

We thank the Reviewer for their positive review and constructive suggestions, which allowed us to improve the clarity of the manuscript. In the following response, reviewer comments are in bold font, author responses in a regular font, and changes made to the text are in italic font.

Authors response to general comments:

I suggest to restructure and shorten section 5.1 and 5.2 of the discussion. It is rather long for the main findings of the study and repeats many points. For example P10 L1-26 discusses preferential flow for the argumentation of snowmelt partitioning. However, preferential flow is again discussed in a similar way on P11 L20 - P12 L19. I suggest to summarize these parts.

We agree with the reviewer that section 5.1 and 5.2 can be written more concisely, Dr. Ireson has also made this point. We have made an effort to reduce the length of these sections and remove redundant points.

Furthermore, I suggest to move the Table 3 and arguments on P11 L20-24 to the results (as it also reads like results). Table 3 is not mentioned in the results and the authors already discuss the relevance of preferential flow without pointing out its evidence in the results.

We thank the reviewer for this point, but we believe the that best place for Table 3 is in its current section of 5.2 which specifically discusses depression and ponded infiltration dynamics. The purpose of Table 3 is to summarise the collective evidence for preferential flow occurring during depression-focused infiltration and groundwater recharge at the study sites, without having to re-discuss points in the results section. Furthermore, we don't discuss why these observations are evidence for preferential flow until section 5.2, so we believe it would be out of context in the results section.

Additionally, the authors should give some more information on the soil properties (if possible). Especially porosity seems to be important in the context of the study, because air-filled porosity is used for the argumentation of preferential flow. E.g. antecedent soil moisture in relation to porosity can be used as an estimate of air-filled porosity available for infiltration.

We agree with the reviewer that information on porosity is important and we have included an additional table including soil information (porosity, bulk density, grain size distribution, K_{sat}). Additionally, to address this comment and those of Dr. Ireson, we have split Figure 1 into two separate figures with additional soil and sediment information. We have also included a new figure with depth- K_{sat} profiles for the 3 sites.

Please make sure that you be consistent with the description in the results. For example, you do not mention MW2 for the Spyhill Upland. Furthermore, sometimes it is not completely clear if you still talk about the upland or the depression (e.g. P6 L27-29)

We thank the reviewer for catching this oversight. We have added some discussion of MW2 to the Spyhill results section 4.3 stating that incomplete snowcover depletion at GP during MW2 prevented the closure of the water balance, but no runoff was observed. We have also gone through all results sections to ensure it is clear in each section which landscape position we are referring to.

Author response to specific comments:

P3 L16: Change “: : of the region surrounding the study sites : : :” to “: : : of the study sites: : :”

Changed.

P4 L13: Please specify why it was not always possible to use pressure transducers.

Transducers were not able to be deployed during MW events due to the possibility of them being damaged by freezing of water in the pond. We have added text to the methods section explaining this.

P5 L10-11: How many soil cores were taken and at which depth? I think you first mention this in the results. Why has the K_{sat} using a permeameter been only determined for this site? Specify that all K_{sat} measures were performed for unfrozen soil.

K_{sat} at depths below the frost zone (2-3 m) were only taken at SE2 because of the water table depth at this site. A piezometer at 3m below the ground surface was installed at this site but the water table is generally more than 5 m below ground surface in the depression and the piezometer remained dry except for a few days during snowmelt. Thus, no slug test was able to be performed at the depth right below the frost zone (i.e., 3 m piezometer) and lab permeameter results were used to fill this gap in the data. We have added text stating this.

We have also specified that all K_{sat} measurements were performed under unfrozen conditions.

P5 L17-18: The slug tests were done for all sites? Please specify this.

At the time of the initial submission, slug tests were not carried out in the shallowest piezometers in depressions SE2 (Stauffer) and W (Triple G) due to the fact that those piezometers tended to be dry (other than during snowmelt) or had water levels within the screened interval of the piezometer, which would make slug test results unreliable. Since submission of the original manuscript, the piezometer in depression W was slug tested when conditions were more favorable, and this value ($3 \times 10^{-7} \text{ m s}^{-1}$) has been included in the updated text and subsequent discussion. See previous comment regarding SE2.

P5 L24: Please clarify that even if R_o is underestimated (P4 L23-24) the catchment wide infiltration rate I is correct. An underestimated R_o leads to an overestimation of I at the uplands, but an underestimation of I in the depressions. From the pure equation one can

think that the calculated I also contains the error resulting from the R0 estimation (hence I would be too high).

We thank the reviewer for bringing up this inciteful point. We have added text discussing this point.

‘In the formulation of equation 1, R_0 is slightly underestimated, which leads to an overestimation of I under uplands, but also underestimates I in the depression, thus we believe this is a reasonable estimate of the catchment wide I.’

P6 L15 and Table 2: Be more consistent for Stauffer MW1 and use “-“ for all components or leave the event out since it was not calculated.

We have included measured values of precipitation and vapour flux measured over the period and left ‘-’ for the components which could not be measured or estimated.

P7 L5-6: Again, I think how you calculated Ksat and that the permeameter tend to underestimate Ksat (with a reference) should be mentioned in the methods.

We have included a statement and reference on the tendency of small core-permeameters to underestimate field K_{sat} .

Reference: Schulze-Makuch, D., Carlson, D. A., Cherkauer, D. S., and Malik, P. 1999. Scale dependency of hydraulic conductivity in heterogeneous media. *Groundwater*, **37**(6): 904-919.

P7 L22 and Table 2: Why did Triple G had no change of SWE during the spring melt event?

A ΔSWE value of zero means that at the start of the time period over which the water balance was calculated before the Spring event, there was no snowcover on the ground surface at the site. During MW2 all snowcover was depleted and so the measured precipitation that fell from the beginning time of the calculation to the end of the snowmelt period would be the total input.

P8 L8-11: Be careful with the comparison of Ksat and recession rates, since a unit gradient assumption is questionable at Triple G with 50 cm ponding head.

We acknowledge that a unit gradient assumption may be questionable under ponding conditions, however the reason we make this comparison is only to note that the rate is the same order of magnitude as measured K_{sat} values. While ponding conditions may create significant hydrostatic pressure at the centre of the pond, gradients at the edge of the pond are likely significantly less than unity. Moreover, it is unlikely that ponding head in saturated partially-frozen soil would create a hydraulic gradient large enough (~ 10) to increase the recession rate to an order of magnitude above K_{sat} . For example, we can compare infiltration in the ponded depression to a much-simplified analog of a single-ring infiltrometer. The increased infiltration flux due to ponding would be roughly ponded head, H , divided by $[0.6 \times \text{pond radius}]$. Thus, a 10 m diameter pond with 0.50 m pond height, would increase flux by $0.5 / (0.6 \times 5) = 0.17$ i.e. 17% or a

factor of 1.17. We acknowledge this is not a perfect comparison, but it shows the effect is relatively small.

P9 L23-28: The runoff ratios also depend on the snowmelt rates observed during a certain event or year. For a comparison of different sites, events or years it is important to mention this.

We thank the reviewer for bringing this point up and agree that it should be mentioned when comparing runoff ratios. We have added a sentence to the discussion bringing up this point.

‘In addition to frozen soil infiltrability, runoff ratios at any given site also depend on snowmelt rates during a melt event, however some useful insight can be made from comparing runoff ratios here with other studies. The snowmelt runoff ratios from all sites ranged between 0 and 57%, with a mean of 21%. These values are lower than those reported by Hayashi et al. (1998) for a...’

P10 L16: Delete “subsurface” in “subsurface infiltration/refreezing”

Change made.

P10 L28: How can you be sure that these lateral pathways are preferential pathways?

This is a very good point, and the reviewer is correct in stating we cannot be certain that the lateral pathways are preferential. However, there were a few factors that we considered when making this argument, which were not included in the manuscript for the sake of brevity. Firstly, the soil under the upland areas at Triple G were relatively dry and never reached close to saturation during overwinter and spring snowmelt periods. Saturated or nearly saturated conditions would be needed to produce diffuse (non-preferential flow) lateral subsurface flow along a topographic gradient in the near surface soils. However our argument is that if most of the infiltrated water flows through previously air-filled macropores surrounded by a matrix of reduced permeability due to freezing conditions, then macropores would require less water to fill (compared to total porosity), and an impeding layer at some shallow depth in the frozen soil (higher soil frost and/or decreased macroporosity) may cause some of the infiltrated water to be laterally deflected along the topographic gradient to depressions.

Secondly, we also performed a simple calculation assuming diffuse saturated flow:

Estimation of volume pond increased by over Spring melt event:

Height pond increased by = 0.1 [m] (Spring event)

Volume of water added to pond = ~74 [m³] (volume calculated from pond height using depth-volume relationship from topographic survey of depression W)

Volumetric flux = K_{sat} × saturated thickness × average slope of catchment × pond circumference

K_{sat} = 1.0×10^{-3} [m s⁻¹] assumed

saturated thickness = 0.1 [m] assumed

Slope = 0.08 [-] estimated from topographic surveys

Diameter of wetted area = ~30 [m] from satellite imagery on March 20, 2017

Circumference of pond = 94.25 [m]

Volumetric flux = ~65 [m³ d⁻¹]

This calculation is very rough and makes several simplifying assumptions but it shows that for saturated diffuse subsurface flow to transport the observed volume of water over the time delays we observed (1 to 2 days), the hydraulic conductivity of the near surface would have to be one to two orders of magnitude higher (10^{-3} m s^{-1}) than the near surface hydraulic conductivity measured with single ring infiltrometers (circa. 10^{-5} - 10^{-6} m s^{-1}). These near surface K_{sat} values are included in the new table and figure mentioned above.

Furthermore, we acknowledge that identifying mechanisms of flow governing subsurface runoff or interflow is very complicated as direct observation of the flowpath is almost impossible. We are currently carrying out chemical tracer experiments at Triple G to better understand this subsurface transmission pathway.

P11 L28: Maybe Graham & Lin (2011) (doi: 10.2136/vzj2010.0119) is a better reference.

Thanks, we have added the reference to text and reference list.

P12 L10-11: Why do you know that infiltration was limited by the hydraulic conductivity of the zone beneath the frozen layer and not by the frozen layer itself? You do not know the saturated hydraulic conductivity of the frozen layer and from the water level recessions you can just estimate an integrated Ksat (with unknown extend) by assuming a unit gradient (what's not always the case).

The reviewer is right that we do not know that hydraulic conductivity of the frozen layer, especially as it changes with time due to changes in total water and ice content at different depths. Our reasoning for stating that the infiltration rate was limited by the hydraulic conductivity of the zone beneath the frozen layer was that the recession/infiltration rate was the same order of magnitude as the K_{sat} measured beneath the frost zone (either by slug tests or permeameters), but was consistently an order of magnitude lower than the K_{sat} of the near surface soil (0-0.3 m) measured by single ring infiltrometers or Guelph permeameters at the sites. Given that ponding conditions saturate the entire soil profile (temporarily), infiltration would be limited by the lowest hydraulic conductivity layer, which in this case was the zone below the frost zone. However, we acknowledge that if macropores are blocked with ice, then the frost layer hydraulic conductivity would be the limiting factor. We have modified the text to make this point more clearly.

P12 L15-16 I would add a reference here, e.g. Schwen et al. (2011) (doi:10.1016/j.still.2011.02.005)

We thank the reviewer for this reference and have added it to the text and references.

P12 L30-31 I would argue that the infiltration is rather dependent on the amount of connected and air-filled macropores in the frozen layer than on the infiltration rate of the subsoil. Without connected macropores the infiltration rate of the topsoil would be much lower since only the frozen soil matrix would conduct water and hence it would take the water a long time to even reach the subsoil.

We completely agree with the reviewer. Our argument that the infiltration rate was dependent on the K_{sat} of the subsoil was precisely because it was not being limited by the frozen layer above because of the presence of conductive macropores. However, as mentioned above, we acknowledge that if macropores are blocked with ice, then the frost layer hydraulic conductivity would be the limiting factor and have modified the text accordingly.

P13 L31: Change to “: : during midwinter snowmelt enhanced surface runoff generation: :”

Change made.

P14 L23: Data availability: Delete “and”. Furthermore, there is no Table in the Supplements.

We apologise for this oversight and will include the table in the updated supplement.