

1 **Comparison of Generalized Non-Data-Driven Lake and Reservoir**
2 **Routing Models for Global-Scale Hydrologic Forecasting of Reservoir**
3 **Outflow at Diurnal Time Steps**

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11 Abstract: Large-scale hydrologic forecasts should account for attenuation through lakes
12 and reservoirs when flow regulation is present. Globally generalized methods for
13 approximating outflow are required but must contend with operational complexity and a
14 dearth of information on dam characteristics at global spatial scales. There is currently no
15 consensus on the best approach for approximating reservoir release rates in large spatial
16 scale hydrologic forecasting, particularly at diurnal time steps. This research compares two
17 parsimonious reservoir routing methods at daily steps; Döll et al. (2003) and Hanasaki et
18 al. (2006). These reservoir routing methods have been previously implemented in large-
19 scale hydrologic modeling applications and have been typically evaluated seasonally.
20 These routing methods are compared across 60 reservoirs operated by the U.S. Army Corps
21 of Engineers. The authors vary empirical coefficients for both reservoir routing methods
22 as part of a sensitivity analysis. The method proposed by Döll et al. (2003) outperformed
23 that presented by Hanasaki et al. (2006) at a daily time step and improved model skill over
24 most run-of-the-river conditions. The temporal resolution of the model influences models
25 performances. The optimal model coefficients varied across the reservoirs in this study and
26 model performance fluctuates between wet years and dry years, and for different
27 configurations such as dams in series. Overall, the method proposed by Döll et al. (2003)
28 could enhance large scale hydrologic forecasting, but can be subject to instability under
29 certain conditions.

30

1. Introduction

31 1.1. Importance of Dams in Hydrologic Simulations

32 Improvements in numerical weather prediction, the increasing abundance of
33 computational power, and greater precision of remotely sensed observations make global
34 hydrologic forecasting and flood warning systems increasingly feasible (Alfieri et al.,
35 2013; Wu et al., 2014; Emerton et al., 2016; Salas et al., 2017). Lack of information
36 concerning anthropogenic influences on runoff is a major deficiency of large-scale flood
37 forecasting systems (Emerton et al., 2016). Reservoir operations tend to distort natural flow
38 patterns, effectively redistributing surface water spatially and temporally (Zhou et al.,
39 2016). Impoundments significantly influence the downstream flow regime at small and
40 large spatial scales (Batalla et al., 2004; Magilligan and Nislow, 2005). Over half of the
41 world's large river systems are now substantially altered by dams (Nilsson et al., 2005)
42 resulting in a seven-fold increase in water storage within the global river system
43 (Vörösmarty et al. 1997). Furthermore, the cumulative alterations from global reservoir
44 impoundments are so significant that it has been suggested that they could buffer global
45 sea-level rise (Chao et al., 2008).

46 Dams primarily impact the hydrologic cycle by changing the magnitude and timing
47 of the discharges downstream (Haddeland et al., 2006; Döll et al., 2009; Biemans et al.,
48 2011; Wu et al., 2014; Zajac et al., 2017), often with the specific intent to mitigate
49 hydrologic extremes (i.e., floods and droughts) (Zajac et al., 2017). Dams reduce peak
50 discharges by roughly a third on average while dampening the daily variation by a similar
51 amount (Graf, 2006). In hydrologic forecasting, accuracy of the timing and magnitude of
52 hydrologic extremes is fundamentally important to the usefulness of the forecasts.

53 Therefore, the significant impacts from dams make inclusion of reservoir operations, or
54 reservoir routing, critical in large scale hydrologic flood forecasting.

55 Integrating dam operations within large-scale river routing and flood forecasting
56 improves model performance downstream of reservoir locations (Snow et al., 2016;
57 Tavakoly et al., 2017; Salas et al., 2017; Zajac et al., 2017). This is often not feasible at
58 large-scales since there may be multiple entities responsible for regulating flow,
59 particularly with respect to transboundary waters. Among other things, operational
60 knowledge, site-specific rule curves, reservoir uses, and local decision-making practices at
61 each individual project dictate dam releases. Thus, dam operations are typically non-linear,
62 complex processes, driven by anthropogenic and environmental influences. This makes
63 generalizing reservoir operations difficult, particularly in the context of predicting dam-
64 induced hydrologic responses at diurnal or sub-diurnal time step. Heuristically accounting
65 for dams within existing routing schemes should improve flood forecast results when
66 scheduled releases are not readily known.

67 Reservoir routing methodologies are generally divided into two basic categories:
68 data-driven and non-data-driven. Machine-learning, artificial intelligence (Coerver et al.,
69 2017; Macian-Sorribes and Pulido-Velazquez, 2017; Ehsani et al., 2016; Mohan and
70 Ramsundram, 2016; Tielavilca and McKee, 2011; Chaves and Chang, 2008; Khalil et al.,
71 2005), and remote sensing (Bonnema et al., 2016; Yoon and Beighley, 2015) are examples
72 of data-driven approaches. Such data-driven methodologies can be effectively applied to
73 dynamic non-linear systems, particularly when the governing influence on the system does
74 not follow any particular deterministic model. These types of approaches require training
75 data or specific knowledge of a particular reservoir to effectively parameterize and apply

76 them. This is often an insurmountable limitation for data-driven approaches. For that
77 reason, the focus of this paper is on non-data-driven reservoir routing methodologies as an
78 incremental improvement over schemes that effectively neglect dams when information is
79 scarce.

80 1.2. Non-Data-Driven Reservoir Storage and Outflow Simulation

81 Non-data-driven approaches to reservoir routing rely on conceptualizing reservoir
82 responses without explicitly observing the actual reservoir operations. The optimal method
83 for a given application depends on a balance between complexity and available information
84 (De Vos, 2015). Therefore, this manuscript focuses on selecting for parsimony.

85 Existing non-data-driven reservoir models range from simple approaches to
86 sophisticated methods. Solander et al. (2016) showed that temperature-based schema best
87 fits the modeling of discharge, $Q_{out,t}$. The Solander et al. (2016) rule is driven by
88 temperature shifts at each model time step above and below the mean temperature. The
89 Solander et al. (2016) method indicates that temperature is the main proxy governing
90 reservoir release, due to the assumption that seasonality drives agricultural production and
91 reservoir operation. However, the Solander et al. (2016) study focuses on long-term
92 climatic forecasting. Diurnal temperature variations will not likely describe day-to-day
93 reservoir operations. Zhao et al., (2016) developed a reservoir routing scheme based on
94 reservoir stage and storage rules. However, real-time insights related to current reservoir
95 stages throughout a region can involve considerable remotely sensed information. The
96 stage information must then be related somehow to storage volume making this a much
97 more data-driven process. Burek et al. (2013) also developed a non-data-driven approach
98 to reservoir routing which was implemented by Zajac et al. (2017). This approach is built

99 into the LISFLOOD model. The Burek et al. (2013) model requires a number of
100 assumptions about storage capacity limits and naturalized streamflow thresholds. For
101 example, the minimum, normal, and maximum storage are assumed to be 0.1, 0.3, and
102 0.97, respectively. To maintain the objective of investigating parsimonious models, the
103 approach by Burek et al. (2013) was not included in this evaluation.

104 Döll et al. (2003), Wada et al. (2014), and Wisser et al. (2010) presented non-data-
105 driven methods to simulate reservoirs operation that can be considered as simple
106 approaches. The Wisser et al. (2010) method follows a simple, rule-based approach to
107 define the reservoir outflow at each time step ($Q_{out,t}$). The rule that Wisser et al. (2010)
108 enacts is that when the inflow at each model time step moves above or below the long-term
109 average inflow, the behavior of the reservoir release changes. De Vos (2015) suggested
110 that this model is too simple to effectively model reservoir outflow. In a similar vein, Wada
111 et al. (2014) introduced a daily estimate of reservoir outflow that is simply the product of
112 the proportion of available reservoir storage and daily inflow, which can be too simplistic
113 to estimate reservoir outflow since no coefficient is introduced into the simulation to
114 account for reservoir heterogeneity.

115 Döll et al. (2003) derived reservoir routing scheme that can be applied to man-made
116 reservoirs and natural water bodies. The Döll et al. (2003) methodology found genesis in
117 the reservoir outflow model proposed by Meigh et al. (1999). Meigh et al. (1999) proposed
118 a simple reservoir release methodology, which intended to mimic outflow at reservoirs
119 from a theoretical rectangular weir. A more substantive version of the Meigh et al. (1999)
120 method is formulated by Döll et al. (2003). Despite its simplicity, the Döll et al. (2003)
121 method demonstrated good performance compared to several other routing methods (De

122 Vos, 2015). The form of the Döll et al. (2003) equation is similar to that proposed by Wada
123 et al. (2014). However, the Döll et al. (2003) methodology incorporates a coefficient that
124 can incorporate a portion of reservoir heterogeneity.

125 Compared to the aforementioned methods, Hanasaki et al. (2006) derived a demand
126 driven approach to reservoir routing, which can be considered a complicated non-data-
127 driven reservoir routing model. They distinguished between irrigation and non-irrigation
128 reservoirs and offered two distinct algorithms for each. Water demands for irrigation,
129 domestic, and industrial uses are considered in the irrigation reservoirs, whereas the
130 releases from non-irrigation reservoirs are simply a proportion of inflow.

131 De Vos (2015) also proposed a within-year/over-year reservoir routing method
132 comprised of two systems of equations, which was considered a non-data-driven approach.
133 Within-year reservoir operations are driven by yearly fill and release cycles and typically
134 have a small storage capacity relative to their total annual demand. Thus, water
135 accumulates during wet periods and decreases during dry periods. Over-year reservoir
136 operation, on the other hand, is based on long-term, multi-year drawdowns. Over-year
137 reservoirs have storage which is sufficiently large, relative to inflow, so that yearly cycles
138 of water storage and release are not necessary (Adeloye and Montaseri, 2000; Vogel et al.,
139 1999). De Vos (2015) compared his methodology to the Hanasaki et al (2006), Döll et al.
140 (2003), and Neitsch et al. (2011). The De Vos (2015) over-year simulation assumes
141 knowledge of the mean and standard deviation of reservoir storage and is still too data-
142 driven for the purposes of this study. Table 1 summarizes each of the inputs required by
143 each non-data-driven approach described above.

144

Table 1. Input requirements for the various reservoir routing methods.

| | <i>Burek et al. (2013)</i> | <i>Zhao et al. (2016)</i> | <i>De Vos (2015)</i> | <i>Solander et al. (2016)</i> | <i>Döll et al. (2003)</i> | <i>Hanasaki et al. (2006) Non-irrigation Method</i> | <i>Wisser et al. (2010)</i> | <i>Wada et al. (2014)</i> |
|--|----------------------------|---------------------------|----------------------|-------------------------------|---------------------------|---|-----------------------------|---------------------------|
| <i>Reservoir Inflow at time step</i> | X | X | | X | X | X | X | X |
| <i>Empirical Coefficients</i> | | X | | X | X | X | X | |
| <i>Minimum Storage/Inactive Storage Limit</i> | X | X | X | | X | X | | X |
| <i>Maximum Storage/Flood Storage Limit</i> | X | X | X | | X | X | | X |
| <i>Average Storage</i> | | | X | | | | | |
| <i>Standard Deviation of Storage</i> | | | X | | | | | |
| <i>Water Stored at model time step</i> | X | X | | X | X | | | |
| <i>Average Inflow</i> | X | | X | | | X | X | |
| <i>Flood Inflow</i> | | X | | | | | | |
| <i>Air Temperature</i> | | | | X | | | | |
| <i>Conservation Storage Limit</i> | | X | | | | | | |
| <i>Normal Storage Limit</i> | X | | | | | | | |
| <i>Normal Outflow</i> | X | | | | | | | |
| <i>Non-Damaging Outflow</i> | X | | | | | | | |
| <i>Precipitation on the Reservoir</i> | X | | | | | | | |
| <i>Evaporation From the Reservoir</i> | X | | | | | | | |
| <i>Fill Fraction</i> | X | | | | | | | |
| <i>Average Total Winter Inflow</i> | | | | X | | | | |
| <i>Pool Elev. at model time step</i> | | X | | | | | | |
| <i>Pool Elev. at top of inactive storage</i> | | X | | | | | | |
| <i>Pool Elev. at the top of conservation storage</i> | | X | | | | | | |
| <i>Pool Elev. at the top of flood storage</i> | | X | | | | | | |
| <i>Flood Seasonality</i> | | | X | | | | | |
| <i>Standardized Precipitation Evapotranspiration Index</i> | | | X | | | | | |

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The Döll et al. (2003) and Hanasaki et al. (2006) require minimal input data to implement: reservoir inflow, average inflow, and storage volume characteristics. Each of these variables are available in existing datasets, such as the Global Reservoir and Dam

150 (GRanD) database (Lehner et al., 2011) or can be generated using climate reanalysis data
151 (Snow et al., 2016). Other non-data-driven methods require data inputs that are not globally
152 available or produced within the hydrologic simulation (De Vos, 2015; Zhao et al., 2016;
153 Burek et al., 2013; Zajac et al., 2017). For example, the Global Flood Awareness System
154 (GloFAS) is the only existing, operational flood forecasting system that accounts for
155 reservoirs at continental to global spatial extents. However, the reservoir routing
156 component of GloFAS requires operational assumptions be made because of a lack of
157 global reservoir operational records (Zajac et al., 2017). Döll et al. (2003) (hereafter
158 referred to as D03) and Hanasaki et al. (2006) (hereafter referred to as H06) do not require
159 that these assumptions be made because of the minimal inputs which they require. Thus,
160 D03 and H06 meet the requirements of being parsimonious with respect to available
161 reservoir information.

162 The Döll et al. (2003) and Hanasaki et al. (2006) methods also provide enough
163 complexity to account for a portion of the model complexity inherent in reservoir
164 operations. De Vos (2015) does not employ the reservoir routing approach of Wisser et al.
165 (2010) because De Vos (2015) and neither does this research, as it does not account for the
166 status of the reservoir at each simulation time step. The approach taken by Wada et al.
167 (2014) is similar to D03 but represents reservoirs with similar inflow and storage
168 characteristics homogeneously.

169 Furthermore, D03 and H06 methods have been implemented in large-scale
170 hydrologic models. D03 was used in the WaterGAP model and the application of H06 was
171 implemented in the TRIP model by the same authors. The main difference in this
172 evaluation and previous evaluations (i.e., Hanasaki et al., 2006; Masaki et al., 2017) of

173 these reservoir routing schemes is that this research evaluates model performance at a
174 diurnal time step.

175 The aim of this study is to assess non-data-driven reservoir routing methods that
176 are parsimonious and align with available information for use in hydrologic forecasting
177 schemes applicable across the global domain at diurnal time steps. Considering these
178 research aims, the non-data driven reservoir routing methods developed by Döll et al.
179 (2003) and Hanasaki et al. (2006) were considered.

180 The following research questions are addressed with respect to the D03 and H06
181 approaches: (1) How well do the selected reservoir routing models improve outflow
182 estimates relative to simulation of naturalized flow (i.e. neglecting dams altogether)? (2)
183 How do reservoir routing coefficients affect model performance? (3) How does the time
184 step affect model performance and stability? This is a critical point for the current regional-
185 to continental-scale forecasting schemes that operate at daily or sub-daily time steps. (4)
186 How sensitive are the reservoir routing schemes to various real-world dam operations and
187 climate variability?

188 To achieve the research objectives of the study, reservoir data including daily
189 inflow and outflow from 2006-2012, for 60 U.S. Army Corps of Engineers (USACE)
190 reservoirs were used to evaluate the reservoir routing schemes. The data were obtained
191 from nine USACE districts: Pittsburg, Nashville, St. Paul, Rock Island, Omaha, Tulsa,
192 Sacramento, Los Angeles, and Vicksburg. The selected dams are representative of a wide
193 range of reservoir sizes, flow regimes, and climatologic settings but are predominately
194 managed for flood control. The results of this analysis will benefit readers in determining

195 if the reservoir routing models implemented within existing, large-scale hydrologic
196 forecasts adequately represent reservoir effects.

197 2. Methodology

198 2.1. Simulation Specifications

199 The storage ratio (Vogel et al., 1999) or Impoundment Ratio is an important metric
200 in previous works examining generalizing reservoir operation (De Vos, 2015; Hanasaki et
201 al., 2006). The impoundment ratio is described as follows:

$$202 \quad IR = \frac{(S_{max}-S_{min})}{Q_{in}*86400*365} \quad (1)$$

204 where S_{max} and S_{min} are the maximum and minimum volumes of the reservoir's active
205 storage [m³], and Q_{in} is the mean annual inflow to the reservoir [m³s⁻¹].

207 A higher impoundment ratio indicates that the capacity of the reservoir is large
208 relative to mean inflows, while the opposite is true of low IR values. De Vos (2015)
209 considered IR values greater than unity “large” reservoirs, as they are capable of storing
210 the average yearly volume of water flowing into them. To utilize H06, the release
211 coefficient (k_r) needs to be determined.

$$212 \quad k_r = \frac{S_{begin}}{\alpha S_{max}} \quad (2)$$

213 where S_{begin} is the storage [m³] at the beginning of each year and α is a dimensionless
214 coefficient, which was set to 0.85 in the Hanasaki et al. (2006) study. In the current study,
215 the α parameter was varied from 0.45-0.95 by increments of 0.10 and solve k_r for each α
216 value.
217

218 Outflow is the quantity of most interest for hydrologic flood forecasting because
 219 these forecasts generally occur over a relatively short 0-10 day lead time. H06 relates
 220 outflow based on the incoming flow. In this study, only the non-irrigation methodology
 221 from H06 was used to simulate reservoir outflow at each time step ($Q_{out,t}$) since one cannot
 222 assume seasonal irrigation demands will be known globally. Further, the primary purpose
 223 of reservoirs selected in this study is not irrigation. the H06 method estimates outflow as
 224 follows:

$$225 \quad Q_{out,t} = \begin{cases} k_r Q_{in,t} & (IR = 0.5) \\ \left(\frac{IR}{0.5}\right)^2 Q_{in,t} + Q_{in,t} \left\{1 - \left(\frac{IR}{0.5}\right)^2\right\} & (0 < IR < 0.5) \end{cases} \quad (3)$$

227 where $Q_{in,t}$ is the inflow [m^3s^{-1}] at time t and k_r is the release coefficient which is
 228 calculated based on Equation 2. The 0.5 threshold value for IR is an empirical condition
 229 derived by Hanasaki et al. (2006).
 230

231 Unlike H06, D03 relates outflow ($Q_{out,t}$) to current available storage capacity of
 232 the reservoir:

$$233 \quad Q_{out,t} = \frac{k_{rd}}{\Delta t} (S_t - S_{min}) \frac{(S_t - S_{min})^{1.5}}{(S_{max} - S_{min})} \quad (4)$$

234 Where Döll empirically derives the release coefficient, $k_{rd} = 0.01$, Δt is the simulation
 235 time step (s), and S_t is the current volume of storage [$m^3 s^{-1}$] at time t. For this study the
 236 D03, k_{rd} was varied using values of 0.01, 0.02, 0.04, 0.06, 0.08, 0.10, 0.20, 0.40, 0.50, 0.60,
 237 0.70, 0.80, and 0.90.
 238

239 The sensitivity analysis of k_r and k_{rd} can provide useful information on how
 240 coefficients may vary based on geographical and reservoir characteristics such as the
 241 impoundment ratio. The two methods were evaluated and results compared to actual

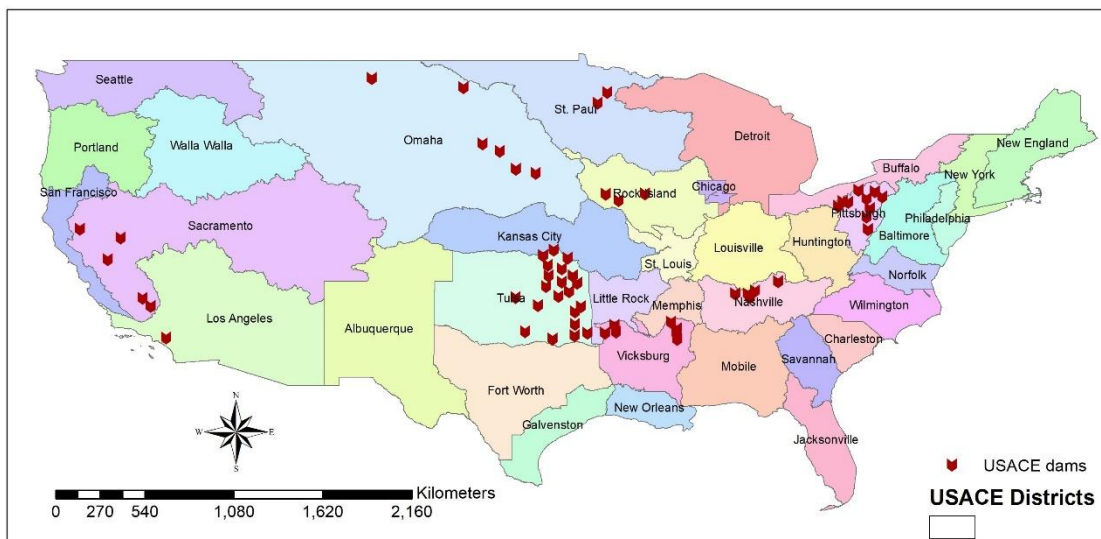
242 outflow records provided by the USACE Districts. Two approaches were used to evaluate
243 model performance: hydrograph assessment of daily and monthly reservoir outflow and
244 statistical evaluation. The statistical evaluation was performed for daily and monthly
245 averaged simulated results vs. observations using the Kling-Gupta efficiency (KGE, Gupta
246 et al., 2009), coefficient of determination (R-Squared), and root mean square error
247 (RMSE). The KGE value ranges from negative infinity to one. Four levels of performance
248 were defined for KGE in this study (Tavakoly et al., 2017): poor performance ($KGE < 0$),
249 acceptable ($0 < KGE < 0.4$), good ($0.4 < KGE < 0.7$), and very good ($0.7 < KGE$).
250 Goodness-of-fit values were evaluated to compare simulated discharge to the actual
251 outflow records provided by the USACE Districts. These are indicators of how well the
252 models perform. The same goodness-of-fit values are calculated to compare actual
253 discharge with inflow to assess baseline performance. The baseline condition represents
254 the treatment of reservoir outflow as naturalized, altogether neglecting reservoir
255 operations. Thus, the baseline condition is that inflow into the reservoir equals outflow
256 from the reservoir. To be viable, the reservoir routing scheme should improve results over
257 the baseline condition in virtually all cases.

258 A true directly measured daily inflow is not available for most reservoirs, including
259 those maintained by the USACE. There are two ways that one can acquire a daily reservoir
260 inflow; estimated using a streamflow model (as in Masaki et al., 2017; Zajac et al., 2017)
261 or estimated using a back calculated inflow based on the known discharge and observed
262 changes in reservoir storage (as in De Vos, 2015). The authors have chosen to utilize a
263 back calculated inflow because this methodology inherently accounts for all other
264 withdraws from the reservoir, such as irrigation, evapotranspiration, seepage, etc. This

265 allows the study to focus exclusively on the reservoir routing methodology. In fact, that
266 would double count withdrawals from the reservoir.

267 2.2. Study Area

268 The model evaluations were conducted on 60 reservoirs in the United States
269 maintained by the U.S. Army Corps of Engineers (USACE). Figure 1 illustrates reservoirs
270 used in this study. The primary purpose of 43 of the reservoirs are flood control, six are
271 hydroelectric, four are recreation, three are water supply, two are classified as other, one is
272 irrigation, and one is a fish and wildlife pond. Despite most reservoirs in the sample being
273 primarily purposed as flood control reservoirs, only three of these reservoirs are exclusively
274 purposed for flood control. Table 1 describes pertinent characteristics of each reservoir in
275 this analysis.



276 Figure 1. USACE districts and location of reservoirs in this study.
277

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Table 2. Select statistical characteristics of reservoirs analyzed in this study.

| Characteristic | Range | Mean | Standard Deviation |
|---|--------------|-------------|---------------------------|
| Minimum Storage ($m^3 \times 10^6$) | 0 - 12,377 | 827 | 2,553 |
| Maximum Storage ($m^3 \times 10^6$) | 25 - 32,070 | 2,695 | 6,184 |
| Annual Inflow (m^3/s) | 0.64 - 780 | 118 | 202 |
| Annual Outflow (m^3/s) | 0.66 - 776 | 113 | 195 |
| Impoundment Ratio | 0.03 - 15.50 | 1.96 | 2.33 |

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3. Results and Discussion

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This section describes the overall results of the study. There is significant improvement in skill over the baseline (the use of inflow as an estimate of outflow) when the optimal D03 coefficient is chosen. D03 tends to outperform the baseline. H06 generally mirrors the results of the baseline. For this reason the discussion largely focuses on D03. The authors examine the distribution of best fitting k_{rd} values. We discuss how dam systems, annual variability, and simulation time step can influence the ability of D03 to estimate reservoir outflow. The authors also discuss the potential for numeric instability in D03 simulations and offer an initial solution to this instability. We also provide an overview of the limitations of this study and suggested future work.

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3.1. Overall Model Performances

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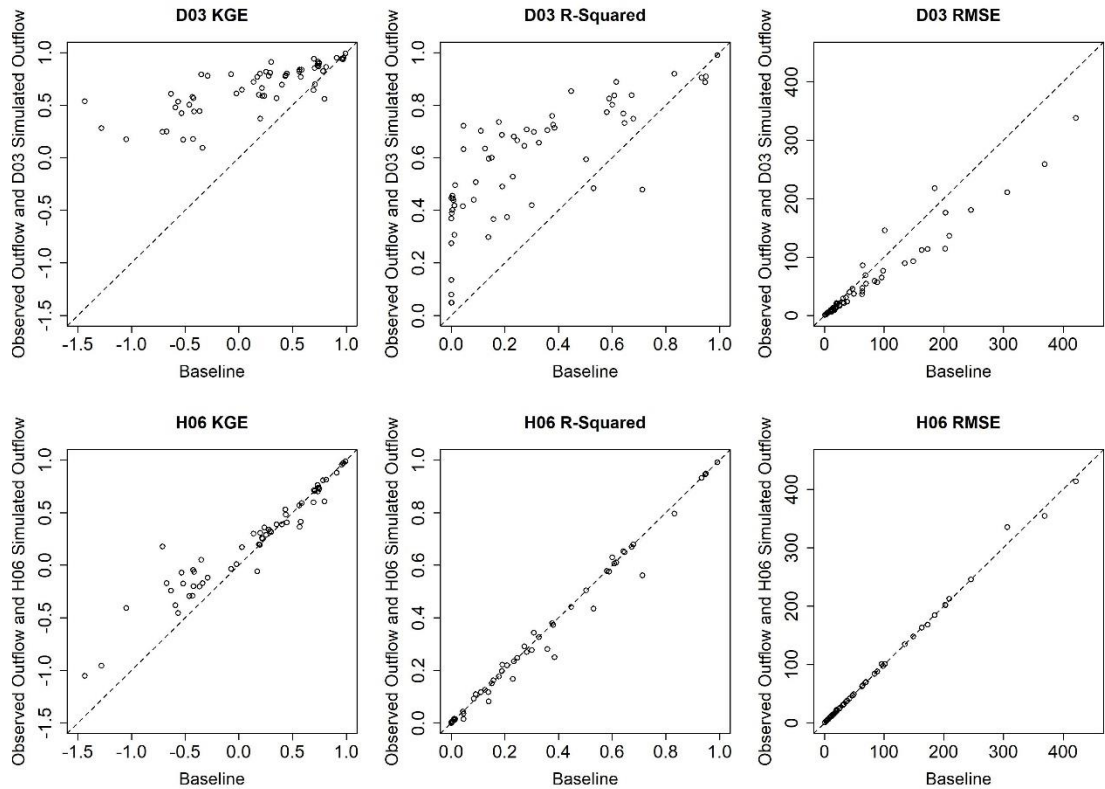
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The goodness-of-fit metrics were calculated for each reservoir in the study. Observed inflow is compared with observed outflow to establish a benchmark used to show whether implementing the two non-data driven reservoir routing schemes improves estimates for reservoir outflow over the use of unregulated flow as the reservoir outflow estimate. Figure 2 illustrates the comparison of skill metrics between baseline and the use of D03 and H06 to simulate outflow. The KGE, R-Squared, and RMSE for D03 and H06

297 in Figure 2 represent the best fit results from the sensitivity study. Data points in Figure 2
298 that fall below the dashed line represent instances where KGE, R-Squared, and RMSE are
299 lower for the reservoir routing method compared to the baseline. Data points falling above
300 the dashed line indicate instances where higher KGE, R-Squared, and RMSE were obtained
301 than the baseline for this study. H06 tends to show minimal utility over the baseline
302 scenario. In general, H06 does not appear to make outflow estimates worse. Estimates that
303 have acceptable KGE values in the baseline scenario tend to produce acceptable results
304 using H06. On the other hand, Figure 2 illustrates that D03 generally tends to increase KGE
305 and R-Squared, and with this increase in goodness-of-fit, decrease RMSE. Thus, the
306 general conclusion is that selecting the optimum D03 release coefficient will ultimately
307 produce an improved estimate of reservoir outflow compared to the baseline. Generally,
308 H06 will produce an estimated reservoir outflow that performs similarly to the baseline
309 scenario.



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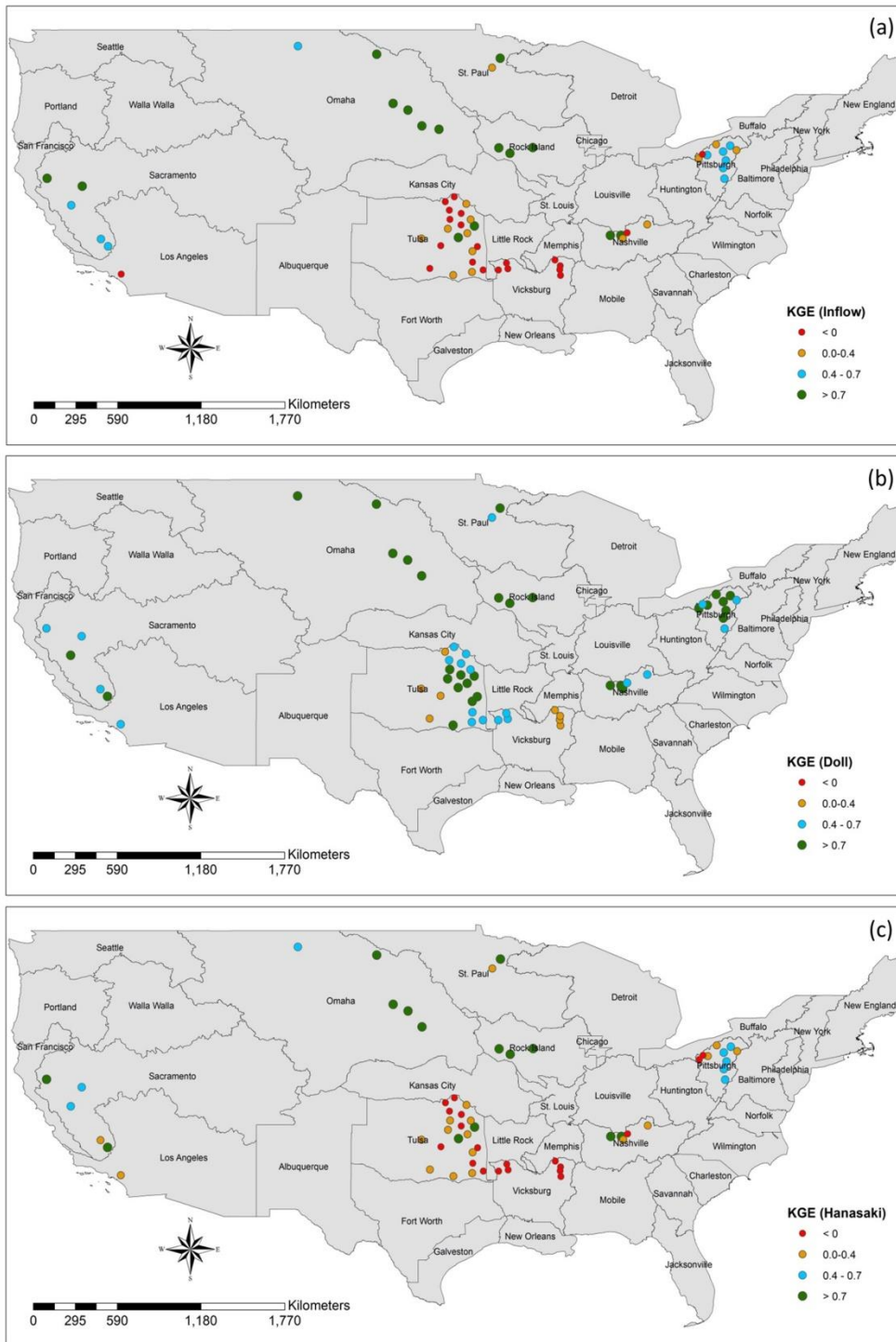
311 Figure 2. Scatter plots of skill metrics between the use of daily observed inflow as outflow (Baseline) and
 312 simulated outflow from best performing D03 and H06 simulations. The dashed line indicates the plane
 313 separating increased and decreased skill that results from using either reservoir routing method.

314 Figure 3 is a geographic representation of the KGE values from the baseline
 315 scenario as well as the best performing implementation of the two routing models for each
 316 reservoir. In general, D03 outperforms the baseline and H06, particularly in the Tulsa and
 317 Pittsburg Districts. H06 tends to provide, at best, minimal improvement in accuracy over
 318 the baseline.

319 D03 tends to improve KGE values at nearly all reservoirs and tends to preserve
 320 high KGE values at locations where the baseline is already a good or very good estimator
 321 of outflow. Only one of the 60 reservoirs in this study demonstrates a significant reduction
 322 in accuracy when D03 is applied. This reservoir, Martis Creek Dam in the Sacramento
 323 District, appears to be an outlier in the reservoir sample. Reservoirs with a similar IR and
 324 average inflow to Martis Creek Dam and in the same USACE district tended to experience

325 improvement in model skill with D03. Overall, when the appropriate k_{rd} value is applied,
326 D03 improves simulation results over the baseline.

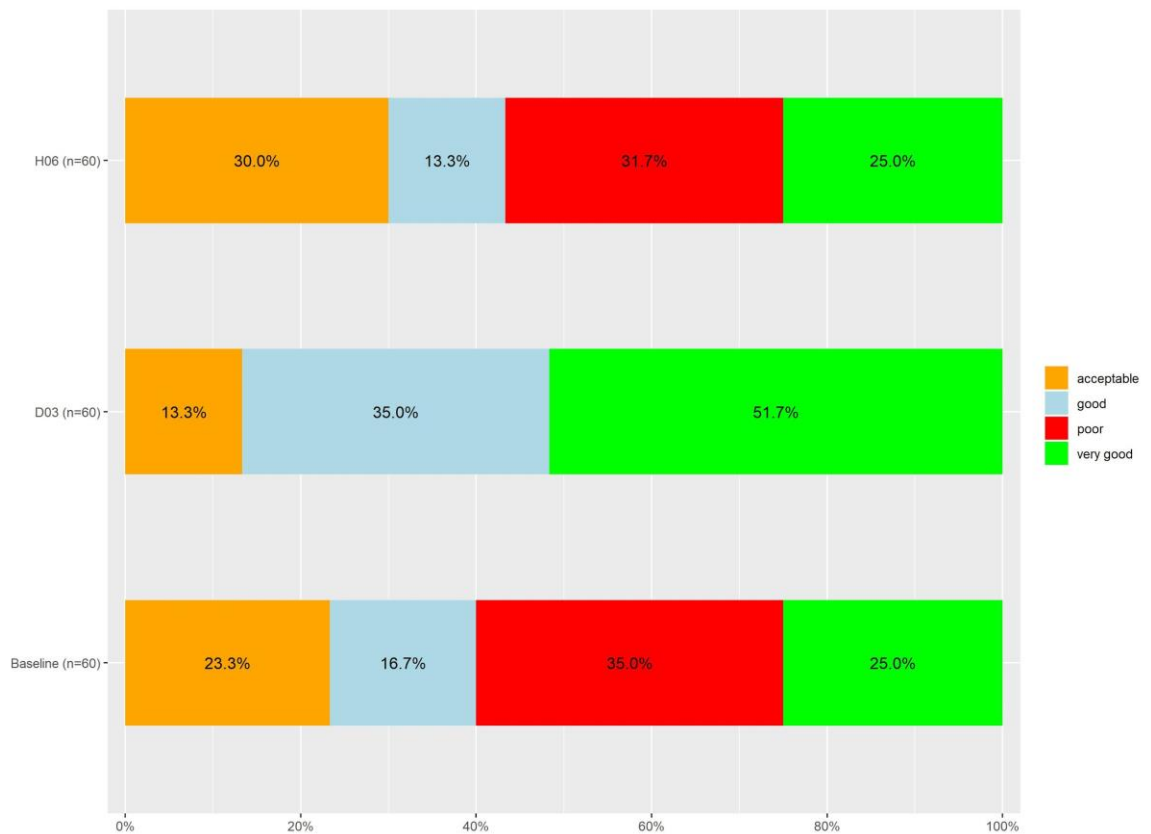
327 Figure 3a illustrates the wide range of reservoir operating conditions present in the
328 study. The reservoir dataset contains reservoirs in which the outflow correlates poorly with
329 the inflow regime as others that correlates well. Figure 3a also portrays significant
330 geographic clustering where reservoirs in certain regions tend to be less correlated with
331 inflow and other clusters where observed inflow and observed outflow correlate strongly.
332 This could indicate that operations at these reservoirs may have a particularly regional
333 context and may bias towards a particular reservoir routing scheme. However, correlation
334 between observed inflow and observed outflow and geographic proximity of the reservoirs
335 does not influence the implementation of either D03 or H06. Thus, the results of this
336 research indicate no significant geographic constraints in the context of this study.



337

338 Figure 3. Spatial distribution of KGE comparing observed daily outflow to the each best estimate of outflow:
 339 a) observed inflow b) Döll Method simulated outflow, c) Hanasaki Method simulated outflow for all
 340 reservoirs in this study. KGE values for the Döll Method and the Hanasaki Method are the maximum KGE
 341 from all coefficient treatments.

342 Figure 4 presents a proportional bar chart comparing baseline KGE and the highest
 343 KGE value for the range D03 and H06 coefficients. This plot categorizes KGE
 344 performance using the same bins as Figure 3. Figure 4 indicates that the best performing
 345 H06 simulation provides only marginal improvement over the baseline condition.
 346 However, the best performing instance of D03 eliminates all poor performing baseline
 347 conditions. Nearly 87% of all best performing D03 simulations are considered to be good
 348 or very good at accurately capturing reservoir outflows, a 22% increase above the baseline
 349 simulation.

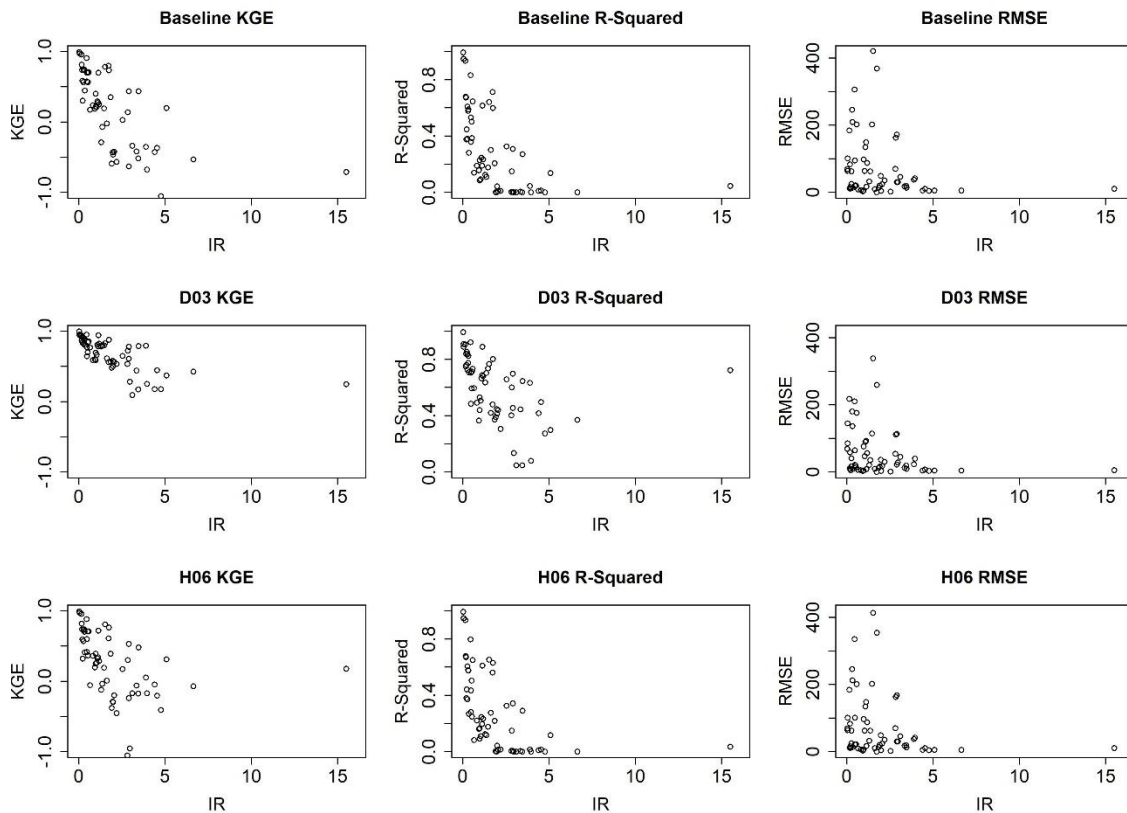


350

351 Figure 4. Proportional bar chart comparing the baseline outflow estimation and the best KGE results for D03
 352 and H06.

353 From multivariate comparison, a negative relationship between two of the best fit
 354 results (KGE and R-Squared) and reservoir IR was found. Figure 5 illustrates this
 355 comparison between IR and each goodness of fit metric for the baseline, D03, and H06.

356 KGE in particular appears non-linearly correlated to IR. A similar, yet less significant,
 357 negative relationship was found between IR and R-Squared. Little statistical correlation
 358 appears to occur between IR and RMSE. However, KGE and R-Squared values in Figure
 359 5 indicate that the ability to predict outflow using the reservoir routing techniques applied
 360 in this study decreases with reservoir with high IR values.



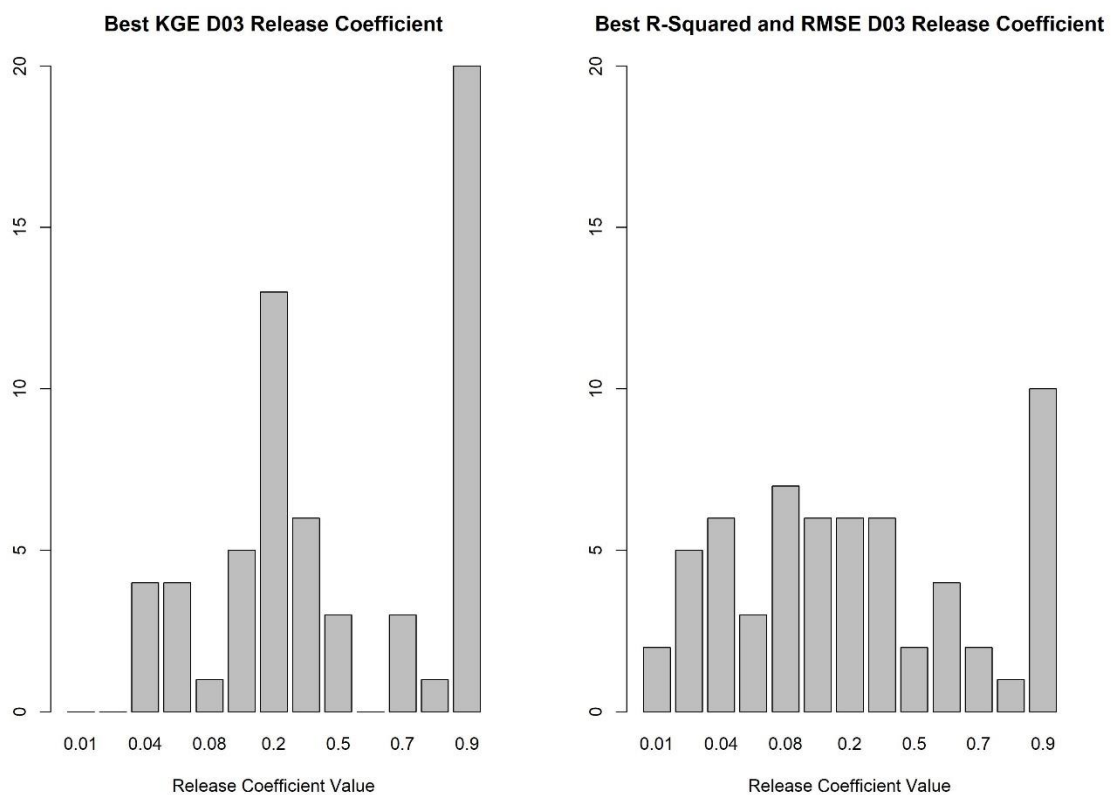
361

362 Figure 5. Comparison of IR and best KGE, R-Squared, and RMSE from goodness of fit metrics for baseline,
 363 D03, and H06.

364 3.2. Sensitivity Analysis of Models

365 Because D03 consistently outperforms H06 at daily time steps, D03 was selected
 366 for the sensitivity analysis at daily time steps. The value of k_{rd} coefficient was introduced
 367 as 0.01 in the Döll et al. (2003) study. In this study, k_{rd} values were varied to obtain
 368 maximum KGE and R-Squared and minimum RMSE. Figure 6 demonstrates the

369 dispersion of k_{rd} values which maximize the model skill for all reservoirs in this study.
 370 For all model skill metrics, $k_{rd}=0.90$ tends to be the most prevalent k_{rd} value that
 371 maximizes model skill. In only two of the 60 reservoirs (Sardis Dam and Enid Dam) $k_{rd} =$
 372 0.01 maximizes R-Squared and minimizes RMSE for the range of k_{rd} coefficients. This
 373 research suggests that the $k_{rd} = 0.01$ is not necessarily the optimum coefficient to
 374 maximize model performance using a daily simulation time step.



375

376 Figure 6. Bar charts of k_{rd} values that maximize KGE and correlation and minimize RMSE.

377 Investigating the linkage between dam characteristics and the best performing k_{rd}
 378 yields no clear relationship. Evaluation of correlation between IR, coefficient of variation
 379 of inflow, ratio of average inflow to average outflow, and geographic location shows low
 380 correlation between each variable and best performing k_{rd} value. However, the range of

381 best performing k_{rd} within this analysis and as demonstrated in Figure 6 suggests that the
382 value is not constant across all reservoirs. Thus, as one implements D03 within their
383 hydrologic forecasting framework, k_{rd} may be adjusted to optimize streamflow estimates
384 to gage observations, like those curated by the Global Runoff Data Centre (GRDC, 2018),
385 when available.

386 3.3. Dam Systems and Reservoir Routing

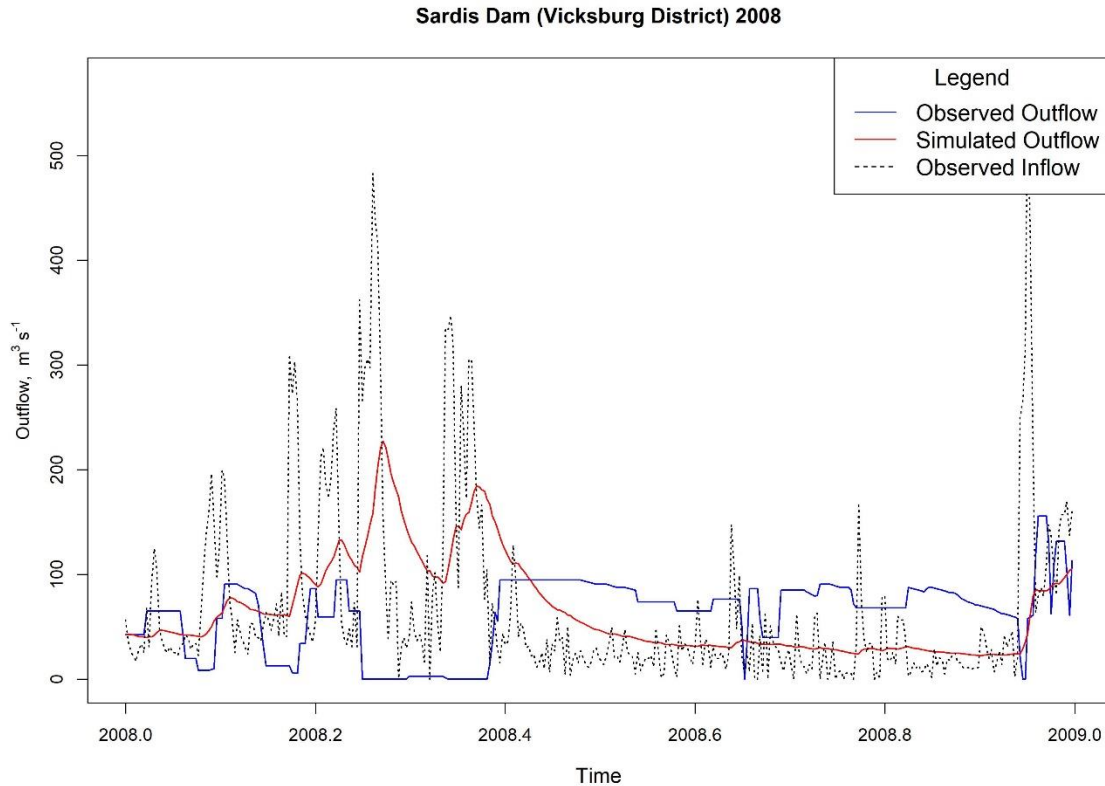
387 Reservoirs in the Vicksburg and Omaha districts were selected to evaluate
388 performance of D03 in environments where reservoirs operate in a coordinated fashion.
389 We broadly refer to these as dam systems. The case of the Vicksburg and Omaha district
390 reservoirs highlights two distinct types of dam systems; one where the dams do not
391 contribute inflow into one another but still coordinate their releases (in parallel) and another
392 where upstream releases flow into downstream reservoirs (in series).

393 A subset of the reservoirs in the Vicksburg District comprise the Yazoo Basin
394 Headwaters Project. Although the reservoirs in the Yazoo Basin Headwaters Project are
395 not directly connected, the reservoir operators coordinate operations in order to minimize
396 flooding in Mississippi's Delta region (Arkabutla Lake History, 2017; USACE, 1987). The
397 operation of these reservoirs presents an interesting case in which the non-date driven
398 models in this study do not characterize the nature of the dam releases well. The modeled
399 results at four Vicksburg District dams yield only minimal improvement over unregulated
400 (i.e. naturalized) flow at these reservoirs. The decrease in reservoir routing performance
401 can be attributed to the large impoundment ratios at these dams indicating the reservoir
402 storage is large relative to annual volume of inflow.

403 The reservoirs of interest in the Vicksburg District include Arkabutla, Sardis, Enid,
404 and Grenada. These dams function in parallel on tributaries of the lower Mississippi River,
405 namely the Coldwater River, Little Tallahatchie River, Yocona River, and Yalobusha
406 River, respectively. Together, these dams control flooding in northern Mississippi as part
407 of the Yazoo Basin Headwaters Project (Arkabutla Lake History, 2017; USACE, 1987).
408 The Yazoo Basin reservoirs discharge directly into the heavily regulated Mississippi River
409 (Meade and Moody, 2010). The reservoirs operate to ensure high releases are not
410 concurrent with large flows upstream on the Mississippi to avoid devastating flooding to
411 the low-lying Louisiana delta regions. This requires a high level of coordination throughout
412 the Yazoo Basin Headwater Project and with regulation upstream on the Mississippi.
413 Additionally, each of the Yazoo Basin reservoirs have a substantial impoundment ratio,
414 ranging from 2.96-3.95. In other words, the reservoirs are capable of containing large
415 volumes of water to mitigate downstream impacts. Thus, current pool levels and forecasted
416 inflow at these four reservoirs do not substantially influence release decisions. The
417 reservoirs also have the capacity to absorb large flood events. As a result, they do not seem
418 to follow the same functional form as the majority of dams in this study.

419 Figure 7 from Sardis Dam in the Yazoo Basin Headwaters Project demonstrates the
420 hydrograph comparing observed inflow and outflow and the modeled outflow that provides
421 the highest KGE (D03, $k_{rd}=0.90$) for the year 2008. Figure 7 demonstrates that peak
422 outflows do not tend to correspond to the time at which peak inflow occurs. In fact, release
423 rates at Sardis Dam are at a minimum during the peak inflow time period. This pattern
424 repeats at each of the reservoirs in the Yazoo Basin Headwaters Project indicating that
425 inflow and consumed storage are not substantial predictors of outflow timing at these

426 reservoirs. This exemplifies the lack of correlation between observed inflow and observed
427 outflow at reservoirs within the Yazoo Basin Headwaters Project.

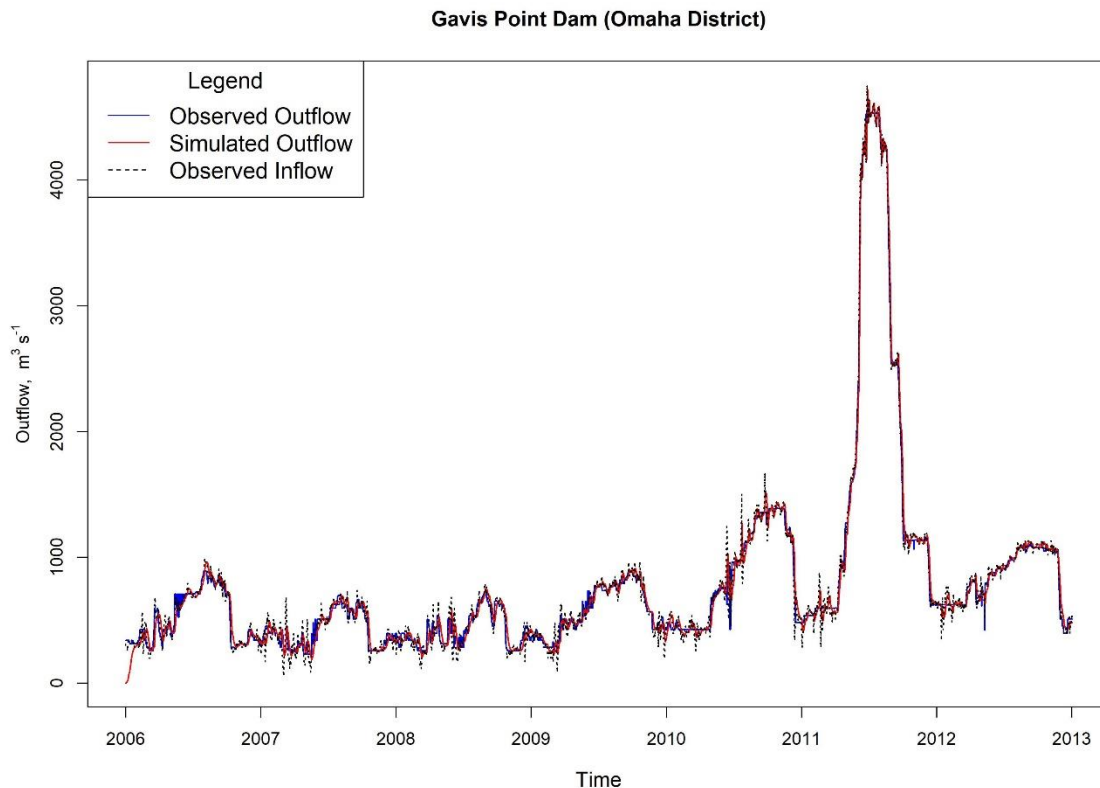


428

429 Figure 7. Hydrographs of observed inflow and outflow versus simulated outflow with the highest KGE value
430 at Sardis Dam (Döll method $k_{rd}=0.90$). KGE comparing observed Inflow and outflow = - 0.34; KGE
431 comparing simulated and observed outflows= 0.095

432 Dams operating in series represent a specific case where compounding model error
433 is a particular concern. USACE operates several large dams in series on the Missouri River.
434 These include Fort Peck, Garrison, Oahe, Big Bend, Fort Randall, and Gavins Point within
435 in the Omaha District (Lund and Ferreira, 1996). For this cascading system on the Missouri
436 River, inflow appears to be a progressively stronger predictor of outflow from upstream to
437 downstream. At the upstream end the baseline yielded a KGE=0.43 at Fork Peck with a
438 KGE=0.99 downstream at Gavins Point Dam. Figure 8 provides a comparison of observed

439 inflow and outflow along with simulated outflow for Gavins Point Dam. D03 tends to
440 provide a slightly better estimate of outflow compared with inflow, except in the instance
441 of Big Bend Dam. At Big Bend Dam, H06 produces an estimate of outflow more consistent
442 with observed outflow than either D03 or inflow alone. However, the differences are almost
443 trivial considering how well inflow alone performed in this case. D03 is particularly
444 accurate during peak inflow conditions, for example the large hydrologic event in mid-
445 2011 at Gavins Point Dam in Figure 8. The performance of non-data driven approaches in
446 this instance is promising since compounding errors are a large concern in this type of
447 system. Other instances involving dams in series should be evaluated to determine out if
448 these findings hold more generally.



449

450 Figure 8. Hydrographs of observed inflow and outflow versus simulated outflow with the highest KGE value
451 at Gavins Point Dam (Döll method $k_{rd}=0.04$). KGE comparing observed Inflow and outflow = 0.99; KGE
452 comparing simulated and observed outflows= 0.99.

453 Reservoir management is unique in both the Yazoo Basin Headwaters Project and
454 the Missouri River. The operators of dams within the Yazoo Basin Headwaters Project tend
455 to regulate outflow in a manner that is more in line with downstream conditions. The
456 attention to downstream conditions is due mainly to the impact that downstream floods will
457 have on the low-lying communities within the Louisiana Delta. The dams in the Yazoo
458 Basin Headwaters Project have among the highest impoundment ratios, which inherently
459 reduces the influence of upstream conditions in discharge decisions. The non-data driven
460 approaches evaluated here do not account for downstream conditions and thus do not
461 perform well in this instance, particularly where large impoundment ratios allow operators
462 considerable leeway.

463 On the other hand, the non-data driven approaches tend to perform well when
464 inflow conditions dictate discharge decisions as we see on the Missouri River system.
465 Reservoirs with smaller impoundment ratios are naturally more responsive to inflow
466 requiring greater consideration for upstream conditions. D03 showed relatively small
467 improvement of outflow estimates compared to inflow as a predictor of outflow in the
468 Yazoo Basin Reservoirs, while the method provided reasonable estimates in dam systems
469 like the Missouri River system. Therefore, it can be inferred that D03 is more applicable
470 for dam systems where reservoir management focuses on upstream hydrologic conditions,
471 while large impoundment ratios may be indicative of reservoirs where downstream
472 conditions are more likely to prevail. This would likely apply for H06 as well since that
473 method links outflow to inflow more directly.

474 3.4. Wet and Dry Year Comparison

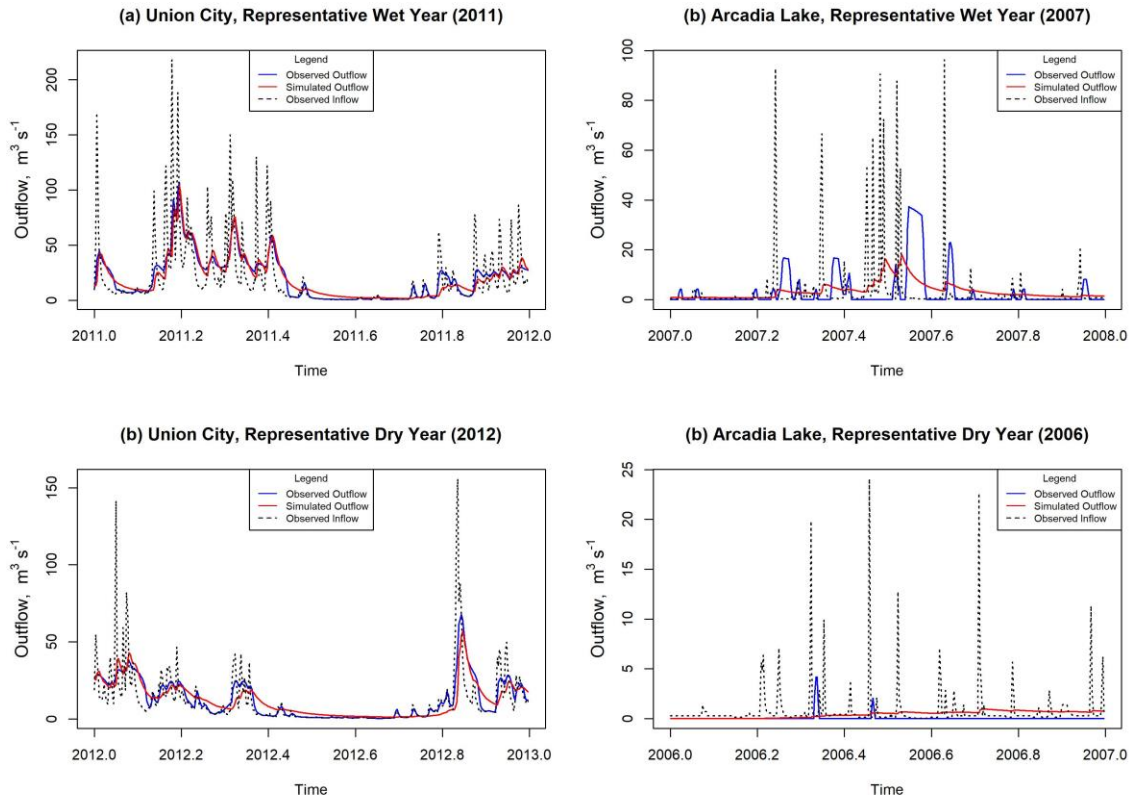
475 Figure 8 shows results for wet and dry years at two reservoirs considered to be
476 representative of this study. D03 provides a relatively good estimate of outflow at Union
477 City Dam (Pittsburg District) in Figure 9a and Figure 9c. D03 performs relatively poorly
478 at Arcadia Lake (Tulsa District) in Figure 9b and Figure 9d. In the case of Union City Dam,
479 D03 tends to produce a noticeable improvement in model skill during both a relatively wet
480 year and a relatively dry year. The performance (Figure 9a and Figure 9c) seems to be
481 independent of wet or dry conditions, at least on an annual basis. This does not hold for
482 Arcadia Lake. The model shows modest skill at Arcadia Lake during the wet year (Figure
483 9b), but almost none during the dry year.

484 There appears to be a difference in the timing discharges between at the two
485 locations in Figure 9. D03 appears to estimate the right amount of volume released during
486 the wet year at Arcadia Lake (Figure 9b). However, the timing of the observed release is
487 delayed until a relatively dry period begins. The lag could indicate that water is being
488 retained, possibly for use in irrigation or domestic supply. In this instance, Arcadia Lake
489 supplies water to the city of Edmond, Oklahoma which may influence release decisions
490 (Arcadia Lake, 2020).

491 D03 performs much more poorly during the 2006 dry year at Arcadia Lake (Figure
492 9d). The model does not predict the sporadic releases throughout the year. The inflow
493 events in that year are not substantial enough to affect storage meaningfully, thus we see
494 almost no response in the modeled output. Observed outflows demonstrate that beyond two
495 relatively high-volume reservoir releases during 2006, the reservoir releases are restricted
496 to practically no outflow the rest of the year. D03 does not anticipate the two large releases,

497 as the reservoir storage does not dramatically shift in either instance. D03 estimates a near
498 constant discharge over the entire year with almost no storage change.

499 Results for wet years and dry years appear to be fairly mixed. Indications are that
500 the performance of D03 could be somewhat site specific. However, reservoirs that tend to
501 be less responsive to storage fluctuations are not represented well in D03 since storage
502 fluctuations drive the model. Arcadia Lake has an IR of about 4.75 which is relatively high.
503 Union City Dam has an IR of about 0.24, which is relatively low. IR is a good indicator of
504 reservoir responsiveness to storage fluctuations. A lack of reservoir responsiveness to
505 storage fluctuations could result in two different types of error when D03 is implemented
506 within a large-spatial-scale hydrologic model. First, forecasted outflow could easily
507 mistime a hydrologic event, particularly during wet years, as Figure 9b demonstrates.
508 Second, the authors anticipate that if the storage does not dramatically fluctuate during a
509 dry year the estimated reservoir release will not anticipate sporadic releases for irrigation
510 and other purposeful discharges. Unaccounted for, these large but short duration releases
511 may lead to a consistent overestimation of reservoir outflow for the entire dry year period.



512

513 Figure 9. Two reservoirs where D03 tends to perform very good and poor: outflow: a) wet year Union City
 514 Dam 2011; b) wet year Arcadia Lake 2007; c) dry year Union City Dam 2012; and d) dry year Arcadia
 515 Lake 2006.

516 3.5. Effects of Time Step on Model Performance

517 Model comparisons are conducted for daily and monthly time steps. Table 2
 518 illustrates the results at Fort Peck, Garrison Dam, Oahe Dam, and Fort Randall Dam, each
 519 of which appears in the Hanasaki et al. (2006) study and this research. Table 2 also contains
 520 Sardis Dam, Mosquito Creek Dam, and Prado Dam, which are not included in Hanasaki et
 521 al. (2006). Results illustrate that the time scale at which comparisons are conducted can
 522 influence simulation results. The monthly comparison amongst Fort Peck, Garrison, Oahe,
 523 and Fort Randall is in agreement with the conclusions of Hanasaki et al. (2006). However,
 524 when the simulation time step changes to a daily time step, the skill of H06 and D03 reverse
 525 and D03 tends to outperform H06. In additional reservoirs (Sardis and Prado), the results

526 indicate that D03 outperformed H06 at both daily and monthly time steps, based upon
527 KGE. However, the results at Mosquito Creek reservoir tend to follow the original
528 Hanasaki et al. (2006) results.

529 The time-scale effect upon model performance may relate to how well observed
530 inflow correlates with observed outflow. Examining Table 2, H06 outperforms D03 when
531 observed inflow and observed outflow are relatively well correlated. The effect is nullified
532 when the inverse is true. H06 estimates outflow as a ratio of inflow, which may be a better
533 estimate of outflow at the monthly time scale, particularly when discharge tracks closely
534 with inflow. However, H06 will fluctuate at the smaller time steps due to inherent
535 variations in inflow. D03 tends to vary less at a daily time step and may be a better estimate
536 of outflow at sub-monthly time steps.

537 The hydrographs from Fort Randall Dam further illustrate the relationships between
538 time step and model skill, particularly during high flow events. Daily and monthly
539 comparisons between observation and simulations for Fort Randall Dam are shown in
540 Figure 10. Figure 10 compares the daily and monthly simulations with observations. Figure
541 10a shows that the H06 simulations perform better than D03 for monthly time steps,
542 particularly during the high inflow periods in 2011. D03 tends to overestimate reservoir
543 outflow, while H06 correlates well with inflow and better matches the peak flow of 2011.
544 At a diurnal time step (Figure 10b), H06 tends to be hypersensitive to inflow variations and
545 overestimates outflow, whereas D03 provides a better approximation of outflow during the
546 2011 high flow event at a daily time step.

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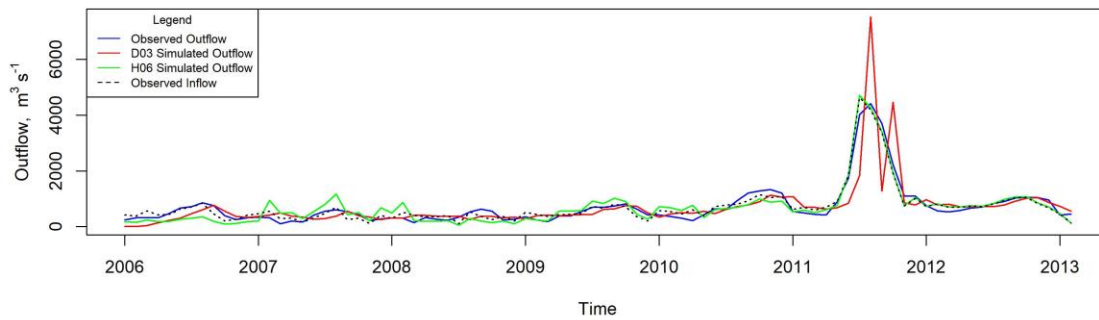
Table 3. Comparison of daily and monthly KGE values at selected reservoirs. The α and k_{rd} values represent the highest KGE values for Hanasaki and Döll methods respectively.

| Reservoir | Daily KGE | | | Monthly KGE | | |
|--|-----------|----------|------|-------------|----------|------|
| | Inflow | Hanasaki | Döll | Inflow | Hanasaki | Döll |
| Fort Peck $\alpha=0.95$ $k_{rd}=0.04$ | 0.43 | 0.53 | 0.78 | 0.54 | 0.62 | 0.51 |
| Garrison Dam $\alpha=0.95$ $k_{rd}=0.06$ | 0.73 | 0.76 | 0.88 | 0.78 | 0.80 | 0.59 |
| Oahe Dam $\alpha=0.95$ $k_{rd}=0.20$ | 0.78 | 0.81 | 0.83 | 0.84 | 0.86 | 0.76 |
| Fort Randall Dam $\alpha=0.95$ $k_{rd}=0.20$ | 0.91 | 0.88 | 0.95 | 0.96 | 0.93 | 0.67 |
| Sardis Dam $\alpha=0.95$ $k_{rd}=0.90$ | -0.34 | -0.17 | 0.09 | 0.06 | -0.03 | 0.16 |
| Mosquito Creek Dam $\alpha=0.45$ $k_{rd}=0.70$ | -0.46 | -0.29 | 0.51 | 0.49 | 0.60 | 0.39 |
| Prado Dam $\alpha=0.95$ $k_{rd}=0.50$ | -0.02 | 0.01 | 0.61 | 0.32 | 0.61 | 0.71 |

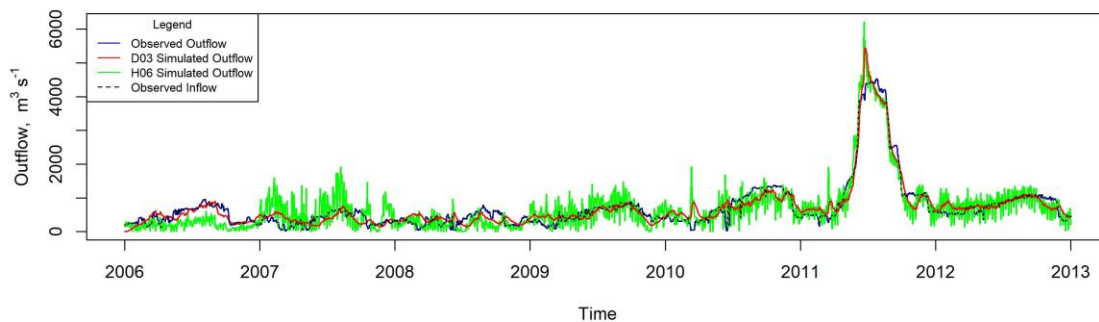
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552

553 It is possible that the conclusions of Hanasaki et al. (2006) suggesting better performance
554 of H06 at the monthly-scale depend on how closely discharge from the dam tracks inflow.
555 D03 may be a better candidate for integration into daily flow forecasting models.

(a) Monthly Comparison of Estimated Reservoir Outflows



(b) Daily Comparison of Estimated Reservoir Outflows



556

557 Figure 10. Comparison of simulated outflow for the Fort Randall Dam with Hanasaki and Döll methods for
558 (a) monthly and (b) daily time steps.

559 3.6. Model Stability

560 Although D03 outperformed H06 when using a daily time step, D03 demonstrated
561 some instability for high k_{rd} values. This instability occurs at three reservoirs in this study.
562 The cause of the instability is a combination of a reservoir having a low IR and a sharp
563 change in the inflow to a reservoir. For instance, inflow into Old Hickory Dam in the
564 Nashville District (IR = 0.04) increased by roughly two orders of magnitude in a matter of
565 a few days in May 2010. During this event, the available storage filled up, necessitating a
566 substantial increase in release flow to prevent overtopping. This occurred within a single
567 time step in the model (D03) and the outflow responded in kind in the next subsequent time

568 step which then drained the reservoir below the specified minimum storage resulting in a
569 non-computable imaginary number as the next solution.

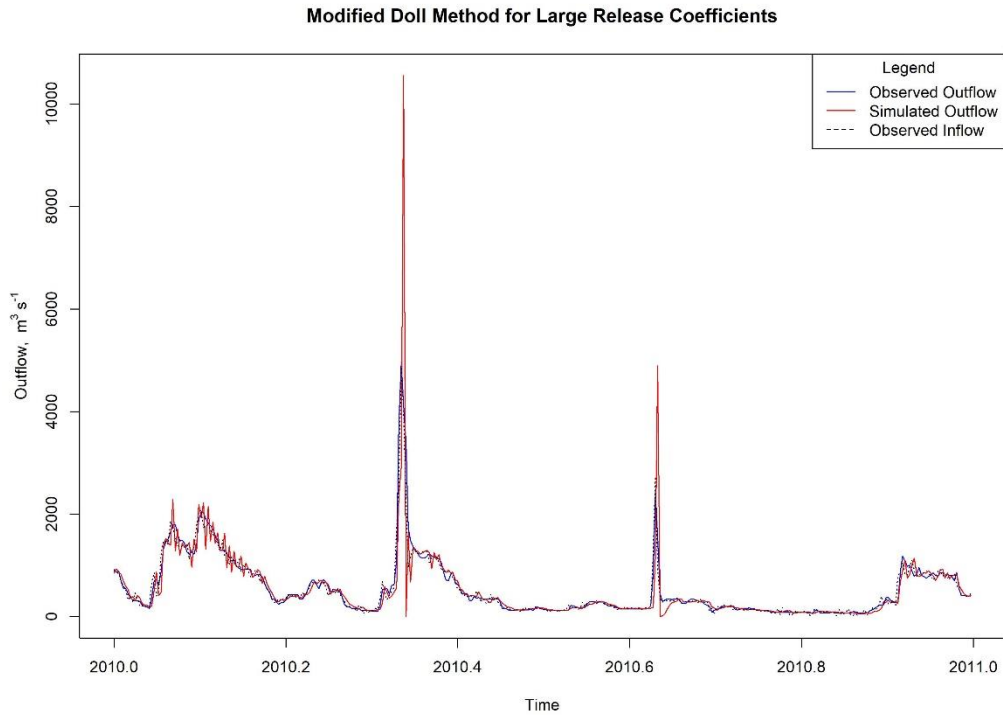
570 Several solutions are posited to address D03 instability. One solution could be to
571 varying k_{rd} values dynamically to mimic reservoir behavior. During large hydrologic
572 events the value of k_{rd} could reduce the peak of the outflow hydrograph, and then increase
573 during normal events. Another solution is the inclusion of rules and an expanded system
574 of equations that govern the solution. Because the intention of D03 is to approximate flow
575 at a free-flowing weir, coupling operational rules with the simulation may better
576 approximate reality. The rules may be as simple as switching behavior or the algorithm
577 when storage approaches either minimum or maximum reservoir storage. A simple
578 condition was tested for when storage drops below the minimum storage during the daily
579 time step:

$$580 \text{ if } S_t \leq S_{\min} \Rightarrow \begin{cases} S_t = S_{\min} \\ Q_{out} = Q_{in} + \frac{S_t - S_{\min}}{\Delta t} \end{cases} \quad (5)$$

581 This condition prevents the reservoir from falling below the minimum storage. Outflow
582 from Old Hickory Dam was re-simulated with $k_{rd} = 0.9$ and the new minimum storage
583 condition (Equation 5). The proposed modification resulted in simulated outflow shown in
584 Figure 11. Outflow is substantially overestimated for one-time step and drops to zero at the
585 next time step. While an oversimplification of actual operations, this condition is similar
586 to an emergency spillway discharge to prevent overtopping. The dam releases tremendous
587 flow for a brief period, when the maximum storage is nearly exceeded and then inhibits the
588 discharge when the storage is at the minimum capacity. The benefit of this modification is

589 that additional reservoir information is not required. However, further testing and
590 evaluation should be performed to validate this refinement.

591



592

593 Figure 11. Outflow simulation for the Old Hickory Dam using the proposed modification of the Doll method
594 for $k_{rd}=0.4$.

595 3.7. Limitations

596 The available sample of dams for this study has some inherent limitations. The vast
597 majority of reservoirs in the sample are primarily purposed as flood control reservoirs with
598 various secondary purposes. They are all commonly operated by USACE. And the dams
599 function within a predominately temperate climate across the United States. These
600 limitations preclude assertions regarding the effect the operating objective, dam ownership,
601 or country of operation on reservoir routing performance.

602 The abbreviated length of the historical records presents another limitation. The
603 evaluation period is limited to a six-year window which may not account for the total range
604 of operational environments for each dam. Thus, this evaluation likely does not capture
605 and evaluate D03 and H06 under absolute extreme circumstances.

606 All inflow utilized in this study is back calculated from observed changes in storage
607 and known discharges. This indirect method can lead to negative inflow values when losses
608 due to seepage, evapotranspiration, or other types of withdrawals are underestimated. De
609 Vos (2015) also noted that they used back-calculated inflow in their study. It is unclear
610 whether Hanasaki et al. (2006) made use of direct observations, but it is worth noting that
611 direct observations of total reservoir inflow are not readily available in most cases.

612 This study is limited to models that only require inputs related to reservoir inflow and
613 storage, primarily to provide insight into the reliability of these measures as indicators of
614 reservoir outflow. Because this study utilizes a back calculated reservoir inflow, inclusion
615 of reservoir withdrawal would also lead to an overestimation of water withdrawals from
616 the reservoir. Both D03 and H06 can account for withdrawals but because of the focus of
617 this study and the data utilized, the authors do not pursue an estimation of reservoir
618 withdrawal in this study. Thus, we have not included more sophisticated approaches,
619 such as Burek et al. (2013) or Zhao et al. (2016) within this study. Beyond this study of
620 sensitivity analysis, no formal calibration procedure was undertaken. A formal calibration
621 of k_{rd} in both D03 and H06 would be better suited for the insertion of the reservoir
622 routing scheme within a hydrologic routing scheme. This study is investigating the
623 feasibility of these methods in 0-10 day lead time, diurnal forecasting and is a precursor
624 to implementation in hydrologic routing schemes. There is limited benefit to standalone

625 calibration of the k_{rd} coefficients, given that reservoir outflow information is rarely
626 available at global scales. Operational calibration of k_{rd} would be challenging without
627 reservoir release records. Zajac et al. (2017) discuss the need for an open access database
628 of daily reservoir records, but no such database is known to be available at this time.
629 Thus, this study does not undertake any standalone, formal calibration of k_{rd} .

630 3.8. Future Work

631 D03 consistently improved simulated, daily streamflow estimates over naturalized
632 flow conditions in the selected reservoirs of this study, suggesting that D03 can potentially
633 improve global streamflow forecasting that do not already account for lakes and reservoirs.
634 D03 performed particularly well at daily time steps commensurate with many large-scale
635 stream routing models. The incorporation of D03 and H06 can be considered as modules
636 in large-scale river routing models such as Routing Application for Parallel computation
637 of Discharge (RAPID, David et al., 2011). The RAPID model is a river routing model that
638 can simultaneously compute streamflow in river networks with thousands of river reaches.
639 This will enable widespread testing and evaluation over large hydrologically diverse areas.

640 The research presented in this article should guide a number of follow-up
641 evaluations that will broaden the scope of this evaluation.

- 642 • We determined that k_{rd} can be varied to improve performance but have no
643 guidance on how to relate k_{rd} to a given reservoir. Future studies should
644 determine how to assign release coefficients to reservoirs.
- 645 • We have chosen parsimonious approaches that minimize assumptions. We
646 have not compared D03 or H06 to more complex models such as Burek et

647 al. (2013) or Zhao et al. (2016) which require these assumptions. Future
648 work will examine tradeoffs between model complexity and performance.

- 649 • Insertion of D03 into large-scale river routing models can facilitate studies
650 of how their results influence overall hydrologic performance, particularly
651 at locations downstream of reservoirs.
- 652 • Three quarters of the sampled dams have their primary purpose for flood
653 control. Efforts to fill the existing dataset with reservoirs that are primarily
654 irrigation, water supply, hydroelectric, recreation, and fish and wildlife
655 habitat and analyze the impacts of use on model performance should be
656 undertaken.
- 657 • The non-data-driven methods considered are conceptualizations of
658 reservoir operations that can be adapted to utilize remotely sensed
659 information, much like the data-driven methods previously mentioned.
660 Non-data-driven methods can be linked to statistical fitting techniques, but
661 they are capable of being employed independent of such pairings. However,
662 the non-data-driven reservoir routing schemes could be enhanced by
663 assimilating remotely sensed data, e.g. near real-time changes in storage
664 resolved from satellite altimetry, and eventually the planned NASA Surface
665 Water and Ocean Topography (SWOT) Mission. This information could
666 constrain reservoir simulations to improve global streamflow forecasts
667 (Yoon and Beighley, 2015).
- 668 • Because D03 skill tends to decline with increases in IR, an over-year
669 simulation capability similar to that proposed by De Vos (2015) may allow

670 for a better means of simulating diurnal reservoirs from reservoirs with large
671 IR. Over-year reservoirs have high IRs and yearly cycles of water storage
672 and release are not necessary (Adeloye and Montaseri, 2000; Vogel et al.,
673 1999).

674 4. Conclusions

675 This research compares two parsimonious reservoir routing methods (D03 and H06)
676 with the intent to determine if these methods can be effective at estimating diurnal reservoir
677 outflow in diurnal, medium-range streamflow forecasting. These methods were compared
678 across 60 USACE operated reservoirs at a daily time step. Results show that D03 tends to
679 outperform H06 at a daily time step. An in depth examination of these results yields the
680 following conclusions.

- 681 • The complexity and data requirements of both D03 and H06 are low and thus
682 computationally inexpensive. Both can be feasibly implemented at large spatial
683 scales at a daily or sub-daily time step.
- 684 • When the best performing k_{rd} is implemented within D03 we find a substantial
685 improvement in the model skill over the baseline for nearly all reservoirs at a
686 daily time step. H06 offers only a minimal improvement over the baseline when
687 the best k_{rd} is implemented for a daily time step. For the categories of KGE
688 specified (Tavakoly et al., 2017), the best performing D03 eliminates all poor
689 performing baseline conditions and increases the proportion of good or very
690 good performing sites by 22%.
- 691 • There is a statistical relationship between reservoir IR and two of the skill
692 metrics applied (KGE and R-Squared). Given that reservoirs with high IR

693 typically are less responsive to short-term fluctuations in inflow and storage,
694 the correlation between these variables is plausible. Further investigation of
695 dam characteristics, such as if the dams operate in series or in parallel and wet
696 and dry year considerations are further evidence of the correlation between the
697 IR and D03 and H06 skill.

698 • Simulation time step appears to be an important component in reservoir routing
699 skill. The comparison of the two methods by Hanasaki et al. (2006) are based
700 on monthly reservoir outflows and conclusions may not hold within diurnal
701 forecasting schemes. At overlapping locations, this study replicates the results
702 reported by Hanasaki et al. for monthly time steps. However, the Hamasaki et
703 al. findings do not hold for a daily time step evaluation.

704 • The best value for the empirical Döll coefficient, k_{rd} , can vary. Optimal values
705 were typically greater than the $k_{rd}=0.01$ value which Döll et al. (2003) derived.
706 This suggests that k_{rd} could be a potential calibration parameter within a large-
707 scale hydrologic modeling framework much like a weir coefficient, which is
708 specific to a particular type of weir.

709 • The Yazoo Basin Headwaters Project (Arkabutla Lake History, 2017; USACE,
710 1987) is an interesting case study in how reservoir system complexity can be
711 difficult to model. The Yazoo Basin Headwaters Project considers downstream
712 flow conditions as the dominant criteria in dam operation. Thus, the inflow and
713 available storage volume are poor predictors for determining reservoir
714 discharge in this type of management scheme. D03 appeared to scale flow

715 correctly at these reservoirs and improve reservoir overall skill, but timing of
716 the releases is not well represented and thus skill improvement is only minimal.

- 717 • Dam discharges in the Missouri River Reservoir System (Lund and Ferreira,
718 1996) are more correlated with storage volume and inflow conditions, which
719 lends itself to the two non-data-driven approaches evaluated here. D03 is
720 particularly capable of accurately modeling daily reservoir outflows in reservoir
721 systems that correlate well with storage and inflow fluctuations. Concerns
722 related to model error being compounded through a series dams may be
723 mitigated somewhat by the fact that inflow appears to be a progressively
724 stronger predictor of outflow further downstream in these types of systems.
- 725 • Numerical stability of D03 is a concern, particularly with higher k_{rd} values.
726 These stability concerns originate at reservoirs with small active storage
727 capacity during high inflow events. Additional model refinement can overcome
728 these stability concerns.
- 729 • D03 showed minimal bias during relatively wet and dry years. Timing of
730 releases can be influenced by wet years and the magnitude appears to be
731 affected during dry years. D03 appears to be most applicable for dam systems
732 where reservoir management focuses on upstream hydrologic conditions. Large
733 IRs could indicate reservoirs where downstream conditions are more likely to
734 influence release decisions at the reservoir.

735

5. Data Availability

736 All input, output, and evaluation data compiled for this study are available on request for
737 scientific purposes from Joseph Gutenson (jlgutenson@gmail.com) or Ahmad Tavakoly
738 (ahmad.a.tavakoly@erdc.dren.mil).

739

6. Author Contributions

740 JLG developed the research questions, analyzed the data, and compiled the paper. AAT
741 compiled the datasets, developed the research questions, took part in critical discussion of
742 the paper, and reviewed the paper. MDW provided input for the formulation of the
743 research questions, took part in critical discussion of the paper, and reviewed the paper.
744 MLF provided input for the formulation of the research questions, took part in critical
745 discussion of the paper, and reviewed the paper.

746

7. Competing Interests

747 The authors declare that they have no conflict of interest.

748

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