



Time-varying copula and design life level-based nonstationary risk
analysis of extreme rainfall events

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Key points:

- The time-varying GEV model and copula models are developed for marginal and multivariate frequency analysis, respectively.
- A design life level-based risk analysis is implemented for hydraulic engineering practice.
- A systematic risk analysis incorporating nonstationarity is emphasized in comparison with stationary models.



Abstract: Due to global climate change and urbanization, more attention has been paid to decipher the nonstationary multivariate risk analysis from the perspective of probability distribution establishment. Because of the climate change, the exceedance probability belonging to a certain extreme rainfall event would not be time invariant any more, which impedes the widely-used return period method for the usual hydrological and hydraulic engineering practice, hence calling for a time dependent method. In this study, a multivariate nonstationary risk analysis of annual extreme rainfall events, extracted from daily precipitation data observed at six meteorological stations in Haihe River basin, China, was done in three phases: (1) Several statistical tests, such as Ljung-Box test, and univariate and multivariate Mann-Kendall and Pettist tests were applied to both the marginal distributions and the dependence structures to decipher different forms of nonstationarity; (2) Time-dependent Archimedean and elliptical copulas combined with the Generalized Extreme Value (GEV) distribution were adopted to model the distribution structure from marginal and dependence angles; (3) A design life level-based (DLL-based) risk analysis associated with Kendall's joint return period (JRP_{ken}) and AND's joint return period (JRP_{and}) methods was done to compare stationary and nonstationary models. Results showed DLL-based risk analysis through the JRP_{ken} method exhibited more sensitivity to the nonstationarity of marginal and bivariate distribution models than that through the JRP_{and} method.

Key words: multivariate risk analysis; time-varying copula; design life level; nonstationarity; Kendall's joint return period



1 1. Introduction

2 Due to climate change and increasing urbanization, heavy rains-induced floods
3 have occurred more frequently all over the world in recent decades, which is becoming
4 a major deterrent to the sustainable development of social economy (Mishra and Singh,
5 2009; Donat et al., 2016; Ali and Mishra, 2018). Various kinds of social activities, such
6 as infrastructure constructions, agricultural irrigation and ecological maintenance
7 would be influenced by hydrometeorological extreme events. A systematic risk analysis
8 of these extreme events would provide sufficient strategies for decision makers.

9 A multitude of studies have addressed the effect of climate change and
10 urbanization on hydrological design to alleviate associated risks. Traditional
11 hydrological frequency analysis or risk analysis is based on the stationary assumption,
12 which recommends that environmental impact indexes, such as climatic factors and
13 land use rate, have a constant mechanism or pattern that affects hydrological variables
14 all the time (Madsen et al., 2017; Milly et al., 2015). The feasibility of hydrological
15 frequency and risk analysis based on stationary assumptions is being challenged
16 because of the multiple effects of climate change, urbanization, and heat island effects.
17 Accordingly, water authorities should amend the present planning, design and
18 management strategies to develop nonstationary distribution models based on the
19 signals of climate change. Therefore, it is urgent to develop an efficient and systematic
20 risk analysis approach from time dependent side to serve for hydraulic design of
21 hydrological infrastructures to cope with the effect of climate change.



22 In recent years, nonstationary hydrological frequency analysis has received a great
23 deal of attention because of increasing attention to climate change (Chen and Sun, 2017;
24 Call et al., 2017; Ghanbari et al. 2019). The time-varying moment approach is widely
25 used to involve time variant probabilistic parameters for mimicking the changing
26 behavior of extreme hydrometeorological variables. Nonstationarity modeling of
27 probability distribution has been conducted for univariate cases in recent years (Zhang
28 et al., 2015; Ganguli and Coulibaly, 2017; Agilan and Umamahesh, 2018).

29 Du et al. (2015) modelled nonstationary low-flow series in Weihe River basin,
30 China, based on the Generalized Additive Models in Location, Scale and Shape
31 (GAMLSS) framework. Results showed that inappropriately estimated statistical
32 parameters would lead to the overstatement of risk corresponding to a low-flow event.
33 Gu et al. (2016) incorporated time, climate indices, precipitation, and temperature into
34 the GAMLSS model to detect nonstationarity in flood frequency. For the univariate
35 case, nonstationary risk analysis, based on the time-varying moment approach, can be
36 decomposed into four steps: (1) Descriptive and exploratory monitoring of hydrological
37 sequences and monitoring of outliers; (2) implementation of the stationarity hypothesis
38 to verify the nonstationarity of hydrological series; (3) development of a hydrological
39 frequency analysis model and estimation of model parameters using different covariates;
40 and (4) risk assessment based on the selected frequency model.

41 The above studies were conducted under nonstationary conditions for univariate
42 cases, while it is known that natural hydrometeorological extreme events are



43 multivariate, characterized by multi-attribute properties which can be statistically
44 correlated. For instance, floods are characterized by volume, peak, and duration, while
45 extreme rainfall events have the attributes of duration, intensity, total amount. As a
46 result, univariate nonstationary risk analysis cannot fully encompass the dependence
47 structure between hydrological attributes. It is therefore desirable to develop a
48 multivariate model to simulate the probabilistic behavior of two or more properties.
49 Copulas, a useful tool for modelling the structure of dependence between hydrological
50 variables regardless of the types of marginal distributions, have been widely used for
51 multivariate frequency analysis of rainfall extreme events (Zhang and Singh, 2007; Kao
52 and Govindaraju, 2008; Rauf and Zeepongsekul, 2014; Vandenberghe et al., 2010);
53 droughts (De Michele et al., 2013; Serinaldi et al., 2009; Shiau, 2006; Song and Singh,
54 2010; Wong et al., 2010); floods (Grimaldi and Serinaldi, 2006; Zhang and Singh, 2006).
55 However, these studies assumed a time invariant dependence pattern, ignoring the
56 influence of climate change and hence did not consider the impact of nonstationarity
57 on the dependence structure.

58 Recently, studies on multivariate distribution fitting have addressed the superiority
59 of dynamic copula-based method to model the nonstationary dependence structure,
60 which are generally caused by complex environment and rapid urbanization (Milly et
61 al., 2015). Former studies have detected nonstationarity in dependence structures (Liu
62 et al. 2017; Assia et al., 2014; Yilmaz and Perera, 2014). Chebana et al. (2013) argued
63 that it was necessary to determine a multivariate distribution model quantifying the



64 time-varying dependence structure of various kinds of hydrological variables. Bender
65 et al. (2014) used a bivariate nonstationary multivariate model with a 50-year moving
66 time window to investigate the time-dependent behavior in bivariate case. Their results
67 showed that the joint probability varied significantly over time for different non-
68 stationary models. Jiang et al. (2015) also did a multivariate risk analysis using the
69 time-varying copula method incorporating time and reservoir index as covariates for
70 low-flow series extracted from two neighboring observed stations.

71 Traditional solutions of hydrological extreme events involve return period-based
72 methods, which are usually calculated as the inverse of annual exceedance probability
73 for a given magnitude under stationary conditions in a univariate case. In a multivariate
74 case, the univariate return period can be extended to joint return periods of hydrological
75 variables. There are three kinds of joint return period methods to quantify the
76 exceedance probability of a multivariate extreme event: the OR method that at least one
77 extreme attribute is larger than the specified threshold; the AND method that all the
78 attributes are larger than the specified threshold; and the Kendall method that the
79 univariate value derived from the Kendall distribution function according to a specified
80 value (Jiang et al., 2015; Salvadori and Michele, 2010; Salvadori et al., 2013). While
81 non-stationary distribution models provide flexibility to analyze the variability of a
82 hydrological variable, they are also incongruent with many of the traditional metrics
83 used in water resources planning. For example, the development of drainage standards
84 are vulnerable to the standard of extreme rainfall return period, which means drainage



85 facilities have been designed to withstand the extreme rainfall event of a specified
86 return period. The multivariate hydrologic and hydraulic design can be influenced by
87 the existence of nonstationarity in both the marginal and joint distributions. The
88 exceedance probability of a given extreme event would be different from year to year,
89 leading to a nonconstant and non-unique value of the conventional return period. Thus,
90 the notion of static return period of an extreme rainfall event (e.g., 100-year extreme
91 rainfall event, 200-year extreme rainfall event) is no longer reliable for hydraulic design
92 under nonstationary conditions (Salas and Obeysekera, 2014; Yan et al., 2017). As a
93 result, Rootzén and Katz (2013) first mentioned the concept of design life level (DLL)
94 to quantify the risk of a given extreme rainfall magnitude over the hydrological
95 structure's life time (Note that following the idea of Rootzén and Katz (2013) we regard
96 the term hydrological risk as the possibility of a certain extreme event occurring and
97 not as a quantification of expected losses). It is a logical extension to handle the
98 nonstationarity of the concept of "risk of failure" (Jakob, 2013), which is more
99 frequently used to quantify the risk of hydrologic extremes under stationarity. Read and
100 Vogel (2015) extended the DLL method to average annual reliability (AAR) method to
101 estimate the hydrologic design value considering nonstationarity. In general, these risk-
102 based methods can provide similar results of hydraulic design for hydrological
103 infrastructure (Yan et al., 2017). However, the cases of multivariate hydrologic designs,
104 especially under nonstationary conditions using the time-varying copula, and design
105 life level-based risk methods have great potentiality in future studies.



106 Therefore, the objective of this study is to do risk analysis of multivariate extreme
107 rainfall events involving the following steps. First, a series of statistical tests, such as
108 Ljung-Box test, and univariate and multivariate Mann-Kendall and Pettist tests are used
109 for both the marginal distributions and the dependence structures to determine different
110 forms of nonstationarity (sudden jump, periodicity, and trend). Second, a nonstationary
111 multivariate probability distribution is developed using a time-varying GEV and
112 copula-based model, which can encompass the nonstationarities probably existed in
113 marginal and joint distributions. Finally, design life level-based risk analysis is
114 extended to multivariate cases through Kendall's joint return period and AND's joint
115 return period methods. In this paper, we investigated two kinds of extreme rainfall
116 attributes: (1) annual extreme rainfall volume (P_s : Annual total precipitation of the
117 daily precipitation more than the 95th percentile threshold) and intensity (Im : Annual
118 maximum daily precipitation), through the nonstationary multivariate risk analysis
119 method. The remainder of this paper is organized as follows. The next section presents
120 the methodology adopted in this study. Section 3 discusses the results of proposed
121 model applied to Haihe River basin, China. Section 4 presents the final conclusion
122 through the proposed model.

123 **2. Methodology**

124 Copulas are tools to build multivariate distribution models of dependence
125 structures between random variables regardless of their marginal distribution types.
126 Detailed information about copulas can be found in Nelson (2007). The present copula-



127 based methods to solve the multivariate risk analysis mostly adopt static parameters for
128 whether the marginal distribution or joint distribution. The changing climate has led to
129 nonstationarity of individual hydrological series or the dependence between
130 hydrological variables. To realize this situation, a time-varying copula-based model can
131 describe the time dependent characteristics for dependence structure of hydrological
132 variables, as inspired by Patton (2006) from the financial field.

133 Let (x, y) represent a hydrological pair. The joint probability distribution of
134 multivariates through time-varying copula model can then be presented as:

$$135 \quad F_{X,Y}(x, y) = C[F_X(x|\theta_X^t), F_Y(y|\theta_Y^t)|\theta_C^t] = C(u, v|\theta_C^t) \quad (1)$$

136 where $C(\cdot)$ denotes the copula function; $F_{X,Y}(\cdot)$ denotes the joint function; $F_X(\cdot)$
137 and $F_Y(\cdot)$ represent the marginal functions of hydrological variable (Ps and Im in this
138 study); θ_X^t and θ_Y^t represent the time-varying marginal distribution parameters; θ_C^t is
139 the dynamic copula parameter which is a linear function of time; and u and v are the
140 marginal probabilities in the time-varying copula in the hypercube unit.

141 In the framework of multivariate risk analysis (**Figure 1**), the property of
142 nonstationarity can be determined not only by one or two marginal variables but also
143 in the dependence structure or vice versa. It is however possible that the nonstationary
144 behavior may exist in both the marginal and joint distribution function. To determine
145 the nonstationarity (mutation, cyclicity and trend) in the synthesized extreme rainfall
146 attribute series, statistical tests, such as Ljung-Box test, and univariate and multivariate
147 Mann-Kendall and Pettist tests are used for both the marginals and the dependence



148 structure. Details of these tests can be found in the references due to (Serinaldi and
149 Kilsby, 2016; Chebana et al., 2013; Rizzo and Székely, 2010).

150 As shown in **Figure 1**, the time-varying copula-based risk analysis model can be
151 decomposed into three main phases: (1) detection of nonstationarity in the marginal
152 variables and dependence structure through a series of nonparametric tests; (2)
153 estimation of the time-varying parameter for the marginal and joint probability
154 distributions; and (3) joint return period and risk analysis by design life level-based risk
155 methodology from the perspectives of Kendall's and AND's return period methods
156 (detailed information can be found in Section 2.3).

157

158 Insert **Figure 1** here.

159

160 *2.1. Time-varying marginal distribution*

161 In this part, the Generalized Extreme Value (GEV) distribution was used to
162 establish time-varying marginal distribution model for the extreme rainfall attributes
163 because it is a good aggregation of the Gumbel, Fréchet, and Weibull distributions and
164 is especially suitable for extreme data sets (Cheng and AghaKouchak, 2014). Let $F(x)$
165 be the cumulative probability distribution function (CDF) of the quantity of interest, P_s
166 or I_m , in this study. The GEV distribution consists of three control parameters, the
167 location, the scale, and the shape, which describe mean value of the sample series,
168 amplitude near the location, and the tail of the distribution, respectively. The cumulative



169 distribution of GEV model under stationary conditions can be expressed as follows:

$$170 \quad F(x) = \begin{cases} \exp \left\{ - \left[1 + \kappa \left(\frac{x - \mu}{\sigma} \right)_+ \right]^{\frac{1}{\kappa}} \right\} & \text{if } \kappa \neq 0 \\ \exp \left\{ - \exp \left(- \frac{x - \mu}{\sigma} \right)_+ \right\} & \text{if } \kappa \rightarrow 0 \end{cases} \quad (2)$$

171 where $z_+ = \max\{y, 0\}$ and

172 $x \in [(\mu - \sigma)/\kappa, +\infty)$ when $\kappa > 0$,

173 $x \in (-\infty, (\mu - \sigma)/\kappa]$ when $\kappa < 0$, and

174 $x \in (-\infty, +\infty)$ when $\kappa = 0$.

175 where μ denotes the location parameter, σ is the scale parameter and κ is the shape
176 parameter. In this study, two kinds of nonstationary GEV models (GEVns-1 and
177 GEVns-2) are developed with the shape parameter being constant. It should be
178 emphasized that modelling the time variance in shape parameter needs long-term
179 observations, which are often not available in practice (Cheng et al., 2014). GEVns-1
180 model considers the time-varying characteristic of the location parameter only, while
181 GEVns-2 model incorporates the time varying features of both location and scale
182 parameter. These two nonstationary models regard significant trends as a linear function
183 of time (in years):

$$184 \quad \mu(t) = \mu_0 + \mu_1 t \quad (3)$$

$$185 \quad \sigma(t) = \exp(\sigma_0 + \sigma_1 t) \quad (4)$$

187 where the scale parameter is always positive throughout, it is usually calculated on the
188 basis of a log link function.

189 In this study, the Bayesian method through the Markov chain Monte Carlo



190 (MCMC) approach (Cheng et al., 2014) was used to estimate the nonstationary GEV
191 model. Simultaneously, the Deviance Information Criterion (*DIC*) and Bayes factors
192 (*BF*) for different stationary and nonstationary models were calculated to select the best
193 fitted marginal model. The minimum *DIC* value yielded the best performance, while
194 *BF* smaller than 1 indicated the best fitting.

195 2.2. Time-varying Copula

196 To model the dependence structure between annual total precipitation (*P_s*) and
197 annual daily maximum precipitation (*Im*) under nonstationary conditions, a time-
198 varying copula was developed. In multivariate hydrological frequency analysis, two
199 kinds of copulas, named elliptical and Archimedean copulas are widely used in
200 hydrological applications. In this study, time-varying elliptical copulas, Student t (St)
201 copula, as well as the widely-used time-varying Archimedean copulas, time-varying
202 Clayton, Gumbel and Frank copula, were selected as candidate models to simulate the
203 time-varying dependence between two extreme rainfall attributes. The Gaussian copula
204 was not used in this study because of its deficiency in describing dependencies of
205 extremes (Renard and Lang, 2007).

206 The copula parameter θ_c^t can be assumed as a linear function of the time (“year”
207 in this study) and can be defined as follows:

$$208 \quad \theta_c^t = \begin{cases} \exp(\beta_o + \beta_1 t) & \theta_c > 0 \\ \beta_o + \beta_1 t & \theta_c \in R \end{cases} \quad (5)$$

209 where $\theta_c > 0$ denotes the Student t (St), Clayton and Gumbel copula, while $\theta_c \in R$
210 represents the Frank copula.



211 The maximum pseudo-likelihood (MPL) method was adopted to estimate the time-
212 varying copula parameter (Genest et al., 1995). The Corrected Akaike Information
213 Criterion (AICc; Hurvich and Tsai, 1989) was employed to make a goodness-of-fit,
214 which is a modified version of AIC for small samples. Obviously, the presence of
215 nonstationarity in the copula parameter was determined by comparison of the AICc
216 value.

217 2.3. Joint return period and risk analysis based on KEN's and AND's methods

218 For hydrological management, engineering administrators focus more on the
219 return period and risk of failure during the design life of hydraulic structures (Condon
220 et al., 2015). Inspired by design life level (DLL) method to present the risk proposed
221 by Rootzén and Katz (2013), we would like to expand the DLL-based risk to the
222 multivariate case.

223 Let $F(X)$ be the cumulative probability distribution function (CDF) of the quantity
224 of interest, in this study, maximum daily precipitation in a year (Im). Conventionally,
225 the T -year return level for certain daily precipitation x_T is equal to the $(1-1/T)$ -th
226 quantile of the marginal distribution of Im (The probability distribution is the same for
227 all years in a stationary situation.). Equivalently, on average, one out of T years has at
228 least one daily rainfall that exceeds x_T , so that $T(1 - F(x_T)) = 1$ (Serinaldi and Kilsby,
229 2015), and the probability of annual maximum daily rainfall exceeds x_T is $1/T$.

230 Then, the hydrological risk R (i.e. risk of failure) of a certain hydraulic structure
231 for a design life of n years can be expressed as the probability that at least one rainfall



232 extreme exceeds the design level x_T in a period of n years. Under stationary conditions,
233 the probability of annual maximum daily rainfall exceeding x_T in every year is the
234 same as $1/T$. In a univariate context, hydrological stationary risk can be defined as
235 (Fernandez and Salas, 1999; Serinaldi and Kilsby, 2015):

$$236 \quad R_s = 1 - F(x_T)^n = 1 - (1 - 1/T)^n \quad (6)$$

237 Considering time-varying exceedance probabilities, the probability of annual
238 maximum daily rainfall exceeding x_T in each year is different. So here we use $F_t(x_T)$
239 to represent the probability of daily rainfall exceeding design level x_T in the t -th year.
240 So the design life level-based nonstationary risk for the univariate case is:

$$241 \quad R_{ns} = 1 - \prod_{t=1}^n F_t(x_T) \quad (7)$$

242 From the perspective of bivariate case, the joint return period (JRP) of extreme
243 rainfall events can be calculated through three methods in a stationary situation
244 (Salvadori et al., 2011). They are AND method corresponding to the probability of
245 $P(X \geq x \cap Y \geq y)$, OR method corresponding to $P(X \geq x \cup Y \geq y)$, and Kendall
246 return period method (KEN). Details of the Kendall return period can be found in
247 Salvadori and De Michele (2004). Since the AND method is widely used and the
248 Kendall method is of great potentiality, we expanded the AND method and the Kendall
249 return period method to the nonstationary case here. Let JRP_{s-and} and JRP_{s-ken}
250 represent the three types of return period in the stationary case; they can be calculated
251 as follows:



$$252 \quad JRP_{s-and} = \frac{1}{P\{X \geq x \cap Y \geq y\}} = \frac{1}{1 - F_X(x) - F_Y(y) + C[F_X(x), F_Y(y)]} \quad (8)$$

$$253 \quad JRP_{s-ken} = \frac{1}{P\{C[F_X(x), F_Y(y)] \geq p_{ken}\}} = \frac{1}{1 - K_c(p_{ken})} \quad (9)$$

254 where $K_c(\cdot)$ is the Kendall distribution function which can be defined as:

$$255 \quad K_c(p_{ken}) = P\{C[F_X(x), F_Y(y)] \leq p_{ken}\} \quad (10)$$

256 Here, $F_X(x)$ and $F_Y(y)$ are the marginal cumulative probability distribution functions
 257 (CDF) for P_s and I_m , respectively, while $C[F_X(x), F_Y(y)]$ is the bivariate copula
 258 function connecting these two extreme attributes. p_{ken} is just the critical probability
 259 level corresponding to $K_c(p_{ken})$.

260 Similar to the JRPs of extreme rainfall events under stationary case, the JPRs of
 261 AND and KEN in nonstationary situations can be achieved by:

$$262 \quad JRP_{ns-and} = \frac{1}{1 - F_X(x|\theta_X^t) - F_Y(y|\theta_Y^t) + C[F_X(x), F_Y(y)|\theta_C^t]} \quad (11)$$

$$263 \quad JRP_{ns-ken} = \frac{1}{1 - K_c^t(p_{ken})} \quad (12)$$

264 where θ_X^t , θ_Y^t and θ_C^t represent the time variant parameters of the marginal and
 265 copula distributions; and $K_c^t(p_{ken})$ is the time-varying Kendall distribution function
 266 corresponding to the time-varying copula.

267 Multivariate extreme value analysis should be focused on the most likely extreme
 268 event with the largest copula density. The most likely event at the T_0 -year level can be
 269 calculated as (Graler et al., 2013):

$$270 \quad (u_m, v_m) = \underset{T_0}{\operatorname{argmax}} c(u, v) \quad (13)$$

271 The most likely design combinations (x_m, y_m) can be computed according to the
 272 inverse of marginal cumulative distribution function:



273 $x_m = F_X^{-1}(u_m)$ and $y_m = F_Y^{-1}(v_m)$ (14)

274 where u, v are the marginal distribution functions of X and Y . Let two pairs of extreme
 275 rainfall attributes $(x_{m_1}, y_{m_1})_{T_0^{and}}$ and $(x_{m_2}, y_{m_2})_{T_0^{ken}}$ be the most likely design
 276 combinations of P_s and Im at the T_0 -year level for JRP_{s-and} and JRP_{s-ken} . Similar
 277 to the nonstationary risk calculation in the univariate case, the hydrological
 278 nonstationary DLL-based risk in the bivariate case can be calculated from two
 279 circumstances:

280 $R_{ns-and} = 1 - \prod_{t=1}^n \{F_X(x_{m_1} | \theta_X^t) + F_Y(y_{m_1} | \theta_Y^t) - C[F_X(x_{m_1}), F_Y(y_{m_1}) | \theta_C^t]\}$ (15)

281 $R_{ns-ken} = 1 - \prod_{t=1}^n K_c^t(p_{ken})_{(x_{m_2}, y_{m_2})}$ (16)

282 where R_{ns-and} , R_{ns-ken} indicate the nonstationary risk for a design life level of n
 283 years in the bivariate case corresponding to two types of joint return period. The
 284 stationary risk can be calculated in the same way with marginal and copula distribution
 285 parameters being constant.

286 In this study, comparison of hydrological risk for the bivariate case between
 287 stationary and nonstationary models can be quantified by the risk changing rate $\Delta R_{T_0}^n$
 288 which can be calculated as:

289 $\Delta R_{T_0}^n = \frac{1}{n} \sum_{i=1}^n \frac{|R_i^{ns} - R_i^s|}{R_i^s}$ (17)

290 where R_i^{ns} and R_i^s are nonstationary risk and stationary risk of a certain hydraulic
 291 structure for a design life of i years. $\Delta R_{T_0}^n$ helps quantify the difference in risk between
 292 stationary and nonstationary models.



293 **3. Application**

294 *3.1. Study area and data collection*

295 The area selected for the study is Haihe River basin, China, which belongs to the
296 temperate East Asian monsoon climate zone (**Figure 2**). In summer, heavy rains take
297 place and temperature and humidity are high caused by marine air masses. The annual
298 rainfall has a great spatial and temporal variability across the basin due to the
299 inconsistency of intensity, retreat time and influence of the Pacific subtropical high over
300 the years. Natural disasters, such as urban floods and mountain torrents induced by
301 extreme rainfall events in the basin have caused huge losses to the social economy and
302 people's lives and property, and have been highly valued by decision-making authorities.
303 As a result, time-varying copula-based multivariate risk analysis of this basin is
304 conducive to providing reliable strategies and alternative options for water resources
305 risk-based decision making.

306 Daily rainfall data from Haihe River basin observed at Wutaishan, Fengning,
307 Zhangjiakou, Beijing, Tianjin, and Nangon were analyzed for the proposed
308 nonstationary model. Detailed information on these six gauges is presented in Table 1.
309 According to various data ranges shown in **Table 1**, the rainfall series from 1958-2017
310 was selected as the final version.

311

312 Insert **Figure 2** Here.

313 Insert **Table 1** Here.



314

315 *4.2. Preprocessing Analysis*

316 Before developing a nonstationary frequency analysis model, it is essential to
317 examine nonstationarities of extreme precipitation attributes (P_s and I_m) as well as the
318 structure of dependence between these two attributes. A series of statistical tests (i.e.
319 Ljung-Box test, univariate and multivariate Man-Kendall tests, and univariate and
320 multivariate Pettitt tests) were performed to detect the nonstationarity in extreme
321 precipitation time series. Trends in the time series can be evaluated using various tests
322 (Lima et al., 2016; Yilmaz et al., 2017; Sarhadi and Soulis, 2017). **Table 2** shows results
323 of tests detecting nonstationarity, while **Figure 3** shows the spatial distribution of trends
324 and change points for two attributes of rainfall extremes (P_s and I_m) as well as the
325 dependence structure between them. First, time series of these two rainfall extremes (P_s
326 and I_m) for all 6 stations can pass the Ljung-Box test with 20 lags (p.value>0.05 in
327 **Table 2**). Extreme observations are mutually independent with no serial autocorrelation,
328 so it is appropriate to apply the standardized Mann-Kendall test to evaluate the
329 statistical significance of trend without any modification (Serinaldi and Kilsby, 2016).

330 As shown in **Figure 3**, concurrences of univariate and bivariate trends, the
331 nonstationarities in rainfall extremes can be detected at several stations (stations 2, 3,
332 and 4). Station 1 exhibits a significant nonstationarity for extreme attribute P_s , while
333 extreme attribute I_m and dependence structure show an insignificant decreasing trend.
334 On the other hand, stations 5 and 6 show a weak decreasing trend. The above tests



335 totally recommend the presence of nonstationarities in extreme series as well as the
336 dependence pattern across three out of 6 sites. According to Porporato and Ridolfi
337 (1998), an insignificant trend should not be ignored because of its effect on the results
338 of hydrological risk analysis. Hence, even if precipitation extremes at a certain station
339 may recommend statistically weak trends, both the nonstationary and stationary models
340 are established for each station in the following section.

341

342 Insert **Table 2** Here.

343 Insert **Figure 3** Here.

344

345 *4.3. Marginal distribution fitting*

346 The nonstationarity can appear either in univariate variables or in dependence
347 structure in the multivariate framework (Bender et al., 2014). Results of trend and
348 change point tests performed in Section 4.1 pointed out the necessity to take the
349 nonstationarity of marginal distributions into consideration. In this study, the
350 Generalized Extreme Value (GEV) distribution which is a good hybrid of the Gumbel,
351 Fréchet, and Weibull distributions fits the block or annual maximum time series better
352 (Cheng et al., 2014). **Table 3(a)-(b)** shows performances of nonstationary vs. stationary
353 models for these six stations. The location parameter (μ) and scale parameter (σ) are
354 regarded as time variant, while the shape parameter κ is time invariant; it should be
355 noted that modeling of time-varying κ requires a sufficiently long record of



356 observations (Cheng et al., 2014). Despite the exception of Im for station 4, the shape
357 parameter κ for most fitted models was in the interval of $[-0.3, 0.3]$ which is in
358 accordance with the previous study (Martins and Stedinger, 2000; Ganguli and
359 Coulibaly, 2017). The best fitted model was selected by performing the minimum DIC
360 criterion combined with the Bayes factor (BF) test. For instance, the GEV_{ns-2} model
361 (nonstationary GEV model with time varying location and scale parameters) was the
362 best selected model for the extreme attribute Im extracted from station 1. That was
363 because the BF values of GEV_{ns-2} and GEV_{ns-1} were both smaller than 1 which meant
364 that these two nonstationary models passed the BF test. Then, the best fitted
365 nonstationary model GEV_{ns-2} for Im of station 1 was achieved following the DIC test.
366 Similarly, the best fitted marginal distribution of two extreme rainfall attributes for all
367 these six stations was selected. Except for stations 4 and 5, the best distributions for the
368 other stations were parallel for nonstationarity tests shown in Section 4.1.

369

370 Insert **Table 3(a)-(b)** here.

371

372 4.4. Copula fitting

373 Elliptical and Archimedean (Clayton, Gumbel, and Frank) copulas have been
374 widely applied in hydrological practice. In this study, time-varying elliptical copulas,
375 Student t (St) copula, as well as Clayton, Gumbel and Frank copulas were selected as
376 alternative models to simulate the dependence structures of extreme attributes. The



377 Gaussian copula was not used in this study because of its deficiency in describing
378 dependencies of extremes (Renard and Lang, 2007). Once a marginal distribution was
379 estimated based on test statistics, the dependence structure for I_m and P_s was described
380 by the time-varying or time invariant copula functions. **Table 4(a)-(b)** illustrates the
381 results of best fitted copula, based on the minimum AICc and maximum log-likelihood
382 value (LL). The time-varying Student t (St) copula exhibited the best performance
383 among the eight candidate copulas (four stationary copulas as well as the corresponding
384 nonstationary copulas) for stations 1, 2, 3, 4, and 6, while the stationary St copula was
385 the best one for station 5 which meant that results of dependence structure modelling
386 for station 5 did not indicate any nonstationarity signal which was reasonable, according
387 to multivariate MK and Pettit tests of station 5 (**Table 2**). Contrary to station 5, the
388 nonstationary St copula fitted better than did the stationary model for stations 1 and 6
389 which was not in accordance with the nonstationarity tests for these two stations (**Table**
390 **2**). Based on the above results, an insignificant trend or weak change point would lead
391 to a nonstationary probability function of dependence patterns to some extent
392 (Porporato and Ridolfi, 1998) which should be dealt with cautiously.

393 With the best fitted marginal distributions and the best copula, the quantiles of
394 extreme rainfall attributes (P_s and I_m) were derived from the pseudo-observations
395 ranging from 0 to 1 in order to provide a benchmark for return period and risk analysis
396 for hydrological and hydraulic design. The method of analysis is presented in the
397 following section.



398

399 Insert **Table 4(a)-(b)** here.

400

401 4.5. *Nonstationary return period and risk analysis for univariate and bivariate cases*

402 (1) Univariate return period: Once parameters of the best fitted models for
403 univariate and bivariate cases have been estimated, the extreme rainfall quantiles for
404 certain return levels (T) can be simulated. In this section, return period and risk analysis
405 was performed by comparing stationary and nonstationary models. The estimated
406 rainfall quantiles (P_s and I_m) versus time in the univariate case are shown in Figures 4
407 for the six stations of Haihe River. I_m and P_s for stations 4 and 6 are not provided,
408 because the best marginal model for the extreme attributes of these two stations was the
409 stationary GEV model (**Table 3**). In the case of P_s for station 1 shown in **Figure 3**, a
410 100-year P_s quantile under stationary circumstances (GEV_s model with dashed red line
411 in **Figure 3**) (355 mm) corresponded to a 35-year P_s under nonstationary conditions
412 (GEV_{ns-2}) in the year 1960 and a 60-year P_s in the year 1970. In other words, an
413 exceedance probability of 0.01 increased to 0.028 and 0.017. On the other hand, the
414 return period associated with a given quantile decreased from 1960 to 2020 for I_m of
415 station 4 and P_s of station 5, while the return period increased for extreme attributes of
416 other stations. Interestingly, the temporal variability between different stations
417 corresponding to the best selected nonstationary model exhibited a significant
418 difference. For example, the nonstationary GEV_{ns-2} model fitted to P_s of stations 1, 2,



419 3, and 5 showed a significant upward or downward trend of extreme quantiles with
420 years. Compared to the temporal variability, the attributes of stations 4 and 6 with
421 GEV_{ns-1} model showed a weaker trend which demonstrated the time variability of scale
422 parameter of the GEV distribution. Finally, it is noteworthy that nonstationary isolines
423 were over stationary isolines for P_s of stations 1 and 5 as well as Im of station 3 (marked
424 in blue star in Figure 4) which meant that the stationary model would underestimate the
425 risk for a certain return period.

426 Insert **Figure 3** here.

427 Insert **Figure 4** here.

428

429 (2) Joint return period (JRP) based on AND and KEN method:

430 After the nonstationary copula and GEV distribution models were selected
431 according to several goodness-of-fit tests, the design values characterizing annual
432 extreme rainfall events were determined through Kendall's (JRP_{ken}) or AND's joint
433 return period (JRP_{and}) expressed by Equations (8)-(10). Although the copula model for
434 station 5 was stationary, it was regarded as a nonstationary model because of the
435 marginal nonstationary GEV_{ns-2} model for P_s or Im , which existed at other stations.

436 Since events with lower exceedance probabilities are of interest for hydrological
437 practice and the joint return period of 50-year level is able to minimize the uncertainties
438 of extrapolation. In this study we focused typically on events with a joint return period
439 of $JRP_{ken} (JRP_{and})=50$ which means the exceedance probability was equal to 0.02.



440 **Figure 5** shows isolines of Kendall return period and AND-based return period at the
441 50-year level for both stationary and nonstationary models. Since the number of isolines
442 of the nonstationary model were 60 (sample size) which might show certain
443 overlapping areas in the isoline map, four isolines corresponding to the year 1960, 1980,
444 2000, and 2020 are presented for simplicity. For comparison, the $JRP_{ken}(JRP_{and})$ -
445 isolines derived from the corresponding stationary model, which was composed of
446 stationary GEV and stationary copula model, are also shown. Observations belonging
447 to each station are also presented. Although the plots for all the years are not shown,
448 the variability of design quantiles over time showed the nonstationary behavior of the
449 dependence structure.

450 From **Figure 5**, JRP_{ken} was larger than JRP_{and} for the dependence structure of
451 the same extreme rainfall attributes, which was caused by Kendall's return period
452 method of generating the same dangerous region, regardless of different realizations
453 (Salvadori et al. 2011). Focusing on JRP_{ken} and JRP_{and} for station 1, the design
454 values of P_s varied over time, while the design values of I_m did not vary with time.
455 From the horizontal direction, both the JRP_{ken} -isolines and JRP_{and} -isolines exhibited
456 a left-moving trend, recommending a descending trend for P_s . The maximum P_s values
457 for the year 1960 were measured as 341.6 mm and 371.5 mm corresponding to JRP_{ken}
458 and $JRP_{and}=50$, respectively, while 246.4 mm and 264.8 mm is calculated as the
459 minimum marginal values. The gap between them reached 100 mm. On the other hand,
460 none of the JRP_{ken} and JRP_{and} -isolines exhibited a variation trend of I_m values for



461 station 1 from the vertical perspective, which can be attributed to the stationary GEV
462 model for Im of station 1 (**Table 3**). Due to sudden changes in the magnitudes of
463 marginal values, the Kendall isolines also crossed each other. In a similar way, the
464 nonstationary behavior of the variables was detected from the variation of design values
465 of extreme attributes at the other five stations compared to the isolines derived from the
466 stationary model (denoted as black line). **Figure 5**. It is noteworthy that the stationary
467 copula model for station 5 also exhibited the variation of design values derived from
468 both JRP_{ken} and JRP_{and} . A weak variation of 22.3 mm for Ps and 4.9 mm for Im was
469 detected because of the corresponding nonstationary GEV model (**Table 3**).

470

471 Insert **Figure 5** here.

472

473 (3) Univariate risk

474 The hydrological risk of a certain design extreme attribute quantile x_{T_0} can be
475 computed using Equations. (6) and (7) on the basis of the initial return period T_0 and
476 design life n . The best marginal distribution model for Im of station 1 as well as Ps of
477 stations 4 and 6 were the stationary GEV model, so these three scenarios were not taken
478 into consideration in this part. Except for the results of Ps of station 5, the risk results
479 of extreme attributes of other five stations were very similar (**Figure 6**). Here, we
480 considered the risk result of the attribute Im of station 2 for detailed illustration.
481 Comparing the risk of stationary and nonstationary models, a definite conclusion can



482 be addressed: risk can increase from the stationary condition to nonstationary condition.
483 For example, when $T_0=50$ and $n=20$, the risks for the stationary and nonstationary
484 conditions were 33.24% and 46.8%, respectively. That is to say an unjustified
485 assumption would lead to an overestimation of the risk under a certain return period
486 and design life. For Im of station 2, the nonstationary risk was higher than the stationary
487 risk when $n \leq 53$. On the other hand, the nonstationary risk was smaller than the
488 stationary risk when $n \geq 53$. This conclusion can be detected from the Ps of station 1.

489 Once T_0 was decided, the risk changing rate was calculated by equation (17).
490 Here, T_0 was set as 50 for illustration. For attribute Im , the risk changing rate $\Delta R_{T_0}^n$
491 corresponding to $T_0=50$ was 45.93%, 5.31%, 18.25%, 39.44% and 37.10% for stations
492 2, 3, 4, 5 and 6, respectively. For attribute Ps , $\Delta R_{T_0}^n$ was 61.26%, 22.47%, 59.51% and
493 20.53% for stations 1, 2, 3 and 5. Generally, Im of station 2 and Ps of station 1 should
494 be paid more attention with the highest risk changing rate in hydrological practice.

495

496 Insert **Figures 6** here.

497

498 (3) Bivariate risk based on JRP_{ken} and JRP_{and}

499 The hydrological nonstationary risk in the bivariate case cannot be calculated until
500 the most likely event at T_0 -year level is generated. In this part, we first focused on the
501 development of the most likely design events where the joint probability density
502 functions had their maximum values on the 50-year level. **Figures 7(a)** and **(b)** illustrate



503 the time dependent development of both variables P_s (upper panel) and Im (lower panel)
504 through the JRP_{ken} and JRP_{and} methods. The attribute P_s for stations 1, 2, and 3
505 showed a positive trend, while attribute Im for stations 4 and 5 exhibited a negative
506 trend through the JRP_{ken} and JRP_{and} methods. On the other hand, the trend of the
507 design value of Im was not significant.

508 The hydrological nonstationary risk based on JRP_{ken} and JRP_{and} was
509 computed using equations (15)-(16). **Figures 8(a) and (b)** show the nonstationary risk
510 R of the most likely design combinations of P_s and Im at the T_0 -year level. The risk
511 results of extreme attributes of stations 1, 3, 4, and 6 were very similar, while the results
512 of stations 2 and 5 exhibited a similar pattern. For stations 1, 3, 4, and 6, the risk
513 increased with design life n under both stationary and nonstationary conditions, but for
514 any T_0 , the nonstationary risk was higher than stationary risk from both JRP_{ken} and
515 JRP_{and} methods. For stations 2 and 5, the nonstationary risk was lower than stationary
516 risk from both the JRP_{ken} and JRP_{and} methods. The corresponding changing rate to
517 quantify the differences in hydrological risk for the bivariate case between stationary
518 and nonstationary models was also calculated by equation (17). Whether it was
519 calculated through the JRP_{ken} and JRP_{and} methods, the changing risk rate increased
520 as T_0 increased, which meant that nonstationarity influenced the risk of lower
521 exceedance probabilities more than that of higher exceedance probabilities.

522 Since a 50-year level with lower exceedance probabilities (0.02) is of great interest
523 in hydrological practice and necessary to control the uncertainties of extrapolation



524 (Bender et al., 2014), in this part, we focused on the risk changing rate under the 50
525 year-level for each station. For JRP_{and} , the risk changing rate $\Delta R_{T_0}^n$ corresponding to
526 $T_0=50$ was 41.83%, 7.96%, 24.27%, 6.94%, 9.21%, and 70.63% for stations 1, 2, 3, 4,
527 5, and 6, respectively. For JRP_{ken} , the risk changing rate $\Delta R_{T_0}^n$ corresponding to
528 $T_0=50$ was 59.93%, 8.44%, 44.19%, 10.69%, 11.96%, and 75.29% for stations 1, 2, 3,
529 4, 5, and 6, respectively. According to the above results of risk changing rate, changing
530 risk rates based on the JRP_{ken} method were higher than those through the JRP_{and}
531 method, which indicated that the JRP_{ken} -based risk was more sensitive to the
532 nonstationarity of marginal and bivariate distribution models.

533

534 Insert **Figures 7(a)-(b)** here.

535 Insert **Figures 8(a)-(b)** here.

536

537 4.6. Further discussion

538 Based on the above analysis, the nonstationary risk analysis over extreme rainfall
539 events using the time-varying GEV and copula-based distribution models were
540 distinguished from those where the partial assumption of stationarity was employed
541 (**Figures 4-8**). These results showed the significance of considering nonstationarity
542 when calculating return period and hydrological risk both in univariate and bivariate
543 cases. There were also certain differences between the results using Kendall's joint
544 return period method and AND's return period method. Im of station 2 and Ps of station



545 1 should be concerned with the highest risk changing rate from the perspective of
546 univariate case, while the dependence structure of station 6 should be paid more
547 attention with the highest risk changing rate from the perspective of bivariate case.

548 According to the results performed by the proposed time-varying models, the
549 following points should be emphasized:

550 ✓ It is necessary to use statistical tests, such as the Ljung-Box test, univariate
551 and multivariate Mann-Kendall test, and univariate and multivariate Pettitt
552 test to evaluate nonstationarities of extreme rainfall attributes (P_s and I_m) as
553 well as the dependence between these two attributes. These two attributes
554 corresponding to six stations showed no serial correlation which rationalized
555 the implementation of traditional multivariate Mann-Kandall tests without
556 any modification.

557 ✓ Nonstationarity in the dependence structure and marginal variable was non-
558 ignorable. The nonstationary (time-varying) GEV and copula-based model
559 not only addressed the abrupt changes and significant trends existed in the
560 marginal variables, but also evaluated the dependence of multivariate
561 hydrological series, which led to the reliable estimation of hydraulic design
562 quantiles.

563 ✓ The traditional hydrological risk under nonstationary conditions in the
564 univariate case was expanded to the bivariate case through the Kendall joint
565 return period method and AND return period method. According to the return



566 period analysis in the univariate case, the scale parameter of the nonstationary
567 GEV distribution demonstrated a significant time variability for uncertainty.
568 The joint return period and risk analysis also showed that the JRP_{ken} -based
569 risk was more sensitive to the nonstationarity of marginal and bivariate
570 distribution models.

571 Moreover, the two indexes used in this study, revealing the characteristics of
572 extreme rainfall events, i.e., Ps and Im , representing rainfall volume and intensity,
573 respectively were extracted from observed daily precipitation datasets. Risk
574 analysis based on these two attributes helped understand extreme rainfall patterns,
575 especially storm events lasting several days, which would be devastating to urban
576 infrastructure and farmlands. In addition, the duration which is another meaningful
577 extreme rainfall attribute should also be incorporated into multivariate risk
578 analysis.

579 **5. Conclusions**

580 In this paper, a nonstationary risk analysis through the time-varying Generalized
581 Extreme Value (GEV) and copula-based distribution model is performed over the
582 extreme rainfall events in Haihe River Basin. The time-dependent copula and GEV
583 models are applied to these two attributes (Ps and Im) extracted from daily rainfall data
584 of six stations in Haihe River basin, China. Nonstationarity and trends in the attribute
585 series were investigated through multivariate Mann-Kendall test and multivariate
586 Pettist test. The best nonstationary GEV model was selected for the attribute of each



587 station through the minimum DIC criterion combined with the Bayes factor (BF) test,
588 while the best-fitted time-varying copula was selected through the minimum Corrected
589 Akaike Information Criterion ($AICc$). Based on frequency analysis by the Kendall joint
590 return period method and the AND return period method, the design values of the two
591 indexes were computed and shown by the JRP_{ken} -isolines and JRP_{and} -isolines. The
592 extended bivariate nonstationary DLL-based risk was calculated through the estimated
593 most likely event (combinations of P_s and I_m) to quantify the risk of each station under
594 nonstationary conditions. Analysis of extreme rainfall occurrence risk based on the
595 observed index series demonstrated that station 6 should be paid more attention with
596 the highest risk changing rate. The following conclusions can be drawn from this study:

- 597 1. A 100-year P_s quantile under stationary conditions (355 mm) can correspond to a
598 35-year P_s under nonstationary conditions. In other words, an exceedance probability
599 of 0.01 can increase to 0.028 and 0.017. On the other hand, the return period associated
600 with a given quantile can decrease for I_m of some stations but can increase for other
601 stations.
- 602 2. The stationary model would underestimate the risk for a certain return period.
- 603 3. From the marginal return period to the joint return period, there can be a significant
604 upward or downward trend in extreme quantiles in the univariate case which can
605 change into a weak trend in the joint return period.
- 606 4. Nonstationarity influences the risk of lower exceedance probabilities more than that
607 of higher exceedance probabilities.



608 5. Changing risk rates based on the JRP_{ken} are higher than those based on the
609 JRP_{and} method, which indicated that the JRP_{ken} -based risk is more sensitive to
610 the nonstationarity of marginal and bivariate distribution models.

611 This study emphasizes the significance of incorporating nonstationarity into
612 multivariate risk analysis through the investigation of univariate and multivariate trend
613 and change points in the attribute series. The Kendall return period is justified as more
614 practical method for hydraulic design than the AND return period method according to
615 the calculation of the design quantiles for the extreme rainfall. The extended bivariate
616 nonstationary DLL-based risk method was applied to both stationary and nonstationary
617 conditions.

618

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Table 1. Information on meteorological gauges of Haihe River basin

	Station ID	Station name	Location		Data range
			Longitude	Latitude	
1	53588	Wutaishan	113 °32'	39 °02'	1952-2017
2	54308	Fengning	116 °32'	41 °12'	1957-2017
3	54401	Zhangjiakou	115 °11'	40 °50'	1958-2017
4	54511	Beijing	116 °19'	39 °57'	1958-2017
5	54527	Tianjin	117 °10'	39 °06'	1958-2017
6	54705	Nangon	115 °23'	37 °22'	1956-2017



Table 2. Detection of trends and change points in extreme rainfall attributes collected from six stations

Station No.	Attribute	L-jung-Box Test		Univariate MK		Univariate Pettitt Test		Multivariate MK		Multivariate Pettitt Test	
		p.value	p.value	p.value	Z	p.value	statistics	p.value	Z statistics	p.value	Z statistics
1	Ps	0.927	0.063	0.063	-1.856*	0.048	0.450	-0.755	0.089	0.089	0.089
	Im	0.798	0.674	0.708	0.421	0.708	0.450	-0.755	0.089	0.089	0.089
2	Ps	0.307	0.105	0.089	-1.652*	0.089	0.098	-1.645*	0.095	0.095	0.095
	Im	0.462	0.132	0.091	-1.649*	0.091	0.098	-1.645*	0.095	0.095	0.095
3	Ps	0.986	0.345	0.131	-0.944	0.131	0.099*	-1.641	0.078	0.078	0.078
	Im	0.575	0.051	0.012	-1.963**	0.012	0.099*	-1.641	0.078	0.078	0.078
4	Ps	0.981	0.072	0.098	-1.799*	0.098	0.055	-1.922*	0.039	0.039	0.039
	Im	0.971	0.054	0.089	-1.926*	0.089	0.055	-1.922*	0.039	0.039	0.039
5	Ps	0.051	0.524	0.801	-0.638	0.801	0.606	0.516	0.589	0.589	0.589
	Im	0.747	0.214	0.678	-1.244	0.678	0.606	0.516	0.589	0.589	0.589
6	Ps	0.815	0.226	0.454	-1.212	0.454	0.115	-1.575	0.359	0.359	0.359
	Im	0.923	0.067	0.024	-1.831*	0.024	0.115	-1.575	0.359	0.359	0.359

**and * represent statistically significant at 5% and 10% significance levels; The standardized Mann-Kendall test statistic (Z statistics) in univariate and multivariate cases indicates positive (negative) with an increasing (decreasing) trend, and statistically significant at 5% and 10% significance levels when $|Z| > 1.96$ and $|Z| > 1.64$ respectively; Change point tests in univariate and multivariate cases are performed at 10% significance level.



Table 3(a) Performance of stationary and nonstationary GEV models fitted to for the marginal distribution corresponding to each attribute (stations 1-3)

Station	Attribute	Model	μ	σ	κ	DIC	BF
1	<i>P_S</i>	GEV _s	144.37	45.32	0.058	1944.47	-
		GEV _{ns-1}	147.82+0.0014t	46.67	0.033	1943.03	0.9999
		GEV_{ns-2}	151.85-0.00042t	exp(3.92-0.000038t)	0.019	1942.08	0.9991
	<i>I_m</i>	GEV_s	50.64	18.13	0.012	1593.71	-
		GEV _{ns-1}	46.69+0.0013t	17.81	0.035	1597.58	1.0011
		GEV _{ns-2}	51.7-0.00062t	exp(2.96-0.000037t)	0.019	1596.53	1.004
2	<i>P_S</i>	GEV _s	106.25	28.84	-0.12	1740.92	-
		GEV _{ns-1}	104.08-0.00018t	28.83	-0.12	1742.66	0.998
		GEV_{ns-2}	99.92+0.0028t	exp(3.25+0.000057t)	-0.13	1742.29	0.996
	<i>I_m</i>	GEV _s	40.71	14.01	-0.053	1497.533	-
		GEV_{ns-1}	42.02+0.00026t	14.30	-0.099	1493.47	0.997
		GEV _{ns-2}	42.90-0.00031t	exp(2.39+0.00012t)	-0.082	1492.82	0.996
3	<i>P_S</i>	GEV _s	86.92	26.92	-0.077	1724.35	-
		GEV _{ns-1}	88.48+0.000096t	27.16	-0.085	1724.306	1.0002
		GEV_{ns-2}	85.78+0.0014t	exp(3.57-0.00014t)	-0.091	1723.76	0.999
	<i>I_m</i>	GEV _s	34.31	11.00	0.13	1455.2	-
		GEV_{ns-1}	37.51-0.00067t	11.38	0.10	1448.41	0.995
		GEV _{ns-2}	41.76-0.0030t	exp(2.37+0.000019t)	0.11	1449.82	0.996

Table 3(b) Performance of stationary and nonstationary GEV models fitted for marginal distribution corresponding to each attribute (Station 2-6)

Station	Attribute	Model	μ	σ	κ	DIC	BF
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4	<i>P_S</i>	GEV _s	141.69	55.87	0.19	2034.265	-
		GEV _{ns-1}	144.73-0.0016t	55.99	0.19	2036.30	1.0014
		GEV _{ns-2}	139.80-0.00014t	exp(3.88+0.000060t)	0.21	2037.05	1.0019
	<i>I_m</i>	GEV _s	56.36	22.29	0.34	1751.59	-
		GEV _{ns-1}	65.73-0.0035t	23.17	0.31	1749.31	0.9995
		GEV _{ns-2}	60.33-0.00055t	exp(3.10+0.000026t)	0.31	1749.48	0.9998
5	<i>P_S</i>	GEV _s	160.34	56.07	-0.21	1963.47	
		GEV _{ns-1}	148.22+0.0016t	54.74	-0.18	1962.35	1.0002
		GEV _{ns-2}	150.25+0.0011t	exp(4.07-0.000038t)	-0.19	1962.12	0.9999
	<i>I_m</i>	GEV _s	67.46	26.37	0.072	1752.36	
		GEV _{ns-1}	77.72-0.0041t	26.88	0.047	1747.30	0.9976
		GEV _{ns-2}	70.19-0.00072t	exp(3.66-0.00019t)	0.067	1749.53	0.9979
6	<i>P_S</i>	GEV _s	137.56	57.44	-0.24	1965.79	-
		GEV _{ns-1}	133.69+0.00084t	58.31	-0.25	1968.09	1.0017
		GEV _{ns-2}	138.26-0.00087t	exp(4.25-0.0001t)	-0.25	1967.70	1.0015
	<i>I_m</i>	GEV _s	63.02	29.43	0.033	1767.72	-
		GEV _{ns-1}	69.26-0.0053t	27.93	0.059	1767.24	1.0001
		GEV _{ns-2}	57.05+0.00083t	exp(3.45-0.000071t)	0.072	1766.92	0.9999



Table 4(a) Performance of stationary and nonstationary copula models fitted to the dependence structure of two attributes (stations 1-3)

Station	Model	Copula	θ	LL	AIC_c	
1	S ^a	St ^c	0.8066	28.83	-55.59	
		Clayton	1.9725	23.81	-45.55	
		Gumbel	2.2677	25.08	-48.08	
		Frank	8.2702	24.88	-46.55	
	NS ^b	St	exp(0.2998-0.0002t)	31.42	-58.71	
		Clayton	exp(-4.8096+0.0023t)	22.73	-43.39	
		Gumbel	exp(-5.268+0.0256t)	23.63	-46.49	
		Frank	7.677+0.00028t	21.14	-42.45	
	2	S	St	0.9000	51.55	-101.05
			Clayton	3.7056	42.85	-83.63
<u>Gumbel</u>			<u>3.4208</u>	<u>52.42</u>	<u>-102.77</u>	
Frank			13.592	45.88	-89.68	
NS		St	exp(0.4247-0.0003t)	53.46	-102.786	
		Clayton	exp(-3.289+0.0017t)	41.38	-80.59	
		Gumbel	exp(-4.344+0.0356t)	45.57	-88.57	
		Frank	10.592+0.0018t	46.07	-93.09	
3		S	<u>St</u>	<u>0.8491</u>	<u>38.25</u>	<u>-74.44</u>
			Clayton	3.0403	35.57	-69.06
	Gumbel		2.6696	33.21	-64.35	
	Frank		9.4212	34.89	-67.15	
	NS	St	exp(-1.6982+0.0008t)	41.49	-78.84	
		Clayton	NaN	NaN	NaN	
		Gumbel	NaN	NaN	NaN	
		Frank	11.385-0.0011t	35.18	-68.59	

^aS is stationary copula model; ^bNS represents the time-varying copula model; ^cSt represents Student's t copula.



Table 4(b) Performance of stationary and nonstationary copula models fitted to the dependence structure of two attributes (stations 2-6)

Station	Model	Copula	θ	LL	$AICc$
4	S	St	0.9001	43.10	-87.13
		Clayton	2.944	36.68	-71.30
		<u>Gumbel</u>	<u>3.273</u>	<u>46.82</u>	<u>-91.57</u>
		Frank	11.462	44.68	-87.67
	NS	St	exp(-1.6101+0.0008t)	51.79	-99.46
		Clayton	exp(-3.369+0.0027t)	38.97	-75.89
		Gumbel	exp(-4.564+0.0289t)	44.68	-85.98
		Frank	12.398-0.00081t	44.68	-87.67
5	S	St	0.9000	47.977	-93.886
		Clayton	3.2678	40.62	-79.17
		Gumbel	3.1054	42.94	-83.82
		Frank	11.023	44.68	-87.67
	NS	St	exp(0.5868-0.0003t)	46.00	-91.359
		Clayton	exp(-2.987+0.0037t)	42.56	-82.69
		Gumbel	exp(-3.898+0.0252t)	41.69	-81.59
		Frank	12.589-0.00036t	42.57	-83.79
6	S	St	0.8889	43.85	-83.55
		<u>Clayton</u>	<u>3.4955</u>	<u>44.77</u>	<u>-84.47</u>
		Gumbel	2.8959	37.84	-73.61
		Frank	11.227	41.89	-81.79
	NS	St	exp(0.7846-0.0005t)	46.02	-87.90
		Clayton	exp(-3.876+0.071t)	39.59	-78.89
		Gumbel	<i>NaN</i>	<i>NaN</i>	<i>NaN</i>
		Frank	11.462	44.68	-87.67

Inf: infinite number, or out of scope of computation

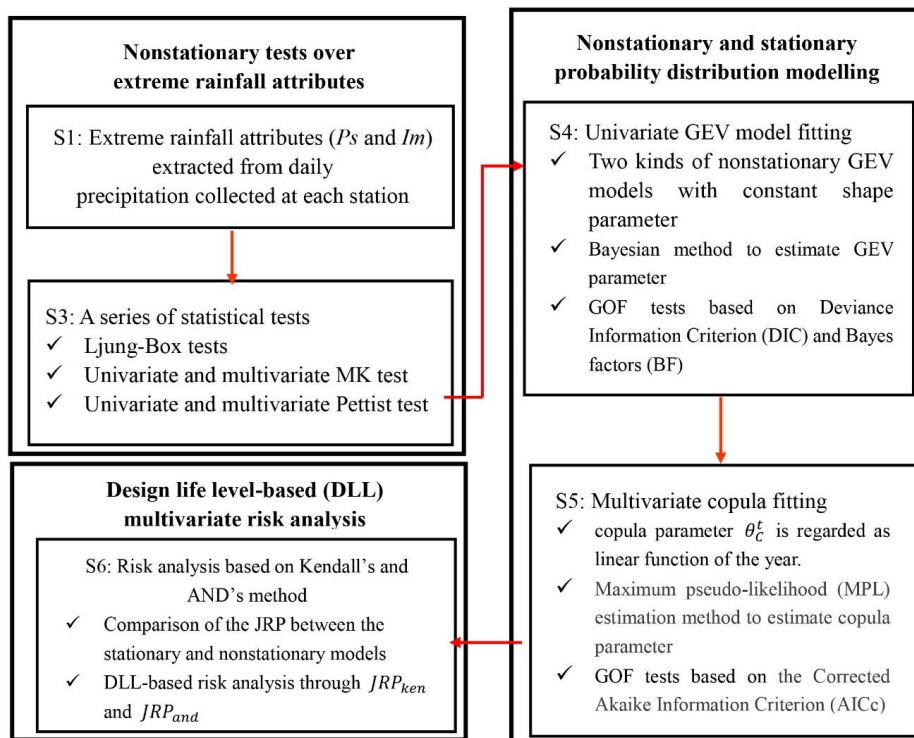


Figure 1. Flowchart of this study

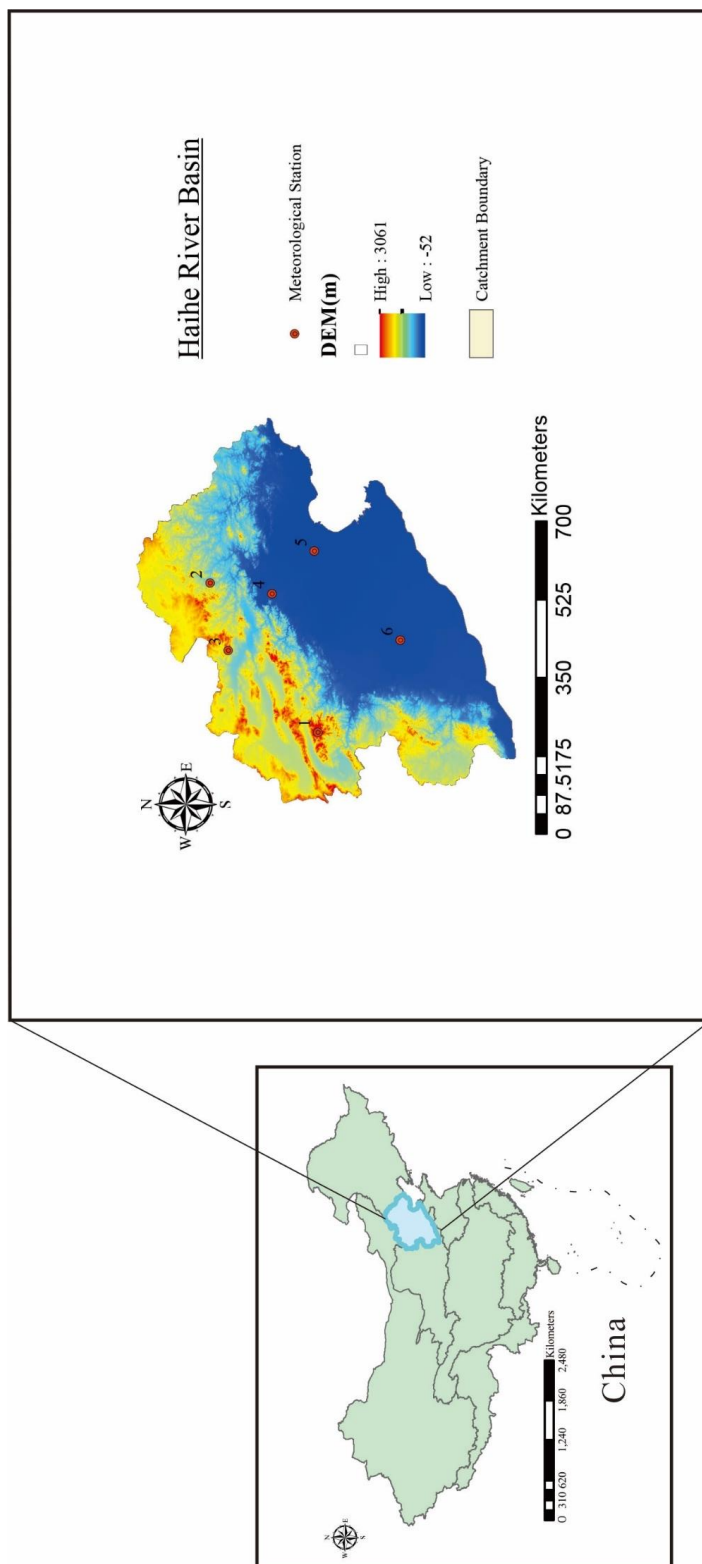


Figure 2. Selected meteorological stations in Haihe River basin

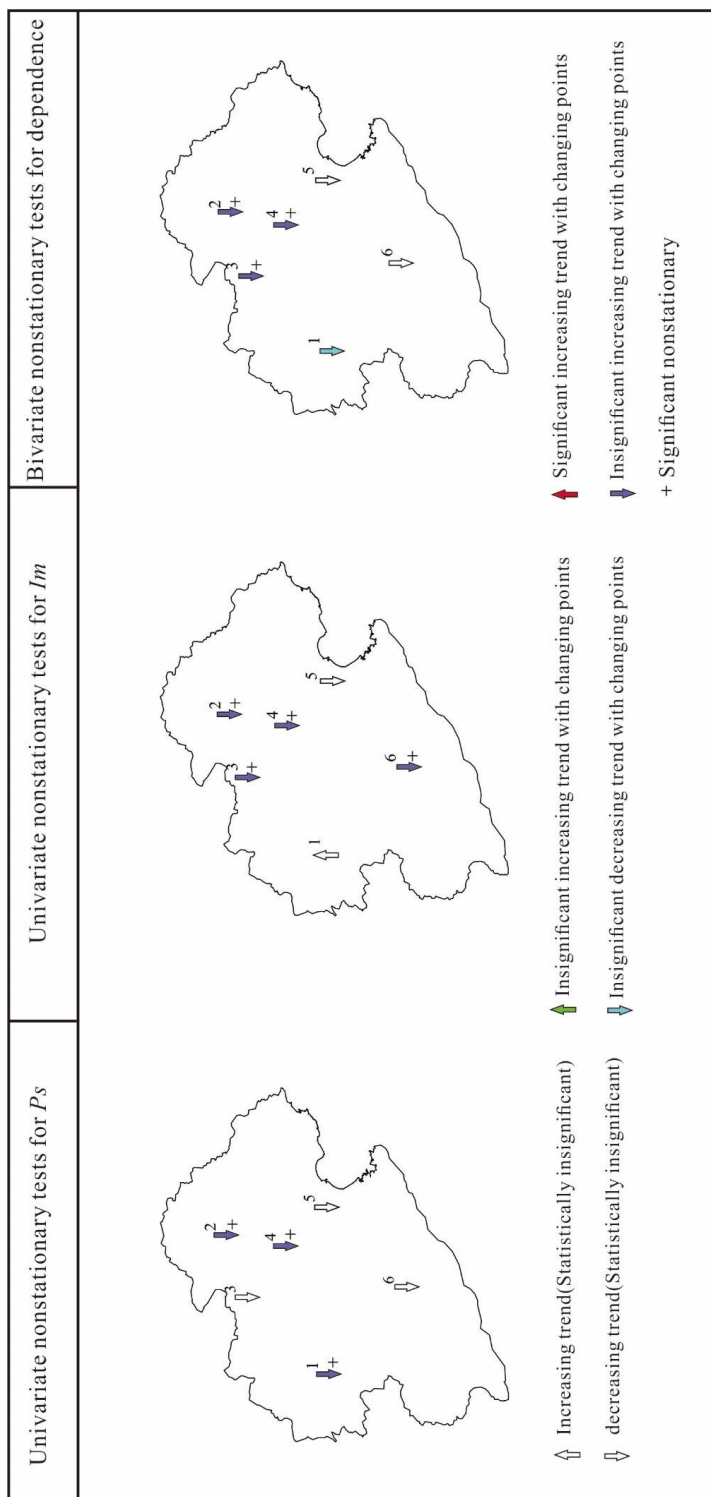


Figure 3. Spatial distribution of the trend and changing points according to univariate and multivariate Mann-Kandall and Pettitt tests shown in Table 2; All above tests are performed at a significance level of 10%, i.e. p value < 0.10 .

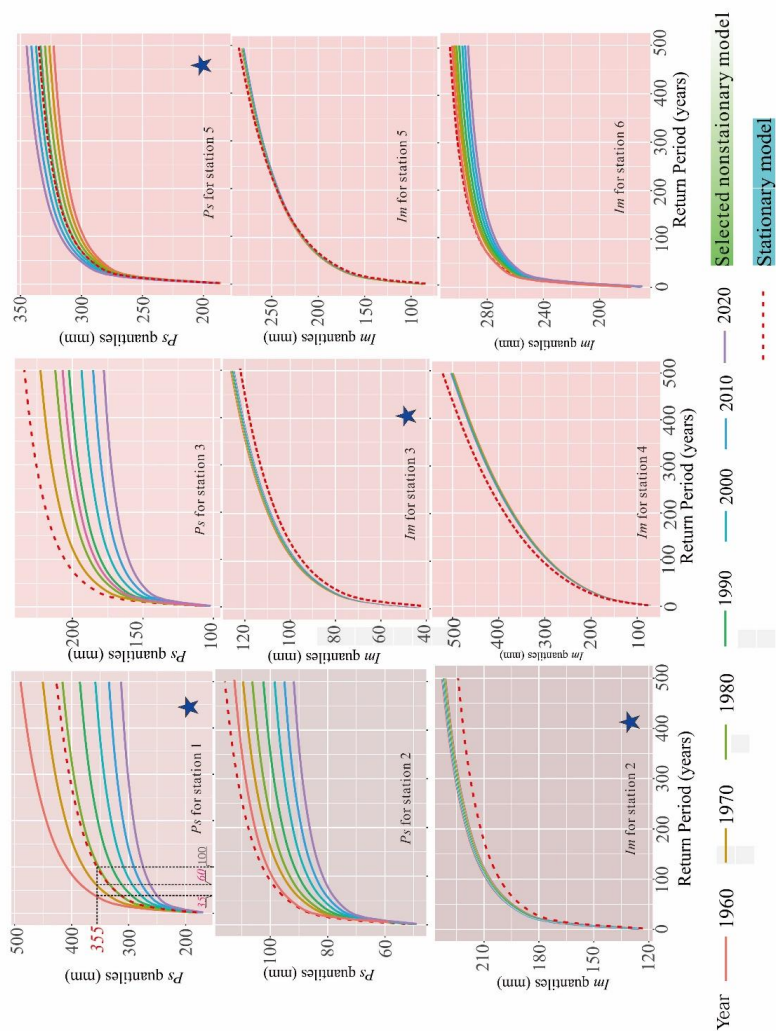


Figure 4. Estimated extreme rainfall attribute (P_s and I_m) quantiles for stationary and selected nonstationary models at six stations.

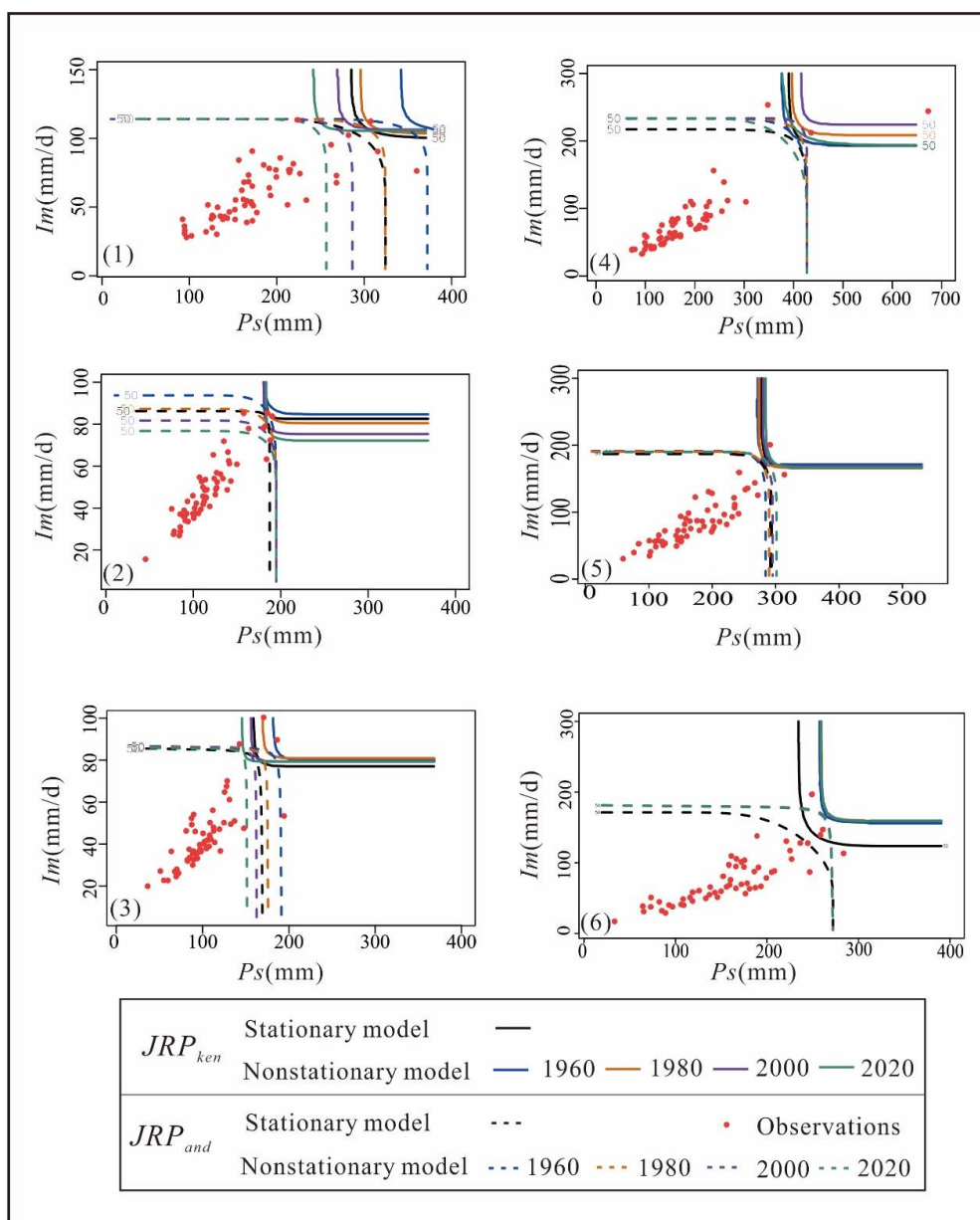


Figure 5. Isolines of Kendall return period and AND-based return period at the 50-year level for both stationary and nonstationary models. (1)-(6) represent the station number.

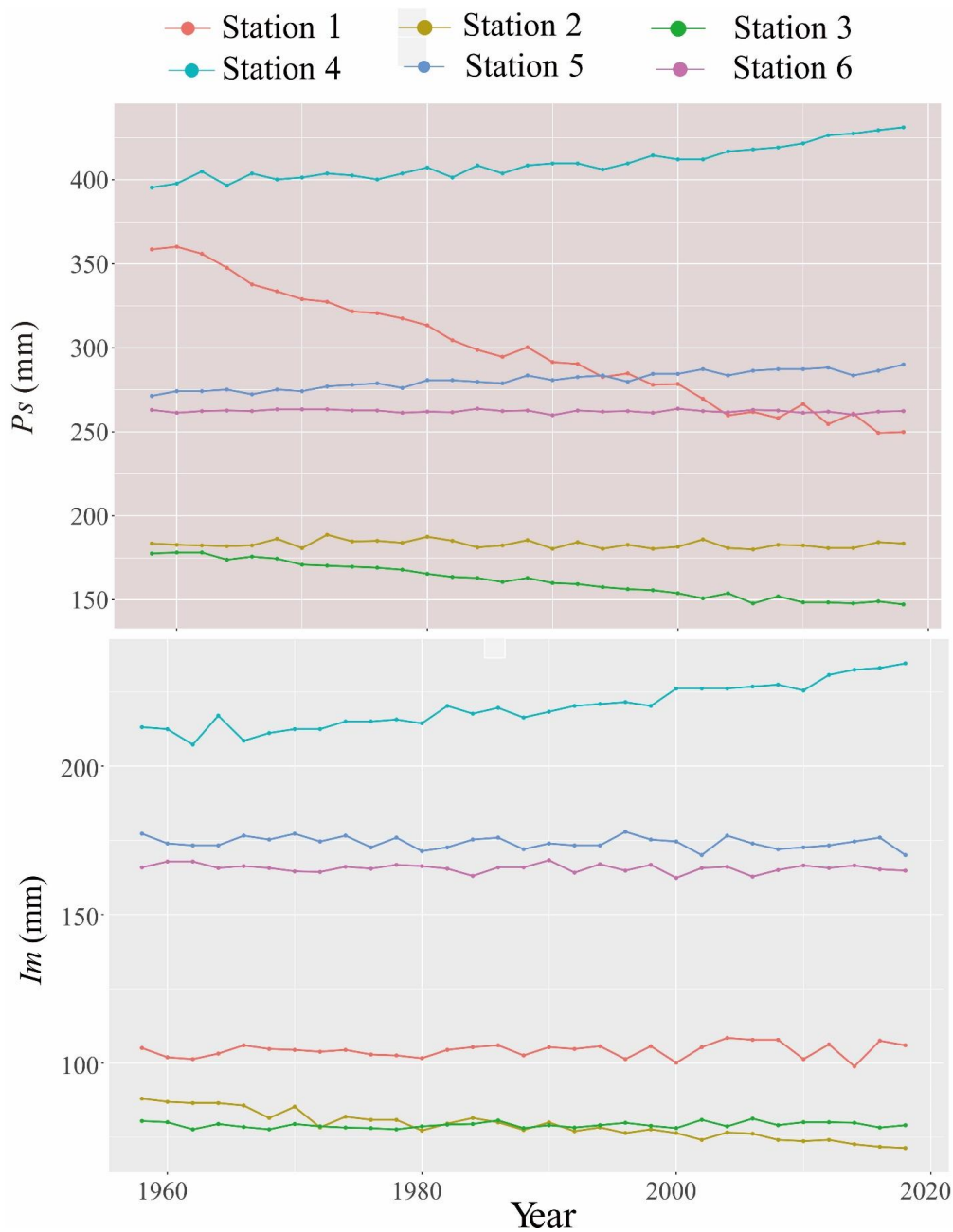


Figure 6(a). The most likely design event of P_s and I_m with $JRP_{and} = 50$ for six stations

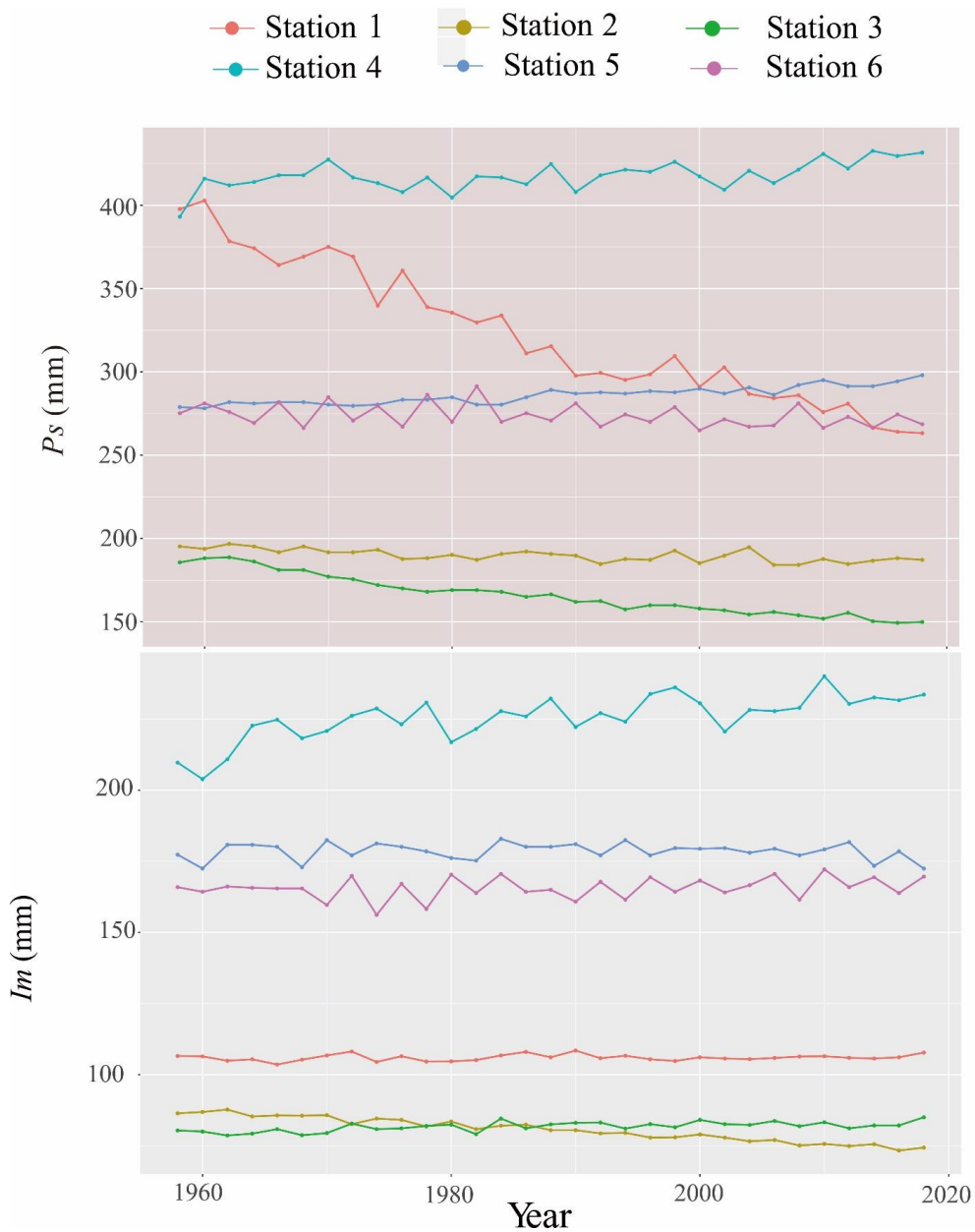


Figure 6(b). The most likely design event of P_s and I_m with $JRP_{ken} = 50$ for six stations

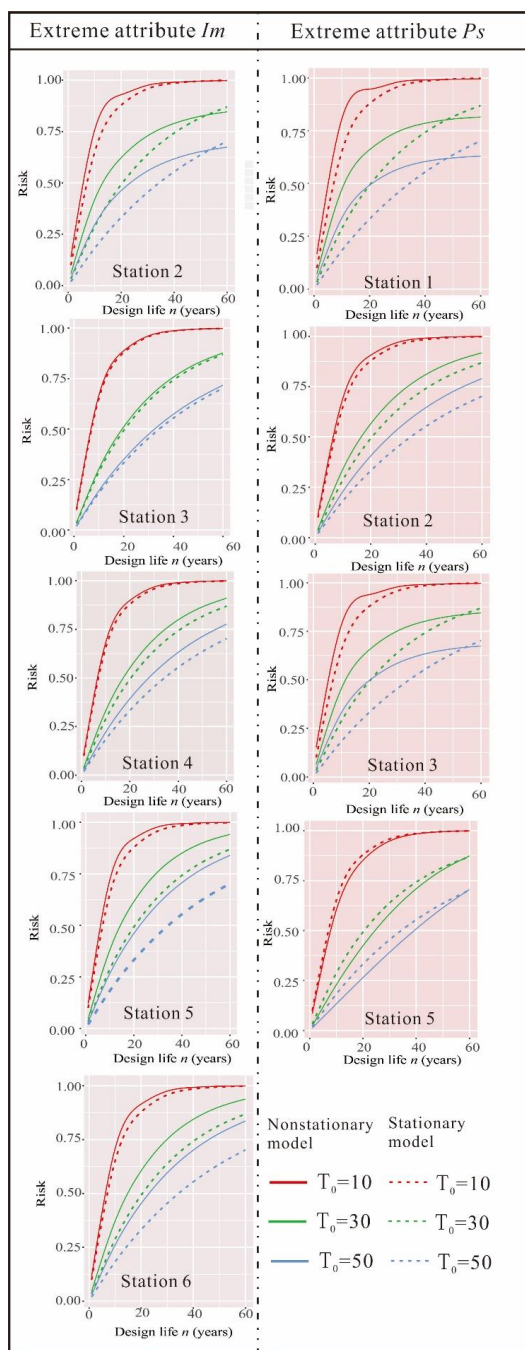


Figure 7. Nonstationary risk R of the Haihe River design extreme rainfall quantile x_{T_0} under the univariate case. The nonstationary design life level-based hydrological risk R is regarded as a function of design life n for x_{T_0} with an initial return period T_0 .

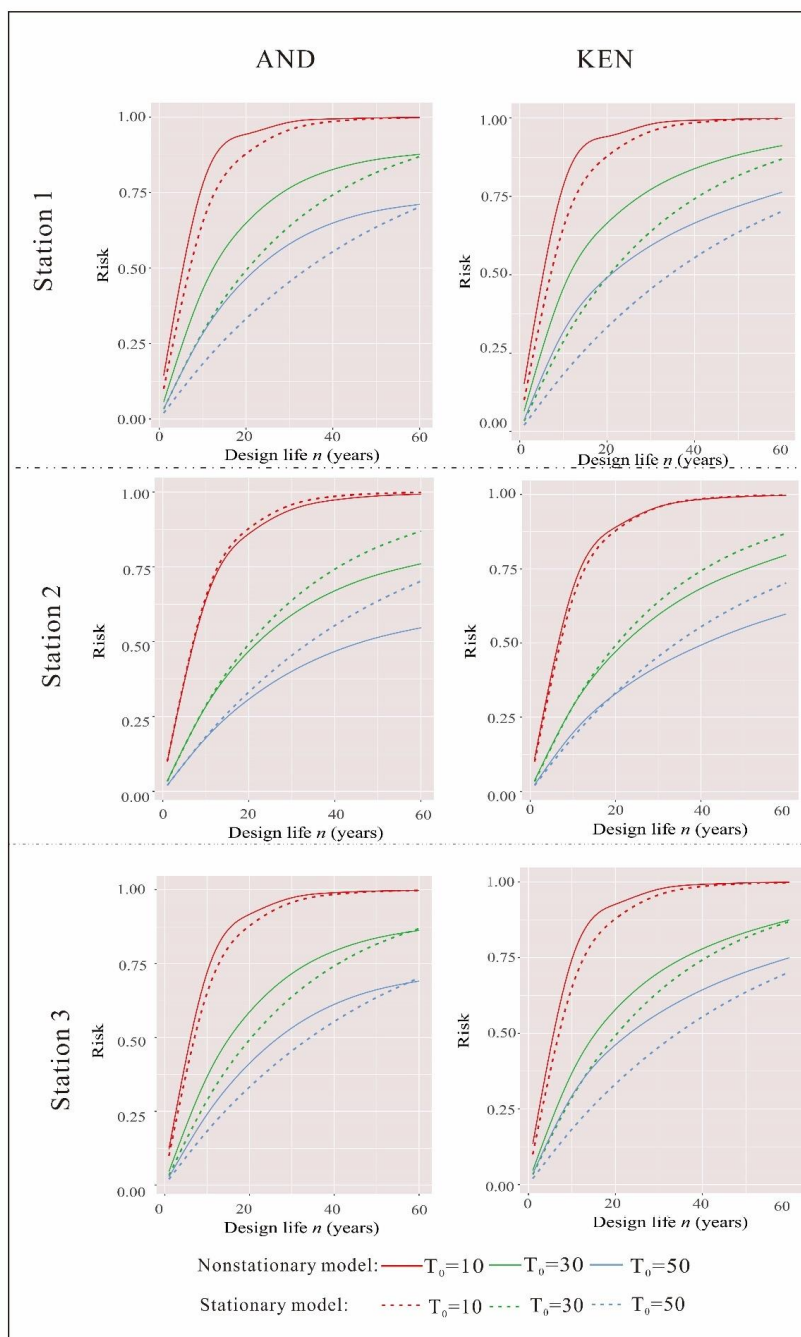


Figure 8(a). Nonstationary risk R of the most likely design combinations of P_s and Im at T_0 -year level based on JRP_{s-and} and JRP_{s-ken} .

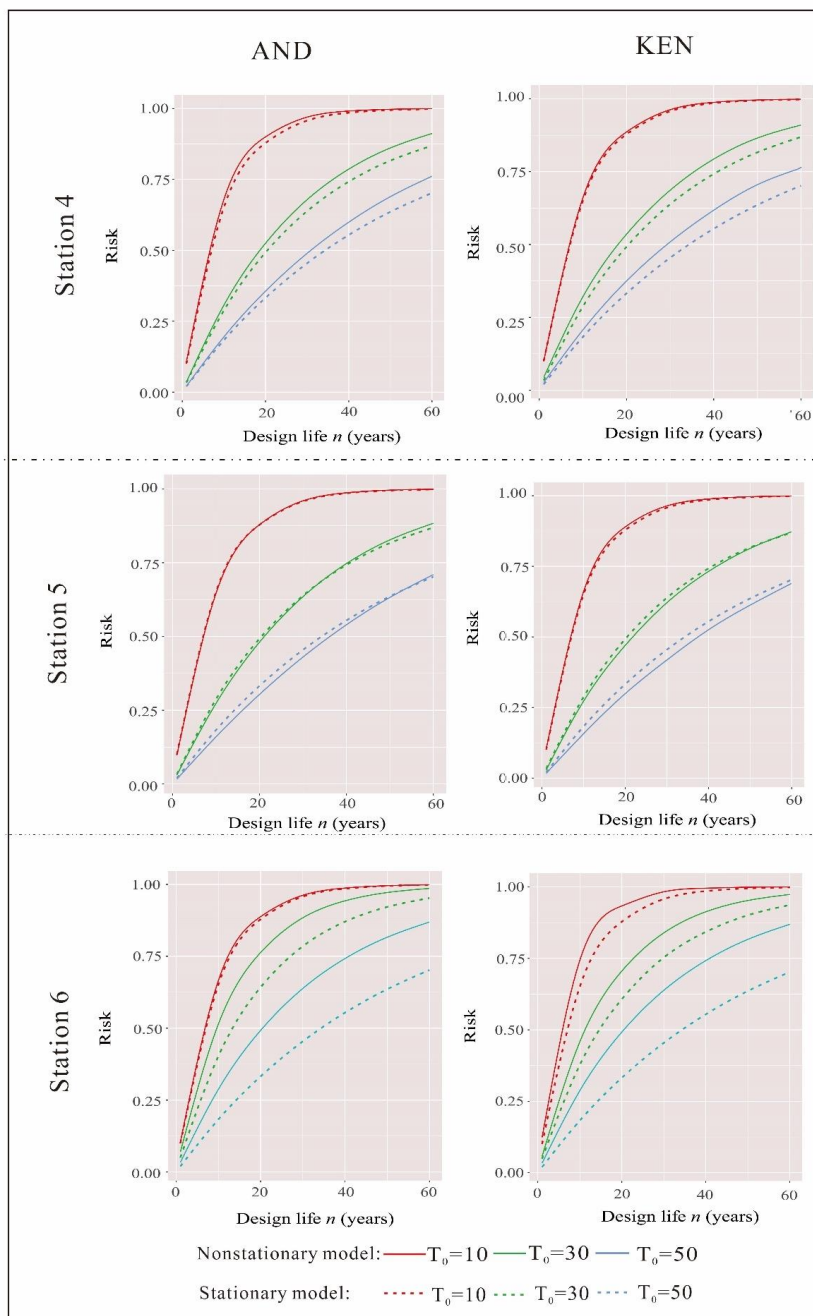


Figure 8(b). Nonstationary risk R of the most likely design combinations of P_s and Im at T_0 -year level based on JRP_{s-and} and JRP_{s-ken} .