

1 **Evaluation of impacts of future climate change and water use**
2 **scenarios on regional hydrology**

3 **Seungwoo Chang¹, Wendy Graham^{1, 2}, Jeffrey Geurink³, Nisai Wanakule³, and**
4 **Tirusew Asefa³**

5 ¹Water Institute, University of Florida, 570 Weil Hall, PO Box 116601, Gainesville, FL 32611,
6 USA

7 ² Department of Agricultural and Biological Engineering, University of Florida, 570 Weil Hall,
8 PO Box 116601, Gainesville, FL 32611, USA

9 ³ Tampa Bay Water, 2575 Enterprise Rd, Clearwater, FL 33763-1102, USA

10 Corresponding author: S. Chang (swjason@ufl.edu)
11

12 **Abstract**

13 General circulation models (GCMs) have been widely used to simulate current and future
14 climate at the global scale. However, the development of frameworks to apply GCMs to assess
15 potential climate change impacts on regional hydrologic systems, the ability to meet future water
16 demand, and compliance with water resource regulations is more recent. In this study eight
17 GCMs were bias-corrected and downscaled using the Bias Correction and Stochastic Analog
18 (BCSA) downscaling method and then used, together with three ET₀ methods, and eight different
19 water use scenarios to drive an integrated hydrologic model previously developed for the Tampa
20 Bay region in west central Florida. Variance-based sensitivity analysis showed that changes in
21 projected streamflow were very sensitive to GCM selection, but relatively insensitive ET₀
22 method or water use scenario. Changes in projections of groundwater level were sensitive to both
23 GCM and water use scenario, but relatively insensitive to ET₀ method. Five of eight GCMs
24 projected a decrease in streamflow and groundwater availability in the future regardless of water
25 use scenario or ET method. For the business as usual water use scenario all 8 GCMs indicated
26 that, even with active water conservation programs, increases in public water demand projected
27 for 2045 could not be met from ground and surface water supplies while achieving current

28 groundwater level and surface water flow regulations. With adoption of 40% wastewater reuse
29 for public supply and active conservation 4 of the 8 GCMs indicate that 2045 public water
30 demand could be met while achieving current environmental regulations; however, drier climates
31 would require a switch from groundwater to surface water use. These results indicate a high
32 probability of a reduction in future freshwater supply in the Tampa Bay region if environmental
33 regulations intended to protect current aquatic ecosystems do not adapt to the changing climate.
34 Broad interpretation of the results of this study may be limited by the fact that all future water
35 use scenarios assumed that increases in water demand would be the result of intensification of
36 water use on existing agricultural, industrial and urban lands. Future work should evaluate the
37 impacts of a range of potential land use change scenarios, with associated water use change
38 projections, over a larger number of GCMs.

39 **1. Introduction**

40 The Intergovernmental Panel on Climate Change (IPCC) along with many other studies
41 have indicated that climate change is likely to alter both the global hydrologic cycle and regional
42 hydrologic cycles (Aalst et al., 2014; Déry et al., 2009; Georgakakos et al., 2014; Hawkins et al.,
43 2014; Milliman et al., 2008). These studies have indicated that climate change is likely to
44 increase the frequency of droughts, as well as the magnitude of floods in many regions
45 (Diffenbaugh and Field, 2013; Georgakakos et al., 2014; Walsh et al., 2014). It is necessary to
46 investigate future climate change and its potential impacts on the natural environment in order to
47 reduce risks and increase resilience for future water resources planning and management (Vano
48 and Lettenmaier, 2013).

49 General Circulation Models (GCMs) and hydrologic models have been widely used to
50 evaluate future climate change and its impact on regional hydrologic cycles (Boé et al., 2007;
51 Maurer and Hidalgo, 2008). However, there are a variety of barriers to direct use of GCMs to
52 drive regional hydrologic models. For example, the current generation of GCMs contain biases
53 that prevent accurate reproduction of historic hydrological conditions when used to drive
54 hydrologic models (Giorgi and Mearns, 2002; Wood et al., 2002). In addition, the coarse
55 resolution of GCMs prevents direct use of their results with regional hydrologic models that
56 require higher resolution climate variables (Solomon et al., 2007). Many bias correction methods
57 and downscaling methods have been developed and evaluated to overcome these limitations

58 (Chen et al., 2013; Ghosh and Mujumdar, 2008; Hwang and Graham, 2013; Langousis et al.,
59 2015; Muerth et al., 2013; Quintana Seguí et al., 2010; Stoll et al., 2011; Zhang and
60 Georgakakos, 2012). Although these bias correction and downscaling methods do not correct
61 problems with large scale synoptic forcing, and are not particularly good at reproducing extreme
62 floods or droughts in the retrospective period, previous research has shown that they are able to
63 simulate broad features of the climate system and are useful for characterizing plausible
64 projections of possible futures (Kundzewicz et al, 2008, 2009). Furthermore, previous work in
65 the study region has shown that hydrologic models driven by bias-corrected downscaled
66 retrospective GCM output adequately reproduce retrospective high stream flows (e.g. 7Q2 and
67 7Q10), as well as the long term mean and standard deviation of monthly flows (Hwang and
68 Graham, 2014).

69 In addition to studies that focus on climate impacts on the hydrological cycle, it is also
70 necessary to evaluate the effects of direct human behavior (Haddeland et al., 2014; Wang and
71 Hejazi, 2011). Human activities such as agricultural production, irrigation (Gupta et al., 2015),
72 municipal pumping (Patterson et al., 2013), deforestation, and urban development alter regional
73 hydrologic behavior (Siriwardena et al., 2006). For robust water resources management and
74 planning better understanding of the influence and relative importance of climate change and
75 human-induced change on hydrology and water resources is essential (Chang et al., 2016; Ma et
76 al., 2008; Tan & Gan, 2015; Ye et al., 2013; Zheng et al., 2009).

77 The relative contributions of climate change and human activities to hydrologic responses
78 have been evaluated using GCM data to drive hydrologic models with plausible future
79 anthropogenic scenarios (Liu et al., 2013; Maurer et al., 2010; Wood et al., 2002). Murray et al.
80 (2012) used the Land-surface Processes and eXchanges (LPX) dynamic global vegetation model
81 and the WaterGAP hydrological model to evaluate the impacts of climate change and socio-
82 economic change on global hydrologic response for the 2070 – 2099 time period. They found
83 that climate change and population growth increased water stress in many regions, and change in
84 runoff was most highly correlated with precipitation change in large global catchments. Harding
85 et al. (2012) applied downscaled outputs of 16 GCMs with the VIC model to investigate the
86 future change in streamflow for the Colorado river basin. They suggested that impact analyses
87 relying on only a few scenarios were unacceptably influenced by the choice of GCM projections.

88 For studies using GCMs to project future hydrologic responses, uncertainties resulting
89 from the choice of GCM, RCP (Representative Concentration Pathways) trajectory, and
90 reference evapotranspiration (ET_0) estimation methods are all significant, and it is important to
91 quantify the relative uncertainties of these factors (Chang et al., 2016; Hawkins & Sutton, 2009,
92 2010; Kingston et al., 2009; Koedyk & Kingston, 2016; McAfee, 2013; Thompson et al., 2014;
93 W. Wang et al., 2015). Furthermore, the effects of climate change on groundwater levels have
94 not explored as extensively as the effects of climate change on surface water flows (Green et al.,
95 2011; Kløve et al., 2014). Kløve et al. (2014) suggested that the uncertainties of groundwater
96 projections attributed to climate models, downscaling techniques, emission scenarios, land use
97 changes and social economic development should be evaluated.

98 This study evaluated the future projections of regional hydrologic response using eight
99 GCMs, three ET_0 estimation methods, and eight human water use scenarios to drive a calibrated
100 regional hydrologic model developed for the Tampa Bay region. A comprehensive evaluation of
101 the relative sensitivity of projections of regional hydrologic response to the choice of GCM, ET_0
102 estimation method, and human water use scenario was conducted. Statistical analyses were
103 performed to determine whether differences in streamflow and groundwater level between
104 retrospective hydrologic and projected future climate were statistically significant given these
105 underlying prediction uncertainties. The ability to satisfy projected increases in future water
106 demand while meeting current groundwater level and surface water flow regulations was
107 evaluated over the suite of GCM and water management scenarios.

108 **2. Materials and Methods**

109 **2.1 Study Region**

110 Tampa Bay Water operates a diverse regional water supply system comprised of a
111 desalination plant, well fields that extract water from the Floridan Aquifer, and surface water that
112 is extracted from the Hillsborough and Alafia Rivers (<https://tampabaywater.org/water-supply-sources-tampa-bay-region>). The fresh groundwater system in the region is composed of two
113 aquifer systems, a thin surficial aquifer and the thick and highly productive carbonate rocks of
114 the Floridan aquifer system (Tihansky & Knochenmus, 2001). Dynamic interacting surface-
115 water and groundwater systems (in which groundwater from in the aquifer used for agricultural
116

117 irrigation and public water supply also feeds the surface springs and rivers) characterize the
118 region and must be considered in the management of water resources (Tihansky, 1999). For
119 example the SWFWMD regulates groundwater pumping for water supply to maintain
120 groundwater levels that promote environmental protection of lakes and wetlands near well-fields.
121 Similarly they regulate the daily volume of flow permitted for extraction from rivers based on
122 maintaining sufficient in-stream flows and spring flows to protect aquatic ecosystems.

123 This study focused on the Integrated Northern Tampa Bay (INTB) model domain
124 (Geurink and Basso, 2013; Hwang and Graham, 2014). Figure 1 shows the INTB model domain,
125 model sub-basins, locations of agricultural, industrial and public water supply wells, two
126 streamflow locations where water is withdrawn for public supply, and three monitoring wells
127 near Tampa Bay Water's consolidated well fields that are used to evaluate compliance with
128 groundwater level regulations. The INTB region land use currently consists of grass/pasture (25
129 %), urban (22 %), forested (15 %), mining/other (7 %), agriculture/irrigated land (6 %), open
130 water (4 %), and wetlands (21 %).

131 2.2 The Integrated Northern Tampa Bay Model

132 Tampa Bay Water and the Southwest Florida Water Management District (SWFWMD)
133 developed the Integrated Hydrologic Model (IHM) simulation engine which integrates the EPA
134 Hydrologic Simulation Program-Fortran (Bicknell et al., 2005) for surface water modeling with
135 the U.S. Geological Survey (USGS) MODFLOW96 (Harbaugh and McDonald, 1996) for
136 groundwater modeling. The IHM simulates the dynamic interaction of surface water and
137 groundwater systems within the INTB region including all processes which affect flow and water
138 levels in uplands, within the unsaturated soil, and within wetlands, rivers and aquifers. In
139 addition, the INTB model can account for variability in climate and anthropogenic stresses such
140 as land use change, groundwater pumping, and diversions to/from rivers, lakes, and wetlands.

141 Tampa Bay Water and the SWFWMD calibrated model parameters to simulate
142 streamflows, groundwater levels, and wetland hydroperiods in the INTB model region. The
143 INTB model was calibrated from 1989 to 1998 and verified from 1999 to 2006 (Geurink and
144 Basso, 2013). Precipitation data for calibrating and validating the model were obtained from 302
145 point gages maintained by National Oceanic and Atmospheric Administration (NOAA), the
146 SWFWMD, and Tampa Bay Water in the model region. Maximum and minimum daily

147 temperature were obtained from six NOAA stations within the INTB region and used to estimate
148 ET_0 using the Hargreaves method. Over the calibration and validation period (1989 to 2006)
149 average annual precipitation input to the model was 1308 mm/year and average annual actual
150 evapotranspiration estimated by the model was 940 mm/year, resulting in net available water
151 (precipitation-actual evapotranspiration) of 368 mm/yr. During this period surface discharge
152 from the domain was 272 mm/year (74 % of net available water), groundwater pumping was 69
153 mm/year (19 %), surface water diversions for water supply were 10 mm/year (3 %), and
154 irrigation applied within the domain was 18 mm/year (5 %). More details about the processes
155 and results of model calibration and validation are described in Geurink and Basso (2013).

156 Streamflow predictions at two United States Geological Survey (USGS) gauging
157 stations, the Hillsborough river (USGS ID: 02303330) and Alafia river (USGS ID: 02301500),
158 were used in this study to evaluate retrospective and future IHM streamflow predictions and
159 quantities of surface water available for public supply. Three Tampa Bay Water monitoring wells
160 (NWH-RMP-08s, CBR-SERW-s, and STK-STARKEY-20s) were used to evaluate retrospective
161 and future groundwater level predictions and compliance with environmental regulations
162 intended to protect nearby wetlands from dewatering as a result of consolidated well field
163 pumping.

164 2.3 Climate Data

165 Forcing data from Phase 2 of the North American Land Data Assimilation System
166 (NLDAS-2) from 1982 to 2005 were used as historical reference climate data for bias correction.
167 Hourly precipitation, air temperature, solar radiation (surface downward longwave radiation and
168 surface downward shortwave radiation), surface pressure and average wind speed were obtained
169 from the NLDAS-2 archive and aggregated to the daily scale at a 1/8th-degree grid spacing over
170 the Tampa Bay region.

171 For retrospective and future climate data, the Coupled Model Intercomparison Project 5
172 (CMIP5) General Circulation Models (GCMs) data set for the 1982-2005 period was used for the
173 retrospective period and 2030-2060 (Future 1) and 2070-2100 (Future 2) were used as future
174 periods. Gridded daily precipitation, air temperature, solar radiation, surface pressure, and
175 average wind speed were obtained for eight GCMs listed in Table 1. These GCMs were chosen
176 because they spanned the range of cool to warm bias and wet to dry bias exhibited by 41 CMIP5

177 GCMs for the southeastern United States (Rupp, 2016), and they had daily values available for
178 all the parameters needed to estimate Penman-Monteith reference evapotranspiration. Mean
179 changes in precipitation projected by these GCMs ranged from -68 mm/year to 293 mm/year
180 over the 2030-2060 period, and from 154 mm/year to 400 mm/year over the 2070-2100 period.
181 Mean changes in ET_0 ranged from 24 mm/year to 137 mm/year over the 2030-2060 period and
182 from 122 mm/year to 351 mm/year over the 2070-2100 period. Mean changes in $P-ET_0$ ranged
183 from -162 mm/year to 220 mm/year over the 2030-2060 period and from -420 mm/year to 159
184 mm/year over the 2070-2100 period (Table 1).

185 Chang et al. (2016) evaluated projected changes in $P - ET_0$ over the continental USA
186 using nine GCMs, ten ET_0 estimation methods, and three RCP scenarios. They showed that the
187 first order sensitivities of water deficit projections ($P-ET_0$) over the Southeast USA were much
188 higher to choice of GCM and ET_0 estimation method than to choice of RCP. First order
189 sensitivities of water deficit projections to RCP scenarios were negligible (<0.01) for the 2030-
190 2060 time period, and averaged 0.2 for the 2070-2100 time period. Therefore for computational
191 efficiency, and to evaluate the influence of the most extreme carbon dioxide forcing on the
192 hydrologic projections, only the RCP 8.5 scenario data was utilized for the future analyses in this
193 study.

194 2.4 BCSA Downscaling Method

195 The BCSA downscaling method, developed by Hwang and Graham (2013), was used in
196 this study. Hwang & Graham (2014) showed that BCSA demonstrated better performance than
197 other statistical downscaling methods (i.e, BCSD (Maurer et al, 2012) or SDBC (Abatzoglou and
198 Brown, 2012)) in reproducing spatiotemporal statistics of both precipitation and daily streamflow
199 in the Tampa Bay region. In particular, the INTB model, when driven by GCMs downscaled
200 using the BCSA method, accurately reproduced frequencies of extreme high and extreme low
201 retrospective streamflows as well as 7Q2 and 7Q10 retrospective streamflows in the Tampa Bay
202 region.

203 The BCSA method preserves both the cumulative frequency distribution of observed
204 daily precipitation as well as the spatial autocorrelation structure of observed daily precipitation
205 fields. BCSA downscaling consists of two separate steps for bias-correction and stochastic
206 analog spatial downscaling. In the first step, a cumulative distribution function (CDF) mapping

207 approach (Block et al., 2009; Hwang et al., 2013, 2014; Hwang & Graham, 2014; Ines &
208 Hansen, 2006; Teutschbein & Seibert, 2012) is used to reduce the biases in raw GCM output at
209 the GCM scale. In this study, NLDAS-2 P and ET_0 were aggregated up to the GCM scale and P
210 and ET_0 from the raw GCMs were bias corrected at the GCM scale using the sequential
211 univariate CDF mapping method (Chang, 2017). NLDAS-2 was selected for bias correction
212 because it includes all the parameters needed to estimate Penman-Monteith reference
213 evapotranspiration. Comparison of the gridded NLDAS-2 data to the precipitation and
214 temperature observations from the weather stations used to calibrate the INTB model showed
215 that the NLDAS-2 data reproduced observed long term monthly means values with biases that
216 ranged from 4 to 12 mm for daily precipitation and 1 to 2°C for daily temperature. Correlations
217 among daily values ranged from 0.75 to 0.87 for rainfall and 0.75 to 0.98 for temperature. The
218 second step in the BCSA method is stochastic analog (SA) spatial downscaling (Hwang &
219 Graham, 2013, 2014) for P. In this method, a synthetic downscaled precipitation field is
220 produced which preserves the GCM-scale daily precipitation amount and the month-specific
221 local-scale spatial correlation structure. For more details on the BCSA method, see (Hwang &
222 Graham, 2013, 2014). ET_0 was not downscaled in this study because observed spatial variability
223 of ET_0 over the INTB region is very small, and the spatial correlation is large compared to P
224 (Chang, 2017).

225 2.5 Reference Evapotranspiration Estimation Methods

226 The Chang et al. (2016) study referenced above found that the projected changes in P –
227 ET_0 were sensitive to both the choice of GCM and the choice of ET_0 method, and that for the
228 Southeast USA the choice of GCM and ET_0 method had approximately equal influence on
229 changes in future P – ET_0 throughout most of the year. However, they noted that not all ten ET_0
230 methods were equally appropriate for use in all US regions, and that regional studies should use
231 methods for which retrospective predictions of ET_0 are generally consistent with historic
232 observations. Several of the ET_0 methods used by Chang et al. (2016) were found to produce
233 unreasonably high or low historic ET_0 estimates for the study region using retrospective and
234 observation data. Therefore in this study three ET_0 estimation methods that are widely used in the
235 Southeast USA, produced retrospective predictions that were consistent with observations, and
236 showed a range of wet to fairly dry projections of future P- ET_0 (Chang et al, 2016) were
237 included in the analysis. These methods include a temperature-based method (Hargreaves;

238 Hargreaves and Allen, 2003), a radiation-based method (Priestley-Taylor; Allen et al., 1998), and
239 a combination method (Penman-Monteith; Allen et al., 1998). All hourly climate variables
240 described above were aggregated to the daily scale and used to calculate daily ET_0 using these
241 three methods.

242 2.6 Retrospective Simulations

243 Water-use in the study region is comprised of five categories; 1) public supply, 2)
244 agricultural, 3) industrial/commercial, 4) mining, and 5) recreational (e.g. golf course irrigation)
245 (Geurink and Basso, 2013). Groundwater sources are used for agricultural,
246 industrial/commercial, mining and recreational water supplies. Public water supply is provided
247 by a combination of groundwater, surface water (Hillsborough and Alafia Rivers), and a 25
248 MGD desalinization plant operated by Tampa Bay Water. The SWFWMD regulates all
249 groundwater pumping and surface water extraction in the study region to protect natural aquatic
250 ecosystems and prevent saltwater intrusion. Over the 1989-2006 calibration-verification period
251 groundwater extractions from the INTB model domain averaged 36 mm/yr for public water
252 supply, 18mm/yr for agricultural irrigation, 9 mm/year for industrial/commercial uses, 6
253 mm/year for mining, and 3 mm/year for recreational uses (Geurink and Basso, 2013).

254 Public Water Supply: Tampa Bay Water has a consolidated permit for its eleven
255 wellfields (the Consolidated Wellfields, hereafter referred to as the CWF). The CWFs are
256 operated as an interconnected system with a combined maximum permitted pumping rate of 90
257 MGD (13 mm/yr over the INTB region). Individual well pumping rates are optimized to
258 maintain minimum groundwater levels near sensitive wetlands to meet regulatory requirements
259 intended to prevent ecological harm. The three monitoring wells evaluated in this study are
260 located near wetlands adjacent to the CWFs (Fig. 1). From 1992-2008 Tampa Bay Water's total
261 water demand averaged ranged from 150-200 MGD. Groundwater is Tampa Bay Water's most
262 inexpensive source for public water supply, therefore for the retrospective simulations the CWFs
263 were assumed to withdraw groundwater continuously at the 90 MGD maximum permitted rate.
264 For the retrospective simulations groundwater extraction for other public water supply (outside
265 of Tampa Bay Water's CWF), industrial/commercial and mining uses were assumed occur
266 continuously at the average pumping rates between years 1989 to 2006 cited above.

267 Maximum available surface water available to Tampa Bay Water for public supply was
268 calculated on a daily basis from retrospective streamflow predictions for both the Hillsborough
269 River and the Alafia River according to site-specific regulations set to maintain sufficient in-
270 stream flows and spring flows to protect aquatic ecosystems. Diversion rates for pumping from
271 the Hillsborough river reservoir by the City of Tampa and from the Tampa Bypass Canal by
272 SWFWMD were set at the historical average daily rate spanning 2003 to 2009 for all
273 retrospective simulations. All other diversion rates were set to zero including the Withlacoochee-
274 Hillsborough overflow. These diversion locations are located either downstream or outside of the
275 watersheds contributing to the surface water gages, and outside the zone of influence of the
276 monitoring wells evaluated in this study so these assumptions do not impact on the results (Fig.
277 1).

278 Agricultural Irrigation Demand: The AFSIRS (Agricultural Field-Scale Irrigation
279 Requirements Simulation) model (Jacobs and Dukes, 2007; Smajstrla, 1990) was used to
280 estimate climate-driven irrigation demand for the retrospective period. The AFSIRS model tracks
281 the water budget in the crop root zone including inputs from rain and irrigation, and outputs from
282 the root zone by drainage and evapotranspiration. The AFSIRS model defines the water storage
283 capacity in the crop root zone as the product of the water-holding capacity of the soil (estimated
284 by the difference between field capacity and wilting point) and the depth of the effective root
285 zone for the crop being grown. Crop evapotranspiration (ET_c) is estimated from the product of
286 potential evapotranspiration (ET_0) and crop water use coefficients. The AFSIRS model
287 subdivides the crop root zone into irrigated and non-irrigated zones and maintains separate water
288 budgets for both zones in order to simulate different types of irrigation systems, such as surface
289 irrigation and subsurface irrigation.

290 The AFSIRS was used as a basis to estimate irrigation demand for the retrospective
291 period using CMIP5 bias-corrected downscaled daily P and bias-corrected ET_0 (using the three
292 ET_0 methods discussed above) and the land use from the calibrated INTB model. Crop
293 coefficients (K_c) for estimating ET_c were obtained from the calibrated INTB model database
294 (Geurink and Basso, 2013) for all vegetative covers except row crops. The crop coefficient for
295 row crops was estimated by the superposition of crop coefficients for tomato and strawberry
296 (Dukes et al., 2012), the two dominant row crops in the region. The relative proportion of these
297 two crops constituting the row crop land use were calculated based on water usage records for

298 the region for 2011 (Jackson and Albritton, 2013). The root zone depth, field capacity, wilting
299 point and other information needed for the AFSIRS model were taken from the calibrated INTB
300 model database. Groundwater pumping required to satisfy the AFSIRS estimated irrigation
301 assumed 85% irrigation efficiency based on Irmak et al. (2011) and Jacobs & Dukes (2007), i.e.,

$$302 \quad \text{agricultural pumping} = \text{irrigation demand} \times \frac{100 \%}{85 \%} \quad (3)$$

303 It should be noted that the AFSIRS model does not predict water demand for bed
304 preparation, freeze protection, crop cooling requirements, or maintenance of irrigation systems.
305 Thus the irrigation demand estimated for the retrospective period only includes crop water
306 demand for evapotranspiration.

307 Boundary Conditions: Lateral boundary conditions are required for aquifers in the model
308 region. A repeating annual cycle of daily General Head Boundary (GHB) time series for the
309 retrospective and future periods IHM simulations was derived using the daily average of the
310 historical daily GHB time series spanning 2000 to 2006. More details about the water
311 withdrawals such as groundwater pumping, agricultural irrigation, CWFs, diversions and
312 boundary conditions during the calibration-verification period are described in Geurink and
313 Basso (2013).

314 2.7 Future Water Use Scenarios

315 In addition to warming temperatures and reduced precipitation due to climate change,
316 increases in water withdrawal for agriculture and other human uses are potentially significant
317 causes of declining river flow and groundwater levels (Alcamo et al., 2003; Vorosmarty et al.,
318 2000). To assess the relative importance of climate change versus anthropogenic impact on the
319 hydrologic system, ability to meet future water demand, and compliance with water resource
320 regulations in the study region, eight future water use scenarios were developed (Table 2). These
321 scenarios were based on discussions with Tampa Bay Water staff, projected increases in public
322 water demand (Tampa Bay Water Water Demand Management Plan Final Report, 2013),
323 projected changes in agricultural land use and agricultural irrigation demand (Florida Statewide
324 Agricultural Irrigation Demand Report, 2017), potential agricultural adaption behaviors, and
325 potential changes in groundwater regulations. For naming simplicity in the future scenarios
326 agricultural and recreational water use categories are combined as agricultural demand and

327 public supply, industrial/commercial and mining are combined as urban demand. The eight water
328 use scenarios included: 1) No groundwater pumping for agriculture or urban demand, 2) No
329 urban groundwater pumping, 3) No agricultural groundwater pumping, 4) Agricultural adaption
330 (increased irrigation efficiency and/or use of reclaimed water), 5) Business as usual, 6) Increased
331 agricultural demand, 7) Relaxed regulatory requirements for CWF pumping (increased CWF
332 pumping), and 8) Relaxed regulatory requirements for all urban groundwater pumping (increased
333 all urban pumping). Details regarding each of these water use scenarios are provided below.

334 The business as usual scenario (scenario 5 in the Table 1) assumed no change in
335 groundwater regulations. Thus the CWF pumping remained at the maximum permitted 90 MGD
336 and all other urban pumping (industrial/commercial, mining and other public water supply)
337 remained at the average pumping rates used in the retrospective simulations. In this case all
338 projected increases in future public water demand must be met by increased surface water
339 extraction (if available), increased conservation, increased wastewater reuse, or desalination
340 capacity. For the business as usual scenario agricultural irrigation demand was estimated using
341 AFSIRS model and assuming 85% irrigation efficiency, as in the retrospective period
342 simulations. However the P and ET_0 used in the AFSIRS model were taken from the bias
343 corrected downscaled future GCM projections for both future 1 (2030-2060) and future 2 (2070-
344 2100).

345 To more clearly separate the impact of human water use versus climate change on the
346 hydrologic system, three extreme groundwater use reduction scenarios were developed. The no
347 agricultural or urban pumping scenario (scenario 1) assumed that there was no groundwater
348 pumping at all in the region. For this scenario agricultural and recreational pumping (and the
349 associated irrigation of the land surface) as well as all urban pumping (including CWF, other
350 public water supply and industrial/mining) were set to zero. For the no urban pumping scenario
351 (scenario 2) all urban pumping including CWF, other public water supplies, industrial/mining
352 was set to zero, however agricultural pumping was assumed to be the same as the business as
353 usual scenario. For the no agricultural pumping scenario (scenario 3) agricultural and
354 recreational pumping were set to zero, however all urban pumping was assumed equal to the
355 business as usual scenario.

356 The agricultural adaption scenario (scenario 4) assumed that increased irrigation
357 efficiency and/or increased use of reclaimed water reduced groundwater pumping for agricultural
358 and recreational irrigation by 40 MGD over climate driven demand (6 mm/year, ~25%). All
359 urban pumping was assumed to be the same as the business as usual scenario. The increased
360 agricultural demand scenario (scenario 6) assumed that irrigation demand increased by 40 MGD
361 over climate driven demand (6 mm/year, ~25%) due to more intensive farming on existing
362 agricultural lands (Florida Statewide Agricultural Irrigation Demand Report, 2017) and that all
363 urban pumping was the same as the business as usual scenario. The relaxed regulatory
364 requirements for CWF pumping (scenario 7) assumed an increase of CWF pumping up to 130
365 MGD (19 mm/year, ~44%) from the current 90 MGD (13 mm/year) to help meet increased
366 public water demand, and that agricultural and recreational pumping followed the business as
367 usual scenario. The relaxed regulatory requirements for all urban pumping (scenario 8) assumed
368 all urban pumping, including CWF pumping, other public water supply, industrial and mining,
369 increased by 44 %, (i.e. the same percentage increase as the CWF pumping for scenario 7) and
370 that agricultural and recreational pumping followed the business as usual scenario. These water
371 use scenarios consist of projected agricultural and urban groundwater pumping volumes that
372 represent from 0 % to 27 % of historic P-ET₀.

373 It should be noted that land use change was not considered in this study. This assumption
374 is consistent with a regional planning strategy that promotes agricultural and urban
375 intensification on existing lands, along with protection of existing conservation lands, wetlands
376 and water supplies (Barnett et al., 2007). This assumption is also consistent with the Florida
377 Statewide Agricultural Irrigation Demand Report (2017) that projects a 2% decline in
378 agricultural land area between 2015-2040, but an 8.5% increase in agricultural water use as a net
379 result of agricultural intensification and increased conservation. Future work will build on this
380 study to evaluate land use change scenarios.

381 2.8 Statistical Analysis

382 Variance-based sensitivity analysis is a global sensitivity analysis (GSA) method (Saltelli
383 et al., 2008, 2010) used to apportion the total model output variance simultaneously onto all the
384 varying input factors, and thus is preferred over the local, one factor at a time, sensitivity
385 analyses (Homma and Saltelli, 1996; Saltelli, 1999). In this research the sensitivity of projected

386 changes between future and retrospective mean monthly streamflow and groundwater levels was
387 evaluated using the variance-based GSA method described in Chang et al. (2016).

388 Using the variance-based GSA method the variance-based first order effect is expressed
389 as:

$$390 \quad V_{X_i} \left(E_{X_{\sim i}}(Y|X_i) \right) \quad (1)$$

391 Where V is the scalar model output (i.e., change in mean monthly streamflow or
392 groundwater level), and X_i are the factors causing variation in the model output (i.e. choice of
393 GCM, ET_0 method, water use scenario). The expectation operator $E_{X_{\sim i}}(Y|X_i)$ indicates that the
394 mean of Y is taken over all possible values of X except X_i (i.e., $X_{\sim i}$) while keeping X_i fixed. The
395 variance, V_{X_i} , is then taken of this quantity over all possible values of X_i . The first-order
396 sensitivity coefficient is

$$397 \quad S_i = \frac{V_{X_i}(E_{X_{\sim i}}(Y|X))}{V(Y)} \quad (2)$$

398 where $V(Y)$ the total variance of Y over all X_i . S_i is a normalized index varying between
399 0 and 1, because $V_{X_i} \left(E_{X_{\sim i}}(Y|X_i) \right)$ varies between 0 and $V(Y)$ according to the identity (Mood et
400 al., 1974):

$$401 \quad V_{X_i} \left(E_{X_{\sim i}}(Y|X_i) \right) + E_{X_i} \left(V_{X_{\sim i}}(Y|X_i) \right) = V(Y) \quad (3)$$

402 The first-order sensitivities of future changes in mean seasonal streamflow and
403 groundwater level to the choice of GCM, ET_0 estimation method, and water use scenario were
404 calculated over the full ensemble of 8 GCMs, 3 ET_0 methods and 8 water use scenarios (192
405 samples) for each future period in order to evaluate the relative contributions of each of these
406 factors on the variation among projections of future changes.

407 In addition to variance-based GSA, differences in future changes of mean projected
408 streamflow and groundwater level across GCMs and across future water use scenarios were
409 evaluated for statistical significance using Tukey's HSD (honest significant difference) test
410 (Zieyel, 1988) that is a single-step multiple statistical test (pairwise comparison). The two-
411 sample t-test was used to test for significant differences between mean projected streamflow and

412 groundwater levels resulting from future climate/water use scenarios and mean retrospective
413 streamflow and groundwater level using the business as usual water use scenario.

414 **3 Results and Discussion**

415 3.1 Global Sensitivity Analysis of Projected Changes

416 The variance-based global sensitivity analysis was conducted for both the wet season
417 (June – September) and the dry season (October – May) to evaluate the relative variation of
418 projected changes in hydrologic response attributed to the choice of GCM, choice of water use
419 scenario, and choice of ET_0 method. Tables 3 and 4 show the first order sensitivity indices of
420 changes in future streamflow and groundwater level (defined as future average seasonal
421 streamflow – retrospective average seasonal streamflow and future average seasonal
422 groundwater level – retrospective average seasonal groundwater level, respectively).

423 Change in streamflow was much more sensitive to choice of GCM than to choice of ET_0
424 method or water use scenario for all river gages, both seasons, and both future periods (Table 3).
425 For example, 94.4% of the variance of the change in wet season Hillsborough river streamflow
426 in Future 1 period (2030-2060) is attributed to differences among GCMs, 0.2% of the variance is
427 attributed to differences among ET_0 method, and 1.6% of the variance is caused by water use
428 scenario, respectively (top row Table 3). Similarly, projected changes in groundwater level were
429 generally more sensitive to the choice of GCM for all monitoring wells and both seasons.
430 However, unlike the projected changes in streamflow, changes in groundwater level were also
431 quite sensitive to the choice of water use scenario (Table 4). The higher sensitivity of
432 groundwater level to groundwater pumping is expected since the monitoring wells are
433 intentionally located near the consolidated wellfields (locations of major groundwater pumping)
434 to detect and mitigate localized impacts of water supply pumping on nearby wetlands. On the
435 other hand, the stream gages are located further from the consolidated well fields and accumulate
436 flow from a large area of the model domain. The first order sensitivity index of groundwater
437 level to water use scenario decreased in future period 2 (2070-2100) over future period 1 (2030-
438 2060), due to the increased variability of GCM precipitation projections in future 2 (2070-2100)
439 versus future 1 (2030-2060).

440 As mentioned previously Chang et al. (2016) evaluated projected changes in $P - ET_0$ over
441 the continental USA using nine GCMs, ten ET_0 estimation methods, and three RCP scenarios
442 and found that for the Southeast USA the choice of GCM and ET_0 method had approximately
443 equal influence on changes in future $P - ET_0$ throughout most of the year. Because this study
444 eliminated several ET_0 estimation methods that produced unreasonably high and low historic ET_0
445 estimates for the study region using the NLDAS-2 data, the first order sensitivity index for ET_0 is
446 significantly lower in this study than in their results. It should be noted that these results do not
447 indicate that the choice of reference ET estimation method does not affect the change in
448 streamflow or groundwater, only that the choice of reference ET estimation method is much less
449 influential than the choice of GCM or choice of water use scenario.

450 3.2 Projections of Streamflow

451 The INTB was run to compare retrospective hydrologic response to historical
452 observations and model predictions generated with the calibrated model using NLDAS-2 data, as
453 well as to future hydrologic response as a result of alternative GCMs, ET_0 methods and water use
454 scenarios. Figure 2 shows observed, NLDAS-2 and retrospective mean monthly streamflow for
455 the Hillsborough river (Fig. 2a) and Alafia river (Fig. 2b), as well as future mean monthly
456 streamflow in future 1 (2030-2060) and future 2 (2070-2100) for the business as usual water use
457 scenario using the Hargreaves ET_0 method originally used to calibrate the INTB model. The
458 boxplots represent the range of mean monthly streamflow projections over eight GCMs for the
459 business as usual water use scenario. Retrospective GCMs (blue box plots) reproduced mean
460 streamflow simulated using NLDAS-2 data quite closely for both river gages with relatively
461 small variation among GCMs. In the dry season (October-May) future 1 (red box plots) and
462 future 2 (green box plots) business as usual mean monthly streamflow values over the 8 GCMs
463 (red box plots) also showed relatively small differences with the retrospective predictions, but
464 larger variation across GCMs. However in the wet season (June through September) future mean
465 monthly streamflows for the business as usual scenario were lower than retrospective, especially
466 in future 2 (2070-2100), and showed much larger variability across GCMs.

3.3 Projections of Groundwater Level

Figure 3 shows observed, NLDAS-2 predicted, and retrospective mean monthly groundwater level for the NWH-RMP-08s (Fig. 3a), CBR-SERW-s (Fig. 3b), and STK-STARKEY-20s wells (Fig. 3c), as well as future mean monthly groundwater level in future 1 (2030-2060) and future 2 (2070-2100) for the business as usual water use scenario and the Hargreaves ET_0 method. Groundwater levels projected by retrospective GCMs showed relatively small variation across GCMs and were very similar to groundwater levels simulated using the historic NLDAS-2 data for all three wells. Although observed seasonal patterns were reproduced accurately for all wells during the retrospective period, NWH-RMP-08s retrospective groundwater level predictions were lower than observed groundwater levels throughout the year (Fig. 3a). In contrast, all CBR-SERW-s and STK-STARKEY-20s retrospective groundwater level predictions were higher than observed groundwater levels throughout the year (Figs. 3b and 3c). These deviations (which are generally less than 0.5m) are consistent with deviations between the observed data and groundwater levels simulated by the original calibrated model using the locally-observed point weather data (Guerink and Basso, 2013). The mean groundwater levels averaged over GCMs for the future period 1 (2030-2060) business as usual scenario were similar to, or slightly lower than, the mean retrospective groundwater levels; however the mean groundwater levels for future 2 (2070-2100) were significantly lower than mean groundwater levels in the retrospective period, especially in the wet season for all wells. Similar to the streamflow results variability in projected groundwater levels among GCMs was larger in future 2 (2070-2100) than in future 1 (2030-2060).

3.4 Changes in Future Surface Water Availability for Public Supply

Tampa Bay Water operates surface-water pumps on the Hillsborough and Alafia rivers to help meet public water demand. The volume of flow permitted for extraction varies daily based on maintaining sufficient in-stream flows and spring flows to protect aquatic ecosystems. In this study, the amount of water that could be withdrawn for public water supply, while meeting current environmental regulations, was analyzed to evaluate projected changes in future water availability for different GCMs and water use scenarios. Boxplots in Fig. 4a show the variation in the projected change in the mean available water that can be withdrawn from the Hillsborough river (the mean available water that can be withdrawn for future streamflow – the mean available

497 water that can be withdrawn for retrospective streamflow) over all GCMs and all ET₀ methods
498 for each water use scenario. The boxplots show large variations due to large differences in future
499 streamflow projections. All boxplots encompass both positive and negative changes for both
500 future periods, but indicate generally lower water availability in future 2 (2070-2100) than future
501 1 (2030-2060). Figure 4b compares the change in the projected mean available water that can be
502 withdrawn from the Hillsborough river over water use scenarios and ET₀ methods for each GCM.
503 While there is some variation across water use scenarios and ET₀ methods, Fig. 4b clearly shows
504 that projected changes in future surface water availability depend strongly on choice of GCM,
505 with 5 GCMs showing less surface water availability in the future regardless of water use
506 scenario. Plots for the Alafia River show very similar behavior both by water use scenario and by
507 GCM (Figure S1 in supplemental materials).

508 The differences between the mean projected changes in available water that can be
509 withdrawn from the Hillsborough and Alafia rivers for individual water use scenarios over
510 GCMs and ET₀ methods (left columns in Table 5), and for individual GCMs over water use
511 scenarios and ET₀ methods (right columns in Table 5), were evaluated for statistical significance
512 using Tukey's HSD (honest significant difference) test. The HSD test confirmed that none of the
513 differences in the mean projected change in available water for different water use scenarios
514 shown in Figure 3a were statistically significant for the Hillsborough river for either future
515 period (In Table 5 scenarios with the same alphabetic subscripts are not statistically significantly
516 different). For the Alafia river the mean projected changes in available water for the extreme
517 groundwater pumping reduction scenario was statistically significantly different from the other
518 water use scenarios in future 1 (2030 – 2060), but no statistically significant changes were
519 detected in future 2 (2070 – 2100). These results imply that due to the large variations in climate
520 projections produced by different GCMs, differences in mean projected changes in streamflow
521 projections due to differences water use scenarios and ET₀ methods cannot be reliably predicted
522 by averaging over GCMs.

523 On the other hand, many of the differences between mean projected changes in available
524 water that can be withdrawn from the Hillsborough and Alafia rivers for individual GCMs over
525 water use scenarios were statistically significant for both future periods (i.e. many of the GCMs
526 on the right side of Table 5 have different alphabetic subscripts). Two GCMs show a distinct
527 increase water availability from these rivers for public supply (GFDL-CM3 and MRI-CGCM3)

528 however, most GCMs show a decrease in water availability (BNU-ESM, GFDL-ESM2G,
529 MIROC-ESM, NorESM1-M, and BCC-CSM). These results underscore the fact that differences
530 in projections of future availability of water from these rivers for public supply are driven more
531 strongly by differences climate models than differences in future human water use scenarios or
532 ET₀ methods. Furthermore manipulating groundwater use to change the amount of available
533 surface water has a very small effect for a given climate. These results are similar to previous
534 studies (Bosshard et al., 2013; Forzieri et al., 2014; Guimberteau et al., 2013; Harding et al.,
535 2012; Kay and Davies, 2008) that showed climate models are a large source of uncertainty for
536 climate-impact projections because of the divergence of GCM projections.

537 In addition, to the HSD test, the two sample t-test was conducted to evaluate statistical
538 significance of differences between the mean available water that can be withdrawn for the
539 retrospective period and the mean available water that can be withdrawn for each future water
540 use scenario calculated over all GCMs and ET₀ methods. The two sample t-test indicated that, at
541 the 0.05 significance level, none of the future scenarios were statistically significantly different
542 from the retrospective business as usual scenario for the Hillsborough river. For the Alafia river
543 only the no pumping and no urban pumping scenarios in future 1 (2030-2060) showed significant
544 differences from the retrospective scenario in the available water that can be withdrawn from the
545 Alafia river (marked as † on the left hand columns of Table 5). In contrast most GCMs projected
546 significantly different mean available water in both future periods compared to the retrospective
547 period when averaged over water use scenarios (marked as † in right hand columns of Table 5).

548 The results that future streamflow projections are relatively insensitive to water use
549 scenarios are contrary to that of Dale et al. (2015). They used historical streamflow and climate
550 data to evaluate the impacts of anthropogenic change on streamflow and found that for an
551 irrigation intensive watershed located in an area with hot summer and limited precipitation
552 (North Central Oklahoma, U.S.) irrigation from groundwater pumping increased antecedent soil
553 moisture and played an equally important role in streamflow variability as climate change. These
554 differences are likely due to that fact that the region studied here is wetter than their study region,
555 the aquifer underlying the study region is large and productive, and land use changes were not
556 considered in this study.

557 3.5 Changes in Compliance with Groundwater Level Regulations

558 Groundwater pumping for water supply in the Tampa Bay region is regulated to maintain
559 groundwater levels that promote environmental protection by preventing dewatering of lakes and
560 wetlands near wellfields. The relative importance of water use scenario and GCM selection on
561 the change in percent of time that future groundwater levels were above the target levels (the
562 percent of the time that groundwater level is above the target level for future scenario – the
563 percent of the time that groundwater level is above the target level for retrospective scenario)
564 was evaluated for three monitoring wells. Boxplots in Fig. 5a show the change in percent of the
565 time that groundwater level was above the target level in the dry season (Oct – May) for the
566 NWH-RMP-08s well over all GCMs for each water use scenario and ET_0 methods. Tukey’s
567 HSD test showed that the two most extreme water use reduction scenarios, i.e. the no pumping
568 scenario and the no urban pumping scenario, showed a statistically significant higher percent of
569 time that groundwater is above the target level in future 1 (2030-2060) compared to the other
570 future water use scenarios for the NWH-RMP-08s well (Table 6). Furthermore the T-test showed
571 a statistically significant difference in the percent of time this well was above the target level in
572 both futures 1 (2030-2060) and 2 (2070-2100) for these two scenarios compared to the
573 retrospective scenario (marked with † in Table 6). Results for the other two wells were more
574 ambiguous with Tukey’s HSD test showing differences among several of the water use scenarios
575 in future 1 for both wells, and among several water use scenarios in future 2 for STK-
576 STARKEY-20s. The T-test for CBR-SERW-s and STK-STARKEY-20s showed statistically
577 significant differences for the two most extreme water use reduction scenarios compared to the
578 retrospective scenario both future 1 and future 2. Collectively these results confirm that future
579 compliance with groundwater levels is sensitive to water use scenario. Scenarios that assume
580 differences in CWF pumping predict statistically significant differences in future groundwater
581 compliance when averaged over possible future climates and ET_0 methods. On the other hand
582 scenarios that assume similar differences in the magnitude of agricultural pumping generally do
583 not show statistically significant differences in future groundwater compliance. These results are
584 largely explained by the concentration of CWF wells near monitoring wells versus the
585 distribution of agricultural pumping wells throughout the model domain.

586 Fig 5b indicates and Tukey’s HSD test (Table 7) confirms that the mean change in
587 percent of time that groundwater is above the target level in the monitoring wells was
588 significantly different for many GCMs in both future periods for all three wells (Figure 5 and

589 Figures S2 – S3 in the supplemental material. Two “wet” GCMs (GFDL-CM3 and MRI-
590 CGCM3) projected statistically significant increases in the mean percent of the time that
591 groundwater is above the target level for both future periods compared to the retrospective period
592 in all three wells when averaged over future water use scenario and ET_0 method (Fig. 5b and
593 marked as † in the Table 7). Three “drier” GCMs (BNU-ESM, MIROC-ESM and BCC-CSM)
594 projected statistically significant decreases in percent of the time that groundwater level is above
595 the target level compared to the retrospective period in future 2 (2070-2100) for all three wells.
596 More GCMs showed significant differences in future period 2 (2070-2100) than in future period
597 1 (2030-2060) compared to the retrospective period because the differences among climate
598 model projections increase in the later future. These results indicate that for drier future climate
599 groundwater level regulations may be difficult to achieve regardless of groundwater pumping
600 scenario, and thus may have to change with the changing climate.

601 3.6 Ability to Meet Future Water demand

602 Future water demand projections for Tampa Bay Water indicate that even with active
603 urban water conservation programs public water supply demand is expected to increase from
604 approximately 220 MGD in 2010 to approximately 278 MGD in 2045 (Tampa Bay Water Water
605 Demand Management Plan Final Report, 2013). At the present time the Tampa Bay water supply
606 system includes 90 MGD groundwater pumping permitted for the CWF, a 25 MGD desalination
607 plant and permitted water withdrawals from the Hillsborough and Alafia rivers that vary daily to
608 maintain ecologically protective in-stream flows. Scenario discovery analysis (Tariq et al., 2017)
609 was used to explore the ability of Tampa Bay Water to meet 2045 water demand with while
610 maintaining or improving existing levels of compliance with surface and groundwater
611 regulations.

612 Figure 6 presents the results of the scenario discovery analyses that evaluates which
613 climate and water use scenarios achieve these objectives in future 1 (2030-2060) using the
614 Hargreaves ET_0 method. In these analyses it was assumed that Tampa Bay Water’s desalination
615 capacity would remain at 25 MGD, surface water would be extracted at the maximum rate that
616 complied with existing regulations, and 0% (current condition), 20%, or 40% of Tampa Bay
617 Water’s public water supply (surface water, groundwater, and desalination) might be reclaimed
618 and reused to satisfy public demand. The axes in figure 6 represent the two most important

619 factors in the climate and water use scenarios that affect achievement of Tampa Bay Water's
620 goals: mean change in precipitation projected by the different GCMs and volume of agricultural
621 and urban groundwater pumping in the water use scenario. Green filled circles indicate futures
622 that meet both 2045 water demand and maintain groundwater compliance levels at or above
623 current conditions in future 1 (2030-2060). Yellow filled circles indicate futures that meet 2045
624 water demand but decrease the level of groundwater compliance. Orange filled circles indicate
625 futures that do not meet 2045 water demand but maintain groundwater compliance levels at or
626 above current conditions. Red filled circles indicate futures that do not meet 2045 water demand
627 and decrease the level of groundwater compliance. The black filled circle indicates the
628 retrospective business as usual condition.

629 Figure 6a shows that, without using reclaimed water to satisfy public water demand only
630 4 scenarios are able to meet 2045 demand and maintain or improve existing levels of compliance
631 with groundwater regulations (filled green circles on Fig 6a). These 4 scenarios assume the 2
632 wettest future climates (projected by GFDL-CM3 and MRI-CGM3) will occur and permitted
633 CWF pumping will increase from 90 MGD to 130 MGD. No other climate-water use scenarios
634 are able to meet 2045 demand without use of reclaimed water (there are no yellow filled circles
635 on Fig. 6a). In fact a significant number of the scenarios, including many that assume the
636 business as usual water use scenario, are not able to meet 2045 demand and also decrease
637 compliance groundwater regulations (red filled circles on Fig 6a).

638 Figure 6b shows that 20% of freshwater withdrawn is reclaimed and used to satisfy
639 public demand the two wettest future climates can meet 2045 demand and maintain or improve
640 existing levels of compliance with groundwater regulations for all water use scenarios. However
641 no other scenarios are able to achieve both goals. If 40% of freshwater withdrawn is reclaimed
642 and used to satisfy public demand more scenarios are able to achieve both goals. These scenarios
643 include the climate scenarios that project that at least the existing average annual rainfall will
644 occur in the future (i.e. projected change in average annual rainfall greater than or equal to zero).
645 However to meet both public water demand and maintain existing compliance with groundwater
646 regulations, scenarios that predict the same rainfall as current climate require a complete switch
647 of public water supply from groundwater to surface water sources (bottom two water use
648 scenarios in Fig 6). This would require Tampa Bay Water to significantly increase their surface
649 water storage and treatment capacity and eliminates the use of their most inexpensive water

650 source (groundwater). If groundwater regulations were relaxed, and 40% freshwater withdrawn
651 in reclaimed, 2045 demand could be met under any climate scenario (yellow circles in Fig. 6c). It
652 should be noted that the Regional Water Supply Planning (2016) reported that in 2015 only
653 about 11.5% of total freshwater withdrawn was reused in Florida. Therefore reclaiming 20% -
654 40% of freshwater withdrawn will be a significant investment.

655 **4 Conclusions**

656 It is important to evaluate possible changes in future streamflow and groundwater levels
657 to evaluate risks in water resources management and planning. This study investigated potential
658 future changes in hydrologic systems, ability to meet future water demand, and compliance with
659 water resource regulation using eight GCMs, eight human water use scenarios and three ET_0
660 methods to drive an integrated hydrologic model developed for the Tampa Bay region.
661 Variance-based sensitivity analysis showed that changes in projected streamflow were very
662 sensitive to GCM selection, but relatively insensitive ET_0 method or water use scenario. Changes
663 in projections of groundwater level were sensitive to both GCM and water use scenario, but
664 relatively insensitive to ET_0 method.

665 The eight GCMs projected diverse changes in streamflow and groundwater level, with
666 most GCMs projecting statistically significant different future streamflow and groundwater
667 levels than the current condition. Five of the 8 GCMs projected a decrease in future streamflow
668 and groundwater level in the INTB region regardless of water use scenario or ET method. None
669 of the 8 GCMs projected that 2045 water demand could be met under the business as usual water
670 use scenario. Two GCMs (GFDL-CM3 and MRI-CGCM3) predicted increased streamflow and
671 groundwater levels and an ability to meet 2045 projected water demand and maintain existing
672 levels of compliance with groundwater standards if permitted CWF pumping were increased
673 from the current 90 MGD to 130 MGD. The GCM that predicted that future annual average
674 rainfall will be approximately equal to current rainfall met 2045 demand maintained existing
675 levels of compliance with groundwater standards only for the water use scenarios that eliminated
676 CWF pumping completely and reclaimed 40% of freshwater withdrawals.

677 These results suggest that it is more likely than not that climate change will reduce the
678 availability of both surface and groundwater for public supply in the Tampa Bay Region. Current
679 regulations on water withdrawals (surface water withdrawal permit thresholds and target levels

680 in monitoring wells near lakes and wetlands) may have to adapt to future climate conditions
681 since only extreme changes human water use (i.e. dramatic increases in use of reclaimed water
682 and a complete switch from groundwater to surface water) may be able to maintain retrospective
683 hydrologic regimes and associated aquatic ecosystems and meet human water demand in the
684 future.

685 It should be noted that the findings of this study are limited by a few major assumptions.
686 For example this study used only 8 GCMs to project future climate which is a relatively small
687 number. However these 8 GCMs spanned the range of cool to warm bias and wet to dry bias
688 exhibited by 41 CMIP5 GCMs for the southeastern United States (Rupp, 2016). In addition land
689 use change was not considered in this study. Instead we assumed the increases in agricultural and
690 urban water demand were the result of intensification of water use on existing land uses. This
691 assumption is consistent with a regional planning strategy that promotes agricultural and urban
692 intensification on existing lands, along with protection of existing conservation lands, wetlands
693 and water supplies (Barnett et al., 2007). However future work should build on this study to
694 evaluate the additional impacts of potential land use change scenarios (Gupta et al., 2015; Lin et
695 al., 2015; Matheussen et al., 2000; Yan et al., 2013).

696

697 **Acknowledgments**

698 This research was supported by Tampa Bay Water and the University of Florida Water
699 Institute. We gratefully acknowledge the modeling groups participating in the Program for
700 Climate Model Diagnosis and Inter-comparison (PCMDI) for their role in making the CMIP5
701 (Coupled Model Intercomparison Project) multi-model data set available.

702

703 **References**

- 704 Aalst, M. Van, Adger, N., Arent, D., Barnett, J., Betts, R., Bilir, E., Birkmann, J., Carmin, J.,
705 Chadee, D., Challinor, A., Chatterjee, M., Cramer, W., Davidson, D., Estrada, Y., Gattuso, J.-P.,
706 Hijioka, Y., Guldberg, O. H.-, Huang, H.-Q., Insarov, G., Jones, R., Kovats, S., Lankao, P. R.,
707 Larsen, J. N., Losada, I., Marengo, J., McLean, R., Mearns, L., Mechler, R., Morton, J., Niang,
708 I., Oki, T., Olwoch, J. M., Opondo, M., Poloczanska, E., Pörtner, H.-O., Redsteer, M. H.,
709 Reisinger, A., Revi, A., Schmidt, D., Shaw, R., Solecki, W., Stone, D., Stone, J., Strzepek, K.,
710 Suarez, A., Tschakert, P., Valentini, R., Vicuna, S., Villamizar, A., Vincent, K., Warren, R.,
711 White, L., Wilbanks, T., Wong, P. P. and Yoh, G.: *Climate Change 2014: Impacts, Adaptation,*
712 *and Vulnerability, Assess. Rep. 5, (October 2013), 1–76, doi:10.1017/CBO9781107415379,*
713 *2014.*
- 714 Abatzoglou, J. T. and Brown, T. J.: A comparison of statistical downscaling methods suited for
715 wildfire applications, *Int. J. Climatol.*, 32(5), 772–780, doi:10.1002/joc.2312, 2012.
- 716 Alcamo, J., Döll, P., Henrichs, T., Kaspar, F., Lehner, B., Rösch, T. and Siebert, S.:
717 Development and testing of the WaterGAP 2 global model of water use and availability, *Hydrol.*
718 *Sci. J.*, 48(3), 317–337, doi:10.1623/hysj.48.3.317.45290, 2003.
- 719 Barnett, J., Dobshinsky, A., Choi, B., Cunningham, A., Dickens, M., Driver, J., Fan, L., Garcia,
720 J., Gibson, N., Graves, J., Henkel, M., Khedhri, S., Lai, J., Lally, J., Lewis, M., Massa, L.,
721 Melusky, A. and Ottoson, L.: *An alternative future Florida in the 21st Century 2020 2040 2060,*
722 *Orlando, Florida., 2007.*
- 723 Bentsen, M., Bethke, I., Debernard, J. B., Iversen, T., Kirkevåg, A., Seland, Ø., Drange, H.,
724 Roelandt, C., Seierstad, I. A., Hoose, C. and Kristjánsson, J. E.: The Norwegian Earth System
725 Model, NorESM1-M – Part 1: Description and basic evaluation of the physical climate, *Geosci.*
726 *Model Dev.*, 6(3), 687–720, doi:10.5194/gmd-6-687-2013, 2013.
- 727 Bicknell, B. R., Imhoff, J. C., Kittle, Jr., J. L., Jobes, T. H. and Donigian, Jr., A. S.: *Hydrological*
728 *Simulation Program-Fortran: HSPF Version 12.2 User’s Manual., 2005.*
- 729 Block, K. and Mauritsen, T.: Forcing and feedback in the MPI-ESM-LR coupled model under
730 abruptly quadrupled CO₂, *J. Adv. Model. Earth Syst.*, 5(4), 676–691, doi:10.1002/jame.20041,
731 2013.
- 732 Block, P. J., Souza Filho, F. A., Sun, L. and Kwon, H. H.: A streamflow forecasting framework
733 using multiple climate and hydrological models, *J. Am. Water Resour. Assoc.*, 45(4), 828–843,
734 doi:10.1111/j.1752-1688.2009.00327.x, 2009.
- 735 Boé, J., Terray, L., Habets, F. and Martin, E.: Statistical and dynamical downscaling of the Seine
736 basin climate for hydro-meteorological studies, *Int. J. Climatol.*, 27(12), 1643–1655,
737 doi:10.1002/joc.1602, 2007.
- 738 Bosshard, T., Carambia, M., Goergen, K., Kotlarski, S., Krahe, P., Zappa, M. and Schär, C.:
739 Quantifying uncertainty sources in an ensemble of hydrological climate-impact projections,

740 Water Resour. Res., 49(3), 1523–1536, doi:10.1029/2011WR011533, 2013.

741 Chang, J., Zhang, H., Wang, Y. and Zhu, Y.: Assessing the impact of climate variability and
742 human activities on streamflow variation, *Hydrol. Earth Syst. Sci.*, 20(4), 1547–1560,
743 doi:10.5194/hess-20-1547-2016, 2016a.

744 Chang, S.: Quantifying the relative uncertainties of changes in climate and water demand for
745 water supply planning, University of Florida., 2017.

746 Chang, S., Graham, W. D., Hwang, S. and Muñoz-Carpena, R.: Sensitivity of future continental
747 United States water deficit projections to general circulation models, the evapotranspiration
748 estimation method, and the greenhouse gas emission scenario, *Hydrol. Earth Syst. Sci.*, 20(8),
749 3245–3261, doi:10.5194/hess-20-3245-2016, 2016b.

750 Chen, J., Brissette, F. P., Chaumont, D. and Braun, M.: Finding appropriate bias correction
751 methods in downscaling precipitation for hydrologic impact studies over North America, *Water*
752 *Resour. Res.*, 49(7), 4187–4205, doi:10.1002/wrcr.20331, 2013.

753 Dale, J., Zou, C. B., Andrews, W. J., Long, J. M., Liang, Y. and Qiao, L.: Climate, water use,
754 and land surface transformation in an irrigation intensive watershed-Streamflow responses from
755 1950 through 2010, *Agric. Water Manag.*, 160, 144–152, doi:10.1016/j.agwat.2015.07.007,
756 2015.

757 Déry, S. J., Hernández-Henríquez, M. A., Burford, J. E. and Wood, E. F.: Observational
758 evidence of an intensifying hydrological cycle in northern Canada, *Geophys. Res. Lett.*, 36(13),
759 1–5, doi:10.1029/2009GL038852, 2009.

760 Diffenbaugh, N. S. and Field, C. B.: Changes in ecologically critical terrestrial climate
761 conditions., *Science* (80-.), 341(August), 486–92, doi:10.1126/science.1237123, 2013.

762 Dukes, M. D., Zotarelli, L., Liu, G. D. and Simonne, E. H.: Principles and Practices of Irrigation
763 Management for Vegetables, , 1–14, 2012.

764 Florida Statewide Agricultural Irrigation Demand Estimated Agricultural Water Demand, 2015 -
765 2040, The Balmoral Group, Winter Park, Florida, 2017.

766 Forzieri, G., Feyen, L., Rojas, R., Flörke, M., Wimmer, F. and Bianchi, A.: Ensemble projections
767 of future streamflow droughts in Europe, *Hydrol. Earth Syst. Sci.*, 18(1), 85–108,
768 doi:10.5194/hess-18-85-2014, 2014.

769 Georgakakos, A., Fleming, P., Dettinger, M., Peters-Lidard, C., Richmond, T., Reckhow, K.,
770 White, K. and Yates, D.: Ch. 3: Water Resources. *Climate Change Impacts in the United States:*
771 *The Third National Climate Assessment.*, 2014.

772 Geurink, J. S. and Basso, R.: Development, Calibration, and Evaluation of the Integrated
773 Northern Tampa Bay Hydrologic Model, Tampa Bay Water/Southwest Florida Water
774 Management District, Clearwater/Brooksville, Florida, 2013.

775 Ghosh, S. and Mujumdar, P. P.: Statistical downscaling of GCM simulations to streamflow using
776 relevance vector machine, *Adv. Water Resour.*, 31(1), 132–146,
777 doi:10.1016/j.advwatres.2007.07.005, 2008.

778 Giorgi, F. and Mearns, L.: Calculation of average, uncertainty range, and reliability of regional
779 climate changes from AOGCM simulations via the “reliability ensemble averaging”(REA)
780 method, *J. Clim.*, 15(10), 1141–1158, doi:http://dx.doi.org/10.1175/1520-
781 0442(2002)015<1141:COAURA>2.0.CO;2, 2002.

782 Green, T. R., Taniguchi, M., Kooi, H., Gurdak, J. J., Allen, D. M., Hiscock, K. M., Treidel, H.
783 and Aureli, A.: Beneath the surface of global change: Impacts of climate change on groundwater,
784 *J. Hydrol.*, 405(3–4), 532–560, doi:10.1016/j.jhydrol.2011.05.002, 2011.

785 Guimberteau, M., Ronchail, J., Espinoza, J. C., Lengaigne, M., Sultan, B., Polcher, J., Drapeau,
786 G., Guyot, J.-L., Ducharne, A. and Ciais, P.: Future changes in precipitation and impacts on
787 extreme streamflow over Amazonian sub-basins, *Environ. Res. Lett.*, 8, 014035,
788 doi:10.1088/1748-9326/8/1/014035, 2013.

789 Guo, H., Golaz, J.-C., Donner, L. J., Ginoux, P. and Hemler, R. S.: Multivariate Probability
790 Density Functions with Dynamics in the GFDL Atmospheric General Circulation Model: Global
791 Tests, *J. Clim.*, 27(5), 2087–2108, doi:10.1175/JCLI-D-13-00347.1, 2014.

792 Gupta, S. C., Kessler, A. C., Brown, M. K. and Zvomuya, F.: Climate and agricultural land use
793 change impacts on streamflow in the upper midwestern United States, *Water Resour. Res.*, 51(7),
794 5301–5317, doi:10.1002/2015WR017323, 2015.

795 Haddeland, I., Heinke, J., Biemans, H., Eisner, S., Flörke, M., Hanasaki, N., Konzmann, M.,
796 Ludwig, F., Masaki, Y., Schewe, J., Stacke, T., Tessler, Z. D., Wada, Y. and Wisser, D.: Global
797 water resources affected by human interventions and climate change, *Proc. Natl. Acad. Sci.*,
798 111(9), 3251–3256, doi:10.1073/pnas.1222475110, 2014.

799 Harbaugh, a. and McDonald, M.: User’s Documentation for MODFLOW-96, an update to the
800 U.S. Geological Survey Modular Finite-Difference Ground-Water Flow Model, Open-File
801 Report, US Geol. Surv., 96–485, 1996.

802 Harding, B. L., Wood, a. W. and Prairie, J. R.: The implications of climate change scenario
803 selection for future streamflow projection in the Upper Colorado River Basin, *Hydrol. Earth
804 Syst. Sci.*, 16(11), 3989–4007, doi:10.5194/hess-16-3989-2012, 2012.

805 Hawkins, E., Anderson, B., Diffenbaugh, N., Mahlstein, I., Betts, R., Hegerl, G., Joshi, M.,
806 Knutti, R., McNeall, D., Solomon, S., Sutton, R., Syktus, J. and Vecchi, G.: Uncertainties in the
807 timing of unprecedented climates, *Nature*, 511(7507), E3–E5, doi:10.1038/nature13523, 2014.

808 Hawkins, E. and Sutton, R.: The potential to narrow uncertainty in regional climate predictions,
809 *Bull. Am. Meteorol. Soc.*, 90(8), 1095–1107, doi:10.1175/2009BAMS2607.1, 2009.

810 Hawkins, E. and Sutton, R.: The potential to narrow uncertainty in projections of regional
811 precipitation change, *Clim. Dyn.*, 37(1–2), 407–418, doi:10.1007/s00382-010-0810-6, 2010.

- 812 Homma, T. and Saltelli, A.: Importance measures in global sensitivity analysis of nonlinear
813 models, *Reliab. Eng. Syst. Saf.*, 52(1), 1–17, doi:10.1016/0951-8320(96)00002-6, 1996.
- 814 Hwang, S. and Graham, W. D.: Development and comparative evaluation of a stochastic analog
815 method to downscale daily GCM precipitation, *Hydrol. Earth Syst. Sci.*, 17(11), 4481–4502,
816 doi:10.5194/hess-17-4481-2013, 2013.
- 817 Hwang, S. and Graham, W. D.: Assessment of Alternative Methods for Statistically
818 Downscaling Daily GCM Precipitation Outputs to Simulate Regional Streamflow, *JAWRA J.*
819 *Am. Water Resour. Assoc.*, 50(4), 1010–1032, doi:10.1111/jawr.12154, 2014.
- 820 Hwang, S., Graham, W. D., Adams, A. and Geurink, J.: Assessment of the utility of
821 dynamically-downscaled regional reanalysis data to predict streamflow in west central Florida
822 using an integrated hydrologic model, *Reg. Environ. Chang.*, 13(S1), 69–80,
823 doi:10.1007/s10113-013-0406-x, 2013.
- 824 Hwang, S., Graham, W. D., Geurink, J. S. and Adams, A.: Hydrologic implications of errors in
825 bias-corrected regional reanalysis data for west central Florida, *J. Hydrol.*, 510, 513–529,
826 doi:10.1016/j.jhydrol.2013.11.042, 2014.
- 827 Ines, A. V. M. and Hansen, J. W.: Bias correction of daily GCM rainfall for crop simulation
828 studies, *Agric. For. Meteorol.*, 138(1–4), 44–53, doi:10.1016/j.agrformet.2006.03.009, 2006.
- 829 Jackson, M. C. and Albritton, B.: 2011 Estimated Water Use Report, Brooksville, FL., 2013.
- 830 Jacobs, J. and Dukes, M.: Revision of AFSIRS crop water simulation model Summary, Palatka,
831 FL., 2007.
- 832 Ji, D., Wang, L., Feng, J., Wu, Q., Cheng, H., Zhang, Q., Yang, J., Dong, W., Dai, Y., Gong, D.,
833 Zhang, R.-H., Wang, X., Liu, J., Moore, J. C., Chen, D. and Zhou, M.: Description and basic
834 evaluation of BNU-ESM version 1, *Geosci. Model Dev. Discuss.*, 7(2), 1601–1647,
835 doi:10.5194/gmdd-7-1601-2014, 2014.
- 836 Kay, a. L. and Davies, H. N.: Calculating potential evaporation from climate model data: A
837 source of uncertainty for hydrological climate change impacts, *J. Hydrol.*, 358(3–4), 221–239,
838 doi:10.1016/j.jhydrol.2008.06.005, 2008.
- 839 Kingston, D. G., Todd, M. C., Taylor, R. G., Thompson, J. R. and Arnell, N. W.: Uncertainty in
840 the estimation of potential evapotranspiration under climate change, *Geophys. Res. Lett.*, 36(20),
841 L20403, doi:10.1029/2009GL040267, 2009.
- 842 Kløve, B., Ala-Aho, P., Bertrand, G., Gurdak, J. J., Kupfersberger, H., Kværner, J., Muotka, T.,
843 Mykrä, H., Preda, E., Rossi, P., Uvo, C. B., Velasco, E. and Pulido-Velazquez, M.: Climate
844 change impacts on groundwater and dependent ecosystems, *J. Hydrol.*, 518(PB), 250–266,
845 doi:10.1016/j.jhydrol.2013.06.037, 2014.
- 846 Koedyk, L. P. and Kingston, D. G.: Potential evapotranspiration method influence on climate
847 change impacts on river flow: a mid-latitude case study, *Hydrol. Res.*, doi:10.2166/nh.2016.152,

- 848 2016.
- 849 Langousis, A., Mamalakis, A., Deidda, R. and Marrocu, M.: Assessing the relative effectiveness
850 of statistical downscaling and distribution mapping in reproducing rainfall statistics based on
851 climate model results, *Water Resour. Res.*, 30(693), n/a-n/a, doi:doi: 10.1002/2015wr017556,
852 2015.
- 853 Lin, B., Chen, X., Yao, H., Chen, Y., Liu, M., Gao, L. and James, A.: Analyses of landuse
854 change impacts on catchment runoff using different time indicators based on SWAT model,
855 *Ecol. Indic.*, 58, 55–63, doi:10.1016/j.ecolind.2015.05.031, 2015.
- 856 Liu, M., Adam, J. C. and Hamlet, A. F.: Spatial-temporal variations of evapotranspiration and
857 runoff/ precipitation ratios responding to the changing climate in the pacific northwest during
858 1921-2006, *J. Geophys. Res. Atmos.*, 118, 380–394, doi:10.1029/2012JD018400, 2013.
- 859 Ma, Z. M., Kang, S. Z., Zhang, L., Tong, L. and Su, X. L.: Analysis of impacts of climate
860 variability and human activity on streamflow for a river basin in arid region of northwest China,
861 *J. Hydrol.*, 352(3–4), 239–249, doi:10.1016/j.jhydrol.2007.12.022, 2008.
- 862 Matheussen, B., Kirschbaum, R. L., Goodman, I. A., O’Donnell, G. M. and Lettenmaier, D. P.:
863 Effects of land cover change on streamflow in the interior Columbia River Basin (USA and
864 Canada), *Hydrol. Process.*, 14(5), 867–885, doi:10.1002/(SICI)1099-
865 1085(20000415)14:5<867::AID-HYP975>3.0.CO;2-5, 2000.
- 866 Maurer, E. P. and Hidalgo, H. G.: Utility of daily vs. monthly large-scale climate data: an
867 intercomparison of two statistical downscaling methods, *Hydrol. Earth Syst. Sci.*, 12(2), 551–
868 563, doi:10.5194/hess-12-551-2008, 2008.
- 869 Maurer, E. P., Hidalgo, H. G., Das, T., Dettinger, M. D. and Cayan, D. R.: The utility of daily
870 large-scale climate data in the assessment of climate change impacts on daily streamflow in
871 California, *Hydrol. Earth Syst. Sci.*, 14(6), 1125–1138, doi:10.5194/hess-14-1125-2010, 2010.
- 872 McAfee, S. A.: Methodological differences in projected potential evapotranspiration, *Clim.*
873 *Change*, 120(4), 915–930, doi:10.1007/s10584-013-0864-7, 2013.
- 874 Milliman, J. D., Farnsworth, K. L., Jones, P. D., Xu, K. H. and Smith, L. C.: Climatic and
875 anthropogenic factors affecting river discharge to the global ocean, 1951-2000, *Glob. Planet.*
876 *Change*, 62(3–4), 187–194, doi:10.1016/j.gloplacha.2008.03.001, 2008.
- 877 Mood, A. M., Graybill, F. A. and Boes, D. C.: Introduction to theory of statistics, McGraw-Hill,
878 Inc., 1974.
- 879 Muerth, M. J., Gauvin St-Denis, B., Ricard, S., Velázquez, J. A., Schmid, J., Minville, M., Caya,
880 D., Chaumont, D., Ludwig, R. and Turcotte, R.: On the need for bias correction in regional
881 climate scenarios to assess climate change impacts on river runoff, *Hydrol. Earth Syst. Sci.*,
882 17(3), 1189–1204, doi:10.5194/hess-17-1189-2013, 2013.
- 883 Murray, S. J., Foster, P. N. and Prentice, I. C.: Future global water resources with respect to

- 884 climate change and water withdrawals as estimated by a dynamic global vegetation model, *J.*
885 *Hydrol.*, 448–449, 14–29, doi:10.1016/j.jhydrol.2012.02.044, 2012.
- 886 Patterson, L. A., Lutz, B. and Doyle, M. W.: Climate and direct human contributions to changes
887 in mean annual streamflow in the South Atlantic , USA, *Water Resour. Res.*, 49(August), 7278–
888 7291, doi:10.1002/2013WR014618, 2013.
- 889 Quintana Seguí, P., Ribes, A., Martin, E., Habets, F. and Boé, J.: Comparison of three
890 downscaling methods in simulating the impact of climate change on the hydrology of
891 Mediterranean basins, *J. Hydrol.*, 383(1–2), 111–124, doi:10.1016/j.jhydrol.2009.09.050, 2010.
- 892 Regional Water Supply Planning, Florida Department of Environmental Protection, Tallahassee,
893 Florida., 2016.
- 894 Saltelli, A.: Sensitivity analysis: Could better methods be used?, *J. Geophys. Res.*, 104(D3),
895 3789–3793, doi:10.1029/1998JD100042, 1999.
- 896 Saltelli, A., Annoni, P., Azzini, I., Campolongo, F., Ratto, M. and Tarantola, S.: Variance based
897 sensitivity analysis of model output. Design and estimator for the total sensitivity index, *Comput.*
898 *Phys. Commun.*, 181(2), 259–270, doi:10.1016/j.cpc.2009.09.018, 2010.
- 899 Saltelli, A., Ratto, M., Andres, T., Campolongo, F., Cariboni, J., Gatelli, D., Saisana, M. and
900 Tarantola, S.: *Global sensitivity analysis: the primer*, John Wiley & Sons, Inc., 2008.
- 901 Siriwardena, L., Finlayson, B. L. and McMahon, T. A.: The impact of land use change on
902 catchment hydrology in large catchments: The Comet River, Central Queensland, Australia, *J.*
903 *Hydrol.*, 326(1–4), 199–214, doi:10.1016/j.jhydrol.2005.10.030, 2006.
- 904 Smajstrla, A. G.: *Technical Manual: Agricultural field scale irrigation requirements simulation*
905 *(AFSIRS) model, Version 5.5*, Gainesville, FL., 1990.
- 906 Solomon, S., Qin, D., Manning, M., Alley, R. B., Berntsen, T., Bindoff, N. L., Chen, Z.,
907 Chidthaisong, A., Gregory, J. M., Hegerl, G. C., Heimann, M., Hewitson, B., Hoskins, B. J.,
908 Joos, F., Jouzel, J., Kattsov, V., Lohmann, U., Matsuno, T., Molina, M., Nicholls, N., Overpeck,
909 J., Raga, G., Ramaswamy, V., Ren, J., Rusticucci, M., Somerville, R., Stocker, T. F., Whetton,
910 P., Wood, D. and Wratt, R. a: *Climate Change 2007 The Physical Science Basis Contribution of*
911 *Working Group I to the Fourth Assessment Report of the Intergovernmental Panel on Climate*
912 *Change, Ipcw Wg1*, 23–78, doi:10.1007/s10894-008-9162-1, 2007.
- 913 Stoll, S., Hendricks Franssen, H. J., Butts, M. and Kinzelbach, W.: Analysis of the impact of
914 climate change on groundwater related hydrological fluxes: a multi-model approach including
915 different downscaling methods, *Hydrol. Earth Syst. Sci.*, 15(1), 21–38, doi:10.5194/hess-15-21-
916 2011, 2011.
- 917 Tampa Bay Water Water Demand Management Plan Final Report., Hazen and Sawyer, 2013.
- 918 Tan, X. and Gan, T. Y.: Contribution of human and climate change impacts to changes in
919 streamflow of Canada., *Sci. Rep.*, 5, 17767, doi:10.1038/srep17767, 2015.

920 Tariq, A., Lempert, R. J., Riverson, J., Schwartz, M. and Berg, N.: A climate stress test of Los
921 Angeles' water quality plans, *Clim. Change*, 144(4), 625–639, doi:10.1007/s10584-017-2062-5,
922 2017.

923 Taylor, K. E., Stouffer, R. J. and Meehl, G. A.: An Overview of CMIP5 and the Experiment
924 Design, *Bull. Am. Meteorol. Soc.*, 93(4), 485–498, doi:10.1175/BAMS-D-11-00094.1, 2012.

925 Teutschbein, C. and Seibert, J.: Bias correction of regional climate model simulations for
926 hydrological climate-change impact studies: Review and evaluation of different methods, *J.*
927 *Hydrol.*, 456–457, 12–29, doi:10.1016/j.jhydrol.2012.05.052, 2012.

928 Thompson, J. R., Green, A. J. and Kingston, D. G.: Potential evapotranspiration-related
929 uncertainty in climate change impacts on river flow: An assessment for the Mekong River basin,
930 *J. Hydrol.*, 510, 259–279, doi:10.1016/j.jhydrol.2013.12.010, 2014.

931 Tihansky, A. B.: Sinkholes, west-central Florida, U.S. Geological Survey, Tampa, FL., 1999.

932 Tihansky, A. B. and Knochenmus, L. a.: Karst features and hydrogeology in west-central Florida
933 — A field perspective, *US Geol. Surv. Karst Interes. Gr. Proc. Water-Resources Investig. Rep.*
934 *01-4011*, 198–211 [online] Available from:
935 http://water.usgs.gov/ogw/karst/kigconference/abt_karstfeatures.htm, 2001.

936 Vano, J. a. and Lettenmaier, D. P.: A sensitivity-based approach to evaluating future changes in
937 Colorado River discharge, *Clim. Change*, 122(4), 621–634, doi:10.1007/s10584-013-1023-x,
938 2013.

939 Vorosmarty, C. J., Green, P., Salisbury, J. and Lammers, R. B.: Global Water Resources:
940 Vulnerability from Climate Change and Population Growth, *Science (80-.)*, 289(5477), 284–
941 288, doi:10.1126/science.289.5477.284, 2000.

942 Walsh, J., Wuebbles, D., Hayhoe, K., Kossin, J., Stephens, G., Thorne, P., Vose, R., Wehner, M.,
943 Willis, J., Anderson, D., Doney, S., Feely, R., Hennon, P., Kharin, V., Knutson, T., Landerer, F.,
944 Lenton, T., Kennedy, J. and Somerville, R.: Ch. 2: Our Changing Climate. *Climate Change*
945 *Impacts in the United States: The Third National Climate Assessment.*, 2014.

946 Wang, D. and Hejazi, M.: Quantifying the relative contribution of the climate and direct human
947 impacts on mean annual streamflow in the contiguous United States, *Water Resour. Res.*, 47(10),
948 n/a-n/a, doi:10.1029/2010WR010283, 2011.

949 Wang, W., Xing, W. and Shao, Q.: How large are uncertainties in future projection of reference
950 evapotranspiration through different approaches?, *J. Hydrol.*, 524, 696–700,
951 doi:10.1016/j.jhydrol.2015.03.033, 2015.

952 Watanabe, S., Hajima, T., Sudo, K., Nagashima, T., Takemura, T., Okajima, H., Nozawa, T.,
953 Kawase, H., Abe, M., Yokohata, T., Ise, T., Sato, H., Kato, E., Takata, K., Emori, S. and
954 Kawamiya, M.: MIROC-ESM: model description and basic results of CMIP5-20c3m
955 experiments, *Geosci. Model Dev. Discuss.*, 4(2), 1063–1128, doi:10.5194/gmdd-4-1063-2011,
956 2011.

957 Wood, A. W., Maurer, E. P., Kumar, A. and Lettenmaier, D. P.: Long-range experimental
958 hydrologic forecasting for the eastern United States, *J. Geophys. Res.*, 107(D20), 4429,
959 doi:10.1029/2001JD000659, 2002.

960 Xian, G., Crane, M. and Su, J.: An analysis of urban development and its environmental impact
961 on the Tampa Bay watershed, *J. Environ. Manage.*, 85(4), 965–976,
962 doi:10.1016/j.jenvman.2006.11.012, 2007.

963 Xiao-Ge, X., Tong-Wen, W., Jiang-Long, L., Zai-Zhi, W., Wei-Ping, L. and Fang-Hua, W.: How
964 well does BCC_CSM1. 1 reproduce the 20th century climate change over China?, *Atmos.*
965 *Ocean. Sci. Lett.*, 6(1), 21–26 [online] Available from:
966 <http://159.226.119.58/aosl/CN/article/downloadArticleFile.do?attachType=PDF&id=332>
967 (Accessed 12 January 2015), 2013.

968 Yan, B., Fang, N. F., Zhang, P. C. and Shi, Z. H.: Impacts of land use change on watershed
969 streamflow and sediment yield: An assessment using hydrologic modelling and partial least
970 squares regression, *J. Hydrol.*, 484, 26–37, doi:10.1016/j.jhydrol.2013.01.008, 2013.

971 Ye, X., Zhang, Q., Liu, J., Li, X. and Xu, C.: Distinguishing the relative impacts of climate
972 change and human activities on variation of streamflow in the Poyang Lake catchment, China, *J.*
973 *Hydrol.*, 494, 83–95, doi:10.1016/j.jhydrol.2013.04.036, 2013.

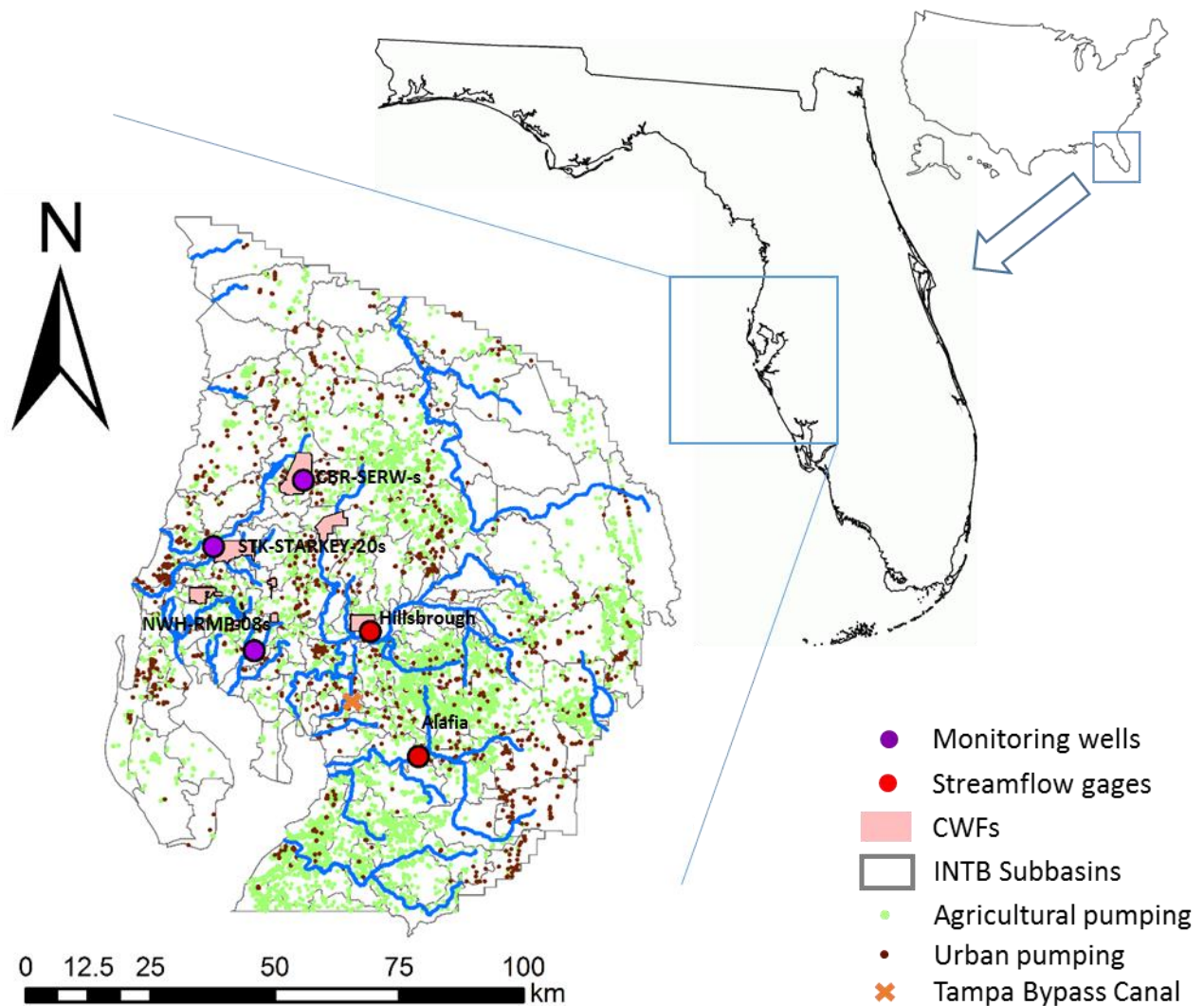
974 Yukimoto, S., Adachi, Y., Hosaka, M., Sakami, T., Yoshimura, H., Hirabara, M., Tanaka, T. Y.,
975 Shindo, E., Tsujino, H., Deushi, M., Mizuta, R., Yabu, S., Obata, A., Nakano, H., Koshiro, T.,
976 Ose, T. and Kitoh, A.: A New Global Climate Model of the Meteorological Research Institute:
977 MRI-CGCM3 -Model Description and Basic Performance-, *J. Meteorol. Soc. Japan*, 90A, 23–
978 64, doi:10.2151/jmsj.2012-A02, 2012.

979 Zhang, F. and Georgakakos, A. P.: Joint variable spatial downscaling, *Clim. Change*, 111(3–4),
980 945–972, doi:10.1007/s10584-011-0167-9, 2012.

981 Zheng, H., Zhang, L., Zhu, R., Liu, C., Sato, Y. and Fukushima, Y.: Responses of streamflow to
982 climate and land surface change in the headwaters of the Yellow River Basin, *Water Resour.*
983 *Res.*, 45(7), doi:10.1029/2007WR006665, 2009.

984 Zieyel, E. R.: *The Collected Works of John W. Tukey*, *Technometrics*, 30(3), 363–363,
985 doi:10.1080/00401706.1988.10488428, 1988.

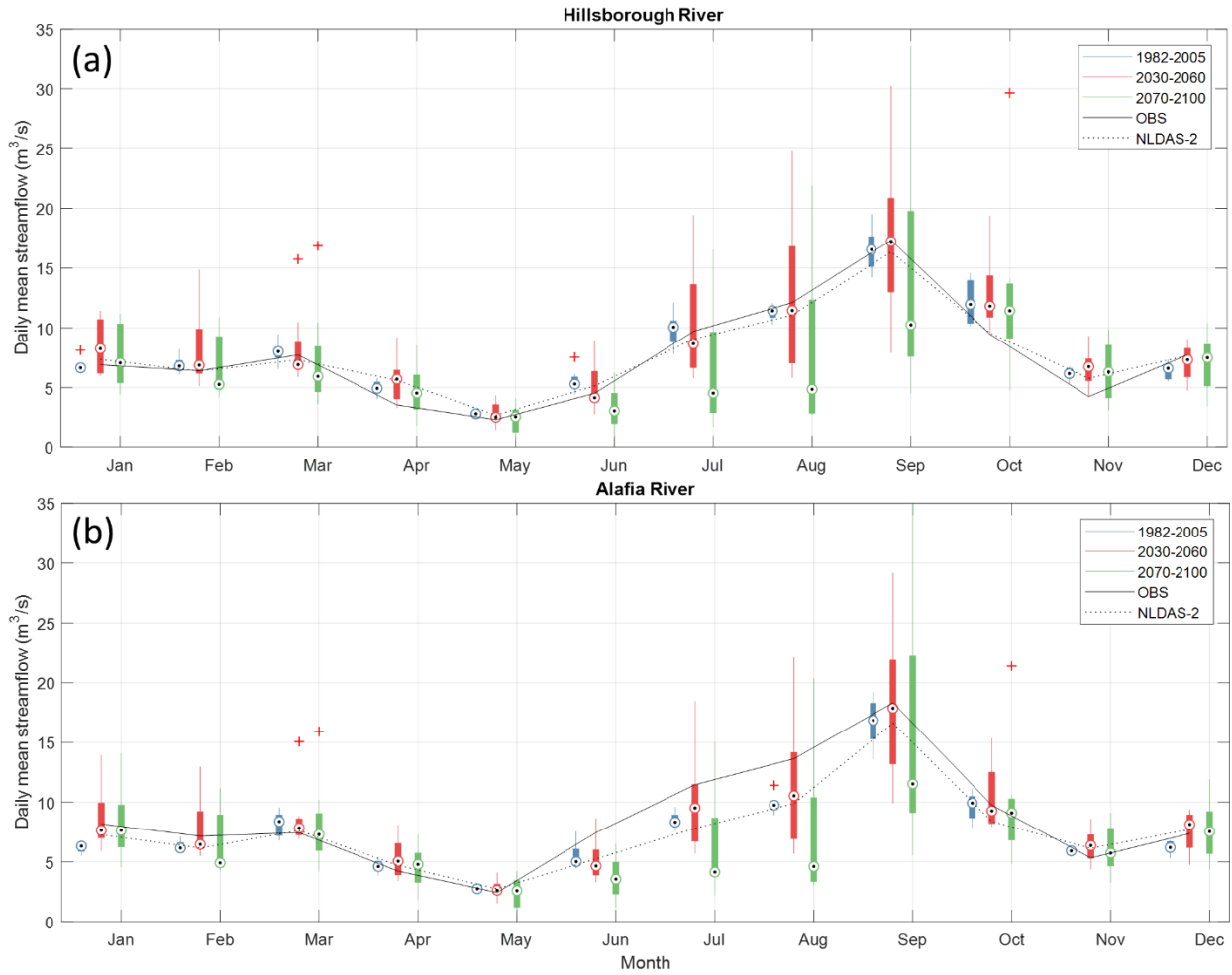
986



987

988 Figure 1. Study region showing the INTB model domain and locations of agricultural, industrial
 989 and public water supply wells, the Tampa Bay Waters Consolidated Wellfields (CWF), two
 990 streamflow locations where water is withdrawn for public supply, the Tampa Bay Bypass Canal,
 991 and three monitoring wells near Tampa Bay Water’s CWFs that are used to evaluate compliance
 992 with groundwater level regulations.

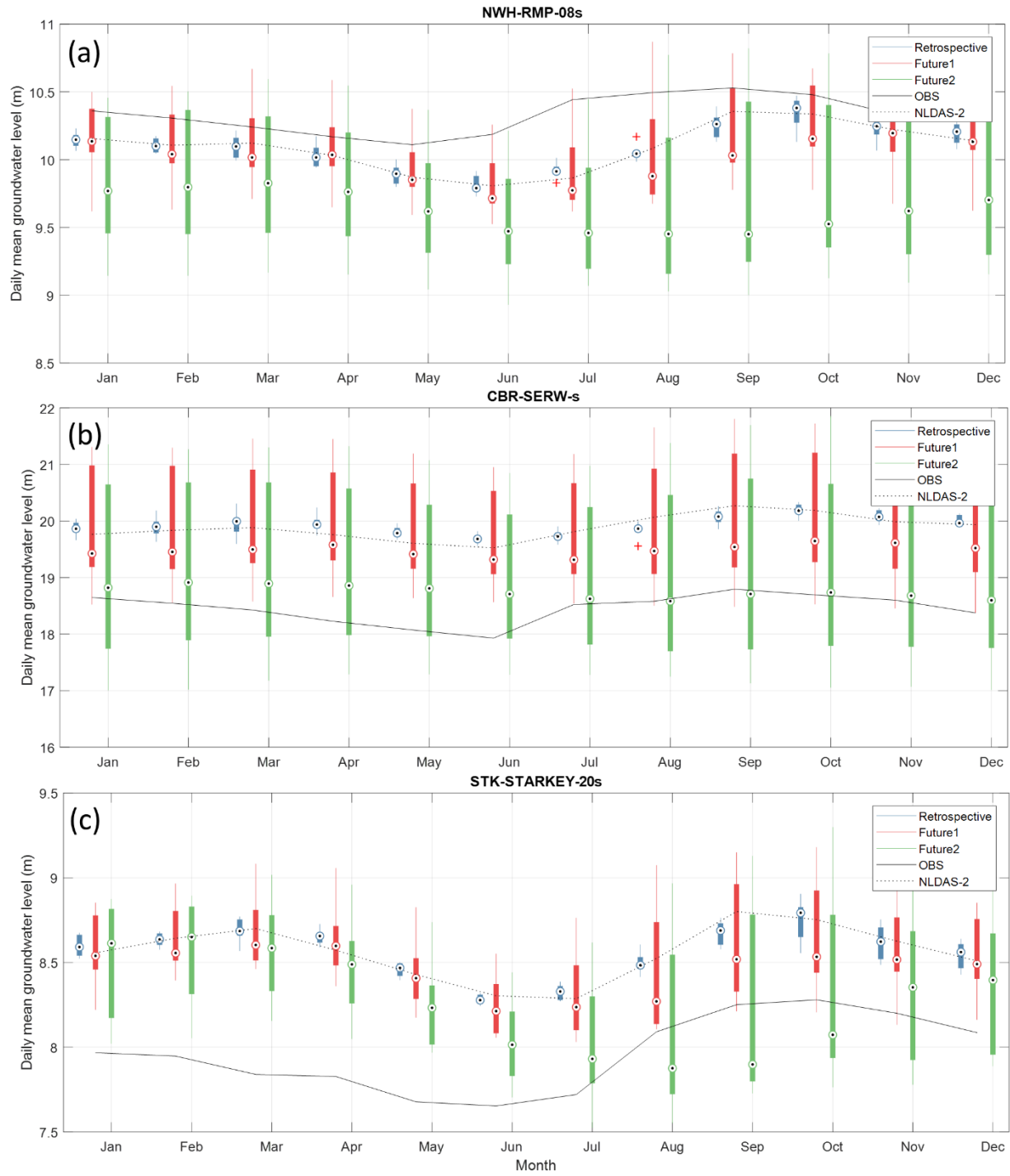
993



994

995 Figure 2. Mean monthly streamflow for the Hillsborough river (top) and Alafia river (bottom) for
 996 business as usual scenario water use and Hargreaves ET_0 method. Box plots indicate range of
 997 prediction over the 8 GCMs.

998

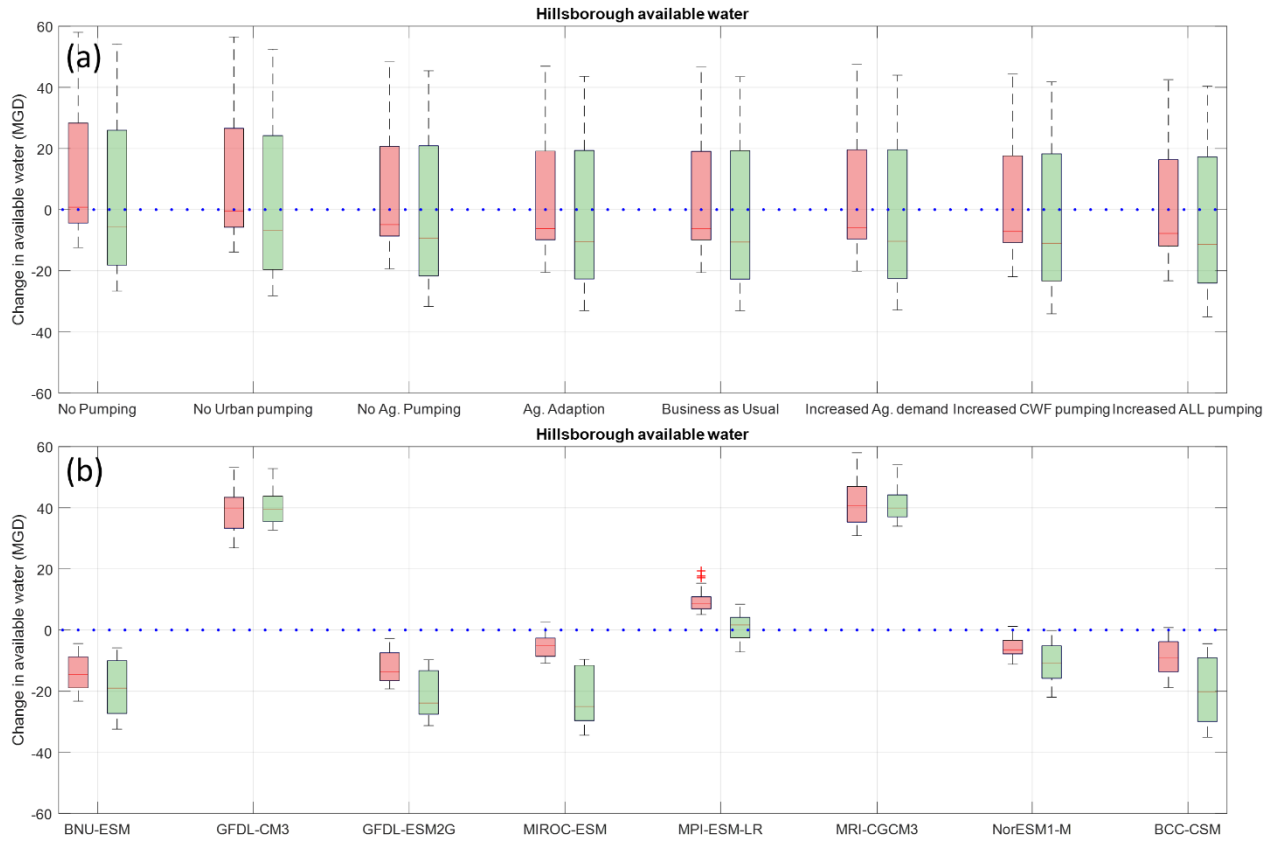


999

1000 Figure 3. Mean monthly groundwater level for the NWH-RMP-08s (top), CBR-SERW-s
 1001 (middle) and STK-STARKEY-20s (bottom) for business as usual water use scenario and
 1002 Hargreaves ET_0 method. Box plots indicate range of prediction over the 8 GCMs.

1003

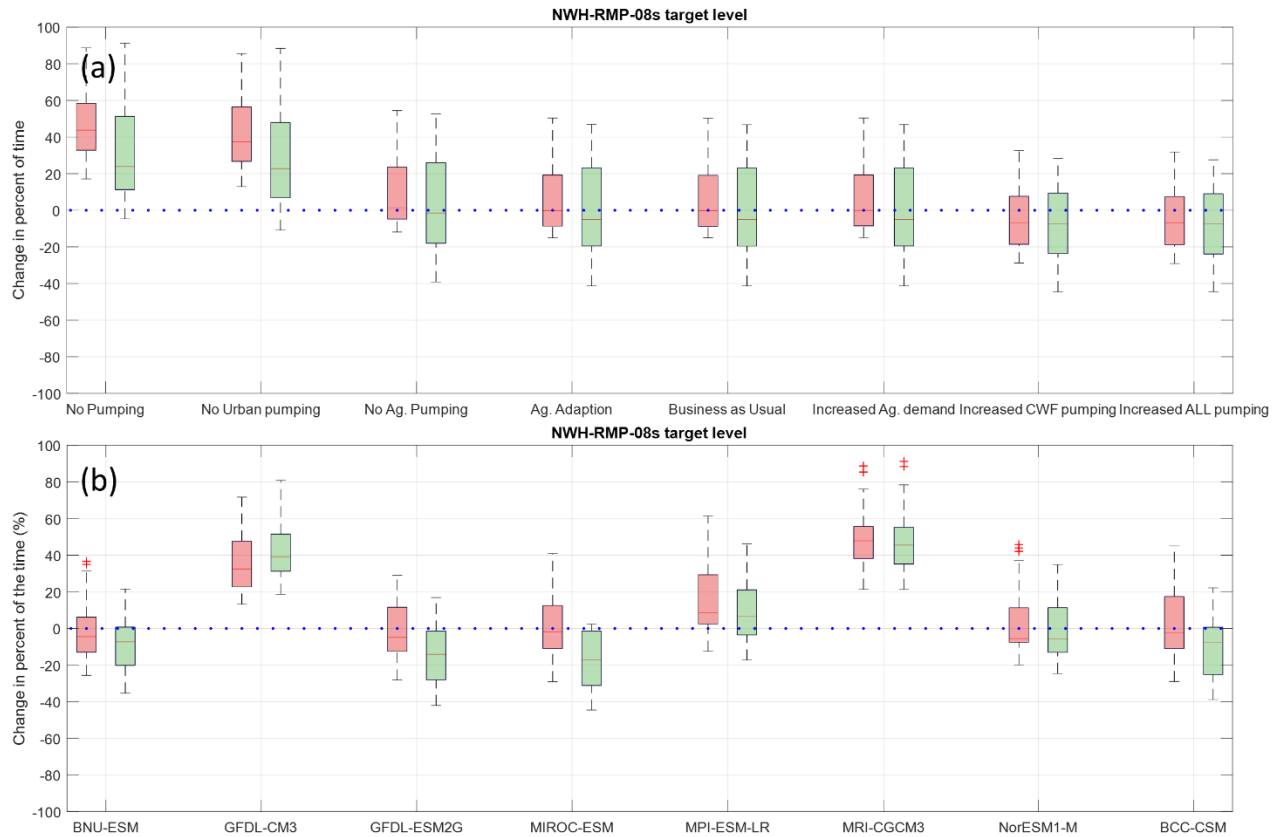
1004



1005

1006 Figure 4. The change in amount of available water can be withdrawn from Hillsborough river by
 1007 (a) different water use scenarios over GCMs and ET₀ methods and by (b) different GCMs over
 1008 water use scenarios and ET₀ methods.

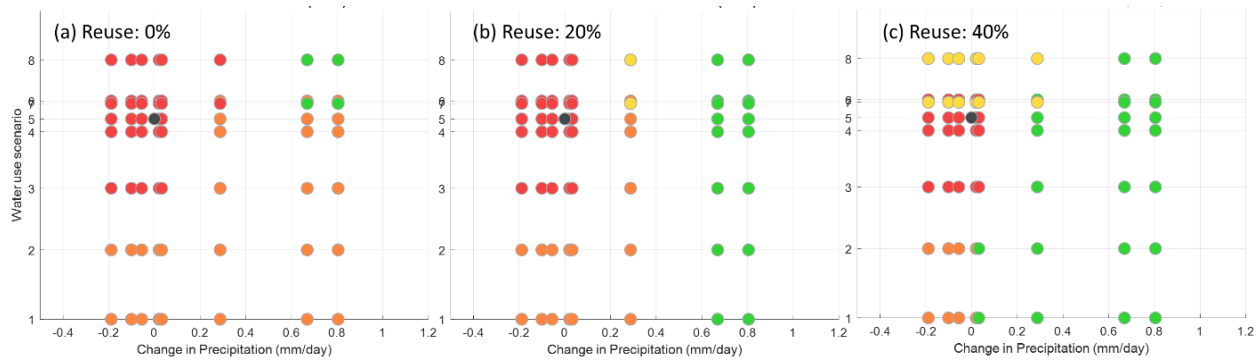
1009



1010

1011 Figure 5. The change in the percent of the time that groundwater level is above the target level
 1012 for NWH-RMP-08s well by (a) different water use scenarios over GCMs and ET_0 methods and
 1013 by (b) different GCMs over water use scenarios and ET_0 methods.

1014



1015

1016

1017

1018

1019

1020

1021

1022

1023

1024

1025

1026

Figure 6. Scatterplot of futures in which the Tampa Bay Water meets 2045 water demands and maintains or improves compliance with groundwater regulations in future 1 (2030-2060) assuming 0%, 20% and 40% of freshwater withdrawn is reclaimed and reused to satisfy urban demand. Green filled circles indicate futures that meet both 2045 water demand and maintain groundwater compliance levels at or above current conditions. Yellow filled circles indicate futures that meet 2045 water demand but decrease the level of groundwater compliance. Orange filled circles indicate futures that do not meet 2045 water demand but maintain groundwater compliance levels at or above current conditions. Red filled circles indicate futures that do not meet 2045 water demand and decrease the level of groundwater compliance. The black filled circle indicates the retrospective business as usual condition.

1027 **Table 1.** Description of the CMIP5 models used in this study.

Model	Institute (country)	Resolutions	Calendar	ΔET_0 (mm/yr)*				Reference
				Δ Precipitation (mm/yr)*				
				2030-2060	2070-2100	2030-2060	2070-2100	
(1) BNU-ESM	College of Global Change and Earth System Science, Beijing Normal University (China)	2.8° lat × 2.8° lon	No leap	-68.9	-57.1	93.3	273.5	Ji et al. (2014)
(2) GFDL-CM3	NOAA/Geophysical Fluid Dynamics Laboratory (USA)	2.0° lat × 2.5° lon	No leap	293.6	400.0	133.1	351.5	Guo et al. (2014)
(3) GFDL-ESM2G	NOAA/Geophysical Fluid Dynamics Laboratory (USA)	2.0° lat × 2.5° lon	No leap	-36.8	-134.6	56.2	133.5	Taylor et al. (2012)
(4) MIROC-ESM	Atmosphere and Ocean Research Institute, National Institute for Environmental Studies, and Japan Agency for Marine-Earth Science and Technology (Japan)	2.8° lat × 2.8° lon	Leap year	7.5	-153.9	99.9	240.8	Watanabe et al. (2011)
(5) MPI-ESM-LR	Max Planck Institute for Meteorology (Germany)	1.87° lat × 1.87° lon	Leap year	105.1	77.8	81.8	230.9	Block and Mauritsen (2013)
(6) MRI-CGCM3	Meteorological Research Institute (Japan)	1.12° lat × 1.12° lon	Leap year	244.2	281.2	24.4	122.1	Yukimoto et al. (2012)
(7) NorESM1-M	Norwegian Climate Centre (Norway)	1.9° lat × 2.5° lon	No leap	11.6	3.0	137.7	324.6	Bentsen et al. (2013)
(8) BCC-CSM1.1	Beijing Climate Center (China)	2.8° lat × 2.8° lon	No leap	-20.4	-117.5	118.1	303.6	Xiao-Ge et al. (2013)

* Change in precipitation (or ET_0) is defined as average of future period minus average of retrospective period.

1029 **Table 2.** Future scenario summary

Scenario Name	Scenario Number	Irrigation Applied to Land Surface	Agricultural pumping	Urban pumping
No pumping	1	No	No	No
No urban pumping	2	AFSIRS*	85 % efficiency	No
No agricultural pumping	3	No	No	RETRO ** CWF 13 mm/yr, Total 51 mm/yr
Agricultural adaption	4	AFSIRS	85 % efficiency Groundwater pumping offset by 6 mm/yr	RETRO CWF 13 mm/yr, Total 51 mm/yr
Business as Usual	5	AFSIRS	85 % efficiency	RETRO CWF 13 mm/yr, Total 51 mm/yr
Increased agricultural demand	6	Increased by 6 mm/yr	85 % efficiency	RETRO CWF 13 mm/yr, Total 51 mm/yr
Relaxed regulatory requirements for urban pumping	7	AFSIRS	85 % efficiency	Increase CWF by 6 mm/yr to 19 mm/yr CWF 19 mm/yr, Total 57 mm/yr
Relaxed regulatory requirements for all pumping	8	AFSIRS	85 % efficiency	Increase all urban pumping by 130/90 CWF 19 mm/yr, Total 74 mm/yr

* AFSIRS: climate driven irrigation water demand estimated by AFSIRS model using GCMs.

** RETRO: groundwater pumping in the future will be equal to retrospective groundwater pumping.

1030

1031

1032 **Table 3.** The first order sensitivity index of change in streamflow (future – retrospective period).

River gage	Season	Period	GCM	ET ₀	Water use scenario
Hillsborough	Wet season	2030-2060	0.944	0.002	0.016
		2070-2100	0.940	0.041	0.006
	Dry season	2030-2060	0.948	0.012	0.029
		2070-2100	0.961	0.001	0.018
Alafia	Wet season	2030-2060	0.928	0.010	0.031
		2070-2100	0.952	0.021	0.012
	Dry season	2030-2060	0.876	0.012	0.072
		2070-2100	0.927	0.001	0.068

1033

1034

1035 **Table 4.** The first order sensitivity index of change in groundwater level (future – retrospective
 1036 period).

Monitoring well	Season	Period	GCM	ET ₀	Water use scenario
NWH-RMP-08s	Wet season	2030-2060	0.442	0.005	0.501
		2070-2100	0.576	0.004	0.278
	Dry season	2030-2060	0.475	0.007	0.435
		2070-2100	0.550	0.002	0.288
CBR-SERW-s	Wet season	2030-2060	0.656	0.000	0.214
		2070-2100	0.755	0.002	0.143
	Dry season	2030-2060	0.639	0.001	0.221
		2070-2100	0.747	0.002	0.146
STK-STARKEY-20s	Wet season	2030-2060	0.604	0.000	0.325
		2070-2100	0.718	0.004	0.198
	Dry season	2030-2060	0.584	0.002	0.330
		2070-2100	0.707	0.001	0.200

1037

1038

1039 **Table 5.** The results of Tukey’s HSD test of mean change in amount of available water (MGD)
 1040 that can be withdrawn from Hillsborough river or Alafia river for each water use scenario over
 1041 GCM and ET₀ method, or for each GCM over water use scenario and ET₀ method (Comparison
 1042 of all possible pairs of means).

By human water use scenario	Hillsborough		Alafia		By GCM	Hillsborough		Alafia	
	2030- 2060	2070- 2100	2030- 2060	2070- 2100		2030- 2060	2070- 2100	2030- 2060	2070- 2100
	mean	mean	mean	mean		mean	mean	mean	mean
No Pumping	11.63 a	3.88 a	4.89 a [†]	2.28 a	BNU-ESM	-14.03 e [†]	-18.76 d [†]	-4.25 d [†]	-5.89 c [†]
No Urban Pumping	10.10 a	2.61 a	4.00 a [†]	1.45 a	GFDL-CM3	39.20 a [†]	40.27 a [†]	8.16 a [†]	9.11 a [†]
No Ag. Pumping	5.57 a	-1.21 a	1.48 a	-0.99 a	GFDL-ESM2G	-12.24 de [†]	-21.68 d [†]	-1.84 cd	-5.70 c [†]
Ag. Adaption	4.22 a	-2.54 a	0.85 ab	-1.60 a	MIROC- ESM2G	-5.01 c	-22.31 d [†]	-0.09 c	-6.26 c [†]
Business as Usual	4.16 a	-2.59 a	0.82 ab	-1.63 a	MPI-ESM-LR	9.71 b [†]	1.07 b	2.01 b	-0.56 b
Increased Ag. Demand	4.56 a	-2.27 a	1.00 ab	-1.47 a	MRI-CGCM3	41.64 a [†]	41.34 a [†]	10.64 a [†]	10.46 a [†]
Increased CWF pumping	2.90 a	-3.66 a	0.81 ab	-1.64 a	NorESM1-M	-5.58 c	-10.71 c [†]	0.78 bc	-2.21 c [†]
Increased All Pumping	1.72 a	-4.65 a	-0.43 b	-2.73 a	BCC-CSM	-8.84 cd [†]	-19.67 d [†]	-1.98 cd	-5.28 c [†]

Means with different subscripts were significantly different in Tukey’s HSD test.

[†]: The results were significantly different than retrospective BAU by two sample t-test at the 0.05 significance level.

1043

1044

1045 **Table 6.** The results of Tukey’s HSD test of mean change in the percent of the time that
 1046 groundwater level is above the target level for monitoring wells over all GCMs and ET₀ methods
 1047 for each water use scenario (Comparison of all possible pairs of means).

By human water use						
scenario	NWH-RMP-08s		CBR-SERW-s		STK-STARKEY-20s	
	2030-2060 mean	2070-2100 mean	2030-2060 mean	2070-2100 mean	2030-2060 mean	2070-2100 mean
No Pumping	46.04 a [†]	32.21 b [†]	31.93 a [†]	22.79 a [†]	27.87 a [†]	18.00 a [†]
No Urban Pumping	41.17 a [†]	28.36 a [†]	31.40 ab [†]	22.45 a [†]	26.91 ab [†]	17.22 ab [†]
No Ag. Pumping	10.28 b	3.69 b	11.00 c [†]	7.21 a	3.92 a [†]	-2.04 bc
Ag. Adaption	6.66 b	0.88 b	10.76 c	7.06 a	3.15 ab	-2.79 c
Business as usual	6.55 b	0.81 b	10.73 c	7.04 a	3.12 ab	-2.80 c
Increased Ag. Demand	6.70 b	0.89 b	11.14 bc [†]	7.32 a	3.21 ab	-2.73 c
Increased CWF pumping	-4.25 b	-7.81 b	5.23 c	3.01 a	-4.31 b	-9.05 c
Increased All Pumping	-4.64 b	-8.13 b	4.08 c	1.93 a	-6.07 b	-10.52 c [†]

1048 Means with different subscripts were significantly different in Tukey’s HSD test.

1049 [†]: The results were significantly different than retrospective BAU by two sample t-test at the 0.05 significance level.

1050

1051 **Table 7.** The results of Tukey’s HSD test of mean change in percent of the time that
 1052 groundwater level is above the target level for monitoring wells over all water use scenarios and
 1053 ET₀ methods for each GCM (Comparison of all possible pairs of means).

By GCM	NWH-RMP-08s		CBR-SERW-s		STK-STARKEY-20s	
	2030-2060 mean	2070-2100 mean	2030-2060 mean	2070-2100 mean	2030-2060 mean	2070-2100 mean
BNU-ESM	-6.39 c	-18.59 bc [†]	-12.08 c [†]	-16.66 c [†]	-12.55 d	-18.30 def [†]
GFDL-CM3	32.35 ab [†]	39.44 a [†]	48.12 ab [†]	56.39 a [†]	19.56 ab [†]	24.50 ab [†]
GFDL-ESM2G	-3.22 bc	-18.93 bc	-7.58 c	-22.84 c [†]	-12.96 d [†]	-16.40 cde [†]
MIROC-ESM	-4.83 c	-35.79 c [†]	4.97 c	-15.52 c [†]	-12.96 d [†]	-39.01 f [†]
MPI-ESM-LR	11.26 abc	3.41 b	29.83 b [†]	14.15 b [†]	12.02 abc [†]	4.06 bc
MRI-CGCM3	41.27 a [†]	39.67 a [†]	62.87 a [†]	56.38 a [†]	34.45 a [†]	26.16 a [†]
NorESM1-M	3.84 bc	-3.47 b	1.18 c	-8.40 c	2.31 bcd	0.17 cd
BCC-CSM	-2.38 bc	-25.30 bc [†]	1.17 c	-12.51 c [†]	-11.45 cd	-28.99 ef [†]

1054 Means with different subscripts were significantly different in Tukey’s HSD test.

1055 [†]: The results were significantly different than retrospective BAU by two sample t-test at the 0.05 significance level.

1056