospheric circulation (Hanna et al., 2004). Since 1980 air temperatures in Iceland have
39 increased on average by 0.5 °C per decade and annual precipitation has increased by 200-300
40 mm/year (Bjornsson et al., 2018). As a result of this warming, Icelandic glaciers have been retreating
41 in recent decades, and are expected to mostly disappear completely over the next two centuries if
42 current climate trends persist (Adalgeirsdottir et al., 2006; Helgi Bjornsson & Pálsson, 2008;
43 Jóhannesson et al., 2004). Since the end of the 19th century glacial melt in Iceland has been estimated
44 to have contributed about 3 mm to global sea level rise (Bjornsson et al., 2013). The observed climate
45 trends have been predicted to impact local snow resources with spring thaws beginning earlier and
46 autumn snow cover to occur later (Jóhannesson et al., 2007), however new research has suggested
47 that the duration of snow cover has increased during the period 2000-2018 (Gunnarsson et al., 2019).

48 Future climate conditions can be estimated using GCM projections (Stocker et al., 2013). To
49 represent future climate we used the ensemble average of the 21 GCM projections in the CMIP5
50 collection for the Representative Concentration Pathway (rcp) 4.5 (Taylor et al., 2012). GCM results
51 have been used to quantify climate change by classifying regional climate and analyzing changes in
52 classifications over time (Chen & Chen, 2013). We used the Köppen-Geiger classification system
53 (Köppen, 1918) which has often been used to study climate change (e.g. Engelbrecht & Engelbrecht,
54 2016; Rube et al., 2017) and to validate GCM results (Lohmann et al., 1993).

55 When estimating the availability of snow resources from remote sensing an important
56 indicator is Snow Cover Frequency (SCF), the number of snow-covered days per year. SCF has been
57 related to growing season length and species habitability (e.g. Callaghan et al., 2011) and is an
58 important parameter in surface energy balance (Cohen, 1994). SCF is an important variable to include
59 in geophysical simulations, like hydrological modelling in cold regions (e.g. Guan et al., 2013; Nahaet
60 al., 2016). Historical changes in the SCF have been used to assess the progression and impacts of
61 climate change on the snow regime (Brown & Mote, 2009). Remote sensing by satellite sensors is a
62 commonly applied way of measuring snow cover (Frei et al., 2003). The launch of the MODIS
63 instruments on board the TERRA and AQUA satellites has improved snow cover information for
64 hydrological model validation (Dietz et al., 2012). Improved understanding and modelling of future
65 changes in snow conditions in cold areas is important, as these are changes that are expected to
66 impact local communities and ecosystems as well as changing the challenges and opportunities for
67 exploiting natural resources in the region (Eliasson et al., 2017).

68 Our hypothesis in this study is that recent warming in Iceland has resulted in decreasing snow
69 cover. We expected the trend to be most apparent at lower altitudes around the coastline but also
70 persistent at higher altitudes. Our results, however, showed that the SCF had increased across the
71 entire country. We observed that the frequency of snow-covered ground has increased and the
72 increase is more prominent at high elevation than at lower elevated areas. The only locations where



73 we consistently observed a significant decreasing SCF trend was at the terminus of the country's
74 glaciers.

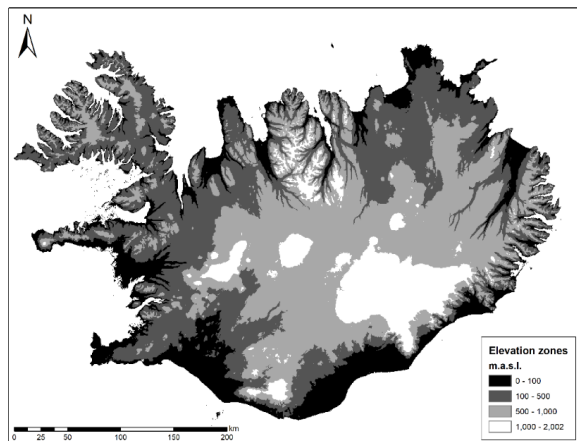
75 2 Methods

76 2.1 Tools and Datasets

77 Google Earth Engine (GEE) (Gorelick et al., 2016) was used to access data and for spatial analysis.
78 Statistical analysis was performed using GEE and the SciPy toolbox (Oliphant, 2007). ArcGis 10.6.1 was
79 used to produce maps showing the results. The MODIS10A1 MODIS/TERRA daily snow cover product
80 (Hall et al. 2006) was used to estimate changes in snow cover. The NASA NEX GDDP dataset (Thrasher
81 et al., 2006) was used to estimate historical and predict future climate conditions. For calculations
82 involving elevation a 20x20m DEM from the National Land Survey of Iceland was used.

83 2.2 Elevation Zones

84 Icelandic climate and ecology vary significantly depending on elevation. Lower elevations are
85 characterized by milder climates and denser vegetation, whereas the central highlands experience
86 much harsher climate and are very sparsely vegetated, the highest elevations are glacierized. We
87 analyzed the climate and snow regimes in various elevation bands. Iceland was divided into four
88 elevation bands: *Coastline* (0-100 m a.s.l.), *Lowland* (100-500 m a.s.l.), *Highland* (500-1000 m. a.s.l.)
89 and *Glaciers/Mountains* (1000 + m a.s.l.). A high resolution (20x20m) digital elevation model (DEM)
90 from the National Land Survey of Iceland was uploaded to GEE and used to calculate the elevation
91 bands. Figure 1 shows the areas of each of the defined elevation zones in Iceland.



92

93

Figure 1. Elevation zones in Iceland used for analysis of climate and snow changes.

94 2.3 SCF trend

95 Annual Snow Cover Frequency (SCF) was calculated with 500m x 500m resolution for Iceland based
96 on the MOD10A1.005 MODIS/TERRA snow cover daily product (Hall et al., 2006). The annual SCF was
97 estimated using the methods described in Eythorsson et al., (2019). The MOD10A1.005 dataset was
98 remapped to provide a binary classification for valid observations. Observations with zenith angles >
99 25° were excluded to decrease the *panoramic bow tie effect* which is a panoramic distortion known
100 to cause systematic errors in snow mapping (Souri & Azizi, 2013). Invalid observations due to cloud



101 cover or polar night, for example, were masked by giving them a null value. The number of days a pixel
102 is covered with snow was counted and divided by the number of valid observations of that pixel, per
103 year. On average we observed an average of 60 valid observations/year per pixel. We did not perform
104 gap-filling of the MODIS snow cover product to avoid introducing another source of uncertainty. The
105 annual SCF was calculated for the period where MODIS observations are available (water years 2001-
106 2016). The trend of annual SCF values in each pixel over the period was estimated by linear regression
107 and Sen's estimator of slope methods. The statistical significance of the observed trend was assessed
108 using both the non-parametric Mann-Kendall trend test and Sen's estimator of slope methods. Maps
109 were produced showing the areas where significant changes in the SCF had occurred at the $\alpha = 0.01$
110 and $\alpha = 0.05$ confidence level.

111 2.3.1 SCF in areas of recent land surface changes.

112 The annual SCF values were extracted for three locations in Iceland where land surface
113 changes had physically impacted the local SCF during the MODIS period. The locations selected were:
114 (a) Holuhraun volcano, which erupted in the winter of 2014, where the SCF was expected to decrease
115 following the eruption; (b) the Háslón area, where a major storage reservoir was commissioned in
116 2007 and an ice-covered lake replaced a deep canyon leading to an expected increase in SCF; and (c)
117 Eystri Hagafellsjökull, where the glacier terminus has receded in recent years, leading us to expect a
118 sudden decrease in the SCF as perennial glacier coverage has given way to reveal the subsurface
119 beneath during the snow-free months of the year.

120 To determine whether the change that has occurred over the historical period (2001-2016) is
121 statistically significant a hypothesis test was developed. As the SCF data was not expected to be
122 normally distributed a Mann-Whitney-Wilcoxon (MWW) test on two sample means was performed.
123 To determine the time of occurrence of the change, historical time series were analyzed by selecting
124 a split in the time series and performing the hypothesis test on the resulting two time series, before
125 and after the split. The null hypothesis was that the means of the two series are the same: $H_0: \mu_1 = \mu_2$
126 and the alternative hypothesis was that the means are not the same: $H_1: \mu_1 \neq \mu_2$. Code was written in
127 python that tests the hypothesis for all possible splits in the historical time series and returns the split
128 with the lowest p-value. We argue that the data are reliable if the tests return the same period of
129 change in the SCF as the known land surface changes.

130 2.4 Köppen-Geiger climate classifications

131 Köppen-Geiger climate classifications were calculated for Iceland in a 0.2-degree horizontal resolution
132 using the methods described in Eythorsson et al., 2019. Climate classifications were assigned to each
133 pixel based on the classification criteria outlined by Kottek et al. (2006) and Peel et al. (2007), as
134 summarized in Appendix 1. The classification scheme contains five main classes, each with two levels
135 of subclasses, in total 30 climate classes. As example, an area that has the main class *D – Cold*, second
136 subclass *w – dry winter* and the third subclass *a – hot summer* would have the code *Dwa*.

137 We classified Icelandic climate for each year in the period 1950-2100. We used the ensemble
138 average of the NASA NEX dataset for both historical and predicted future climate conditions (Thrasher
139 et al., 2006). The dataset contains an ensemble of 21 Global Circulation Models (GCM's) used in the
140 CMIP5 model intercomparison project of the International Panel on Climate Change (IPCC) (Taylor et
141 al., 2012).

142 The dataset includes model runs for two Representative Concentration Pathways (RCP) - RCP
143 4.5 and RCP 8.5. We selected RCP 4.5 as a more conservative prediction of future climate change. RCP
144 4.5 is a stabilization scenario where total radiative forcing is stabilized before 2100 by employment of
145 a range of technologies and strategies for reducing greenhouse gas emissions, whereas RCP 8.5 is
146 characterized by increasing greenhouse gas emissions over time and is representative for scenarios in
147 the literature leading to high greenhouse gas concentration levels (van Vuuren et al., 2011). We



148 calculated the proportion of each climate class within each study area for each year and estimated
149 climate change as the change in KG classification in each grid cell.

150 2.5 Comparison of SCF and Climate trends

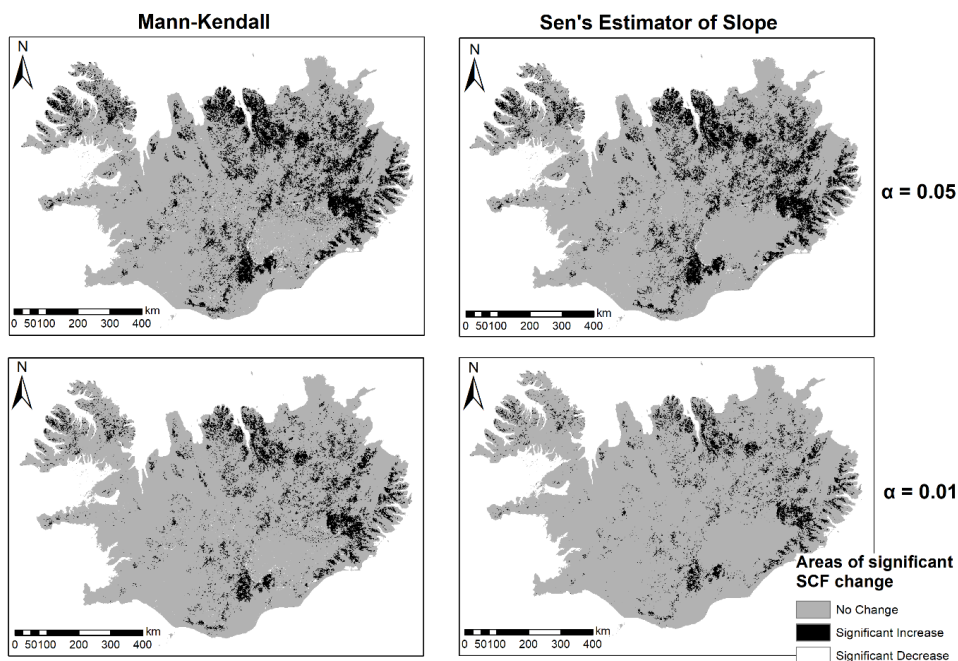
151 The evidence for statistical significance of the changes in the snow and climate regimes was
152 estimated by performing the non-parametric Mann-Kendall and Sen's estimator of slope methods on
153 the timeseries of the SCF and the proportional KG class coverage. Both of these tests have been used
154 often to assess the significance of the trends in hydro-meteorological time series (Drapela &
155 Drapelova, 2011; Gocic & Trajkovic, 2013). In the significance tests the null hypothesis is that there is
156 no monotonic trend present in the data $H_0: s = 0$, while the alternative hypothesis is that the data has
157 a monotonic trend $H_1: |s| > 0$. Both tests were performed at $\alpha = 0.01$ and $\alpha = 0.05$ significance levels.

158 The significance of the trend calculated by linear regression was assessed using the Mann-Kendall
159 trend test. We applied the test using the *stats.kendalltau()* function in SciPy and the
160 *ee.Reducer.kendallsCorrelation()* function in GEE. We applied Sen's slope test on the significance of
161 the trend using the *stats.mstats.theilslopes()* function in SciPy and the *ee.Reducer.sensSlope()* function
162 in GEE. We developed maps of areas showing evidence for significant changes in the SCF during the
163 period 2001-2016. We then calculated the ratio of area with significant SCF changes during the period
164 within each study area. We extracted the time series of KG climate class coverage within each study
165 area and assessed the significance of the trend of the two most common climate classes, using both
166 the Mann-Kendall and Sen's slope tests.

167 3 Results

168 3.1 SCF Trend Analysis

169 The statistical significance of the SCF trend in Iceland over the period 2001-2016 was analyzed with
170 both the non-parametric Mann-Kendall trend test and Sen's estimator of slope methods as described
171 in Section 2.5. Figure 3 shows the results from these analyses at two confidence levels ($\alpha = 0.05$ and
172 $\alpha = 0.01$). Black areas indicate a significant positive SCF trend and white areas a significant negative
173 SCF trend, while gray areas do not support rejection of the null hypothesis.



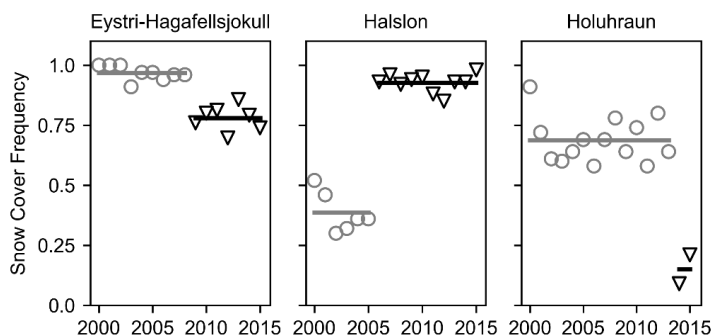
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Figure 2. Areas of significant change in the SCF identified using Mann-Kendall and Sen's Slope hypothesis tests at confidence levels $\alpha = 0.05$ and $\alpha = 0.01$

177 Figure 2 shows that large areas of Iceland experienced a significant change in the SCF during the period
178 2001-2016. The figure clearly shows that the SCF increased specifically in the eastern and northeastern
179 parts of the country. This is an interesting result as these are relatively arid areas as they are in the
180 rain shadow of the Vatnajökull Glacier. Increasing SCF trends were also observed on the north-facing
181 peninsulas, e.g. the Trollaskagi and Gjögraskagi peninsulas. A clear increase in SCF was also observed
182 southwest of the Vatnajökull Glacier. Few areas of decreasing SCF trends were observed, mostly at
183 the terminus of outlet glaciers which have receded in recent years. This is especially evident for
184 example at the terminus of outlet glaciers along the south-western edge of the Vatnajökull Glacier,
185 Tungnárjökull, Síðujökull and Skeiðarárjökull.

186 3.2 Verification of SCF trends

187 The SCF data were analyzed in areas where recent land surface changes were expected to have had
188 an impact on the snow regime during the MODIS record. Three locations were analyzed: Holuhraun
189 volcano, which erupted in 2014 and the hot lava has melted snow for two consecutive years;
190 Karahnjúkar canyon, where a hydropower reservoir was formed in 2006 by inundation of a deep
191 canyon and now has ice cover for most of the year; and Eystri Hagafellsjökull, an outlet glacier that
192 receded during the study period. One 500x500m pixel at each location was selected for analysis. Figure
193 1 shows the time series of the SCF in the three ground truth locations as well as mean values for the
194 periods before and after the expected SCF change.



195
 196

Figure 3. SCF data from verification locations of recent land surface changes (horizontal bars show mean values)

197 The time series for the SCF at each location was analyzed for evidence of statistically significant change
 198 during the MODIS record (2001-2016) by an MWW hypothesis test, as described in Section 2.3.1.

199 Table 1 shows the results of the hypothesis tests. The results show that for all three locations
 200 the null hypothesis was rejected. Hence, the SCF record captures physical land surface changes that
 201 occurred during the period and that the timing of these changes can be identified by the MWW test.
 202 We also note that, like the Háslón reservoir, all other hydropower reservoirs constructed during the
 203 MODIS era, Sporðöldulón, Ufsarlón and Kelduárlón, could be clearly identified from the SCF maps.

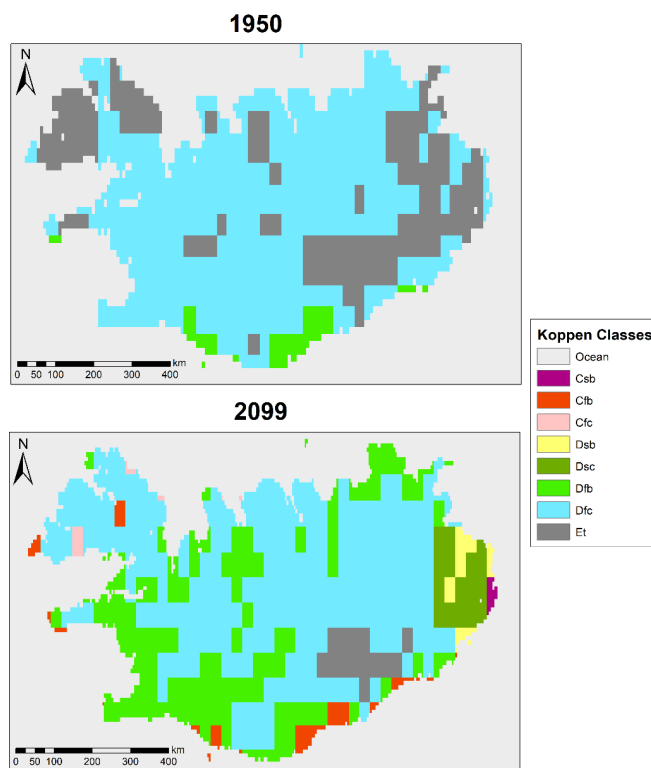
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Table 1. Results of SCF verification experiment in Iceland

Location:	Time of change	MWW p-value	SCF trend
Holuhraun	2014	0.017	Significant decrease
Háslón	2007	0.00025	Significant increase
Eystri Hagafellsjökull	2008	0.00031	Significant decrease

206 3.3 Köppen-Geiger classifications

207 Köppen-Geiger (KG) classifications were calculated for Iceland in 0.2x0.2-degree resolution. The four
 208 most common KG classes in Iceland in this period were: *ET* – Polar Tundra, *Dfc* – Cold climate with cold
 209 summers and no dry season, *Dfb* – Cold climate with warm summer and no dry season, and *Dsc* – Cold
 210 climate with cold summers and dry summers. Figure 4 shows the KG classification for Iceland for the
 211 years 1950 (upper) and 2099 (lower).



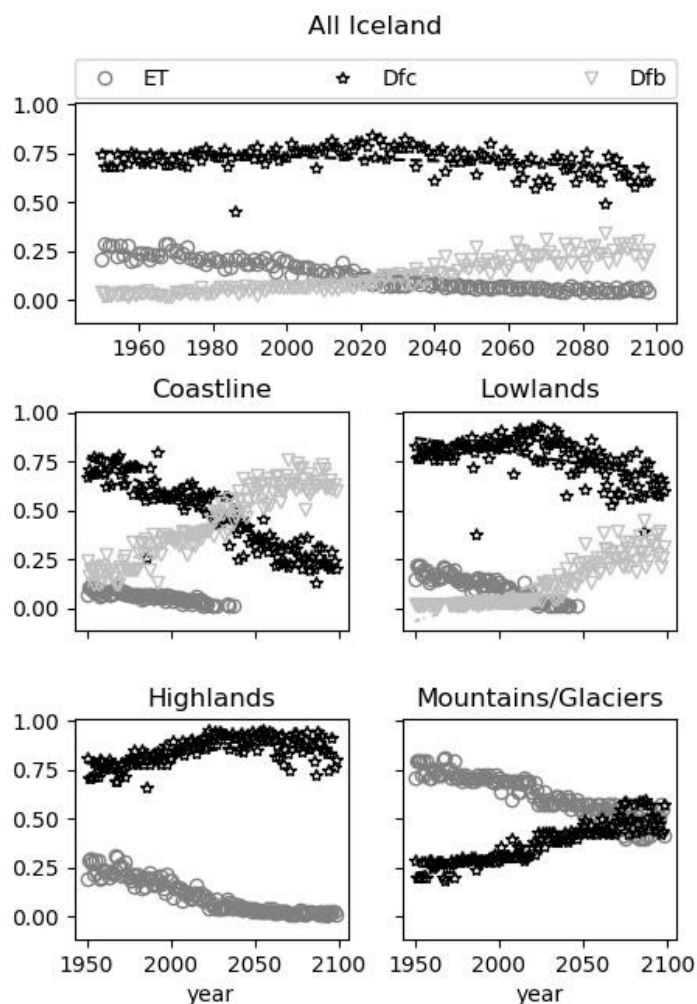
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213 Figure 4. Annual Köppen-Geiger (KG) classification maps calculated for Iceland in the study for the years 1950 (upper) and
214 2099 (lower). (*Dfc* – Cold climate with cold summers and no dry season and *Dfb* – Cold climate with warm summer and no
215 dry season)

216

217 The results show that between 1950 and 2099 the polar tundra climate that initially covered a large
218 portion of the highlands will mostly disappear, except on the Vatnajökull Glacier. Here *ET* is replaced
219 by a cold climate with the cold summer classes *Dfc* and *Dsc*. A warm summer climate (*Dfb*), that in the
220 beginning of the period was mostly limited to small areas in the southern lowlands on each side of the
221 Mýrdalsjökull Glacier, will by the end of the period have spread almost around the entire country and
222 stretched far into the highlands. At the middle of the current century temperate climate classes (*Cfb*,
223 *Cfc* and *Csb*) start appearing consistently in coastal areas. This would be the first time that such a
224 climate classification would be experienced in Iceland.

225 Figure 5 shows the proportional coverage of the top climate classes for the period 1950-2099.
226 The uppermost graph shows the results for the whole of Iceland and the lower graphs show the main
227 climate classes within each elevation zone. The results in Figure 5 show that by the end of the current
228 century the polar tundra climate (Class *ET*) in Iceland will decrease from about 20% coverage in 1950
229 to about 5%, by the middle of the current century and the *ET* class will disappear altogether in the
230 coastal and lowland regions. Over the same period, warm summers (class *Dfb*) will increase by about
231 the same amount. The net coverage of the most common climate class, *Dfc*, will not change much
232 over the period. However, as seen in Figures 4 and 6, class *Dfc* is replaced by class *Dfb* in coastal areas
233 while it replaces class *ET* in the highlands; thus we expect the spatial distribution of class *Dfc* to change
234 significantly during the period.



235

236 *Figure 5. Proportional coverage of the three main KG Classes in each elevation zone in Iceland for the period 1950-2099*

237 3.4 Comparison of SCF and Climate trends

238 We calculated the ratio of the area where significant SCF changes were observed within each of the
239 elevation zones described in section 2.2 for the MODIS period (2001-2016). We also calculated the
240 proportional annual coverage of the main KG climate classes in each elevation zone for both the
241 MODIS period and for the full historical climate time series (1950-2016). The significance of the trend
242 of the proportional annual KG coverage was assessed using the Mann-Kendall and Sen's estimator of
243 slope hypothesis tests, as described in section 2.5.

244 Table 3 summarizes the evidence for the statistical significance of these changes. The results
245 in Table 3 show that the biggest SCF changes have occurred in the highlands, where 27.4% of the area
246 has already seen a significant change in the SCF. The smallest SCF changes were observed by the



247 coastline and at the highest elevations in mountain peaks and on glaciated terrain, where 2.4 and 5.4%
 248 of the areas has undergone a significant change in SCF, respectively. Areas with significant SCF changes
 249 covered 14.3 % in the lowlands, and 16.3% across Iceland.

250 The results in Table 2 show that for the MODIS period there have been no significant changes
 251 to any of the main climate changes in Iceland nor within any of the elevation zones considered.
 252 However, for the full historical period 1950-2016, we observed significant changes in every one of the
 253 climate classes within all elevation zones. For the full historical period the changes in climate classes
 254 were significant to a confidence level of $\alpha = 0.001$. These results suggest that climate change is
 255 underway in Iceland, but that during the MODIS period, these changes were not identifiable by
 256 statistical means.

257
 258

Table 1. Changes in proportional annual KG coverage and percentage of area with significant SCF changes

Area	KG classes 1950-2016				KG classes 2001-2016			SCF Significant Area [%]		Mean SCF change [%/yr]
	Name	Slope [%/yr]	MK p-value	Sens's slope	Slope [%/yr]	MK p-value	Sens's slope	$\alpha = 0.05$	$\alpha = 0.01$	
All Iceland	Dfc	0.10	< 0.001	Sig	0.15	> 0.05	InSig	16.3	7.4	0.64
	ET	-0.18	< 0.001	Sig	-0.24	> 0.05	InSig			
	Dfb	0.07	< 0.001	Sig	0.02	> 0.05	InSig			
Coastline	Dfc	-0.27	< 0.001	Sig	-0.06	> 0.05	InSig	2.4	0.7	0.08
	ET	-0.08	< 0.001	Sig	-0.12	> 0.05	InSig			
	Dfb	0.33	< 0.001	Sig	0.1	> 0.05	InSig			
Lowlands	Dfc	0.11	< 0.001	Sig	0.1	> 0.05	InSig	14.3	5.9	0.67
	ET	-0.18	< 0.001	Sig	-0.34	> 0.05	InSig			
	Dfb	0.04	< 0.001	Sig	0.02	> 0.05	InSig			
Highlands	Dfc	0.22	< 0.001	Sig	0.17	> 0.05	InSig	27.4	13.3	0.95
	ET	-0.23	< 0.001	Sig	-0.26	> 0.05	InSig			
Glacier/ Mountain	Dfc	0.1	< 0.001	Sig	0.003	> 0.05	InSig	5.4	2.3	0.23
	ET	-0.1	< 0.001	Sig	0	> 0.05	InSig			

259

260 5 Discussion

261 In this study we assessed the impact of climate change on snow cover in Iceland. Our results
 262 showed that warmer climate classes have been replacing colder across the country. In the highlands,
 263 cold climate with cold summers, *Dfc*, has been replacing polar tundra, and at lower elevations cold
 264 climate with warm summer, *Dfb*, has been replacing class *Dfc*. This warming trend has been
 265 progressing since at least the 1950's and is unaffected by the reported midcentury cooling trend
 266 (Hanna et al., 2004). However, while the Icelandic climate has been warming, large areas have
 267 experienced a significant increase in the SCF in the MODIS period 2001-2016, most notably the inland
 268 regions in the highlands and lowlands where SCF has increased by about 2.4 and 3.5 days/year in the
 269 study period, respectively. These results are inconsistent with our original hypothesis that the
 270 warming trend observed would result in a reduced frequency of snow cover.

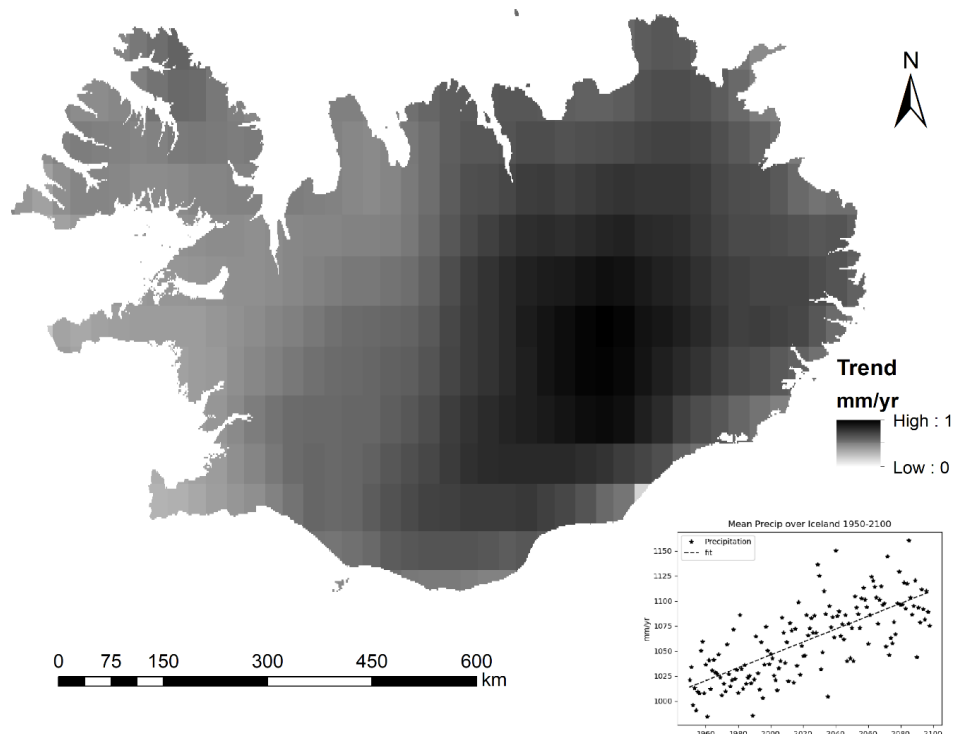
271 The results clearly identify local changes to the snow cover regime during the MODIS era. Land
 272 surface changes like reservoir construction, volcanic eruptions and glacial recession are clearly
 273 observed in the SCF data and can be related to changes in snow cover. Since 2000 the area of the
 274 glaciers in Iceland has been reduced by 600 km² (IMO, 2017). This recent recession of the country's
 275 main glaciers was clearly observed in our results, as in the areas in front of the retreating termini of
 276 all the major outlet glaciers in Iceland areas of significantly decreasing SCF were observed. These



277 results are somewhat counterintuitive, as most of these areas are in the highlands, where our results
278 showed very large areas of increasing SCF trends.

279 Several recent studies have analyzed snow cover changes in the Northern Hemisphere (NH),
280 most of whom observed an **increase** in snow cover in the past few decades. Yunlong et al., 2018
281 estimated a decrease in NH snow covered days by 5.3 days/decade since 2001, using MODIS, IMS and
282 AMSR-E data. Hori et al., 2017 used MODIS and AVHRR data to estimate changes in Snow Cover Extent
283 (SCE) across the NH and found an average decrease of 10 days/decade in large areas of the NH during
284 the period 1978-2015. Liston & Hiemstra, 2011 used MERRA reanalysis data to model NH snow cover
285 and observed a 2.5 days/decade decrease in the number of snow covered days for the period 1979-
286 2009. Fontrodona et al, 2018 analyzed in situ snow depth data across Europe, including sites in Iceland,
287 and found an average decrease in mean snow depth of – 12.2% per decade. Eythorsson et al., 2019
288 used MODIS data to estimate a 9.1 days / decade decrease in the Arctic SCF during the period 2001 –
289 2016, their results however showed large areas in the far north where SCF is increasing during the
290 period. The increasing SCF at high latitudes is consistent with studies that suggest that decreasing sea
291 ice concentrations in the Arctic seas have resulted in increasing precipitation in these areas (e.g. Kopec
292 et al., 2016; Singarayer et al., 2006). The observed pattern of increasing SCF trends in Iceland is
293 consistent with the results of Gunnarsson et al., (2019) which found an increasing snow cover duration
294 in all months except October and November using a gap filled MODIS product.

295 We suggest that the increasing SCF trends observed in Iceland could be associated with
296 increased precipitation, which would lead to a more frequent and thicker snowpack which would
297 persist longer in the spring. This is consistent with Bjornsson et al., (2018) who found annual
298 precipitation to have increased by about 10% during the period 1980-2015. Figure 6 shows Sen's slope
299 of the average annual precipitation during the period 1950-2100, according to the ensemble mean of
300 the NEX GDDP projections for RCP 4.5. Figure 6 shows that the average annual precipitation is
301 expected to increase during this time across all regions in Iceland, but most in the eastern highlands.
302 The small graph in Figure 6 shows the average annual precipitation across Iceland for the period 1950-
303 2100. The results show that precipitation in Iceland has shown an upward trend since the 1950's which
304 is expected to continue throughout the present century (Thrasher et al., 2006).



305
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Figure 6 Sens slope of precipitation changes and time series of mean annual precipitation in Iceland, 1950-2100

307

308 These results deserve further investigation. We note that the MODIS period, 2001-2016, is
309 rather short and trends during this short period can be induced by low frequency cyclical climate
310 patterns, or by a small amount of extreme weather events. We suggest that the physical reasons for
311 the SCF trends we observed in our study be analyzed further. The impact of increased precipitation on
312 the boundaries of the melt season in spring and autumn should be assessed. Changes in the intensity
313 of the summer melt season as the climate warms should be investigated. Lastly, we note that the
314 changes in the content of water stored in the snowpack cannot be inferred from the SCF data as they
315 only measure the frequency of snow cover. However, it is not unreasonable assume that water storage
316 in snow correlates positively with SCF, but this hypothesis deserves further analysis.



317 5 Conclusion

318 All over Iceland we observed a trend of warmer climate classes reaching further inland and
319 into the central highlands, replacing colder climates like polar tundra. Our study showed these changes
320 occurring since the 1950's based on the CMIP5 projections. Our results show that these trends can be
321 expected to continue until at least the end of the current century. We found these changes to be
322 statistically significant during the period 1950-2016, but during the MODIS period (2001-2016) we did
323 not find evidence for any statistically significant changes in climate classifications anywhere in Iceland.
324 However, during the MODIS period, our results showed evidence for changes in the SCF over large
325 areas of the country, especially in the central highlands. In most of these areas the SCF had increased
326 during the period. These results provide clear evidence of ongoing climate change in Iceland.

327 Projected until the end of the present century our results showed that warmer climate classes
328 that have until now been restricted to small areas in the southern lowlands will continue to spread
329 north and into the highlands. Well into the current century we observed temperate climate classes
330 appearing consistently in coastal areas for the first time in recorded history. If these climate trends
331 continue, we expect that the increasing SCF trends we observed will halt or reverse in the future,
332 especially in the lowland regions. The evolution of the snow regime in the Icelandic highlands deserves
333 further attention as these are the sources of a large portion of the country's water resources.
334

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342



343 References

- 344 Adalgeirsdottir, G., Johannesson, T., Bjornsson, H., Palsson, F., & Sigurdsson, O. : Response of
345 Hofsjokull and southern Vatnajokull, Iceland, to climate change. *Journal of Geophysical*
346 *Research-Earth Surface*, 111(F3). <https://doi.org/Artn F03001>\rDoi 10.1029/2005jf000388,
347 2006
- 348 Bjornsson, Halldor, Sigurdsson, B., Davidsdottir, B., Olafsson, J., Astthorsson, O., Olafsdottir, S., ...
349 Jonsson, T.: *Climate Change and it's impact on Iceland - Report of the scientific committee on*
350 *Climate Change*, 2018
- 351 Bjornsson, Helgi, & Palsson, F.: Icelandic glaciers. *Jökull*, 58(58), 365–386.
- 352 Bjornsson, Helgi, Palsson, F., Gudmundsson, S., Magnusson, E., Adalgeirsdottir, G., Johannesson, T.,
353 Berthier, E., Sigurdsson O., Thorsteinsson, T. (2013). Contribution of Icelandic ice caps to sea
354 level rise: Trends and variability since the Little Ice Age. *Geophysical Research Letters*, 40(8),
355 1546–1550. <https://doi.org/10.1002/grl.50278>, 2008
- 356 Brown, R. D., & Mote, P. W.: The response of Northern Hemisphere snow cover to a changing
357 climate. *Journal of Climate*, 22(8), 2124–2145. <https://doi.org/10.1175/2008JCLI2665.1>, 2009
- 358 Callaghan, T. V., Johannsson, M., Brown, R. D., Groisman, P. Y., Labba, N., Radionov, V., Bradley R. S.,
359 Blangy S., Bulygina O. N., Christensen T. R., Colman J. E., Essery R. L. H., Forbes B. C.,
360 Forchhammer M. C., Golubev V. N., Honrath R. E., Juday G. P., Meshcherskaya A. V., Phoenix G.
361 K., Pomeroy J., Rautio A., Robinson D. A., Schmidt N. M., Serreze M. C., Shevchenko V. P.,
362 Shiklomanov A. I., Shmakin A. B., Sköld P., Strum M., Woo M., Wood, E. F.: Multiple effects of
363 changes in arctic snow cover. *Ambio*, 40(SUPPL. 1), 32–45. [https://doi.org/10.1007/s13280-](https://doi.org/10.1007/s13280-011-0213-x)
364 011-0213-x, 2011
- 365 Chen, D., & Chen, H. W.: Using the Koppen classification to quantify climate variation and change: An
366 example for 1901-2010. *Environmental Development*, 6(1), 69–79.
367 <https://doi.org/10.1016/j.envdev.2013.03.007>, 2013
- 368 Cohen, J.: Snow cover and climate. *Weather*, 49(5), 150–156. [https://doi.org/10.1002/j.1477-](https://doi.org/10.1002/j.1477-8696.1994.tb05997.x)
369 8696.1994.tb05997.x, 1994
- 370 Dietz, A. J., Kuenzer, C., Gessner, U., & Dech, S.: Remote sensing of snow – a review of available
371 methods. *International Journal of Remote Sensing*, 33(13), 4094–4134.
372 <https://doi.org/10.1080/01431161.2011.640964>, 2012
- 373 Drapela, K., & Drapelova, I.: Application of Mann-Kendall test and the Sen's slope estimates for trend
374 detection in deposition data from Bily Kriz (Beskydy Mts., the Czech Republic) 1997--2010.
375 *Beskydy*, 4(2), 133–146. 2011
- 376 Eliasson, K., Ulfarsson, G. F., Valsson, T., & Gardarsson, S. M.: Identification of development areas in
377 a warming Arctic with respect to natural resources, transportation, protected areas, and
378 geography. *Futures*, 85, 14–29. <https://doi.org/10.1016/j.futures.2016.11.005>, 2017
- 379 Engelbrecht, C. J., & Engelbrecht, F. A.: Shifts in Köppen-Geiger climate zones over southern Africa in
380 relation to key global temperature goals. *Theoretical and Applied Climatology*.
381 <https://doi.org/10.1007/s00704-014-1354-1>, 2016
- 382 Eythorsson, D., Gardarsson, S. M., Ahmad, S. K., Hossain, F., & Nijssen, B.: Arctic climate and snow
383 cover trends - Comparing Global Circulation Models with remote sensing observations.
384 *International Journal of Applied Earth Observation and Geoinformation*, 80, 71–81.
385 <https://doi.org/https://doi.org/10.1016/j.jag.2019.04.003>, 2019
- 386 Eythorsson, D., Gardarsson, S. M., Gunnarsson, A., & Hrafnkelsson, B.: Statistical summer mass-
387 balance forecast model with application to Brúarjökull glacier, South East Iceland. *Journal of*
388 *Glaciology*. <https://doi.org/10.1017/jog.2018.22>, 2018.
- 389 Fontrodona Bach, A., van der Schrier, G., Melsen, L. A., Klein Tank, A. M. G., & Teuling, A. J.:
390 Widespread and Accelerated Decrease of Observed Mean and Extreme Snow Depth Over
391 Europe. *Geophysical Research Letters*. <https://doi.org/10.1029/2018GL079799>, 2018
- 392 Frei, A., Miller, J. A., & Robinson, D. A.: Improved simulations of snow extent in the second phase of



- 393 the Atmospheric Model Intercomparison Project (AMIP-2). *Journal of Geophysical Research*,
394 108(D12), 4369. <https://doi.org/10.1029/2002JD003030>, 2003
- 395 Gocic, M., & Trajkovic, S.: Analysis of changes in meteorological variables using Mann-Kendall and
396 Sen's slope estimator statistical tests in Serbia. *Global and Planetary Change*, 100, 172–182.
397 <https://doi.org/10.1016/j.gloplacha.2012.10.014>, 2013
- 398 Gorelick, N., Hancher, M., Dixon, M., Ilyushchenko, S., Thau, D., & Moore, R.: Google Earth Engine:
399 Planetary-scale geospatial analysis for everyone. *Remote Sensing of Environment*.
400 <https://doi.org/10.1016/j.rse.2017.06.031>, (2016).
- 401 Guan, B., Molotch, N. P., Waliser, D. E., Jepsen, S. M., Painter, T. H., & Dozier, J.: Snow water
402 equivalent in the Sierra Nevada: Blending snow sensor observations with snowmelt model
403 simulations. *Water Resources Research*, 49(8), 5029–5046.
404 <https://doi.org/10.1002/wrcr.20387>, 2013
- 405 Gunnarsson, A., Garðarsson, S. M., & Sveinsson, Ó. G. B.: Icelandic snow cover characteristics
406 derived from a gap-filled MODIS daily snow cover product. *Hydrol. Earth Syst. Sci.*, 23(7), 3021–
407 3036. <https://doi.org/10.5194/hess-23-3021-2019>, 2019
- 408 H Souri, A., & Azizi, A.: Removing Bowtie Phenomenon by Correction of Panoramic Effect in MODIS
409 Imagery. *International Journal of Computer Applications, Published by Foundation of Computer
410 Science, New York, USA.*, 68, 12–16, 2013
- 411 Hall, D. K., Riggs, G. A., & Salomonson, V. V. MODIS/Terra Snow Cover Daily L3 Global 500m Grid
412 V005 2000-01-01 to 2016-12-31., 2006
- 413 Hanna, E., Jónsson, T., & Box, J. E.: An analysis of Icelandic climate since the nineteenth century.
414 *International Journal of Climatology*, 24(10), 1193–1210. <https://doi.org/10.1002/joc.1051>,
415 2004
- 416 Hori, M., Sugiura, K., Kobayashi, K., Aoki, T., Tanikawa, T., Kuchiki, K., Niwano M., Enomoto, H.: A 38-
417 year (1978–2015) Northern Hemisphere daily snow cover extent product derived using
418 consistent objective criteria from satellite-borne optical sensors. *Remote Sensing of
419 Environment*, 191, 402–418. <https://doi.org/10.1016/j.rse.2017.01.023>, 2017
- 420 IMO.: *Overview of Icelandic Glaciers at the End of 2017*. Retrieved from
421 http://en.vedur.is/media/Eplicanámскеið/VAT_newsletter_2018_06.pdf, 2017
- 422 Johannesson, T., Adalgeirsdottir, G., Björnsson, H., Crochet, P., Eliasson, B. E., Gudmundsson, S.,
423 Jonsdottir J. F., Olafsson H., Pálsson F., Rognvaldsson O., Sigurdsson O., Snorrason A., Sveinsson
424 O. G. B., Thorsteinsson, T.: *Effect of climate change on hydrology and hydro-resources in Iceland*.
425 Reykjavik, 2007
- 426 Jóhannesson, T., Adalgeirsdottir, G., Björnsson, H., Pálsson, F., & Sigurðsson, O.: *RESPONSE OF
427 GLACIERS AND GLACIER RUNOFF IN ICELAND TO CLIMATE CHANGE*, 2004
- 428 Kopec, B. G., Feng, X., Michel, F. A., & Posmentier, E. S.: Influence of sea ice on Arctic precipitation.
429 *Proceedings of the National Academy of Sciences*, 113(1), 46–51.
430 <https://doi.org/10.1073/pnas.1504633113>, 2016
- 431 Köppen, W.: Klassifikation der Klimate nach Temperatur, Niederschlag und Jahresablauf
432 (Classification of climates according to temperature, precipitation and seasonal cycle).
433 *Petermanns Geographische Mitteilungen*, 1918
- 434 Kottek, M., Grieser, J., Beck, C., Rudolf, B., & Rubel, F.: World map of the Köppen-Geiger climate
435 classification updated. *Meteorologische Zeitschrift*, 15(3), 259–263.
436 <https://doi.org/10.1127/0941-2948/2006/0130>, 2006
- 437 Liston, G. E., & Hiemstra, C. A.: The changing cryosphere: Pan-Arctic snow trends (1979–2009).
438 *Journal of Climate*, 24(21), 5691–5712. <https://doi.org/10.1175/JCLI-D-11-00081.1>, 2011
- 439 Lohmann, U., Sausen, R., Bengtsson, L., Cubasch, U., Perlwitz, J., & Roeckner, E.: The Koppen climate
440 classification as a diagnostic tool for general circulation models. *CLIM.RES.*, 3(3), 177–193.
441 <https://doi.org/10.3354/cr003177>, 1993
- 442 Massé, G., Rowland, S. J., Sicre, M. A., Jacob, J., Jansen, E., & Belt, S. T.: Abrupt climate changes for
443 Iceland during the last millennium: Evidence from high resolution sea ice reconstructions. *Earth*



- 444 *and Planetary Science Letters*. <https://doi.org/10.1016/j.epsl.2008.03.017>, 2008
- 445 Naha, S., Thakur, P. K., & Aggarwal, S. P.: Hydrological Modelling and data assimilation of Satellite
446 Snow Cover Area using a Land Surface Model, VIC. *ISPRS - International Archives of the*
447 *Photogrammetry, Remote Sensing and Spatial Information Sciences, XLI-B8*(July), 353–360.
448 <https://doi.org/10.5194/isprsarchives-XLI-B8-353-2016>, 2016
- 449 Oliphant, T. E.: Python for scientific computing. *Computing in Science and Engineering*, 9(3), 10–20.
450 <https://doi.org/10.1109/MCSE.2007.58>, 2007
- 451 Peel, M. C., Finlayson, B. L., & McMahon, T. A.: Updated world map of the Köppen–Geiger climate
452 classification. *HESD Earth Syst. Sci. Discuss*, 4(4), 439–473. [https://doi.org/10.1127/0941-](https://doi.org/10.1127/0941-2948/2006/0130)
453 [2948/2006/0130](https://doi.org/10.1127/0941-2948/2006/0130). 2007
- 454 Rubel, F., Brügger, K., Haslinger, K., & Auer, I.: The climate of the European Alps: Shift of very high
455 resolution Köppen–Geiger climate zones 1800–2100. *Meteorologische Zeitschrift*.
456 <https://doi.org/10.1127/metz/2016/0816>, 2017
- 457 Singarayer, J. S., Bamber, J. L., Valdes, P. J., Singarayer, J. S., Bamber, J. L., & Valdes, P. J.: Twenty-
458 First-Century Climate Impacts from a Declining Arctic Sea Ice Cover. *Journal of Climate*, 19(7),
459 1109–1125. <https://doi.org/10.1175/JCLI3649.1>, 2006
- 460 Stocker, T. F., Qin, D., Plattner, G. K., Tignor, M. M. B., Allen, S. K., Boschung, J., Nauels A., Xia Y., Bex
461 V., Midgley, P. M.: *Climate change 2013 the physical science basis: Working Group I*
462 *contribution to the fifth assessment report of the intergovernmental panel on climate change.*
463 *Climate Change 2013 the Physical Science Basis: Working Group I Contribution to the Fifth*
464 *Assessment Report of the Intergovernmental Panel on Climate Change.*
465 <https://doi.org/10.1017/CBO9781107415324>, 2013
- 466 Taylor, K. E., Stouffer, R. J., & Meehl, G. A.: An overview of CMIP5 and the experiment design.
467 *Bulletin of the American Meteorological Society*. <https://doi.org/10.1175/BAMS-D-11-00094.1>,
468 2012
- 469 Thrasher, B., Melton, F., & Weile, W.: NASA Earth Exchange Global Daily Downscaled Projections
470 (NEX-GDDP) | CDS. Retrieved May 1, 2017, from <https://cds.nccs.nasa.gov/nex-gddp/>, 2006
- 471 van Vuuren, D. P., Edmonds, J., Kainuma, M., Riahi, K., Thomson, A., Hibbard, K., ... Rose, S. K.: The
472 representative concentration pathways: An overview. *Climatic Change*, 109(1), 5–31.
473 <https://doi.org/10.1007/s10584-011-0148-z>, 2011
- 474 Yunlong, W., Huang, X., Hui, L., Sun, Y., Qisheng, F., & Tiangang, L.: Tracking Snow Variations in the
475 Northern Hemisphere Using Multi-Source Remote Sensing Data (2000–2015). *Remote Sensing*,
476 10(1)(136). <https://doi.org/doi:10.3390/rs10010136>, 2018
- 477
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479 Appendix 1. Köppen-Geiger climate classification criteria
 480

1 st	2 nd	3 rd	Description	Criteria*
A			Tropical	$T_{\text{cold}} \geq 18$
	f		-Rainforrest	$P_{\text{dry}} \geq 60$
	m		-Monsoon	Not Af and $P_{\text{dry}} \geq 100\text{-MAP}/25$
	w		-Savannah	Not Af and $P_{\text{dry}} < 100\text{-MAP}/25$
B			Arid	$\text{MAP} < 10 * P_{\text{threshold}}$
	w		-Desert	$\text{MAP} < 5 * P_{\text{threshold}}$
	s		-Steppe	$\text{MAP} \geq 5 * P_{\text{threshold}}$
		H	-Hot	$\text{MAT} \geq 18$
		K	-Cold	$\text{MAT} < 18$
C			Temperate	$T_{\text{hot}} > 10$ & $0 < T_{\text{cold}} < 18$
	s		-Dry Summer	$P_{\text{sdry}} < 40$ & $P_{\text{sdry}} < P_{\text{wwet}}/3$
	w		-Dry Winter	$P_{\text{wdry}} < P_{\text{swet}}/10$
	f		-Without dry season	Not Cs or Cw
		A	-Hot Summer	$T_{\text{hot}} \geq 22$
		B	-Warm Summer	Not a & $T_{\text{mon}10} \geq 4$
		C	-Cold Summer	Not a or b & $1 \leq T_{\text{mon}10} \leq 4$
D			Cold	$T_{\text{hot}} > 10$ and $T_{\text{cold}} \leq 0$
	s		-Dry Summer	$P_{\text{sdry}} < 40$ & $P_{\text{sdry}} < P_{\text{wwet}}/3$
	w		-Dry Winter	$P_{\text{wdry}} < P_{\text{swet}}/10$
	f		-Without dry season	Not Ds and Not Dw
		A	-Hot Summer	$T_{\text{hot}} \geq 22$
		B	-Warm Summer	Not a & $T_{\text{mon}10} \geq 4$
		C	-Cold Summer	Not a, b or d
		D	-Very cold Winter	Not a or b & $T_{\text{cold}} < -38$
E			Polar	$T_{\text{hot}} < 10$
	T		-Tundra	$T_{\text{hot}} > 0$
	F		-Frost	$T_{\text{hot}} \leq 0$

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* MAP = Mean annual precipitation, MAT = mean annual temperature, T_{hot} = temperature of the hottest month, T_{cold} = temperature of the coldest month, $T_{\text{mon}10}$ = number of months with mean temperatures above 10, P_{dry} = precipitation of the driest month, P_{sdry} = precipitation of the driest month in summer, P_{wdry} = precipitation of the driest month in winter, P_{swet} = precipitation of the wettest month in summer, P_{wwet} = precipitation of the wettest month in winter, $P_{\text{threshold}}$ = varies according to: (if 70% of MAP occurs in winter then $P_{\text{threshold}} = 2 * \text{MAT}$, if 70% of MAP occurs in summer then $P_{\text{threshold}} = 2 * \text{MAT} + 28$, otherwise $P_{\text{threshold}} = 2 * \text{MAT} + 14$). Summer (winter) is defined as the warmer (cooler) six-month period of ONDJFM and AMJJAS.
 From Peel et al., 2007

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