Hydrol. Earth Syst. Sci. Discuss., https://doi.org/10.5194/hess-2018-32
Manuscript under review for journal Hydrol. Earth Syst. Sci.
Discussion started: 13 February 2018

Hydrology and Earth System Sciences
Open Access
Discussions

A new probability density function for spatial distribution of soil water storage capacity leads to SCS curve number method

Dingbao Wang
Department of Civil, Environmental, and Construction Engineering, University of Central Florida, Orlando, Florida, USA
Correspondence to: D. Wang, dingbao.wang@ucf.edu

Abstract

Following the Budyko framework, soil wetting ratio (the ratio between soil wetting and precipitation) as a function of soil storage index (the ratio between soil wetting capacity and precipitation) is derived from the SCS-CN method and the VIC type of model. For the SCS-CN method, soil wetting ratio approaches one when soil storage index approaches infinity, due to the limitation of the SCS-CN method in which the initial soil moisture condition is not explicitly represented. However, for the VIC type of model, soil wetting ratio equals soil storage index when soil storage index is lower than a certain value, due to the finite upper bound of the power distribution function of storage capacity. In this paper, a new distribution function, supported on a semi-infinite interval $x \in [0, \infty)$, is proposed for describing the spatial distribution of storage capacity. From this new distribution function, an equation is derived for the relationship between soil wetting ratio and storage index. In the derived equation, soil wetting ratio approaches zero as storage index approaches zero; when storage index tends to infinity, soil wetting ratio approaches a certain value (≤1) depending on the initial storage. Moreover, the derived equation leads to the exact SCS-CN method when initial water storage is zero. Therefore, the new distribution function for soil water storage capacity explains the SCS-CN method as a saturation excess runoff model and unifies the surface runoff modeling of SCS-CN method and VIC type of model.

Hydrol. Earth Syst. Sci. Discuss., https://doi.org/10.5194/hess-2018-32
Manuscript under review for journal Hydrol. Earth Syst. Sci.
Discussion started: 13 February 2018

Keywords: SCS curve number method, VIC, Xinanjiang, saturation excess, distribution function, soil water storage capacity, soil wetting

1. Introduction

The Soil Conservation Service Curve Number (SCS-CN) method [*Mockus*, 1972] has been popularly used for direct runoff estimation in engineering communities. Even though the SCS-CN method was obtained empirically [*Ponce*, 1996; *Beven*, 2011], it is often interpreted as an infiltration excess runoff model [*Bras*, 1990; *Mishra and Singh*, 1999]. *Yu* [1998] showed that partial area infiltration excess runoff generation on a statistical distribution of soil infiltration characteristics provided similar runoff generation equation as the SCS-CN method. Recently, *Hooshyar and Wang* [2016] derived an analytical solution for Richards' equation for ponded infiltration into a soil column bounded by a water table; and they showed that the SCS-CN method, as an infiltration excess model, is a special case of the derived general solution. The SCS-CN method has also been interpreted as a saturation excess runoff model [*Steenhuis et al.*, 1995; *Lyon et al.*, 2004; *Easton et al.*, 2008]. During an interview, Mockus, who developed the proportionality relationship of the SCS-CN method, stated that "saturation overland flow was the most likely runoff mechanism to be simulated by the method" [*Ponce*, 1996]. Recently, *Bartlett et al.* [2016a] developed a probabilistic framework, which provides a statistical justification of the SCS-CN method and extends the saturation excess interpretation of the event-based runoff of the method.

Since the 1970s, various saturation excess runoff models have been developed based on the concept of probability distribution of soil storage capacity [*Moore*, 1985]. TOPMODEL is a well-known saturation excess runoff model based on spatially distributed topography [*Beven and Kirkby*, 1979; *Sivapalan et al.*, 1987]. To quantify the dynamic change of saturation area during rainfall events, the spatial variability of soil moisture storage capacity is described by a cumulative probability distribution function in the Xinanjiang model [*Zhao*, 1977; *Zhao et al.*, 1992] and the Variable Infiltration Capacity (VIC) model [*Wood et al.*, 1992; *Liang et al.*, 1994]. The distribution of storage capacity is described by a power function in these models, which have been used for catchment scale runoff prediction and large scale land surface hydrologic simulations. *Bartlett et al.* [2016b] unified TOPMODEL, the VIC type of model, and the SCS-CN method by an event-based probabilistic storage framework, which includes a spatial description of the runoff concept of "prethreshold" and "threshold-excess" runoff [*Bartlett et al.*, 2016a].

By applying the generalized proportionality hypothesis from the SCS-CN method to mean annual water balance, *Wang and Tang* [2014] derived a one-parameter Budyko equation [*Budyko*, 1974] for mean annual evaporation ratio (i.e., the ratio of evaporation to precipitation) as a function of climate aridity index (i.e., the ratio of potential evaporation to precipitation). As an analogy to the Budyko framework, the SCS-CN method and the VIC type of model at the event scale can be represented by the relationship between soil wetting ratio, defined as the ratio between soil wetting and precipitation, and soil storage index which is defined as the ratio between soil wetting capacity and precipitation.

In this paper, the functional forms for soil wetting ratio versus soil storage index are compared between the SCS-CN model and the VIC/Xinanjiang type of model. Based on the comparison, a new distribution function is proposed for describing the soil water storage capacity in the VIC type of model so that the SCS-CN method and VIC type of model are unified. In section 2, the SCS-CN method is presented in the form of Budyko-type framework with two parameterization schemes. In section 3, the VIC type of model is presented in the form of

Budyko-type framework. In section 4, the SCS-CN method is then compared with the VIC type of model from the perspectives of number of parameters and boundary conditions (i.e., the lower and upper bounds of soil storage index). In section 5, the proposed new distribution function is introduced and compared with the power distribution of VIC type of model; and a modified SCS-CN method considering initial storage explicitly is derived from the new distribution function. Conclusions are drawn in section 6.

2. SCS curve number method

In this section, the SCS-CN method is described in the form of surface runoff modeling and then is presented for infiltration modeling in the Budyko-type framework. The initial storage at the beginning of a time interval (e.g., rainfall event) is denoted by S_0 [mm], and the maximum value of average storage capacity over the catchment is denoted by S_b [mm]. The storage capacity for soil wetting for the time interval, S_p [mm], is computed by:

$$S_p = S_b - S_0 \tag{1}$$

The total rainfall during the time interval is denoted by P [mm]. Before surface runoff is generated, a portion of rainfall is intercepted by vegetation and infiltrates into the soil. This portion of rainfall is called initial abstraction or initial soil wetting denoted by W_i [mm]. The remaining rainfall $(P - W_i)$ is partitioned into runoff and continuing soil wetting. This competition is captured by the proportionality relationship in the SCS-CN method:

$$\frac{W-W_i}{S_p-W_i} = \frac{Q}{P-W_i} \tag{2}$$

where W [mm] is the total soil wetting; $W - W_i$ is continuing wetting and $S_p - W_i$ is its potential value; Q [mm] is surface runoff; and $P - W_i$ is the available water and interpreted as the potential value of Q. Since rainfall is partitioned into total soil wetting and surface runoff, i.e., $P = W + Q$, surface runoff is computed by substituting $W = P - Q$ into equation (2):

Hydrol. Earth Syst. Sci. Discuss., https://doi.org/10.5194/hess-2018-32
Manuscript under review for journal Hydrol. Earth Syst. Sci.
Discussion started: 13 February 2018

$$Q = \frac{(P-W_i)^2}{P+S_p-2W_i} \quad (3)$$

This equation is used for computing direct runoff in the SCS-CN method.

The SCS-CN method can also be represented in terms of soil wetting ratio ($\frac{W}{P}$). Substituting equation (3) into $W = P - Q$ and dividing P on both sides, the soil wetting ratio equation is obtained:

$$\frac{W}{P} = \frac{\frac{S_p}{P}-\frac{W_i^2}{P^2}}{1+\frac{S_p}{P}-2\frac{W_i}{P}} \quad (4)$$

Climate aridity index is defined as the ratio between potential evaporation and precipitation. In climate aridity index, both available water supply and water demand are determined by climate.

$$\Phi_{sc} = \frac{S_p}{P} \quad (5)$$

A similar dimensionless parameter for the ratio between the maximum soil storage capacity and mean rainfall depth of rainfall events was defined in *Porporato et al.* [2004]. In soil storage index, water demand is determined by soil and available water supply is determined by climate. Substituting equation (5) into equation (4), the soil wetting equation for the SCS-CN method is obtained:

$$\frac{W}{P} = \frac{\Phi_{sc}-\frac{W_i^2}{P^2}}{1+\Phi_{sc}-2\frac{W_i}{P}} \quad (6)$$

Two potential schemes for parameterizing the initial wetting in equation (6) are discussed in the following sections.

2.1. Parameterization scheme 1: ratio between initial wetting and storage capacity

The initial wetting is usually parameterized as the ratio between initial wetting and storage capacity in the SCS-CN method. The potential for continuing wetting is called potential maximum retention and is denoted by $S_m = S_p - W_i$. S_m is computed as a function of curve number which is dependent on land use/land cover and soil permeability. The ratio between W_i and S_m in the SCS curve number method is denoted by $\lambda = \frac{W_i}{S_p - W_i}$, and then the ratio between initial soil wetting and storage capacity is computed by:

$$\frac{W_i}{S_p} = \frac{\lambda}{1+\lambda} \tag{7}$$

The value of λ varies in the range of $0 \le \lambda \le 0.3$, and a value of 0.2 is usually used [*Ponce and Hawkins*, 1996]. Substituting equation (7) into equation (6) leads to:

$$\frac{W}{P} = \frac{1-\left(\frac{\lambda}{1+\lambda}\right)^2 \Phi_{sc}}{1-\frac{2\lambda}{1+\lambda}+\Phi_{sc}{}^{-1}} \tag{8}$$

Equation (8) is plotted in Figure 1 for λ = 0.1 and 0.3. As we can see, the range of Φ_{sc} is dependent on the parameter λ. Since $W_i \le P$, Φ_{sc} is in the range of $\left[0, 1+\frac{1}{\lambda}\right]$. Equation (8) satisfies the following boundary conditions: $\frac{W}{P} \to 0$ as $\Phi_{sc} \to 0$; and $\frac{W}{P} \to 1$ as $\Phi_{sc} \to \frac{\lambda+1}{\lambda}$. When $\lambda \to 0$, equation (8) becomes:

$$\frac{W}{P} = \frac{1}{1+\Phi_{sc}{}^{-1}} \tag{9}$$

Equation (9) is the lower bound for $\frac{W}{P}$ based on this parameterization scheme.

2.2. Parameterization scheme 2: ratio between initial wetting and total wetting

In order to avoid the situation that the range of Φ_{sc} is dependent on the parameter λ, we can use the following parameterization scheme [*Chen et al.*, 2013; *Tang and Wang*, 2017]:

$$\varepsilon = \frac{W_i}{W} \tag{10}$$

Substituting equation (10) into equation (6), we can obtain the following equation:

$$\frac{W}{P} = \frac{\Phi_{sc} - \varepsilon^2 \frac{W^2}{P^2}}{1+\Phi_{sc} - 2\varepsilon \frac{W}{P}} \tag{11}$$

We can solve for $\frac{W}{P}$ from equation (11):

$$\frac{W}{P} = \frac{1+\Phi_{sc}-\sqrt{(1+\Phi_{sc})^2-4\varepsilon(2-\varepsilon)\Phi_{sc}}}{2\varepsilon(2-\varepsilon)} \quad (12)$$

Equation (12) has the same functional form as the derived Budyko equation for long-term evaporation ratio [*Wang and Tang*, 2014; *Wang et al.*, 2015]. Equation (12) satisfies the following boundary conditions: $\frac{W}{P} \to 0$ as $\Phi_{sc} \to 0$; and $\frac{W}{P} \to 1$ as $\Phi_{sc} \to \infty$. Based on equation (10), the range of ε is $[0, 1]$, and $\varepsilon = 1$ corresponds to the upper bound (Figure 1). Equation (12) becomes equation (9) as $\varepsilon \to 0$, and it is the lower bound. Figure 1 plots equation (12) for $\varepsilon =$ 0.1 and 0.3. Due to the dependence of the range of Φ_{sc} on the parameter λ in the first parameterization scheme, the second parameterization scheme is focused on in the following sections.

In the SCS-CN method, the soil wetting ratio is a function of soil storage index with a parameter for describing initial wetting. The average wetting capacity at the catchment scale is used for computing soil storage index; but the spatial variability of wetting capacity is not represented in the SCS-CN method.

3. Saturation excess runoff model

The spatial variability of soil water storage capacity is explicitly represented in the saturation excess runoff models such as VIC and Xinanjiang. In these models, the spatial variation of point-scale storage capacity (C) is represented by a power function:

$$F(C) = 1 - \left(1 - \frac{C}{C_m}\right)^{\beta} \quad (13)$$

where $F(C)$ is the cumulative probability, i.e., the fraction of catchment area for which the storage capacity is less than C [mm]; and C_m [mm] is the maximum value of point-scale storage capacity over the catchment. The water storage capacity includes vegetation interception, surface retention, and soil moisture capacity; β is the shape parameter of storage capacity distribution and is usually assumed to be a positive number. β ranges from 0.01 to 5.0 as suggested by *Wood et al.* [1992]. The storage capacity distribution curve is concave down for $0 < \beta < 1$ and concave up for $\beta > 1$. The average value of storage capacity over the catchment is equivalent to S_b in the SCS-CN method, and it is obtained by integrating the exceedance probability of storage capacity $S_b = \int_0^{C_m}(1 - F(x))\,dx$:

$$S_b = \frac{C_m}{\beta+1} \tag{14}$$

Similarly, for a given C, the catchment-scale storage S [mm] can be computed [*Moore*, 1985]:

$$S = S_b\left[1 - \left(1 - \frac{C}{C_m}\right)^{\beta+1}\right] \tag{15}$$

To derive wetting ratio as a function of soil storage index, the initial storage at the catchment scale is parameterized by the degree of saturation:

$$\psi = \frac{S_0}{S_b} \tag{16}$$

Recalling equation (1) and the definition of soil storage index (i.e., equation (5)), we obtain:

$$\frac{S_b}{P} = \frac{\Phi_{sc}}{1-\psi} \tag{17}$$

The value of C corresponding to the initial storage S_0 is denoted as C_0, and $S_0 = S_b\left[1 - \left(1 - \frac{C_0}{C_m}\right)^{\beta+1}\right]$ is obtained by substituting S_0 and C_0 into equation (15). When $P + C_0 \geq C_m$, each point within the catchment is saturated and soil wetting reaches its maximum value, i.e., $W = S_p$. Substituting $C_0 = C_m - C_m\left(1 - \frac{S_0}{S_b}\right)^{\frac{1}{\beta+1}}$ into $P + C_0 \geq C_m$, we obtain:

$$\Phi_{sc} \leq b \text{ where } b = (\beta+1)^{-1}(1-\psi)^{\frac{\beta}{\beta+1}} \tag{18}$$

Therefore, this condition is equivalent to:

$$\frac{W}{P} = \Phi_{sc} \text{ when } \Phi_{sc} \leq b \tag{19}$$

Next, we will derive $\frac{W}{P}$ for the condition of $\Phi_{sc} > b$. The storage at the end of the modeling period (e.g., rainfall-runoff event) is denoted as S_1, which is computed by:

$$S_1 = S_b\left[1 - \left(1 - \frac{P + C_0}{C_m}\right)^{\beta+1}\right] \tag{20}$$

Since $W = S_1 - S_0$, wetting is computed by:

$$W = S_b\left[1 - \left(1 - \frac{P + C_0}{C_m}\right)^{\beta+1}\right] - S_0 \tag{21}$$

From equation (21), we obtain (see Appendix A for details):

$$\frac{W}{P} = \Phi_{sc}\left[1 - \left(1 - b{\Phi_{sc}}^{-1}\right)^{\beta+1}\right] \text{ when } \Phi_{sc} > b \tag{22}$$

The limit of equation (22) for $\Phi_{sc} \to \infty$ can be obtained (see Appendix B for details):

$$\lim_{\Phi_{sc}\to\infty} \frac{W}{P} = (1-\psi)^{\frac{\beta}{\beta+1}} \tag{23}$$

Equations (19) and (22) provide $\frac{W}{P}$ as a function of Φ_{sc} with two parameters (ψ and β). Figure 2 plots equations (19) and (22) for ψ = 0 and 0.5 when β = 0.2 and 2. As we can see, $\frac{W}{P}$ decreases as β increases for given values of ψ and Φ_{sc}; and $\frac{W}{P}$ decreases as ψ increases for given values of β and Φ_{sc}, implicating that soil wetting ratio decreases with the degree of initial saturation under a given soil storage index.

4. Comparison between SCS-CN model and VIC type of model

The SCS-CN model with the parameterization of ratio between initial wetting and total wetting is compared with the VIC type of saturation excess runoff model. In sections 2 and 3, we derived $\frac{W}{P}$ as a function of Φ_{sc} based on the SCS-CN method and the VIC type of model, which uses a power function to describe the spatial distribution of storage capacity. The SCS-CN method is a function of storage capacity S_p; but the VIC type of model is a function of storage capacity S_p

Hydrol. Earth Syst. Sci. Discuss., https://doi.org/10.5194/hess-2018-32
Manuscript under review for journal Hydrol. Earth Syst. Sci.
Discussion started: 13 February 2018

and the degree of initial saturation $\frac{S_0}{S_b}$. As a result, the function of $\frac{W}{P} \sim \frac{S_p}{P}$ for the SCS-CN method has only one parameter (ε), but it has two parameters (β and ψ) for the VIC type of model.

Table 1 shows the boundary conditions for the relationships between $\frac{W}{P}$ and Φ_{sc} from the SCS-CN method and the VIC type of model. The lower boundary of the SCS-CN method with parameter ε is $\frac{W}{P} \to 0$ as $\Phi_{sc} \to 0$. However, for the VIC type of model, $\frac{W}{P} = \Phi_{sc}$ when $\Phi_{sc} \leq b$. For the SCS-CN method, W reaches its maximum (S_p) when rainfall reaches infinity; while for the VIC type of model, W reaches its maximum value (S_p) when rainfall reaches a finite number ($C_m - C_0$). In other words, for the SCS-CN method, the entire catchment becomes saturated when rainfall reaches infinity; while for the VIC type model, the entire catchment becomes saturated when rainfall reaches a finite number.

As shown in Table 1, the upper boundary of the SCS-CN method (with parameter ε) is 1. However, for the VIC type of model, the upper boundary is $(1-\psi)^{\frac{\beta}{\beta+1}}$ instead of 1. This is due to the effect of initial storage in the VIC type of model. When initial storage is zero (i.e., $\psi = 0$), the wetting ratio $\frac{W}{P}$ for the VIC type of model has the same upper boundary condition as the SCS-CN method.

5. Unification of SCS-CN method and VIC type of model

Based on the comparison between the SCS-CN method and VIC type of model, a new distribution function is proposed in this section for describing the spatial distribution of soil water storage capacity, which unifies the SCS-CN method and VIC type of model. As discussed in section 4, the upper boundary condition of the SCS-CN model (i.e., $\frac{W}{P} \to 1$ as $\Phi_{sc} \to \infty$) does not depend on the initial storage. This upper boundary condition needs to be modified by including the effect of initial storage so that the limit of $\frac{W}{P}$ as $\Phi_{sc} \to \infty$ is dependent on the degree of initial saturation like the VIC type of model. However, the lower boundary condition of the VIC model needs to be modified so that the lower boundary condition follows that $\frac{W}{P} \to 0$ as $\Phi_{sc} \to 0$ like the SCS-CN method. Through these modifications, the SCS-CN method and the VIC type of saturation excess runoff model can be unified from the functional perspective of soil wetting ratio.

Based on the comparison one may have the following questions: 1) Can the SCS-CN method be derived from the VIC type of model by setting initial storage to zero? 2) If yes, what is the distribution function for soil water storage capacity? Once we answer these questions, a modified SCS-CN method considering initial storage explicitly can be derived as a saturation excess runoff model based on a distribution function of water storage capacity, and it unifies the SCS-CN method and VIC type of model. In this section, a new distribution function is proposed for describing the spatial variability of soil water storage capacity, from which the SCS-CN method is derived as a VIC type of model.

5.1. A new distribution function

The probability density function (PDF) of the new distribution for describing the spatial distribution of water storage capacity is represented by:

$$f(C) = \frac{(2-a)\mu^2}{[(C+\mu)^2 - 2a\mu C]^{3/2}} \tag{24}$$

where C is point-scale water storage capacity and supported on a positive semi-infinite interval $(C \geq 0)$; a is the shape parameter and its range is $0 < a < 2$; and μ is the mean of the distribution (i.e., the scale parameter). Figure 3a plots the PDFs for five sets of shape and scale parameters. When $a \leq 1$, the PDF monotonically decreases with the increase of C, i.e., the peak of PDF occurs at $C = 0$; while when $a > 1$, the peak of PDF occurs at $C > 0$ and the location of the peak depends on the values of a and μ. For comparison, Figure 3b plots the PDF for VIC model:

$$f(C) = \frac{\beta}{C_m}\left(1 - \frac{C}{C_m}\right)^{\beta-1} \quad (25)$$

As shown by the solid black curve in Figure 3b, when $0 < \beta < 1$, $f(C)$ approaches infinity as $C \to C_m$. It is a uniform distribution when $\beta = 1$. The peak of PDF occurs at $C = 0$ when $\beta > 1$. Therefore, the peak of PDF for VIC model occurs at $C = 0$ or C_m.

The cumulative distribution function (CDF) corresponding to the proposed PDF is obtained by integrating equation (24):

$$F(C) = 1 - \frac{1}{a} + \frac{C + (1-a)\mu}{a\sqrt{(C+\mu)^2 - 2a\mu C}} \quad (26)$$

Figure 4a plots the CDFs corresponding to the PDFs in Figure 3a. For comparison, Figure 4b plots the CDFs corresponding to the PDFs in Figure 3b. The storage capacity distribution curve for the proposed distribution is concave up for $a \leq 1$ and *S*-shape for $a > 1$ (Figure 4a); while the storage capacity distribution curve for VIC model is concave up for $\beta > 1$ and concave down for $0 < \beta < 1$ (Figure 4b). Therefore, the proposed distribution can fit the *S*-shape of cumulative distribution for storage capacity which is observed from soil data [*Huang et al.*, 2003], but the power distribution of VIC type of model is not able to fit the *S*-shape of CDF.

5.2. Deriving SCS-CN method from the proposed distribution function

The soil wetting and surface runoff can be computed when equation (26) is used to describe the spatial distribution of soil water storage capacity in a catchment. The average value of storage capacity over the catchment is the mean of the distribution:

$$\mu = S_b \quad (27)$$

Hydrol. Earth Syst. Sci. Discuss., https://doi.org/10.5194/hess-2018-32
Manuscript under review for journal Hydrol. Earth Syst. Sci.
Discussion started: 13 February 2018

For a given C, the catchment-scale storage S can be computed by $S = \int_0^C [1 - F(x)]dx$ [*Moore*, 1985]. From equation (26), we obtain:

$$S = \frac{C + S_b - \sqrt{(C + S_b)^2 - 2aS_bC}}{a} \tag{28}$$

For a rainfall-runoff event, the average initial storage at the catchment scale is denoted as S_0 and the corresponding value of C is denoted as C_0. Substituting S_0 and C_0 into equation (28), we obtain:

$$S_0 = \frac{C_0 + S_b - \sqrt{(C_0 + S_b)^2 - 2aS_bC_0}}{a} \tag{29}$$

Dividing S_b in both-hand sides of equation (29), we obtain:

$$m = \frac{\psi(2 - a\psi)}{2(1 - \psi)} \tag{30}$$

where $\psi = \frac{S_0}{S_b}$ is defined in equation (16), and m is defined as:

$$m = \frac{C_0}{S_b} \tag{31}$$

The average storage at the catchment scale after infiltration is computed by substituting $C = C_0 + P$ into equation (28):

$$S_1 = \frac{C_0 + P + S_b - \sqrt{(C_0 + P + S_b)^2 - 2aS_b(C_0 + P)}}{a} \tag{32}$$

The soil wetting is computed as the difference between S_1 and S_0:

$$W = \frac{P + \sqrt{(C_0 + S_b)^2 - 2aS_bC_0} - \sqrt{(C_0 + P + S_b)^2 - 2aS_b(C_0 + P)}}{a} \tag{33}$$

Dividing P on the both-hand sides of equation (33) and substituting equation (31), we obtain:

$$\frac{W}{P} = \frac{1 + \frac{S_b}{P}\sqrt{(m+1)^2 - 2am} - \sqrt{\left(1 + (m+1)\frac{S_b}{P}\right)^2 - 2am\left(\frac{S_b}{P}\right)^2 - 2a\frac{S_b}{P}}}{a} \tag{34}$$

Substituting equation (17) into equation (34), we obtain:

Hydrol. Earth Syst. Sci. Discuss., https://doi.org/10.5194/hess-2018-32
Manuscript under review for journal Hydrol. Earth Syst. Sci.
Discussion started: 13 February 2018

$$\frac{W}{P}=\frac{1+\frac{\sqrt{(m+1)^2-2am}}{1-\psi}\Phi_{sc}-\sqrt{\left(1+\frac{m+1}{1-\psi}\Phi_{sc}\right)^2-2am\left(\frac{\Phi_{sc}}{1-\psi}\right)^2-\frac{2a}{1-\psi}\Phi_{sc}}}{a} \tag{35}$$

Figure 5 plots equation (35) for ψ = 0, 0.4, and 0.6 when a = 0.6 and 1.8. As we can see, $\frac{W}{P}$ increases with a for given values of ψ and Φ_{sc}; and $\frac{W}{P}$ decreases with ψ for given values of a and Φ_{sc}, which is consistent with the VIC model and implicates that soil wetting ratio decreases with the degree of initial saturation under a storage index. As shown in Figure 5, equation (35) satisfies the lower boundary of SCS-CN method and the upper boundary of the VIC model. Specifically, equation (35) satisfies the following boundary conditions (see Appendix C for details) shown in Table 1:

$$\lim_{\Phi_{sc}\to 0}\frac{W}{P}=0 \tag{36-1}$$

$$\lim_{\Phi_{sc}\to \infty}\frac{W}{P}=\frac{\sqrt{(m+1)^2-2am}+a-m-1}{a\sqrt{(m+1)^2-2am}} \tag{36-2}$$

When the effect of initial storage is negligible (i.e., $\psi=0$), $\frac{S_b}{P}=\Phi_{sc}$ from equation (17) and $m=0$ from equation (30). Then, equation (35) becomes:

$$\frac{W}{P}=\frac{1+\frac{S_b}{P}-\sqrt{\left(1+\frac{S_b}{P}\right)^2-2a\frac{S_b}{P}}}{a} \tag{37}$$

Equation (37) is same as equation (12) with $a=2\varepsilon(2-\varepsilon)$. We can obtain the following equation from equation (37) (see Appendix D for detailed derivation):

$$\frac{Q}{P-\varepsilon W}=\frac{W-\varepsilon W}{S_b-\varepsilon W} \tag{38}$$

where εW is defined as initial abstraction (W_i) in the SCS-CN method. Since $S_b=S_p$ when $\psi=0$, equation (38) is same as equation (2), i.e., the proportionality relationship of SCS-CN method.

Equation (35) is derived from the VIC type model by using equation (26) to describe the spatial distribution of soil water storage capacity. From this perspective, equation (35) is a saturation excess runoff model. Since equation (35) becomes the SCS-CN method when initial storage is negligible, equation (35) is the modified SCS-CN method which considers the effect of initial storage on runoff generation explicitly. Therefore, the new distribution function represented by equation (26) unifies the SCS-CN method and VIC type of model.

Bartlett et al. [2016a] developed an event-based probabilistic storage framework including a spatial description of "prethreshold" and "threshold-excess" runoff; and the framework has been utilized for unifying TOPMODEL, VIC and SCS-CN [*Bartlett et al.*, 2016b]. The extended SCS-CN method (SCS-CNx) from the probabilistic storage framework is derived given the following assumptions: 1) the spatial distribution of rainfall is exponential; 2) the spatial distribution of soil moisture deficit is uniform; and 3) the spatial distribution of storage capacity is exponential. When "prethreshold" runoff is zero (i.e., there is only threshold-excess or saturation excess runoff), the SCS-CNx method leads to the SCS-CN method without the initial abstraction term (i.e., there is no εW term in equation (38)). In this paper, the new probability distribution function is used for storage capacity in the VIC model in which the spatial distribution of precipitation is assumed to be uniform. The obtained equation for saturation excess runoff leads to the exact SCS-CN method as shown in equation (38).

5.3. Surface runoff of unified SCS-CN and VIC model

From the unified SCS-CN and VIC model (i.e., equation (34)), surface runoff (Q) can be computed as:

$$Q = \frac{(a-1)P - S_b\sqrt{(m+1)^2 - 2am} + \sqrt{[P+(m+1)S_b]^2 - 2amS_b{}^2 - 2aS_bP}}{a} \quad (39)$$

The parameter m is computed by equation (30) as a function of ψ and a. Equation (39) represents surface runoff as a function of precipitation (P), average soil water storage capacity (S_b), shape parameter of storage capacity distribution (a), and initial soil moisture (ψ). Figure 6 plots equation (39) under different values of P, S_b, a, and ψ. Figure 6a shows the effects of S_b and ψ on rainfall-runoff relationship with given shape parameter of a=1.9. The solid lines show the rainfall-runoff relations with zero initial storage (ψ=0); and the dashed lines show the rainfall-runoff relations with ψ=0.2. Given the same amount of precipitation and storage capacity, wetter soil (ψ=0.2) generates more surface runoff than drier soil (ψ=0); and the difference of runoff is higher for watersheds with larger average storage capacity. Figure 6b shows the effects of S_b and a on rainfall-runoff relationship with given initial soil moisture (ψ=0.2). The solid lines show the rainfall-runoff relations for a=1.9; and the dashed lines show the rainfall-runoff relations for a=1.2. As we can see, the shape parameter affects the runoff generation significantly for watersheds with larger average storage capacity.

In the SCS-CN method, surface runoff is computed as $Q = \frac{(P-0.2S_b)^2}{P+0.8S_b}$. The effect of initial soil moisture on runoff is considered implicitly by varying the curve number for normal, dry and wet conditions depending on the antecedent moisture condition. In the unified SCS-CN model shown in equation (39), the effect of initial soil moisture is explicitly included through ψ, which is the ratio between average initial water storage and average storage capacity. In the SCS-CN method, the value of initial abstraction W_i is parameterized as a function of average storage capacity, i.e., $W_i = 0.2S_b$. In the unified SCS-CN model shown in equation (39), W_i is dependent on the shape parameter a. Therefore, the unified SCS-CN model extends the original SCS-CN method for including the effect of initial soil moisture explicitly and estimating the parameter for initial abstraction.

Hydrology and Earth System Sciences
Open Access
Discussions

6. Conclusions

In this paper, the SCS-CN method and the saturation excess runoff models based on distribution functions (e.g., VIC model) are presented in terms of soil wetting (i.e., infiltration). Like the Budyko framework, the relationship between soil wetting ratio and soil storage index is obtained for the SCS-CN method and the VIC type of model. It is found that the boundary conditions for the obtained functions do not fully match. For the SCS-CN method, soil wetting ratio approaches 1 when soil storage index approaches infinity, and this is due to the limitation of the SCS-CN method, i.e. the initial soil moisture condition is not explicitly represented in the proportionality relationship. However, for the VIC type of model, soil wetting ratio equals soil storage index when soil storage index is lower than a certain value, and this is due to the finite bound of the distribution function of storage capacity.

In this paper, a new distribution function, which is supported by $x \in [0, \infty)$ instead of a finite upper bound, is proposed for describing the spatial distribution of soil water storage capacity. From this new distribution function, an equation is derived for the relationship between soil wetting ratio and storage index, and this equation satisfies the following boundary conditions: when storage index approaches zero, soil wetting ratio approaches zero; when storage index approaches infinity, soil wetting ratio approaches a certain value (≤ 1) depending on the initial storage. Meanwhile, the model becomes the exact SCS-CN method when initial storage is negligible. Therefore, the new distribution function for soil water storage capacity explains the SCS-CN method as a saturation excess runoff model, and unifies the SCS-CN method and the VIC type of model for surface runoff modeling.

Future potential work could test the performance of the proposed new distribution function for quantifying the spatial distribution of storage capacity by analyzing the spatially

Hydrol. Earth Syst. Sci. Discuss., https://doi.org/10.5194/hess-2018-32
Manuscript under review for journal Hydrol. Earth Syst. Sci.
Discussion started: 13 February 2018

Hydrology and Earth System Sciences
Open Access
Discussions

distributed soil data. On one hand, the distribution functions of probability distributed model [*Moore*, 1985], VIC model, and Xinanjiang model could be replaced by the new distribution function and the model performance would be further evaluated. On the other hand, the extended SCS-CN method (i.e., equation (35)), which includes initial storage explicitly, could be used for surface runoff modeling in SWAT model, and the model performance would be evaluated.

Acknowledgements

The author would like to acknowledge the support of NSF grant CBET-1665343. The author would like to thank University of Central Florida Faculty Excellence for their support during his sabbatical leave. This paper is theoretical and does not contain any supplementary data.

Appendix A

The following equation is obtained by dividing P on both sides of equation (21):

$$\frac{W}{P} = \frac{S_b - S_0}{P} - \frac{S_b}{P}\left(1 - \frac{P + C_0}{C_m}\right)^{\beta+1} \tag{A1}$$

Substituting $\frac{C_0}{C_m} = 1 - \left(1 - \frac{S_0}{S_b}\right)^{\frac{1}{\beta+1}}$ into equation (A1), we obtain:

$$\frac{W}{P} = \frac{S_b - S_0}{P} - \frac{S_b}{P}\left(1 - \frac{P}{C_m} - \left[1 - \left(1 - \frac{S_0}{S_b}\right)^{\frac{1}{\beta+1}}\right]\right)^{\beta+1} \tag{A2}$$

Substituting equation (14) into equation (A2),

$$\frac{W}{P} = \frac{S_b - S_0}{P} - \left(\left(\frac{S_b - S_0}{P}\right)^{\frac{1}{\beta+1}} - \frac{\left(\frac{S_b}{P}\right)^{-\frac{\beta}{\beta+1}}}{\beta+1}\right)^{\beta+1} \tag{A3}$$

Substituting equations (5) and (17) into (A3), we obtain:

$$\frac{W}{P} = \Phi_{sc} - \left(\Phi_{sc}{}^{\frac{1}{\beta+1}} - \frac{\left(\frac{\Phi_{sc}}{1-\psi}\right)^{-\frac{\beta}{\beta+1}}}{\beta+1} \right)^{\beta+1} \tag{A4}$$

which leads to:

$$\frac{W}{P} = \Phi_{sc}\left[1 - \left(1 - b\Phi_{sc}{}^{-1}\right)^{\beta+1}\right] \tag{A5}$$

where b is defined in equation (18).

Appendix B

$$\lim_{\Phi_{sc}\to\infty} \frac{W}{P} = \lim_{\Phi_{sc}\to\infty} \Phi_{sc}\left[1 - \left(1 - b\Phi_{sc}{}^{-1}\right)^{\beta+1}\right] \tag{B1}$$

The right hand side of equation (B1) is re-written as:

$$\lim_{\Phi_{sc}\to\infty} \Phi_{sc}\left[1 - \left(1 - b\Phi_{sc}{}^{-1}\right)^{\beta+1}\right] = \lim_{\Phi_{sc}\to\infty} \frac{1-\left(1-b\Phi_{sc}{}^{-1}\right)^{\beta+1}}{\Phi_{sc}{}^{-1}} \tag{B2}$$

Since $\lim_{\Phi_{sc}\to\infty} 1 - \left(1 - \mathrm{b}\Phi_{sc}{}^{-1}\right)^{\beta+1} = 0$ and $\lim_{\Phi_{sc}\to\infty} \Phi_{sc}{}^{-1} = 0$, we apply the L'Hospital's Rule,

$$\lim_{\Phi_{sc}\to\infty} \frac{\left[1-\left(1-\mathrm{b}\Phi_{sc}{}^{-1}\right)^{\beta+1}\right]'}{\left(\Phi_{sc}{}^{-1}\right)'} = \lim_{\Phi_{sc}\to\infty} \mathrm{b}(\beta+1)\left(1 - \mathrm{b}\Phi_{sc}{}^{-1}\right)^{\beta} \tag{B3}$$

Since $\lim_{\Phi_{sc}\to\infty} \left(1 - \mathrm{b}\Phi_{sc}{}^{-1}\right)^{\beta} = 1$, the limit for $\frac{W}{P}$ is obtained:

$$\lim_{\Phi_{sc}\to\infty} \frac{W}{P} = b(\beta+1) \tag{B4}$$

Substituting equation (18) into (B4), we obtain:

$$\lim_{\Phi_{sc}\to\infty} \frac{W}{P} = (1-\psi)^{\frac{\beta}{\beta+1}} \tag{B5}$$

Appendix C

$$\lim_{\Phi_{sc}\to\infty}\frac{W}{P}=\lim_{\Phi_{sc}\to\infty}\frac{1+\frac{\sqrt{(m+1)^2-2am}}{1-\psi}\Phi_{sc}-\sqrt{\left(1+\frac{m+1}{1-\psi}\Phi_{sc}\right)^2-2am\left(\frac{\Phi_{sc}}{1-\psi}\right)^2-\frac{2a}{1-\psi}\Phi_{sc}}}{a} \tag{C1}$$

Multiplying $1+\frac{\sqrt{(m+1)^2-2am}}{1-\psi}\Phi_{sc}+\sqrt{\left(1+\frac{m+1}{1-\psi}\Phi_{sc}\right)^2-2am\left(\frac{\Phi_{sc}}{1-\psi}\right)^2-\frac{2a}{1-\psi}\Phi_{sc}}$ to the denominator and numerator of the right hand side, equation (C1) leads to:

$$\lim_{\Phi_{sc}\to\infty}\frac{W}{P}=\frac{1}{a}\lim_{\Phi_{sc}\to\infty}\frac{\frac{2\sqrt{(m+1)^2-2am}}{1-\psi}\Phi_{sc}-\frac{2(m+1)}{1-\psi}\Phi_{sc}+\frac{2a}{1-\psi}\Phi_{sc}}{1+\frac{\sqrt{(m+1)^2-2am}}{1-\psi}\Phi_{sc}+\sqrt{\left(1+\frac{m+1}{1-\psi}\Phi_{sc}\right)^2-2am\left(\frac{\Phi_{sc}}{1-\psi}\right)^2-\frac{2a}{1-\psi}\Phi_{sc}}} \tag{C2}$$

Dividing Φ_{sc} in the denominator and numerator, we obtain:

$$\lim_{\Phi_{sc}\to\infty}\frac{W}{P}=\frac{1}{a(1-\psi)}\lim_{\Phi_{sc}\to\infty}\frac{2\sqrt{(m+1)^2-2am}-2(m+1)+2a}{\frac{1}{\Phi_{sc}}+\frac{\sqrt{(m+1)^2-2am}}{1-\psi}+\sqrt{\left(\frac{1}{\Phi_{sc}}+\frac{m+1}{1-\psi}\right)^2-2am\left(\frac{1}{1-\psi}\right)^2-\frac{2a}{(1-\psi)\Phi_{sc}}}} \tag{C3}$$

Therefore, the limit of $\frac{W}{P}$ as $\Phi_{sc}\to\infty$ is:

$$\lim_{\Phi_{sc}\to\infty}\frac{W}{P}=\frac{\sqrt{(m+1)^2-2am}+a-m-1}{a\sqrt{(m+1)^2-2am}} \tag{C4}$$

Appendix D

Substituting $a=2\varepsilon(2-\varepsilon)$ into equation (37), one can obtain:

$$\frac{W}{P}=\frac{1+\frac{S_b}{P}-\sqrt{\left(1+\frac{S_b}{P}\right)^2-4\varepsilon(2-\varepsilon)\frac{S_b}{P}}}{2\varepsilon(2-\varepsilon)} \tag{D1}$$

Equation (D1) is the solution of the following quadratic function:

$$\varepsilon(2-\varepsilon)\left(\frac{W}{P}\right)^2-\left(1+\frac{S_b}{P}\right)\frac{W}{P}+\frac{S_b}{P}=0 \tag{D2}$$

Multiplying P^2 at the both-hand sides of equation (D2), equation (D2) becomes:

$$\varepsilon(2-\varepsilon)W^2-(P+S_b)W+S_bP=0 \tag{D3}$$

Equation (D3) can be written as the following one:

$$\frac{P-W}{P-\varepsilon W}=\frac{W-\varepsilon W}{S_b-\varepsilon W} \tag{D4}$$

Substituting $Q = P - W$ into equation (D4), we obtain the proportionality relationship of SCS-CN method:

$$\frac{Q}{P-\varepsilon W} = \frac{W-\varepsilon W}{S_b-\varepsilon W} \tag{D5}$$

References

Bartlett, M. S., A. J. Parolari, J. J. McDonnell, and A. Porporato (2016a), Beyond the SCS-CN method: A theoretical framework for spatially lumped rainfall-runoff response, *Water Resour. Res.*, 52, 4608–4627, doi:10.1002/2015WR018439.

Bartlett, M. S., A. J. Parolari, J. J. McDonnell, and A. Porporato (2016b), Framework for event-based semidistributed modeling that unifies the SCS-CN method, VIC, PDM, and TOPMODEL, *Water Resour. Res.*, 52, 7036–7052, doi:10.1002/2016WR019084.

Beven, K. J. (2011), *Rainfall-runoff modelling: the primer*, John Wiley & Sons.

Beven, K., and M. J. Kirkby (1979), A physically based, variable contributing area model of basin hydrology, *Hydrol. Sci. J.*, 24(1), 43-69.

Bras, R. L. (1990), *Hydrology: an introduction to hydrologic science*, Addison Wesley Publishing Company.

Budyko, M. I. (1974), *Climate and Life*, 508 pp., Academic Press, New York.

Chen, X., N. Alimohammadi, and D. Wang (2013), Modeling interannual variability of seasonal evaporation and storage change based on the extended Budyko framework, *Water Resour. Res.*, 49, doi:10.1002/wrcr.20493.

Easton, Z. M., D. R. Fuka, M. T. Walter, D. M. Cowan, E. M. Schneiderman, and T. S. Steenhuis (2008), Re-conceptualizing the soil and water assessment tool (SWAT) model to predict runoff from variable source areas, *J. Hydrol.*, 348(3), 279-291.

Huang, M., X. Liang, and Y. Liang (2003), A transferability study of model parameters for the variable infiltration capacity land surface scheme, *J. Geophys. Res.*, 108(D22), 8864, doi:10.1029/2003JD003676.

Hooshyar, M., and D. Wang (2016), An analytical solution of Richards' equation providing the physical basis of SCS curve number method and its proportionality relationship, *Water Resour. Res.*, 52(8), 6611-6620, doi: 10.1002/2016WR018885.

Liang, X., D. P. Lettenmaier, E. F. Wood, and S. J. Burges (1994), A simple hydrologically based model of land surface water and energy fluxes for general circulation models, *J. Geophys. Res.: Atmospheres*, 99(D7), 14415-14428.

Lyon, S. W., M. T. Walter, P. Gérard-Marchant, and T. S. Steenhuis (2004), Using a topographic index to distribute variable source area runoff predicted with the SCS curve - number equation, *Hydrol. Process.*, 18(15), 2757-2771.

Mishra, S. K., and V. P. Singh (1999), Another look at SCS-CN method, *J. Hydrol. Eng.*, 4(3), 257-264.

Mockus, V. (1972), *National Engineering Handbook Section 4, Hydrology*, NTIS.

Moore, R. J. (1985), The probability-distributed principle and runoff production at point and basin scales, *Hydrol. Sci. J.*, 30, 273-297.

Ponce, V. (1996), Notes of my conversation with Vic Mockus, Unpublished material. Available from: http://mockus. sdsu. edu/[Accessed 29 September 2017].

Ponce, V. M. and R. H. Hawkins (1996), Runoff curve number: has it reached maturity? *J. Hydrol. Eng.*, 1(1), 9-20.

Porporato, A., E. Daly, and I. Rodriguez-Iturbe (2004), Soil Water Balance and Ecosystem Response to Climate Change, *Am. Nat.*, 164(5), 625-632.

Sivapalan, M., K. Beven, E. F. Wood (1987), On hydrologic similarity: 2. A scaled model of storm runoff production, *Water Resour. Res.*, 23(12), 2266–2278.

Steenhuis, T. S., M. Winchell, J. Rossing, J. A. Zollweg, and M. F. Walter (1995), SCS runoff equation revisited for variable-source runoff areas, *J. Irrig. Drain. Eng.*, 121(3), 234-238.

Tang, Y., and D. Wang (2017), Evaluating the role of watershed properties in long-term water balance through a Budyko equation based on two-stage partitioning of precipitation, *Water Resour. Res.*, 53, 4142–4157, doi:10.1002/2016WR019920.

Wang, D. and Y. Tang (2014), A one-parameter Budyko model for water balance captures emergent behavior in Darwinian hydrologic models, *Geophys. Res. Lett.*, *41*, 4569–4577, doi:10.1002/2014GL060509.

Wang, D., J. Zhao, Y. Tang, and M. Sivapalan (2015), A thermodynamic interpretation of Budyko and L'vovich formulations of annual water balance: Proportionality hypothesis and maximum entropy production, *Water Resour. Res.*, 51, 3007–3016, doi:10.1002/2014WR016857.

Wood, E. F., D. P. Lettenmaier, and V. G. Zartarian (1992), A land‐surface hydrology parameterization with subgrid variability for general circulation models, *J. Geophys. Res.: Atmospheres*, 97(D3), 2717-2728.

Yu, B. (1998), Theoretical justification of SCS method for runoff estimation, *J. Irrig. Drain. Eng.*, 124(6), 306-310.

Zhao, R. (1977), *Flood forecasting method for humid regions of China*, East China College of Hydraulic Engineering, Nanjing, China.

Zhao, R. (1992), The Xinanjiang model applied in China, *J. Hydrol.*, 135, 371-381.

Figure captions:

Figure 1: Wetting ratio $\left(\frac{W}{P}\right)$ versus soil storage index $\left(\frac{S_p}{P}\right)$ from the SCS-CN method based on two parameterization schemes: $\lambda = \frac{W_i}{S_p - W_i}$ (scheme 1) and $\varepsilon = \frac{W_i}{W}$ (scheme 2).

Figure 2: The impact of β and the degree of initial storage $(\psi = S_0/S_b)$ on soil wetting ratio (W/P).

Figure 3: The probability density functions (PDF) with different parameter values: (a) the proposed PDF represented by equation (24); and (b) the power distribution of VIC model, i.e., equation (25).

Figure 4: The cumulative distribution functions (CDF) with different parameter values: (a) the proposed distribution function represented by equation (26); and (b) the power distribution of VIC model represented by equation (13).

Figure 5: The effects of the degree of initial storage (ψ=0, 0.4, and 0.6) and shape parameter (a=0.6 and 1.8) on soil wetting in the modified SCS-CN method derived from the proposed distribution function for soil water storage capacity.

Hydrol. Earth Syst. Sci. Discuss., https://doi.org/10.5194/hess-2018-32
Manuscript under review for journal Hydrol. Earth Syst. Sci.
Discussion started: 13 February 2018

Table 1: The boundary conditions of the functions for relating wetting ratio $\left(\frac{W}{P}\right)$ to soil storage index (Φ_{sc}): 1) the SCS-CN method; 2) the VIC type of model; and 3) the modified SCS-CN method based on the proposed new distribution for VIC type of model.

Event Scale Model	Lower Boundary Condition	Upper Boundary Condition
SCS-CN, parameterization of initial wetting, $\varepsilon = \frac{W_i}{W}$	$\frac{W}{P} \to 0$ as $\Phi_{sc} \to 0$	$\frac{W}{P} \to 1$ as $\Phi_{sc} \to \infty$
Power function for storage capacity distribution (VIC type of model)	$\frac{W}{P} = \Phi_{sc}$ when $\Phi_{sc} \leq a$	$\frac{W}{P} \to (1-\psi)^{\frac{\beta}{\beta+1}}$ as $\Phi_{sc} \to \infty$
Modified SCS-CN method based on the proposed distribution for storage capacity	$\frac{W}{P} \to 0$ as $\Phi_{sc} \to 0$	$\frac{W}{P} \to \frac{\sqrt{(m+1)^2-2am}+a-m-1}{a\sqrt{(m+1)^2-2am}}$ as $\Phi_{sc} \to \infty$

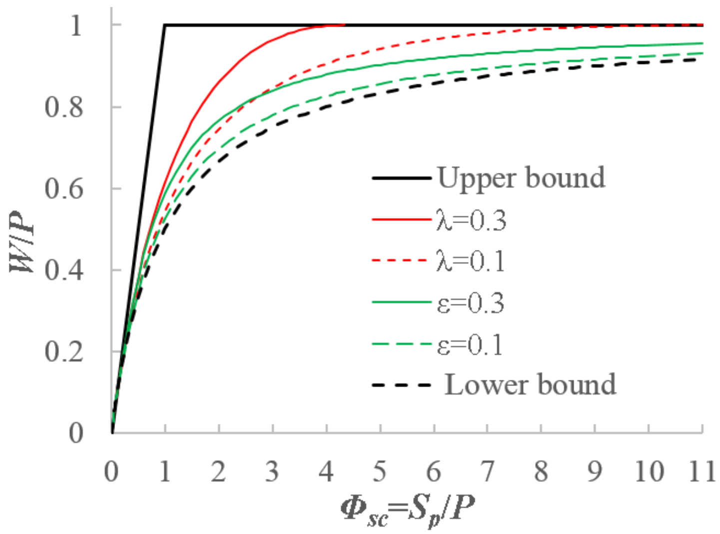

Figure 1: Wetting ratio $\left(\frac{W}{P}\right)$ versus soil storage index $\left(\frac{S_p}{P}\right)$ from the SCS-CN method based on two parameterization schemes: $\lambda = \frac{W_i}{S_p - W_i}$ (scheme 1) and $\varepsilon = \frac{W_i}{W}$ (scheme 2).

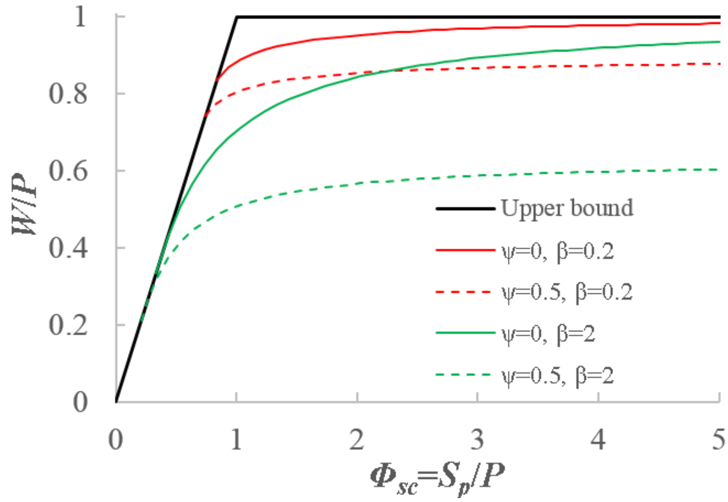

Figure 2: The impact of β and the degree of initial storage ($\psi = S_0/S_b$) on soil wetting ratio (W/P).

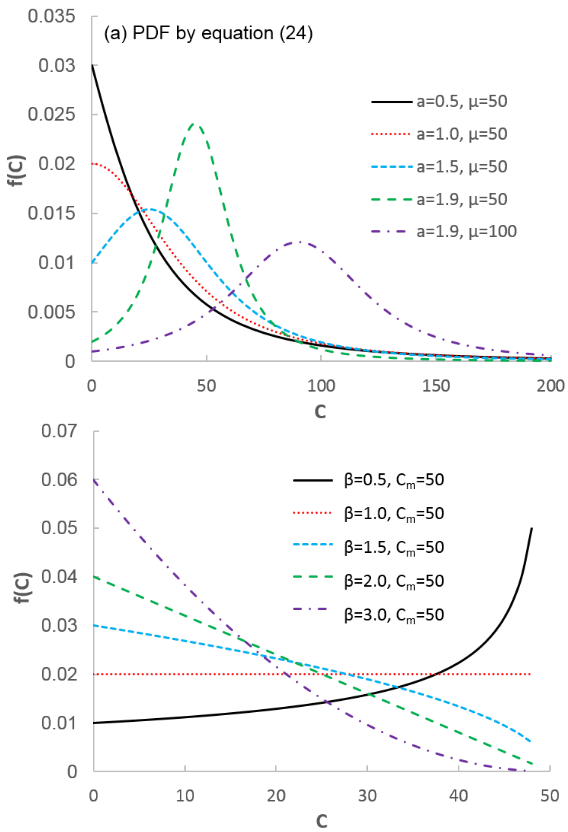

Figure 3: The probability density functions (PDF) with different parameter values: (a) the proposed PDF represented by equation (24); and (b) the power distribution of VIC model, i.e., equation (25).

Hydrol. Earth Syst. Sci. Discuss., https://doi.org/10.5194/hess-2018-32
Manuscript under review for journal Hydrol. Earth Syst. Sci.
Discussion started: 13 February 2018

Hydrology and Earth System Sciences
Open Access
Discussions

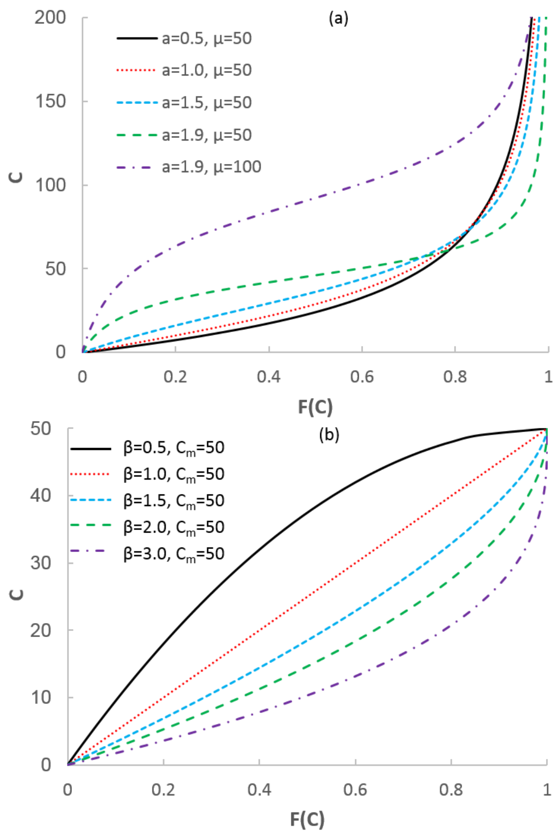

Figure 4: The cumulative distribution functions (CDF) with different parameter values: (a) the proposed distribution function represented by equation (26); and (b) the power distribution of VIC model represented by equation (13).

Hydrol. Earth Syst. Sci. Discuss., https://doi.org/10.5194/hess-2018-32
Manuscript under review for journal Hydrol. Earth Syst. Sci.
Discussion started: 13 February 2018

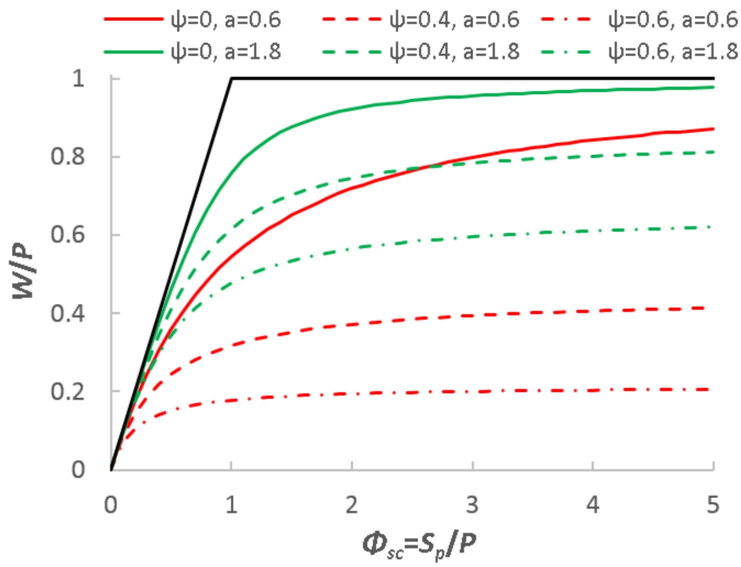

Figure 5: The effects of the degree of initial storage (ψ=0, 0.4, and 0.6) and shape parameter (a=0.6 and 1.8) on soil wetting in the modified SCS-CN method derived from the proposed distribution function for soil water storage capacity.

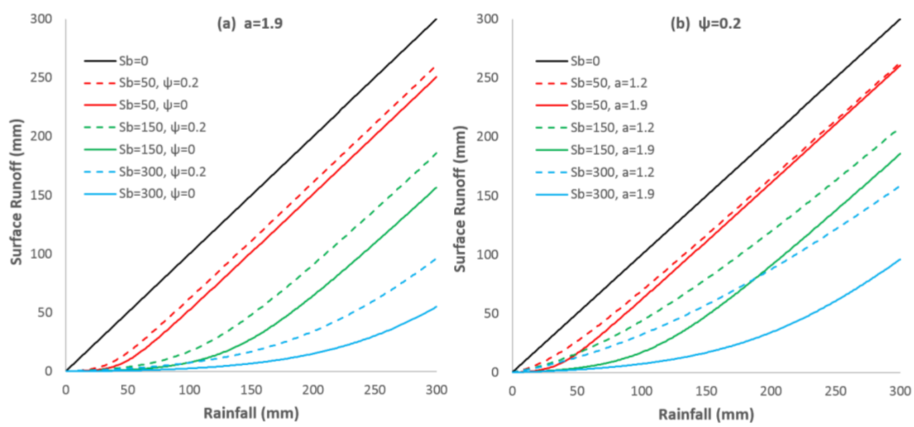

Figure 6: (a) The effects of average storage capacity and initial storage on rainfall-runoff relation; and (b) The effects of average storage capacity and shape parameter on rainfall-runoff relation.