

1 **HESS Opinions: Beyond the Long-term Water Balance: Evolving Budyko's**
2 **Supply-Demand Framework for the Anthropocene towards a Global**
3 **Synthesis of Land-surface Fluxes under Natural and Human-altered**
4 **Watersheds**

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15
16 **Abstract**

17 Global hydroclimatic conditions have been significantly altered over the past century by
18 anthropogenic influences that arise from the warming global climate and also from local/regional
19 anthropogenic disturbances. Traditionally, studies have used coupling of multiple models to
20 understand how land-surface water fluxes vary due to changes in global climatic patterns and
21 local land-use changes. We argue that because the basis of the Budyko framework relies on the
22 supply and demand concept, the framework could be effectively adapted and extended to
23 quantify the role of drivers – both changing climate and local human disturbances – in altering

24 the land-surface response across the globe. We review the Budyko framework, along with these
25 potential extensions, with the intent of furthering the applicability of the framework to emerging
26 hydrologic questions. Challenges in extending the Budyko framework over various spatio-
27 temporal scales and the use of global data sets to evaluate the water balance at these various
28 scales are also discussed.

29

30 **The historical evolution of the Budyko framework in hydroclimatology**

31 The traditional Budyko formulation provides the long-term water balance as a simple but
32 effective partitioning of precipitation into runoff and evapotranspiration; and it has been verified
33 over numerous natural watersheds around the globe (Sankarasubramanian and Vogel, 2003;
34 Zhang et al., 2004; Yang et al., 2007; Li et al., 2013; Padrón et al., 2017). Besides the aridity
35 index, which is defined as the ratio of the mean annual potential evapotranspiration to the mean
36 annual precipitation, Milly et al. (1994) and Sankarasubramanian and Vogel (2002) proposed
37 additional controls on the long-term water balance including seasonality and soil moisture
38 holding capacity. These additional controls enhance the ability of the Budyko framework to
39 explain the spatial variability in mean annual runoff at the continental scale. Studies have also
40 extended the Budyko framework for capturing the interannual variability in runoff (Koster and
41 Suarez, 1999; Sankarasubramanian and Vogel, 2002, 2003). More recently, the Budyko
42 framework has been extended for explaining the seasonal hydroclimatology of basins (Petersen
43 et al., 2012; Chen et al. 2013; Petersen et al. 2018). Similarly, the Budyko framework has been
44 extended for quantifying the non-dimensional sensitivity (also termed elasticity) of land-surface
45 response to changes in climatic controls under different hydroclimatic regimes (Dooge, 1992;
46 Dooge et al., 1999; Sankarasubramanian et al., 2001).

47 Perhaps the most unique aspect of the Budyko framework lies in its Darwinian approach,
48 which enables us to view the entire hydroclimatic system without focusing on each physical
49 process in isolation (Harman and Troch, 2014; Wang and Tang, 2014). The Darwinian approach
50 seeks to document patterns of variation in populations of hydrologic systems and develop
51 theories that explain these observed patterns in terms of the mechanisms and conditions that
52 determine their historical development (Harman and Troch, 2014). Even though most studies
53 which employed the Budyko framework have focused on natural basins, the original monograph
54 (Budyko, 1974), *Climate and Life*, considered the role of human influence on climate, including
55 impacts of reservoir storage and irrigation on evapotranspiration. As hydroclimatic regimes
56 evolve in the Anthropocene, it is critical to understand how land-surface fluxes change due to
57 changes in local watershed conditions and global climate change. Given the Budyko
58 framework’s emphasis on describing patterns of variation across differing hydrogeologic and
59 hydroclimatic regimes – and by extension, its emphasis on an integrative, Darwinian approach –
60 coupled with its ability to capture the fundamental dimensions of land-surface fluxes, a global
61 synthesis addressing the variability in these fluxes across natural and human-altered watersheds
62 should provide insights on the sensitivity of the critical hydroclimatic processes to local and
63 global changes in the Anthropocene.

64

65 **Budyko Framework for the Anthropocene**

66 We are at a critical time in which the hydroclimate, particularly land-surface fluxes, have
67 been significantly altered over the past century by anthropogenic disturbances (Entekhabi et al.,
68 1999; Vogel et al., 2015). For instance, both annual precipitation and streamflow have increased
69 during the period of 1948–1997 across the eastern United States, and those trends appear to arise

70 primarily from increases in autumn precipitation (Small et al., 2006; Rice et al., 2015). Similarly,
71 the frequency of floods is increasing in many regions, while magnitudes of flooding appear only
72 to be systematically increasing in certain spatially cohesive regions (Hirsch and Archfield, 2015;
73 Malikpour and Villarini, 2015; Archfield et al., 2016) particularly in urban areas (Vogel et al.,
74 2011; Barros et al., 2014 and Prosdocimi et al., 2015). Irrigation in the U.S. high plains leads to
75 increases in summer rainfall and streamflow in the Midwest due to land-surface and atmosphere
76 feedback (Kustu et al., 2011). Based on hydroclimatic observations from 100 large basins
77 globally, Jaramillo and Destouni (2015) found consistent and dominant effects of increasing
78 relative evapotranspiration (evapotranspiration relative to precipitation) from flow regulation and
79 irrigation and decreasing temporal runoff variability from flow regulation. Development of
80 irrigation networks and man-made reservoirs is also associated with increased surface water and
81 groundwater withdrawals and land-use change (Dieter et al., 2018; Sankarasubramanian et al.,
82 2017; Das et al., 2018). Similarly, construction of large dams has significantly altered the
83 downstream flow variations impacting downstream ecology (Gao et al., 2008; Wang et al.,
84 2017). Changes in land use and land cover also impact the local energy balance, creating urban
85 heat islands (Memon et al., 2008), affecting recharge and baseflow (Price, 2011), which in turn
86 impacts a very broad range of streamflows (Allaire et al., 2015) with particularly significant
87 increases in high flows (Vogel et al., 2011; Barros et al. 2014; Prosdocimi et al. 2015). Thus,
88 anthropogenic influences arising from global climate change and local to regional disturbances
89 can significantly impact the land-surface response from the watershed. Anthropogenic influences
90 including changes in climate, land use, and water use exhibit complex interactions which must be
91 considered jointly, to understand their impact on hydrologic flow alteration (Allaire et al. 2015).
92 Performing a synthesis on how the spatio-temporal variability of land-surface fluxes – runoff,

93 evapotranspiration, net radiation, and hydrologic flow alteration – differ globally in natural and
94 human-altered watersheds is a critical need to enable a complete understanding of global
95 hydroclimate during the Anthropocene. The Budyko framework provides an ideal approach for
96 such inquiry, because it has been used to decompose changes in observed land-surface fluxes due
97 to both natural variability and human influence (e.g., Roderick and Farquhar, 2011; Wang and
98 Hejazi, 2011; Yang et al., 2014; Jiang et al., 2015).

99 **Budyko Framework Adaptation in Watershed Modeling**

100 Figure 1 provides the general setup of the Budyko framework for explaining the spatio-
101 temporal variability of land-surface fluxes in natural watersheds and human-altered landscapes.
102 The framework relies on conservation of mass and energy to model and predict the “actual”
103 hydroclimatic variable of interest based on the available “demand” and “supply” of mass and
104 energy (Figure 1). The term “demand” is defined as the upper bound of the “actual” variable if
105 the “supply” variable is unlimited. The rationale for using the Budyko framework for
106 understanding the spatial variability in land-surface fluxes over natural/human-altered
107 watersheds lies in its ability to capture the hydroclimatic dimensions of “supply” and “demand”,
108 thereby providing a low-dimensional parsimonious approach (Figure 1) to understand the spatial
109 variability in the “actual” hydroclimatic variable of interest . Even though Budyko’s framework
110 is commonly used to understand the long-term water balance, the supply-demand framework
111 could be used for understanding the spatial variability of different land-surface fluxes. This
112 study evaluates and extends the Budyko supply-demand framework for understanding the spatio-
113 temporal variability of two hydrologic fluxes namely global evapotranspiration and water
114 balance (Figure 2), infiltration (Figure 3) and three human-altered fluxes namely reservoir
115 releases using linear hedging (Figure 4) and environmental flow alteration (Figure 5).

116 *Long-term Water Balance*

117 The most commonly used framework for modeling the long-term water balance is to
118 estimate the mean annual evapotranspiration (“actual”) based on the ratio of mean annual
119 potential evapotranspiration (“demand”) to the mean annual precipitation (“supply”). Thus, the
120 upper limit for mean annual evapotranspiration is potential evapotranspiration (precipitation) in a
121 humid (arid) region. The family of Budyko curves estimates the evapotranspiration ratio
122 (“actual”/ “supply”) based on the aridity index (“demand”/“supply”). For additional details, see
123 Sankarasubramanian and Vogel (2001). Most studies have focused on evaluating the long-term
124 water balance at regional and continental scale (see Wang et al., 2016 for a detailed review).
125 Studies have also focused on the impact of land cover and climate on long-term water yield using
126 global data (Li et al., 2013; Wang and Tang, 2014). Here, we evaluate the Budyko framework to
127 the global scale using the data from the Global Land-Surface Data Assimilation System, version
128 2 (GLDAS2) (Rodell et al., 2004). Data points of mean annual evapotranspiration and aridity
129 index are obtained using the Penman-Monteith method from the Noah Land Surface Model in
130 the GLDAS2 dataset with a spatial resolution of 0.25 ° for the period 1948-2010 (Rui, 2011).
131 Figure 2 shows the performance of the Budyko curve in estimating the mean annual
132 evapotranspiration based on aridity index data between -60° S to 60° N. The Budyko curve
133 provides a first-order approximation of the spatial variability in the evapotranspiration ratio
134 (Figure 2); however, the scatter around the curve is quite considerable. Studies have shown that
135 seasonality in moisture and energy and their co-availability (i.e., phase difference between
136 moisture and energy availability within the year) and soil moisture holding capacity partially
137 control the scatter around the Budyko curves (Milly et al., 1994, Sankarasubramanian and Vogel,
138 2003). Another question of interest is to understand the lower bound on the evapotranspiration

139 ratio, which is typically limited by the moisture availability in a region (Wang and Tang, 2014).
140 Numerous studies on long-term balance have employed fitting the observed long-term water
141 balance by parameterizing the Budyko curves (see the review paper of Wang et al., 2016).
142 However, little to no effort has been undertaken on how this cloud of the long-term water
143 balance is expected to change under potential climate change, and how this interplay between
144 moisture and energy is expected to affect the long-term water balance in different type of
145 watersheds (Creed et al., 2014). Similarly, recent studies have extended Budyko’s steady-state
146 supply-to-demand framework for modeling land-surface fluxes over fine (daily and monthly)
147 time scales (Zhang et al., 2008). Validating these emerging frameworks with global hydrologic
148 data will provide an understanding of the critical process controls in estimating land-surface
149 fluxes. This validation effort will also help in understanding the advantages and limitations of
150 such a parsimonious modeling approach towards estimating evapotranspiration and streamflow
151 at various spatio-temporal scales.

152 *Extension of Budyko’s “supply and demand” concept for infiltration*

153 The upper bounds on the Budyko framework arise from the conservation of mass and
154 energy. Hence, in principle, it could be applied to other hydrological processes. Zhang et al.
155 (2008) applied the Budyko monthly supply and demand attributes to estimate the catchment
156 retention and the overland runoff from the soil moisture zone. Wang (2018) developed the
157 infiltration equation for saturation excess in the Budyko supply and demand framework, i.e.,
158 modelling the ratio of infiltration to rainfall depth as a function of the ratio between storage
159 (infiltration) capacity and rainfall depth (Figure 3). The cumulative infiltration depth during a
160 rainfall event is defined as the “actual” variable of interest, and the cumulative rainfall depth
161 during an event is defined as the “supply”. The effective soil water storage capacity for the event

162 is defined as the “demand”, which is dependent on the initial soil moisture condition. Alternate
163 definitions of storage capacity based on soil hydraulic equilibrium storage could also be
164 considered (Zehe et al., 2019).

165 Figure 3 represents the initial soil moisture condition by the degree of saturation, ψ ,
166 which is defined as the ratio of initial soil water storage and storage capacity (Wang, 2018). For
167 a dry soil with low ψ , infiltration is expected to be higher with lower surface runoff potential.
168 The upper bounds of these curves (Figure 3) are similar to the Budyko’s asymptotes
169 corresponding to infiltration capacity-limited and rainfall depth-limited conditions. In this
170 illustration, the Budyko framework is extended to estimate the temporal variability of infiltration
171 into the soil based on soil water storage capacity and antecedent conditions (ψ). Thus, the
172 parsimonious framework derived from the Budyko’s supply and demand concept is used to
173 develop the asymptotes, and then those asymptotes are used to identify and explain various
174 critical process controls (e.g., infiltration in Figure 3).

175 Although the above extensions of the Budyko framework demonstrate the potential for
176 developing a low-dimensional parsimonious modeling strategy, data-based validation efforts
177 have focused primarily on the long-term hydroclimatic attributes (i.e., mean, variance and
178 elasticity) of observed land-surface fluxes in natural basins (Figure 2) (Sankarasubramanian and
179 Vogel, 2001; Abatzoglou and Ficklin, 2017). Representing a hydroclimatic variable of interest
180 (i.e., “actual”) as a ratio to the “supply” and explaining its spatio-temporal variability based on
181 the demand/supply ratio and other variables (e.g., soil moisture holding capacity for long-term
182 water balance) provides a simple, non-dimensional form for understanding the process controls.
183 For instance, in the long-term water balance context, defining the demand/supply relationship
184 explains the predominant controls on the spatio-temporal variability of mean annual runoff and

185 mean annual evapotranspiration based on the basin aridity, seasonality of demand and supply
186 attributes (i.e., in-phase or out-of-phase between moisture and energy), and soil moisture holding
187 capacity (Milly, 1991). Synthesizing relevant process controls and representing them within the
188 Budyko low-dimensional framework will also help us in catchment classification and in
189 understanding how different hydroclimatic processes of interest vary across wider regimes and
190 landscapes.

191

192 **Extending Budyko Framework for Human-altered Watersheds and Landscapes**

193 Figures 4 and 5 extend the Budyko framework to explain the spatio-temporal variability
194 in land-surface fluxes in human-altered watersheds and landscapes. A synthesis that extends and
195 evaluates the Budyko framework for estimating land-surface fluxes in human-altered watersheds
196 will help us understand the role of key anthropogenic disturbances (e.g., reservoir storage, land-
197 use and land-cover changes) in altering land-surface fluxes at various spatio-temporal scales.

198

199 *Extension of Budyko’s “supply and demand” Framework for Reservoir Operation and* 200 *Hedging*

201 We extend the Budyko framework for reservoir operation to meet the target demand
202 based on the standard operating policy (SOP) and linear hedging policy (Draper and Lund,
203 2004). A hedging policy in reservoir operation aims to conserve water for future use by
204 curtailing the current demand (Draper and Lund, 2004). Given an initial storage (S_{t-1}), inflow
205 (I_t), demand (D_t) and evaporation (E_t) over a given time step (t), one could obtain the actual
206 release (R_t), and ending storage (S_t) along with spill (SP_t) using a mass balance (equation 1).

207
$$S_t = S_{t-1} + I_t - E_t - R_t - SP_t \quad \dots (1)$$

208 By defining available water, $AW_t = S_{t-1} + I_t - E_t$, we obtain release (as “actual”) under a given
 209 hedging fraction ($0 \leq \alpha \leq 1$) for three reservoir storage conditions using equation 2. The SOP of a
 210 reservoir simply corresponds to $\alpha=1$ by supplying available water or demand at a given time.

$$\begin{aligned}
 & S_t = S_{\max}, R_t = D_t, SP_t = AW_t - D_t - S_{\max} \quad \text{if } AW_t - D_t \geq S_{\max} \\
 211 \quad & S_t = AW_t - R_t, R_t = \alpha D_t, SP_t = 0 \quad \quad \quad \text{if } S_{\min} < AW_t - D_t < S_{\max} \quad \quad \quad \dots(2) \\
 & S_t = S_{\min}, R_t = AW_t, SP_t = 0 \quad \quad \quad \text{if } S_{\min} \geq AW_t - D_t
 \end{aligned}$$

212 Rewriting AW_t as “supply”, D_t as “demand” and R_t (“actual”), we develop the Budyko
 213 framework for the reservoir operation under SOP and hedging policy (Figure 4). The SOP
 214 simply provides the asymptotes, the upper bounds, for the R_t / AW_t (“actual”/ “supply”) ratio.
 215 Figure 4 also demonstrates the developed framework for a hypothetical system for estimating the
 216 monthly releases (see supporting information (SI) Tables 1-2 for data and details). Increased
 217 hedging reduces the release and increases the storage and spill from the system. For
 218 demonstration, a linear hedging policy is applied. But the real-world system operation will have
 219 a complex non-linear release policy, still the data points are expected to lie within the bounds.
 220 For systems with a small storage-to-demand ratio, the spill portion on the left asymptote is
 221 expected to be much longer than a system with large storage-to-demand ratio. Similarly, for
 222 systems with large (small) storage-to-demand ratio, most data points are expected to lie below
 223 (on) the asymptotes portion of the framework. Given that this framework in Figure 4 is non-
 224 dimensional, we could analyze release to demand characteristics for reservoirs with competing
 225 purposes (e.g., hydroelectric vs flood control) and synthesize how release patterns vary based on
 226 the demand-to-available water ratio across different type of systems. Similarly, one can also
 227 formulate the functional forms for a non-linear hedging policy as Budyko equations, because the
 228 upper bounds are specified by the “supply and demand” relationship.

229

230 ***Representing Human Demand and Environmental Flows in from Reservoir Operation***

231 Reservoir storages reduce runoff variability to meet the human demand, thereby resulting
232 in significant flow alterations (Wang et al., 2014). By adding a dedicated term, environmental
233 flow, EF_t , we rewrite the reservoir mass balance in equation (3).

234
$$S_t = AW_t - R_t - EF_t - SP_t \quad \dots (3)$$

235 Given our variable of interest here is EF_t (“actual”), we represent the “demand” as $R_t + EF_t$ and
236 available water, AW_t , as “supply”, which gives us a simple framework to visualize the ratio,
237 environmental flow allocation EF_t / AW_t , has the upper bound AW_t , which is specified by the
238 1:1 line. The term, $1 - EF_t / AW_t$, simply represents the alteration ratio at a given time step. The
239 lower bound specifies only allocation ($R_t / EF_t = 0$) for human demand and a slope of 0.5
240 indicates equal allocation for human need and ecological demand. For instance, if R_t / EF_t falls
241 below the slope of 0.5, it indicates significant flow alteration to meet human demand.

242 In the case of Falls Lake (Figure 5), a major water supply reservoir in the triangle area in
243 North Carolina (see SI Table 3 for data and additional details), it is evident that flow alteration is
244 significant due to increased allocation for human demand since more data points lie below the
245 equal allocation line. Using the proposed framework in Figure 5, one could synthesize how a
246 reservoir system with a large residence time (commonly known as degree of regulation) could
247 impact flow alteration under arid and humid conditions. The negative linear trend indicates
248 (Figure 5) increased allocation for human use results in decreased environmental flow allocation.
249 For instance, reservoirs in arid (humid) climates are typically larger to reduce the larger (smaller)
250 interannual variability in runoff, hence such systems are expected to have higher (lower) degree
251 of regulation. However, this synthesis of reservoir systems across different climatic regimes

252 needs to be evaluated in the context of withdrawals for human use, their purpose, and the
253 consumptive use associated with them. We argue the proposed framework could be useful for
254 understanding the trade-off between water allocation for human use and downstream ecological
255 requirements.

256 We argue that the Budyko supply and demand framework should not be considered only
257 for the long-term water balance. As the supply and demand framework is based on conservation
258 equations, it could be exploited for understanding and quantifying the spatial variability in land-
259 surface fluxes under natural and human-altered landscapes. It is important to mention that the
260 relationship between “Actual/Supply” to “Demand/Supply” could arise due to spurious
261 correlation as they are shared by a common denominator term “Supply” (Benson, 1965; Bretts;
262 2014). Hence, the relationship between the “Actual” and the “Demand” should be evaluated
263 carefully by avoiding the spurious correlation. In Figures 4-5, we provide an extension of the
264 Budyko framework for understanding how land-surface fluxes are modified due to human
265 influence. Understanding the key drivers that alter the spatial variability of land-surface fluxes
266 using the modified and extended the Budyko framework should help in identifying the relevant
267 low-dimensional attributes that control the regional hydroclimate of human-altered watersheds
268 and landscapes. For long-term ET, the aridity index is the control. For infiltration, it is the ratio
269 of infiltration capacity to rainfall depth. For reservoir operation, it is the ratio of human water
270 demand to available water in reservoir. For environmental flows, it is the competition with
271 human demand and available water. Thus, the low-dimensional attribute varies for each
272 environmental issue. Further, extending the Budyko framework to include anthropogenic causes
273 should enable the explicit decomposition and attribution of changes in land-surface fluxes at
274 various temporal scales resulting from changes in local/regional hydroclimate or watershed-level

275 modification. To refine existing hydroclimatologic models and datasets developed at the
276 regional, continental, and global scale, a synthesis study is needed to understand how the land-
277 surface response varies across natural and human-altered watersheds. Such a synthesis effort
278 should also enable a systematic decomposition of watershed-scale anthropogenic influences and
279 large-scale climate impacts in modulating land-surface fluxes at the global scale, providing a
280 tribute to Budyko’s legacy.

281 **Opportunities, challenges, and relevance to other hydrologic synthesis studies**

282 Emphasis on understanding the complex interactions and feedback between human and
283 hydrological systems has renewed focus on “Socio-hydrology” (Sivapalan et al., 2012). The
284 impact of water use, land use and land cover and other anthropogenic influences on watershed
285 runoff and associated issues of non-stationarity have been referred to as the study of “Hydro-
286 morphology” (Vogel, 2011). Vogel et al. (2015) argue that “to resolve the complex water
287 problems that the world faces today, nearly every theoretical hydrologic model introduced
288 previously is in need of revision to accommodate how climate, land, vegetation, and
289 socioeconomic factors interact, change, and evolve over time.” Study of interactions between
290 humans and the earth system has also received considerable support from various agencies such
291 as the National Science Foundation, the National Institute of Food and Agriculture and the U.S.
292 Geological Survey, including targeted programs (e.g., NSF Solicitation 13-535, “Water
293 Sustainability and Climate”, NSF Solicitation 18-545, “Coupled Human-Natural Systems and
294 Innovations in Food-Energy-Water Systems. Thus, evolving the Budyko framework to better
295 understand how land-surface responses vary under natural and human-altered landscapes will
296 also support various ongoing studies on the impact of humans on hydrological systems.

297 Enhancements to the Budyko framework will also support other ongoing activities that
298 focus on improving the ability to predict the hydrologic behavior of natural and ungauged
299 watersheds. As competition for water increases, there has been increasing attention placed on the
300 need for water availability information at ungauged locations, even in regions where water has
301 not been considered in the past to be a limited resource. For these reasons, the decade from 2003
302 to 2012 was recognized by the International Association of Hydrological Sciences as the
303 Prediction in Ungauged Basins (PUB) Decade (Sivapalan et al., 2003). Blöschl et al. (2013;
304 tables A7-A10) showed that several methods to predict streamflow in ungauged watersheds have
305 been proposed; however, no one method has been universally accepted or demonstrated to work
306 in all hydrologic settings. Other studies have evaluated predictability in ungauged basins at the
307 global scale (Hrachowitz et al., 2013). Since the Budyko framework provides an approach for
308 improving our understanding of ungauged basins, there is potential cross-fertilization in various
309 ongoing studies for evaluating the extended Budyko framework and associated datasets, in order
310 to support a range of global- and continental-scale hydrologic initiatives.

311 Another exciting aspect of the extension of the Budyko framework for considering
312 anthropogenic influences, concerns the development of hydrologic indicators for a wide range of
313 purposes related to human impacts. These include watershed classification, environmental
314 permitting and a variety of water management activities. There is a continuing need for new
315 hydrologic indicators, founded in the science of hydrology, for the purpose of watershed
316 classification--as expressed so nicely by Wagener et al. (2007). The development and plotting of
317 nondimensional variables, analogous to the nondimensional variables proposed in Figure 1, has a
318 close association with the development of nondimensional hydroclimatologic indicators for both
319 natural (Weiskel et al. 2014) and human-dominated (Weiskel et al., 2007) hydrologic systems.

320 For example, the aridity and runoff ratios, two commonly used nondimensional hydroclimatic
321 indicators arise naturally from the Budyko framework for natural watersheds. We anticipate that
322 a wide range of new hydrologic indicators, founded on the science of hydrology, yet useful for
323 water management and watershed classification, will arise from the types of studies envisioned
324 here which extend the Budyko framework to accommodate anthropogenic influences.

325 One significant challenge in evaluating the Budyko framework under human-altered
326 landscapes would be the availability of data on hydroclimate, storages, and human influences,
327 including water withdrawal and land use changes, reservoir storages and releases, at different
328 spatio-temporal scales. For example, the monthly change in total water storage is a critical
329 component of accurate assessments of land-surface fluxes, particularly in regions of high
330 anthropogenic influence where storage is impacted by pumping of groundwater resources, or
331 conversion of surface water to evapotranspiration through diversion for irrigation. In addition to
332 the tremendous challenges relating to data availability, there is the open research question of how
333 we can capture the complexity of human-water systems with a low dimensional, parsimonious
334 modeling approach. One approach involves a gradual refinement of model features – a top-down
335 approach – as needed (Zhang et al., 2008; Sivapalan et al., 2003). Another strategy involves
336 development of critical data sets, and the subsequent addition of model features as the spatio-
337 temporal scale of the data permits. Such a global synthesis effort will require global-scale data
338 sets from a variety of sources, including remotely sensed data. The selection of appropriate data
339 at this scale presents challenges in balancing spatial resolution, uncertain data accuracy, and
340 consistency among the considered data sets. Findings from another synthesis study titled,
341 “Water Availability for Ungaged Basins” revealed that, as various hydrologic modeling
342 communities converge towards continental-domain hydrologic models, these communities will

343 encounter similar limitations and challenges (Archfield et al., 2015). It is our hope and
344 contention that the Budyko framework can provide a unifying perspective for bridging gaps in
345 hydrologic data availability and model resolution over a wide range of spatial and temporal
346 scales. As shown in this opinion article, the Budyko framework can also be modified beyond the
347 traditional long-term water balance, to improve basic understanding of how the land-surface
348 response—runoff and evapotranspiration—vary across natural and human-altered landscapes.

349

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355

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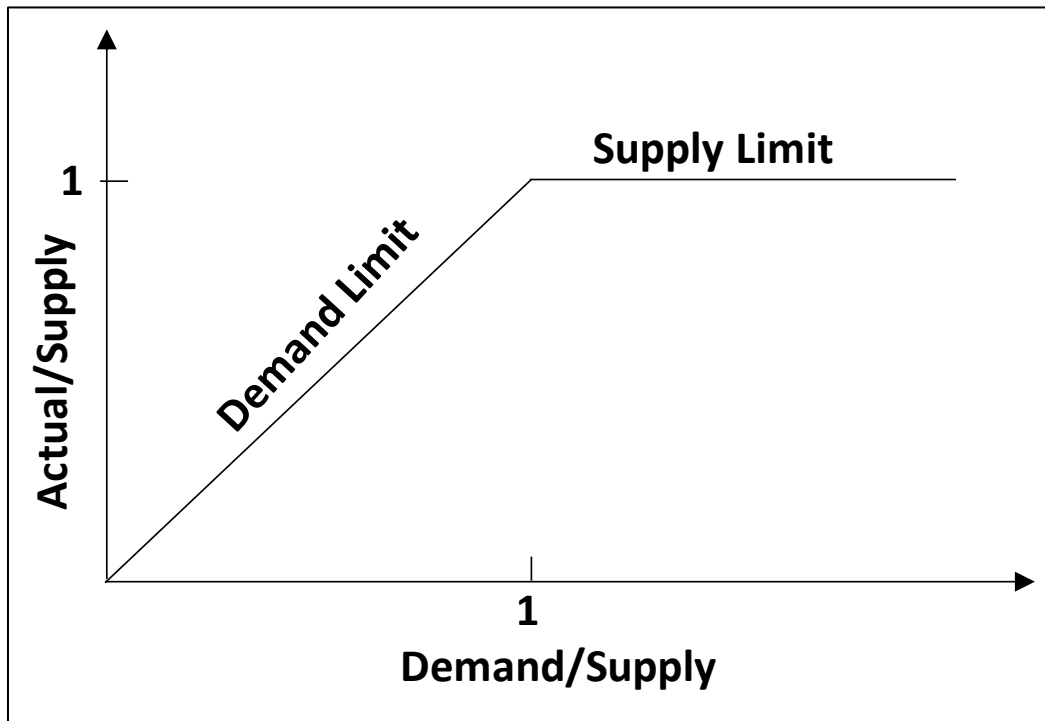
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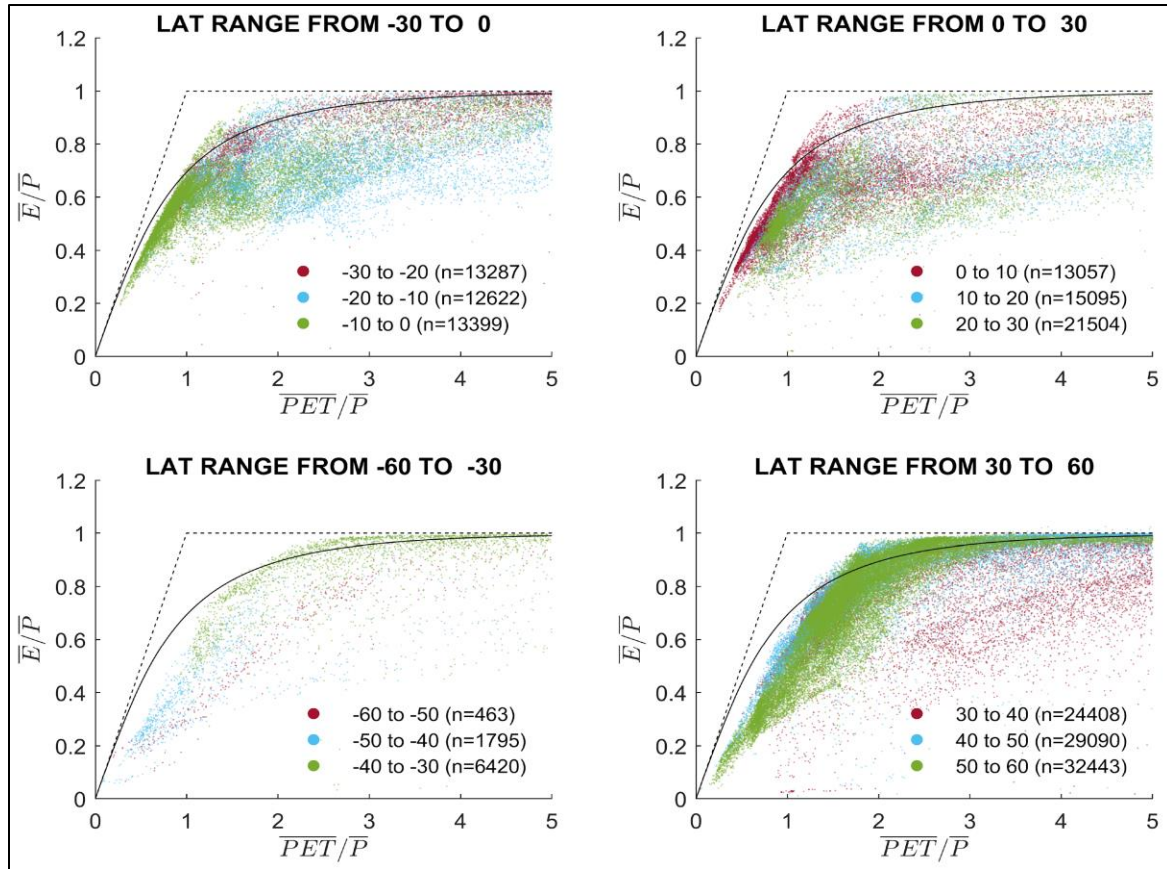
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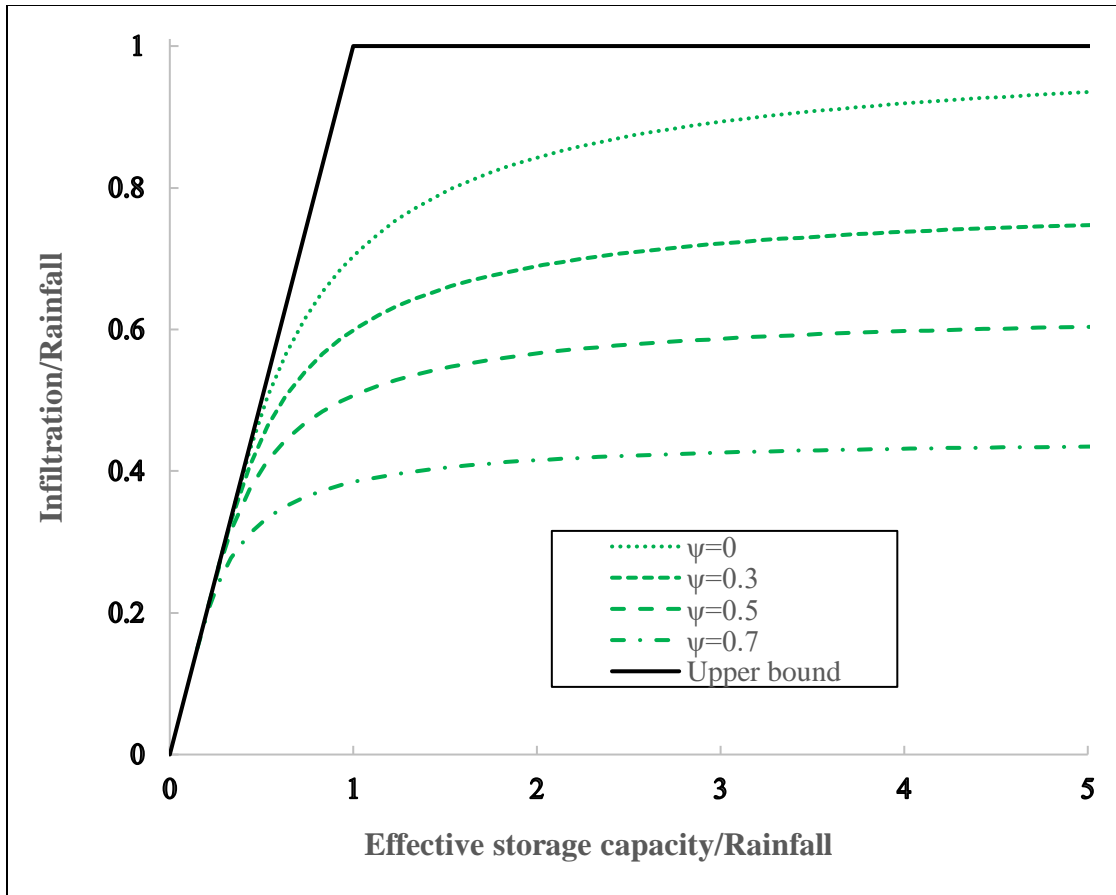
539 Figure 1: An overview of the Budyko supply and demand framework for understanding the land-
540 surface flux response (actual) over natural and human-altered watersheds. The “limits” concept
541 as suggested by Budyko (1958) quantifies the actual response (Y axis) based on the physical
542 demand-to-supply ratio of energy/moisture over the control volume or the watershed.
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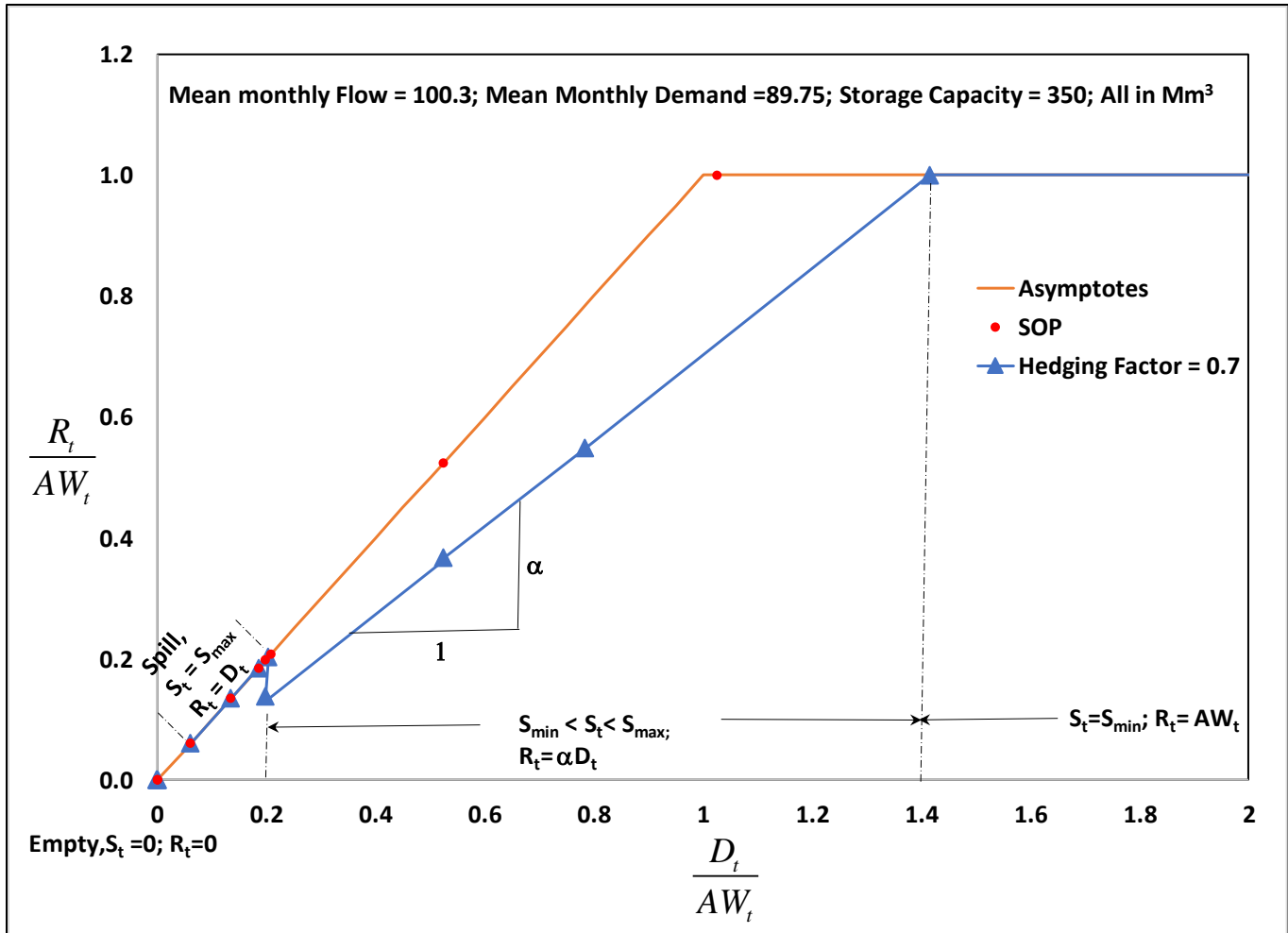
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546 Figure 2: The traditional Budyko framework for long-term water balance along with the
 547 asymptotes and the Budyko curve ($\overline{ET} / \overline{P} = \left[(1 - \exp(-\overline{PET} / \overline{P})) * \overline{PET} / \overline{P} * \tanh(\overline{PET} / \overline{P}) \right]^{0.5}$). The ratio
 548 of mean annual potential evapotranspiration (\overline{PET} , demand) to mean annual precipitation (\overline{P} ,
 549 supply) explains the ratio of mean annual evapotranspiration (\overline{ET} , actual) and \overline{P} , and the data
 550 points are from GLDAS-2 estimates at the pixel level (0.25°) for the period 1948-2010 over the
 551 northern (top row, 0° - 30° and 30° - 60° latitudes) and southern (bottom row, 0° - 30° and -30° to -
 552 60° latitudes) hemispheres.

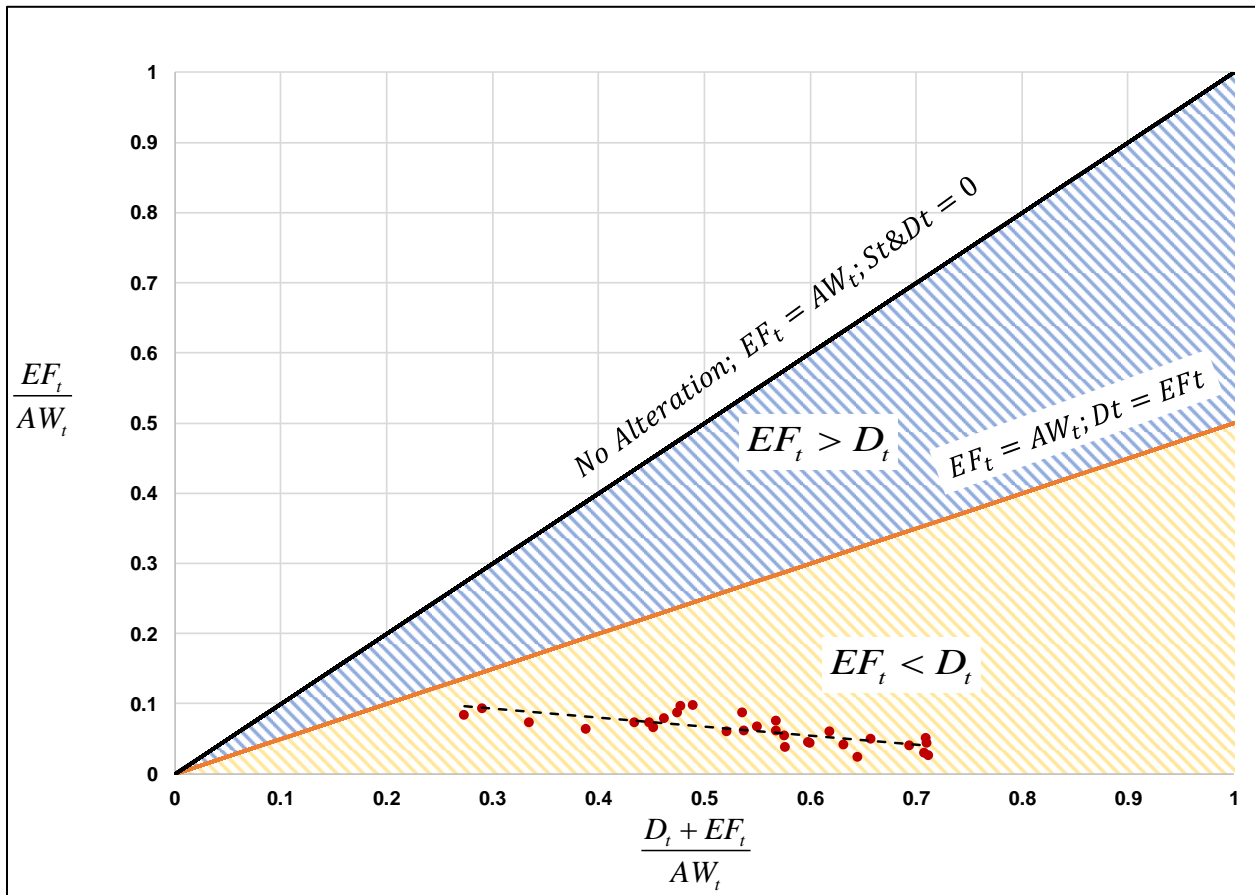
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 556 Figure 3: Modeling infiltration in the Budyko supply and demand framework: the ratio of
 557 infiltration (actual) and rainfall depth is a function of the ratio of infiltration capacity (demand)
 558 and rainfall depth (supply) as well as the initial soil moisture condition represented by the degree
 559 of saturation (ψ) (Reproduced from Wang (2018)).
 560



561
 562 Figure 4: Modeling hedging policy of reservoir operations in the Budyko supply and demand
 563 framework. The standard operating policy (SOP) is corresponding to the asymptotes. For the
 564 hedging rule, delivery or release is “actual”, available water is “supply”, and human use is
 565 “demand”. For demonstration purposes, a linear function is assumed for the hedging rule (i.e.,
 566 $R_t = \alpha D_t$). The storage conditions are indicated for the hedging policy alone.
 567



569

570 Figure 5: Modeling synthesizing flow alteration in the Budyko supply and demand framework:
 571 the ratio of environmental flow (“actual”) and the available water as a function of the ratio the
 572 total demand for human and environmental flow (“demand”) and the available water (“supply”).
 573 Annual flows from Falls Lake (red dots) show human withdrawal for water supply is more than
 574 the downstream environmental flow release.

575