

Interactive comment on “Mountain water cellars: a chemical characterization and quantification of the hydrological processes and contributions from snow, glaciers and groundwater to the Upper Mendoza River basin ($\sim 32^\circ$ S), Argentina” by Sebastián A. Crespo et al.

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Received and published: 6 October 2018

Responses to Referee 1, identified as follows: (1) comments from Referee, (2) author's response, (3) author's changes in manuscript.

Answer to Referee, comment 1.

(1) An English proofreading must be performed for the manuscript, also including the

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figure captures. It is out of scope of this review to address the frequent grammatical deficits or the necessity to rephrase sentences (e.g. “Cuevas, Vacas and Tupungato rivers when join in Punta de Vacas, form the Mendoza River.” (P22L7-8) should be “Cuevas, Vacas and Tupungato rivers form the Mendoza River in Punta de Vacas”. It is unclear, what “The logo of Copernicus Publications” means in this figure capture.). There are many examples which could be provided here.

(2) We agree. The entire manuscript is being reviewed and corrected by a technical reviewer and native English speaker.

(3) New Figure 1 and caption added. The figure captions were streamlined and the specific indication (P22L7-8) was framed by the legend, becoming no longer necessary in the figure caption. The "The logo of Copernicus Publications" sentence was deleted.

Answer to Referee, comment 2.

(1) The manuscript is generally quite long and it would be helpful to streamline the text and to leave out parts which are not essential.

(2) We agree. The entire manuscript was rewritten. We considered the referee’s suggestions, taking out extensive non essential parts of the manuscript.

Answer to Referee, comment 3.

(1) Many measurements were performed and are used in the analysis presented in the manuscript. It would therefore be good to add a section “Study Area and Data Basis” to the manuscript. Here, the authors should add a table summarizing all the measurements performed, which will help the reader to keep an overview. Additionally, the general hydrological characteristics and setting should be described (i.e. long-term mean values of precipitation, discharge, evapotranspiration, temperature etc.). After this overview, it is easier to describe the methods applied, without having to refer to settings of the measurements (as is the case now).

(2) The data description was better explained. Tables with the measured variables were

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already included within the Supplementary data (Tables S2 to S6). According to the previous recommendation of "streamline the manuscript", including them in the main text would extend more the paper. We added the long term variables of precipitation, temperature and Cuevas River streamflow (Fig. S1 and Tables S8 and S9) in the Supplementary data.

(3) See Tables S2 to S6, Figure S1 and tables S8 and S9.

Answer to Referee comment 4.

(1) All locations and regions mentioned in the manuscript, figures and figure captions should be consistent (what is currently not the case, e.g. in Fig. 5 and 6; in Fig. 4, for the first time, the Colorado river and Uspallata Stream are mentioned, without further references in the text).

(2) The consistency between text references, figures and figure captions was optimized. In order to give streamline the manuscript, Figures 4, 5 and 6 were not incorporated in the new version.

Answer to Referee comment 5.

(1) In Fig. 1 an overview map is given. Here, substantial improvements are necessary, also in the context of giving the reader a better overview. In the map, measurement locations presented in the table (see (3)) should be displayed. From the map, the topography is not easily understandable – adding a hillshade layer and using different elevation colours would help in this context. The colours and symbology used should be adopted for easier readability (e.g. The boundary colour of the Horcocones River Basin is basically the same as the ice bodies). "References" should be "Legend". Other relevant information and locations, which are mentioned in the text, should be included in the map (e.g. Punta de Vacas).

(2) Corrected. The locations and information mentioned in the text were incorporated in Fig. 1. "References" was changed to "Legend".

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(3) See new figure 1.

Answer to Referee comment 6.

(1) The quality of the figures needs to be improved. The font sizes are frequently too small and cannot be read easily. The x-axis labels in Fig. 5 & 6 should be consistent. Adding vertical grid lines, e.g. at every month, would help to analyse the temporal dynamics (e.g. begin of snow melt, glacier melt) described in the text. Using colours in Fig. 10, 11 and 12 would be helpful. Also adding a legend in the Fig. 10 and 11 is necessary.

(2) The figures were optimized as recommended. Figures 5 and 6 are not included in the new version.

(3) See new figures 4(ex-10), 5 (ex-11) and 6 (ex-12).

Answer to Referee comment 7.

(1) Figure 3: It is not clear, what is meant with “Ice covered basin efficiency” and “Efficiency related to Cuevas river in %”

(2) Basin efficiency referred to basin yield. “The yield indicates how many times more (or less) water produces a basin per area unit, with respect to the average production of the basin that drains (described before as: "efficiency related to Cuevas River in %"). Figure 3 was deleted. The basin yield regarding the basin area was changed, as suggested in comment 17, to the water contribution, not weighted by area (Table 4).

(3) See Table 4.

Answer to Referee comment 8.

(1) Tab. 1 & 2: No significant variables are marked in bold.

(2) Corrected.

(3) See Tables 1 and 2.

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Answer to Referee comment 9.

(1) Tab. 4: Table caption is one of the examples where a rephrasing is needed. The term “rate per month” is not appropriate.

(2) Corrected.

(3) See new Table 5 (ex 4).

Answer to Referee comment 10.

(1) Methodology: Are all sub-sections necessary?

(2) The sub-sections were streamlined.

Answer to Referee comment 11.

(1) The term for “Height-discharge calibration curves” in Hydrology is “rating curves”. The authors present on P4L25 and P5L3 equations for Mt. Tolosa rock glacier and the Horcones Inferior Glacier. The equations should be consistently formatted as shown in Eq. 1. In general, I think that these equations do not have an added values for the reader.

(2) The term “Height-discharge calibration curves” was modified to “rating curves”. The equations are not longer in the manuscript.

(3) Section 2.1: Streamflow and environmental variables analysis

Two different procedures for each analyzed site were applied for the construction of the rating curves, which links the height measured continuously by the sensor and the volumetric flow. One was by constant saline flow (Gordon et al., 2013) and the other by the velocity–area method (Francou and Poyaud, 2004). Mt. Tolosa rock glaciers conglomerate, composed of four rock glaciers is located between the peaks "Leñas del Tolosa" and the "Morro El Paso", at 32.80°S – 70.01°W (Fig. S2). The rock glaciers range from 3509 to 3749 m a.s.l., with an average altitude of 3614 m a.s.l., a mean

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length of 378 m and a total covered area of 0.16 km². This site was chosen for the simplicity provided by the cryogenic origin of the rock glaciers covering this sub-basin, accessibility throughout most of the year and because it is near (~ 6 km), and in the same geological province, to the Horcones Inferior Glacier. Furthermore, the similarity in orientation (south) and average altitude of both glaciated basins reduces the environmental noise when comparing water delivery dynamics of these different kinds of ice bodies. Because the Mt. Tolosa rock glaciers conglomerate has a low flow, saline tracer constant flow measurements were performed periodically (Gordon et al., 2013). These dissolution methods involve introducing a tracer substance as a chemical tracer (in this case table salt) in the stream and then monitor concentration changes downstream. This method is especially useful in turbulent mountain streamflow (Moore, 2004; Gordon et al., 2013). The salt solution was injected at a constant flow rate using a Mariotte bottle built for that purpose. Using a digital conductivity meter, calibration curves were constructed with measurements of salinity downstream, after pouring the saline solution. The Horcones Inferior Glacier is located in the Aconcagua Mt. Provincial Park (32.73°S and 69.97°W), drains to the Horcones River, a tributary of the Cuevas River, which drains to the Mendoza River (Fig. 1). The glacier is distributed from 3472 to 5460 m a.s.l., with an average altitude of 4151 m a.s.l. It presents southeast orientation, a total length of 12.7 km and 6.95 km² area (Fig. S3). To calculate the glacier streamflow, speed–area rating curves were performed (Francoud and Poyaud, 2004).

Answer to Referee comment 12.

(1) P4L26-27: “The determination coefficient between calculated and observed streamflow was 0.95, for $y = 0.999x - 2E-05$ ”. It is unclear what $y = 0.999x - 2E-05$ means in this context.

(2) In order to give streamline the manuscript as was recommended, these equations are no longer in the manuscript.

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(1) P5L4: For what is this equation? What is the difference between measured flow (Q_m) and calculated flow (Q_{calc}), since Q_m is also calculated as a function of h ?

(2) The equation described the fitted values of the simulation with the measured flow, but since the equation is not considered relevant, we eliminated it from the revised draft.

Answer to Referee comment 14.

(1) P6L3-6: Was the data of these totalizer used in the study?

(2) Yes. These data were used to describe the punctual summer storms isotopic enrichment of the Horcones Inferior Glacier streamflow (Fig. 10, now Fig. 4). The isotopic results are shown in Table S3 and the Table S10 was added for more detail. New references were added in the text.

(3) Text references, Figure 4 and Tables S3 and S10.

Text references:

Section 3.2.2 Temporal variability of snow and glacier contributions

The isotopic composition changes observed in the Horcones Inferior Glacier streamflow, from the beginning to the end of the melting season, responds to the contributions of sporadic enriched summer storm events (Fig. 4 and Tables S3 and S10). For example, samples from January 19th and 22nd show enrichments due to summer precipitations on January 16th to 18th, which felt as light snow in the Horcones Inferior Glacier area (Fig. 4 and Table S7). The January 29th sample shows the enrichment due to the January 26th and 27th summer storms. Then MODIS imagery indicates that all the snow melted, and glacier contribution yielded streamflow stable isotope impoverishment until February 13rd. Less successive snow melt contributions from January 27th storm were detected until February 19th, due to the February 14th to 18th storm. On February 26th, an enrichment is again registered due to February 24th and 25th storms, with successive impoverishing until the next sampling (3/15/2014), to the mean

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values of the proglacial stream. These summer rain or light snow events (Fig. 4 and Tables S3 and S10) increased stable isotope values of streamflow, showing the sensitivity of the measurements.

Answer to Referee comment 15.

(1) P9L6-7: The equations should be formatted as shown in Eq. 1. The naming of the variables should be improved, e.g. TMD air (France) is an odd naming for a variable.

(2) Corrected. All the equations were formatted as recommended and the naming were unified and improved.

(3) New equation:

$$\text{MDS HI} = 0.2221 * \text{MDAT2} + 0.9243 \text{ (Eq. 8)} \quad \text{MDS HI} = 0.2318 * \text{MDAT1} - 0.6489 \text{ (Eq. 9)}$$

For:

MDS: Horcones Inferior Glacier mean daily streamflow (m³ s⁻¹)

MDAT2: mean daily air temperature (°C) in station 2

MDAT1: mean daily air temperature (°C) in station 1

Answer to Referee comment 16.

(1) Section 3.1: What other variables were analysed apart from air temperature? I am surprised that the authors do not mention any relationship between glacier melt and solar radiation.

(2) The influence of environmental variables on the streamflow was modeled with the generalized linear effects model, with significant probabilities. The generalized linear effect model was carried out for all the environmental variables measured in stations 1 and 2 (Table S2, already showed in point 3 in C, referee comment 3), expressed in section 2.1. The results for the Horcones Inferior Glacier were expressed in section 3.1.1. The results for the Tolosa rock glaciers conglomerate are expressed in 3.1.2.

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(see 3). We also expected solar radiation statistical significance. Solar radiation may influence the ice melting but, during the period studied in the austral summer, the effect was not significant. The temperature (which probably includes the solar radiation effect), was significant.

(3) Text references:

Section 2.1: Streamflow and environmental variables analysis

The Irrigation General Department of Mendoza weather station (labelled as 1 in Fig. 1), located at 3043 m a.s.l. in “Laguna de Horcones” (32.80°S – 69.95°W), measured: air temperature, soil temperature, wind speed and direction, relative humidity, incident radiation and snow water equivalent, hereinafter “station 1”. Air and soil temperature HOBO sensors were also installed in the Horcones Inferior Glacier (labelled as 2 in Fig. 1), at 4016 m a.s.l. (32.69°S – 69.97°W), hereinafter “station 2” (Table S2). Both stations covered an altitude gradient of 973 m.

Generalized linear models (nlme package, Pinheiro et al., 2013) in the R program (R Core Team, 2013) were conducted. The streamflow was considered the response variable, as a function of environmental data (predictor variables).

Section 3.1.1: Horcones Inferior Glacier

Streamflow of the Horcones Inferior Glacier showed a similar variability as that of temperatures (Fig. 2a). The best overall linear model obtained for the total measured variables in both, stations 1 and 2 (Table S2), includes as significant variability only for the mean daily air temperature ($p < 0.01$). The Horcones Inferior Glacier average daily streamflow variability, fits linearly with respect to the mean daily air temperature for both stations, following equations 8 ($R^2 = 0.7$) and 9 ($R^2 = 0.6$), respectively:

$$\text{MDS HI} = 0.2221 * \text{MDAT2} + 0.9243 \text{ (Eq. 8)}$$

$$\text{MDS HI} = 0.2318 * \text{MDAT1} - 0.6489 \text{ (Eq. 9)}$$

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For:

MDS: Horcones Inferior Glacier mean daily streamflow (m³ s⁻¹)

MDAT2: mean daily air temperature (°C) in station 2

MDAT1: mean daily air temperature (°C) in station 1

Section 3.1.2.: Mt. Tolosa rock glaciers conglomerate

From all the environmental variables measured in stations 1 and 2 (Table S2), the most influential variables in the emergent flow of the analyzed rock glacier cluster, were mean daily air temperature ($p < 0.01$) and mean daily maximum air temperature ($p = 0.027$), both corresponding to station 2 ($R^2 = 0.56$). A generalized linear modeling was performed based on this result, considering the response variable (average daily flow) regarding just those significant variables ($R^2 = 0.49$). Subsequently, an inference of models with elimination of variables according to their significance was followed to determine the relative importance of each predictor variable, in order to simplify the model to the minimum number of variables explaining the streamflow. The most significant predictor variable was the mean daily air temperature measured in station 2 ($R^2 = 0.62$), adjusted through a third-order polynomial equation (Eq 11):

MDS T= $4E-5x^3 - 9E-5x^2 - 0x + 0.007$ (Eq. 11)

For:

MDS T: Tolosa rock glacier conglomerate mean daily streamflow (m³ s⁻¹)

x: mean daily air temperature (°C) in station 2

Certain threshold, around 6 °C (Fig. 2b), is needed for a higher flow delivery rate, as the air temperature increases. The isolation created by the debris layer, which in turn makes a more delayed thermal inertia for the glacier ice thawing (Østrem, 1965; Buk, 2002; Trombotto and Ahumada, 2005), could explain this behavior.

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Answer to Referee comment 17.

(1) Section 3.2.1: What does “Glacier covered basins performance” mean? Does the mentioned Cuevas River Basin here refer to Punta de Vacas? Some interesting results, e.g. what is the percentage (not weighted by area) of the contribution of the glacier to the total runoff, is not given here, but can only be found in the discussion. Maybe it would also make sense to merge the results and discussion part.

(2) Glacier covered basins performance indicates how many times more (or less) water produces a basin per area unit regarding the average production of the basin that drains. The basin yield regarding the basin area was changed, as suggested, to the water contribution not weighted by area (Table 4).

The reference to the Cuevas River measured in Punta de Vacas was now incorporated to the manuscript.

The percentages not weighted by area were now included (Table 4) and referred in section 3.1.3 and 3.2. Merging these results with EMMA (Table 5) was very synergetic and the discussion (section 3.2) was enriched (text parts added in point 3).

(3) References in text. See Tables 4 and 5.

Section 3.1.3 Glaciated basins streamflow contributions

The Horcones Inferior Glacier basin streamflow contributes 22 Hm³ along the melting period (Dec-Mar). Tolosa rock glacier conglomerate basin, contributes 0.15 Hm³ for the same period. These basins absolute contributions represent 38 and 0.25% of the Cuevas River streamflow in Punta de Vacas (58 Hm³), respectively (Table 4) for this melting period. The contribution of the Horcones Inferior Glacier is relatively high, considering that it represents only one of the 190 crioforms that discharge to the Cuevas River basin (IANIGLA-ING, 2018a).

Section 3.2

When analyzing the relative contribution from each glaciated basin to the Cuevas River for each month, the streamflow (Table 4) and the EMMA results (Table 5) yield different information. The streamflow measured at the Tolosa rock glaciers conglomerate represents 0.21 and 0.11% of the Cuevas streamflow for January and February, respectively (Table 4), while the EMMA shows 47 and 34% input from rock glacier source for the same months (Table 5). The small area of this conglomerate may explain the low contributions estimated with direct streamflow measurements. The larger estimates of rock glaciers input estimated with EMMA may be attributed to other rock glaciers draining the Cuevas River. For glacier contributions estimates, considering only the Horcones Inferior Glacier, just one of the 190 circoforms in the basin (IANIGLA-ING, 2018a), the measured streamflow represents 42.43 and 34.10% (for January and February, respectively) of the Cuevas River streamflow measured in Punta de Vacas (Table 5). For the same months, the EMMA estimates 8 and 9% input from the glacier source (Table 5). In this case, contributions estimated with direct streamflow measurements of one glacier are much larger than the estimated with EMMA, pointing to other processes, not considered with any of the approaches.

The water delivered by glaciers might infiltrate through deep fractured aquifers and to the soil, generating a large groundwater matrix draining in the lower basin area, where the Cuevas River flows (Fig. 7). This process may change ion chemistry during water transport to the Cuevas River sampling site, increasing salinity and emerging in EMMA as groundwater source. In addition, between the glacier and the Punta de Vacas Cuevas River sampling site (in Puente del Inca), deep thermal groundwater flow to the river, probably changing its ion concentration. These results imply that the 32, 45 and 56% of groundwater contributions obtained with EMMA for December, January and February, respectively (Table 5), are composed of old water infiltrated to aquifers. These old water sources may have originated from snow, glacier or permafrost melting, during unknown periods of time. In EMMA with two tracers and three components, the three components form the vertices of a triangle and all the river samples must be framed by the triangle. If samples are located outside the triangle, as the Cuevas

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River in March, it means either that the tracers are not conservative, or there may be contributions from additional sources (Fig. 6). Puente del Inca geothermal waters may represent this additional source (Fig. 1 and 6).

The Puente del Inca geothermal waters were confined in a very narrow region of the scatter plot (Fig. 3). This stable isotopes low dispersion may indicate the isolation of water, compared to the surface waters, without being affected by the fluctuations in precipitation or water melting from snow or ice bodies. The deviation to the right of the global meteoric line indicates an enrichment in ^{18}O , probably due to prolonged isotopic exchange with the rocks at temperatures between 25 and 100 °C (Craig, 1963; Aggarwal et al., 2007). Puente del Inca geothermal water presented stable temperatures of 33°C in all samples. According to this hypothesis, the EMMA results for March, where the Cuevas River waters are outside of the triangle, can be explained by the oxygen enrichment caused by the Puente del Inca geothermal waters input (Fig. 4, 6 and 7).

In the Aconcagua River of Chile, situated at the same latitude as the Mendoza River basin, an hydrograph separation, resulting from an EMMA analysis for the 2011-2012 melting period, was done by Rodriguez et al. (2014). They show a December snow contribution of 19-25% to the Juncal River, different from the 48% obtained in this work. In the Chilean work they do not discriminate between the glaciers from the periglacial (rock glaciers) and from the glacial environments. For that large group of glacialized sources considered in that work, they calculate a contribution of 51-55% in the spring, while for our study it was 19.7% (December month, Table 5). A closer result was observed for subsurface sources, where the Juncal River basin contribution was around 20-30% to the spring flow, and 32% in the Cuevas River. Glacial contributions during the summer increased to 58-66% of the seasonal flow in the Rodriguez et al. study, and in our results were 43-55% (considering glaciers and rock glaciers). The underground sources for the summer were calculated for the Rio Juncal 2011-2012 study at 34-42%, while in our study it was between 45-56%.

Although the estimates for both basins differ because of a different time period and

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geographic location of the studies, both point to glaciers as important contributor to river flow.

Answer to Referee comment 18.

(1) Section: 3.2.2.: The naming of the section needs rephrasing. For what is Fig. 4d, since it is not referred to in the text?

(2) In order to streamline the paper, this old section 3.2.2 was deleted, as Figure 4 and many others.

Answer to Referee comment 19.

(1) Section 3.4: Fig. 13 is not in the manuscript.

(2) Corrected, it referred to figure 11 (now Figure 5).

(3) See Figure 5 (ex-11).

Answer to Referee comment 20.

(1) It would be good to change the “Conclusions” into “Summary and Conclusion” and to add here the main findings and quantitative results from the manuscript, e.g. the contributions of the different water sources to the total flow. These numbers are also missing in the abstract.

(2) Changed “Conclusions” into “Summary and Conclusion” as recommended. Main findings and quantitative results were added in "Summary and Conclusion" and in the abstract.

(3)

Abstract. The Mendoza River flow provides fresh water for more than 1.1 million inhabitants in an agriculture based arid region in the Monte desert. Most of the Mendoza River streamflow derives almost exclusively from winter snow precipitation fell in Cordillera Principal, originated from the Pacific Ocean moisture. In addition to the snow

that precipitates in this area of 3023 km², there are 951 glaciers, covering an area of 404 km². Given the high inter-annual variability of snowfall (ranging from 5 to 240% of the long term mean records), strongly affected by ENSO events, and the aridity of the region, it is crucial to quantify the contribution from different water sources to the Mendoza River flow. Glaciers play an important role regulating water availability, with mass accumulation in wet and cold years, and melting in hot, dry years. Combining instrumental records of streamflow from glaciers and rivers, meteorological data, remote sensing of snow covered area and chemical analysis of different water sources, this study attempts to understand the hydrological contribution of different water sources to the Cuevas River, the original tributary of the Mendoza River. Isotopic and ion composition allowed us to differentiate snowmelt from glacier ice melt. In addition, it was possible to detect contributions of summer storms from Atlantic origin that occasionally reach the Cordillera Principal. Finally, with end member mixing analysis, the relative contribution of different water sources was estimated over time, showing the contribution of glacial and periglacial environments as the melting season progressed. Groundwater input to the total flow showed relatively large contributions, 32, 45 and 56% for December, January and February, respectively, pointing to the importance of this water source on maintaining Cuevas River flow.

Section 5: Summary and conclusions

Distinct water sources differ in composition along a melting season for both, stable water isotopes and ion chemistry. These findings reinforce what had been observed at a seasonal scale by Crespo et al. (2016) and allows us to distinguish and estimate the relative contribution of snow, groundwater, and ice bodies to rivers streamflow.

The convective summer precipitation events were detected in glaciated basin streams by a distinct water stable isotope composition, allowing to pair singular flow increases to particular summer storm events. The snow to ice contribution transition could also be detected in a sub weekly resolution. Air temperature significantly modulates the glaciated basins streamflow along the melting period, presenting a delayed thermal

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inertia for rock glaciers.

Periglacial (rock glaciers) and glacial environments contributions were relevant during this analyzed dry year, with ice bodies and groundwater contributing most to the Cordillera Principal rivers streamflow in this relatively dry period analyzed. Natural tracers indicated at least, 8-9% of the water was from glacial sources, 34-47% from rock glacier melting and between 32 to 56% from groundwater system. Direct streamflow measurements indicated large discharges for a particular glacier, Horcones Inferior, which was not detected as a proportional glacier contribution downstream at the Cuevas River, with natural tracers. This discrepancy suggests that an important proportion of water derived from glacier melting is infiltrated to groundwater, where it increases ionic composition, and then discharges downstream with the chemical signal of groundwater. Thus, the major contributor to the Cuevas River obtained with natural tracers is groundwater, which may reflect delayed water inputs incoming from glacial and periglacial environments, with some contributions of geothermal groundwater. To estimate groundwater residence times, and validate the conceptual model developed here, other tracers could be used, such as radioactive isotopes (tritium, ^{14}C) and CFCs, SF₆, etc. analysis. Those kinds of analysis were out of the scope within this investigation.

The present investigation is the first scientific work estimating glacial, periglacial and groundwater contributions to the Mendoza River basin, reinforcing previous assumptions about the importance of ice bodies to maintain river flows. Furthermore, this study points to the importance of glacier-groundwater-river relations, and the need of additional groundwater studies to better map strategic water source areas to be protected, in addition to those included in the National Glacier Inventory.

Answer to Referee comment 21.

(1) In section 5, the authors write “By deepening our understanding about the delivery and depletion dynamics of different water sources influencing the hydrological

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processes of the Mendoza River basin, the vital artery of this arid and fragile territory, these tools are expected to serve to decision makers and to generate the necessary mitigation policies for an improved water resources administration along the actual and future climate change scenarios.” (By the way - Word is giving me a warning: “Long sentence (consider revising)”). It is unclear, what is meant with “tools” in this context. What would be an example, how decision makers can generate the necessary mitigation policies for an improved water resources administration along the actual and future climate change scenarios from the results? Are these not empty phrases, which the world anyway has too many?

(2) Modified and simplified as recommended. The whole section was rewritten.

(3) See Section 5 in point 3, Referee comment 20.

Please also note the supplement to this comment:

<https://www.hydrol-earth-syst-sci-discuss.net/hess-2018-212/hess-2018-212-AC2-supplement.pdf>

Interactive comment on Hydrol. Earth Syst. Sci. Discuss., <https://doi.org/10.5194/hess-2018-212>, 2018.

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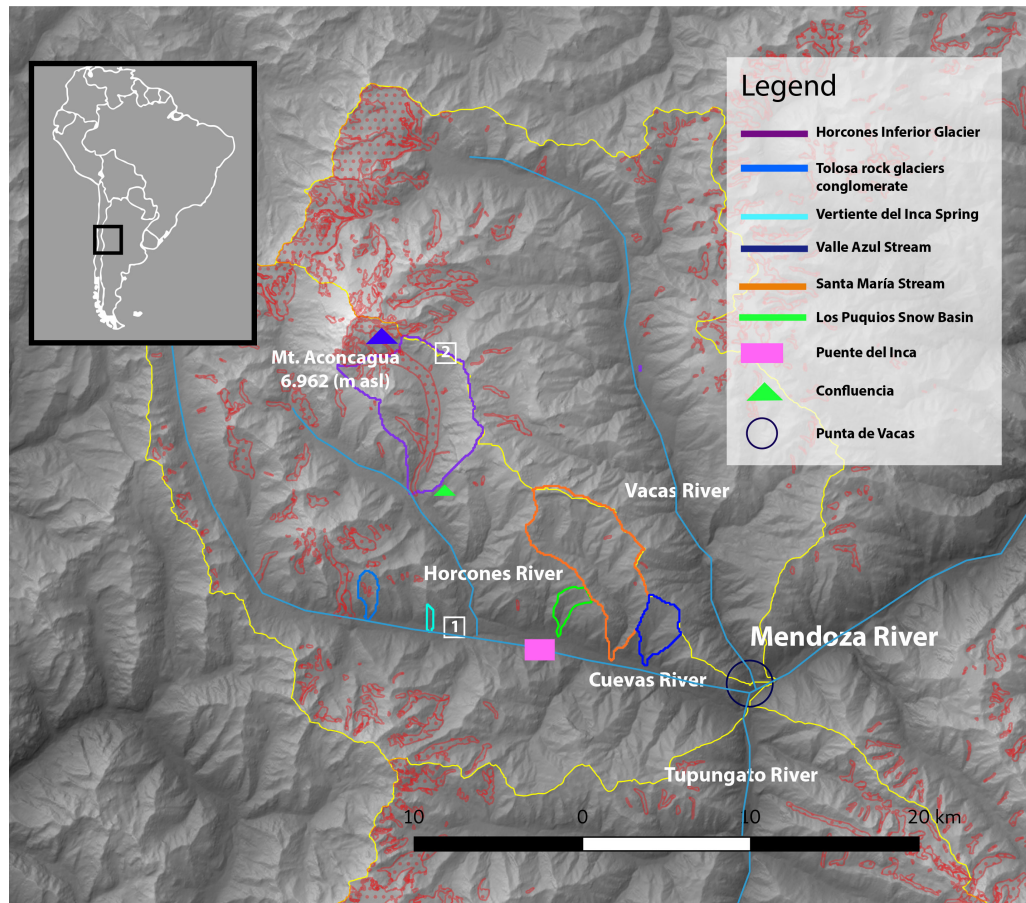


Fig. 1. Map with the digital elevation model, sampling sites and ice bodies. Glacier shapes (marked with red contour) were taken from the glacier official inventory (IANIGLA–ING, 2018a and 2018b).

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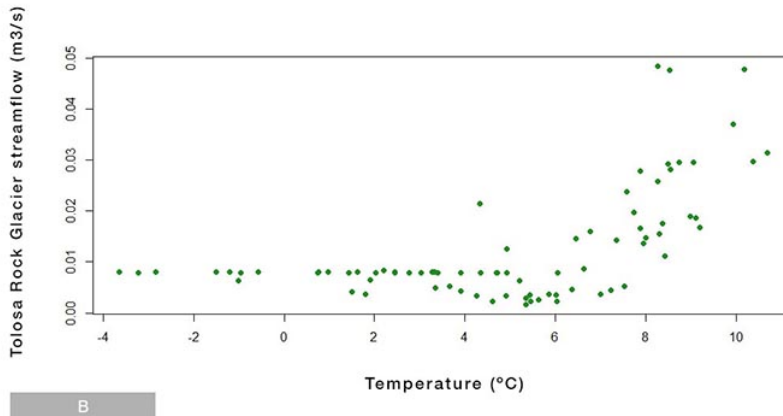
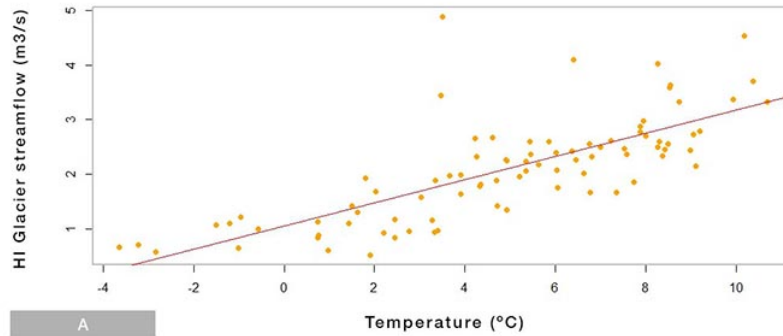


Fig. 2. Scatter plot of streamflow and mean daily air temperature for Horcones Inferior Glacier (A) and for Mt. Tolosa rock glaciers conglomerate (B).

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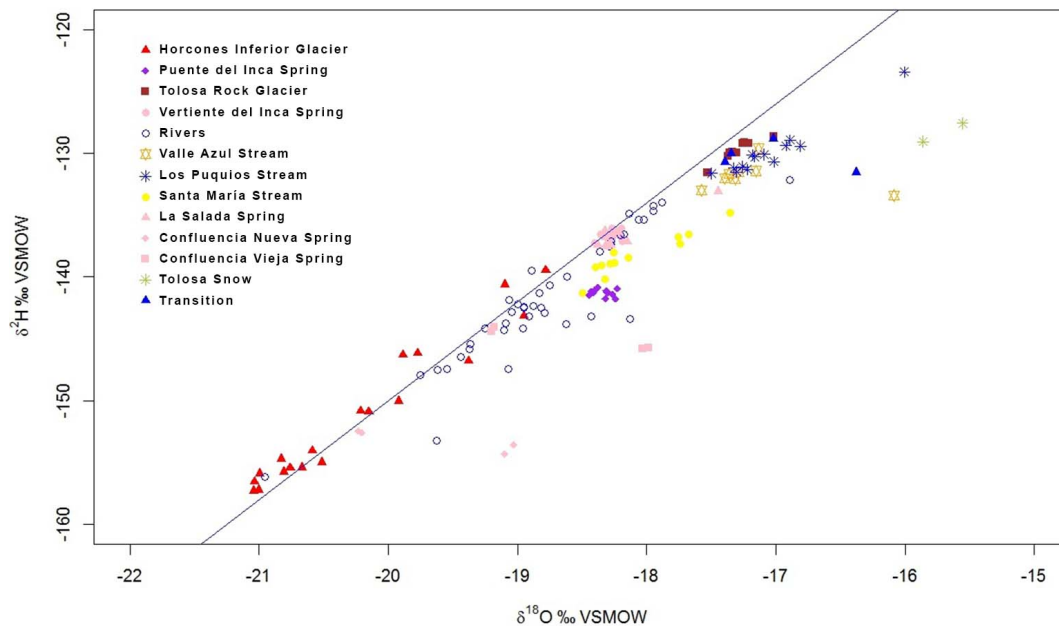


Fig. 3. $\delta^{18}\text{O}$ and $\delta^2\text{H}$ values scatter plot of the analyzed samples. The adjusted line is the global meteoric water line (Craig, 1961).

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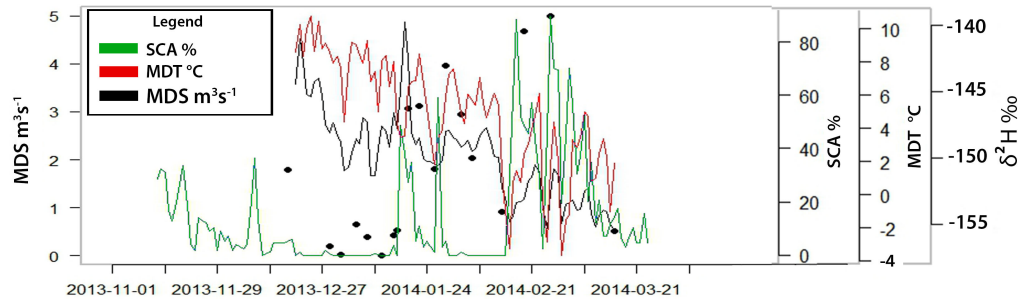


Fig. 4. Time-series showing the snow covered area (SCA), mean daily streamflow (MDS), mean daily air temperature (MDT) and $\delta^2\text{H}$ composition of the Horcones Inferior Glacier (bold black dots).

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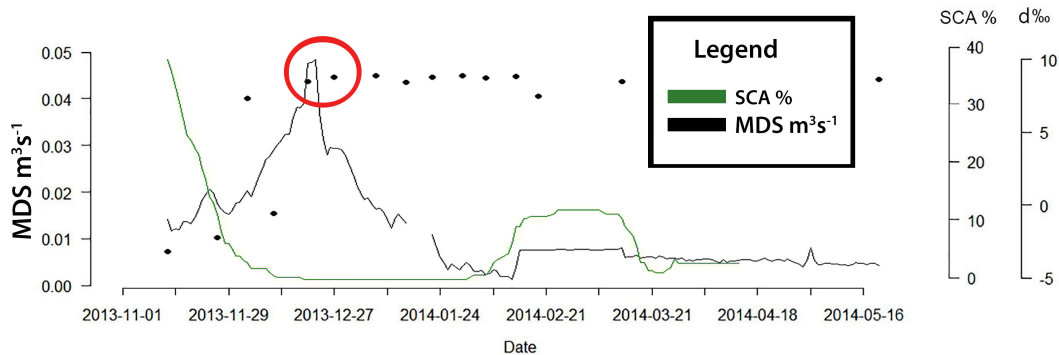


Fig. 5. Mean daily streamflow (MDS) of Mt. Tolosa rock glaciers conglomerate and snow covered area percentage (SCA). The dots are the deuterium excess values (d‰).

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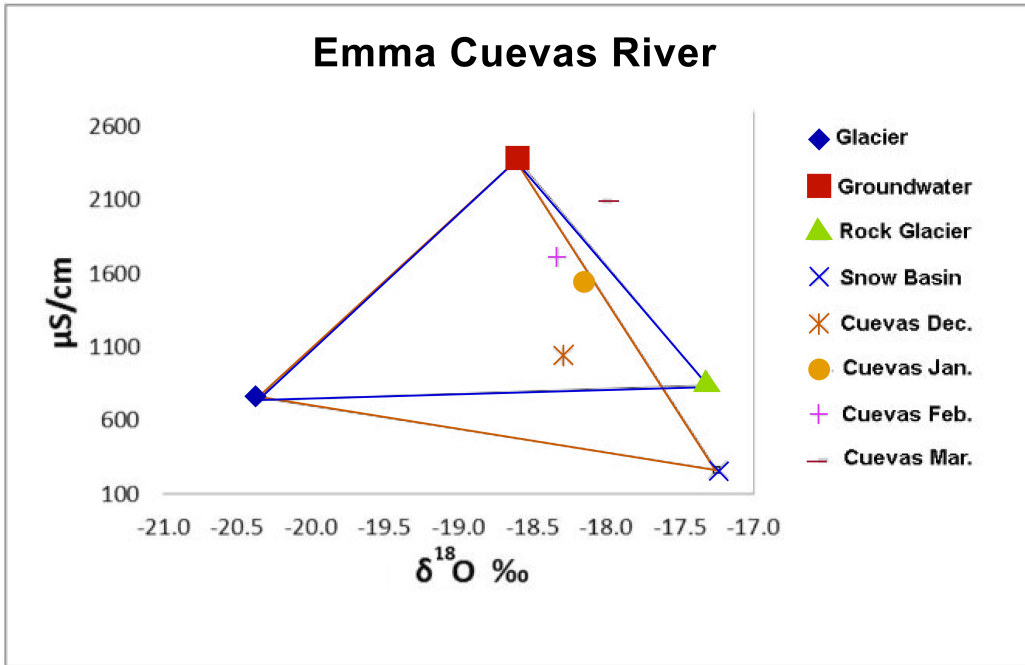


Fig. 6. Dispersion plot of mean oxygen stable water isotope composition ($\delta^{18}\text{O}$) and electrical conductivity of different water sources draining to the Cuevas River.

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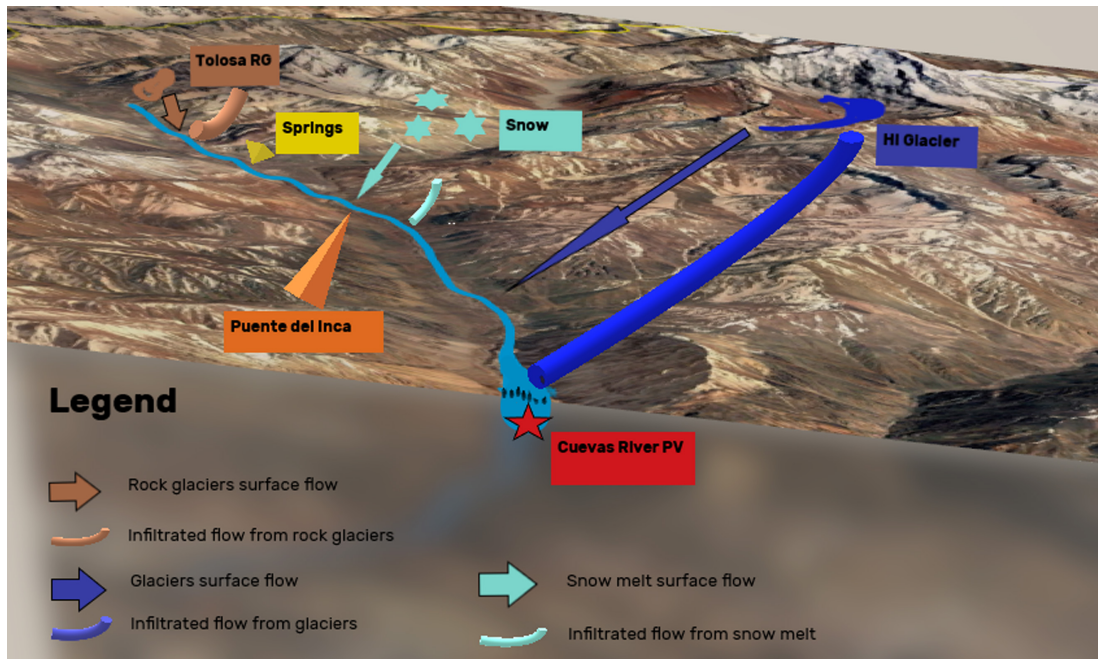


Fig. 7. Hydrological conceptual model.

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Table 1: Mean, standard deviation (SD) and confidence interval (CI) for electrical conductivity of different water sources in comparison with the intercept (Horcones Inferior Glacier). The significant codes are: 0; **** 0.001; ** 0.01; * 0.05; . 0.1. The significant variables are marked in bold.

Variable	Mean	SD	CI (2.5–97.5%)	p
Intercept	674.1	458.9	-225.41 & 1573.62	0.14
Summer precipitation	-289.1	775.1	-1808.22 & 1230.01	0.71
Rock glacier	144.3	742.2	-1310.32 & 1598.97	0.85
River	447.6	561.7	-653.23 & 1548.49	0.43
Valle Azul snow basin	-243.5	744.1	-1701.92 & 1214.96	0.74
Los Puquios snow basin	-527	742.4	-1982.06 & 928.02	0.48
Sta. María basin	-440.3	742.6	-1895.76 & 1015.22	0.55
Groundwater	1495	587.6	343.29 & 2646.65	0.01 *

Fig. 8. Table 1

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Table 2: Mean, standard deviation and confidence interval $\delta^{18}\text{O}$ composition for water source and the intercept (Horcones Inferior Glacier). The significant codes are: 0; '****' 0.001; '***' 0.01; '**' 0.05; '.' 0.1. The significant variables are marked in bold.

Variable	Mean	SD	CI (2.5–97.5%)	p	
Intercept	-20.23	0.56	-21.33-19.13	$<2e^{-16}$	
Puente del Inca	1.89	0.79	0.34 & 3.45	0.0168	*
Summer precipitation	7.63	0.92	8.83 & 9.43	$<2e^{-16}$	***
Rock Glacier	3.22	0.89	1.48 & 4.96	0.0003	***
Rivers and streams	1.43	0.68	0.09 & 2.76	0.0353	*
Valle Azul snow basin	3.07	0.91	1.28 & 4.86	0.0008	***
Los Puquios snow basin	3.18	0.89	1.43 & 4.93	0.0004	***
Santa María basin	2.15	0.89	0.39 & 3.90	0.0165	*
Groundwater	1.61	0.71	0.21 & 3.01	0.0238	*

Fig. 9. Table 2

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Table 3: Water sources and Cuevas River mean $\delta^{18}\text{O}$ ‰ and electrical conductivity.

Component/Tracer	$\delta^{18}\text{O}$ ‰	CE $\mu\text{S}/\text{cm}$
Glacier	-20.4	762
Groundwater	-18.6	2382
Rock glacier	-17.3	873
Snow basin	-17.2	265
Cuevas River December	-18.3	1037
Cuevas River January	-18.1	1540
Cuevas River February	-18.3	1710
Cuevas River March	-18.02	2094

Fig. 10. Table 3

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Table 4: Percentage contribution from different kind of ice covered basins to the Cuevas River since December 2013 to March 2014. % of Cuevas River refers to the % of streamflow regarding the Cuevas River. Sources: Cuevas River in Punta de Vacas streamflow: Secretariat of Water Resources, 32.86° S and 69.77°W), Horcones Inferior Glacier (HIG) and Tolosa Rock glacier conglomerate (Tolosa RGC) streamflow were measured in this study.

	Cuevas River			HIG		Tolosa RGC
	m ³ month ⁻¹	Hm ³ month ⁻¹	Hm ³ month ⁻¹	% of Cuevas River	Hm ³ month ⁻¹	% of Cuevas River
Dec	19,316,621	19.32	8.94	46.28	0.081	0.42
Jan	15,574,896	15.57	6.61	42.43	0.033	0.21
Feb	12,369,370	12.37	4.22	34.10	0.014	0.11
Mar	10,984,118	10.98	2.50	22.77	0.019	0.17
Total	58,245,005	58	22		0.15	

Fig. 11. Table 4

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Table 5: Percentage contribution from different water sources to the Cuevas River since December 2013 to February 2014, estimated with natural tracers (EMMA).

Water Source	Dec	Jan	Feb
Glacier	19.7	8	9
Groundwater	32	45	56
Rock Glacier	0	47	34
Snow	48	0	0

Fig. 12. Table 5

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