1	A data-based predictive model for spatio-temporal
2	variability in stream water quality
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10 Abstract

Our current capacity to model stream water quality is limited particularly at large spatial scales across 11 12 multiple catchments. To address this, we developed a Bayesian hierarchical statistical model to simulate the spatio-temporal variability in stream water quality across the state of Victoria, Australia. The model 13 was developed using monthly water quality monitoring data over 21 years, across 102 catchments, which 14 span over 130,000 km². The modelling focused on six key water quality constituents: total suspended 15 16 solids (TSS), total phosphorus (TP), filterable reactive phosphorus (FRP), total Kjeldahl nitrogen (TKN), nitrate-nitrite (NO_x), and electrical conductivity (EC). The model structure was informed by 17 knowledge of the key factors driving water quality variation, which had been identified in two preceding 18 studies using the same dataset. Apart from FRP, which is largely unexplainable, the model explains 19 21.6% (NO_x) to 90.7% (EC) of total spatio-temporal variability in water quality. Across constituents, 20 the model generally captures over half of the observed spatial variability; temporal variability remains 21 largely unexplained across all catchments, while long-term trends are well captured. The model is best 22 used to predict proportional changes in water quality in a Box-Cox transformed scale, but can have 23 24 substantial bias if used to predict absolute values for high concentrations. This model can assist catchment management by (1) identifying hot-spots and hot moments for waterway pollution; (2) 25 predicting effects of catchment changes on water quality e.g. urbanization or forestation; and (3) 26 27 identifying and explaining major water quality trends and changes. Further model improvements should 28 focus on: (1) alternative statistical model structures to improve fitting for truncated data, for constituents where a large amount of data below the detection-limit; and (2) better representation of non-conservative 29 constituents (e.g. FRP) by accounting for important biogeochemical processes. 30

31 Keywords

stream water quality; spatio-temporal variability; sediments; nutrients; statistical modeling; Bayesian
 hierarchical model

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35 **1. Introduction**

Deteriorating water quality in aquatic systems such as rivers and streams can have significant environmental, economic and social ramifications (e.g. Whitworth et al., 2012;Vörösmarty et al., 2010;Qin et al., 2010;Kingsford et al., 2011). Reducing these impacts requires effective management and mitigation of poor water quality; however, high variability in water quality both across space and time reduces our ability to accurately assess the status of water quality and to develop effective management strategies. Thus, improved modelling frameworks to predict and interpret this variability would be useful for water quality management (Chang, 2008;Ai et al., 2015;Zhou et al., 2012).

43 Water quality conditions can vary across individual events, as well as at daily, seasonal and inter-annual scales at an individual location (Arheimer and Lidén, 2000; Kirchner et al., 2004; Larned et al., 2004; 44 Pellerin et al., 2012; Saraceno et al., 2009). Water quality conditions also typically differ substantially 45 across locations (Meybeck and Helmer, 1989; Chang, 2008; Varanka et al., 2015; Lintern et al., 2018a). 46 47 These variabilities in stream water quality are driven by three key mechanisms: (1) source, which defines the total amount of constituents being available in a catchment; (2) mobilization, which detaches 48 constituents (both in particulate and dissolved forms) from their sources via processes such as erosion 49 and biogeochemical processing; and (3) delivery of mobilized constituents from catchments to receiving 50 51 waters via multiple hydrologic pathways including surface and subsurface flow (Granger et al., 2010).

52 Spatial variability in stream water quality is driven by human activities within catchments (e.g., land use and management, vegetation cover etc.) (Lintern et al., 2018a;Carey and Migliaccio, 2009;Giri and Qiu, 53 54 2016;Heathwaite, 2010), along with natural catchment characteristics such as climate, geology, soil type, topography and hydrology (Hrachowitz et al., 2016; Poulsen et al., 2006; Sueker et al., 55 2001;Onderka et al., 2012). At the same time, temporal shifts in water quality are also influenced by 56 changes in pollutant sources, such as land use and land management including urbanization, agriculture 57 and vegetation clearing (Ren et al., 2003;Smith et al., 2013;Ouyang et al., 2010). In addition, water 58 quality can also vary in time with variations in the mobilization and delivery processes, which are largely 59 driven by the hydro-climatic conditions at a catchment, such as streamflow (Ahearn et al., 60 2004; Mellander et al., 2015; Sharpley et al., 2002; Zhang and Ball, 2017), the timing and magnitude of 61

62 rainfall events (Fraser et al., 1999; Miller et al., 2014) and temperature (Bailey and Ahmadi, 2014).

As abovementioned, we have good understanding of the key controls for variations in water quality. 63 albeit in an isolated, idealized context. We still lack a sound understanding of how relationships between 64 65 specific landscape characteristics and water quality can shift with influences from other landscape characteristics, and how the drivers of temporal variability in water quality can interact and vary across 66 large spatial scales (Musolff et al., 2015;Lintern et al., 2018a;Ali et al., 2017). In contrast, current 67 detailed understanding have been primarily based on field studies at small scales with detailed 68 69 information on specific temporal drivers ranging from hydrologic conditions to detailed management decisions such as fertilizer rates and application timing (Smith et al., 2013;Poudel et al., 2013;Adams et 70 al., 2014). While operational weather observation networks, stream gauging networks and remote 71 sensing can provide some of this information, developing a large-scale understanding of water quality 72 73 patterns across catchments would ideally also involve an extensive suite of management information that substantially exceeds what is currently available. 74

75 Due to the limited understanding of large-scale water quality patterns, we currently lack the capacity to 76 model spatio-temporal variabilities in water quality at large scales across multiple catchments. This hinders our ability to inform the development of effective policy and mitigation strategies over large 77 regions. Specifically, conceptual or physically-based water quality models are typically limited by the 78 simplification of physical processes such as flow pathways (Hrachowitz et al., 2016). Furthermore, 79 practical implementation of these models can be also limited by the intensive data requirements for 80 calibration and validation, particularly for large regions with highly heterogeneous catchment conditions 81 (Fu et al., 2018; Abbaspour et al., 2015). In contrast, when performed over large geographical regions, 82 statistical water quality models are generally more capable of simulating water quality variability while 83 84 requiring less detailed information and thus effort for implementation. However, existing statistical models often focus only on either the spatial variation of time-averaged water quality conditions 85 (Tramblay et al., 2010; Ai et al., 2015) or the temporal variation at individual locations (Kisi and Parmar, 86 2016;Kurunc et al., 2005;Parmar and Bhardwaj, 2015), which often limits their value as practical 87 88 management tools. Modelling the spatio-temporal variability simultaneously remains challenging over long time periods and large regions. 89

Accordingly, this research attempts to bridge the gap between fully-distributed physically-based water 90 91 quality models and data-driven statistical approaches. We aim to develop a process-informed, datadriven model to predict spatio-temporal changes in stream water quality over a large region consisting 92 of multiple catchments. Specifically, this model was established using long-term (21 years) stream water 93 quality observations across 102 catchments in Australia, with an aggregate catchment area of 130,000 94 km². To obtain the necessary understanding of process drivers required to develop this model, two 95 preceding studies were conducted on the same dataset to identify the key drivers for the spatial and 96 temporal variability of water quality, respectively (Lintern et al., 2018b; Guo et al., 2019). The aim of 97 this study is to develop an integrated spatio-temporal model using the previously-identified spatial and 98 temporal predictors, and to then assess the performance of this model. Spatio-temporal variability of 99 100 water quality was modelled using a novel Bayesian hierarchical approach which can jointly account for 101 both variability components, including accounting for varying temporal water quality dynamics between 102 catchments. This modelling approach also has relatively low requirement for input data, which keeps 103 the modelling detail commensurate with the level of data availability. During the model development, we also obtained additional understanding on the patterns of spatial variations in the effects of each 104 temporal predictor. The model can potentially provide useful information for large-scale catchment 105 management, assessment and policy making, such as testing major changes in land use patterns, 106 107 informing pollution hot-spots, as well as identification and attribution of water quality trends and changes over time. 108

109 **2. Method**

We first discuss the process used to develop the integrated spatio-temporal model (Section 2.1). Sections
2.1.1 and 2.1.2 introduces the statistical modelling framework and the data used for model development,
respectively. The approaches to determine model structure was then introduced, which include the
choice of key predictors (Section 2.1.3) and the calibration for model parameters (Section 2.1.4). Finally,
the approaches to evaluate model performance and robustness are described in Section 2.2.

115 **2.1 Model development**

116 2.1.1 Spatio-temporal modelling framework

A Bayesian hierarchical approach was used to model the spatio-temporal variability in stream water 117 118 quality. The Bayesian approach enables the inherent natural stochasticity of water quality to be incorporated into the model (Clark, 2005). A key strength of applying the hierarchical model structure 119 to analyze spatio-temporal variability is that this structure enables the key controls of temporal 120 variability in water quality to vary across locations (Webb and King, 2009;Borsuk et al., 2001). This 121 variability has been found to be important in other study regions where the (temporal) solute export 122 regime varies with catchment characteristics such as climate and land use (Musolff et al., 2015:Poor and 123 124 McDonnell, 2007).

125 The structure of the Bayesian hierarchical model is presented below in Eq. 1 to 6. Eq. 1 formulates the 126 transformed constituent concentration (see Section 2.1.2 for justification) at time *i* and site *j* (C_{ij}) as a 127 normally distribution with a mean μ_{ij} and standard deviation σ representing inherent randomness.

$$\boldsymbol{C}_{ij} \sim \boldsymbol{N}(\boldsymbol{\mu}_{ij}, \boldsymbol{\sigma}) \tag{1}$$

128 To represent spatio-temporal variability, μ_{ij} is modelled as the sum of the site-level mean constituent 129 concentration (\bar{C}_i) and the deviation from that mean at time $i(\Delta_{ii})$ (Eq. 2).

$$\boldsymbol{\mu}_{ij} = \boldsymbol{C}_j + \boldsymbol{\Delta}_{ij} \tag{2}$$

To describe spatial variability, the site-level mean concentration at site j (\bar{C}_j) is modelled as a linear function of a global intercept (*intC*), and the sum of *m* catchment characteristics $S_{1,j}$ to $S_{m,j}$ (e.g. land use, topography) weighted by their relative contributions to spatial variability (βS_1 to βS_m) (Eq. 3).

$$\overline{C}_{j} = intC + \beta S_{1} \times S_{1,j} + \beta S_{2} \times S_{2,j} + \dots + \beta S_{m} \times S_{m,j}$$
(3)

133 The temporal variability, represented by the deviation from the mean (Δ_{ij}) , is a linear combination of *n* 134 temporal variables, $T_{1,ij}$ to $T_{n,ij}$ (e.g., climate condition, streamflow, vegetation cover) (Eq. 4), at time 135 *i* and site *j*.

$$\Delta_{ij} = \beta T_{1,j} \times T_{1,ij} + \dots + \beta T_{n,j} \times T_{n,ij}$$
(4)

The selection of key spatial and temporal predictors for the model has been performed in our two
preceding studies (Lintern et al., 2018b; Guo et al., 2019) and is briefly described in Section 2.1.3. Eq.
1 to 4 enable the model to separately represent the spatial and temporal variability in water quality;

however, there is still a further step required to make the model fully spatio-temporal (i.e. being able to 139 140 predict over both time and location). Specifically, in Guo et al. (2019), clear spatial variation was observed in the relationships between water quality and its key temporal predictors (i.e. in the $\beta T_{N,j}$ in 141 Eq. 4). To be able to model multiple catchments across a large spatial area simultaneously, we must 142 143 account for differences in these temporal influences across sites. To do this, the effect of each temporal variable at site j ($\beta T_{N,i}$ with N in 1,2, ... n) is drawn from a distribution with a mean of $\mu\beta T_{N,i}$ (Eq. 5), 144 which is then modelled with a linear combination of two additional chatchment characteristics, $ST_{N1,i}$ 145 and $ST_{N2,i}$ (Eq. 6). Details of the selection for these two additional predictors are presented in Section 146 2.1.3. 147

$$\beta T_{N,j} \sim N(\mu \beta T_{N,j}, \sigma \beta T), for N in 1, 2, \dots n$$
⁽⁵⁾

$$\mu \beta T_{N,j} = int \beta T_N + \beta S T_{N1} \times S T_{N1,j} + \beta S T_{N2} \times S T_{N2,j}$$
(6)

148 2.1.2 Data collection and processing

The Bayesian hierarchical model was developed with 21 years of monthly stream water quality 149 observations at 102 catchments in the state of Victoria, Australia (aggregate catchment area > 130,000 150 151 km²). The collection and processing of the data are detailed in previous publications that worked with the same dataset (Lintern et al., 2018b; Guo et al., 2019). Briefly, stream water quality data were 152 extracted from the Victorian Water Measurement Information System (Department of Environment 153 Land Water and Planning (DELWP) Victoria, 2016b), which contains monthly grab samples of water 154 quality at approximately 400 sites across Victoria. Water quality data sampled between 1994 and 2014 155 at 102 sites were used to develop the model (Fig. 1). These sites and time period were chosen because 156 they provided the longest consistent period of continuous records over the greatest number of monitoring 157 sites. The catchments corresponding to these water quality monitoring sites were delineated using the 158 Geofabric tool (Bureau of Meteorology, 2012), and have areas ranging from 5 km² to 16,000 km². The 159 water quality parameters of interest were: total suspended solids (TSS), total phosphorus (TP), filterable 160 reactive phosphorus (FRP), total Kjeldahl nitrogen (TKN), nitrate-nitrite (NO_x) and electrical 161 conductivity (EC). These parameters represent sediments, nutrients and salts, which are some of the key 162 concerns for water quality managers in Australia and around the world. These water quality samples 163 were collected following standard DELWP protocols (Australian Water Technologies, 1999) and 164

- analysed in National Association of Testing Authorities accredited laboratories. Note that in the
- sampling protocol, FRP is defined as '*Reactive Phosphorus for a filtered sample to a defined filter size*
- 167 (e.g. $RP(<0.45 \ \mu m))$ ', which is equivalent to the more widely-used terminology, SRP i.e. Soluble
- 168 Reactive Phosphorus (Jarvie et al., 2002).



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Figure 1. Map of (a) the 102 selected water quality monitoring sites and their catchment
 boundaries, with inserts showing the location of the state of Victoria within Australia; (b) annual
 average temperature and (c) annual precipitation and (d) elevation across Victoria.

To compile a dataset for the potential spatial explanatory variables (i.e. predictors to explain spatial 173 174 variability in water quality), a comprehensive literature review was conducted (Lintern et al., 2018a), which summarized the key catchment landscape characterisitics that are widely known to influence 175 176 water quality. Further, as part of Lintern et al. (2018b), fifty potential explanatory catchment 177 characteristics were selected, which included catchment land use, land cover, topographic, climatic, geological, lithological and hydrological catchment characteristics. These variables were derived using 178 datasets obtained from Geoscience Australia (2004, 2011), the Bureau of Meteorology (2012), the 179 180 Bureau of Rural Sciences (2010), Department of Environment Land Water and Planning Victoria (2016) 181 and the Terrestrial Ecosystem Research Network (2016) (see Table S1 in the Supplementary Material 182 for detailed variable names and data sources). We used a static set of land use data from 2005-2006 to represent the entire study period, as a preliminary analysis between 1996 and 2011 suggested less than 183 1% changes in the key land uses in these catchments (i.e. agricultural, grazing, conservation). 184

185 Nineteen potential temporal explanatory variables were included. Firstly, data of discharge (originally

in ML d⁻¹) and water temperature (°C) corresponding to the same timestamps for water quality 186 187 observations were also extracted for each monitoring site over the study period (Department of Environment Land Water and Planning Victoria, 2016). Discharge was converted to runoff depth (mm 188 d^{-1}) for each catchment, and the average streamflows over 1, 3, 7, 14 and 30 days preceding the water 189 quality sampling dates were calculated. In addition, we extracted gridded dataset from the Australian 190 Water Availability Project (AWAP) (Frost et al., 2016; Raupach et al., 2009, 2012) and Australian Water 191 Resources Assessment Landscape (AWRA-L) model (Frost et al., 2016). These datasets were used to 192 calculate catchment averaged values of daily average temperature (°C), daily rainfall (mm), antecedent 193 rainfall (1, 3, 7, 14 and 30 days preceding sampling), dry spell (> 0.1mm rainfall) length in the antecedent 194 14 days, daily actual evapotranspiration (ET) (mm), as well as soil moisture for the root-zone and the 195 196 deep-zone (averaged volumetric content for shallower and deeper than 1m, respectively). In addition, catchment averaged monthly NDVI data were extracted from Advanced Very High Resolution 197 198 Radiometer (AVHRR) Product (Eidenshink, 1992) and Moderate Resolution Imaging Spectroradiometer MOD13A3 (NASA LP DAAC, 2017). A summary of these datasets of temporal 199 200 variables and their corresponding sources are in Table S2 in the Supplementary Material and details are 201 provided in Guo et al. 2019.

202 The raw input data were filtered and transformed to increase the data reliability, continuity and symmetry, making them more suitable for use in the linear spatio-temporal model structure (Eq. 3, 4 203 204 and 6). For the filtering process, we first removed all water quality records with flags indicating quality 205 issues. We also removed any values below the detection limit (DL), which was defined as the 'minimum concentration detected for which there is 95% confidence of accuracy and therefore is accurate enough 206 207 to report' in the monitoring protocols for this dataset (Australian Water Technologies, 1999). This was because the uncertainty in values below the DL would be amplified after transformation, which would 208 209 influence the subsequent model fitting. Furthermore, those undetectable low concentrations were of less 210 interest for management purposes. Water quality records corresponding to days with zero flows were 211 also excluded from further analyses.

The transformation process was performed for each of the spatial catchment characteristics, temporalexplanatory variables, as well as each water quality constituent to improve the symmetry of individual

distributions. The log-sinh transformation (Wang et al., 2012) (Eq. 7) was used for all catchment characteristics, due to its ability to resolve the presence of zero values in several of the catchment characteristics (e.g., percentage area of individual land uses). The GA package in R (Luca Scrucca, 2019) was used to identify the log-sinh transformation parameters (*a* and *b*) for each spatial explanatory variable that minimized the data skewness (i.e. symmetry is maximized) across all 102 catchments.

219
$$y_{log-sinh} = \frac{1}{b} \log(\sinh[a + by_{raw}])$$
(7)

In addition, all observed constituent concentrations and temporal explanatory variables were Box-Cox
transformed (Box and Cox, 1964) (Eq. 8).

222
$$y_{Box-Cox} = \begin{cases} \frac{y_{Raw}^{\lambda}-1}{\lambda}, & for \ \lambda \neq 0\\ logy, & for \ \lambda = 0 \end{cases}$$
(8)

223 For each variable, the optimal Box-Cox transformation parameter λ was identified using the *car* R 224 package and a maximum likelihood-like approach. We first identified the optimal Box-Cox parameter λ 225 using the data at each site (i.e. 21-year time-series). The averaged λ across all sites was then used to 226 transform the data across all catchments together. This transformation approach ensured that all sites 227 used a consistent transformation parameter. All transformation parameters used are summarized in 228 Tables S3 and S4 in the Supplementary Material. The transformation process has greatly improved the 229 data symmetry and thus suitability for use in a linear model (the quality of the transformations was assessed via visual inspection in Lintern et al., 2018b; Guo et al., 2019; and summarized in Figures S2, 230 231 S4 and S6 in the Supplementary Material).

232 2.1.3 Selection of key model predictors

Key predictors for the model were selected in a process-informed and data-driven manner based on our two preceding studies (Lintern et al., 2018b; Guo et al., 2019). Lintern et al. (2018b) identified the best spatial predictors (S_1 to S_m in Eq. 3) for the model, while the best temporal predictors across all sites (T_1 to T_n in Eq. 4) have been identified in Guo et al., (2019). In both studies, the best predictors were selected using an exhaustive search approach (May et al., 2011;Saft et al., 2016), which considered all possible combinations of the potential predictors introduced earlier in this section. This selection approach required firstly fitting an individual model to all possible candidate predictor sets, and then 240 comparing all fitted models to select a single best set of predictors. Alternative models were evaluated

241 based on the Akaike Information Criterion (AIC) (Akaike, 1974) and Bayesian Information Criterion

(BIC) (Schwarz, 1978) to ensure optimal balance between model performance and complexity. 242

243 The best predictors to explain the spatial and temporal variabilities in each constituent are listed in Table 244 1. Generally speaking, the key factors controlling the spatial variability in river water quality were land-245 use and long-term climate conditions (Lintern et al., 2018b). Temporal variability was mainly explained 246 by temporal changes in streamflow conditions, water temperature and soil moisture (Guo et al., 2019). 247 The potential mechanisms via which these key drivers influence water quality are discussed in details in these two previous studies. 248

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Table 1. Key factors affecting the spatial and temporal variability for each of six constituents, as identified in Lintern et al. (2018) and Guo et al. (2019b), respectively. 250

252 Whilst the previous studies (Lintern et al. 2018b, Guo et al. 2019) identified the predictors for spatial and temporal variability respectively, they did not provide guidance on the predictors for spatial 253 254 variability in the relationships between drivers of temporal variability and temporal water quality response (i.e. βT in Eq 4). As such, the final step of the predictor selection process to develop the 255 combined spatio-temporal model was to identify the key catchment characteristics that affect spatial 256 variability in the hydroclimatic parameters driving temporal changers in water quality (βT_1 to βT_n in Eq. 257 258 4, also right column in Table 1). This is achieved by selecting two spatial characteristics that are most closely related to the coefficient for each temporal predictor (ST_{N1} and ST_{N2} , Eq. 6) across all sites, 259 where only two spatial characteristics were used to avoid over-fitting. Selection of these two spatial 260 characteristics were based on a Spearman correlation analysis between the fitted parameter values of 261 each temporal predictor variable and the fifty potential spatial explanatory variables (as mentioned 262 263 earlier in this section), following three steps:

- 1. from the 50 candidate spatial predictors, the one with the highest Spearman correlation with βT_N is 264 selected as ST_{N1} , provided the correlation is statistically significant (p<0.05); 265
- the subset of remaining spatial predictors with spearman correlation with $ST_{N1} < 0.7$ is found; and 266 2.
- 267 3. from this subset, the spatial predictor with the highest spearman correlation with βT_N is selected as

268 ST_{N2} , provided the correlation has p<0.05;

Steps 2 and 3 intended to avoid cross-correlations between ST_{N1} and ST_{N2} . The selected spatial characteristics that influence the temporal relationships in our model are presented and interpreted in Section 3.1. Note that the entire process to select ST_{N1} and ST_{N2} was performed with the fitted parameters for each predictor of the temporal variability obtained from Guo et al. (2019).

273 2.1.4 Model calibration

After identifying the spatial and temporal predictors for each constituent, as well as the spatial 274 275 characteristics which affect the strengths of each temporal predictor, the Bayesian hierarchical spatio-276 temporal model was fitted for each constituent across all monitoring sites simultaneously. To achieve this, we used the R package *rstan* (Stan Development Team, 2018), which enabled both the sampling of 277 278 parameter values from posterior distributions with Markov chain Monte Carlo (MCMC) and model 279 evaluation. Constituent standard deviation (σ) was assumed to be drawn from a minimally informative 280 prior half-normal of N(0,10) distribution truncated to only positive values (Gelman, 2006; Stan 281 Development Team, 2018). The regression coefficient of each spatial predictor ($\beta S_1, \beta S_2, ..., \beta S_m$ in Eq. 3) was independently drawn from hyper-parameter distributions of $N(\mu\beta S_M, \sigma\beta S_M)$. The site-level 282 regression coefficients of the temporal predictors ($\beta T_{1,j}$, $\beta T_{2,j}$, ..., $\beta T_{n,j}$ in Eq. 4, respectively) were 283 sampled from the corresponding hyper-parameter distribution of $N(\mu\beta T_N, \sigma\beta T_N)$. The hyper-parameters 284 285 were further assumed to be drawn from minimally informative prior distributions, following 286 recommendations in Gelman (2006) and Stan Development Team (2019): for all the hyper-parameter 287 means, a normal prior distribution of N(0,5) was used; for all the hyper-parameter standard deviations, a half-normal prior distribution of N(0,10) was used, which was truncated to only positive values. In 288 289 each model run there were four independent Markov chains. A total of 20,000 iterations were used for each chain. Convergence of the chains was ensured by checking the *Rhat* value (Sturtz et al., 2005), 290 291 which is a summary statistic on the convergence of the Bayesian models from the four Markov chains used in model calibration (Stan Development Team, 2018). Specifically, an Rhat value much greater 292 than 1 indicates that the independent Markov chains have not been mixed well, and a value of below 1.1 293 294 is recommended (Stan Development Team, 2018).

295 2.2 Model performance evaluation and sensitivity analyses

Performance evaluation of the model was undertaken on several aspects of the model results (Section.
3.2). Since the model was calibrated in a Box-Cox transformation scale (see justification in Section
2.1.2), the Box-Cox transformation scale was used for model evaluation to enable a clear investigation
on the influences of a wide range of factors that can influence model performance. Detailed performance
evaluations include:

- Ability to capture total spatio-temporal variability. Firstly, the simulations from the fitted model 301 1. and the corresponding observed concentrations were compared at 102 sites altogether to 302 understand how the overall spatio-temporal variabilities were captured. For each constituent, 303 this evaluation was performed with: 1) these above-DL data to focus only on data used for 304 calibration (as detailed in Section. 2.1.2); and 2) the full dataset including the below-DL data 305 306 (set to half of the DL of the specific constituent), to understand how well the model represents the full distribution of constituent concentrations. A good model performance when including 307 the below-DL data would suggest that the calibrated model is transferable to below-DL data 308 too. All performance assessments were based on both visual inspection of model fitting as well 309 310 as the Nash-Sutcliffe efficiency (NSE), which quantified the proportion of variability that was 311 explained by the model (Nash and Sutcliffe, 1970).
- Proportions of spatial and temporal variability explained. This involved a decomposition of the
 total observed variability using Eq. 2., into proportions contributed by spatial variability
 (variations in all site-mean concentrations from the grand average of site-mean concentrations)
 and temporal variability (variations in all concentrations from the corresponding site-mean
 concentrations). The corresponding modelled values were then used to calculate NSE for each
 variability component of each constituent.
- 3. *Ability to capture variation in ambient conditions across space, and temporal variation* 3. *Ability to capture variation in ambient conditions across space, and temporal variation* 3. *(including trends) across multiple catchments.* These were evaluated by a) comparing all 3. *simulated and observed site-averaged long-term mean concentrations; and b) comparing the* 3. *simulated and observed time-series and long-term trends at representative sites.* Further to a),

performance was also evaluated on a real measurement scale by first back-transforming all modelled sample concentrations, calculating the back-transformed site-level means and then compared those to the corresponding observations. A further analysis to b) was also performed by comparing the estimated Sen's slope (Akritas et al., 1995) for the observations and simulations at all sites, and then computing the percentage of sites where the observed trends as indicated by the Sen's slope have been correctly represented by the model.

Additional evaluations of model sensitivity were conducted with calibration and validation on subsetsof the full data (Section. 3.3), to understand model transferability and stability:

Model sensitivity to the monitoring sites used for calibration. We randomly selected 80% of the
 sites for calibration and used the remaining 20% for validation, and repeated this validation
 process 50 times. We compared all calibration and validation performances of these 'partial
 models' with each other, as well as with the performance of the full model, to obtain a
 comprehensive evaluation of the sensitivity of model performance to calibration sites.

2. Model sensitivity to calibration data period. Since the study region was greatly influenced by a 335 prolonged drought from 1997 to 2009 - known as the Millennium Drought (van Dijk et al., 336 2013), we also investigated model robustness for before, during and after this drought period. 337 Specifically, we calibrated the model to each pre-, during- and post-drought period (1994-1996, 338 1997-2009 and 2010-2014, respectively) with model validation on the remaining data. For 339 example, when calibrating to the pre-drought period (1997-2009), validation was performed on 340 the merged during and post-drought period (1994-1996 plus 2010-2014). The corresponding 341 342 calibration and validation performances were compared with each other as well as against that of the full model, to identify potential impacts of the drought on model robustness. 343

344 3. Results

345 **3.1 Spatial variation in the impact of temporal factors**

The key controls of the spatial and temporal variations in water quality have been identified in our two preceding studies (Lintern et al. 2018b, Guo et al. 2019) and briefly summarized in Section 2.1.3. and are thus not discussed here. As also detailed in Section 2.1.3, to achieve full spatio-temporal predictive capacity, the model developed in this study considers the spatial variation in the strength of each temporal predictor by using two additioal catchment spatial characteristics ($ST_{N1,j}$ and $ST_{N2,j}$ in Eq. 6). on the Spearman's correlations. Here we focus on the most important temporal predictor for each constituent, streamflow, where Table 2 shows the two spatial characteristics identified that are most closely related to the spatial variation of the effects of impact of streamflow on water quality. The full list of the selected key catchment characteristics for all temporal predictors of each constituent is summarized in Table S5 and visualized in Figure S4.

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Table 2. The key catchment landscape characteristics that are related to the varying relationships of water quality and same-day streamflow across space, which were selected as the two predictors for the streamflow effect in our model. The corresponding Spearman's correlation (ρ at p<0.05) between the effect of streamflow and each catchment characteristic is presented.

360 TSS, TP and TKN show consistent patterns of the spatial variation in the effects of streamflow on water quality, which are strongly driven by the differences in average rainfall conditions across catchments. 361 Specifically, streamflow generally has a larger effect on water quality in catchments with higher average 362 annual rainfall. Since the streamflow effects are positive for the majority of catchments (as shown in 363 Figure S5), these correlations indicate that for the same increase in transformed streamflow, a greater 364 increase in transformed concentrations of TSS, TP and TKN will occur at a catchment with higher annual 365 366 average rainfall. Given that the Box-Cox lambda values (Table S4) are close to zero, the transformation is log-like and hence changes in transformed flow and concentration approximately correspond to 367 proportional changes in the real values of flow and concentration. In contrast, for FRP, NO_x and EC, the 368 369 spatial patterns of streamflow effects are specific to each constituent. This difference in the model results 370 between TSS, TP and TKN against the other constituents might be related to the distinct transport pathways of particulate and dissolved constituents. The former is predominantly related to surface flow 371 and thus relies heavily on rainfall contribution. Dissolved constituents are likely transported along the 372 subsurface pathway. Apart from streamflow, the spatial patterns in other key temporal drivers of water 373 quality (e.g. antecedent streamflow, soil moisture etc.) are less consistent across different constituents 374 375 (Figure S4).

376 3.2 Model performance evaluation

377 The spatio-temporal water quality models show varying performances between the constituents. When

378 assessed with only the above-DL data (Fig. 2), the best performing models are those for EC and TKN, 379 which capture 90.7% and 65.8% of the total observed spatio-temporal variability. The modelling performance is lowest for FRP, NO_x and TSS, with NSE values of -1.92, 0.216 and 0.225, respectively. 380 381 When evaluated against the entire dataset (i.e., including both below- and above DL data), the models explain 19.9% (FRP) to 88.6% (EC) of spatio-temporal variability (Table 3). Model performances for 382 FRP, NO_x and TSS improve notably compared with the previous evaluation of above-DL data, however, 383 they remain as the three constituents that are most difficult to predict. We further discuss the possible 384 385 factors influencing their model performance in Section 4.1.



386

387 388

Figure 2. Performance of the spatio-temporal models for each of the six constituents, represented by the simulated median concentrations and corresponding observations of above-DL records across all 102 calibration sites, in Box-Cox transformed space. Darker regions 389 390 represent denser distribution of simulation and observation points. Dashed red lines show the 1:1 lines whereas dashed blue lines show the DL levels. For each constituent, the percentage of 391 392 data below the DL and the model performance (NSE) are also specified.

393 394

Table 3. Comparison of model performance for all records and only the above-LOR records for each constituent.

The model performance to predict spatial and temporal variability is summarized in Figure 3, which 395 396 compares the observed and explainable variability for each of the spatial and temporal components 397 (detailed in Section 2.1.4). Regarding the observed variability (lighter colours), EC is strongly dominated by spatial variability (91.8%), highlighting that within-site variation in water quality is minimal compared to between-site variation. To a lesser extent, spatial variability also contributes to major proportions of total variability for TP and TKN (60.8% and 66.6%, respectively). TSS, FRP and NO_x are more influenced by temporal variability (57.4%, 56.6%, 60.5%, respectively).

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Figure 3. Observed spatial and temporal variabilities as proportions of the total variability (total width of each bar, 100%). The dashed line differentiates temporal variability (left side) with
spatial variability (right side), and the darker colours highlight the proportions of spatial and
temporal variabilities that are explainable by the model. All values were estimated in Box-Cox transformed space.

409 The explained variability (darker colours) shows that, across all catchments, temporal variability is much 410 more difficult to model compared with spatial variability. It also appears that a substantial part of the model's overall performance is driven by its ability to capture spatial variability in ambient water quality 411 conditions. For example, the models for TSS, FRP and NO_x show poorer overall performance (Fig. 2, 412 with NSE values of 0.225, -1.92 and 0.216, respectively)), because the total variability for each of these 413 414 is dominated by temporal variability (57.4%, 56.6%, 60.5%, respectively), which largely remains unexplained by the model (Fig. 3). In contrast, the EC model shows a very good fit with 90.7% of total 415 variability explained - 91.8% of the total observed variability is due to spatial variability, of which 416 94.7% is explained by the model. Therefore, although the EC model can only explain a small portion of 417 temporal variability (20% out of 8.2% of total variability), the overall model performance remains high. 418 As highlighted in Fig. 3, the model has good capacity to capture spatial variability in water quality. 419

420 This is further evaluated in Fig. 4 by comparing the simulated and observed site-level mean

concentrations. The highest model performance is for EC and lowest performance is for FRP
(explaining 94.7% and 44.2% spatial variability, respectively). At the back-transformed scale, the
model shows greater biases for sites with higher concentrations (approximately the highest 10% sites
for each constituent) (Fig. 5). This is not surprising as the model was fitted to a Box-Cox transformed
space that reduces focus on high values and increases the focused on low values. This compromised its
ability to represent sites with unusually high concentrations. The implications of the model having
higher predictive capacity in the transformed scale is further discussed in Section. 4.1.



428

Figure 4. Model fit for site-level mean concentration at the 102 calibration sites for six
constituents, with the 95% lower and upper bounds of posterior simulations shown in vertical
grey lines. All simulations and observations are presented in in Box-Cox transformed space. The
NSE for each constituent is also shown and red dash lines show the 1:1 lines.

433



Figure 5. Back-transformation of the model simulations to the measurement scale emphasizes lack of fit
for the highest concentrations, illustrated by simulated against observed site-level mean concentrations of
each constituent in a back-transformed scale. The 95% lower and upper bounds of all posterior
simulations shown in vertical grey lines. The NSE for each constituent is also shown and red dash lines
show the 1:1 lines.

434



relatively limited. This is further explored in Fig. 6, where the observed and simulated time-series are

443 presented for one monitoring site for each constituent, at which the model performance (NSE) was the

444 highest. These results show that even for catchments where the model has the highest ability to capture

temporal variability, the model consistently underestimated temporal variability for all constituents.



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Figure 6. Model fit of the within-site (temporal) water quality variability, illustrated with the
observed and simulated time-series for the best-performing site for each constituent. All values
are presented in Box-Cox transformed space. The NSE for each constituent is also shown. The
red line indicates the corresponding mean of all posterior simulations, while the pink bands
show the corresponding 95% lower and upper bounds (only visible for FRP).

452 Fig. 6 also illustrates that, although the model shows substantial underestimation of temporal

- 453 variability within site, long-term temporal trends in the time-series are well captured at the best sites
- 454 (except for FRP). Table 4 summarizes the ability of the model to capture observed trends across all

455 102 catchments for each constituent. In general, the model is able to capture observed trends in most

- 456 sites for NO_x and EC and for both positve and negative trends. For TP and TKN, positive trends are
- 457 well captured while for TSS the negative trends are better captured.

Table 4. Model ability to capture observed water quality trends across all monitoring sites for
each constituent. The percentages of sites where observed positive and negative trends are
captured by the model are presented separately. Values in brackets indicate numbers of sites
where corresponding positive or negative trends are observed. For detailed estimation of these
percentages please refer to Sect. 2.2.

463 **3.3 Model sensitivity analyses**

We first compare the performance of each spatio-temporal model fitted with the full dataset with those obtained from the 50 corresponding "partial" models that were calibrated to only 80% of the monitoring

- sites. Note that in this comparison, the FRP model was not assessed due to its poor performance (Section
- 467 3.2). The calibration and validation results for the 50 partial models are summarized in Table 5 along

with the performance of the full model calibrated to all 102 sites (see Figs. S6 and S7 in the Supplementary Material for detailed comparison of model residuals of the partial calibration/validation). Across constituents, the calibration performance of the full model was comparable with the 50 partial models. In addition, model performance is highly consistent between corresponding calibration and validation, with most differences in NSEs less than 0.1. These suggest that the spatio-temporal model performance is highly robust and unaffected by the choice of calibration sites.

Table 5. Comparison of model performances (as NSE) of the full model (Column 2) and the 50
partial models (Columns 3 to 5) with each calibrated to 80% randomly selected monitoring sites.
Columns 3 to 5 summarize the mean, minimum and maximum NSE values across the 50 runs,
where for each constituent, the top row showing calibration performance and the bottom row
showing the validation performance (i.e. at the 20% sites that were not used for calibration).

480 The performance of the full model for each constituent is also compared with that of the three models 481 calibrated to the pre-, during and post-drought periods. In general, we observe consistent performance for each constituent, across calibrations to the three periods of contrasting hydrological conditions 482 (Table 6, see Figs. S8 to S13 in the Supplementary Material for detailed model fittings). One notable 483 484 common pattern is that the performance for calibration and validation is more consistent during the drought period than either the pre- and post-drought periods. However, this is most likely explained by 485 relative sizes of the calibration data sets, which are 3, 13 and 5 years for the pre-, during and post-486 drought periods respectively. 487 Of all constituents (excluding FRP), TSS shows greater differences in model performances across 488 periods – especially when comparing the pre-drought calibration with its validation for the site-level 489

490 mean concentrations (Fig. 7). Notably, when calibrated to the pre-drought period and validated on both

491 the during- and post-drought periods, the validated model over-estimates most of the data (Fig. 7 (b));

- 492 and when calibrated to the during-drought period, the validated model slightly under-estimates pre-
- 493 and post-drought period TSS (Fig. 7 (d)).

Table 6. Comparison of model performances (as NSE) of the full model and the three models
that were calibrated to the pre-drought (1994-1996), drought (1997-2009) and the post-drought
(2010-2014) periods. For each of the models, the calibration performance is shown on the top
row and the validation performance (i.e. over the periods that were not used for calibration) is

shown on the bottom row. See Section 2.1.4 for details of the calibration and validation approach.



Figure 7. Comparison of the TSS model performance, as the simulated against observed site level mean concentrations in Box-Cox transformed space. The left column shows calibration
 performance for the model calibrated to the pre-drought (1994-1996), drought (1997-2009) and
 the post-drought (2010-2014) periods, respectively; the right column shows the corresponding
 validation performance for each period. The 95% lower and upper bounds of simulations shown
 in vertical grey lines and red dash lines show the 1:1 lines.

The potential impacts of drought on TSS dynamics are further illustrated with the performance of the
spatio-temporal model (calibrated to the full dataset with all sites and all data from 1994 to 2014) over

the pre-, during and post-drought periods (Fig. 8). Both the during- and post-drought periods have
consistently good performances, while the model underestimates most sites for the pre-drought period.
This is consistent with Fig. 7 in suggesting a systematic decrease in TSS concentration since the
drought began. The better performance of the full model during and after drought (Fig. 8) can be a
result of the calibration period of the full spatio-temporal model – between 1994 and 2014 – which
was dominated by the during- and post-drought periods.

516 In summary, Figs 7 and 8 together with Figs. S13-S17 suggest that whilst model performance for most

517 constituents are not affected by the hydrological periods used for calibration and validation, the

calibration period did have notable impact on TSS. Some possible causes are discussed in Section 4.3.



519

Figure 8. Comparison of the performance of the full spatio-temporal TSS model calibrated to all
data across a) pre-drought (1994-1996), b) during drought (1997-2009) and c) post-drought
(2010-2014) periods, as represented by the simulated against observed site-level mean
concentrations in Box-Cox transformed space. The 95% lower and upper bounds of simulations
shown in vertical grey lines and red dash lines show the 1:1 lines.

525 **4. Discussion**

526 4.1 Implications for statistical water quality modelling

In this study, we developed the first process-informed statistical model that is capable of explaining a reasonable proportion of water quality variability for a large spatial area of over 130,000km². Although the calibration data have relatively low sampling frequency (i.e. monthly), our model generally performs satisfactorily in explaining the total variability in water quality. This demonstrates the effectiveness of the Bayesian hierarchical modelling framework in predicting spatio-temporal variability in water quality across large scales. The Bayesian hierarchical model is: a) more advantageous than other simpler statistical water quality models with its more comprehensive and process-informed approach, and capacity to represent varying temporal relationships across large-scale regions; b) less demanding for input data compared with those required by fully-distributed, processes-based models. From a practical perspective, this model has the potential to contribute to a number of management activities including catchment planning, management and policy-making activities, specifically:

- The spatial predictive capacity can be used to identify pollution hot-spots and the catchment
 conditions that are likely causes of high concentrations. This can be used to help identify target
 catchment(s) to prioritize future water quality monitoring and management (Figs. 4 and 5);
- 541 2) Further to 1), since water quality has been linked with catchment characteristics in this model,
 542 it can also be used to assess potential impacts of alternative options of land use and land cover
 543 change, as well as potential effects of climate change, on ambient water quality conditions;
- 544 3) The model's temporal predictive capacity can identify changes in water quality due to changes 545 in hydro-climatic conditions and vegetation cover, and thus enabling attribution of detected 546 trends. On the other hand, any 'unexpected' trends can be identified to prompt further 547 investigation to identify causes (Figure 6 and Table 4). The model could also be used for 548 assessing the impacts of long-term catchment changes on water quality (Figures 7 and 8).

549 Despite the opportunities highlighted above, the model's performance also suggests some current 550 limitations of the modelling framework in the following situations:

1) *High within-site temporal variability*. In Section 3.2 we have identified a general lack of
predictive power for temporal variability. The potential impacts of high temporal variability on
model performance is particularly evident for results of TSS, NO_x and FRP in Fig. 3. Since our
model has already included hydro-climatic conditions and vegetation cover to explain temporal
variability, the unexplained temporal variability is likely due to other uncaptured temporal
drivers. These could be: changes in land use and land management, bio-geochemical processes,
or transit time of water through catchments.

558 2) Presence of high proportions of below-DL data. The full datasets for the three poorly modelled constituents (FRP, TSS and NO_x) all have higher proportions of data below the detection limit 559 560 (38.2% 17.3% and 15% of all data, respectively) compared with other constituents. As 561 illustrated in Fig. 2, for each of these constituents, removal of below-DL data before model calibration had created clear a truncation on the left-hand side of the distribution. This 562 563 substantially increases the degrees of skewness and discontinuity of the data, essentially violating the assumption of normally distributed residuals and thus limiting model performance. 564 565 The model capacity to handle truncated data might be improved by model fitting approaches explicitly designed for this issue. For example, Wang and Robertson (2011) and Zhao et al. 566 (2016) illustrated an approach to resolving the discontinuity of the likelihood estimation in 567 model fitting to data with presence of a lower bound such as zero rainfall values. 568

569 3) Non-conservativeness of constituents. The results indicate that the reactivity of the constituent is broadly associated with performance, which suggest that bio-geochemical processes (e.g. 570 phosphorus cycling, nitrification/de-nitrification) can make water quality dynamics more 571 572 difficult for the model to capture. To better capture changes in reactive constituents, the model may require greater consideration of and more extensive spatial and temporal data to represent 573 bio-geochemical processes. Examples include improvements on the process representation for 574 nitrogen cycling and the desorption and adsorption of phosphorus (Granger et al., 2010;Smyth 575 et al., 2013; Tian and Zhou, 2007). 576

As previously noted, our model was developed in a Box-Cox transformed scale to ensure the validity of the statistical assumptions (see details on data transformation in Sect. 2.1.2), which shows limited performance for high constituent concentrations when simulations are back-transformed to the measurement scale (Figs. 4 and 5). However, our model approximately represents proportional changes in water quality¹, which can thus help managers to understand proportional changes to inform practical catchment management.

583 For future implementations, the established model structure and parameterization would be best suited

584 to within the study region. Before performing new simulations (e.g. for new monitoring sites or for 585 current study sites over a different time-period), the statistical properties of the new input datasets should 586 be checked to ensure that they are similar to the calibration datasets. To model new catchments outside 587 of the study region, a re-calibration of the model is required. This would involve extensive selection of key predictors and model calibration, much as performed in this study and the two preceding ones 588 589 (Lintern et al., 2018b; Guo et al., 2019). A sufficiently long record length (e.g. 20 years) is ideal for such 590 modelling, as it ensures a reasonable understanding of the temporal variability to be obtained.

4.2 Implications for water quality monitoring programs 591

The current spatio-temporal model extracts water quality temporal variability from monthly data. 592 Utilizing data with higher temporal resolution may further strengthen the model capacity to explain 593 594 temporal variability, especially by capturing more information on water quality dynamics during flow events. This may be possible into the future; however, current high-frequency water quality sensors 595 (Bende-Michl and Hairsine, 2010;Outram et al., 2014;Lannergård et al., 2019;Pellerin et al., 2016) still 596 have very high resourcing requirements that limits widespread deployment in operational networks. 597

598 Furthermore, changes in land use and management over time are currently not considered here as predictors of temporal variability in water quality, which include but not limit to land clearing, 599 600 urbanization, tillage, fertiliser application and irrigation. This is due to a complete lack, or inconsistency 601 of available data. However, changes in land use/land management practices can occur over short time periods, which can lead to increases in pollutant sources and changes to runoff generation processes 602 603 (e.g. Tang et al., 2005;DeFries and Eshleman, 2004;Smith et al., 2013). Therefore, our modelling framework can potentially be improved by having additional monitoring data on the temporal patterns 604 605 of land use/land management to better capture their impacts on water quality.

606

4.3 Potential impacts of long-term drought on water quality dynamics

607 Results of model calibration and validation to different time periods suggest a systematic decrease in 608 TSS concentrations during and after the prolonged drought, in comparison with the pre-drought period under the same spatial and temporal conditions. Such a shift is not observed for any other five 609 610 constituents analyzed (nutrients and salts) (Section 3.3).

A further analysis of the calibrated model parameters for pre-, during and post-drought periods suggest 611 612 that the effects of key spatial predictors do not vary much across periods (Figure S14). In contrast, the effects of key temporal predictors highlight a clear shift in the role of antecedent flow (prior 7-day flow) 613 across different time periods (Figure 9). Specifically, the antecedent flow effects are mostly positive 614 615 across catchments before the drought, and shift to mostly negative during the drought. After the drought, the antecedent flow effects have mixed directions among different catchments. Considering the limited 616 617 performance of the TSS model (i.e. substantial under-estimation of temporal variability in Section 3.1), 618 these changing relationships suggested in the calibrated parameters might be unreliable. However, this 619 should not affect the reliability of the observed change in TSS since the drought (Section 3.3), which 620 was based on the systematic differences of model fitting between different periods, revealing a broadscale patterns across the state on the drought influences. 621



Figure 9. Effects of the five key predictors for the temporal variability in TSS across 102 sites,
 summarized by the posterior mean of the calibrated parameter values for each predictor (box
 shows values across all sites), from left: flow, 7-day antecedent flow, water temperature, root zone soil moisture and deep soil moisture.

In the literature, impacts of the Millennium Drought on the hydrology and runoff regimes of south-627 eastern Australia are well understood (van Dijk et al., 2013;Leblanc et al., 2012;Saft et al., 2015). 628 However, less is known about how this major and prolonged drought event has impacted water quality 629 (Bond et al., 2008). Previous studies on other drought events around the world mainly focused on 630 changes in water quality as responses to the reduced streamflow during drought. For example, reduction 631 in sediment levels during drought has been reported and attributed to lower erosion from the contributing 632 catchment, together with lower rates of solid transport associated with reduced flows (Murdoch et al., 633 634 2000; Caruso, 2002). At a more local scale, increasing sediment concentrations during drought have also been observed in streams adjacent to land with high densities of livestock and bushland, which both 635

constantly contribute to sediment load during drought, leading to elevated concentrations with lower 636 637 dilution rate (Caruso, 2002). Similar to sediments, the impact of droughts on stream nutrient and salt concentrations have also commonly been understood as responses to reduced runoff generation and 638 streamflow. In catchments with no significant point-source pollution, nutrient concentrations typically 639 decreased during droughts (Mosley, 2015) with less nutrient leaching and overland flow, but may also 640 increase due to increasing livestock inputs at more local scales (Caruso, 2002). In contrast, catchments 641 with significant point-source pollution generally experience water quality deterioration during drought 642 due to reduced dilution (van Vliet and Zwolsman, 2008:Mosley, 2015). For salinity, concentration often 643 644 increases during drought with reduced dilution and increased evaporation (Caruso, 2002). This is particularly evident for catchments that are more influenced by saline groundwater input as the relative 645 contribution of groundwater increased during drought (Costelloe et al., 2005). 646

647 In contrast to these previous studies, our findings suggest additional possible pathways along which drought can affect stream water quality, that prolonged drought might have altered the relationships 648 between sediments and its predictors (Figs. 7 and 8). In contrast to sediments, our model suggests no 649 clear shifts in the dynamics of nutrients and salts in a regional scale. Our findings are in line with a few 650 651 previous studies which reported temporal changes in the concentration-discharge relationships for sediments and nutrients, specifically, when comparing high- and low-flow conditions (Zhang, 652 653 2018; Moatar et al., 2017), as well as drought and recovery period (Burt et al., 2015). Our findings provide extra dimensions to what would be offered by simple trend analyses using approaches such as 654 Mann Kendall test or Sen's slope (e.g. Smith et al., 1987; Chang, 2008; Hirsch et al., 1991; Bouza-Deaño 655 et al., 2008). Those approaches are only capable of indicating direction and magnitude of observed 656 trends. In contrast, our model was able to attribute the consistent upward shift in TSS concentration to 657 change in relationships between water quality and its key driving factors since the start of drought. 658

In addition, we also acknowledge that our ability to represent the pre- and post-drought conditions in this study may be limited by the record length, since only 2 years of pre-drought and 4 years of postdrought data were available. Once longer records build up, they will enable us to update our understanding of the impact of this prolonged drought. We would be also able to conduct more sophisticated investigations, such as comparing the impacts of long-term droughts versus individual dry and wet years and events (e.g. Saft et al., 2015;Outram et al., 2014;Burt et al., 2015).

665 **5.** Conclusions

This study aims to address the current lack of water quality models that operate at large scales across 666 667 multiple catchments. To achieve this, we used long-term stream water quality data collected from 102 sites in south-eastern Australia, and developed a Bayesian hierarchical statistical model to simulate the 668 spatio-temporal variabilities in six key water quality constituents: TSS, TP, FRP, TKN, NO_x and EC. 669 The choice of model predictors was guided by previous studies on the same dataset (Lintern et al., 670 2018b; Guo et al., 2019). The model generally well captures the spatio-temporal variability in water 671 quality, where spatial variability between catchments is much better represented than temporal 672 variability. The model is best used to predict proportional changes in water quality in a Box-Cox 673 674 transformed scale, and can have substantial bias if used to predict absolute values for high concentrations. Cross-validation shows that the spatio-temporal model can predict water quality in non-675 monitored locations under similar conditions to the historical period and the calibration catchments that 676 we investigated. This can assist management by (1) identifying hot-spots and key temporal periods for 677 678 waterway pollution; (2) testing effects of catchment changes e.g. urbanization or afforestation; and (3) 679 identifying and attributing major water quality trends and changes.

Based on the above model evaluations, we discussed potential ways to further enhance the model 680 681 performance. In improving the modelling framework, alternative statistical approaches could be considered to reduce the impact of below detection limit data on model performance. In addition, the 682 models could be extended to consider some key bio-geochemical processes to better dynamics in non-683 conservative constituents (e.g., FRP or NO_x). Regarding data availability, the current models could 684 potentially benefit from improved monitoring of changes in land use intensity and management to be 685 able to include these drivers in the model. The inclusion of high-frequency water quality sampling data 686 may also extend the model's ability to represent temporal variability. However, high-frequency water 687 quality data are also typically highly variable with large noise. Therefore, the implication of such data 688 for the spatio-temporal modelling framework remains an open question, which needs further 689 investigation in future applications of this modeling framework. 690

691 Data availability

All data used in this study were extracted from public domain. All stream water quality data were extracted from the Victorian Water Measurement Information System (via <u>http://data.water.vic.gov.au/,</u> provided by the Department of Environment Land Water and Planning Victoria). The catchments corresponding to these water quality monitoring sites were delineated using the Geofabric tool provided by the Bureau of Meteorology, via <u>ftp://ftp.bom.gov.au/anon/home/geofabric/</u>. We have listed the sources of all other data for the spatial and temporal predictors of our models in Tables S1 and S2 in the Supplementary Materials.

699 Author contribution

All authors contributed to the conceptualization the models and the design of methodology. A. Lintern and S. Liu contributed to the data curatiom. D.Guo carried out the formal analyses, visualization and validation. J.A. Webb, D. Ryu, U. Bende-Michl and A.W. Western contributed to the funding acquisition. D. Guo, A. Lintern, J.A. Webb, D. Ryu, S. Liu and A.W. Western contributed to the investigation. D. Guo carried out project administration and coding to run the experiments. J.A. Webb, D. Ryu, and A.W. Western contributed to the supervision. D.Guo prepared the manuscript with contributions from all co-authors.

707 Competing interests

The authors declare that they have no conflict of interest.

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965 Tables

in Lintern et al. (2018) and Guo et al. (2019b), respectively. Constituent Key factors that affect spatial variability Kev factors that affect TSS Hottest month maximum temperature Same-day streamflow Percentage area covered by grass 7-day antecedent streamflow Percentage area covered by shrub Water temperature Percentage cropping area Soil moisture root Maximum elevation Soil moisture deep Dam storage Percentage clav area TP Erosivity Same-day streamflow Percentage area covered by grass 30-day antecedent streamflow Percentage area covered by shrub Water temperature Percentage area made up of roads Soil moisture root Percentage cropping area Soil moisture deep Average soil TP content FRP Percentage area covered by shrub Same-day streamflow Percentage cropping area Water temperature Catchment area Soil moisture deep Average soil TP content Mean channel slope Same-day streamflow TKN Percentage clay area Warmest quarter mean temperature 30-day antecedent streamflow NDVI Coldest quarter rainfall Percentage cropping area Water temperature Percentage pasture area Soil moisture root Average soil TP content Soil moisture deep NO_x Annual radiation Same-day streamflow 30-day antecedent streamflow Warm quarter rainfall NDVI Hottest month maximum temperature Average soil TP content Water temperature Mean channel slope Soil moisture root Soil moisture deep Same-day streamflow EC Annual radiation Annual rainfall 14-day antecedent streamflow Wettest quarter rain Water temperature Hottest month maximum temperature Soil moisture root Percentage agriculture area Soil moisture deep Percentage cropping area Percentage area covered by shrub Average soil TN content

Table 1. Key factors affecting the spatial and temporal variability for each of six constituents, as identified in Lintern et al. (2018) and Guo et al. (2019b), respectively.

968

969 Table 2. The key catchment landscape characteristics that are related to the varying relationships of water

970 quality and same-day streamflow across space, which were selected as the two predictors for the

971 streamflow effect in our model. Two characteristics were selected to summary the variability of

972 streamflow effects across space for each constituent, see Section 2.3 for details of the selection method. The

973 corresponding Spearman's correlation (R, at p<0.05) between the effect of streamflow and each

974 catchment characteristic is presented.

Constituent	Key factors that affect spatial variability in temporal effects	Spearman's ρ (p<0.05)
TSS	Annual rainfall	0.722
	Hottest month maximum temperature	-0.575
ТР	Annual rainfall	0.695
	Percentage area used for cropping	-0.556
FRP	Percentage agriculture area	0.392
	Percentage area underlain by mixed igneous bedrock	0.314
TKN	Annual rainfall	0.713
	Hottest month maximum temperature	-0.618
NOx	Total storage capacity of dams in catchment	-0.493
	Mean soil TN content	0.458
EC	Percentage area covered by grassland	-0.347

Percentage area covered by woodland	-0.317

976	Table 3. Comparison of model performance for all records and only the above-DL records for
977	each constituent.

Constituent	Above-DL records only	All records
TSS	0.225	0.397
ТР	0.433	0.445
FRP	-1.920	0.199
TKN	0.658	0.630
NO _x	0.216	0.382
EC	0.907	0.886

978

979

980 Table 4. Model ability to capture observed water quality trends across all monitoring sites for 981 each constituent. The percentages of sites where observed positive and negative trends are 982 captured by the model are presented separately. Values in brackets indicate numbers of sites 983 where corresponding positive or negative trends are observed. For detailed estimation of these 984 percentages please refer to Sect. 2.2.

Constituent	% positive trends captured	% negative trends captured
TSS	33.3 (12)	85.0 (20)
ТР	82.1 (28)	16.7 (12)
FRP	47.1 (17)	55.6 (9)
TKN	81.1 (37)	40.0 (10)
NOx	68.6 (35)	66.7 (27)
EC	82.6 (23)	77.3 (22)

985

986 Table 5. Comparison of model performances (as NSE) of the full model (Column 2) and the 50

- 987 partial models (Columns 3 to 5) with each calibrated to 80% randomly selected monitoring sites.
- 988 Columns 3 to 5 summarize the mean, minimum and maximum NSE values across the 50 runs,
- 989 where for each constituent, the top row showing calibration performance and the bottom row

Constituent	Full model	50 CV mean	50 CV min	50 CV max
TSS	0.225	0.413	0.376	0.439
		0.382	0.292	0.513
ТР	0.433	0.461	0.427	0.501
		0.411	0.151	0.575
FRP	-1.92	0.168	0.067	0.232
		0.129	-0.078	0.272
TKN	0.658	0.654	0.622	0.670
	_	0.622	0.468	0.691
NOx	0.216	0.453	0.414	0.489
		0.397	0.258	0.563
EC	0.907	0.893	0.882	0.903
		0.875	0.809	0.924

|--|

992	Table 3. Comparison of model performances (as NSE) of the full model and the three models
993	that were calibrated to the pre-drought (1994-1996), drought (1997-2009) and the post-drought

(2010-2014) periods. For each of the models, the calibration performance is shown on the top 994

row and the validation performance (i.e. over the periods that were not used for calibration) is shown on the bottom row.

Constituent	Full model	Pre-drought calibration	During drought calibration	Post-drought calibration
TSS	0.225	0.495	0.399	0.499
		0.208	0.402	0.390
TP	0.433	0.477	0.438	0.525
		0.421	0.474	0.411
FRP	-1.92	-1.336	0.187	0.204
		-1.406	0.197	0.024
TKN	0.658	0.649	0.650	0.711
		0.566	0.648	0.610
NOx	0.216	0.443	0.426	0.509
		0.394	0.471	0.393
EC	0.907	0.854	0.901	0.901
		0.887	0.873	0.884