



1 **Manuscript title**

2 Coffee and shade trees show complementary use of soil water in a traditional agroforestry
3 ecosystem

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19

20 **Abstract**

21 On a global scale, coffee has become one of the most sensitive commercial crops that will be
22 affected by climate change. The majority of Arabica coffee (*Coffea arabica*) grows in traditionally
23 shaded agroforestry systems and accounts for ~70% of the coffee production worldwide.
24 Nevertheless, the interaction between plant and soil water sources in these coffee plantations
25 remains poorly understood. To investigate the functional response of dominant shade trees species
26 and coffee (*C. arabica* var. *typica*) plants to different soil water availability conditions, we
27 conducted a study during a normal and more pronounced dry season (2014 and 2017, respectively)
28 and the 2017 wet season in a traditional agroecosystem in central Veracruz, Mexico. For the
29 different periods, we specifically investigated the variations in water sources and root water uptake
30 via MIXSIAR mixing models using $\delta^{18}\text{O}$ and $\delta^2\text{H}$ stable isotopes of rainfall, plant xylem and soil
31 water, along with micrometeorological and soil moisture measurements. To further increase our
32 mechanistic understanding about root activity, the distribution of belowground biomass and soil
33 macronutrients were also examined and considered in the model. Results showed that, over the
34 course of the two dry seasons investigated, all shade tree species (*Lonchocarpus guatemalensis*,



35 *Inga vera* and *Trema micrantha*) relied on water sources from deeper soil layers (>15 to 120 cm
36 depth; 86%), while the use of much shallower water sources (<15 cm depth; 60%) was observed in
37 coffee plants. In addition, in these same periods, coffee water uptake was strongly influenced by
38 antecedent precipitation conditions, whereas trees showed little sensitiveness to short-term wetness
39 status. Our findings also showed that during the wet season coffee plants substantially increased the
40 use of near surface water (+48% from <5 cm depth), while shade trees extended the water
41 acquisition to much shallower soil layers (+32% from <15 cm depth) in comparison to drier periods.
42 Despite the plasticity in soil water uptake observed among canopy trees and coffee plants, a spatial
43 segregation of the main water source prevailed during the dry and wet seasons studied. However,
44 more variability in plant-soil water uptake was observed among species in the rainy season when
45 higher soil moisture conditions were present and water limitation was largely absent.

46

47 *Key words:* Shade trees, *Coffea arabica*; water stable isotopes, roots, nutrients, MixSIAR, Mexico



48 **1. Introduction**

49 Coffee agroforestry systems are highly valued because of their ecological, environmental,
50 economic and social benefits (Mas and Dietsch, 2004; Perfecto et al., 2007; Tschardt et al.,
51 2011). Moreover, shade coffee of the species Arabica (*Coffea arabica*) accounts for ~ 70% of the
52 total coffee production (USDA, 2017). Although Arabica coffee is mainly grown in tropical
53 montane regions, it is cultivated under a wide range of climatic and soil conditions (Jha et al.,
54 2014). Coffee Arabica plantations can be broadly classified as traditional or modern coffee systems,
55 according to vegetation composition and structure and management practices (Moguel and Toledo,
56 1999). In the traditional systems, coffee plants are cultivated under a diverse canopy of native
57 and/or introduced shade tree species. In contrast, monoculture coffee plantations exemplify the
58 modern cultivation scheme, in which the shade is provided by a single commercial tree species. The
59 use of agrochemicals is also typically required in this type of plantation (Moguel and Toledo, 1999).

60 Until recently, the vast majority of Arabica coffee was cultivated in traditionally managed
61 shaded coffee plantations, which have lower production costs and enhanced biodiversity, carbon
62 sequestration, soil fertility and biological pest control in comparison to modern systems (Greenberg
63 et al., 1997; Perfecto et al., 2002; Kellermann et al., 2008). However, coffee management practices
64 worldwide have increasingly become more intensive promoting the replacement of native trees with
65 fast-growing monospecific timber species (i.e. *Cedrela Odorata*, *Eucalyptus deplupta*, *Hevea*
66 *brasiliensis*) (Nath et al., 2011).

67 Growing a crop in association with shade trees inevitably leads to some degree of
68 competition for the above-ground (light) and below-ground (water and nutrients) resources
69 (Monteith et al., 1991). In an agroforestry system, the outcome of competition for light is relatively
70 predictable due to the hierarchical structure of the canopy (i.e., shade trees intercept part of the
71 sunlight, thereby reducing the amount available for the understory crop). Conversely, competitive
72 interactions for soil resources can be much more diverse and complex. The central hypothesis of
73 agroforestry underscores that crops and trees are complementary in their use of soil resources
74 (Cannell et al., 1996), however the degree to which this occurs will be largely controlled by the
75 spatial and temporal patterns of resource availability, root distribution and root activity, which in
76 turn depend on factors such as climate, soil conditions, crop and tree species, and plantation age,
77 density and management practices (Beer et al., 1998; Lehmann, 2003; van Noordwijk et al., 2015).
78 In addition, below-ground competitive interactions for water and/or nutrients are much more
79 difficult to elucidate than above-ground relationships. So far, the most common approach is to
80 measure the distribution of root abundance of crops and trees, and examine to what extent they
81 overlap or are separated (e.g., Schaller et al., 2003; van Kanten et al., 2005). An important



82 limitation of this method is, however, that the spatial distribution of roots does not always mirror
83 the actual resource capture along the soil profile (Dawson et al., 2002; Lehmann, 2003). Another
84 approach is to examine the vertical patterns of soil water content (Cannavo et al., 2011; Padovan et
85 al., 2015) or nutrient (Schroth et al., 2000, cited in Lehmann, 2003) depletion. These methods are
86 problematic because they cannot provide information on whether resource depletion is caused by
87 the crop, the trees, or both (Cannavo et al., 2011; Padovan et al., 2015). However, the use of
88 hydrogen ($\delta^2\text{H}$) and oxygen ($\delta^{18}\text{O}$) water stable isotope techniques in combination with Bayesian
89 mixing models provides a powerful tool for quantifying the proportions of different water sources to
90 plant uptake across different ecosystems and regions (Barbeta et al., 2015; Muñoz-Villers et al.,
91 2018; Penna et al., 2018) and can largely overcome the above-mentioned limitations (Dawson et al.,
92 2002; Lehmann, 2003; van Noordwijk et al., 2015). Although rarely implemented, including
93 nutrient and root distribution data along the soil profile to inform these models could provide more
94 comprehensive insights into depth of plant water uptake.

95 To date, research into plant-soil interactions and plant water source partitioning in coffee
96 agroforestry systems is extremely scarce. To our knowledge, only five studies have investigated the
97 water sources of shade trees and coffee shrubs using either information on the isotopic composition
98 of plant xylem and bulk soil water (Wu et al., 2016), soil water depletion (Cannavo et al., 2011;
99 Padovan et al., 2015) or root distribution (Schaller et al., 2003; van Kanten et al., 2005). Moreover,
100 all of these studies have been carried out in intensive monospecific plantations characterized by
101 high coffee planting densities (4000–5000 shrubs ha⁻¹), low density (~150–280 trees ha⁻¹) and
102 very low diversity (1-2 species) of shade trees. While recognizing the limitations of some of the
103 methods used in these previous studies, the available information suggests that competition for
104 water between coffee and trees can be strong at sites with a pronounced seasonal dry period (Wu et
105 al., 2016; Padovan et al., 2015), while it seems to be virtually absent at sites with no or a relatively
106 short dry season (Schaller et al., 2003; Cannavo et al., 2011). Further, although most coffee roots
107 are usually located in the upper soil layers (< 30 cm depth; van Kanten et al., 2005, and references
108 therein), the plant and soil interactions for water during the dry season seem to occur below the
109 main crop rooting zone (> 30 cm depth) (Wu et al., 2016). The latter reflects the ability of coffee to
110 develop an extensive root system, and to increase the root water uptake at greater soil depths once
111 the available water has been depleted in shallower layers (Huxley et al., 1974, cited in Lehmann,
112 2003).

113 Currently, we lack of information on plant water sources in traditional shade coffee
114 plantations. In these agroforestry systems, the higher density and diversity of shade trees could
115 potentially lead to stronger and more diverse tree-crop interactions (van Noordwijk et al., 2015). On



116 the other hand, the dense tree canopy reduces light availability limiting coffee water use. This could
117 lead to a lower soil water demand and thus increased plant water availability during the dry season.

118 Further, ecohydrological research in these shade coffee systems is becoming increasingly
119 important since trees have been promoted as a strategy for mitigating and adapting to future climate
120 (Schroth et al., 2009; Vaast et al., 2016; Rice, 2018). Shaded coffee plantations store more carbon
121 than sun-grown coffee systems, thereby contributing to the reduction of greenhouse gases (Vaast et
122 al., 2016; Rice, 2018, and references therein). In addition, the tree canopy provides some level of
123 protection against the rising mean and maximum air temperatures (Baker and Hagggar, 2007;
124 Schroth et al., 2009; Vaast et al., 2016), which in recent modeling studies have been pointed out as
125 the key climatic changes affecting coffee growth, yield and quality (Schroth et al., 2009; Baca et al.,
126 2014; Bunn et al., 2015). Although there are important differences across sites, rainfall is also
127 predicted to decrease and become more variable in many of the world's coffee-growing regions. For
128 example, Giorgi (2006) estimated that rainfall will decrease by about 17% (per 100 years) during
129 the dry season and by about 9% during the wet season in Mexico and Central America. Similarly,
130 predictions by Karmalkar et al. (2011) for the same region pointed out changes in rainfall of -24%
131 to +8% (per 100 years) during the dry season and of -39% to -1% during the wet season (with the
132 range reflecting variability among regions). As such, if warming is accompanied by decreases in
133 rainfall, this could lead to, or exacerbate, competition for water sources between coffee shrubs and
134 shade trees (Baker and Hagggar, 2007), which in turn could affect the long-term sustainability of
135 these agroecosystems.

136 Mexico is among the largest shade coffee producers in the world, and the central region of
137 Veracruz constitutes the second most important coffee zone in the country. In this area, we selected
138 a representative traditional shade coffee plantation to investigate plant water sources of dominant
139 shade trees species and coffee (*C. arabica* var. *typica*) shrubs under different conditions of soil
140 water availability. Hence, during a normal and more pronounced dry season (2014 and 2017,
141 respectively) and the 2017 wet season, variations in depth of plant water uptake were examined
142 using the stable isotopic composition ($\delta^{18}\text{O}$ and $\delta^2\text{H}$) of rainfall, plant xylem and soil water in
143 combination with a Bayesian mixing model (MixSIAR), along with micrometeorological and soil
144 moisture measurements. To further increase our understanding about root activity and water uptake,
145 the distribution of roots and macronutrients along the soil profile were also examined and
146 considered in the mixing model as prior information. Specifically, we addressed the following
147 questions:

148



- 149 1. Does a complementary water use strategy between shade trees and coffee shrubs prevail in
150 a traditional shaded agroforestry system?
151 2. Does competition exist for water sources among tree and coffee species during more
152 pronounced dry periods?
153 3. What are the seasonal patterns in plant-water source partitioning?
154

155 **2. Materials and methods**

156 *2.1 Study site*

157 The research was carried out in the “La Orduña” coffee plantation (~100 ha) located on a
158 flat plateau at an elevation of 1210 m a.s.l. on the eastern slopes of the Cofre de Perote mountain
159 (19°28' N, 96°56' W) in central Veracruz State, Mexico (Fig. 1). The coffee plantations in this
160 region occur between elevations of 1000 and 1350 m a.s.l. (Hernández-Martínez et al., 2013;
161 Marchal and Palma, 1985).

162 The climate is classified as temperate humid with abundant rains during the summer
163 (García, 1988). Two distinct seasons can be distinguished: (1) a wet season (May–October), during
164 which rainfall is associated primarily with cumulus and cumulonimbus clouds formed during
165 convective and orographic uplift of the moist maritime air masses brought in by the easterly trade
166 winds; and (2) a (relatively) dry season (November–April), during which most rainfall falls from
167 stratus clouds associated with the passage of cold fronts (Báez et al., 1997). Mean annual rainfall
168 measured nearby the study site during the period 1971–2000 was 1765 mm, with on average
169 monthly rainfall of 389 mm falling during the dry season and 1376 mm falling during the wet
170 season (SMN, 2018). Mean annual temperature over this period was 19.5 °C, with a minimum and
171 maximum monthly average value of 15.5 and 22.5°C observed in January and May, respectively
172 (SMN, 2018). Annual potential evapotranspiration (ET₀) is about 1120 mm (Holwerda et al., 2013).

173 The investigated shade coffee plantation is a so-called traditional commercial polyculture
174 system (*sensu* Moguel and Toledo, 1999), which was established more than 80 years ago. The tree
175 canopy is diverse and consists predominantly of the species *Inga spp.*, *Citrus spp.*, *Lonchocarpus*
176 *guatemalensis*, *Trema micrantha* and *Enterolobium cyclocarpum* (Holwerda et al., 2016). The
177 shade trees were planted at a density of ca. 500 ha⁻¹, and currently form a canopy of about 14 m
178 high. The Arabica coffee plants were of the variety *typica*. *Typica* –a tall cultivar of *Coffea*
179 *arabica*– was the first coffee variety that arrived from Ethiopia to Mexico (Renard, 2010); it has
180 bronze-tipped young leaves and the berries are large. Plants of *typica* variety are tolerant to
181 conditions of low soil fertility and drought, but vulnerable to most pests and diseases (Escamilla et
182 al., 2005). In the study site, this cultivar was planted approximately 20 years ago at a density of



183 about 1700 shrubs ha⁻¹, currently having an average height of ~ 2 m. In this region, the coffee
184 flowering occurs in March or April, fruit development between May and October, and ripening and
185 harvest between October and February (Villers et al., 2009). The management of the plantation
186 involves weed control practices and selective pruning of mature coffee plants and shade trees at
187 irregular times once every ~ 7 years (cf. Hernández-Martínez et al., 2009). No pruning activities
188 occurred during or in between our study periods. A photograph of the coffee plantation is shown in
189 the Supplementary Material.

190 The soil type is an Andic Acrisol derived from volcanic ashes. Soil profiles (~150 cm) are
191 multilayered (A, B1/BT and BC) and have clay (~ 65%) as the dominant texture across all layers. A
192 general description of the soil profile showed a dark brown to dark yellowish brown, clay silty
193 organic A horizon (0–20 cm) overlying a dark yellowish brown, clay silty sand B1/BT horizon (20–
194 135 cm), followed by a dark yellowish brown, clay sandy BC horizon (>135 cm). Average soil bulk
195 densities and porosities were 1.2 gr cm⁻³ and 63%, respectively, along the A and B horizons
196 (Holwerda et al., 2013). The underlying material consists of deeply weathered old lava and sandy-
197 gravelly pyroclastic flow deposits (Rodríguez et al., 2010). Soils are mostly covered by a thin (1-2
198 cm) but continuous layer of litter.

199

200 **2.2 Hydrometeorological measurements**

201 During the study period, rainfall and microclimate conditions were continuously monitored
202 above the canopy in an 18 m high tower, located in the southwestern part of the coffee plantation.
203 Rainfall (P , mm) was measured using a TR-525 M tipping bucket rain gauge (Texas Electronics,
204 USA). Temperature (T , °C) and relative humidity (RH, %) were measured using a HC2-S3 probe
205 (Rotronic, USA). Data were recorded every 30 s, accumulated (P) or averaged values (all other
206 parameters) were stored at 5 min intervals using a CR1000 datalogger (Campbell Scientific Ltd.,
207 USA).

208

209 **2.3 Isotope sampling**

210 To examine the water sources of overstory shade trees and understory coffee shrubs, plant
211 tissue and soil samples were collected for isotope analysis at the middle (Jan. 23) and end (Apr. 11
212 and 26) of the 2014 dry season. In 2017, the dry season was warmer and drier offering the
213 opportunity to examine the vegetation responses to more pronounced dry conditions. Therefore, a
214 second sampling campaign was performed to collect plant and bulk soil samples at the middle (Feb.
215 27), end (Apr. 5) and late end (May. 20) of the 2017 dry season. Another sampling was carried out



216 in the middle of the 2017 wet season (Aug. 4) to evaluate plant-soil water uptake patterns at higher
217 soil water availability conditions.

218 In all seven samplings, xylem samples were obtained from three individuals of each of the
219 three dominant shade tree species (*Lonchocarpus guatemalensis*, *Inga vera* and *Trema micrantha*)
220 by extracting ~5-6 cm cores using a Pressler increment borer inserted at 1.2 m above ground. On
221 each occasion, xylem samples were taken from the same individuals but from various aspects of the
222 trunk. The bark was immediately removed after core extraction to avoid contamination of phloem
223 water. For the coffee plants, samples were obtained from ~6 cm segments of mature branches that
224 were cut near the main stem of several shrubs each time. All coffee plants were sampled randomly.
225 During the 2014 and 2017 dry seasons, sampling of coffee shrubs involved 5-6 individuals each
226 time. Since only one sampling occasion was performed during the 2017 wet season, a larger number
227 of individuals (10) was sampled to reduce the uncertainties associated with different sampling sizes
228 between wet and dry seasons respectively. For each tree, we measured diameter at breast height
229 (DBH) and height, and for the coffee plants the diameter of the main stem was measured below its
230 bifurcation in small branches (Table 1).

231 Bulk soil samples were collected at three locations and at 5, 15, 30, 60, 90 and 120 cm
232 depth each, using a hand auger. Auger sampling points were located so that each of the sampled
233 shade trees and coffee plants had a total of three soil sampling points within their 3 m radius.

234 Samples of xylem and bulk soil were collected during the morning and early afternoon
235 (between 8:30 to 13:30 hrs), and each sampling campaign was preceded by at least 6 days up to 22
236 days without or with minimum accumulated rainfall (< 5 mm). All xylem and soil samples were
237 collected quickly and carefully and contained in water-tight vials to avoid any evaporation (see
238 section below).

239 To establish the local meteoric water line and compare soil water sources with recent
240 rainfall, bulk samples of rainfall were collected weekly at a nearby (~ 5 km) meteorological station
241 over the course of four years (December 2013 to December 2017) as part of a long-term isotope
242 sampling of precipitation (cf. Muñoz-Villers et al., 2018).

243

244 **2.4 Isotope collection and analysis**

245 Samples of precipitation, plant xylem and bulk soil for isotope analysis were collected in
246 30-ml borosilicate glass vials sealed with polycone caps to prevent evaporation. All samples were
247 refrigerated until extraction and analysis at the Center of Stable Isotope Biogeochemistry (CSIB) at
248 the University of California-Berkeley, USA.



249 Xylem and soil samples were extracted using cryogenic vacuum distillation (temperature:
250 $100 \pm 1.1^\circ\text{C}$, vacuum: 3 ± 1.5 Pa and time: 60-70 min) following the method of West et al. (2006).
251 The $\delta^2\text{H}$ and $\delta^{18}\text{O}$ isotopic compositions of extracted water samples were then determined using an
252 isotope-ratio mass spectrometer (Thermo Delta Plus XL, Thermo Fisher Scientific, USA). The
253 analytical precision of the instrument was $\pm 0.60\text{‰}$ (1 SD) for $\delta^2\text{H}$ and $\pm 0.12\text{‰}$ (1 SD) for $\delta^{18}\text{O}$.
254 Samples of precipitation were analyzed for $\delta^2\text{H}$ and $\delta^{18}\text{O}$ using a laser water isotope analyzer
255 (L2140-i) from Picarro Inc. (Santa Clara, CA, USA) in high precision and without Micro-
256 Combustion Module mode. The analytical precision was $\pm 0.65\text{‰}$ (1 SD) and $\pm 0.20\text{‰}$ (1 SD) for
257 $\delta^2\text{H}$ and $\delta^{18}\text{O}$, respectively.

258 The isotope values are expressed in delta notation (‰) relative to Vienna Standard Mean
259 Ocean Water (VSMOW). To evaluate evaporative enrichment in the soil and xylem water isotopes
260 relative to rainfall, we calculated the deuterium-excess parameter ($d = \delta^2\text{H} - 8 * \delta^{18}\text{O}$; Dansgaard,
261 1964).

262

263 **2.5 Soil sampling and laboratory determinations**

264 To determine volumetric soil water content (SWC), samples were collected at 5, 15, 30, 60,
265 90 and 120 cm depth from each of the three boreholes excavated during the soil isotope samplings.
266 Soil moisture content was determined gravimetrically and converted to volumetric values by using
267 bulk density of the soil sample. In addition, to determine the antecedent moisture conditions for the
268 15 days prior to each sampling date, an antecedent precipitation index (API) was calculated
269 following Viessman et al. (1989).

270 To examine pH and N, P and K macronutrient concentrations along the soil profile, soil
271 samples were collected at 5, 15, 30, 60, 90 and 120 cm depth from each borehole ($n = 3$ samples per
272 soil depth) during three sampling campaigns: Apr. 11, 2014 (dry season), Feb. 27, 2017 (dry
273 season) and Aug. 4, 2017 (wet season). Samples were first air-dried and then sieved using 2 mm
274 screens. Soil pH was determined using a glass electrode pH meter in a 1:2 soil: water ratio. Total
275 carbon (C) and total nitrogen (N) were measured using a TruSpec dry combustion CN analyzer
276 (LECO, USA). Extractable phosphorus (P) was determined by the Bray I method (Bray and Kurtz,
277 1945). Exchangeable K^+ was determined by extracting soil with 1 M NH_4OAc at pH 7.0, and then
278 analyzed using flame photometry.

279

280 **2.6 Root biomass**

281 To examine the root biomass distribution along the soil profile in the study plot, 33 soil
282 cores were obtained using 5 cm diameter and 10 cm long samplers. Soil cores were extracted at



283 5, 20, 40, 60 and 90 cm depth (from 5 to 40 cm: $n = 9$ for each depth, and from 60 to 90 cm: $n = 3$
284 for each depth). All cores were processed immediately in the laboratory. Soil samples were first
285 sieved using 2 mm screens to separate the bigger roots. Next, the samples were washed using a fine
286 nylon mesh sieve, and then separated into diameter classes (< 1 mm, 1–2 mm and > 2 mm) and
287 dried at 70 °C for 48 hours. Root biomass (g m^{-3}) was calculated from the dry weight of the roots
288 and the volume of the core sampler for each class and soil depth.

289

290 ***2.7 Plant water uptake sources and temporal patterns***

291 The MixSIAR Bayesian mixing model framework (Moore and Semmens, 2008; Stock et
292 al., 2018) was used to determine the most likely contributions of water sources for the shade tree
293 species and coffee shrubs sampled over the course of the 2014 (Jan. 23, Apr. 11 and 26) and 2017
294 (Feb. 27, Apr. 5, May. 20) dry seasons and the 2017 wet season (Aug. 4). To assess temporal
295 changes of the different plant water sources, the seven sampling occasions were modeled separately.
296 The mixture data for the model was the mean xylem water isotopic composition of the shade tree
297 species and coffee shrubs, changing accordingly with the sampling date. The relative contribution of
298 three potential xylem end-member water sources was evaluated. These included rainfall as surrogate
299 for near surface water (< 5 cm depth), shallow soil water (average of 5–15 cm depth) and deep soil
300 water (average of 30–120 cm depth). For each sampling date, the mean and standard deviation of the
301 isotopic signatures of the water sources were introduced into the model as follows: rain water
302 isotope data from a month prior to the xylem sampling and soil water isotope data from the two
303 different grouped soil depths, all corresponding to the date of xylem tissue collection.

304 Further, we also considered the use of additional data such as soil macronutrients (N, P, K)
305 and root biomass information to constrain model estimates by specifying an ‘informative’ prior
306 distribution of the soil source proportions (Stock et al., 2018). These data were grouped into two
307 classes based on the depth of the soil samplings and corresponding largely with the grouping for
308 soil water: shallow (0–20 cm) and deep (40–120 cm). In addition, the nearest corresponding dry or
309 wet season dataset of soil macronutrients were used according to the date of sampling. The effect of
310 using these priors (i.e. a weight proportion before considering the isotope data) on the water sources
311 distribution was then examined by comparing these with the results of ‘non-informative’ (i.e. all the
312 combinations of proportions of water sources were equally likely) simulations. The results of each
313 of these model runs were accepted based on the examination of Markov Chain Monte Carlo
314 convergence using the Gelman-Rubin and Geweke diagnostic tests (Gelman et al., 2014).

315 Lastly, the relative contributions of the water sources were compared among shade trees
316 and coffee shrubs across all sampling dates using factorial ANOVA and Tukey’s HSD post-hoc



317 tests. The analyses were carried out in R Statistical Software version 3.2.4 (R Core Development
318 Team, 2016).

319

320 **3. Results**

321 **3.1 Hydrometeorological conditions**

322 Precipitation (P) was 1650 mm in the first study year (Nov. 2013 – Oct. 2014) and 1423
323 mm in the second study year (Nov. 2016 – Oct. 2017). During the 2013-2014 dry season (Nov –
324 Apr.), rainfall was 323 mm, and mean daily values of temperature (T) and vapor pressure deficit
325 (VPD) were $17.6 \pm 3.0^\circ\text{C}$ and 0.65 ± 0.39 kPa, respectively. The lowest monthly amounts of P and
326 the highest values of T and VPD were observed in April at the end of the dry season (Fig. 2a,b).
327 During the 2016-2017 dry season, rainfall was 235 mm, with lowest monthly values registered in
328 January and February at the middle of the season (Fig. 2b). Mean daily T was $18.3 \pm 2.6^\circ\text{C}$, with
329 highest values observed at the end of the dry period. Overall, high VPD values prevailed over the
330 course of this dry season (0.78 ± 0.46 kPa on average), although maximum values were particularly
331 observed in February and May.

332 Compared to long-term (1971–2000) climatic records of the region, rainfall in the first year
333 of study was very close to the mean annual precipitation of 1765 mm (SMN, 2018). In contrast, the
334 second year was drier (~ 300 mm less); the difference was particularly observed during the dry
335 season, which had about 40% lower precipitation than the average value of 389 mm. Also, higher
336 mean monthly temperatures ($+ 0.54^\circ\text{C}$) prevailed across the 2017 dry season in comparison with the
337 1971–2000 period.

338 Rainfall during the 2017 wet season (May – Oct.) was lower in comparison to 2014 (1188
339 mm vs. 1326 mm, respectively) (Fig. 2b). Further, the mean air temperature and vapor pressure
340 deficit were slightly higher in the 2017 wet season than in the 2014 wet season ($20.7 \pm 1.6^\circ\text{C}$ and
341 0.67 ± 0.25 kPa vs. $20.1 \pm 1.5^\circ\text{C}$ and 0.60 ± 0.21 kPa, respectively) (Fig. 2a).

342

343 **3.2 Soil moisture and antecedent precipitation during sampling campaigns**

344 During the 2014 dry season campaign (Jan. – Apr.), mean soil water content (SWC) was on
345 average $39.8 \pm 6.7\%$ in the shallower layers (5-15 cm depth) and $46.0 \pm 5.2\%$ in the deeper layers
346 (30-120 cm depth) (Fig. 2b). In comparison, SWC in the 2017 dry season campaign (Feb. – May.)
347 was lower in the shallower layers ($32.5 \pm 5.8\%$), meanwhile water content in the deeper layers was
348 similar ($48.9 \pm 2.9\%$) with respect to the 2014 dry period. In 2014, lowest SWC values were
349 observed at the end of the dry season (April), whereas the greatest soil moisture depletion in 2017
350 was registered at the middle of the dry season (February) (Fig. 2b).



351 During the wet season sampling in August 2017, higher SWC values in the shallower (35.0
352 $\pm 3.7\%$) and deeper ($48.9 \pm 7.2\%$) soil layers were generally observed in comparison to the 2017
353 dry period (Fig. 2b). For the different samplings, antecedent wetness conditions were, respectively,
354 4, 30 and 13 mm for Jan. 23, Apr. 11 and 26, 2014 and 1, 12, 9 and 43 mm for Feb. 27, Apr. 5,
355 May. 20 and Aug. 4, 2017.

356

357 **3.3 Stable isotope composition of waters**

358 Over the study periods, a greater range of variation was found in the rainfall isotope
359 composition of the 2013-2014-year (from -126.7 to 14.4% $\delta^2\text{H}$; from -17.7 to 0.01% $\delta^{18}\text{O}$) in
360 comparison to the 2016-2017-year (from -113.3 to 15.5% $\delta^2\text{H}$; from -15.9 to 0.01% $\delta^{18}\text{O}$) ($p >$
361 0.05)(Fig. 3). Overall, mean dry season rainfall was significantly more enriched than the mean wet
362 season rainfall in $\delta^2\text{H}$ and $\delta^{18}\text{O}$ ($p \leq 0.001$) (Table 2 and 3). In the second study year, the isotopic
363 compositions of the dry and wet season rainfall were both on average more depleted than the first
364 study year; thus the local meteoric water line of 2016-2017 had a slightly steeper slope in
365 comparison to the one for 2013-2014 (Fig. 3). Nevertheless, the range of variation of deuterium
366 excess values was similar between years (9 – 29% for the first year vs. 9 – 31% for the second year;
367 Fig. 3), and deuterium excess values of rainfall between the dry and wet seasons were not
368 statistically different ($p \geq 0.05$).

369 For all sampling dates, hydrogen and oxygen isotope composition of bulk soil water showed
370 a consistent pattern of increasing isotope depletion with depth (Supplementary Materials), in which
371 shallower (5-15 cm) soil water was significantly more enriched than deeper (30-120 cm) soil water
372 ($p \leq 0.001$) (Table 2 and 3; Fig. 3). In correspondence, lower values of deuterium excess generally
373 characterized the shallower soil water pool.

374 For the 2014 dry season samplings, bulk soil ranged from -83.3 to -11.9% for $\delta^2\text{H}$ and
375 from -11.1 to -0.9% for $\delta^{18}\text{O}$ (Fig. 3a). For the 2017 dry season samplings, bulk soil water showed
376 a narrower range of variation and more enriched isotope values (from -54.8 to -19.1% $\delta^2\text{H}$ and
377 from -7.5 to -1.5% $\delta^{18}\text{O}$) in comparison to 2014 (Fig. 3b). However, statistical differences were
378 only suggested for the deeper soil layers in both water isotopes between the two dry seasons
379 investigated ($p \leq 0.001$).

380 In the 2017 wet season sampling, bulk soil isotope composition ranged from -70.5 to $-$
381 37.5% for $\delta^2\text{H}$ and from -8.4 to -4.1% for $\delta^{18}\text{O}$ (Fig. 3c), showing significant differences in the
382 shallow and deep soil water pools in comparison to 2017 dry season ($p \leq 0.001$). In all sampling
383 periods, bulk soil water across depths was isotopically distinct from rainfall during the 2014 and the
384 2017 dry seasons ($p \leq 0.001$ for both water isotopes).



385 Across all sampling periods, xylem water of coffee shrubs was more enriched than that of
386 shade trees ($p \leq 0.001$) (Table 2 and 3), and values of $\delta^2\text{H}$ and $\delta^{18}\text{O}$ plant xylem ($-40.8 \pm 15.0\text{‰}$
387 and $-4.6 \pm 1.6\text{‰}$, respectively) were on average more positive in comparison to bulk soil water ($-$
388 $46.7 \pm 16.4\text{‰}$ and $-6.0 \pm 2.3\text{‰}$, respectively) ($p > 0.05$) (Fig. 3).

389 In the 2014 dry season, xylem water isotope values of shade trees ranged from -65.5 to $-$
390 32.1‰ for $\delta^2\text{H}$ and from -7.6 to -3.6‰ for $\delta^{18}\text{O}$, meanwhile a larger variation was observed in the
391 xylem water of coffee shrubs (from -46.5 to -9.6‰ $\delta^2\text{H}$ and from -6.3 to -0.6‰ $\delta^{18}\text{O}$) ($p \leq 0.001$)
392 (Fig. 3a). Among tree species, *Lonchocarpus guatemalensis* showed the most depleted xylem water
393 isotope signature ($-58.1 \pm 4.8\text{‰}$ $\delta^2\text{H}$ and $-6.8 \pm 0.5\text{‰}$ $\delta^{18}\text{O}$), whereas *Inga vera* reported the most
394 enriched values with a greater range of variation ($-51.0 \pm 10.2\text{‰}$ $\delta^2\text{H}$ and $-5.3 \pm 1.1\text{‰}$ $\delta^{18}\text{O}$).
395 Intermediate $\delta^2\text{H}$ and $\delta^{18}\text{O}$ values were observed in *Trema micrantha* ($-57.1 \pm 5.4\text{‰}$ and $-6.6 \pm$
396 0.6‰ , respectively) (Fig. 4a). Statistical tests showed that *Inga vera* was significantly different
397 from *L. guatemalensis* and *T. micrantha* in $\delta^{18}\text{O}$ ($p < 0.05$).

398 In the 2017 dry season, the isotopic composition of shade trees varied from -56.7 to $-$
399 34.5‰ for $\delta^2\text{H}$ and from -6.0 to -3.2‰ for $\delta^{18}\text{O}$; corresponding values for coffee shrubs varied
400 from -39.6 to -7.8‰ for $\delta^2\text{H}$ and from -4.4 to -1.1‰ for $\delta^{18}\text{O}$ ($p \leq 0.001$) (Fig. 3b). *L.*
401 *guatemalensis* showed the most enriched isotope value ($-41.3 \pm 5.7\text{‰}$ $\delta^2\text{H}$ and $-4.6 \pm 0.5\text{‰}$ $\delta^{18}\text{O}$),
402 and *I. vera* reported the most depleted values ($-48.5 \pm 5.1\text{‰}$ $\delta^2\text{H}$ and $-4.8 \pm 0.8\text{‰}$ $\delta^{18}\text{O}$), with
403 differences being suggested for $\delta^2\text{H}$ ($p < 0.05$). Intermediate $\delta^2\text{H}$ and $\delta^{18}\text{O}$ values were observed in
404 the xylem water of *T. micrantha* ($-45.9 \pm 3.6\text{‰}$ and $-3.9 \pm 0.6\text{‰}$, respectively), showing
405 differences in $\delta^{18}\text{O}$ with the other two species ($p < 0.05$).

406 Overall, more enriched isotope values of plant xylem water were observed in the 2017 dry
407 season in comparison to those in 2014 ($p \leq 0.001$) (Fig. 3a,b; Fig. 4). Also, lower deuterium excess
408 values were obtained in shade trees and coffee shrubs in 2017 as sign of a more evaporative
409 signature (Table 2 and 3; Fig. 3). Plots of $\delta^2\text{H}$ xylem water against height for the individual shade
410 trees and coffee shrubs sampled in both dry seasons are shown in Figure 4, in which a similar $\delta^2\text{H}$
411 pattern was displayed between trees and coffee shrubs in the 2014 and 2017 years.

412 During the 2017 wet season sampling, more depleted $\delta^2\text{H}$ and $\delta^{18}\text{O}$ values in xylem water of
413 trees and coffee shrubs were observed in comparison to the 2017 dry season ($p < 0.05$) (Fig. 3c).
414 The range of variation was from -60.6 to -45.6‰ $\delta^2\text{H}$ and -6.2 to -5.4‰ $\delta^{18}\text{O}$ for trees, and from
415 -42.2 to -34.4‰ $\delta^2\text{H}$ and -5.4 to -4.4‰ $\delta^{18}\text{O}$ for coffee shrubs ($p \leq 0.001$).

416

417 **3.4 Root biomass and macronutrients along soils profile**



418 Overall, most roots were concentrated in the first 5 cm of soil with a sharp decline in
419 biomass at 20 cm depth (Fig. 5a). Fine roots (< 1mm) followed by bigger roots (> 2 mm) dominated
420 the shallower soil layers (< 20 cm), meanwhile roots in general were scarce at deeper depths (> 60
421 cm). Soil acidity was higher at the near surface layers decreasing gradually as soil depth increased
422 (Table 4). Organic matter (OM) was also greatest between 5 and 15 cm depth, however values
423 decreased rapidly below ~30 to 60 cm depth. Highest concentrations of nitrogen were found in the
424 first 15 cm, although values remain relatively high and constant at deeper layers (Fig. 5b).
425 Phosphorus showed its highest concentration at the topsoil with values decreasing sharply below 30
426 cm depth. In contrast, lowest concentrations of potassium were found at the near surface layers (<
427 15 cm depth) while maximum values were observed at 120 cm depth.

428

429 **3.5 Plant water sources**

430 In general, there was a good agreement between the MixSIAR Bayesian mixing model
431 results using a non-informative and an informative prior distribution (on average 5% difference
432 across all xylem water contributing sources; $p > 0.05$). This indicates that the independent
433 distribution (soil macronutrients and root data) set *a priori* to optimize model source proportion
434 estimates (informative approach) was not influential enough to significantly modify the results
435 obtained using the isotope signatures of the xylem end-member water sources alone (non-
436 informative approach). Having this agreement between models, we present the results of the water
437 source contribution based on the informative prior distribution. Results of the non-informative
438 approach have been provided in the Supplementary Materials.

439 The model results showed that the deep soil water pool was the main source for the shade
440 trees over the course of the 2014 dry season (Fig. 6 and Supplementary Materials). Across this
441 period, *L. guatemalensis* and *T. micrantha* showed on average higher proportion of deep soil water
442 uptake ($91 \pm 12\%$ and $88 \pm 10\%$, respectively) in comparison with *I. vera* ($74 \pm 14\%$) ($p > 0.05$). For
443 the coffee plants, the analysis showed that water uptake was mainly sustained by shallow soil water
444 sources ($65 \pm 23\%$) (Fig. 6), having significant differences with *L. guatemalensis* ($p < 0.001$), *T.*
445 *micrantha* ($p < 0.001$) and *I. vera* ($p < 0.05$) tree species across the 2014 dry period studied.

446 During the 2017 dry season, all shade tree species were tapping high proportions of deep
447 soil water ($86 \pm 13\%$ for *L. guatemalensis*; $85 \pm 19\%$ for *T. micrantha* and $92 \pm 12\%$ for *I. vera*;
448 Fig. 7a,b,c and Supplementary Materials). Among samplings, differences were only suggested
449 between *L. guatemalensis* and *I. vera* at the end of the dry period (Apr. 5) ($p < 0.05$). Coffee water
450 sources were mainly obtained from shallow soil layers ($54 \pm 29\%$) (Fig. 7a,b,c), although in 2017
451 the contribution of this water source was slightly smaller (9%) in comparison to 2014



452 (Supplementary Materials), statistical differences were not suggested between dry periods ($p >$
453 0.05).

454 Across the individual samplings throughout the two dry seasons, we observed that
455 antecedent precipitation had a stronger effect in the water uptake sources of coffee plants than trees
456 (Fig. 8). For example, when dry antecedent wetness prevailed ($API_{15} < 5$ mm; Fig. 2b) coffee water
457 sources were mainly composed by deep ($46 \pm 23\%$) and shallow ($38 \pm 35\%$) soil water, meanwhile
458 when wetter antecedent conditions were present ($API_{15} > 10$ mm) the shallower soil water layer (63
459 $\pm 22\%$) was the main contributing source. On the contrary, tree water uptake was essentially
460 sustained by deep soil sources ($91 \pm 13\%$ and $80 \pm 15\%$, respectively) at low and relatively high
461 antecedent wetness conditions (Fig. 8). For all species investigated, the relationships between API
462 and the contribution of deep soil water sources were not found statistically significant ($p > 0.05$).

463 In the 2017 wet season, sources of tree water uptake differed significantly among species
464 (Fig. 7d and Supplementary Materials). During this period, *L. guatemalensis* and *I. vera* still
465 showed the greatest use of deep soil water ($64 \pm 37\%$ and $70 \pm 36\%$, respectively) ($p > 0.05$),
466 meanwhile shallow soil water was the main source for *T. micrantha* ($72 \pm 39\%$; Fig. 7d), having
467 differences with the other two tree species ($p < 0.001$). Coffee consistently showed the use of near
468 surface water sources ($69 \pm 22\%$) (Fig. 7d and Supplementary Materials), being significantly
469 different in relation to all shade tree species ($p < 0.001$).

470

471 **4. Discussion**

472 **4.1 Complementary water use strategy between shade trees and coffee shrubs**

473 Our ecohydrological research consistently showed that all shade tree species (*L.*
474 *guatemalensis*, *I. vera* and *T. micrantha*) relied mainly on water sources from deep soil layers (> 15
475 to 120 cm depth), while the use of much shallower water sources (< 15 cm) was observed in the
476 coffee (*C. arabica* var. *typica*) over the course of the normal and more pronounced dry seasons
477 studied. These findings suggest a spatial and temporal partitioning of belowground water resources
478 between trees and coffee plants during drier periods and water-resource complementary in the
479 mixed species plantings. Although comparisons with other traditional shade coffee plantations are
480 difficult because studies are essentially lacking in this type of agroecosystems, there are a handful
481 of other investigations carried out in shade coffee monospecific plantations in the humid tropics in
482 which complementary rather than competitive water use strategies prevailed. For example, Cannavo
483 et al. (2011) compared the water use and soil water availability of an unshaded coffee vs. a shaded
484 monoculture (*Inga densiflora*) coffee plantation in Costa Rica, both of 7-8 years old, using soil
485 moisture measurements and water balance calculations. Their results showed that soil water content



486 in the deeper soil layers (> 120 cm depth) was lower in the shaded coffee than in the full sun coffee
487 system, while water content in the shallower layers was similar. This suggested that associated
488 shade trees preferentially used water from deeper soil horizons providing some evidence of
489 complementarity water use between coffee plants and native *Inga* trees during the dry season.
490 However, the authors acknowledged that they were unable to separate roots from coffee than those
491 of trees in the soil profiles, so they could not be certain whether trees were the only individuals
492 extracting water from deeper sources. In this respect, our study showed that there was always a
493 mixture in water uptake from different sources (soil depths), but a separation between the main
494 sources of water for shade trees and coffee shrubs clearly prevailed.

495 Other investigations in Costa Rica have examined the belowground resource competition of
496 Arabica coffee in association with fast-growing timber species using data of plant growth, root
497 distribution and density, and soil moisture and nutrients patterns. For example, the study of Schaller
498 et al. (2003) carried out in a commercial (*Eucalyptus deplupta*) shade coffee plantation where soils
499 are highly fertilized, showed that coffee had a relatively even root distribution along the first 40 cm
500 of soil depth with a higher root density in the proximity of the coffee rows. Conversely, the root
501 system of *E. deplupta* was much shallower having most roots concentrated in the upper 10 cm of
502 soil. In this case, the tree root density was found highest in the alleys between the coffee rows. The
503 authors explained that the apparent complementary resource exploitation of this tree-crop system
504 was mainly attributed to high soil resources availability and the high competitiveness of the coffee
505 limiting the expansion of tree roots (cf. Lehmann, 2003). Although in our study we did not
506 determine the depth distribution of coffee and tree roots, our findings showed that all shade tree
507 species were tapping water from deeper soil layers than coffee, suggesting that trees are deep rooted
508 and being able to explore larger soil volumes causing little competition with coffee.

509 In Nicaragua, Padovan et al. (2015) compared the root distribution, soil moisture,
510 transpiration and leaf water potential patterns in an unshaded coffee plantation and an agroforestry
511 system of coffee planted with two timber trees (deciduous *Tabebuia rosea* and evergreen
512 *Simarouba glauca*). Their findings showed that coffee roots were more abundant than tree roots and
513 mainly concentrated in the shallower soil layers (0–80 cm depth). Most roots of both tree species
514 were observed in deeper layers (>100 cm) suggesting a clear niche differentiation with coffee.
515 During the 3-year study period, volumetric water content along a 2 m soil profile was higher in the
516 full sun grown coffee than in the shaded coffee, which was explained by greater soil water uptake
517 from trees below the crop rooting zone (Padovan et al., 2015). Moreover, coffee shrubs in the
518 shaded plantation were more water stressed (i.e. lowest midday leaf water potentials) during the
519 pronounced dry season studied (Padovan et al., 2018). These results suggest that despite the clear



520 hydrological niche segregation, competition between coffee and shade trees may occur if the
521 season is long and severe enough.

522 Our findings also showed that during the wet season coffee plants substantially increased
523 the use of near surface water (~70%) in comparison to the dry season, while all shade trees also
524 extended their water acquisition to much shallower soil water pools. This is largely explained by the
525 increases in soil moisture in the first 30 cm depth due to frequent rainfall inputs that characterize the
526 wet season in our study area. This suggests that coffee had a higher root activity in surface soil
527 layers during the wet season in comparison to the dry season, as has been documented in other
528 studies (Huxley et al., 1974). Regarding the shade tree species, we observed that *T. micrantha*
529 showed the greatest response to the wet season conditions by drawing most water from the first 15
530 cm of soil (72%), whereas this was much less evident in *L. guatemalensis* (30%) and *I. vera* (29%).
531 Although we did not determine the vertical distribution of roots for each of the shade tree species
532 studied, these findings suggest that *T. micrantha* has a shallower rooting system than the other tree
533 species. The fact that the *T. micrantha* trees were more recently planted (i.e. younger with less
534 developed root system) than the *L. guatemalensis* and *I. vera* trees supports this idea. On other
535 hand, the high temperature and rainfall that characterize the wet season at our study site may favor
536 rapid mineralization of nutrients and their subsequent leaching to deeper soil layers (i.e. potassium;
537 Table 4). Hence, for the larger trees studied (*L. guatemalensis*), the water and nutrients available at
538 deeper depths could have been an important resource for plant growth in this period, partly
539 explaining the lower activity of their shallow roots. Despite the changes and the higher variability in
540 depth of water uptake observed among canopy trees and coffee shrubs, a complementary use of soil
541 water prevailed during the wet season. Future work should be focused on the distribution and
542 dynamics of tree and crop roots and their seasonal variation in relation to the availability of
543 nutrients and water in the soil. Also, it would be desirable to relate these dynamics to crop and
544 shade tree phenology to elucidate temporal synergistic or competitive water requirements.

545

546 ***4.2 The role of antecedent wetness in coffee water uptake***

547 Our study showed that antecedent wetness strongly influenced the water uptake patterns of
548 coffee plants (cf. Huxley et al., 1974). We found that under relatively wet antecedent conditions
549 prevailing after small rainfall events during the dry season, coffee substantially increased the use of
550 shallower soil water sources, possibly as an opportunistic strategy to overcome the soil water
551 deficits in this period and taking advantage of their much shallower rooting system compared to
552 trees. Conversely, tree water uptake was mainly sourced by deep soil water showing less
553 sensitiveness to higher antecedent wetness. In this respect there are no comparative studies in shade



554 coffee agroecosystems evaluating short-term dynamics of plant water sources at different wetness
555 conditions. Nevertheless, plant and soil water interactions under dry and relatively wet conditions
556 have been examined in other types of agroforestry systems. For example, in the study of Gao et al.
557 (2018) carried out in a semiarid region in China, the authors evaluated the seasonal variations in
558 water use of jujube (*Ziziphus jujuba*) trees planted with annual (*Brassica napus*) and perennial
559 (*Hemerocallis fulva*) crops under various soil wetness status. Using stable isotope techniques and
560 Bayesian mixing modelling, their results showed that jujube trees generally tapped water (> 58%)
561 from deep soil layers (60-200 cm depth) at low antecedent wetness, while *B. napus* and *H. fulva*
562 crops primarily extracted water (> 65%) from intermediate (20-60 cm) and shallow (0-20 cm) soil
563 layers. This exhibits a complementary water use strategy between trees and crops. However, at
564 higher antecedent precipitation conditions both the jujube trees and the inter-row crops extracted
565 most water from the first 0-60 cm of soil depth (> 65%). This indicated that both species exhibited
566 an opportunistic strategy for accessing resources at shallower soil depths. In this case, contrary to
567 our findings, tree roots rather than crop roots showed the stronger capacity to switch rapidly from
568 deep to shallow sources in response to increased soil water availability.

569

570 **4.3 Implications and future directions**

571 The consistent complementarity in plant water use strategies observed under different
572 hydrometeorological conditions in the coffee plantation studied provides support to the central tenet
573 of agroforestry systems (Cannel et al., 1996). Since the trees obtained their water from deeper soil
574 layers than the coffee, this could mean that they utilize nutrients leaching beyond the reach of the
575 coffee plants, and so contribute to improved nutrient cycling and increased overall productivity of
576 the system (van Noordwijk et al., 2015).

577 Nevertheless, the current outcome may change given the new coffee management practices
578 that consist on replacing traditional coffee varieties (e.g. *C. arabica* var. *typica*) with others (*C.*
579 *arabica* var. *costa rica*; *C. canephora*) that may exhibit deeper roots systems and perhaps different
580 water (and nutrient) uptake strategies, in response to prevalent diseases such as leaf rust or root
581 nematodes. Therefore, future research should be focused on evaluating the water source partitioning
582 of traditional vs. new coffee varieties and their relation to shade tree water use. In this respect, there
583 are further questions with regard to strategic use of shade tree species, whereby fast-growing
584 species might be more (commercially) productive but also more competitive. Some evidence from
585 elsewhere has shown that such management practices do not necessarily increase competition and
586 may even enhance the water use efficiency as part of drought-avoidance mechanisms. For example,
587 in southeast China, Wu et al. (2016) used $\delta^2\text{H}$ and $\delta^{18}\text{O}$ stable isotope methods to examine the



588 seasonal water use of a fast-growing rubber tree species (*Hevea brasiliensis*) planted with Arabica
589 coffee. Their findings showed that rubber trees were mostly accessing water from intermediate (15-
590 50 cm depth) and deep soil layers (50-110 cm), meanwhile coffee was mostly tapping water from
591 the topsoil (< 15 cm). Additionally, rubber trees showed strong root plasticity in soil water uptake
592 avoiding competition with coffee during the rainy and relatively dry seasons. However, more
593 research is needed since these results depend largely on tree-crop specie combinations and local
594 climatic and soil conditions.

595 In addition to effects of changing management practices, climate warming may induce
596 changes in plant transpiration throughout the year (e.g. Karmalkar et al., 2011). In our study, we
597 used a water stable isotope approach along with root and macronutrients data to estimate the
598 relatively contribution of the plant water sources. However, for a more complete assessment of the
599 plant and soil interactions, seasonal plant water fluxes need to be quantified. Our results so far have
600 made the first steps towards serving coffee producers to make better decisions on sustainable coffee
601 and water management, as well as providing new insights into water resources in general, which are
602 urgently required for implementing efficient and equitable management programs in humid tropical
603 environments (Hamel et al., 2018). However, future work should be focused on water use of
604 individual trees and coffee shrubs using ecophysiological and hydrological techniques in order to
605 better understand how much water is used and where from.

606 Finally, in our methodology we used prior information alongside the stable water isotope
607 approach to better understand plant water uptake dynamics. Even though our results did not change
608 significantly by including or excluding the root and nutrient data, exploring plant water source
609 partitioning using these two approaches provided more confidence in our results. We would
610 recommend that other authors also consider using nutrient and root data in combination with plant
611 xylem water end members to better understand water uptake patterns, especially to explore the
612 additional value of this information in different environments.

613

614 **5. Conclusions**

615 This study provides the first baseline information on plant water sources for a traditional
616 shade coffee plantation in the humid tropics. Our results showed that coffee water uptake was
617 mainly sustained from shallow soil sources (<15 cm depth) while all shade trees relied on water
618 sources from deeper soil layers (>15 to 120 cm depth). This complementary strategy in
619 belowground resource use between crops and trees was consistent over the course of the normal and
620 more pronounced dry seasons investigated. Across these same periods, we observed that antecedent
621 precipitation had a strong influence in coffee plants increasing their water uptake to shallower soil



622 water sources as an opportunistic strategy to overcome the reduced water availability. In the wet
623 season, coffee plants substantially increased the use of near surface water (<5 cm depth), whereas
624 shade trees expanded their water acquisition to the first 15 cm soil depth. Overall, a greater soil
625 water partitioning prevailed among tree and coffee species when higher soil moisture conditions
626 were present. Nevertheless, despite such variability in plant-soil water interactions across seasons, a
627 clear spatial segregation of the main water source prevailed between trees and crops during the
628 rainy and dry periods investigated.

629

630 **Author contributions.** LEMV designed the experiment. LEMV, MSAB and FH collected the field
631 data. MSAB performed the Bayesian mixing model analysis. JG contributed in the data analysis.
632 LEMV prepared the first draft of the manuscript. FH, MSAB and JG edited and commented on the
633 manuscript several times. TED carried out the final revision.

634

635 **Competing interests.** The authors declare that they have no conflict of interest.

636

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653

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Table 1. Characteristics of the shade trees and coffee plants sampled for water isotope analysis during 2014 and 2017. Numbers between parentheses are the standard deviation.

Family	Species	Canopy layer	2014		2017		<i>n</i>
			DBH cm	Height m	DBH cm	Height m	
Fabaceae	<i>Lonchocarpus guatemalensis</i>	Overstory	101.5 (12.6)	20.3 (1.3)	119.8 (12.1)	21.0 (1.2)	3
Fabaceae	<i>Inga vera</i>	Overstory	39.3 (15.7)	10.7 (4.8)	48.1 (13.3)	9.6 (1.2)	3
Cannabaceae	<i>Trema micrantha</i>	Overstory	13.16 (6.8)	8.15 (3.1)	23.3 (7.2)	15.2 (2.2)	3
Rubiaceae	<i>Coffea arabica</i> var. <i>typica</i>	Understory	12.7 (2.1)	2.83 (0.7)	n.a.	n.a.	5* 6** 10***

* Number of individuals sampled each time in the 2014 dry season
 ** Number of individuals sampled each time in the 2017 dry season
 *** Number of individuals sampled in the 2017 wet season



Table 2. Mean \pm (SD) H and O stable isotope composition of 2013-2014 precipitation, tree xylem water and bulk soil water of the 2014 dry season sampling, and corresponding deuterium excess values (‰)

Precipitation <i>n</i> = 41						Bulk soil water <i>n</i> = 54						Shade trees xylem water <i>n</i> = 27			Coffee shrubs xylem water <i>n</i> = 14		
Dry season			Wet season			5-15 cm depth			30-120 cm depth								
$\delta^2\text{H}$	$\delta^{18}\text{O}$	<i>d</i> -excess	$\delta^2\text{H}$	$\delta^{18}\text{O}$	<i>d</i> -excess	$\delta^2\text{H}$	$\delta^{18}\text{O}$	<i>d</i> -excess	$\delta^2\text{H}$	$\delta^{18}\text{O}$	<i>d</i> -excess	$\delta^2\text{H}$	$\delta^{18}\text{O}$	<i>d</i> -excess	$\delta^2\text{H}$	$\delta^{18}\text{O}$	<i>d</i> -excess
1.6 \pm 8.5	-1.9 \pm 1.4	17.0 \pm 5.1	-42.4 \pm 36.1	-7.2 \pm 4.3	14.9 \pm 2.8	- 25.6 \pm 9.9	-3.0 \pm 1.2	-1.3 \pm 5.2	-63.8 \pm 10.4	-8.3 \pm 1.4	2.0 \pm 5.0	-55.4 \pm 7.6	-6.2 \pm 1.0	-5.8 \pm 4.1	-25.5 \pm 10.8	-3.4 \pm 1.8	1.7 \pm 5.0



Table 3. Mean \pm (SD) H and O stable isotope composition of 2016-2017 precipitation, tree xylem water and bulk soil water of 2017 dry season sampling, and corresponding deuterium excess values (‰)

Precipitation <i>n</i> = 39						Bulk soil water <i>n</i> = 54						Shade trees xylem water <i>n</i> = 24			Coffee shrubs xylem water <i>n</i> = 18		
Dry season			Wet season			5-15 cm depth			30-120 cm depth								
$\delta^2\text{H}$	$\delta^{18}\text{O}$	<i>d</i> -excess	$\delta^2\text{H}$	$\delta^{18}\text{O}$	<i>d</i> -excess	$\delta^2\text{H}$	$\delta^{18}\text{O}$	<i>d</i> -excess	$\delta^2\text{H}$	$\delta^{18}\text{O}$	<i>d</i> -excess	$\delta^2\text{H}$	$\delta^{18}\text{O}$	<i>d</i> -excess	$\delta^2\text{H}$	$\delta^{18}\text{O}$	<i>d</i> -excess
-2.9 ± 16.0	-3.0 ± 1.8	21.5 ± 4.3	-47.8 ± 34.4	-7.9 ± 4.1	15.2 ± 3.3	-28.6 ± 5.9	-3.0 ± 0.9	-4.9 ± 5.4	-45.3 ± 4.7	-6.2 ± 0.6	4.6 ± 3.6	-44.9 ± 5.6	-4.4 ± 0.7	-9.7 ± 5.4	-21.3 ± 7.2	-2.8 ± 1.0	1.3 ± 6.2



Table 4. Soil characteristics (average values) determined at the different depths

Soil depth (cm)	pH (H ₂ O)	OM (%)	C (%)	N (%)	K (cmol _c kg ⁻¹)	P (mg kg ⁻¹)
5	4.07	5.18	2.54	0.38	0.60	33.33
15	4.12	2.89	1.90	0.30	0.47	4.60
30	4.34	1.55	1.31	0.23	0.77	n.d.
60	4.95	1.02	0.69	0.22	0.93	n.d.
90	5.10	0.48	0.50	0.20	1.11	n.d.
120	5.16	0.41	0.51	0.20	1.45	n.d.

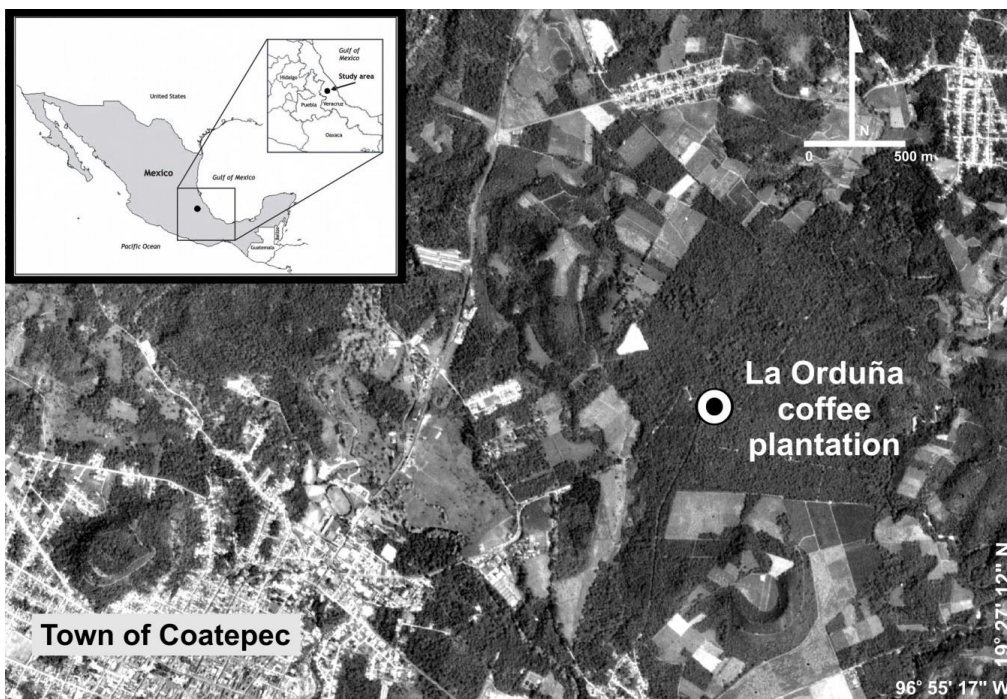


Figure 1. Study site location in the municipality of Coatepec, Veracruz, Mexico. Source: QuickBird Satellite Image (2010). Copyright DigitalGlobe, Inc.

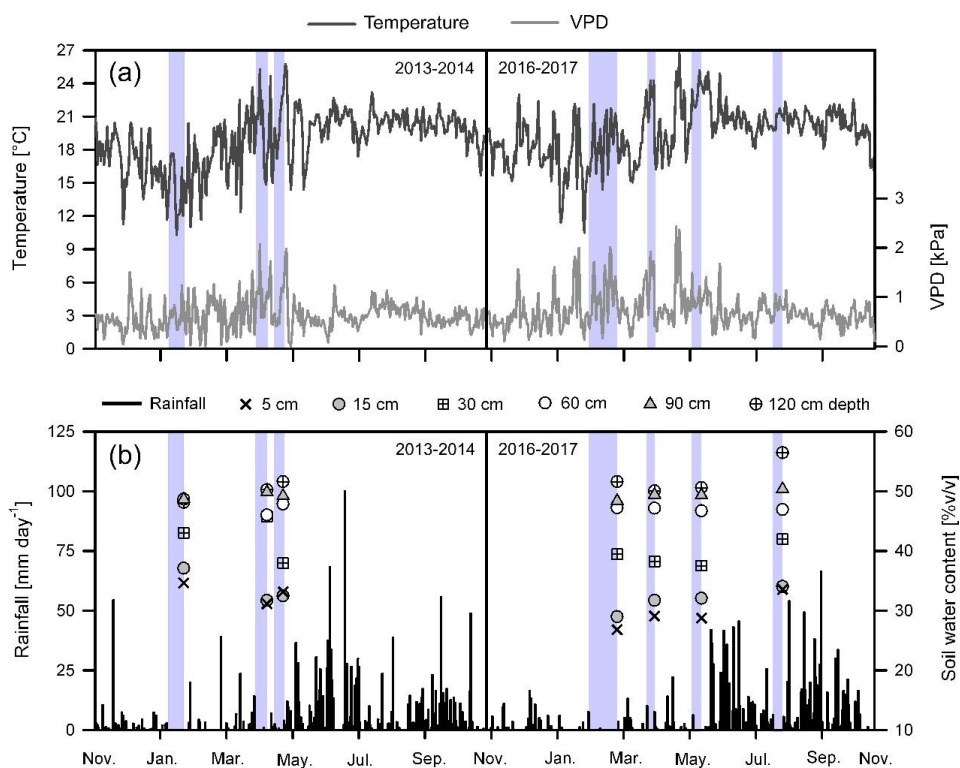


Figure 2. (a) Daily mean air temperature and vapor pressure deficit (VPD) and (b) and daily total rainfall (P), as measured from November 2013 to October 2014 and from November 2016 to October 2017, and volumetric soil water content (SWC) measured at different depths during the sampling campaigns in the study area; different depths are indicated by the unique symbols shown in the lower panels (the key to the symbols is at top). The blue-colored areas indicate the 6- to 22-day period of minimum rainfall (< 5 mm) preceding the dates of isotope sampling in January (mid dry season) and April (late dry season) of 2014, and in February (mid dry season), April and May (late and end of dry season), and August (mid wet season) of 2017.

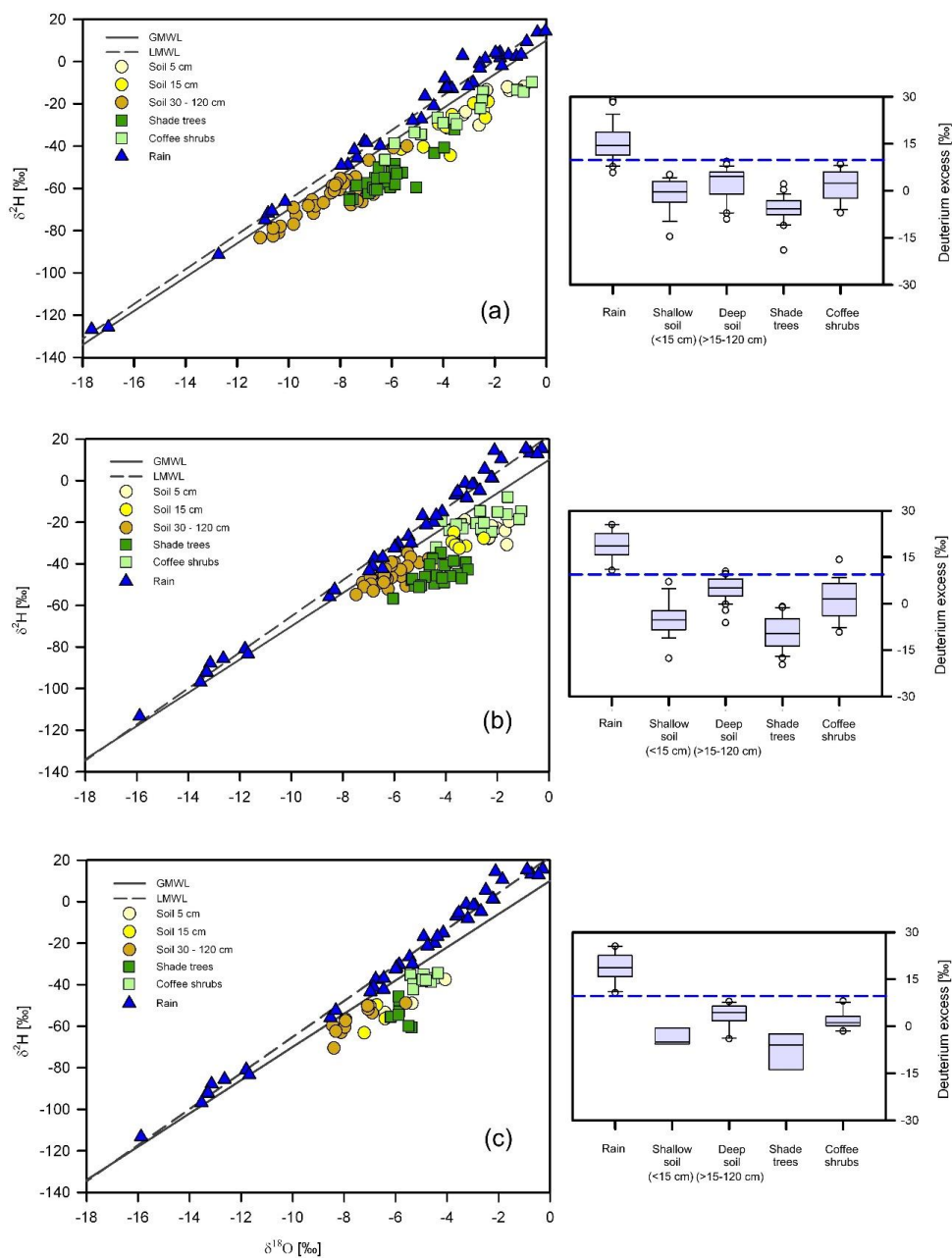


Figure 3. (a) Isotope composition of xylem water for shade trees and coffee shrubs, bulk soil at different depths as observed during the three sampling dates (Jan. 23, Apr. 11 and Apr. 26, 2014),



and isotope values of rainfall during the period December 2013 to November 2014. The dashed line represents the 2013–2014 local meteoric water line (LMWL; $\delta^2\text{H} = 17.82 + 8.26 * \delta^{18}\text{O}$), (b) Isotope composition of xylem water for shade trees and coffee shrubs, bulk soil at different depths during the three sampling dates (Feb. 27, Apr. 5 and May. 20, 2017) and isotope values of rainfall during the period December 2016 to November 2017. The dashed line represents the 2016–2017 local meteoric water line (LMWL; $\delta^2\text{H} = 21.0 + 8.36 * \delta^{18}\text{O}$), and (c) Isotope composition of xylem water for shade trees and coffee shrubs, bulk soil at different depths during the middle of the 2017 wet season (Aug. 4) and isotope values of rainfall during the period December 2016 to November 2017. The solid line represents the global meteoric water line (GMWL; $\delta^2\text{H} = 10 + 8 * \delta^{18}\text{O}$). The panels on the right show the deuterium excess values for the plants and soil water sources and rainfall preceding the sampling campaigns. The dashed blue line represents the deuterium excess value of the GMWL.

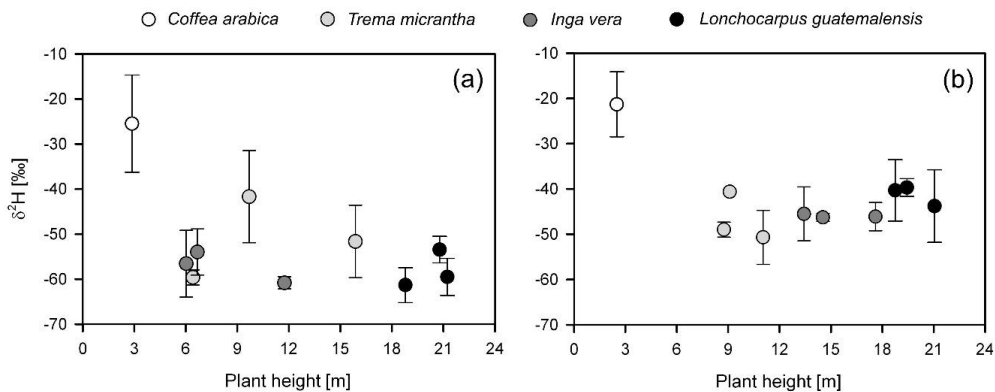


Figure 4. Plant height vs $\delta^2\text{H}$ xylem water for coffee plants and shade tree species corresponding to (a) the 2014 and (b) 2017 dry season samplings.

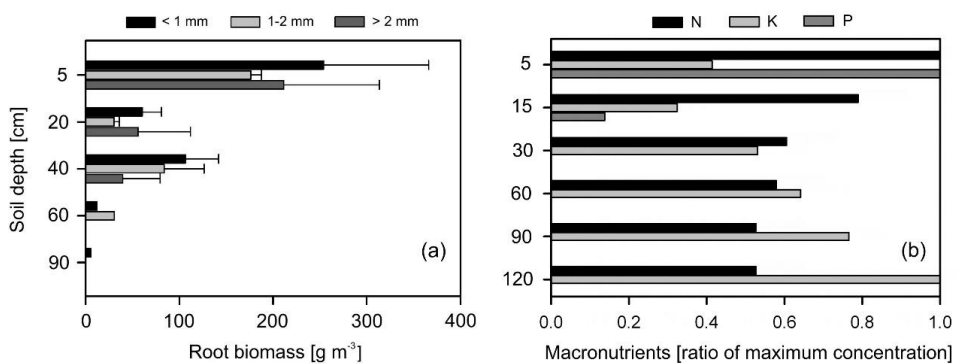


Figure 5. (a) Distribution of root biomass for three size classes of roots (different color bars), and (b) macronutrients distribution along the soil profile, here normalized and expressed as in ratio to their maximum values (absolute values in Table 4).

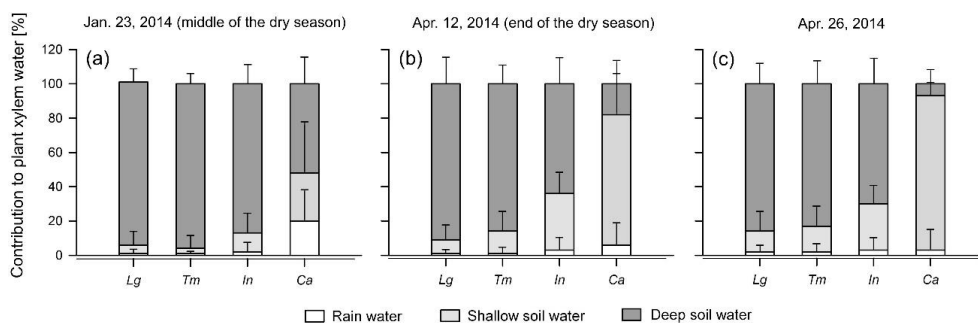


Figure 6. MixSIAR Bayesian mixing model results showing the mean likely contribution of each water source to the xylem water of shade canopy trees and coffee shrubs. (a), (b) and (c) show results for the sampling dates of Jan. 23, Apr. 12 and Apr. 26, 2014 respectively, using the informative prior distribution. *Lg*: *L. guatemalensis*; *Tm*: *T. micrantha*; *In*: *I. vera* and *Ca*: *Coffea Arabica*. Error bars represent one standard deviation of uncertainty.

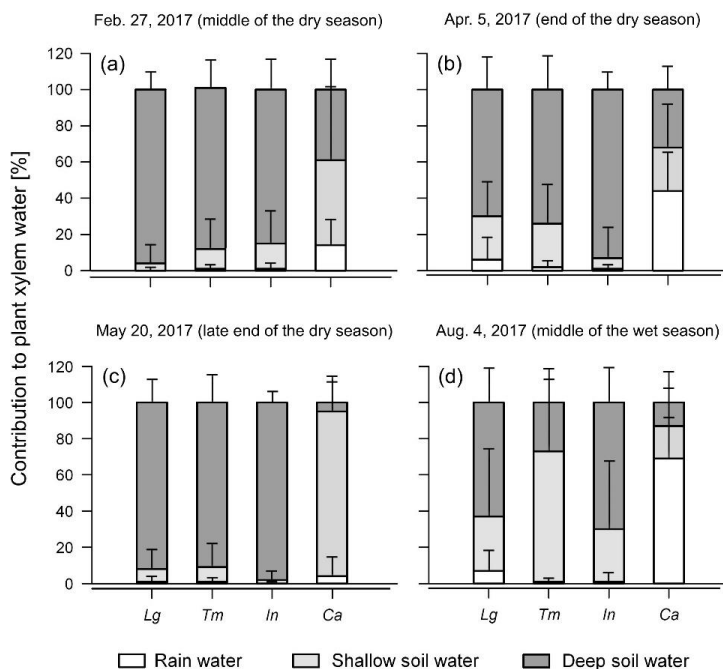


Figure 7. MixSIAR Bayesian mixing model results showing the mean likely contribution of each water source to the xylem water of shade canopy trees and coffee shrubs. (a), (b), (c) and (d) show results for the sampling dates of Feb. 27, Apr. 5, May. 20 and Aug. 4, 2017 respectively, using the informative prior distribution. *Lg*: *L. guatemalensis*; *Tm*: *T. micrantha*; *In*: *I. vera* and *Ca*: *Coffea Arabica*. Error bars represent one standard deviation of uncertainty.

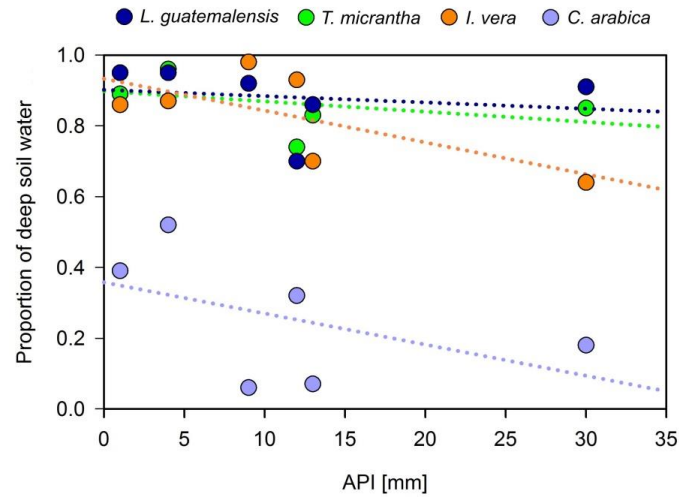


Figure 8. Contribution of deep soil water to plant uptake at different antecedent precipitation conditions across the 2014 and 2017 dry seasons.