

Please find below our reply to the corrections:

1.) Should it be "two investigated dry seasons"?

Reply: The change has been made (L33)

2.) Line 313: Is this a typo or why were the grouping of isotope ratio (Line 305) and macronutrients and roots done using different depth groups? How can this be then used in the same mixing analysis?

Reply: Yes, this is a typo. Thank you for the observation. The correction has been made (L313-314)

.) Line 469: Should it probably be "sampling dates"?

Reply: Correct. The change has been made (L469).

.) Line 522: references are not correct formatted

Reply: The changes have been made (L522).

.) 621: Change to "On the other hand"

Reply: The change has been made (L621)

.) Line 649: Is there a ", respectively" missing at the end of this sentence? If not, consider adding "both" in this sentence to clarify that intermediate and shallow applies to both species (*B. napus* and *H. fulva*).

Reply: We added "both" to make this clarification (L648).

.) Line 653: You refer to crops, but use "contrary to our findings". This can be a bit misleading since you did not study crops. Please rephrase.

Reply: We studied the coffee, which is a perennial crop, planted with shade trees. Therefore, we do not think that the sentence is misleading and requires to be rephrased.

.) Table 4: Two decimals shown for Clay at 5 cm, while all other have one decimal points.

Reply: The correction has been made.

.) You refer to the findings of Poca et al. (2019) when discussing the potentially fractionated coffee xylem water in Figure 3b. However, the mycorrhiza seem to cause a depletion in $\delta^2\text{H}$ according to Poca et al. (2019, see their Fig. 3). So, can you really support your findings by Poca's findings?

Reply: We referred to the work of Poca et al. (2019) as an example of the potential influence of mycorrhizal fungi on xylem isotope fractionation. However, we agree that their findings showed a depletion effect rather than an enrichment effect in the deuterium isotopic composition. Therefore, we have decided to remove this citation from the text (L549-550).

.) Figure 8: I recognize that you changed the y-axis to "Proportion of near surface soil water", but the caption is still with "deep soil water".

Reply: The correction has been made.

.) Figure 8: Reconsider if regression lines should be shown. I suggest to leave them out, since the relationship shown is not significant.

Reply: Despite the correlation was not statistically significant, we would like to keep the regression lines in the figure because it helps to highlight the differences across the investigated species.

1 **Manuscript title**

2 Coffee and shade trees show complementary use of soil water in a traditional agroforestry
3 ecosystem

4
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18
19
20 **Abstract**

21 Globally, coffee has become one of the most sensitive commercial crops being affected by climate
22 change. Arabica coffee (*Coffea arabica*) grows in traditionally shaded agroforestry systems in
23 tropical regions and accounts for ~70% of the coffee production worldwide. Nevertheless, the
24 interaction between plant and soil water sources in these coffee plantations remains poorly
25 understood. To investigate the functional response of dominant shade trees species and coffee (*C.*
26 *arabica* var. *typica*) plants to different soil water availability conditions, we conducted a study during
27 a near normal and a more pronounced dry season (2014 and 2017, respectively) and a wet season
28 (2017) in a traditional coffee plantation in central Veracruz, Mexico. For the different periods, we
29 specifically investigated the variations in water sources and root water uptake via MixSIAR mixing
30 models that use $\delta^{18}\text{O}$ and $\delta^2\text{H}$ stable isotope composition of rainfall, plant xylem and soil water. To
31 further increase our mechanistic understanding about root activity, the distribution of belowground
32 biomass and soil macronutrients were also examined and considered in the model as prior
33 information. Results showed that, over the course of the two investigated dry seasons ~~investigated~~,
34 all shade tree species (*Lonchocarpus guatemalensis*, *Inga vera* and *Trema micrantha*) relied on

35 average, on water sources from intermediate (>15 to 30 cm depth: $58 \pm 18\%$ (SD)) and deep soil
36 layers (> 30 to 120 cm depth: $34 \pm 21\%$), while coffee plants used much shallower water sources (<
37 5 cm depth: $42 \pm 37\%$ and 5-15 cm depth: $52 \pm 35\%$). In addition, in these same periods, coffee water
38 uptake was influenced by antecedent precipitation, whereas trees showed little sensitiveness to
39 antecedent wetness. Our findings also showed that during the wet season coffee plants substantially
40 increased the use of near surface water (+56% from < 5 cm depth), while shade trees extended the
41 water acquisition to much shallower soil layers (+19% from < 15 cm depth) in comparison to drier
42 periods. Despite the plasticity in root water uptake observed between canopy trees and coffee plants,
43 a complementary use of soil water prevailed during the dry and wet seasons investigated. However,
44 more variability in plant water sources was observed among species in the rainy season when higher
45 soil moisture conditions were present and water stress was largely absent.

46

47 *Key words: Coffea arabica; water stable isotopes, roots, nutrients, clay-rich soils, MixSIAR,*

48 *Mexico*

49 **1. Introduction**

50 Coffee agroforestry systems are highly valued because of their ecological, environmental,
51 economic and social benefits (Mas and Dietsch, 2004; Perfecto et al., 2007; Tschardt et al., 2011).
52 Moreover, shade coffee of the species Arabica (*Coffea arabica*) accounts for ~ 70% of the total coffee
53 production (USDA, 2017). Although Arabica coffee is mainly grown in tropical montane regions, it
54 is cultivated under a wide range of climatic and soil conditions (Jha et al., 2014). Coffee Arabica
55 plantations can be broadly classified as traditional or modern coffee systems, according to vegetation
56 composition and structure and management practices (Moguel and Toledo, 1999). In the traditional
57 systems, coffee plants are cultivated under a diverse canopy of native and/or introduced shade tree
58 species. In contrast, monoculture coffee plantations exemplify the modern cultivation scheme, in
59 which the shade is provided by a single commercial tree species. The use of agrochemicals is also
60 typically required in this type of plantation (Moguel and Toledo, 1999).

61 Until recently, the vast majority of Arabica coffee was cultivated in traditionally managed
62 shaded coffee plantations, which have lower production costs and enhanced biodiversity, carbon
63 sequestration, soil fertility and biological pest control in comparison to modern systems (Greenberg
64 et al., 1997; Perfecto et al., 2002; Kellermann et al., 2008). However, coffee management practices
65 have become more intensive promoting the replacement of native trees with fast-growing
66 monospecific timber species (i.e. *Cedrela odorata*, *Eucalyptus deplupta*, *Hevea brasiliensis*) (Nath et
67 al., 2011).

68 Growing a crop in association with shade trees inevitably leads to some degree of competition
69 for the above-ground (light) and below-ground (water and nutrients) resources (Monteith et al., 1991).
70 In an agroforestry system, the outcome of competition for light is relatively predictable due to the
71 hierarchical structure of the canopy (i.e., shade trees intercept part of the sunlight, thereby reducing
72 the amount available for the understory crop). Conversely, competitive interactions for below-ground
73 resources can be much more diverse and complex. The central hypothesis of agroforestry underscores
74 that crops and trees are complementary in their use of soil water (Cannell et al., 1996), however the
75 degree to which this occurs will be largely controlled by the spatial and temporal patterns of resource
76 availability, root distribution and root activity, which in turn depend on factors such as climate, soil
77 conditions, crop and tree species, and plantation age, density and management practices (Beer et al.,
78 1998; Lehmann, 2003; van Noordwijk et al., 2015). In addition, below-ground competitive
79 interactions for water and/or nutrients are much more difficult to elucidate than above-ground
80 relationships. So far, the most common approach is to measure the distribution of root abundance of
81 crops and trees, and examine to what extent they overlap or are separated (e.g., Schaller et al., 2003;
82 van Kanten et al., 2005). An important limitation of this method is, however, that the spatial

83 distribution of roots does not always mirror the actual resource capture along the soil profile (Dawson
84 et al., 2002; Lehmann, 2003). Another approach is to examine the vertical patterns of soil water
85 (Cannavo et al., 2011; Padovan et al., 2015) or nutrient (Schroth et al., 2000, cited in Lehmann, 2003)
86 depletion. However, these methods are problematic because they cannot provide information on
87 whether resource depletion is caused by the crop, the trees, or both (Cannavo et al., 2011; Padovan et
88 al., 2015). Recently, the use of hydrogen ($\delta^2\text{H}$) and oxygen ($\delta^{18}\text{O}$) water stable isotope techniques in
89 combination with mixing models based on Bayesian theory has proved to be a powerful tool for
90 quantifying the proportions and probability distributions of different water sources to plant uptake
91 across different ecosystems and regions (Barbeta et al., 2015; Beyer et al., 2018; Penna et al., 2018),
92 with the potential to largely overcome the above-mentioned limitations (Dawson et al., 2002;
93 Lehmann, 2003; van Noordwijk et al., 2015). Although rarely implemented, including nutrient and
94 root distribution data along the soil profile to inform these models could provide more comprehensive
95 insights into depth of plant water uptake (cf. Muñoz-Villers et al., 2018).

96 To date, research into plant-soil interactions and plant water source partitioning in coffee
97 agroforestry systems is extremely scarce. To our knowledge, only five studies have investigated the
98 water sources of shade trees and coffee shrubs using either information on the isotopic composition
99 of plant xylem and bulk soil water (Wu et al., 2016), soil water depletion (Cannavo et al., 2011;
100 Padovan et al., 2015) or root distribution (Schaller et al., 2003; van Kanten et al., 2005). Moreover,
101 all these studies have been carried out in intensive monospecific plantations characterized by high
102 coffee planting densities ($\sim 4000\text{--}5000$ shrubs ha^{-1}), low density ($\sim 150\text{--}280$ trees ha^{-1}) and very
103 low diversity (1-2 species) of shade trees. While recognizing the limitations of some of the methods
104 used in these previous studies, the available information suggests that competition for water between
105 coffee and trees can be strong at sites with a pronounced seasonal dry period (Padovan et al., 2015;
106 Wu et al., 2016), while it seems to be virtually absent at sites with no or a relatively short dry season
107 (Schaller et al., 2003; Cannavo et al., 2011). Further, although most coffee roots are usually located
108 in the upper soil layers (< 30 cm depth; van Kanten et al., 2005, and references therein), the plant and
109 soil interactions for water during the dry season seem to occur below the main crop rooting zone ($>$
110 30 cm depth) (Wu et al., 2016). The latter reflects the ability of coffee to develop an extensive root
111 system, and to increase the root water uptake at greater soil depths once the available water has been
112 depleted in shallower layers (Huxley et al., 1974, cited in Lehmann, 2003).

113 Currently, we lack of information on plant water sources in traditional shade coffee
114 plantations. In these agroforestry systems, the higher density and diversity of shade trees could
115 potentially lead to stronger and more diverse tree-crop interactions (van Noordwijk et al., 2015). On
116 the other hand, the dense tree canopy reduces light availability and hence limits coffee water use. This

117 could lead to a lower soil water demand and thus increased plant water availability during the dry
118 season.

119 Further, ecohydrological research in these shade coffee systems is becoming increasingly
120 important since trees have been promoted as a strategy for mitigating and adapting to future climate
121 (Schroth et al., 2009; Vaast et al., 2016; Rice, 2018). Shaded coffee plantations store more carbon
122 than sun-grown coffee systems, thereby contributing to the reduction of greenhouse gases (Vaast et
123 al., 2016; Rice, 2018, and references therein). In addition, the tree canopy provides some level of
124 protection against the rising mean and maximum air temperatures (Baker and Hagggar, 2007; Schroth
125 et al., 2009; Vaast et al., 2016), which in recent modeling studies have been pointed out as the key
126 climatic changes affecting coffee growth, yield and quality (Schroth et al., 2009; Baca et al., 2014;
127 Bunn et al., 2015). Although there are important differences across sites, rainfall is also predicted to
128 decrease and become more variable in many of the world's coffee-growing regions. For example,
129 Giorgi (2006) estimated that rainfall will decrease by about 17% (per 100 years) during the dry season
130 and by about 9% during the wet season in Mexico and Central America. Similarly, predictions by
131 Karmalkar et al. (2011) for the same regions pointed out changes in rainfall of -24% to $+8\%$ (per 100
132 years) during the dry season and of -39% to -1% during the wet season. As such, if warming is
133 accompanied by decreases in rainfall, this could lead to, or exacerbate, competition for water sources
134 between coffee shrubs and shade trees (Baker and Hagggar, 2007), which in turn could affect the long-
135 term sustainability of these agroecosystems.

136 Mexico is among the largest shade coffee producers in the world, and the central region of
137 Veracruz constitutes the second most important coffee zone in the country. In this area, we selected
138 a representative traditional shade coffee plantation to investigate plant water sources of dominant
139 shade trees species and coffee (*C. arabica* var. *typica*) shrubs under different conditions of soil water
140 availability. During a near normal and a more pronounced dry season (2014 and 2017, respectively)
141 and a wet season (2017), variations in depth of plant water uptake were examined using the stable
142 isotopic composition ($\delta^{18}\text{O}$ and $\delta^2\text{H}$) of rainfall, plant xylem and soil water in combination with a
143 Bayesian mixing model (MixSIAR), along with microclimatic and soil moisture measurements. To
144 further increase our understanding about root activity and water uptake, the distribution of roots and
145 macronutrients along the soil profile were also examined and considered in the mixing model as prior
146 information. Specifically, we addressed the following questions:

147

- 148 1. Does a complementary water use strategy between shade trees and coffee shrubs prevail
149 over competition in a traditional shaded agroforestry system?

- 150 2. Does competition exist for water sources among tree and coffee species during more
151 pronounced dry periods?
152 3. What are the seasonal patterns in plant-water source partitioning?
153

154 **2. Materials and methods**

155 *2.1 Study site*

156 The research was carried out in the “La Orduña” coffee plantation (~100 ha) located on a flat
157 plateau at an elevation of 1210 m a.s.l. on the eastern slopes of the Cofre de Perote mountain (19°28'
158 N, 96°56' W) in central Veracruz State, Mexico (Fig. 1). The coffee plantations in this region occur
159 between elevations of 1000 and 1350 m a.s.l. (Marchal and Palma, 1985; Hernández-Martínez et al.,
160 2013).

161 The climate is classified as temperate humid with abundant rains during the summer (García,
162 1988). Two distinct seasons can be distinguished: (1) a wet season (May–October), during which
163 rainfall is associated primarily with cumulus and cumulonimbus clouds formed during convective
164 and orographic uplift of the moist maritime air masses brought in by the easterly trade winds; and (2)
165 a (relatively) dry season (November–April), during which most rainfall falls from stratus clouds
166 associated with the passage of cold fronts (Báez et al., 1997). Mean annual rainfall measured nearby
167 the study site during the period 1971–2000 was 1765 mm, with on average 389 mm falling during
168 the dry season and 1376 mm falling during the wet season (SMN, 2018). Mean annual temperature
169 over this period was 19.5°C, with a minimum and maximum monthly average value of 15.5 and
170 22.5°C observed in January and May, respectively (SMN, 2018). Annual potential evapotranspiration
171 (ET₀) is about 1120 mm (Holwerda et al., 2013).

172 The investigated shade coffee plantation is a so-called traditional commercial polyculture
173 system (*sensu* Moguel and Toledo, 1999), which was established more than 80 years ago. The tree
174 canopy was diverse and consisted predominantly of the species *Inga spp.*, *Citrus spp.*, *Lonchocarpus*
175 *guatemalensis*, *Trema micrantha* and *Enterolobium cyclocarpum* (Holwerda et al., 2016). The shade
176 trees were planted at a density of ca. 500 ha⁻¹, and currently form a canopy of about 14 m high. The
177 Arabica coffee plants were of the variety *typica*. *Typica* –a tall cultivar of *Coffea arabica*– was the
178 first coffee variety that arrived from Ethiopia to Mexico (Renard, 2010); it has bronze-tipped young
179 leaves and the berries are large. Plants of *typica* variety are tolerant to conditions of low soil fertility
180 and drought, but vulnerable to most pests and diseases (Escamilla et al., 2005). In the study site, this
181 cultivar was planted approximately 20 years ago at a density of about 1700 shrubs ha⁻¹, currently
182 having an average height of ~ 2 m. In this region, the coffee flowering occurs in March or April, fruit
183 development between May and October, and ripening and harvest between October and February

184 (Villers et al., 2009). The management of the plantation involves weed control practices and selective
185 pruning of mature coffee plants and shade trees at irregular times once every ~ 7 years (cf. Hernández-
186 Martínez et al., 2009). No pruning activities occurred during or in between our study periods. A
187 photograph of the coffee plantation is provided in the Supplementary Material.

188 The soil type is an Andic Acrisol derived from volcanic ashes. Soil profiles (~150 cm) are
189 multilayered (A, B1/BT and BC) and have clay (~ 65%) as the dominant texture across all layers. A
190 general description of the soil profile showed a dark brown to dark yellowish brown, clay silty organic
191 A horizon (0–20 cm) overlying a dark yellowish brown, clay silty sand B1/BT horizon (20–135 cm),
192 followed by a dark yellowish brown, clay sandy BC horizon (>135 cm). Average soil bulk densities
193 and porosities were 1.2 gr cm⁻³ and 63%, respectively, along the A and B horizons (Holwerda et al.,
194 2013). The underlying material consists of deeply weathered old lava and sandy-gravelly pyroclastic
195 flow deposits (Rodríguez et al., 2010). Soils were mostly covered by a thin (1-2 cm) but continuous
196 layer of litter.

197

198 **2.2 Hydrometeorological measurements**

199 During the study period, rainfall and microclimate conditions were continuously monitored
200 above the canopy in an 18 m high tower, located in the southwestern part of the coffee plantation.
201 Rainfall (P , mm) was measured using a TR-525 M tipping bucket rain gauge (Texas Electronics,
202 USA). Temperature (T , °C) and relative humidity (RH, %) were measured using a HC2-S3 probe
203 (Rotronic, USA). Data were recorded every 30 s, accumulated (P) or averaged values (all other
204 parameters) were stored at 5-min intervals using a CR1000 datalogger (Campbell Scientific Ltd.,
205 USA).

206

207 **2.3 Isotope sampling**

208 To examine the water sources of overstory shade trees and understory coffee shrubs, plant
209 tissue and soil samples were collected for isotope analysis at the middle (Jan. 23) and end (Apr. 11
210 and 26) of the 2014 dry season. In 2017, the dry season was warmer and drier offering the opportunity
211 to examine the vegetation responses to more pronounced dry conditions. Therefore, a second
212 sampling campaign was carried out to collect plant and bulk soil samples at the middle (Feb. 27), end
213 (Apr. 5) and late end (May. 20) of the 2017 dry season. Another sampling was carried out in the
214 middle of the 2017 wet season (Aug. 4) to evaluate plant-soil water uptake patterns at higher soil
215 water availability conditions.

216 In all seven samplings, xylem samples were obtained from three individuals of each of the
217 three dominant shade tree species (*Lonchocarpus guatemalensis*, *Inga vera* and *Trema micrantha*) by

218 extracting ~5-6 cm cores using a Pressler increment borer inserted at 1.2 m above ground ($n = 60$
219 samples of trees in total). On each occasion, xylem samples were taken from the same individuals but
220 from various aspects of the trunk. The bark was immediately removed after core extraction to avoid
221 contamination of phloem water. For the coffee plants, samples were obtained from ~6 cm segments
222 of mature suberized branches that were cut near the main stem of several shrubs each time. The bark
223 (~1mm thick) and cambium were not stripped from the coffee branches, to avoid exposure of the
224 samples to evaporation. All coffee plants were sampled randomly ($n = 40$ samples of coffee shrubs
225 in total). During the 2014 and 2017 dry seasons, sampling of coffee shrubs involved 5-6 individuals
226 each time. Since only one sampling occasion was performed during the 2017 wet season, a larger
227 number of individuals (10) was sampled to reduce the uncertainties associated with different sampling
228 sizes between wet and dry seasons respectively. For each tree, we measured diameter at breast height
229 (DBH) and height, and for the coffee plants the diameter of the main stem was measured below its
230 bifurcation in small branches (Table 1).

231 Bulk soil samples were collected at three locations and at depths of 5, 15, 30, 60, 90 and 120
232 cm, using a hand auger ($n = 126$ samples of soil in total). Auger sampling points were located so that
233 each of the sampled shade trees and coffee plants had one soil sampling point within a 3 m radius.

234 Samples of xylem and bulk soil were collected during the morning and early afternoon
235 (between 8:30 to 13:30 hrs), and each sampling campaign was preceded by at least 6 days up to 22
236 days without or with minimum accumulated rainfall (< 5 mm). All xylem and soil samples were
237 collected quickly and carefully and stored in water-tight vials to avoid any evaporation (see section
238 below).

239 To establish the local meteoric water line and compare soil water sources with recent rainfall,
240 bulk samples of rainfall ($n = 80$ in total) were collected weekly at a nearby (~ 5 km) meteorological
241 station over the course of the two years studied (Nov. 2013 – Oct. 2014 and Nov. 2016 – Oct. 2017)
242 as part of a long-term isotope sampling of precipitation (cf. Muñoz-Villers et al., 2018).

243

244 ***2.4 Isotope collection and analysis***

245 Samples of precipitation, plant xylem and bulk soil for isotope analysis were collected in 30-
246 ml borosilicate glass vials sealed with polycone caps to prevent evaporation. All samples were
247 refrigerated until extraction and analysis at the Center of Stable Isotope Biogeochemistry (CSIB) at
248 the University of California-Berkeley, USA.

249 Xylem and soil samples were extracted using cryogenic vacuum distillation (temperature:
250 $100 \pm 1.1^\circ\text{C}$, vacuum: 3 ± 1.5 Pa and time: 60-70 min) following the method of West et al. (2006).
251 The $\delta^2\text{H}$ and $\delta^{18}\text{O}$ isotopic compositions of extracted water samples were ~~then~~ determined using an

252 isotope-ratio mass spectrometer (Thermo Delta Plus XL, Thermo Fisher Scientific, USA). The
253 analytical precision of the instrument was $\pm 0.60\text{‰}$ (1 SD) for $\delta^2\text{H}$ and $\pm 0.12\text{‰}$ (1 SD) for $\delta^{18}\text{O}$.
254 Samples of precipitation were analyzed for $\delta^2\text{H}$ and $\delta^{18}\text{O}$ using a laser water isotope analyzer (L2140-
255 i) from Picarro Inc. (Santa Clara, CA, USA) in high precision and without Micro-Combustion Module
256 mode. The analytical precision was $\pm 0.65\text{‰}$ (1 SD) and $\pm 0.20\text{‰}$ (1 SD) for $\delta^2\text{H}$ and $\delta^{18}\text{O}$,
257 respectively.

258 The isotope values are expressed in delta notation (‰) relative to Vienna Standard Mean
259 Ocean Water (VSMOW). To evaluate evaporative enrichment in the soil and xylem water isotopes
260 relative to rainfall, we calculated the deuterium-excess parameter ($d = \delta^2\text{H} - 8 * \delta^{18}\text{O}$; Dansgaard,
261 1964).

262

263 ***2.5 Soil sampling and laboratory determinations***

264 To determine volumetric soil water content (SWC), samples were collected at 5, 15, 30, 60,
265 90 and 120 cm depth from each of the three boreholes excavated during the soil isotope samplings.
266 Soil moisture content was determined gravimetrically and converted to volumetric values by using
267 bulk density of the soil sample. In addition, to determine the antecedent moisture conditions for the
268 15 days prior to each sampling date, an antecedent precipitation index (API) was calculated following
269 Viessman et al. (1989).

270 To examine pH and N, P and K macronutrient concentrations along the soil profile, soil
271 samples were collected at 5, 15, 30, 60, 90 and 120 cm depth from each borehole ($n = 3$ samples per
272 soil depth) during three isotope sampling campaigns: Apr. 11, 2014 (dry season), Feb. 27, 2017 (dry
273 season) and Aug. 4, 2017 (wet season). Samples ($n = 18$) for determining other chemical properties
274 were collected at the same depths in soil profiles. All samples were first air-dried and then sieved
275 using 2 mm screens. Soil pH was determined using a glass electrode pH meter in a 1:2 soil: water
276 ratio. Organic matter (OM) was determined by the Walkley-Black method. Total carbon (C) and total
277 nitrogen (N) were measured using a TruSpec dry combustion CN analyzer (LECO, USA). Extractable
278 phosphorus (P) was determined by the Bray I method (Bray and Kurtz, 1945). Exchangeable cations
279 (Ca^+ , Mg^+ , K^+ , Na^+) were determined by extracting soil with 1 M NH_4OAc (pH 7.0). Ca^+ and Mg^+
280 were analyzed using atomic absorption spectrometry and K^+ and Na^+ were analyzed using flame
281 photometry. Soil cation exchange capacity (CEC) was determined by the ammonium acetate 1N (pH
282 7.0) method (Van Reeuwijk, 2002) and base saturation (BS) was calculated as the portion of CEC
283 that is occupied by exchangeable bases: $(\text{Ca}^+, \text{Mg}^+, \text{K}^+, \text{Na}^+)/\text{CEC}$.

284

285 ***2.6 Root biomass***

286 To examine the root biomass distribution along the soil profile in the study plot, 33 soil cores
287 were collected using 5 cm diameter and 10 cm long samplers. Soil cores were extracted at 5, 20, 40,
288 60 and 90 cm depth (from 5 to 40 cm: $n = 9$ for each depth, and from 60 to 90 cm: $n = 3$ for each
289 depth). All cores were processed immediately in the laboratory. Soil samples were first sieved using
290 2 mm screens to separate the bigger roots. Next, the samples were washed using a fine nylon mesh
291 sieve, and then separated into diameter classes (< 1 mm, 1–2 mm and > 2 mm) and dried at 70°C for
292 48 hours. Root biomass (g m^{-3}) was calculated from the dry weight of the roots and the volume of
293 the core sampler for each class and soil depth. No differentiation between roots of coffee shrubs and
294 shade trees was made.

295

296 ***2.7 Plant water uptake sources and temporal patterns***

297 The MixSIAR Bayesian mixing model framework (Moore and Semmens, 2008; Stock et al.,
298 2018) was used to determine the most likely contributions of water sources for the shade tree species
299 and coffee shrubs sampled over the course of the 2014 (Jan. 23, Apr. 11 and 26) and 2017 (Feb. 27,
300 Apr. 5, May. 20) dry seasons and the 2017 wet season (Aug. 4). To assess temporal changes of the
301 different plant water sources, the seven sampling occasions were modeled separately. The mixture
302 data for the model was the mean xylem water isotopic ($\delta^2\text{H}$ and $\delta^{18}\text{O}$) composition of the shade tree
303 species and coffee shrubs, changing accordingly with the sampling date. Based on statistical tests, the
304 relative contributions of four potential plant water sources were evaluated and restricted to the
305 following soil groups: near surface water (< 5 cm), shallow (5 to 15 cm), intermediate (> 15 to 30 cm)
306 and deep soil water (> 30 to 120 cm). For each sampling date, the mean and standard deviation of the
307 soil water isotope ($\delta^2\text{H}$ and $\delta^{18}\text{O}$) signatures from the four different grouped soil depths were
308 introduced into the model, all corresponding to the date of xylem tissue collection.

309 Further, we also considered the use of additional data such as soil macronutrients (N, P, K)
310 and root biomass information to constrain model estimates by specifying an ‘informative’ prior
311 distribution of the soil source proportions (Stock et al., 2018). These data were also grouped into four
312 classes based on the depth of the soil samplings and corresponding largely with the grouping for soil
313 water: near surface (< 5 cm) shallow (5 to ~~1520~~ cm), intermediate (> 1520 to ~~3040~~ cm) and deep ($>$
314 ~~3040~~ to 120 cm). In addition, the nearest corresponding dry or wet season dataset of soil
315 macronutrients were used according to the date of sampling. More details on the informative prior
316 parametrization are provided in the Supplementary Materials. The effect of using these priors (i.e. a
317 weight proportion before considering the isotope data) on the water sources distribution was then
318 examined by comparing these with the results of ‘non-informative’ (i.e. all the combinations of
319 proportions of water sources were equally likely) simulations. The results of each of these model runs

320 were accepted based on the examination of Markov Chain Monte Carlo convergence using the
321 Gelman-Rubin and Geweke diagnostic tests (Gelman et al., 2014).

322 Furthermore, the effect of isotope fractionation on the quantification of plant water sources
323 was specifically explored by comparing the results of the informed two-isotope mixing model with
324 those from a mixing model using only one water stable isotope ratio in the MixSIAR Bayesian
325 framework. This approach has been used elsewhere (e.g. Evaristo et al., 2017; Barbeta et al., 2019)
326 to provide some initial insights. Nevertheless, we are aware that the use of a single isotope ratio
327 approach in a multiple water source model could lead to erroneous results due to the overlap of
328 feasible solutions with poor constrained of uncertainties (see Parnell et al., 2010).

329 Lastly, the relative contributions of the water sources were compared among shade trees and
330 coffee shrubs across all sampling dates using factorial ANOVA and Tukey's HSD post-hoc tests. The
331 analyses were carried out in R Statistical Software version 3.2.4 (R Core Development Team, 2016).

332

333 **3. Results**

334 **3.1 Hydrometeorological conditions**

335 Precipitation (P) was 1650 mm in the first study year (Nov. 2013 – Oct. 2014) and 1423 mm
336 in the second study year (Nov. 2016 – Oct. 2017). During the 2013-2014 dry season (Nov – Apr.),
337 rainfall was 323 mm, and mean daily values of temperature (T) and vapor pressure deficit (VPD)
338 were $17.6 \pm 3.0^\circ\text{C}$ and 0.65 ± 0.39 kPa, respectively. The lowest monthly P and the highest T and
339 VPD were observed in April at the end of the dry season (Fig. 2a,b). During the 2016-2017 dry season,
340 rainfall amounted to 235 mm, with lowest monthly values registered in January and February at the
341 middle of the season (Fig. 2b). Mean daily T was $18.3 \pm 2.6^\circ\text{C}$, with the highest values observed at
342 the end of the dry period. Generally, VPD was high during the entire dry season (0.78 ± 0.46 kPa on
343 average), and reached maximum values in February and May.

344 Compared to long-term (1971–2000) climatic records of the region, rainfall in the first study
345 year was very close to the mean annual precipitation of 1765 mm (SMN, 2018). In contrast, the second
346 year was drier (~ 300 mm less; -20%), especially during the dry season, which had about 40% lower
347 precipitation than the average value of 389 mm. Also, higher mean monthly temperatures ($+ 0.54^\circ\text{C}$)
348 prevailed across the 2017 dry season in comparison with the 1971–2000 period. Although rainfall
349 during the 2013-2014 dry season was also about 20% lower than normal, this season was considered
350 as near average.

351 Rainfall during the 2017 wet season (May – Oct.) was lower in comparison to 2014 (1188
352 mm vs. 1326 mm, respectively) (Fig. 2b). Further, the mean air temperature and vapor pressure deficit

353 were slightly higher in the 2017 wet season than in the 2014 wet season ($20.7 \pm 1.6^\circ\text{C}$ and $0.67 \pm$
354 0.25 kPa vs. $20.1 \pm 1.5^\circ\text{C}$ and 0.60 ± 0.21 kPa, respectively) (Fig. 2a).

355

356 ***3.2 Soil moisture and antecedent precipitation during sampling campaigns***

357 During the 2014 dry season campaign (Jan. – Apr.), mean soil water content (SWC) was on
358 average $33.8 \pm 1.7\%$ at 5 cm depth, $40.2 \pm 14.5\%$ at 15 cm depth, $38.9 \pm 6.4\%$ at 30 cm depth and
359 $48.3 \pm 1.4\%$ at 60 to 120 cm depth (Fig. 2b). In comparison, SWC in the 2017 dry season campaign
360 (Feb. – May.) was lower in the first 30 cm ($32.5 \pm 3.9\%$), meanwhile water content in the deeper
361 layers was similar ($49.0 \pm 2.9\%$) with respect to the 2014 dry period. In 2014, lowest SWC values
362 were observed at the end of the dry season (April), whereas the greatest soil moisture depletion in
363 2017 was registered at the middle of the dry season (February) (Fig. 2b).

364 During the wet season sampling in August 2017, SWC values at 5 cm ($28.2 \pm 2.6\%$), 15 cm
365 ($30.9 \pm 4.3\%$), 30 cm ($38.4 \pm 4.8\%$) and 60 to 120 cm ($49.0 \pm 2.9\%$) depths were generally higher in
366 comparison to the 2017 dry period (Fig. 2b). Although the 2017 wet season sampling showed slightly
367 lower SWC values in the shallower soil layers in comparison to the 2014 dry season, the SWC values
368 in the deeper layers were higher. For the different samplings, antecedent precipitation conditions
369 (API) were, respectively, 4, 30 and 13 mm for Jan. 23, Apr. 11 and 26, 2014 and 1, 12, 9 and 43 mm
370 for Feb. 27, Apr. 5, May. 20 and Aug. 4, 2017.

371

372 ***3.3 Stable isotope composition of waters***

373 Over the study periods, a greater range of variation was found in the rainfall isotope
374 composition of the 2013-2014-year (from -126.7 to 14.4% for $\delta^2\text{H}$; from -17.7 to 0.0% for $\delta^{18}\text{O}$) in
375 comparison to the 2016-2017-year (from -113.3 to 15.5% for $\delta^2\text{H}$; from -15.9 to 0.0% for $\delta^{18}\text{O}$) (p
376 > 0.05) (Fig. 3). Overall, mean dry season rainfall was significantly more enriched than the mean wet
377 season rainfall in $\delta^2\text{H}$ and $\delta^{18}\text{O}$ ($p \leq 0.001$) (Table 2 and 3). On average, the isotopic compositions of
378 the dry and wet season rainfall were both more depleted during the second study year than during the
379 first study year; thus, the local meteoric water line of 2016-2017 had a slightly steeper slope in
380 comparison to the one for 2013-2014 (Fig. 3). Nevertheless, the range of variation of deuterium excess
381 values was similar between years (9 – 29% for the first year vs. 9 – 31% for the second year; Fig. 3),
382 and deuterium excess values of rainfall within the dry and wet seasons were not statistically different
383 ($p \geq 0.05$).

384 For all sampling dates, hydrogen and oxygen isotope composition of bulk soil water showed
385 a consistent pattern of increasing isotope depletion with soil depth (Supplementary Materials), in
386 which shallower (5-15 cm) soil water was significantly more enriched than intermediate (15-30 cm)

387 and deeper (30-120 cm) soil water layers ($p \leq 0.001$) (Table 2 and 3; Fig. 3). In correspondence,
388 lowest values of deuterium excess generally characterized the near surface soil water pool.

389 For the 2014 dry season samplings, bulk soil ranged from -83.3 to -11.9‰ for $\delta^2\text{H}$ and from
390 -11.1 to -0.9‰ for $\delta^{18}\text{O}$ (Fig. 3a). For the 2017 dry season samplings, bulk soil water showed a
391 narrower range of variation and more enriched isotope values (from -54.8 to -19.1‰ for $\delta^2\text{H}$ and
392 from -7.5 to -1.5‰ for $\delta^{18}\text{O}$) in comparison to 2014 (Fig. 3b). However, statistical differences were
393 only suggested for the intermediate and deeper soil layers in both water isotopes between the two dry
394 seasons investigated ($p \leq 0.001$).

395 In the 2017 wet season sampling, bulk soil isotope composition ranged from -70.5 to -37.5‰
396 for $\delta^2\text{H}$ and from -8.4 to -4.1‰ for $\delta^{18}\text{O}$ (Fig. 3c), showing significant differences in the shallow,
397 intermediate and deep soil water pools in comparison to the 2017 dry season ($p \leq 0.001$). In all
398 sampling periods, bulk soil water across the different depth groups was isotopically distinct from
399 rainfall during the 2014 and the 2017 dry seasons ($p \leq 0.001$ for both water isotopes).

400 Across all sampling periods, xylem water of coffee shrubs was more enriched than that of
401 shade trees ($p \leq 0.001$) (Table 2 and 3; Figure 3). In the 2014 dry season, xylem water isotope values
402 of shade trees ranged from -65.5 to -32.1‰ for $\delta^2\text{H}$ and from -7.6 to -3.6‰ for $\delta^{18}\text{O}$, meanwhile a
403 larger variation was observed in the xylem water of coffee shrubs (from -46.5 to -9.6‰ for $\delta^2\text{H}$ and
404 from -6.3 to -0.6‰ for $\delta^{18}\text{O}$) ($p \leq 0.001$) (Fig. 3a). Among tree species, *Lonchocarpus guatemalensis*
405 showed the most depleted xylem water isotope signature ($-58.1 \pm 4.8\text{‰}$ for $\delta^2\text{H}$ and $-6.8 \pm 0.5\text{‰}$ for
406 $\delta^{18}\text{O}$), whereas *Inga vera* had the most enriched values with a greater range of variation ($-51.0 \pm$
407 10.2‰ for $\delta^2\text{H}$ and $-5.3 \pm 1.1\text{‰}$ for $\delta^{18}\text{O}$). Statistical tests showed that *Inga vera* was different from
408 the other tree species in $\delta^{18}\text{O}$ ($p < 0.05$).

409 In the 2017 dry season, the isotopic composition of shade trees varied from -56.7 to -34.5‰
410 for $\delta^2\text{H}$ and from -6.0 to -3.2‰ for $\delta^{18}\text{O}$; corresponding values for coffee shrubs varied from -39.6
411 to -7.8‰ for $\delta^2\text{H}$ and from -4.4 to -1.1‰ for $\delta^{18}\text{O}$ ($p \leq 0.001$) (Fig. 3b). Contrary to 2014, *L.*
412 *guatemalensis* showed the most enriched isotope value ($-41.3 \pm 5.7\text{‰}$ for $\delta^2\text{H}$ and $-4.6 \pm 0.5\text{‰}$ for
413 $\delta^{18}\text{O}$), and *I. vera* had the most depleted values ($-48.5 \pm 5.1\text{‰}$ for $\delta^2\text{H}$ and $-4.8 \pm 0.8\text{‰}$ for $\delta^{18}\text{O}$),
414 with differences being statistically significant for $\delta^2\text{H}$ ($p < 0.05$).

415 Overall, isotope values of plant xylem water were more enriched during the 2017 dry season
416 than during the 2014 dry season ($p \leq 0.001$) (Fig. 3a,b; Fig. 4). Deuterium excess values were also
417 lower in shade trees and coffee shrubs during 2017, indicating a more evaporative signature (Table 2
418 and 3; Fig. 3). Plots of $\delta^2\text{H}$ xylem water against height for the individual shade trees and coffee shrubs
419 sampled in both dry seasons are shown in Figure 4, in which a similar $\delta^2\text{H}$ pattern was displayed
420 between trees and coffee shrubs in the 2014 and 2017 years.

421 During the 2017 wet season sampling, $\delta^2\text{H}$ and $\delta^{18}\text{O}$ values in xylem water of trees and coffee
422 shrubs were more depleted in comparison to the 2017 dry season ($p < 0.05$) (Fig. 3c). The range of
423 variation was from -60.6 to -45.6 ‰ in $\delta^2\text{H}$ and -6.2 to -5.4 ‰ in $\delta^{18}\text{O}$ for trees, and from -42.2 to $-$
424 34.4 ‰ in $\delta^2\text{H}$ and -5.4 to -4.4 ‰ in $\delta^{18}\text{O}$ for coffee shrubs ($p \leq 0.001$).

425 It was observed that the xylem isotopic composition of all shade trees and coffee plants fell
426 within the range of the soil water sources during the 2014 dry season samplings (Fig. 3a). For the
427 2017 dry season, we again observed a good isotopic match between the shade tree xylem water and
428 soil water. However, for the coffee plants, the xylem water was more enriched in $\delta^2\text{H}$ in comparison
429 to soil water (Fig. 3b). During the 2017 wet season sampling, a slight enrichment in $\delta^2\text{H}$ was again
430 observed in the xylem water of coffee, while trees showed a good overlap with soil water (Fig. 3c).
431 Based on these results, tests were carried out to specifically evaluate the effects of deuterium
432 fractionation on coffee water sources by running a simple mixing model using only hydrogen isotope
433 ratios in the MixSIAR framework.

434 435 **3.4 Root biomass and macronutrients along soils profile**

436 Overall, most roots were concentrated in the first 5 cm of soil with a sharp decline in biomass
437 at 20 cm depth (Fig. 5a). Fine roots ($< 1\text{mm}$) followed by bigger roots ($> 2\text{ mm}$) dominated the
438 shallower soil layers ($< 20\text{ cm}$), meanwhile roots in general were scarce at deeper depths ($> 60\text{ cm}$).
439 Soil acidity was highest near the surface and decreased gradually with depth (Table 4). Organic matter
440 (OM) and total carbon were also greatest between 5 and 15 cm depth, while values decreased rapidly
441 below ~ 30 to 60 cm depth. Although highest concentrations of nitrogen were found in the first 15 cm
442 of soil, values remained relatively high and constant at deeper layers (Fig. 5b). Phosphorus showed
443 its highest concentration at the topsoil with values decreasing sharply below 30 cm depth. In contrast,
444 concentrations of potassium, sodium and magnesium were lowest in the first 15 cm, while maximum
445 values were observed below 90 cm depth. Base saturation (BS) was very low along the soil profile,
446 indicating poor availability of soil macronutrients. Soil cation exchange capacity (CEC) was generally
447 low across depths, indicating little potential for interaction between clay particles and cations.

448 449 **3.5 Plant water sources**

450 We found a good agreement between the MixSIAR Bayesian mixing model results using a
451 non-informative and an informative prior distribution (on average 5% difference across all xylem
452 water contributing sources; $p > 0.05$). This indicates that the independent distribution (soil
453 macronutrients and root data) set a *priori* to optimize source proportion estimates (informative
454 approach) in the model was not influential enough to significantly modify the results obtained using

455 the isotope signatures of the xylem water sources alone (non-informative approach). Having this
456 agreement between models, we present the results of the water source contribution based on the
457 informative model runs. Results of the non-informative approach are provided in the Supplementary
458 Materials.

459 The model results showed that the intermediate and deep soil water pools (> 15 to 120 cm
460 soil depth) were the main sources for the shade trees over the course of the 2014 dry season ($91 \pm$
461 37% on average; Fig. 6 and Supplementary Materials). Across this period, *L. guatemalensis* showed
462 on average the highest proportion of water uptake between 30 and 120 cm soil depth ($49 \pm 26\%$),
463 while *T. micrantha* and *I. vera* depended strongly on soil water sources between 15 and 30 cm ($54 \pm$
464 18% and $67 \pm 6\%$) ($p < 0.001$). In contrast, the water uptake of coffee plants was mainly sustained by
465 sources from the first 15 cm of soil ($94 \pm 27\%$ on average; Fig. 6 and Supplementary Materials),
466 having significant differences with all shade tree species ($p < 0.001$).

467 During the 2017 dry season, the same trend with most water extracted from intermediate and
468 deep soil layers was observed in the shade trees ($91 \pm 39\%$ on average; Fig. 7a,b,c and Supplementary
469 Materials). Among sampling dates, differences between tree species only appeared to occur at the
470 end of the dry period (Apr. 5) ($p < 0.05$). Coffee water sources were again restricted to much shallower
471 soil layers (0–5cm: $53 \pm 44\%$ and 5–15 cm: $42 \pm 41\%$; Fig. 7a,b,c and Supplementary Materials)
472 compared to shade trees.

473 Overall, we did not find any statistically significant difference among main plant water
474 sources between the dry periods investigated ($p > 0.05$). Across the individual samplings throughout
475 the two dry seasons, we observed that antecedent precipitation had a stronger effect on the water
476 uptake sources of coffee plants than trees (Fig. 8). For example, when dry antecedent wetness
477 prevailed ($API_{15} < 5$ mm; Fig. 2b) coffee water sources were mainly composed of soil water from >
478 5 to 15 cm depth ($91 \pm 3\%$). Alternatively, when wetter antecedent conditions were present ($API_{15} >$
479 10 mm), the near surface soil water layer ($58 \pm 31\%$) was the main contributing source. On the
480 contrary, tree water uptake was essentially sustained by deeper soil water sources at low and relatively
481 high antecedent wetness conditions ($94 \pm 23\%$ and $87 \pm 23\%$, respectively) (Fig. 8). Nevertheless, for
482 all species investigated, the relationships between API and the contribution of near surface soil water
483 sources were not statistically significant ($p > 0.05$).

484 During the 2017 wet season, water source partitioning differed among shade tree species (Fig.
485 7d and Supplementary Materials). During this period, *L. guatemalensis* and *I. vera* showed the
486 greatest use of deep soil water ($74 \pm 37\%$ and $69 \pm 41\%$, respectively), while shallower soil water
487 was the main source for *T. micrantha* ($91 \pm 23\%$), having significant differences with the other tree
488 species ($p < 0.001$). Coffee consistently showed the use of near surface water sources ($98 \pm 5\%$; Fig.

489 7d and Supplementary Materials), which was significantly different from all shade tree species ($p <$
490 0.001).

491

492 **3.6 Fractionation effects on coffee water sources**

493 To evaluate the effects of xylem deuterium fractionation on our results for coffee water source
494 uptake, we compared the relative contribution of each soil water source obtained via the single-
495 isotope ($\delta^2\text{H}$) mixing model with those obtained via the informative two-isotope mixing model. In
496 general, we observed that the $\delta^2\text{H}$ model consistently estimated a lower contribution of the shallow
497 soil water source and a higher contribution of the near surface soil water source (Supplementary
498 Materials). On average, the reduction in the shallow soil water source ($-25.7 \pm 29.0\%$) coincided very
499 well with the increase in the near surface soil water source ($+28.1 \pm 30.6\%$). These differences were
500 most pronounced for the 2017 dry season samplings ($p > 0.05$; Supplementary Materials), during
501 which the differences in $\delta^2\text{H}$ between coffee xylem water and soil water were greatest. However,
502 there were no significant differences between the relative contributions of the intermediate and deep
503 soil water sources estimated by the two models ($p > 0.05$). In summary, the results of the $\delta^2\text{H}$ mixing
504 model suggested an even more pronounced soil water partitioning between coffee and shade tree
505 species than those obtained with the informative two-isotope mixing model.

506

507 **4. Discussion**

508 **4.1 Methodological aspects**

509 To our knowledge, the ecohydrological study presented here is one of the first that
510 incorporates biophysical properties as prior information alongside plant water source information
511 from stable isotope ($\delta^{18}\text{O}$ and $\delta^2\text{H}$) data into a MixSIAR Bayesian mixing model framework, as a way
512 to improve our understanding of the processes that lead to differences in the depth of plant water
513 uptake. Even though our findings did not change significantly by including or excluding the prior
514 information such as soil macronutrients and root data, exploring plant water source partitioning using
515 these two model approaches provided more confidence in our results. Therefore, we call for more
516 studies that combine soil nutrient and root biomass distribution with plant water source information
517 from $\delta^{18}\text{O}$ and $\delta^2\text{H}$ data, to explore the additional value of these biophysical parameters elucidating
518 plant-soil interactions in different regions and environments.

519 In recent years, some plant, soil and/or deep subsurface water source studies that have used
520 stable isotopes have identified isotope variation that could be the result of isotope fractionation
521 processes caused by water molecules interacting with clay surfaces, partially filled pore spaces or
522 salts ([Oerter et al., 2014](#); [Oshun et al., 2015](#); [Chen et al., 2016](#); [Lin et al., 2017](#); Gaj and McDonnell,

2019; ~~Lin et al., 2017~~ ~~Oerter et al., 2014~~; ~~Oshun et al., 2015~~). Our soils were rich in clay content and according to some studies this type of soil structure can impart isotope fractionation (~~Lin et al., 2017~~; ~~Meißner et al., 2014~~; ~~Oerter et al., 2014~~; ~~Orlowski et al., 2016a~~; ~~Lin et al., 2017~~). Thus far, however, these isotope effects have been more evident in clay-rich soils having high cation exchange capacities (CEC ~ 30 to 70 cmol_c kg⁻¹; ~~Oerter et al., 2014~~; ~~Orlowski et al., 2016b~~) in combination with low soil water contents (SWC < 20% ~~Meißner et al., 2014~~; ~~Orlowski et al., 2016b~~). In this respect, the soils in our study area are characterized by low CEC (< 21 cmol_c kg⁻¹; Table 4). This reflects relatively little interaction between cations adsorbed and clay mineral particles, which indirectly suggests minimal impacts of interlayer water bound in the soil structure (cf. Vidal and Dubacq, 2009). In addition, our soil samples were collected at relatively high SWC across the different sampling periods (~ 30% to 60%; Figure 1). As such, we have assumed that the probability of fractionation due to soil properties that may impact water extraction efficiency, was very small or completely absent and therefore, the extracted soil water was the same the plants had access to.

With regard to our plant samples, we specifically observed enrichment in the deuterium composition of the xylem water in the coffee plants in comparison to bulk soil water. It is not surprising that fractionation was evident for $\delta^2\text{H}$ and not $\delta^{18}\text{O}$, given the higher fractionation factor of ^2H relative to ^{18}O (Rundel et al., 2012). Some possible explanations for this xylem water enrichment could be related to bark evaporation (Ellsworth and Sternberg, 2015) and/or xylem-phloem water exchange (Cernusak et al., 2005), since we did not remove the bark and cambium from our coffee branch samples. On the other hand, like many other crops, coffee plants associate symbiotically with arbuscular mycorrhizal fungi (López-Andrade et al., 2009; Perea-Rojas et al., 2019). Studies in our coffee growing region of Veracruz have documented the presence of mycorrhizal structures in coffee roots (~~Muleta et al., 2008~~; ~~Arias et al., 2012~~; ~~Muleta et al., 2008~~), which can promote increases in plant water and nutrient uptake (~~Scheneiger and Jakobsen, 2000~~; ~~Augé, 2004~~; ~~Scheneiger and Jakobsen, 2000~~). Although no research has been carried out yet to test the influence of mycorrhizal fungi on isotope fractionation during coffee root water uptake, this effect could have been present and being also responsible for the isotopic mismatch between the coffee xylem water and soil water sources, ~~as it has been reported elsewhere (Poca et al., 2019)~~.

We did evaluate the effects of these isotope enrichments in the coffee xylem water on the relative contributions of the coffee water sources using a single-isotope ($\delta^2\text{H}$) mixing model. Consistently, the model results estimated a higher near surface water and a lower shallow soil water source contribution in comparison to the dual isotope informative prior mixing model. In contrast, the estimated proportions of the intermediate and deep soil water sources were similar between models. Thus, the effect of fractionation was translated into a more pronounced spatial separation

557 between the main soil water sources of the coffee plants and shade trees, but our overall results were
558 not different.

559

560 **4.2 Complementary water use strategy between shade trees and coffee shrubs**

561 Our findings showed that all shade tree species (*L. guatemalensis*, *I. vera* and *T. micrantha*)
562 relied mainly on water sources from deep soil layers (> 15 to 120 cm depth), while the use of much
563 shallower water sources (< 15 cm) was observed in the coffee (*C. arabica* var. *typica*) over the course
564 of the near normal and the more pronounced dry seasons studied. These findings suggest a spatial and
565 temporal partitioning of soil water sources between shade trees and coffee plants during drier periods
566 and water-resource complementary in this coexistence species environment.

567 Although comparisons of our findings with other traditional shade Arabica coffee plantations
568 are difficult because studies are essentially lacking in this type of agroecosystems, there are a handful
569 of other investigations carried out in shade coffee monospecific plantations in the humid tropics in
570 which complementary rather than competitive water use strategies prevailed. For example, Cannavo
571 et al. (2011) compared the water use and soil water availability of an unshaded coffee vs. a shaded
572 monoculture (*Inga densiflora*) coffee plantation in Costa Rica, both of 7-8 years old, using soil
573 moisture measurements and water balance calculations. Their results showed that soil water content
574 in the deeper soil layers (> 120 cm depth) was lower in the shaded coffee than in the sun-grown coffee
575 system, while water content in the shallower layers was similar. This suggested that associated shade
576 trees preferentially used water from deeper soil horizons providing some evidence of
577 complementarity water use between coffee plants and native *Inga* trees during the dry season.
578 However, the authors acknowledged that they were unable to separate roots of coffee from those of
579 trees in the soil profiles, so they could not be certain whether trees were the only individuals extracting
580 water from deeper sources. In this respect, our study showed that there was always a mixture in water
581 uptake from different sources (soil group depths), but a separation between the main sources of water
582 for shade trees and coffee shrubs clearly prevailed.

583 Other investigations in Costa Rica have examined the belowground resource competition of
584 Arabica coffee in association with fast-growing timber species using data of plant growth, root
585 distribution and density, and soil moisture and nutrients patterns. For example, the study of Schaller
586 et al. (2003) carried out in a commercial (*Eucalyptus deplupta*) shade coffee plantation where soils
587 are highly fertilized, showed that coffee had a relatively even root distribution along the first 40 cm
588 of soil depth with a higher root density in the proximity of the coffee rows. Conversely, the root
589 system of *E. deplupta* was much shallower having most roots concentrated in the upper 10 cm of soil.
590 In this case, the tree root density was found highest in the alleys between the coffee rows. The authors

591 explained that the apparent complementary resource exploitation of this tree-crop system was mainly
592 attributed to high availability of soil resources and the high competitiveness of the coffee limiting the
593 expansion of tree roots (cf. Lehmann, 2003). Although in our study we did not determine the depth
594 distribution of coffee and tree roots, our findings showed that all shade tree species were tapping
595 water from deeper soil layers than coffee, suggesting that trees are deep rooted and being able to
596 explore larger soil volumes causing little competition with coffee.

597 In Nicaragua, Padovan et al. (2015) compared the root distribution, soil moisture,
598 transpiration and leaf water potential patterns in a sun-grown coffee system and an agroforestry of
599 coffee planted with two timber trees (deciduous *Tabebuia rosea* and evergreen *Simarouba glauca*).
600 Their findings showed that coffee roots were more abundant than tree roots and mainly concentrated
601 in the shallower soil layers (0–80 cm depth). Most roots of both tree species were observed in deeper
602 layers (>100 cm) suggesting a clear niche differentiation with coffee. During the 3-year study period,
603 volumetric water content along a 2 m soil profile was higher in the sun-grown coffee than in the
604 shaded coffee, which was explained by greater soil water uptake from trees below the crop rooting
605 zone (Padovan et al., 2015). Moreover, coffee shrubs in the shaded plantation were more water
606 stressed (i.e. lowest midday leaf water potentials) during the pronounced dry season studied (Padovan
607 et al., 2018). Their results suggest that despite the clear hydrological niche segregation, competition
608 between coffee and shade trees may occur if the dry season is long and severe enough.

609 Our findings also showed that during the wet season coffee plants substantially increased the
610 use of near surface water (+56%) in comparison to the dry season, while all shade trees also extended
611 their water acquisition to much shallower soil water pools (+19%). This is largely explained by the
612 increases in soil moisture in the first 30 cm depth due to frequent rainfall inputs that characterize the
613 wet season in our study area. This also suggests that coffee had a higher root activity in the top soil
614 layers during the wet season in comparison to the dry season, as has been documented in other studies
615 (Huxley et al., 1974). Regarding the shade trees, we observed that *T. micrantha* showed the greatest
616 response to wetter conditions by drawing most water from the first 15 cm of soil (92%), whereas this
617 was much less evident in *L. guatemalensis* (21%) and *I. vera* (27%). Although we did not determine
618 the vertical distribution of roots for each of the shade tree species studied, these findings suggest that
619 *T. micrantha* has a shallower rooting system than the other tree species. The fact that the *T. micrantha*
620 trees were more recently planted (i.e. younger with less developed root system) than the *L.*
621 *guatemalensis* and *I. vera* trees supports this idea. On the other hand, the high temperature and rainfall
622 that characterize the wet season at our study site may favor rapid mineralization of nutrients and their
623 subsequent leaching to deeper soil layers (i.e. potassium, calcium and magnesium; Table 4). Hence,
624 for the larger trees studied (*L. guatemalensis*), the availability of water and nutrients at deeper depths

625 could have been an important resource for plant growth in this period, partly explaining the lower
626 activity of their shallower roots. Despite the changes and the higher variability in depth of water
627 uptake observed among canopy trees and coffee shrubs, a complementary use of soil water prevailed
628 during the wet season. Future work should be focused on the distribution and dynamics of tree and
629 crop roots and their seasonal variation in relation to the availability of nutrients and water in the soil.
630 Also, it would be desirable to relate these dynamics to crop and shade tree phenology to elucidate
631 temporal synergistic or competitive water requirements.

632

633 ***4.3 The role of antecedent wetness in coffee water uptake***

634 Despite the relatively small sample size, our study showed that antecedent wetness strongly
635 influenced the water uptake patterns of coffee plants (cf. Huxley et al., 1974). We found that under
636 relatively wet antecedent conditions prevailing after small rainfall events during the dry season, coffee
637 substantially increased the use of near surface soil water sources, possibly as an opportunistic strategy
638 to overcome the soil water deficits in this period and taking advantage of their much shallower rooting
639 system compared to trees. Conversely, tree water uptake was mainly sourced by deeper soil water
640 layers showing less sensitiveness to higher antecedent wetness. In this respect there are no
641 comparative studies in shade coffee agroecosystems evaluating the functional response of plant water
642 uptake over a range of antecedent wetness. Nevertheless, plant and soil water interactions under dry
643 and relatively wet conditions have been examined in other types of agroforestry systems. For
644 example, in the study of Gao et al. (2018) carried out in a semiarid region in China, the authors
645 evaluated the seasonal variations in water use of jujube (*Ziziphus jujuba*) trees planted with annual
646 (*Brassica napus*) and perennial (*Hemerocallis fulva*) crops. Using stable isotope techniques and
647 Bayesian mixing modelling, their results showed that jujube trees generally tapped water (> 58%)
648 from deep soil layers (60-200 cm depth) at low antecedent wetness, while **both** *B. napus* and *H. fulva*
649 crops primarily extracted water (> 65%) from intermediate (20-60 cm) and shallow (0-20 cm) soil
650 layers. This exhibits a complementary water use strategy between trees and crops. However, at higher
651 antecedent wetness both the jujube trees and crops extracted most water from the first 0-60 cm of soil
652 depth (> 65%). This indicated that both species exhibited an opportunistic strategy for accessing
653 resources at shallower soil depths. In this case, contrary to our findings, tree roots rather than crop
654 roots showed the stronger capacity to switch rapidly from deep to shallow sources in response to
655 increased soil water availability.

656

657 ***4.4 Implications and future directions***

658 The consistent complementarity in plant water use strategies observed under different
659 hydrometeorological conditions in the coffee plantation studied provides support to the central tenet
660 of agroforestry systems (Cannel et al., 1996). Based on our findings, *L. guatemalensis*, *I. vera* and *T.*
661 *micrantha* provide good choices for coffee shade trees due to their complementarity in soil water use.
662 Since these tree species obtained their water from deeper soil layers than the coffee, this could mean
663 that they utilize nutrients leaching beyond the reach of the coffee plants, and so contribute to improved
664 nutrient cycling and increased overall productivity of the system (van Noordwijk et al., 2015).

665 Nevertheless, the current outcome may change given the new coffee management practices
666 that consist on replacing traditional coffee varieties (e.g. *C. arabica* var. *typica*) with others (*C.*
667 *arabica* var. *costa rica*; *C. canephora*) that may exhibit deeper roots systems and perhaps different
668 water (and nutrient) uptake strategies, in response to prevalent diseases such as leaf rust or root
669 nematodes. Therefore, future research should be focused on evaluating the water source partitioning
670 of traditional vs. new coffee disease-resistant varieties and their relation to shade tree water use. In
671 this respect, there are further questions with regard to strategic use of shade tree species, whereby
672 fast-growing species might be more (commercially) productive but also more competitive. Some
673 evidence from elsewhere has shown that such management practices do not necessarily increase
674 competition and may even enhance the water use efficiency as part of drought-avoidance
675 mechanisms. For example, in southeast China, Wu et al. (2016) used $\delta^2\text{H}$ and $\delta^{18}\text{O}$ stable isotope
676 methods to examine the seasonal water use of a fast-growing rubber tree species (*Hevea brasiliensis*)
677 planted with Arabica coffee. Their findings showed that rubber trees were mostly accessing water
678 from intermediate (15-50 cm depth) and deep soil layers (50-110 cm), meanwhile coffee was mostly
679 tapping water from the topsoil (< 15 cm). Additionally, rubber trees showed strong root plasticity in
680 soil water uptake avoiding competition with coffee during the rainy and relatively dry seasons.
681 However, more research is needed since these results depend largely on tree-crop specie combinations
682 and local climatic and soil conditions.

683 In addition to effects of changing management practices, climate warming may induce
684 changes in plant transpiration throughout the year (e.g. Karmalkar et al., 2011). In our study, we used
685 a water stable isotope approach along with root and soil macronutrients data to estimate the relatively
686 contribution of the plant water sources. However, for a more complete assessment of the plant and
687 soil interactions, seasonal plant water fluxes need to be quantified. Our results so far have made the
688 first steps towards serving coffee producers to make better decisions on sustainable coffee and water
689 management, as well as providing new insights into water resources in general, which are urgently
690 required for implementing efficient and equitable management programs in humid tropical
691 environments (Hamel et al., 2018). However, future work should be focused on water use of

692 individual trees and coffee shrubs using ecophysiological and hydrological techniques in order to
693 know how much water is used from the different soil water pools.

694

695 **5. Conclusions**

696 This study provides the first baseline information on plant water sources for a traditional
697 shade coffee plantation in the humid tropics. Our results showed that coffee water uptake was mainly
698 sustained from shallow soil sources (< 15 cm depth) while all shade trees relied on water sources
699 from deeper soil layers (>15 to 120 cm depth). This complementary strategy in soil water use between
700 crops and trees was consistent over the course of the near normal and the more pronounced dry
701 seasons investigated. Across these same periods, we observed that antecedent precipitation had a
702 strong influence in coffee plants increasing their water uptake to near surface soil water sources as an
703 opportunistic strategy to overcome the reduced water availability. In the wet season, coffee plants
704 substantially increased the use of near surface water (< 5 cm depth), whereas shade trees expanded
705 their water acquisition to the first 15 cm of soil depth. Overall, a greater soil water partitioning
706 prevailed among tree and coffee species when higher soil moisture conditions were present.
707 Nevertheless, despite such variability in plant-soil water interactions across seasons, a clear spatial
708 segregation of the main water source prevailed between shade trees and coffee plants during the rainy
709 and dry periods investigated.

710

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712 data. MSAB performed all the Bayesian mixing model analysis. JG contributed in the data analysis.
713 LEMV prepared the first draft of the manuscript. FH, MSAB and JG edited and commented on the
714 manuscript several times, and TED carried out the final revision. Later, all the co-authors
715 contributed with revisions.

716

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718

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736

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Table 1. Characteristics of the shade trees and coffee plants sampled for water isotope analysis during 2014 and 2017. Numbers between parentheses are the standard deviation.

Family	Species	Canopy layer	2014		2017		<i>n</i>
			DBH cm	Height m	DBH cm	Height M	
Fabaceae	<i>Lonchocarpus guatemalensis</i>	Overstory	101.5 (12.6)	20.3 (1.3)	119.8 (12.1)	21.0 (1.2)	3
Fabaceae	<i>Inga vera</i>	Overstory	39.3 (15.7)	10.7 (4.8)	48.1 (13.3)	9.6 (1.2)	3
Cannabaceae	<i>Trema micrantha</i>	Overstory	13.16 (6.8)	8.15 (3.1)	23.3 (7.2)	15.2 (2.2)	3
Rubiaceae	<i>Coffea arabica</i> var. <i>typica</i>	Understory	12.7 (2.1)	2.83 (0.7)	n.a.	n.a.	5* 6** 10***

* Number of individuals sampled each time in the 2014 dry season

** Number of individuals sampled each time in the 2017 dry season

*** Number of individuals sampled in the 2017 wet season

Table 2. Mean \pm (SD) H and O stable isotope composition of 2013-2014 precipitation, tree xylem water and bulk soil water of the 2014 dry season sampling, and corresponding *d*-excess values (‰)

Precipitation <i>n</i> = 41						Bulk soil water <i>n</i> = 54												Shade trees xylem water <i>n</i> = 27			Coffee shrubs xylem water <i>n</i> = 14		
Dry season			Wet season			0-5 cm depth			>5-15 cm depth			>15-30 cm depth			>30-120 cm depth								
δ^2 H	δ^{18} O	<i>d</i> - exce ss	δ^2 H	δ^{18} O	<i>d</i> - exce ss	δ^2 H	δ^{18} O	<i>d</i> - exce ss	δ^2 H	δ^{18} O	<i>d</i> - exce ss	δ^2 H	δ^{18} O	<i>d</i> - exce ss	δ^2 H	δ^{18} O	<i>d</i> - exce ss	δ^2 H	δ^{18} O	<i>d</i> - exce ss	δ^2 H	δ^{18} O	<i>d</i> - exce ss
1. 6 \pm 8. 5	- 1.9 \pm 1.4	17.0 \pm 5.1	- 42. 4 \pm 36. 1	- 7.2 \pm 4.3	14.9 \pm 2.8	- 20. 5 \pm 7.8	- 2.4 \pm 1.0	-1.5 \pm 4.1	- 30. 8 \pm 9.4	- 3.7 \pm 1.1	-1.2 \pm 6.3	- 54. 7 \pm 10. 3	- 7.0 \pm 0.9	1.2 \pm 6.6	- 66. 8 \pm 8.6	- 8.7 \pm 1.3	3.0 \pm 4.7	- 55. 4 \pm 7.6	- 6.2 \pm 1.0	-5.8 \pm 4.1	- 25. 5 \pm 10. 8	- 3.4 \pm 1.8	1.7 \pm 5.0

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Table 3. Mean \pm (SD) H and O stable isotope composition of 2016-2017 precipitation, tree xylem water and bulk soil water of 2017 dry season sampling, and corresponding *d*-excess values (‰)

Precipitation <i>n</i> = 39						Bulk soil water <i>n</i> = 54												Shade trees xylem water <i>n</i> = 24			Coffee shrubs xylem water <i>n</i> = 18		
Dry season			Wet season			0-5 cm depth			>5-15 cm depth			>15-30 cm depth			>30-120 cm depth								
δ^2 H	δ^{18} O	<i>d</i> - exce ss	δ^2 H	δ^{18} O	<i>d</i> - exce ss	δ^2 H	δ^{18} O	<i>d</i> - exce ss	δ^2 H	δ^{18} O	<i>d</i> - exce ss	δ^2 H	δ^{18} O	<i>d</i> - exce ss	δ^2 H	δ^{18} O	<i>d</i> - exce ss	δ^2 H	δ^{18} O	<i>d</i> - exce ss	δ^2 H	δ^{18} O	<i>d</i> - exce ss
- 2.9 \pm 16. 0	- 3.0 \pm 1.8	21.5 \pm 4.3	- 47. 8 \pm 34. 4	- 7.9 \pm 4.1	15.2 \pm 3.3	- 24. 3 \pm 3.9	- 2.2 \pm 0.5	-6.9 \pm 6.6	- 32. 1 \pm 5.3	- 3.6 \pm 0.5	-3.4 \pm 4.1	- 41. 9 \pm 5.7	- 5.7 \pm 0.6	3.4 \pm 4.8	- 47. 3 \pm 3.8	- 6.5 \pm 0.5	5.0 \pm 3.2	- 44. 9 \pm 5.6	- 4.4 \pm 0.7	-9.7 \pm 5.4	- 21. 3 \pm 7.2	- 2.8 \pm 1.0	1.3 \pm 6.2

Table 4. Soil characteristics (average values) determined at the different depths

Soil depth	pH (H ₂ O)	P	Na	K	Ca	Mg	CEC	BS	OM	C	N	Clay	Loam	Sand
(cm)		(mg kg ⁻¹)	(cmol _c kg ⁻¹)					(%)						
5	4.07	33.33	1.47	0.60	3.86	0.87	16.10	0.42	5.18	3.01	0.38	60.83	25.1	13.9
15	4.12	4.60	1.08	0.47	0.95	0.12	13.27	0.20	2.89	1.90	0.30	63.8	24.3	11.9
30	4.34	n.d.	2.22	0.77	1.92	0.54	14.65	0.37	1.55	1.31	0.23	70.9	18.6	10.5
60	4.95	n.d.	2.36	0.93	3.81	1.21	20.35	0.41	1.02	0.69	0.22	66.9	16.3	16.8
90	5.10	n.d.	2.75	1.11	3.78	1.27	18.85	0.47	0.48	0.50	0.20	66.1	14.9	19.1
120	5.16	n.d.	3.00	1.45	3.76	1.20	17.60	0.53	0.41	0.51	0.20	65.1	14.0	20.9

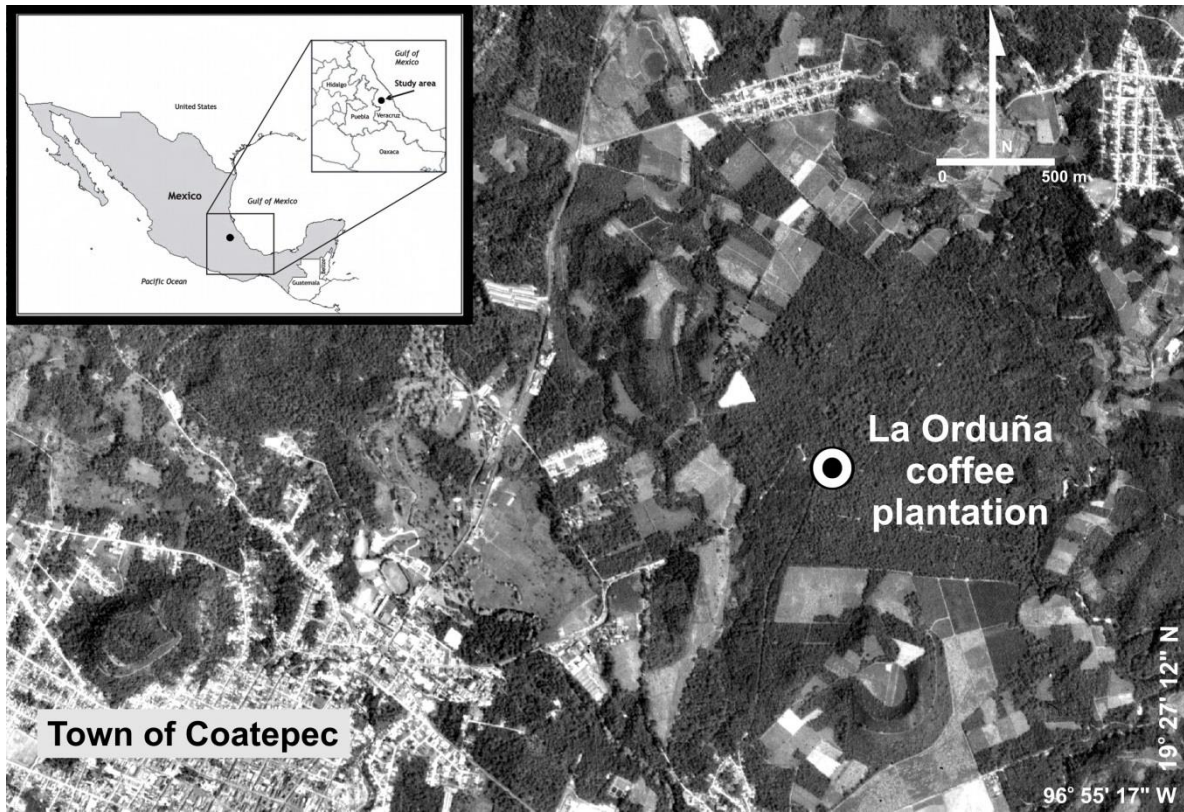


Figure 1. Study site location in the municipality of Coatepec, Veracruz, Mexico. Source: QuickBird Satellite Image (2010). Copyright DigitalGlobe, Inc.

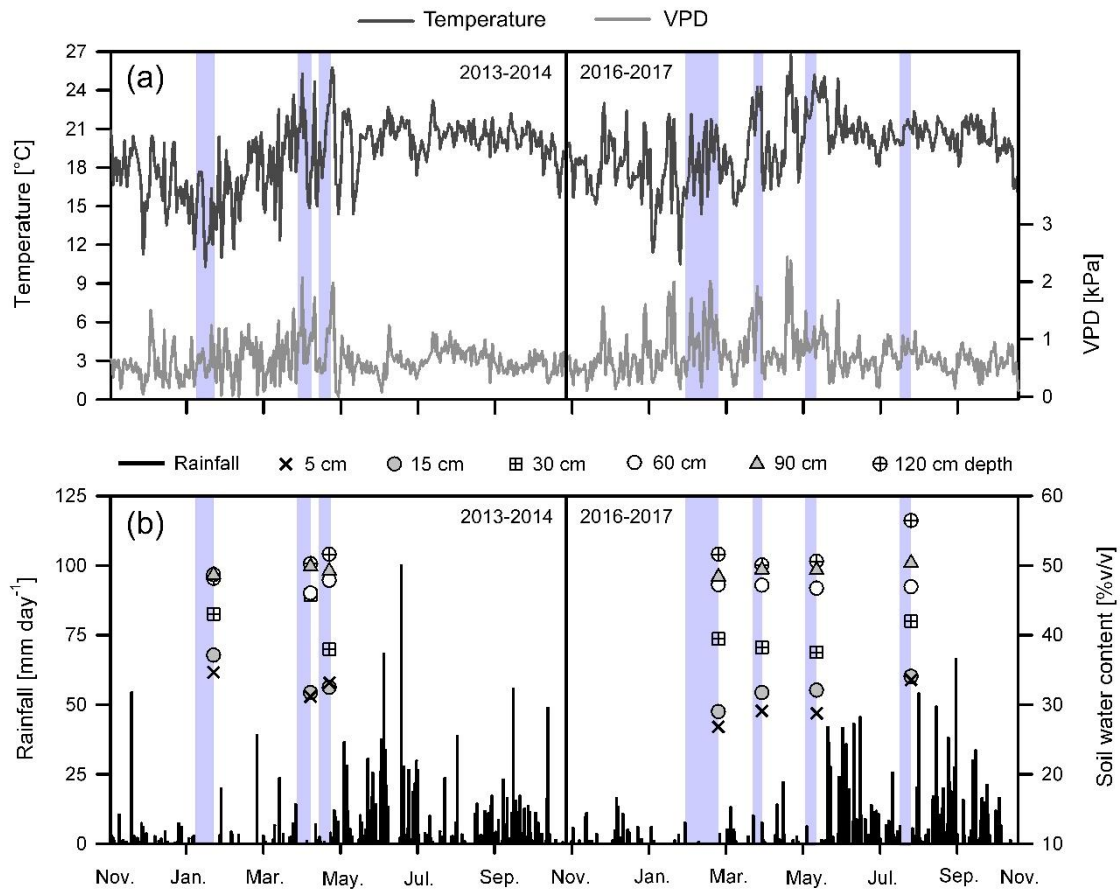


Figure 2. (a) Daily mean air temperature and vapor pressure deficit (VPD) and (b) and daily total rainfall (P), as measured from November 2013 to October 2014 and from November 2016 to October 2017, and volumetric soil water content (SWC) measured at different depths during the sampling campaigns in the study area; different depths are indicated by the unique symbols shown in the lower panels (the key to the symbols is at top). The blue-colored areas indicate the 6- to 22-day period of minimum rainfall (< 5 mm) preceding the dates of isotope sampling in January (mid dry season) and April (late dry season) of 2014, and in February (mid dry season), April and May (late and end of dry season), and August (mid wet season) of 2017.

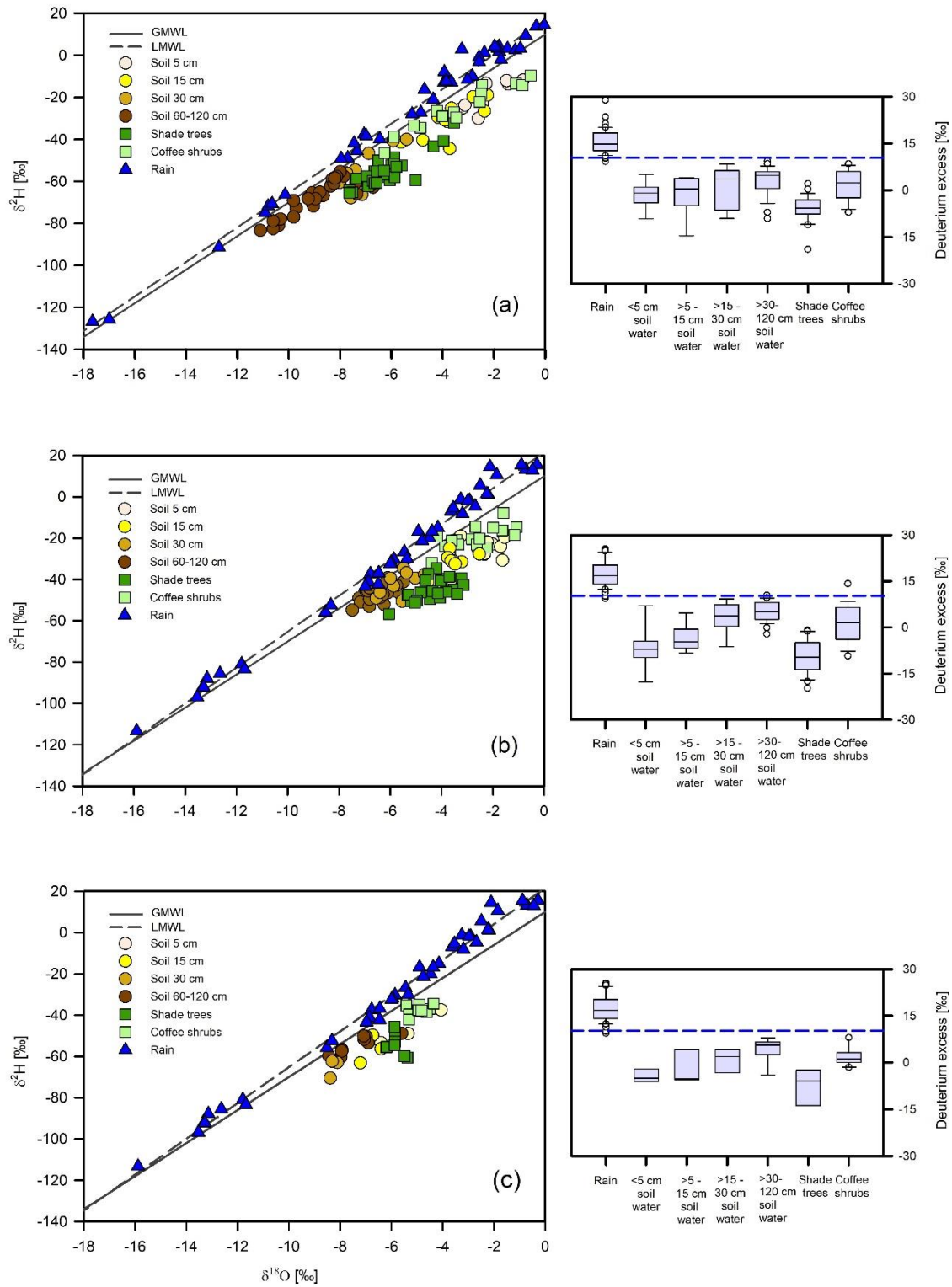


Figure 3. (a) Isotope composition of xylem water for shade trees and coffee shrubs, bulk soil at different depths as observed during the three sampling dates (Jan. 23, Apr. 11 and Apr. 26, 2014),

and isotope values of rainfall during the period December 2013 to November 2014. The dashed line represents the 2013–2014 local meteoric water line (LMWL; $\delta^2\text{H} = 17.82 + 8.26 * \delta^{18}\text{O}$), (b) Isotope composition of xylem water for shade trees and coffee shrubs, bulk soil at different depths during the three sampling dates (Feb. 27, Apr. 5 and May. 20, 2017) and isotope values of rainfall during the period December 2016 to November 2017, and (c) Isotope composition of xylem water for shade trees and coffee shrubs, bulk soil at different depths during the middle of the 2017 wet season (Aug. 4) and isotope values of rainfall during the period December 2016 to November 2017. The dashed lines in panels (b) and (c) represent the 2016–2017 local meteoric water line (LMWL; $\delta^2\text{H} = 21.0 + 8.36 * \delta^{18}\text{O}$). The solid line in all panels represents the global meteoric water line (GMWL; $\delta^2\text{H} = 10 + 8 * \delta^{18}\text{O}$). The panels on the right show the deuterium excess values for the plants and soil water sources and rainfall preceding the sampling campaigns. The dashed blue line represents the deuterium excess value of the GMWL.

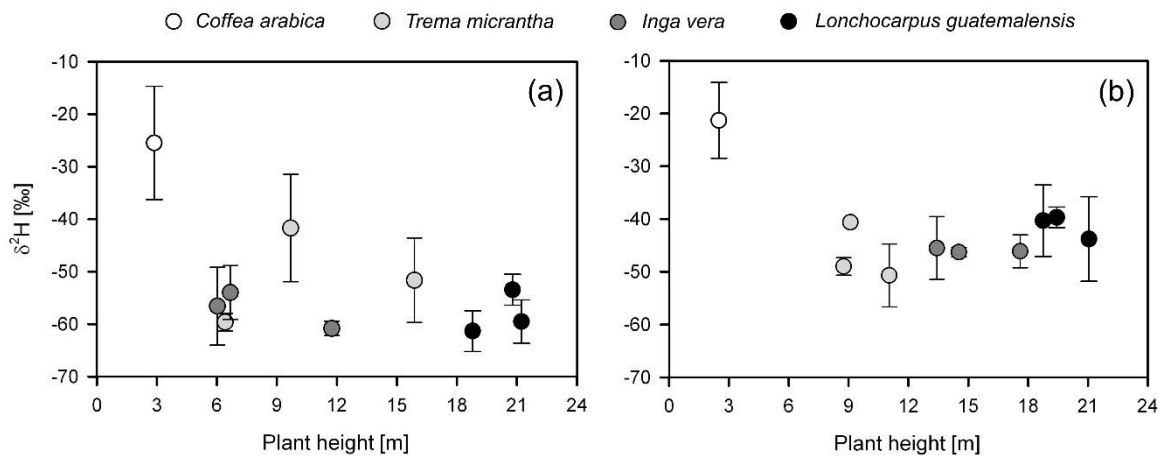


Figure 4. Plant height vs $\delta^2\text{H}$ xylem water for coffee plants and shade tree species corresponding to (a) the 2014 and (b) 2017 dry season samplings.

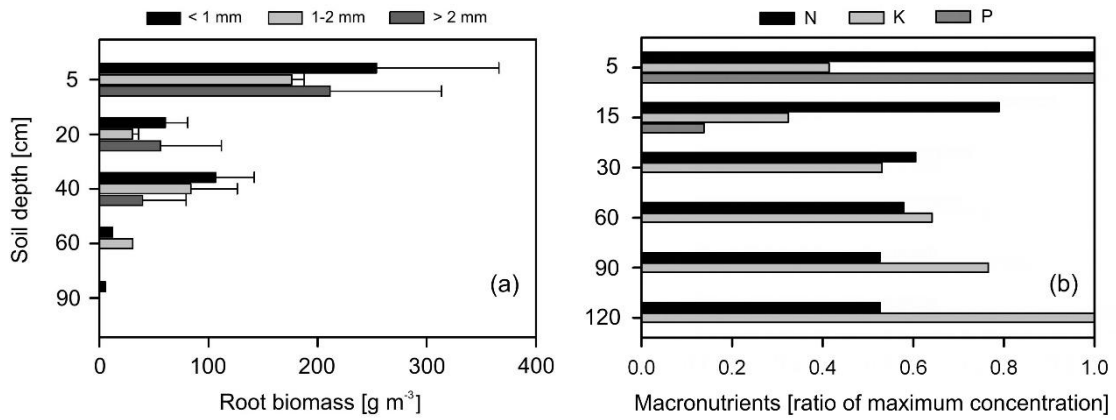


Figure 5. (a) Distribution of root biomass for three size classes of roots (different color bars), the error bars in represent one standard deviation of uncertainty and (b) macronutrients distribution along the soil profile, here normalized and expressed as in ratio to their maximum values (absolute values in Table 4).

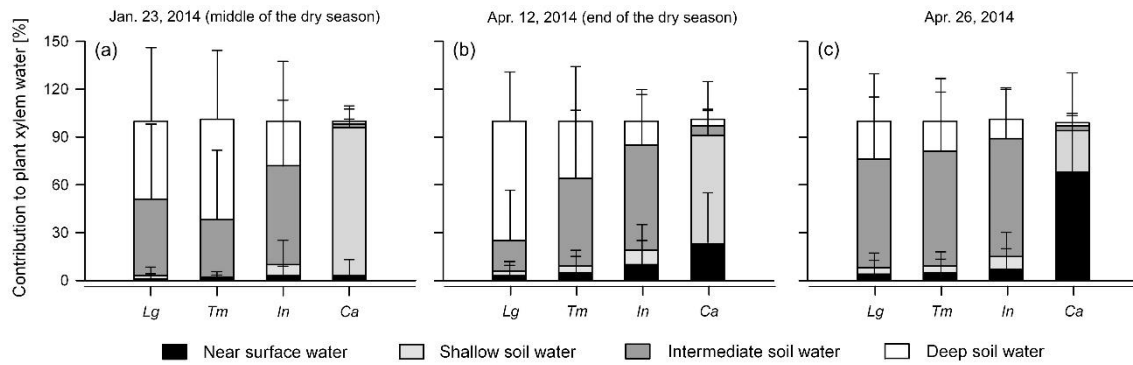


Figure 6. MixSIAR Bayesian mixing model results showing the mean likely contribution of each water source to the xylem water of shade canopy trees and coffee shrubs. (a), (b) and (c) show results for the sampling dates of Jan. 23, Apr. 12 and Apr. 26, 2014 respectively, using the informative prior distribution. *Lg*: *L. guatemalensis*; *Tm*: *T. micrantha*; *In*: *I. vera* and *Ca*: *Coffea arabica*. Error bars represent one standard deviation of uncertainty.

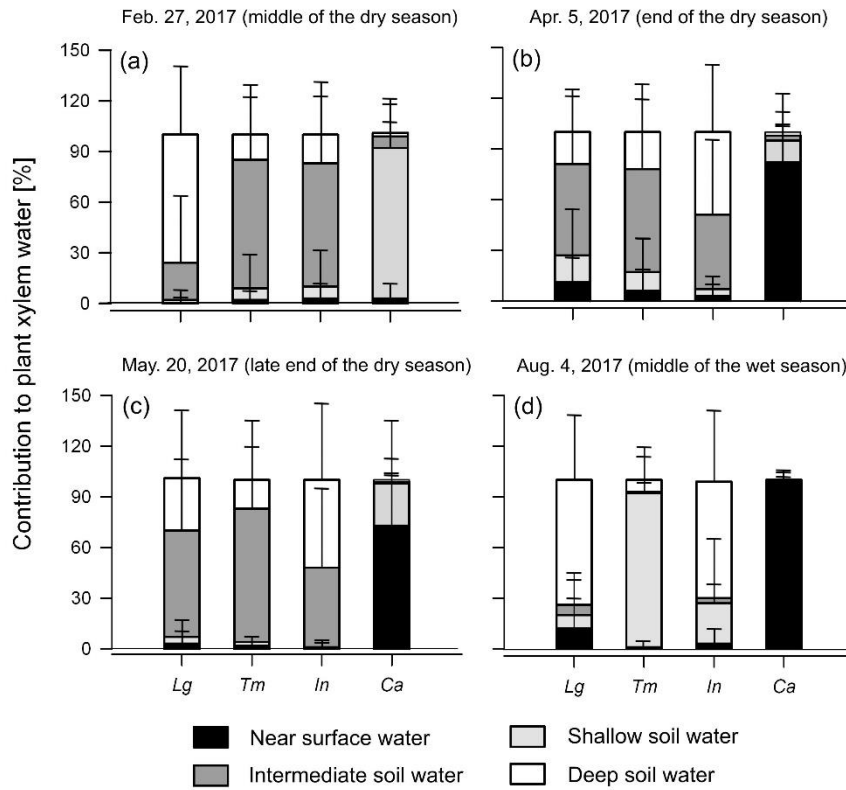


Figure 7. MixSIAR Bayesian mixing model results showing the mean likely contribution of each water source to the xylem water of shade canopy trees and coffee shrubs. (a), (b), (c) and (d) show results for the sampling dates of Feb. 27, Apr. 5, May. 20 and Aug. 4, 2017 respectively, using the informative prior distribution. *Lg*: *L. guatemalensis*; *Tm*: *T. micrantha*; *In*: *I. vera* and *Ca*: *Coffea arabica*. Error bars represent one standard deviation of uncertainty.

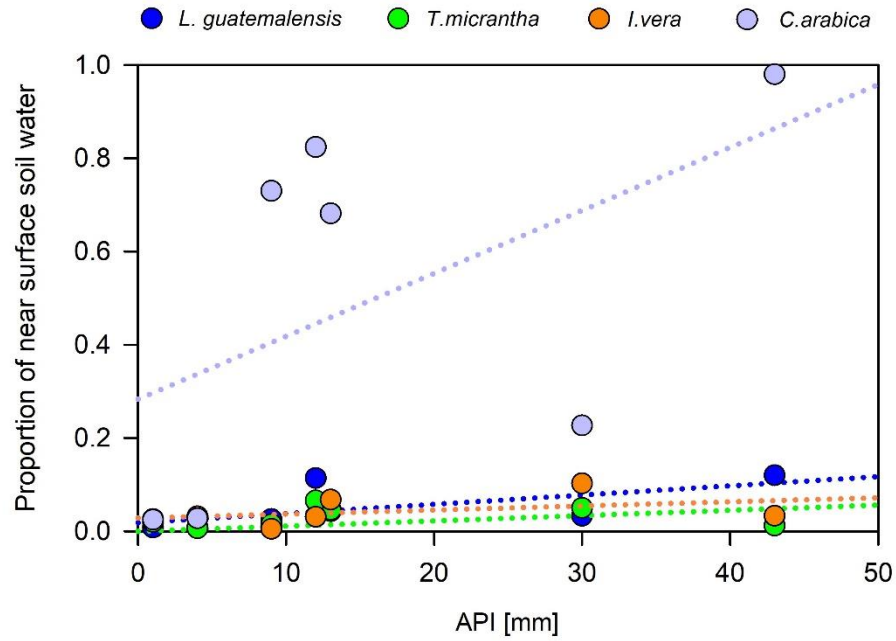


Figure 8. Contribution of near surface~~deep~~ soil water to plant uptake at different antecedent precipitation conditions across the 2014 and 2017 dry seasons.