

Dear Editor and Reviewers:

My coauthors and I thank you for your thoughtful insight into how we can improve our manuscript, which is now entitled, “Comparison of Generalized Non-Data-Driven Lake and Reservoir Routing Models for Global-Scale Medium-Range Hydrologic Forecasting of Reservoir Outflow at Diurnal Time Steps”. Generally, our edits have improved the flow of ideas and clarified the discussion and insights gained from this analysis.

Specific improvements that have been made include:

1. The authors updated the manuscript to refer to the Döll Method and Hanasaki Method as D03 and H06, respectively.
2. The authors adapt the title and manuscript to better reflect the application of the paper to hydrologic forecast models at daily time steps.
3. The authors evaluated Masaki et al. (2017) to determine if their results at reservoirs along the Missouri River were comparable to those in this study. Because the study is more focused on intermodel comparison at seasonal time steps, there is little overlap with the intentions of our study and no comparison of the manuscripts deemed necessary by the authors.
4. We have verified that all inflow estimates in our reservoir sample are a back calculated inflow.
5. To better describe why a back calculated inflow was used in our study, Section 2.1 now describes why a back calculated inflow was chosen in this study. Section 3.7 describes the limitations of this study, based upon the use of a back calculated inflow.
6. To better describe our study’s objectives, clarification of why the D03 and H06 methods were chosen was provided in Section 1.2.
7. The manuscript was altered in Section 3.8 to better describe that non-data-driven methods can be linked to statistical fitting techniques and remote sensing data.
8. We investigate the reservoir routing methodology employed by Wada et al. (2014) but do not include this method because we deem it to be too simple and too similar to the Döll et al. (2003) approach. Section 1.2 describes this investigation in the manuscript.
9. In Section 1.2, we alter the manuscript to more clearly describe the rationale for comparing D03 and H06.
10. Units and dimensions were added to the descriptions of the equations in Section 2.1
11. Added the reference Macian-Sorribes and Pulido-Velazquez (2017) to the listed references.
12. A statement was added to Section 3.1 to explain why RMSE decreases and R-Squared and KGE increase.
13. Figure 6, Figure 7, Figure 8, Figure 9, and Figure 10 were altered to reference discharge as  $\text{m}^3 \text{s}^{-1}$ .
14. We have reviewed the document for spelling and grammatical errors.
15. A stacked proportional bar graph (Figure 4) and analysis were added to Section 3.1 to better describe the improvement that D03 provides over the baseline and H06 simulations.
16. We added verbiage to Section 2.2 make it clear in the manuscript that the reservoirs in this study are almost exclusively multipurpose and perform more than flood control.

17. An analysis of best performing  $k_{rd}$  in relation to IR was conducted and no significant statistical or visual relationship was found.
18. The authors found only one instance where model accuracy was substantially worse than the baseline condition. We consider this to be an outlier in our study because this reservoir behaves much differently than reservoirs of a similar IR and average inflow. We note this in Section 3.1 of the manuscript.
19. In Section 3.7, we added a discussion concerning the lack of diversity in reservoir operational purposes in our study's sample and how this inhibits the study's ability to determine the effect purpose has on reservoir routing performance.
20. Clarification was added to Figure 2, Figure 3, and Figure 5 to ensure that the description captured that these simulations depicted describe the best performing form of D03 and H06.

We look forward to your feedback on this version of the manuscript. Thank you again for your time and patience.

Best,  
Joseph Gutenson

# Comparison of ~~Outflow Estimation Using~~ Generalized Non-Data-Driven Lake and Reservoir Routing Models for Global-Scale ~~Medium-Range~~ Hydrologic ~~Modeling~~Forecasting of Reservoir Outflow at Diurnal Time Steps

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Abstract: Large-scale hydrologic ~~simulations-forecasts~~ should account for attenuation through lakes and reservoirs when flow regulation is present. ~~G~~Globally generalized methods for approximating outflow are required ~~since-but must contend with operational reservoir operation is~~ complexity and a dearth of information on dam characteristics at global spatial scales. ~~and specific real-time release information is typically unavailable at global scales.~~ There is currently no consensus on the best approach for approximating reservoir release rates in large spatial scale hydrologic forecasting, ~~particularly at diurnal time steps.~~ This research compares two parsimonious reservoir routing methods ~~at daily steps.~~ ~~The methods considered are those proposed by~~ Döll et al. (2003) and Hanasaki et al. (2006). ~~These reservoir routing methods have been~~ previously implemented in large-scale hydrologic modeling applications ~~and have typically-been typically evaluated seasonally.-requiring minimal data so as not to limit their usage. The methods considered are those proposed by~~ Döll et al. (2003) and Hanasaki et al. (2006). ~~This paper~~These routing methods are ~~compareds the two methodologies-~~across 60 reservoirs operated ~~from 2006-2012~~ by the U.S. Army Corps of Engineers. The authors vary empirical coefficients for both reservoir routing methods as part of a sensitivity analysis. The ~~Döll method~~method proposed by ~~-Döll et al. (2003)~~ outperformed ~~generally-outperformed the Hanasaki method~~that presented by Hanasaki et al. (2006) at a daily time step, ~~improving- and improved~~ model skill ~~over most in most cases beyond~~ run-of-the-river conditions. The temporal resolution of the model influences ~~models~~ performances. The optimal model coefficients varied across the reservoirs in this study and model performance fluctuates between wet years and dry years, and for different configurations such as dams in series. Overall, ~~the method proposed by Döll et al. (2003)~~ the Döll and Hanasaki Methods could enhance large scale hydrologic forecasting, but can be subject to instability under certain conditions.

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## 1. Introduction

### 38 1.1. Importance of Dams in Hydrologic Simulations

39 Improvements in numerical weather prediction, the increasing abundance of  
40 computational power, and greater precision of remotely sensed observations make global  
41 hydrologic forecasting and flood warning systems increasingly feasible (Alfieri et al.,  
42 2013; Wu et al., 2014; Emerton et al., 2016; Salas et al., 2017). Lack of information  
43 concerning anthropogenic influences on runoff is a major deficiency of large-scale flood  
44 forecasting systems (Emerton et al., 2016). Reservoir operations tend to distort natural flow  
45 patterns, effectively redistributing surface water spatially and temporally (Zhou et al.,  
46 2016). Impoundments significantly influence the downstream flow regime at small and  
47 large spatial scales (Batalla et al., 2004; Magilligan and Nislow, 2005). Over half of the  
48 world's large river systems are now substantially altered by dams (Nilsson et al., 2005)  
49 resulting in a seven-fold increase in water storage within the global river system  
50 (Vörösmarty et al. 1997). Furthermore, the cumulative alterations from global reservoir  
51 impoundments are so significant that it has been suggested that they could buffer global  
52 sea-level rise (Chao et al., 2008).

53 Dams primarily impact the hydrologic cycle by changing the magnitude and timing  
54 of the discharges downstream (Haddeland et al., 2006; Döll et al., 2009; Biemans et al.,  
55 2011; Wu et al., 2014; Zajac et al., 2017), often with the specific intent to mitigate  
56 hydrologic extremes (i.e., floods and droughts) (Zajac et al., 2017). Dams reduce peak  
57 discharges by roughly a third on average while dampening the daily variation by a similar  
58 amount (Graf, 2006). In hydrologic forecasting, accuracy of the timing and magnitude of  
59 hydrologic extremes is fundamentally important to the usefulness of the forecasts.

Therefore, the significant impacts from dams make inclusion of reservoir operations, or reservoir routing, critical in large scale hydrologic flood forecasting.

~~At continental scales, no current forecasting operations systematically account for dam and reservoir influences (Emerton et al., 2016).~~ Integrating dam operations within ~~large-scale hydrologic models~~ large-scale river routing and flood forecasting is shown to improve model performance downstream of reservoir locations (Snow et al., 2016; Tavakoly et al., 2017; Salas et al., 2017; Zajac et al., 2017). This is often not feasible at large-scales since there may be multiple entities responsible for regulating flow, particularly with respect to transboundary waters. Among other things, operational knowledge, site-specific rule curves, reservoir uses, and local decision-making practices at each individual project dictate dam releases. Thus, dam operations are typically non-linear, complex processes, driven by anthropogenic and environmental influences. This makes generalizing reservoir operations difficult, particularly in the context of predicting dam-induced hydrologic responses at diurnal or sub-diurnal time step. Heuristically accounting for dams within existing routing schemes should improve flood forecast results when scheduled releases are not readily known.

Reservoir routing methodologies are generally divided into ~~the~~ two basic categories: data-driven and non-data-driven. Machine-learning, artificial intelligence (Coerver et al., 2017; Macian-Sorribes and Pulido-Velazquez, 2017; Ehsani et al., 2016; Mohan and Ramsundram, 2016; Ticlavilca and McKee, 2011; Chaves and Chang, 2008; Khalil et al., 2005), and remote sensing (Bonnema et al., 2016; Yoon and Beighley, 2015) are examples of data-driven approaches. Such data-driven methodologies can be effectively applied to dynamic non-linear systems, particularly when the governing

influence on the system does not follow any particular deterministic model. These types of approaches require training data or specific knowledge of a particular reservoir to effectively parameterize and apply them. This is often an insurmountable limitation for data-driven approaches. For that reason, the focus of this paper is on non-data-driven reservoir routing methodologies as an incremental improvement over schemes that effectively neglect dams when information is scarce.

## 1.2. Non-Data-Driven Reservoir Storage and Outflow Simulation

Non-data-driven approaches to reservoir routing rely on conceptualizing reservoir responses without explicitly observing the actual reservoir operations. The optimal method for a given application depends on a balance between complexity and available information (De Vos, 2015). Therefore, this manuscript focuses on selecting for parsimony.

Existing non-data-driven reservoir models range from simple approaches to sophisticated methods. Solander et al. (2016) showed that temperature-based schema best fits the modeling of discharge,  $Q_{out,t}$ . The Solander et al. (2016) rule is driven by temperature shifts at each model time step above and below the mean temperature. The Solander et al. (2016) method indicates that temperature is the main proxy governing reservoir release, due to the assumption that seasonality drives agricultural production and reservoir operation. However, the Solander et al. (2016) study focuses on long-term climatic forecasting. Diurnal temperature variations will not likely describe day-to-day reservoir operations. Zhao et al., (2016) developed a reservoir routing scheme based on reservoir stage and storage rules. However, real-time insights related to current reservoir stages throughout a region can involve considerable remotely sensed information. The stage information must then be related somehow to storage volume making this a much

more a data-driven process. Burek et al. (2013) also developed a non-data-driven approach to reservoir routing which was implemented by Zajac et al. (2017). This approach is built into the LISFLOOD model. The Burek et al. (2013) model requires a number of assumptions about storage capacity limits and naturalized streamflow thresholds. For example, the minimum, normal, and maximum storage are assumed to be 0.1, 0.3, and 0.97, respectively. To maintain the objective of investigating parsimonious models, the approach by Burek et al. (2013) was not included in this evaluation.

Döll et al. (2003), [Wada et al. \(2014\)](#), and ~~and~~ [Wisser et al. \(2010\)](#) ~~were~~ presented non-data-driven methods to simulate reservoirs operation that can be considered as simple approaches.

The Wisser et al. (2010) method follows a simple, rule-based approach to define the reservoir outflow at each time step ( $Q_{out,t}$ ). The rule that Wisser et al. (2010) enacts is that when the inflow at each model time step moves above ~~or~~ [and](#) below [the](#) long-term average inflow, the behavior of the reservoir release changes. De Vos (2015) suggested that this model is too simple to effectively model reservoir outflow. [In a similar vein, Wada et al. \(2014\) introduced a daily estimate of reservoir outflow that is simply the product of the proportion of available reservoir storage and daily inflow, which we can be consider to be too simplistic to estimate reservoir outflow since as inflow no coefficient is introduced into the simulation to account for reservoir heterogeneity.](#)

Döll et al. (2003) derived ~~a natural lake~~ reservoir routing scheme. ~~Hence, this but this methodology that can be applied is applicable~~ to man-made reservoirs and natural water bodies. The Döll et al. (2003) methodology found genesis in the reservoir outflow model proposed by Meigh et al. (1999). Meigh et al. (1999) proposed a simple reservoir release

methodology, which intended to mimic outflow at reservoirs from a theoretical rectangular weir. A more substantive version of the Meigh et al. (1999) method is formulated by Döll et al. (2003). Despite its simplicity, the Döll [et al. \(2003\)](#) method demonstrated good performance compared to several other [routing methods](#) ~~previously mentioned~~ (De Vos, 2015). The form of the Döll et al. (2003) equation is similar to that proposed by Wada et al. (2014). However, the Döll et al. (2003) methodology incorporates a coefficient that can incorporate a portion of reservoir heterogeneity.

Compared to the aforementioned methods, Hanasaki et al. (2006) derived a demand driven approach to reservoir routing, which can be considered ~~as a~~ complicated non-data-driven reservoir routing model. They distinguished between irrigation and non-irrigation reservoirs and offered two distinct algorithms for each. Water demands for irrigation, domestic, and industrial uses are considered in the irrigation reservoirs, whereas the releases from non-irrigation reservoirs are simply a ~~ratio-proportion~~ of inflow.

De Vos (2015) also proposed a within-year/over-year reservoir routing method comprised of two systems of equations, which ~~they was~~ considered a non-data-driven approach. Within-year reservoir operations are driven by yearly fill and release cycles and typically have a small storage capacity relative to their total annual demand. Thus, water accumulates during wet periods and decreases during dry periods. Over-year reservoir operation, on the other hand, is based on long-term, multi-year drawdowns. Over-year reservoirs have storage which is sufficiently large, relative to inflow, so that yearly cycles of water storage and release are not necessary (Adeloye and Montaseri, 2000; Vogel et al., 1999). De Vos (2015) compared his methodology to the Hanasaki et al (2006), Döll et al. (2003), and Neitsch et al. (2011). The De Vos (2015) over-year simulation assumes



knowledge of the mean and standard deviation of reservoir storage and is still too data-driven for the purposes of this study.

~~The goal of this research is to evaluate reservoir routing schemes that are parsimonious and align with available information for use in diurnal hydrologic forecasting across a global domain. Considering these research aims, the non data driven reservoir routing methods developed by Döll et al. (2003) (referred to as D03) and Hanasaki et al. (2006) (referred to as H06), which will be referred to as Döll and Hanasaki methods, were considered in this research for several reasons.~~

~~The Döll et al. (2003) D03 and Hanasaki et al. (2006) H06 Both models require minimal input data to implement: reservoir inflow, average inflow, and storage volume characteristics, i.e. current, minimum, and maximum storage volume that can be estimated when detailed reservoir information is not available. Each of these variables are available in existing datasets, such as the Global Reservoir and Dam (GRanD) database (Lehner et al., 2011) or can be generated produced using climate reanalysis data (Snow et al., 2016). Other non-data-driven methods require data inputs that are not globally available or produced within the hydrologic simulation (De Vos, 2015; Zhao et al., 2016; Burek et al., 2013; Zajac et al., 2017). For example, the Global Flood Awareness System (GloFAS) is the only existing, operational flood forecasting system that accounts for reservoirs at continental to global spatial extents. However, the reservoir routing component of GloFAS requires operational assumptions be made because of a lack of global reservoir operational records (Zajac et al., 2017). D03 and H06 do not require that these assumptions be made because of the minimal inputs which they require. Thus,~~

D03 and H06 meet the requirements of being both parsimonious with respect to available reservoir information.

The Döll et al. (2003) and Hanasaki et al. (2006) methods D03 and H06 also provide enough complexity to account for a portion of the model complexity inherent in reservoir operations. De Vos (2015) does not employ the reservoir routing approach of Wisser et al. (2010) because De Vos (2015) contends that this method is overly simplistic. The approach taken by Wada et al. (2014) is similar to D03 but represents reservoirs with similar inflow and storage characteristics homogeneously.

Furthermore, ~~Additionally, both models~~ Döll et al. (2003) (hereafter referred to as D03) and Hanasaki et al. (2006) (hereafter referred to as H06) ~~D03 and H06 methods~~ have been implemented in large-scale hydrologic models. ~~The Döll method D03~~ was used in the WaterGAP model and the application of ~~the Hanasaki method H06~~ was implemented in the TRIP model by the same authors. The main difference in this evaluation and previous evaluations (i.e., Hanasaki et al., 2006; Masaki et al., 2017) of these reservoir routing schemes is that this research evaluates model performance at a diurnal time step.

The aim of this study is to assess non-data-driven reservoir routing methods that are parsimonious and align with available information for use in hydrologic forecasting schemes applicable across the global domain at diurnal time steps.– Considering these research aims, the non-data driven reservoir routing methods developed by Döll et al. (2003) and Hanasaki et al. (2006) were considered.

~~The Döll and Hanasaki methods were found to be sufficiently parsimonious for wide scale implementation.~~ The following research questions are addressed with respect to the ~~two chosen~~ D03 and H06 approaches: (1) How well do the selected ~~chosen~~ reservoir

routing models improve outflow estimates relative to simulation of naturalized flow (i.e. neglecting dams altogether)? (2) How do reservoir routing coefficients affect model performance? (3) How does the time step affect model performance and stability? This is a critical point for the current regional- to continental-scale forecasting schemes that operate at daily- or sub-daily- time steps. (4) How sensitive are the reservoir routing schemes to various real-world dam operations and climate variability?

To ~~achieve the research~~~~achieve research~~ objectives of the study, reservoir data including daily inflow and outflow from 2006-2012, for 60 [U.S. Army Corps of Engineers](#) (USACE) reservoirs were used to evaluate the reservoir routing schemes. The data were obtained from nine USACE districts: Pittsburg, Nashville, St. Paul, Rock Island, Omaha, Tulsa, Sacramento, Los Angeles, and Vicksburg. The selected dams are representative of a wide range of reservoir sizes, flow regimes, and climatologic settings but are predominately managed for flood control. The results of this analysis will benefit readers in determining if the reservoir routing models implemented within existing- large-scale hydrologic ~~models-forecasts~~ adequately represent reservoir effects.

## 2. Methodology

### 2.1. Simulation Specifications

The storage ratio (Vogel et al., 1999) or Impoundment Ratio (~~impoundment ratio~~) is an important metric in previous works examining generalizing reservoir operation ~~(by De Vos, (2015); and Hanasaki et al., (2006))~~. The impoundment ratio is described as follows:

$$IR = \frac{(S_{max} - S_{min})}{Q_{in} * 86400 * 365} \quad (1)$$

where  $S_{max}$  and  $S_{min}$  are the maximum and minimum volumes of the reservoir's active storage [m<sup>3</sup>], and  $Q_{in}$  is the mean annual inflow to the reservoir [m<sup>3</sup>s<sup>-1</sup>].

A higher impoundment ratio indicates that the capacity of the reservoir is large relative to mean inflows, while the opposite is true of low IR values. De Vos (2015) considered IR values greater than unity “large” reservoirs, as they are capable of storing the average yearly volume of water flowing into them. To utilize the Hanasaki method H06, the release coefficient ( $k_r$ ) needs to be determined.

$$k_r = \frac{S_{begin}}{\alpha S_{max}} \quad (2)$$

where  $S_{begin}$  is the storage [m<sup>3</sup>] at the beginning of each of the each year and  $\alpha$  is a dimensionless coefficient, which was set to 0.85 in the Hanasaki et al. (2006) study. In the current study, the  $\alpha$  parameter was varied from 0.45-0.95 by increments of 0.10 and solve  $k_r$  for each  $\alpha$  value.

Outflow is the quantity of most interest for hydrologic flood forecasting because these forecasts gthese simulations ggenerally occur over a relatively short 0-10 day lead time. The Hanasaki Method H06 relates outflow based on the incoming flow. In this study, only the non-irrigation methodology from the Hanasaki Method H06 was used to simulate reservoir outflow at each time step ( $Q_{out,t}$ ) since one cannot assume seasonal irrigation demands will be known globally. Further, the primary purpose of selected reservoirs selected in this study is not irrigation. Hanasaki the H06 method estimates outflow as follows:

$$Q_{out,t} = \begin{cases} k_r Q_{in,t} & (IR = 0.5) \\ \left(\frac{IR}{0.5}\right)^2 Q_{in,t} + Q_{in,t} \left\{1 - \left(\frac{IR}{0.5}\right)^2\right\} & (0 < IR < 0.5) \end{cases} \quad (3)$$

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where  $Q_{in,t}$  is the inflow [ $m^3s^{-1}$ ] at time  $t$  and  $k_r$  is the release coefficient which is calculated based on Equation 2. The 0.5 threshold value for IR is an empirical condition derived by Hanasaki et al. (2006).

Unlike Hanasaki method H06, the Döll method D03 relates outflow ( $Q_{out,t}$ ) to current available storage capacity of the reservoir:

$$Q_{out,t} = \frac{k_{rd}}{\Delta t} (S_t - S_{min}) \frac{(S_t - S_{min})^{1.5}}{(S_{max} - S_{min})} \quad (4)$$

Where Döll empirically derives the release coefficient,  $k_{rd} = 0.01$ ,  $\Delta t$  is the simulation time step (s), and  $S_t$  is the current volume of storage [ $m^3s^{-1}$ ] at time “ $t$ ”. For this study analysis of the Döll method D03ology,  $k_{rd}$  was varied using values of 0.01, 0.02, 0.04, 0.06, 0.08, 0.10, 0.20, 0.40, 0.50, 0.60, 0.70, 0.80, and 0.90 in this study. The results for the sensitivity analysis are discussed in the section 3.3.

The sensitivity analysis of  $k_r$  and  $k_{rd}$  can provide useful information on how coefficients may vary based on geographical and reservoir characteristics such as the impoundment ratio. The two methods were evaluated and results compared to actual outflow records provided by the USACE Districts. Two approaches were used to evaluate model performances: hydrograph assessment of daily and monthly reservoir outflow and statistical evaluation. The statistical evaluation was performed for daily and monthly averaged simulated results vs. observations using the Kling-Gupta efficiency (KGE, Gupta et al., 2009), coefficient of determination (R-Squared), and root mean square error (RMSE). The KGE value ranges from negative infinity to one. Four levels of performance were defined for KGE in this study (Tavakoly et al., 2017): poor performance ( $KGE < 0$ ), acceptable ( $0 < KGE < 0.4$ ), good ( $0.4 < KGE < 0.7$ ), and very good ( $0.7 < KGE$ ).

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Goodness-of-fit values were evaluated to compare simulated discharge to the actual outflow records provided by the USACE Districts. These are indicators of how well the models perform. The same goodness-of-fit values are calculated to compare actual discharge with ~~observed~~ inflow to assess baseline performance. The baseline condition represents the treatment of reservoir outflow as naturalized, altogether neglecting reservoir operations. Thus, the baseline condition is that inflow into the reservoir equals outflow from the reservoir. To be viable, the reservoir routing scheme should improve results over the baseline condition in virtually all cases.

A true directly measured ~~observed~~ daily inflow is not available for most nearly all reservoirs, including those maintained by the USACE. There are two ways that one can acquire a daily reservoir inflow; estimated using a streamflow model (as in Masaki et al., 2017; Zajac et al., 2017) or ~~use a~~ estimated using a back calculated inflow based on the known discharge ~~derived from observed reservoir outflow and observed changes in reservoir storage fluctuation~~ (as in De Vos, 2015). The authors have chosen to utilize a back calculated inflow because this methodology inherently accounts for all other withdraws from the reservoir, such as irrigation, ~~evapotranspiration, seepage, etc.~~ This allows the study to focus exclusively on the reservoir routing methodology. ~~utilized with no need to. This is also the reason that we do not the need to account for such withdraws.~~ In fact, that ~~in this study; as this would be double counting~~ withdrawals from the reservoir.

## 2.2. Study Area

The model ~~tests and~~ evaluations were conducted on 60 reservoirs in the United States maintained by the U.S. Army Corps of Engineers (USACE). Figure 1 illustrates

reservoirs used in this study. The primary purpose of 43 of the reservoirs are flood control, six are hydroelectric, four are recreation, three are water supply, two are classified as other, one is irrigation, and one is a fish and wildlife pond. Despite most reservoirs in the sample being primarily purposed as flood control reservoirs, only three of these reservoirs are exclusively purposed for flood control. Table 1 describes pertinent characteristics of each reservoir in this analysis.

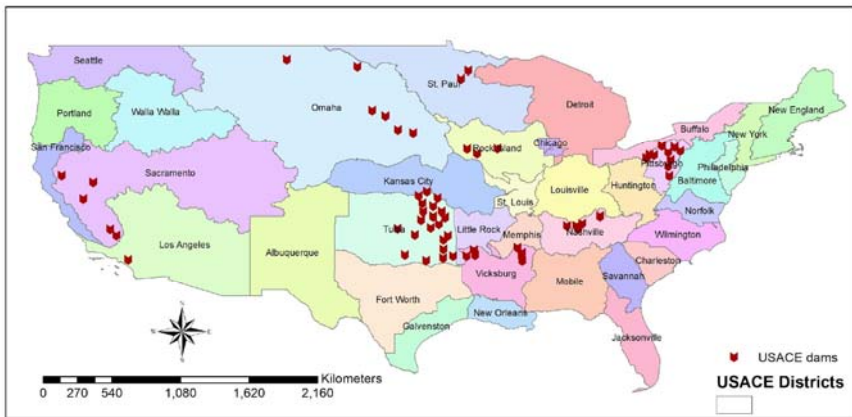


Figure 1. USACE districts and location of reservoirs in this study.

Table 1. Select statistical characteristics of reservoirs analyzed in this study.

Characteristic	Range	Mean	Standard Deviation
Minimum Storage ( $\text{m}^3 \cdot 10^6 \text{MCM}$ )	0 - 12.377	827	2,553
Maximum Storage ( $\text{m}^3 \cdot 10^6 \text{MCM}$ )	25 - 32,070	2,695	6,184
Annual Inflow ( $\text{m}^3 \cdot \text{s}^{-1} \text{cms}$ )	0.64 - 780	118	202
Annual Outflow ( $\text{m}^3 \cdot \text{s}^{-1} \text{cms}$ )	0.66 - 776	113	195
Impoundment Ratio	0.03 - 15.50	1.96	2.33

### 3. Results and Discussion

This section ~~describes~~ ~~reviews~~ the overall results of the study. There is significant improvement in skill over the baseline ~~run of the river condition~~ ~~(the use of inflow as an estimate of outflow)~~ when the optimal ~~best~~ D03 coefficient is chosen. ~~Because~~ D03 tends to outperform the baseline, ~~and~~ H06 generally mirrors the results of the baseline. ~~For this reason~~ ~~beyond the initial review of the results,~~ the discussion largely focuses on D03. The authors examine the distribution of best fitting  $k_{rd}$  values. We discuss how dam systems, annual variability ~~wet and dry years,~~ and simulation time step can influence the ~~cap~~ ability of D03 to estimate reservoir outflow. The authors also discuss the potential for numeric instability in D03 simulations and offer an initial solution to this instability. We also provide an overview of the limitations of this study and suggested future work.

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#### 3.1. Overall Model Performances

The goodness-of-fit metrics were calculated for each reservoir in the study. Observed inflow is compared with observed outflow to establish a benchmark used to show whether implementing the two non-data driven reservoir routing schemes improves estimates for reservoir outflow over ~~the use of unregulated flow as the reservoir outflow estimate~~ ~~simply treating as unregulated flow~~. Figure 2 illustrates the comparison of skill metrics between baseline ~~(the use of inflow as an estimate of outflow)~~ and the use of ~~D03~~ ~~the Döll and Hanasaki methods~~ H06 to simulate outflow. The KGE, R-Squared, and RMSE for ~~the Döll~~ ~~D03~~ and ~~Hanasaki methods~~ H06 in Figure 2 represent the best fit results from the sensitivity study. Data points in Figure 2 that fall below the dashed line represent



324 instances where KGE, R-Squared, and RMSE are lower for the reservoir routing method  
325 compared to the baseline. Data points falling above the dashed line indicate instances  
326 where higher KGE, R-Squared, and RMSE were obtained than the baseline for this study.  
327 ~~The Hanasaki MethodH06~~ tends to ~~show produce~~ minimal utility over the baseline  
328 scenario. In general, ~~the Hanasaki MethodH06~~ does not appear to make outflow estimates  
329 worse. Estimates that have acceptable KGE values in the baseline scenario tend to produce  
330 acceptable results using ~~the Hanasaki MethodH06~~. On the other hand, Figure 2 illustrates  
331 that ~~the Döll MethodD03~~ generally tends to increase KGE and R-Squared, and with this  
332 increase in goodness-of-fitaccuracy, decrease RMSE. Thus, the general conclusion is that  
333 selecting the optimum ~~Döll D03~~ release coefficient will ultimately produce an improved  
334 estimate of reservoir outflow compared to the baseline. Generally, ~~the Hanasaki~~  
335 ~~MethodH06~~ will produce an estimated reservoir outflow that performs similarly to the  
336 baseline scenario.

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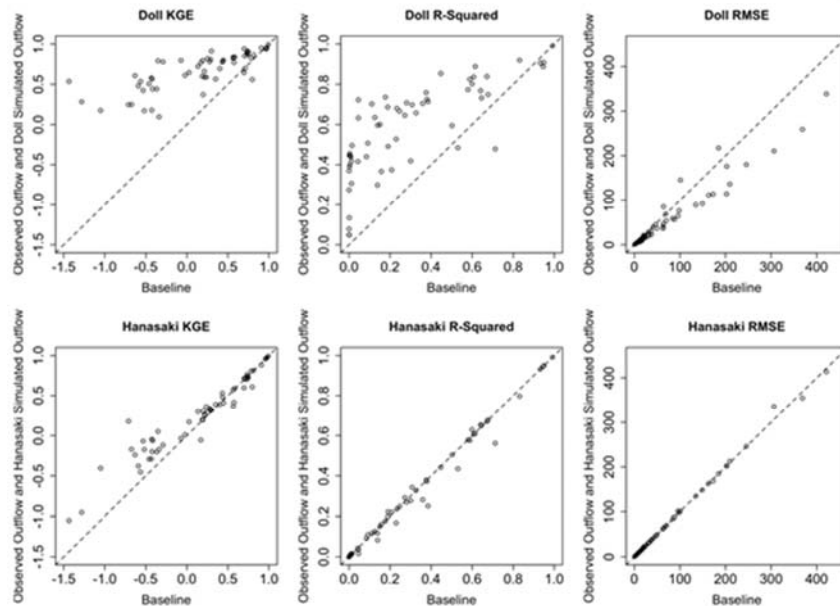


Figure 2. Scatter plots of skill metrics between the use of daily observed inflow as outflow (Baseline) and simulated outflow from best performing D03 and H06 simulations. The dashed line indicates the plane separating increased and decreased skill that results from using either reservoir routing method.

Figure 3 is a geographic representation of the KGE values from the baseline scenario as well as the best performing implementation of the two routing models for each reservoir. In general, ~~the Doll Method~~D03 outperforms the baseline and ~~Hanasaki Method~~H06, particularly in the Tulsa and Pittsburg Districts. H06 tends to provide, at best, minimal improvement in accuracy over the baseline.

~~Furthermore, the Doll Method~~D03 tends to improve KGE values at nearly all reservoirs and tends to preserve high KGE values at locations where the baseline is already a good or very good estimator of outflow. Only one of the 60 reservoirs in this study demonstrates a significant reduction in accuracy when D03 is applied. This reservoir, Martis Creek Dam in the Sacramento District, appears to be an outlier in the reservoir sample. Reservoirs with a similar IR and average inflow to Martis Creek Dam and in the

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same USACE district tended to experience improvement in model skill with D03. Overall, when the appropriate  $k_{rd}$  value is applied, D03 improves simulation results over the baseline.

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Figure 3a illustrates the wide range of reservoir operating conditions present in the study. The reservoir dataset contains reservoirs in which the outflow correlates poorly with the inflow regime as others that correlates well. Figure 3a also portrays significant geographic clustering where reservoirs in certain regions tend to be less correlated with inflow and other clusters where observed inflow and observed outflow correlate strongly. This could indicate that operations at these reservoirs may have a particularly regional context and may bias towards a particular reservoir routing scheme. However, ~~it can be seen that~~ correlation between observed inflow and observed outflow and geographic proximity of the reservoirs does not influence the implementation of either ~~the D04~~ D03 or Hanasaki method H06. Thus, the results of this research indicate no significant geographic constraints in the context of this study.

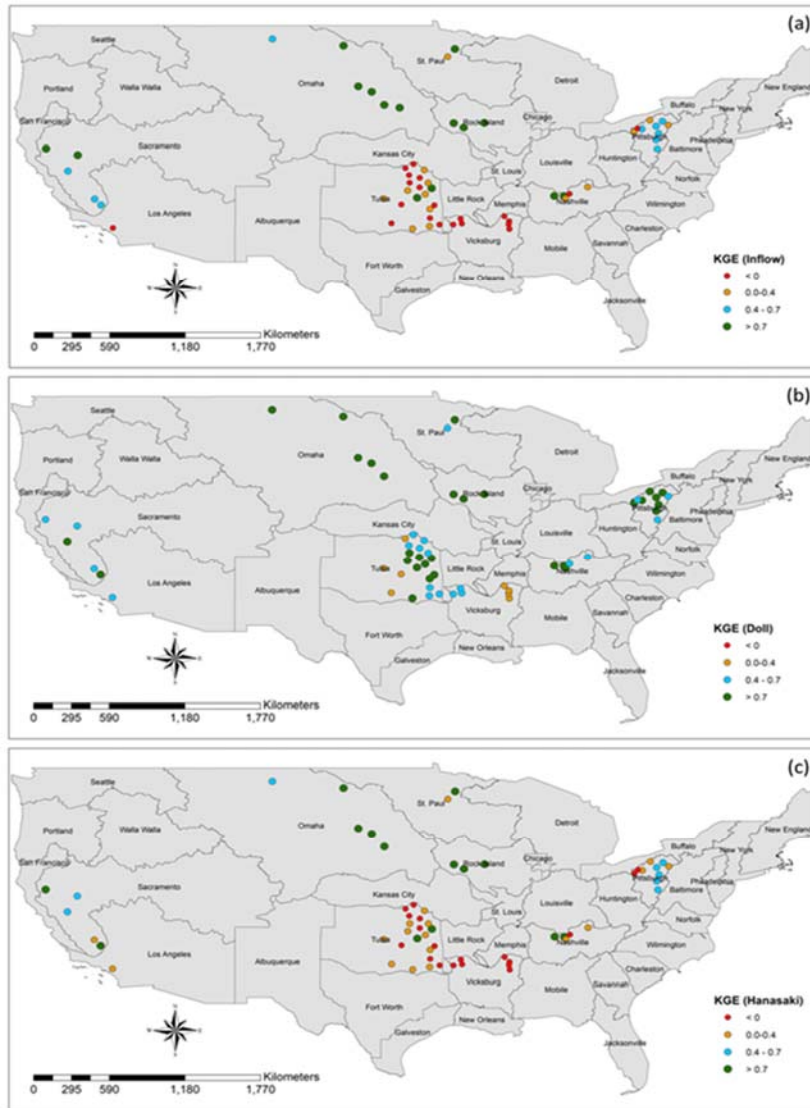


Figure 3. Spatial distribution of KGE comparing observed daily outflow to the each best estimate of outflow: a) observed inflow b) Döll Method simulated outflow, c) Hanasaki Method simulated outflow for all reservoirs in this study. KGE values for the Döll Method and the Hanasaki Method are the maximum KGE from all coefficient treatments.

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Figure 4 presents a proportional bar chart comparing baseline KGE and the highest KGE value for the range D03 and H06 coefficients. This plot categorizes KGE performance using the same bins as Figure 3. Figure 4 indicates that the best performing H06 simulation provides only marginal improvement over the baseline condition. However, the best performing instance of D03 eliminates all poor performing baseline conditions. Nearly 87% of all best performing D03 simulations are considered to be good or very good at accurately capturing reservoir outflows, a 22% increase above the baseline simulation.

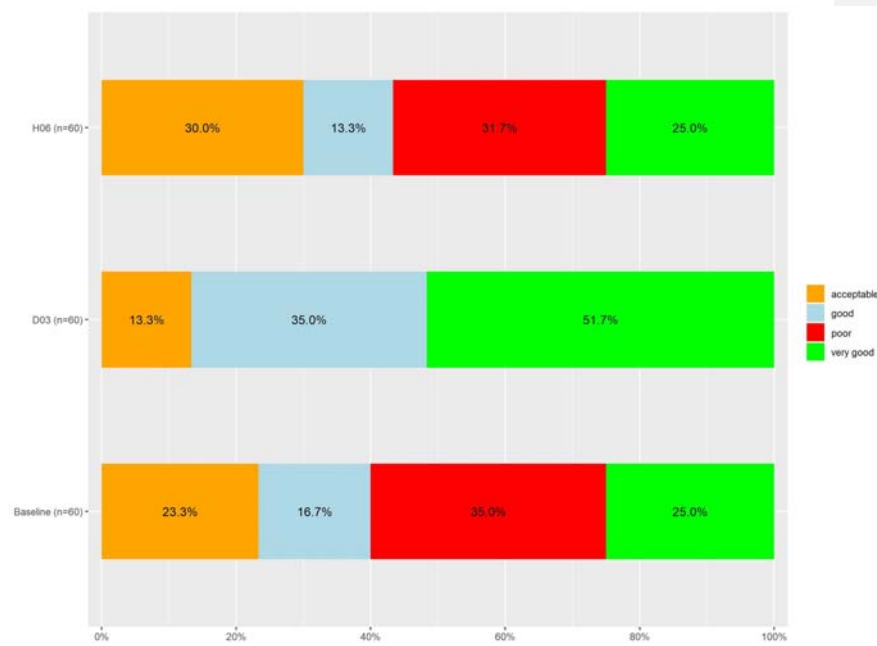


Figure 4. Proportional bar chart comparing the baseline outflow estimation and the best KGE results for D03 and H06.

From multivariate comparison, a ~~substantial~~ negative relationship between two of the best fit results (KGE and R-Squared) and reservoir IR was found. Figure 54 illustrates this comparison between IR and each goodness of fit metric for the baseline, ~~D03~~, and

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Hanasaki method H06. Based upon Figure 54, KGE in particular appears to non-linearly correlated to IR. A similar, yet less significant, negative relationship was found between IR and R-Squared. Little statistical correlation appears to occur between IR and RMSE. However, KGE and R-Squared values in Figure 54 indicate that the ability to predict outflow using the reservoir routing techniques applied in this study decreases with reservoir with high IR values. Proceeding sections investigate some of the possible reasons for this relationship between reservoir routing model performance and IR.

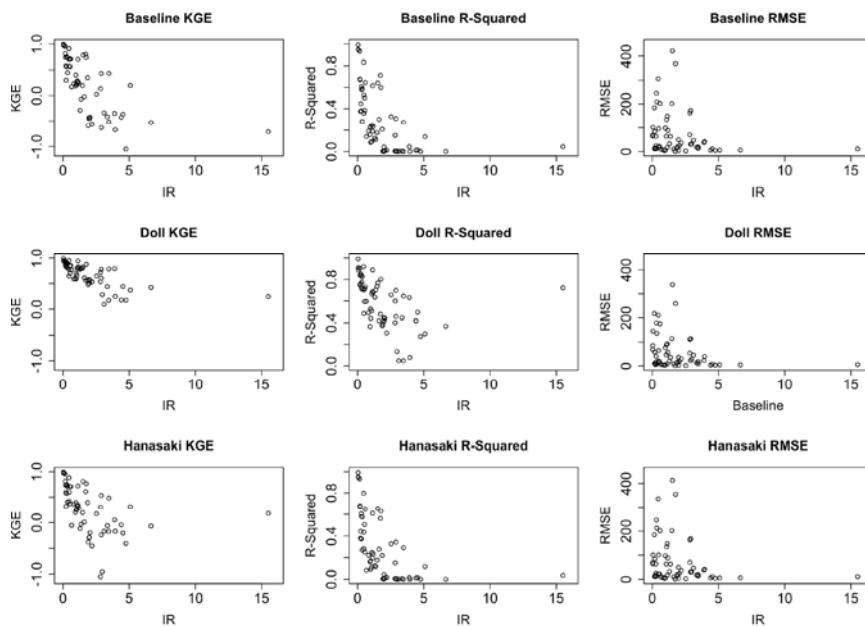


Figure 54. Comparison of IR and best KGE, R-Squared, and RMSE from goodness of fit metrics for baseline, D03, and H06.

### 3.2. Sensitivity Analysis of Models

Because the Doll method D03 consistently outperforms the Hanasaki method H06 at daily time steps, the Doll Method D03 was selected for the sensitivity analysis at daily

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time steps. The value of  $k_{rd}$  coefficient was introduced as 0.01 in the Döll et al. (2003) study. In this study,  $k_{rd}$  values were varied to obtain maximum KGE and R-Squared and minimum RMSE. Figure 65 demonstrates the dispersion of  $k_{rd}$  values which maximize the model skill to simulate reservoir routing for all selected reservoirs in this study. For all model skill metrics,  $k_{rd}=0.90$  tends to be the most prevalent  $k_{rd}$  value that maximizes model skill. In only two of the 60 reservoirs (Sardis Dam and Enid Dam)  $k_{rd} = 0.01$  maximizes R-Squared and minimizes RMSE for the range of  $k_{rd}$  coefficients. This research suggests that the  $k_{rd} = 0.01$  is not necessarily the optimum coefficient to maximize model performance using a daily simulation time step.

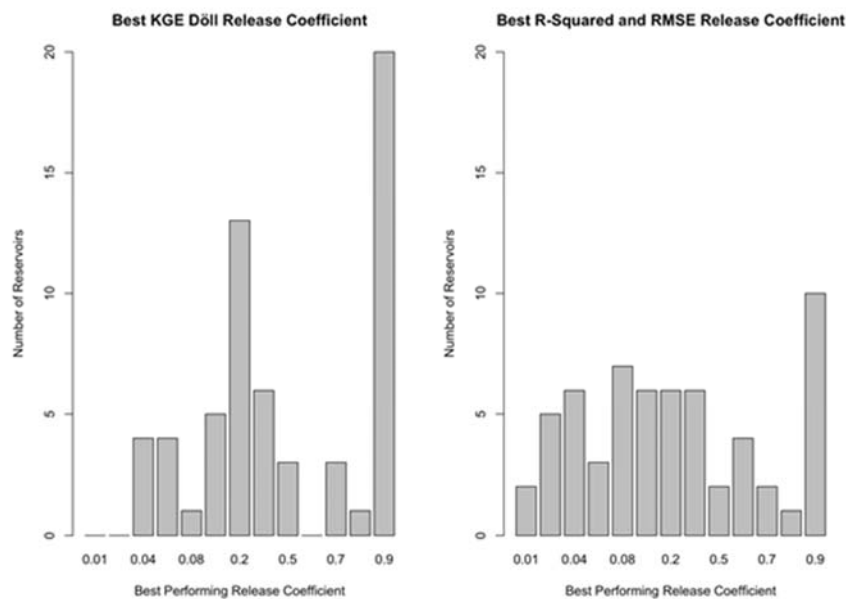


Figure 56. Bar charts of  $k_{rd}$  values that maximize KGE and correlation and minimize RMSE.

Investigating the linkage between dam characteristics and the best performing  $k_{rd}$  yields no clear relationship. Evaluation of correlation between ~~IR~~~~impoundment ratio~~, coefficient of variation of inflow, ratio of average inflow to average outflow, and geographic location shows low correlation between each variable and best performing  $k_{rd}$  value. However, the range of best performing  $k_{rd}$  within this analysis and as demonstrated in Figure 56 suggests that the value is not constant across all reservoirs. Thus, as one implements ~~the Döll Method D03~~ within their hydrologic ~~forecasting~~~~modeling~~ framework,  $k_{rd}$  may be adjusted ~~to optimize by~~~~when comparing~~ streamflow estimates to gage observations, like those curated by the Global Runoff Data Centre (GRDC, 2017), ~~when~~ [available](#).

### 3.3. Dam Systems and Reservoir Routing

Reservoirs in the Vicksburg and Omaha districts were selected to evaluate performance of ~~the Döll Method D03~~ ~~in environments where a complex drainage systems~~~~reservoirs operate in a coordinated fashion.~~ We broadly refer to these as dam systems. The case of the Vicksburg and Omaha district reservoirs highlights two distinct types of dam systems; one where the dams do not contribute inflow into one another but still coordinate their releases (in parallel) and another where upstream releases flow into downstream reservoirs (in series).

A subset of the reservoirs in the Vicksburg District comprise the Yazoo Basin Headwaters Project. Although ~~these the~~ reservoirs in the Yazoo Basin Headwaters Project are not directly connected, the reservoir operators coordinate operations in order to minimize flooding in ~~the Louisiana Delta regions near the mouth of the Mississippi River~~Mississippi's Delta region (USACE, 2017; USACE, 1987). The operation of these



reservoirs presents an interesting case in which the non-date driven models in this study do not characterize the nature of the dam releases well. The modeled results at four Vicksburg District dams yield only minimal improvement over unregulated (i.e. naturalized) flow at these reservoirs. The decrease in reservoir routing performance can be attributed to the large impoundment ratios at these dams indicating the reservoir storage is large relative to annual volume of inflow.

The reservoirs of interest in the Vicksburg District include Arkabutla, Sardis, Enid, and Grenada. These dams function in parallel on tributaries of the lower Mississippi River, namely the Coldwater River, Little Tallahatchie River, Yocona River, and Yalobusha River, respectively. Together, these dams control flooding in northern Mississippi as part of the Yazoo Basin Headwaters Project (USACE, 2017; USACE, 1987). The Yazoo Basin reservoirs discharge directly into the heavily regulated Mississippi River (Meade and Moody, 2010). The reservoirs operate to ensure high releases are not concurrent with large flows upstream on the Mississippi to avoid devastating flooding to the low-lying Louisiana delta regions. This requires a high level of coordination throughout the Yazoo Basin Headwater Project and with regulation upstream on the Mississippi. Additionally, each of the Yazoo Basin reservoirs have a substantial impoundment ratio, ranging from 2.96-3.95. In other words, the reservoirs are capable of containing large volumes of water to mitigate downstream impacts. Thus, current pool levels and forecasted inflow at these four reservoirs do not substantially influence release decisions. The reservoirs also have the capacity to absorb large flood events. As a result, they do not seem to follow the same functional form as the majority of other dams in this study.

Figure 76 from Sardis Dam in the Yazoo Basin Headwaters Project demonstrates the hydrograph comparing observed inflow and outflow and the modeled outflow that provides the highest KGE (Döll method D03,  $k_{rd}=0.90$ ) for the year 2008. Figure 76 demonstrates that peak outflows do not tend to correspond to the time at which peak inflow occurs. In fact, release rates at Sardis Dam are at a minimum during the peak inflow time period. This pattern repeats at each of the reservoirs in the Yazoo Basin Headwaters Project indicating that inflow and consumed storage are not substantial predictors of outflow timing at these reservoirs. This exemplifies the lack of correlation between observed inflow and observed outflow at reservoirs within the Yazoo Basin Headwaters Project.

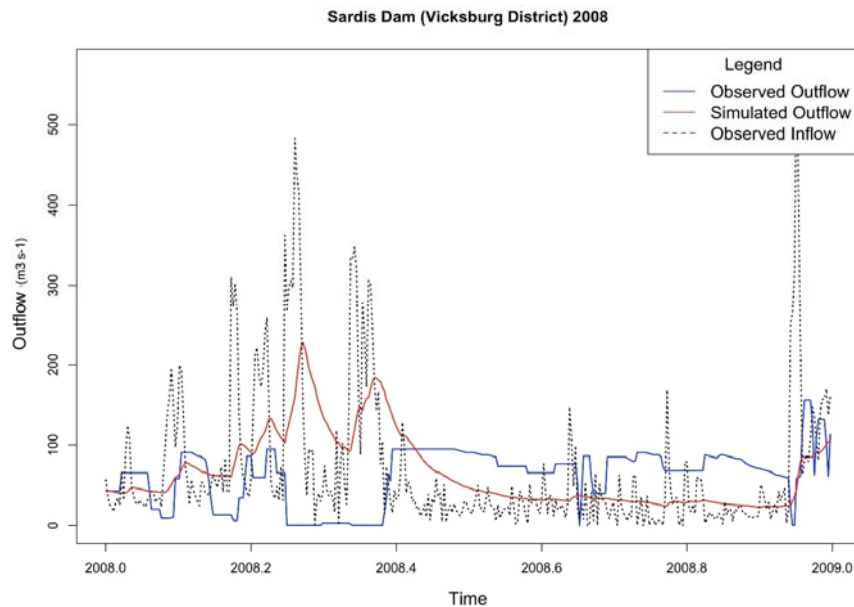


Figure 76. Hydrographs of observed inflow and outflow versus simulated outflow with the highest KGE value at Sardis Dam (Döll method  $k_{rd}=0.90$ ). KGE comparing observed Inflow and outflow = - 0.34; KGE comparing simulated and observed outflows= 0.095

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468  
469 Dams operating in series represent a specific case where compounding model error  
470 is a particular concern. USACE operates several large dams in series on the Missouri River.  
471 These include Fort Peck, Garrison, Oahe, Big Bend, Fort Randall, and Gavins Point within  
472 in the Omaha District (Lund and Ferreira, 1996). For this cascading system on the Missouri  
473 River, inflow appears to be a progressively stronger predictor of outflow from upstream to  
474 downstream. At the upstream end ~~inflow-the baseline~~ yielded a KGE=0.43 at Fork Peck  
475 with a KGE=0.99 downstream at Gavins Point Dam. Figure 87 provides a comparison of  
476 observed inflow and outflow along with simulated outflow for Gavins Point Dam. ~~The Döll~~  
477 ~~methodD03~~ tends to provide a slightly better estimate of outflow compared with inflow,  
478 except in the instance of Big Bend Dam. At Big Bend Dam, ~~the Hanasaki-methodH06~~  
479 produces an estimate of outflow more consistent with observed outflow than either ~~the Döll~~  
480 ~~methodD03~~ or inflow alone. However, the differences are almost trivial considering how  
481 well inflow alone performed in this case. ~~The Döll-methodD03~~ is particularly accurate  
482 during peak inflow conditions, for example the large hydrologic event in mid-2011 at  
483 Gavins Point Dam in Figure 87. The performance of non-data driven approaches in this  
484 instance is promising since compounding errors are a large concern in this type of system.  
485 Other instances involving dams in series should be evaluated to determine ~~find~~ out if these  
486 findings hold more generally.

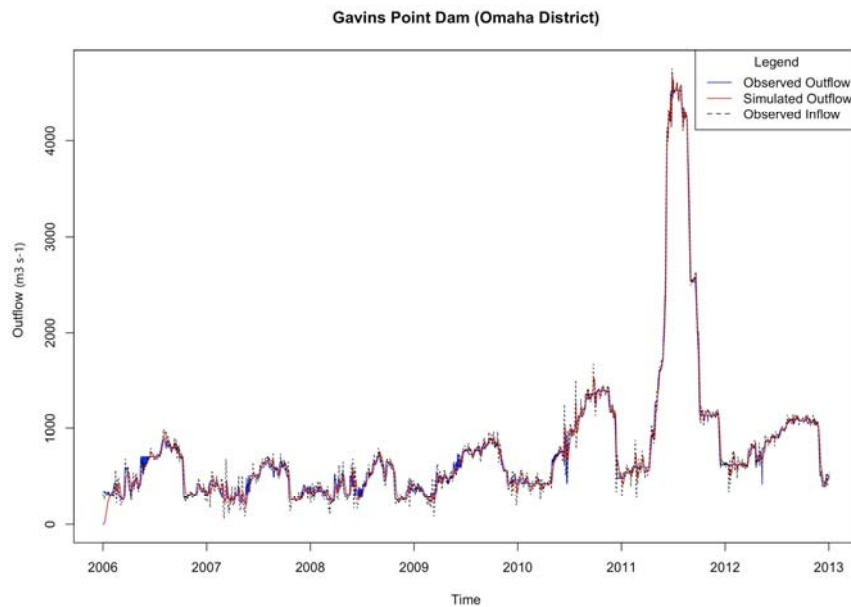


Figure 8. Hydrographs of observed inflow and outflow versus simulated outflow with the highest KGE value at Gavins Point Dam (Döll method  $k_{rd}=0.04$ ). KGE comparing observed Inflow and outflow = 0.99; KGE comparing simulated and observed outflows= 0.99.

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Figure SEQ Figure \\* ARABIC 7. Hydrographs of observed inflow and outflow versus simulated outflow with the highest KGE value at Gavins Point Dam (Döll method  $k_r=0.04$ ). KGE comparing observed Inflow and outflow = 0.99; KGE comparing simulated and observed outflows= 0.99.

491  
492 The reservoir management is unique in both the Yazoo Basin Headwaters Project  
493 and the Missouri River. The operators of dams within the Yazoo Basin Headwaters Project  
494 tend to regulate outflow in a manner that is more in line with downstream conditions. The  
495 attention to downstream conditions is due mainly to the impact that downstream floods will  
496 have on the low-lying communities within the Louisiana Delta. The dams in the Yazoo  
497 Basin Headwaters Project have among the highest impoundment ratios, which inherently  
498 reduces the influence of upstream conditions in discharge decisions. The non-data driven  
499 approaches evaluated here do not account for downstream conditions and thus do not  
500 perform well in this instance, particularly where large impoundment ratios allow operators  
501 considerable leeway.

On the other hand, the non-data driven approaches tend to perform well when inflow conditions dictate discharge decisions as we see on the Missouri River system. Reservoirs with smaller impoundment ratios are naturally more responsive to inflow requiring greater consideration for upstream conditions. ~~The Döll Method~~D03 showed relatively small improvement of outflow estimates compared to inflow as a predictor of outflow in the Yazoo Basin Reservoirs, while the method provided reasonable estimates in dam systems like the Missouri River system. Therefore, it can be inferred that ~~the Döll method~~D03 is more applicable for dam systems where reservoir management focuses on upstream hydrologic conditions, while large impoundment ratios may be indicative of reservoirs where downstream conditions are more likely to prevail. This would likely apply for ~~the Hanasaki Method~~H06 as well since that method links outflow to inflow more directly.

#### 3.4. Wet and Dry Year Comparison

Figure 8 shows results for wet and dry years at two reservoirs considered to be representative of this study. ~~The Döll Method~~D03 provides a relatively good estimate of outflow at Union City Dam (Pittsburg District) in Figure 98a and Figure 98c. D03 ~~It~~ performs relatively poorly at Arcadia Lake (Tulsa District) in Figure 98b and Figure 98d. In the case of Union City Dam, ~~the Döll Method~~D03 tends to produce a noticeable improvement in model skill during both a relatively wet year and a relatively dry year. The performance (Figure 98a and Figure 98c) seems to be independent of wet or dry conditions, at least on an annual basis. This does not hold for Arcadia Lake. The model shows modest skill at Arcadia Lake during the wet year (Figure 98b), but almost none during the dry year.

There appears to be a difference in the timing discharges between at the two locations in Figure 98. ~~The Döll MethodD03~~ appears to estimate the right amount of volume released during the wet year at Arcadia Lake (Figure 98b). However, the timing of the observed actual release is delayed until a relatively dry period begins~~from the estimate given by the model~~. The lag could indicate that water is being retained, possibly for use in irrigation or domestic supply. In this instance, Arcadia Lake supplies water to the city of Edmond, Oklahoma which may influence release decisions (Arcadia Lake Park Office, 2018).

~~The Döll MethodD03~~ performs much more poorly during the 2006 dry year at Arcadia Lake (Figure 98d). The model does not predict the sporadic releases throughout the year. The inflow events in that year are not substantial enough to affect storage meaningfully, thus we see almost no response in the modeled output. Observed outflows demonstrate that beyond two relatively high-volume reservoir releases during 2006, the reservoir releases are restricted to practically no outflow the rest of the year. ~~The Döll MethodD03~~ does not anticipate the two large releases, as the reservoir storage does not dramatically shift in either instance. ~~D03Arcadia Dam appears to be operating in a conservation mode for nearly the entire year. The Döll MethodD03 does not account for this. Instead, it~~ estimates a near constant discharge over the entire year with almost no storage change.

Results for wet years and dry years appear to be fairly mixed. Indications are that the performance of ~~the Döll MethodD03~~ could be somewhat site specific. However, reservoirs that tend to be less responsive to storage fluctuations are not represented well in ~~the Döll MethodD03~~ since storage fluctuations drive the model. Arcadia Lake has an IR of

about 4.75 which is relatively high. Union City Dam has an IR of about 0.24, which is relatively low. IR is a good indicator of reservoir responsiveness to storage fluctuations. A lack of reservoir responsiveness to storage fluctuations could result in two different types of error when ~~the DöH Method~~D03 is implemented within a large-spatial-scale hydrologic model. First, forecasted outflow could easily mistime a hydrologic event, particularly during wet years, as Figure ~~89~~b demonstrates. Second, the authors anticipate that if the storage does not dramatically fluctuate during a dry year the estimated reservoir release ~~likely~~ will not anticipate sporadic releases for irrigation and other purposeful discharges. Unaccounted for, these large but short duration releases may lead to a consistent overestimation of reservoir outflow for the entire dry year period.



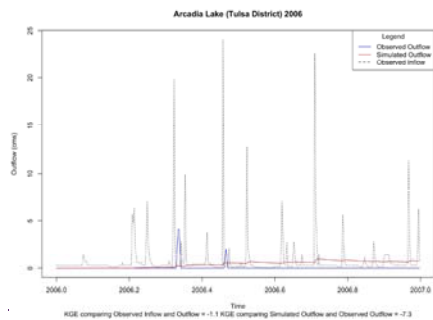
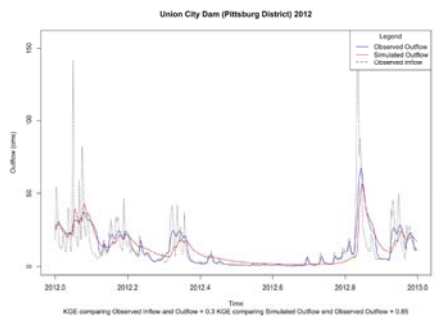
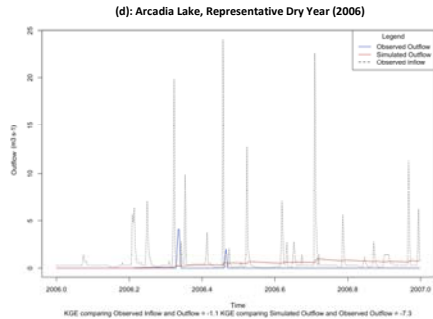
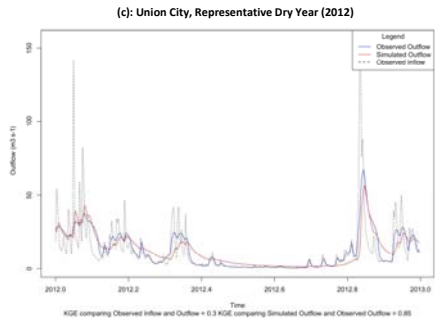
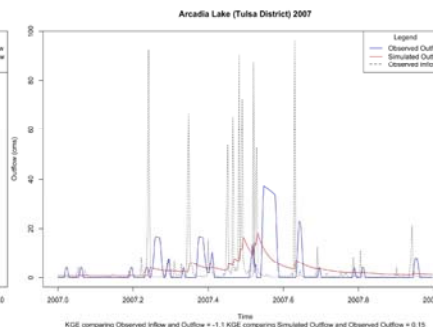
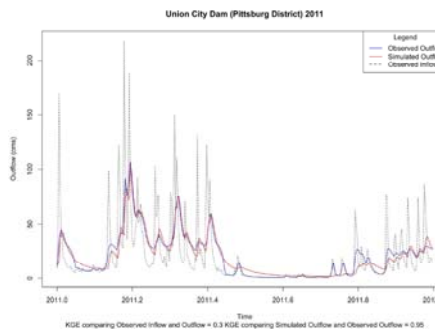
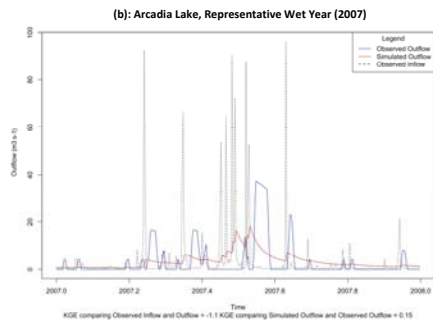
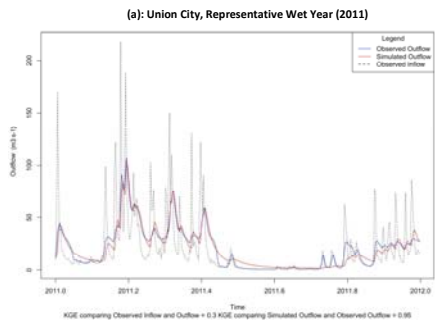


Figure 98. Two reservoirs where ~~the Döll Method D03~~ tends to perform very good and poor: outflow: a) wet year Union City Dam 2011; b) wet year Arcadia Lake 2007; c) dry year Union City Dam 2012; and d) dry year Arcadia Lake 2006.

### 3.5. Effects of Time Step on Model Performance

Model comparisons are conducted for daily and monthly time steps. Table 2 illustrates the results at Fort Peck, Garrison Dam, Oahe Dam, and Fort Randall Dam, each of which appears in the Hanasaki et al. (2006) study and this research. Table 2 also contains Sardis Dam, Mosquito Creek Dam, and Prado Dam, which are not included in Hanasaki et al. (2006). Results illustrate that the time scale at which comparisons are conducted can influence simulation results. The monthly comparison amongst Fort Peck, Garrison, Oahe, and Fort Randall is in agreement with the conclusions of Hanasaki et al. (2006). However, when the simulation time step changes to a daily time step, the skill of ~~Hanasaki Method H06~~ and ~~the Döll method D03~~ reverse and ~~the Döll method D03~~ tends to outperform ~~the Hanasaki Method H06~~. In additional reservoirs (Sardis and Prado), the results indicate that ~~the Döll method D03~~ outperformed ~~the Hanasaki Method H06~~ at both daily and monthly time steps, based upon KGE. However, the results at Mosquito Creek reservoir tend to follow the original Hanasaki et al. (2006) results.

The time-scale effect upon model performance may relate to how well observed inflow correlates with observed outflow. Examining Table 2, ~~Hanasaki Method H06~~ outperforms ~~the Döll Method D03~~ when observed inflow and observed outflow are relatively well correlated. The effect is nullified when the inverse is true. ~~The Hanasaki Method H06~~ estimates outflow as a ratio of inflow, which may be a better estimate of outflow at the monthly time scale, particularly when discharge tracks closely with inflow. However, ~~the Hanasaki Method H06~~ will fluctuate at the smaller time steps due to inherent

variations in inflow. ~~The Döll MethodD03~~ tends to vary less at a daily time step and may be a better estimate of outflow at sub-monthly time steps.

The hydrographs from Fort Randall Dam further illustrate the relationships between time step and model skill, particularly during high flow events. Daily and monthly comparisons between observation and simulations for Fort Randall Dam are shown in Figure 910. ~~Figure 10~~ This figure compares the daily and monthly simulations with observations. Figure 910a shows that the ~~H06 Hanasaki~~ simulations perform better than ~~the Döll MethodD03~~ for monthly time steps, particularly during the high inflow periods events in 2011. ~~The Döll methodD03~~ tends to overestimate reservoir outflow, while ~~the Hanasaki MethodH06~~ correlates well with inflow and better matches the peak flow of 2011. At a diurnal time step (Figure 109b), ~~the Hanasaki MethodH06~~ tends to be hypersensitive to inflow variations and overestimates outflow, whereas ~~the Döll methodD03~~ provides a better approximation of outflow during the 2011 high flow event at a daily time step.

Table 1. Comparison of daily and monthly KGE values at selected reservoirs. The  $\alpha$  and  $k_{rd}$  values represent the highest KGE values for Hanasaki and Döll methods respectively.

Reservoir	Daily KGE			Monthly KGE		
	Inflow	Hanasaki	Döll	Inflow	Hanasaki	Döll
Fort Peck $\alpha=0.95$ $k_{rd}=0.04$	0.43	0.53	0.78	0.54	0.62	0.51
Garrison Dam $\alpha=0.95$ $k_{rd}=0.06$	0.73	0.76	0.88	0.78	0.80	0.59
Oahe Dam $\alpha=0.95$ $k_{rd}=0.20$	0.78	0.81	0.83	0.84	0.86	0.76
Fort Randall Dam $\alpha=0.95$ $k_{rd}=0.20$	0.91	0.88	0.95	0.96	0.93	0.67
Sardis Dam $\alpha=0.95$ $k_{rd}=0.90$	-0.34	-0.17	0.09	0.06	-0.03	0.16
Mosquito Creek Dam $\alpha=0.45$ $k_{rd}=0.70$	-0.46	-0.29	0.51	0.49	0.60	0.39

<b>Prado Dam</b> $\alpha=0.95$ $k_{rd}=0.50$	<u>-0.02</u>	<u>0.01</u>	<u>0.61</u>	<u>0.32</u>	<u>0.61</u>	<u>0.71</u>
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Table 2. Comparison of daily and monthly KGE values at selected reservoirs. The  $\alpha$  and  $k_{rd}$  values represent the highest KGE values for Hanasaki H06 and Döll D03 methods respectively.

Reservoir	Daily KGE			Monthly KGE		
	Inflow	Hanasaki	Döll	Inflow	Hanasaki	Döll
Fort Peck $\alpha=0.95$ $k_{rd}=0.04$	0.43	0.53	0.78	0.54	0.62	0.51
Garrison Dam $\alpha=0.95$ $k_{rd}=0.06$	0.73	0.76	0.88	0.78	0.80	0.59
Oahe Dam $\alpha=0.95$ $k_{rd}=0.20$	0.78	0.81	0.83	0.84	0.86	0.76
Fort Randall Dam $\alpha=0.95$ $k_{rd}=0.20$	0.91	0.88	0.95	0.96	0.93	0.67
Sardis Dam $\alpha=0.95$ $k_{rd}=0.90$	-0.34	-0.17	0.09	0.06	-0.03	0.16
Mosquito Creek Dam $\alpha=0.45$ $k_{rd}=0.70$	-0.46	-0.29	0.51	0.49	0.60	0.39
Prado Dam $\alpha=0.95$ $k_{rd}=0.50$	-0.02	0.01	0.61	0.32	0.61	0.71

It is possible that the conclusions of Hanasaki et al. (2006) suggesting better performance of the Hanasaki Method H06 at the monthly-scale depend on how closely discharge from the dam tracks inflow. The Döll method D03 may be a better candidate for integration into daily flow forecasting models.

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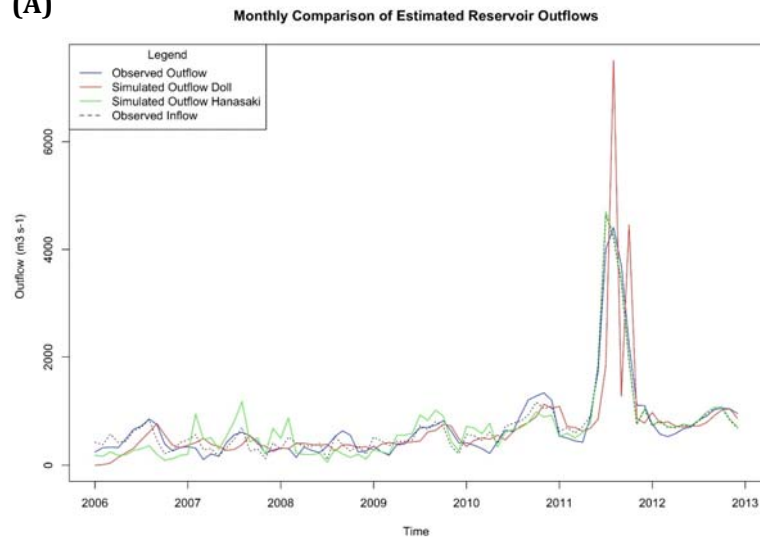
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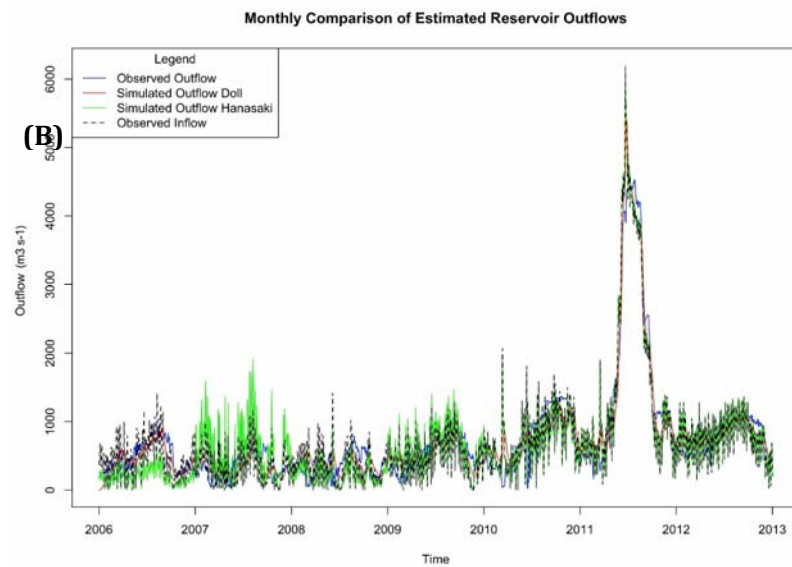
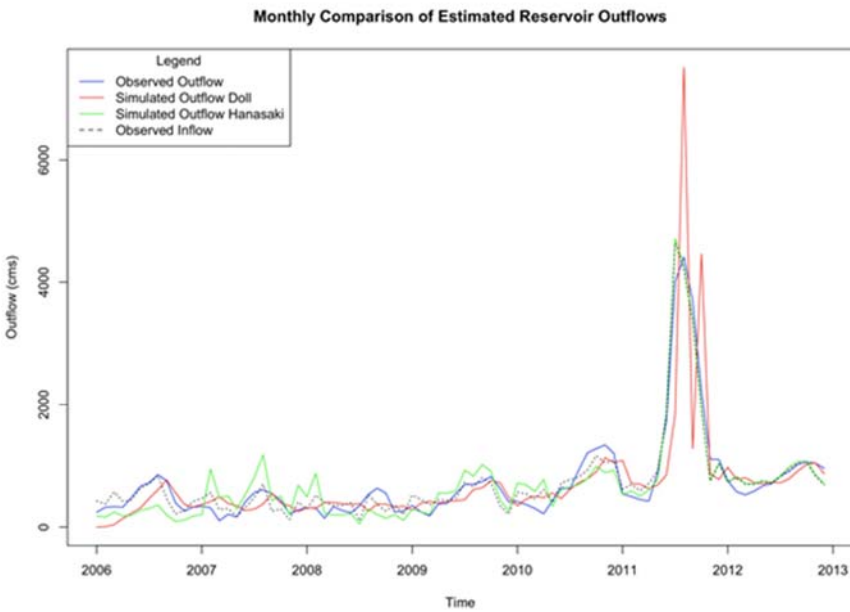


Figure 9. Comparison of simulated outflow for the Fort Randall Dam with Hanasaki and Doll methods for (a) monthly and (b) daily time steps

Figure 10. Comparison of simulated outflow for the Fort Randall Dam with Hanasaki and Döll methods for (a) monthly and (b) daily time steps.



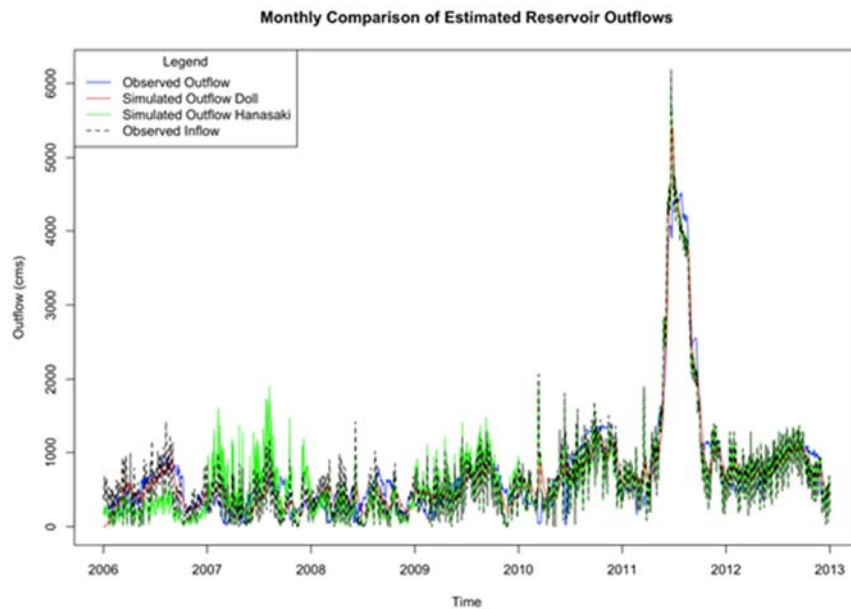


Figure 910. Comparison of simulated outflow for the Fort Randall Dam with Hanasaki and Doll methods for (a) monthly and (b) daily time steps.

### 3.6. Model Stability

Although the Doll Method D03 outperformed the Hanasaki Method H06 when using a daily time step, the Doll Method D03 demonstrated some instability for high  $k_{rd}$  values. This instability occurs at three reservoirs in this study. The cause of the instability is a combination of a reservoir having a low  $IR_{impoundment\ ratio}$  and a sharp change in the inflow to a reservoir. For instance, inflow into Old Hickory Dam in the Nashville District ( $IR = 0.04$ ) increased by roughly two orders of magnitude in a matter of a few days in May 2010. During this event, the available storage filled up, necessitating a substantial increase in release flow to prevent overtopping. This occurred within a single time step in the model

(~~Döll MethodD03~~) and the outflow responded in kind in the next subsequent time step which then drained the reservoir below the specified minimum storage resulting in a non-computable imaginary number as the next solution.

Several solutions are posited to address ~~Döll MethodD03~~ instability. One solution could be to varying  $k_{rd}$  values dynamically to mimic reservoir behavior. During large hydrologic events the value of  $k_{rd}$  could reduce the peak of the outflow hydrograph, and then increase during normal events. Another solution is the inclusion of rules and an expanded system of equations that govern the solution. Because the intention of ~~the Döll MethodD03~~ is to approximate flow at a free-flowing weir, coupling operational rules with the simulation may better approximate reality. The rules may be as simple as switching behavior or the algorithm when storage approaches either minimum or maximum reservoir storage. A simple condition was tested for when storage drops below the minimum storage during the daily time step:

$$\text{if } S_t \leq S_{\min} \Rightarrow \begin{cases} S_t = S_{\min} \\ Q_{out} = Q_{in} + \frac{S_t - S_{\min}}{\Delta t} \end{cases} \quad (5)$$

This condition prevents the reservoir from falling below the minimum storage. Outflow from Old Hickory Dam was re-simulated with  $k_{rd} = 0.9$  and the new minimum storage condition (Equation 5). The proposed modification resulted in simulated outflow shown in Figure 119. Outflow is substantially overestimated for one-time step and drops to zero at the next time step. While an oversimplification of actual operations, this condition is similar to an emergency spillway discharge to prevent overtopping. The dam releases tremendous flow for a brief period, when the maximum storage is nearly exceeded and then inhibits the discharge when the storage is at the minimum capacity. The benefit of this



modification is that additional reservoir information is not required. However, further testing and evaluation should be performed to validate this refinement.

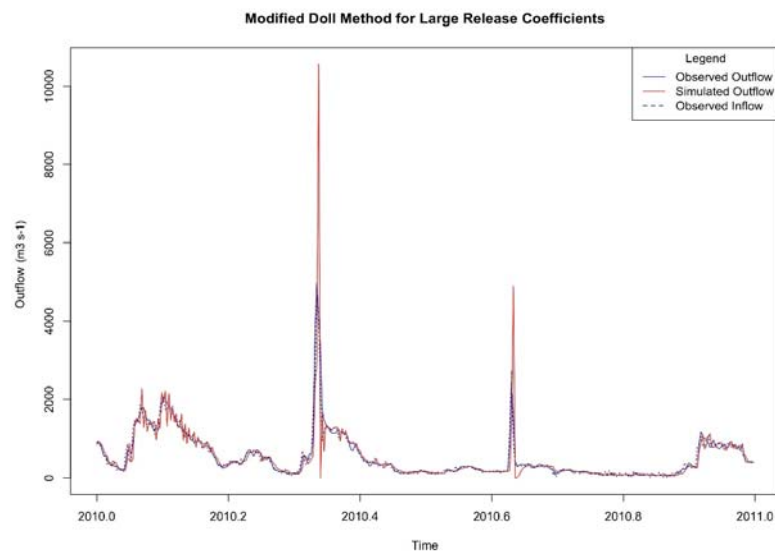


Figure 11. Outflow simulation for the Old Hickory Dam using the proposed modification of the Doll method for  $k_{rd}=0.4$ .

### 3.7. Limitations

~~This study is limited to models that require only reservoir inflow and storage, primarily to provide insight into the reliability of these measures as indicators of reservoir outflow. The inclusion of additional demand and evapotranspiration parameters could improve the results, but could also add considerable uncertainty. Of the two models, only Hanasaki et al. (2006) currently includes an estimate for withdrawals of any nature.~~

~~Another limitation of this study is the inflow that drives the simulations. All inflow utilized in this study, except for the Nashville district, is back-calculated from observed~~

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changes in storage and known discharges. This indirect method can lead to negative inflow values when losses due to seepage, evapotranspiration, or other types of withdrawals are underestimated. De Vos (2015) also noted that they used back-calculated inflow in their study. It is unclear whether Hanasaki et al. (2006) made use of direct observations, but it is worth noting that direct observations of total reservoir inflow are not readily available in most cases difficult to acquire.

This study is limited to models that only require inputs related to only-reservoir inflow and storage, primarily to provide insight into the reliability of these measures as indicators of reservoir outflow. The inclusion of additional demand and evapotranspiration parameters could improve the results, but could also add considerable uncertainty. Because this study utilizes a back calculated reservoir inflow, inclusion of reservoir withdrawal would also lead to an overestimation over-estimation of water withdrawals from the reservoir. Both D03 and H06 can account for withdrawals with draws but because on the basis of the focus of this study and the data utilized, the authors do not pursue an estimation of reservoir withdrawal in this study. Of the two models, only Hanasaki et al. (2006) currently includes an estimate for withdrawals of any nature.

Beyond this studies sensitivity analysis, no formal calibration procedure was undertaken. A formal calibration of  $k_{rd}$  in both D03 and H06 would be better suited for the insertion of the reservoir routing scheme within a hydrologic routing scheme. This study is investigating the feasibility of these methods in 0-10 day lead time, medium range, diurnal forecasting and is a precursor to implementation in hydrologic routing schemes. There is limited benefit to standalone calibration of the  $k_{rd}$  coefficients, given that

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reservoir outflow information is rarely available at global scales. Operational calibration of  $k_{rd}$  would be challenging without reservoir release records. Zajac et al. (2017) discuss the need for an open access database of daily reservoir records, but no such database is known to be available at this time. Thus, this study does not undertake any standalone, formal calibration of  $k_{rd}$ .

Because the vast majority of reservoirs in the sample we considered are primarily purposed as flood control reservoirs with secondary purposes, we are unable to make an assertion about the effect the operating objective has on reservoir routing performance.

### 3.8. Future Work

~~The non-data-driven approaches evaluated~~ consistently improved simulated, daily, streamflow estimates over naturalized flow conditions, suggesting that ~~these approaches~~ can potentially improve global streamflow forecasting that do not already account for lakes and reservoirs. ~~The Döll Method~~ D03 performed particularly well at daily time steps commensurate with many large-scale stream routing models. The incorporation of ~~the Döll Method~~ D03 into the RAPID code, a large-scale river routing model for simulating streamflow throughout distributed stream networks over large spatial extents (David et al., 2011), is under development. This will enable widespread testing and evaluation over large hydrologically diverse areas.

The non-data-driven methods we consider are conceptualizations of reservoir operations that can be adapted to utilize remotely sensed information, much like the data-driven methods previously mentioned. Non-data-driven methods can be linked to statistical fitting techniques, but they are capable of being employed independent of such pairings. However, the non-data-driven reservoir routing schemes could be enhanced by

assimilating remotely sensed data, e.g. near real-time changes in storage resolved from satellite altimetry, and eventually the planned NASA Surface Water and Ocean Topography (SWOT) Mission. This information could constrain reservoir simulations to improve global streamflow forecasts (Yoon and Beighley, 2015). ~~These simulations could provide the training data necessary for more data-intensive reservoir routing approaches, e.g. applying Artificial Intelligence and Machine Learning techniques to infer reservoir rule-curves.~~

Because D03 skill tends to decline with increases in IR, an over-year simulation capability similar to that proposed by De Vos (2015) may allow for a better means of simulating diurnal reservoirs from reservoirs with large IR. Over-year reservoirs have high IRs and yearly cycles of water storage and release are not necessary (Adeloye and Montaseri, 2000; Vogel et al., 1999). ~~Eventually, global streamflow forecasting models should leverage all available data to account for anthropogenic influence, utilizing techniques that range from simple to extremely complex.~~

#### 4. Conclusions

This research compares two parsimonious reservoir routing methods (D03 and H06) with the intent to determine if these methods can be effective at estimating diurnal reservoir outflow in diurnal, medium-range streamflow forecasting. ~~that have previously been implemented in large-scale hydrologic modeling applications, namely the Döll D03 and Hanasaki Methods H06.~~ These methods were compared across 60 USACE operated reservoirs at a daily time step. Results show that ~~the Döll Method D03~~ tends to outperform ~~the Hanasaki Method H06~~ at a daily time step. An in depth examination of these results yields the following conclusions.

- The complexity and data requirements of both ~~Döll-D03~~ and ~~Hanasaki MethodsH06~~ are low and thus computationally inexpensive. Both can be feasibly implemented at large spatial scales at a daily or sub-daily time step.
- When the best performing  $k_{rd}$  is implemented within D03 we find a substantial improvement in the model skill over the baseline for nearly all reservoirs in this study when compared at a daily time step. H06 offers only a minimal improvement over the baseline when the best  $k_{rd}$  is implemented for a daily time step. For the categories of KGE specified (Tavakoly et al., 2017), the best performing D03 eliminates all poor performing baseline conditions and increases the proportion of good or very good performing sites by 22%.
- There is a ~~statistical significant~~ relationship between reservoir IR and two of the skill metrics applied (KGE and R-Squared). Given that reservoirs with high IR typically are less responsive to short-term fluctuations in inflow and storage, the correlation between these variables is plausible. Further investigation of dam characteristics, such as if the dams operate in series or in parallel and wet and dry year considerations are further evidence of the correlation between the IR and ~~Döll D03~~ and ~~Hanasaki MethodsH06~~ skill.
- Simulation time step ~~appears to be an~~ plays an important component part in reservoir routing skill. The comparison of the two methods by Hanasaki et al. (2006) are based on monthly reservoir outflows and conclusions may not hold within diurnal forecasting schemes. At overlapping locations, this study replicates the results reported by Hanasaki et al. for monthly time steps. However, the Hanasaki et al. findings do not hold for a daily time step evaluation.

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- The best value for the empirical Döll coefficient,  $k_{rd}$ , can vary. Optimal values were typically greater than the  $k_{rd}=0.01$  value which Döll et al. (2003) derived. This suggests that  $k_{rd}$  could be a potential calibration parameter within a large-scale hydrologic modeling framework much like a weir coefficient, which is specific to a particular type of weir.
- The Yazoo Basin Headwaters Project (USACE, 2017; USACE, 1987) is an interesting case study in how reservoir system complexity can be difficult to model. The Yazoo Basin Headwaters Project considers downstream flow conditions as the dominant criteria in dam operation. Thus, the inflow and available storage volume are poor predictors for determining reservoir discharge in this type of management scheme. ~~The Döll Method D03~~ appeared to scale flow correctly at these reservoirs and improve reservoir overall skill, but timing of the releases is not well represented and thus skill improvement is only minimal.
- Dam discharges in the Missouri River Reservoir System (Lund and Ferreira, 1996) are more correlated with storage volume and inflow conditions, which lends itself to the two non-data-driven approaches evaluated here. ~~The Döll Method D03~~ is particularly capable of accurately modeling daily reservoir outflows in reservoir systems that correlate well with storage and inflow fluctuations. Concerns related to model error being compounded through a series dams may be mitigated somewhat by the fact that inflow appears to be a progressively stronger predictor of outflow further downstream in these types of systems.
- Numerical stability of ~~the Döll Method D03~~ is a concern, particularly with higher  $k_{rd}$  values. These stability concerns originate at reservoirs with small active storage

capacity during high inflow events. Additional model refinement can overcome these stability concerns.

- ~~The Döll Method~~D03 showed minimal bias during relatively wet and dry years. Timing of releases can be influenced by wet years and the magnitude appears to be affected during dry years. ~~The Döll Method~~D03 appears to be most applicable for dam systems where reservoir management focuses on upstream hydrologic conditions. Large IRimpoundment-ratios could indicate reservoirs where downstream conditions are more likely to influence release decisions at the reservoir.

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- 797 Arcadia Lake Park Office. (2018). Arcadia Lake. Retrieved from  
798 <http://edmondok.com/338/Arcadia-Lake>
- 799 Adeloye, A. J., & Montaseri, M. (1999). Predicting critical period to characterise over-  
800 year and within-year reservoir systems. *Water Resources Management*, 13(6), 383–  
801 407. <https://doi.org/10.1023/A:1008185304170>
- 802 Alfieri, L., Burek, P., Dutra, E., Krzeminski, B., Muraro, D., Thielen, J., & Pappenberger,  
803 F. (2013). GloFAS-global ensemble streamflow forecasting and flood early warning.  
804 *Hydrology and Earth System Sciences*, 2013, 17(3), 1161–1175.  
805 <https://doi.org/10.5194/hess-17-1161-2013>
- 806 Batalla, R. J., Gómez, C. M., & Kondolf, G. M. (2004). Reservoir-induced hydrological  
807 changes in the Ebro River basin (NE Spain). *Journal of Hydrology*, 290(1–2), 117–  
808 136. <https://doi.org/10.1016/j.jhydrol.2003.12.002>
- 809 Biemans, H., Haddeland, I., Kabat, P., Ludwig, F., Hutjes, R. W. A., Heinke, J., ...  
810 Gerten, D. (2011). Impact of reservoirs on river discharge and irrigation water  
811 supply during the 20th century. *Water Resources Research*, 47(3), 1–15.  
812 <https://doi.org/10.1029/2009WR008929>
- 813 Bonnema, M., Sikder, S., Miao, Y., Chen, X., Hossain, F., Ara Pervin, I., ... Lee, H.  
814 (2016). Understanding satellite-based monthly-to-seasonal reservoir outflow  
815 estimation as a function of hydrologic controls. *Water Resources Research*, 52(5),  
816 4095–4115. <https://doi.org/10.1002/2015WR017830>
- 817 Burek, P., Knijff, J. v. d., & Roo, A. de. (2013). *LISFLOOD: Distributed Water Balance*  
818 *and Flood Simulation Model*. Luxembourg, Belgium. <https://doi.org/10.2788/24719>
- 819 Chaves, P., & Chang, F.-J. (2008). Intelligent reservoir operation system based on  
820 evolving artificial neural networks. *Advances in Water Resources*, 31(6), 926–936.  
821 <https://doi.org/10.1016/j.advwatres.2008.03.002>
- 822 Coerver, H. M., Rutten, M. M., & van de Giesen, N. C. (2017). Deduction of Reservoir  
823 Operating Rules for Application in Global Hydrological Models. *Hydrology and*  
824 *Earth System Sciences Discussions*, (January), 1–27. [https://doi.org/10.5194/hess-](https://doi.org/10.5194/hess-2016-660)  
825 2016-660
- 826 David, C. H., Maidment, D. R., Niu, G., Yang, Z., Habets, F., & Eijkhout, V. (2011).  
827 River Network Routing on the NHDPlus Dataset.  
828 <https://doi.org/10.1175/2011JHM1345.1>

Formatted: Font: Italic

Formatted: Font: Italic

829 De Vos, J. (2015). *Non data-driven reservoir outflow and storage simulations in*  
830 *hydrological models*. TU Delft.

831 Döll, P., Fiedler, K., & Zhang, J. (2009). Global-scale analysis of river flow alterations  
832 due to water withdrawals and reservoirs. *Hydrology and Earth System Sciences*  
833 *Discussions*, 6(4), 4773–4812. <https://doi.org/10.5194/hessd-6-4773-2009>

834 Döll, P., Kaspar, F., & Lehner, B. (2003). A global hydrological model for deriving water  
835 availability indicators: Model tuning and validation. *Journal of Hydrology*, 270(1–  
836 2), 105–134. [https://doi.org/10.1016/S0022-1694\(02\)00283-4](https://doi.org/10.1016/S0022-1694(02)00283-4)

837 Ehsani, N., Fekete, B. M., Vörösmarty, C. J., & Tessler, Z. D. (2016). A neural network  
838 based general reservoir operation scheme. *Stochastic Environmental Research and*  
839 *Risk Assessment*, 30(4), 1151–1166. <https://doi.org/10.1007/s00477-015-1147-9>

840 Emerton, R. E., Stephens, E. M., Pappenberger, F., Pagano, T. C., Weerts, A. H., Wood,  
841 A. W., Salamon, P., Brown, J. D., Hjerdt, N., Donnelly, C., Brown, J. D., Hjerdt, N.,  
842 Donnelly, C., Baugh, C. A., & Cloke, H. L. (2016). et al. Continental and global  
843 scale flood forecasting systems. *Wiley Interdisciplinary Reviews: Water*,  
844 2016;3(3), 391–418. <https://doi.org/10.1002/wat2.1137>

845 Escobar, R., & Boulder, C. (2017). *Remote Sensing of Hydrological Extremes*.  
846 <https://doi.org/10.1007/978-3-319-43744-6>

847 Giacomoni, M. H., Kanta, L., & Zechman, E. M. (2013). Complex Adaptive Systems  
848 Approach to Simulate the Sustainability of Water Resources and Urbanization.  
849 *Journal of Water Resources Planning and Management*, 139(June), 554–564.  
850 [https://doi.org/10.1061/\(ASCE\)WR.1943-5452](https://doi.org/10.1061/(ASCE)WR.1943-5452)

851 Global Runoff Data Centre (GRDC). (2017). Welcome to the Global Runoff Data Centre.  
852 Retrieved January 8, 2018, from  
853 [http://www.bafg.de/GRDC/EN/Home/homepage\\_node.html](http://www.bafg.de/GRDC/EN/Home/homepage_node.html)

854 Graf, W. L. (2006). Downstream hydrologic and geomorphic effects of large dams on  
855 American rivers. *Geomorphology*, 79(3–4), 336–360.  
856 <https://doi.org/10.1016/j.geomorph.2006.06.022>

857 Gupta, H. V., Kling, H., Yilmaz, K. K., & Martinez, G. F. (2009). Decomposition of the  
858 mean squared error and NSE performance criteria: Implications for improving  
859 hydrological modelling. *Journal of Hydrology*, 377(1–2), 80–91.  
860 <https://doi.org/10.1016/j.jhydrol.2009.08.003>

861 Haddeland, I., Skaugen, T., & Lettenmaier, D. P. (2006). Anthropogenic impacts on  
862 continental surface water fluxes. *Geophysical Research Letters*.  
863 <https://doi.org/10.1029/2006GL026047>

864 Hanasaki, N., Kanae, S., & Oki, T. (2006). A reservoir operation scheme for global river  
865 routing models. *Journal of Hydrology*, 327(1–2), 22–41.  
866 <https://doi.org/10.1016/j.jhydrol.2005.11.011>

867 Hejazi, M. I., Cai, X., & Ruddell, B. L. (2008). The role of hydrologic information in  
868 reservoir operation - Learning from historical releases. *Advances in Water*  
869 *Resources*, 31(12), 1636–1650. <https://doi.org/10.1016/j.advwatres.2008.07.013>

870 Khalil, A., McKee, M., Kemblowski, M., & Asefa, T. (2005). Sparse Bayesian learning  
871 machine for real-time management of reservoir releases. *Water Resources Research*,  
872 41(11), 1–15. <https://doi.org/10.1029/2004WR003891>

873 Lehner, B., Liermann, C. R., Revenga, C., Vörösmarty, C., Fekete, B., Crouzet, P., ...  
874 Frenken, K. (2011). *Global Reservoir and Dam (GRanD) database*. *European*  
875 *Environment*.

876 Lund, J. R., & Ferreira, I. (1996). Operating Rule Optimization for Missouri River  
877 Reservoir System. *Journal of Water Resources Planning and Management*, 122(4),  
878 287–295. [https://doi.org/10.1061/\(ASCE\)0733-9496\(1996\)122:4\(287\)](https://doi.org/10.1061/(ASCE)0733-9496(1996)122:4(287))

879 [Macian-Sorribes, H., & Pulido-Velazquez, M. \(2017\). Integrating historical operating](#)  
880 [decisions and expert criteria into a DSS for the management of a multireservoir](#)  
881 [system. Journal of Water Resources Planning and Management. 143\(1\).](#)  
882 [\[https://doi.org/10.1061/\\(ASCE\\)WR.1943-5452.0000712\]\(https://doi.org/10.1061/\(ASCE\)WR.1943-5452.0000712\)](#)

883 [Masaki, Y., Hanasaki, N., Biemans, H., Müller Schmied, H., Tang, Q., Wada, Y.,](#)  
884 [Gosling, S. N., Takahashi, K., & Hijikawa, Y. \(2017\). Intercomparison of global river](#)  
885 [discharge simulations focusing on dam operation at multiple models analysis in two](#)  
886 [case-study river basins, Missouri–Mississippi and Green–Colorado. \*Environ. Res.\*](#)  
887 [\*Lett.\* 12\(5\), 1–16. <https://iopscience.iop.org/article/10.1088/1748-9326/aa57a8/pdf>](#)

888 Magilligan, F. J., & Nislow, K. H. (2005). Changes in hydrologic regime by dams.  
889 *Geomorphology*, 71(1–2), 61–78. <https://doi.org/10.1016/j.geomorph.2004.08.017>

890 McManamay R. A. (2014). Quantifying and generalizing hydrologic responses to dam  
891 regulation using a statistical modeling approach. *Journal of Hydrology*, 519(Part A):  
892 2014;519(PA):1278–1296. <https://doi.org/10.1016/j.jhydrol.2014.08.053>

893 Meade, R. H., & Moody, J. A. (2010). Causes for the decline of suspended-sediment  
894 discharge in the Mississippi River system, 1940–2007. *Hydrological Processes*, 24,  
895 35–49. <https://doi.org/10.1002/hyp>

896 Meigh, J. R., McKenzie, A. A., & Sene, K. J. (1999). A grid-based approach to water  
897 scarcity estimates for eastern and southern Africa. *Water Resources Management*,  
898 13(2), 85–115. <https://doi.org/10.1023/A:1008025703712>

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899 Mohan, S., & Ramsundram, N. (2016). Predictive Temporal Data-Mining Approach for  
 900 Evolving Knowledge Based Reservoir Operation Rules. *Water Resources*  
 901 *Management*, 30(10), 3315–3330. <https://doi.org/10.1007/s11269-016-1351-5>

902 Neitsch, S., Arnold, J., Kiniry, J., & Williams, J. (2011). *Soil & Water Assessment Tool*  
 903 *Theoretical Documentation Version 2009*. Texas Water Resources Institute.  
 904 <https://doi.org/10.1016/j.scitotenv.2015.11.063>

905 Nilsson C, Reidy CA, Dynesius M, & Revenga C. (2005). Fragmentation and Flow  
 906 Regulation of the World's Large River Systems. *Science*, 308(5720), 7  
 907 2005;308(5720):405–408. <https://doi.org/10.1126/science.1107887>

908 ~~Patra, A., & Debbarma, N. (n.d.). Prediction of Reservoir Release using Genetic~~  
 909 ~~Programming and ANFIS Models Coupled with Wavelet Transform, 101–108.~~

910 Salas, F. R., Somos-Valenzuela, M. A., Dugger, A., Maidment, D. R., Gochis, D. J.,  
 911 David, C. H., Yu, W., Ding, D., Clark, E. P., & Noman, N. (2017). Towards Real-  
 912 Time Continental Scale Streamflow Simulation in Continuous and Discrete Space.  
 913 *Journal of the American Water Resources Association*, 54(1), 7-27.  
 914 <https://doi.org/10.1111/1752-1688.12586>. 2017;92373(February 2018):7–27.

915 Snow, ~~A. Alan D., Scott D.~~ Christensen S. D., ~~Nathan R.~~ Swain, N. R., ~~James Nelson, E.~~  
 916 ~~J., Daniel P.~~ Ames, D. P., ~~Norman L.~~ Jones, N. L., ~~Deng Ding, D., Nawajish~~  
 917 ~~Noman, N., & Cédric H.~~ David, F. P. C. H. (2016). A High-Resolution National-  
 918 Scale Hydrologic Forecast System from a Global Ensemble Land Surface Model.  
 919 *Journal of the American Water Resources Association*, *Special Issue: Open Water*  
 920 *Data Initiative*, 950-964. ~~in press (4)~~. <https://doi.org/10.1111/1752-1688.12434>

921 Solander, K. C., Reager, J. T., Thomas, B. F., David, C. H., & Famiglietti, J. S. (2016).  
 922 Simulating Human Water Regulation: The Development of an Optimal Complexity,  
 923 Climate-Adaptive Reservoir Management Model for an LSM. *Journal of*  
 924 *Hydrometeorology*, 17(3), 725–744. <https://doi.org/10.1175/JHM-D-15-0056.1>

925 Tavakoly, A. A., Snow, A. D., David, C. H., Follum, M. L., Maidment, D. R., & Yang,  
 926 Z.-L. (2017). Continental-Scale River Flow Modeling of the Mississippi River Basin  
 927 Using High-Resolution NHD Plus Dataset. *JAWRA Journal of the American Water*  
 928 *Resources Association*, 78712. <https://doi.org/10.1111/1752-1688.12456>

929 Ticlavilca, A. M., & McKee, M. (2011). Multivariate Bayesian Regression Approach to  
 930 Forecast Releases from a System of Multiple Reservoirs. *Water Resources*  
 931 *Management*, 25(2), 523–543. <https://doi.org/10.1007/s11269-010-9712-y>

932 U.S. Army Corps of Engineers. (2017). Arkabutla Lake History. Retrieved July 13, 2017,  
 933 from [http://www.mvk.usace.army.mil/Missions/Recreation/Arkabutla-Lake/History-](http://www.mvk.usace.army.mil/Missions/Recreation/Arkabutla-Lake/History-and-Mission/)  
 934 [and-Mission/](http://www.mvk.usace.army.mil/Missions/Recreation/Arkabutla-Lake/History-and-Mission/)

Formatted: Font: Italic

Formatted: Font: Italic

- 935 U.S. Army Corps of Engineers. (1987). *Yazoo Basin Delta Flood Control: Environmental*  
936 *Impact Statement*. Vicksburg, MS.
- 937 Van Beek, L. P. H., Wada, Y., & Bierkens, M. F. P. (2011). Global monthly water stress:  
938 1. Water balance and water availability. *Water Resources Research*.  
939 <https://doi.org/10.1029/2010WR009791>
- 940 Vogel, R. M., Lane, M., Ravindiran, R. S., & Kirshen, P. (1999). Storage Reservoir  
941 Behavior in the United States. *Journal of Water Resources Planning and*  
942 *Management*, 125(5), 245–254. Retrieved from  
943 [http://ascelibrary.org/doi/pdf/10.1061/\(ASCE\)0733-9496\(1999\)125:5\(245\)](http://ascelibrary.org/doi/pdf/10.1061/(ASCE)0733-9496(1999)125:5(245))
- 944 Vörösmarty, C. J., Sharma, K. P., Fekete, B. M., Copeland, A. H., Holden, J., Marble, J.,  
945 & Lough, J. A. (1997). The Storage and Aging of Continental Runoff in Large  
946 Reservoir Systems of the World. *Ambio*, 1997, 26(4), 210–219.
- 947 Wada, Y., Wisser, D., & Bierkens, M. F. P. (2014). Global modeling of withdrawal,  
948 allocation and consumptive use of surface water and groundwater resources. *Earth*  
949 *Syst. Dynam.*, 5, 15–40. <https://doi.org/10.5194/esd-5-15-2014>
- 950 Widén-Nilsson, E., Halldin, S., & Xu, C. (2007). Xu C Yu. Global water-balance  
951 modelling with WASMOD-M: Parameter estimation and regionalisation. *Journal of*  
952 *Hydrology*, 2007, 340(1–2), 105–118. <https://doi.org/10.1016/j.jhydrol.2007.04.002>
- 953 Wisser, D., Fekete, B. M., Vörösmarty, C. J., & Schumann, A. H. (2010). Reconstructing  
954 20th century global hydrography: a contribution to the Global Terrestrial Network-  
955 Hydrology (GTN-H). *Hydrol. Earth Syst. Sci*, 14, 1–24. Retrieved from  
956 [www.hydrol-earth-syst-sci.net/14/1/2010/](http://www.hydrol-earth-syst-sci.net/14/1/2010/)
- 957 Wu, H., Adler, R. F. R., Tian Y., Huffman G. J., Li H., & Wang J. (2014). Real-time  
958 global flood estimation using satellite-based precipitation and a coupled land surface  
959 and routing model. *Water Resources Research*, 50(3), 2693–2717.  
960 <https://doi.org/10.1002/2013WR014710>..... 2014 [accessed 2014 Sep  
961 15]; 50(3):2693–2717.  
962 <http://onlinelibrary.wiley.com/doi/10.1002/2013WR014710/full>
- 963 Wu, Y., & Chen, J. (2012). An Operation-Based Scheme for a Multiyear and  
964 Multipurpose Reservoir to Enhance Macroscale Hydrologic Models. *Journal of*  
965 *Hydrometeorology*, 13(1), 270–283. <https://doi.org/10.1175/JHM-D-10-05028.1>
- 966 Yoon, Y., & Beighley, E. (2015). Simulating streamflow on regulated rivers using  
967 characteristic reservoir storage patterns derived from synthetic remote sensing data.  
968 *Hydrological Processes*, 29(8), 2014–2026. <https://doi.org/10.1002/hyp.10342>

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Formatted: Font: Italic

Formatted: Font: Italic

Formatted: Font: Italic

Formatted: Font: Italic

- 969 Zajac, Z., Revilla-Romero, B., Salamon, P., Burek, P., Hirpa, F. A., & Beck, H. (2017).  
970 The impact of lake and reservoir parameterization on global streamflow simulation.  
971 *Journal of Hydrology*, 548, 552–568. <https://doi.org/10.1016/j.jhydrol.2017.03.022>
- 972 Zhao, G., Gao, H., Naz, B. S., Kao, S. C., & Voisin, N. (2016). Integrating a reservoir  
973 regulation scheme into a spatially distributed hydrological model. *Advances in*  
974 *Water Resources*, 98, 16–31. <https://doi.org/10.1016/j.advwatres.2016.10.014>