1 Extending seasonal predictability of Yangtze River summer floods

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Abstract. Extreme pluvial floods across China's Yangtze River basin in the summer of 2016 werewas 10 strongly connected with intense atmospheric moisture transport, and resulted in vast loss of properties 11 after a strong El Niño winter. Predicting such extreme floods in advance is essential for hazard 12 mitigation, but the flood forecast skill is relatively low due to the limited predictability of summer 13 precipitation. By using a "perfect model" assumption, here we show that atmospheric moisture flux has 14 a higher potential predictability than precipitation over the Yangtze River at seasonal time scales. The 15 predictability of precipitation and moisture flux are-is higher in post-El Niño summers than those-in 16 post-La Niñas, especially for flooding events. As compared with extreme precipitation, the potential 17 detectability of extreme moisture flux increases by 20% in post-El Niño summers, which suggests that 18 atmospheric moisture flux could be crucial for early warning of Yangtze River summer floods. 19

21 1. Introduction

Located in eastern China with dense population and major agricultural and industrial productions, the 22 Yangtze River basin suffers from frequent flooding due to large interannual variability of the East Asian 23 summer monsoon. In June-July of 2016, extreme pluvial floods hit the middle and lower reaches of the 24 Yangtze River, caused severe inundations over many big cities, and resulted in direct economic loss of 25 70 billion RMB (about 10 billion U.S. dollars) (Yuan et al., 2018). Effective early warning of upcoming 26 extreme flood events is urgent to mitigate the potential damages, and which strongly depends on 27 accurate precipitation forecasts not only at synoptic- but also subseasonal-to-seasonal scales (Yang et al., 28 29 2008; Tian et al., 2017). However, predicting flood at seasonal time scales is still a grand challenge due to limited forecast skill in precipitation at long leads (Alfieri et al., 2013; Yuan et al., 2015). This raises 30 the interests to explore other relevant variables that are more predictable than precipitation for flood 31 early warning. 32

Predictability is an inherent property of the climate system, and it represents the "ability of the model to 33 "predict itself" be predicted" (Boer et al., 2013). As for a numerical prediction model, it is widely 34 accepted that we cannot improve the (precipitation) predictability without improving its dynamical 35 framework, data assimilation and/or physical parameterizations, etc (e.g., Barnston et al., 2012). 36 However, most of the heavy precipitation and flood events in many mid-latitude regions, especially in 37 coastal areas, are strongly related to intense horizontal atmospheric moisture transport (Banacos and 38 Schultz, 2005; Ralph et al., 2006; Lavers et al., 2014). The atmospheric moisture flux is supposed to be 39 better predicted by large-scale climate models than precipitation that is not only connected to mesoscale 40 (or more local scale) circulation but also influenced by the vertical convection and the localized 41

convections (Lavers et al., 2014, 2016b). This 42 orography provides a potential to use atmospheric moisture flux to extend the predictability of floods. Recently, a 43 series of studies (Lavers et al., 2014, 2016a, 2016b) have assessed the varying predictability of 44 precipitation and moisture flux in winter, and shown that moisture flux yields a higher predictability 45 than precipitation at synoptic-scales (less than two weeks) across northwest Europe and western U.S. 46 that are known as affected by atmospheric rivers. At sub-seasonal to seasonal time scales, however, 47 whether such moisture flux and precipitation predictability relation also applies in China's monsoonal 48 summer seasons where convection is active, such as the Yangtze River summer flood, is still unclear. 49

The middle and lower reaches of the Yangtze region-River basin in eastern China is one of the most 50 strongly El Niño-Southern Oscillation (ENSO)-affected regions in the world (e.g., Wang, 2000; Wu et 51 al., 2003; Ding and Chan, 2005). The persistent Sea Surface Temperature (SST) anomalies in the 52 equatorial eastern Pacific can alter the tropical and subtropical circulations via local air-sea interaction 53 and/or teleconnections, and thus affect the East Asia summer climate significantly, including the 54 summer precipitation in the Yangtze region. Such ENSO-related climate anomaly in the Yangtze region 55 is not concurrent with the ENSO cycle, but hasat a seasonal lag. A possible mechanism for this lag-56 impact of ENSO on East Asia summer climate is the Indo-western Pacific ocean capacitor (IPOC), 57 where the North Indian occan warming after El Niño plays a crucial role (e.g., Xie et al., 2016). 58 Therefore, the precipitation predictability over the Yangtze River is closely associated with the 59 atmospheric and oceanic conditions, which is similar to other regions (Gershunov, 1998; Kumar and 60 Hoerling, 1998; Lavers et al., 2016a). For instance, Kumar and Hoerling (1998) indicated that the North 61 American climate is most predictable during the late winter and early spring seasons of the warm ESNO 62

events. Lavers et al (2016a) showed that the moisture flux and extreme precipitation have different
prediction skill predictability during different North Atlantic Oscillation (NAO) phases. In short, the
weather or climate forecasts initialized at different atmospheric/oceanic conditions can have varying
levels of predictability, so understanding how the Yangtze River rainfall predictability varies during
different ENSO phases is also a concern.

In present study, we aim to address the above questions by evaluating the seasonal predictability of precipitation and moisture flux for the middle and lower reaches of Yangtze River (110°-123°E, 27°-34°N) based on multisource observational data, and ensemble hindcasts and real-time forecasts from a dynamical seasonal forecast model Climate Forecast System version 2 (CFSv2; Saha et al., 2014) for the period of 1982-2016.

73 2. Data and Method

74 2.1 Observation and reanalysis data

Monthly mean precipitation data at $1^{\circ} \times 1^{\circ}$ resolution over <u>the</u> Yangtze River basin was obtained from NOAA's precipitation reconstruction over land (PREC/L), which agrees well with gauge-based datasets (Chen et al., 2002). Monthly mean atmospheric fields including geopotential height, u-wind, v-wind, and specific humidity at <u>different pressure levels300, 400, 500, 700, 850, 925 and 1000 hPa</u> were derived from the ERA-Interim reanalysis (Dee et al., 2011). Herein, the mean June-July zonal and meridional atmospheric moisture fluxes between 300 and 1000 hPa were calculated separately, and their magnitudes were combined as the total moisture flux (Lavers et al., 2016a).

82 NINO3.4 (5°S–5°N, 120°–170°W) SST anomaly based on ERSSTv4 monthly data (Huang et al., 2014)

83 during 1948–2016 was used to analyze the impact of ENSO on the seasonal predictability of rainfall

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and moisture <u>flux</u> over the Yangtze River. An ENSO event was defined as the averaged NINO3.4 SST
 anomaly during preceding December-January-February (DJF) exceeding its 0.5 standard deviation (σ).

86 2.2 CFSv2 seasonal hindcast and real-time forecast data

The ensemble hindcast and real-time forecast datasets including the monthly specific humidity and wind 87 field at different levels and monthly precipitation from Climate Forecast System version 2 (CFSv2) 88 (Saha et al., 2014), were used here to quantify the potential predictability. The predicted moisture flux 89 was calculated the same as the observation mentioned in Section 2.1. CFSv2 has 24 ensemble members 90 with different initial conditions (Yuan et al., 2011) and has been widely used for subseasonal to 91 92 seasonal forecasting (e.g., Kirtman et al., 2014; Yuan et al., 2015; Tian et al., 2017). All monthly anomalies were calculated based on the climatology from the entire hindcast period (1982-2010). The 93 0.5-month lead forecast ensembles started from mid-May to early June (Saha et al., 2014), and predicted 94 through June-July. Similarly, the 1.5-month lead forecasts for the June-July started from the mid of 95 96 April, and so on.

97 In order to investigate the predictability at finer temporal resolution (e.g., weekly mean fields), the 98 CFSv2 daily reforecasts were also obtained from the Subseasonal to Seasonal (S2S) prediction project 99 for the period of 1999-2010, with the forecast lead times up to 45-days (Vitart et al. 2017). As for the 90 June 1-7 weekly mean fields, the reforecasts started from May 18 were used as the first ensemble 91 member, the reforecasts started from May 19 were used as the second, and so on. This resulted in 14 92 ensemble members, with forecast lead times from 1-day to 14-days. The above process was repeated for 93 other weekly averaged fields during June and July. This is called as the first group of ensemble subseasonal forecasts, with lead times of 1-14days. The second group of ensemble reforecasts started

105 from 17 May, 18 May ..., and 30 May were formed similarly, with lead times of 2-15 days, and so on.

106 2.3 The potential predictability approach

The potential predictability was quantified by using a "perfect model" assumption (Koster et al., 2000, 2004; Luo and Wood, 2006; Becker et al., 2013; Kumar et al., 2014; Lavers et al., 2016b). For the predictions of June-July mean precipitation and moisture <u>flux</u> over each grid cell within <u>the</u> Yangtze River basin (110°-123°E, 27°-34°N) at a given lead time, ensemble member 1 was considered as <u>the</u> observation and the average of members 2–24 was taken as the prediction, which resulted in two time series with 35 years of record (1982-2016). The skill of this forecast was then calculated by using the anomaly correlation (AC; Becker et al., 2013) between these two time series, which is defined as

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$$\underline{AC} = \frac{\sum X' Y'}{\left[\sum (X')^2 (Y')^2\right]^{1/2}} \underline{AC} = \frac{\sum X' Y'}{\left[\sum (X')^2 (Y')^2\right]^{1/2}}, \text{ where } X' \text{ is the "observed" precipitation/moisture flux}$$

anomaly and Y is was the predicted counterparts. Here, the 95% (90%) significant level is 0.33 (0.22) 115 for AC according to a two-tailed Student's t-test. Figure 1 gives an example of the potential 116 predictability calculation at a grid near Wuhan city, where the ensemble member 1 was taken as the 117 truth and the mean of the members 2-24 was the prediction. Result shows that moisture flux has a 118 higher predictability (AC) than precipitation at 0.5- and 1.5-month lead for member 1. This method was 119 repeated 24 times, with each member being considered as the observation, so as to obtain 24 AC values; 120 the average of these 24 values was the final estimate of the potential predictability. In addition to the 121 calculation for individual grid cells, AC value was also calculated by using both spatial and temporal 122

samples for the Yangtze River basin with 72 CFSv2 grid cells. Here, an AC higher than 0.05 would be
considered as significant at 95% confidence level, both for ENSO events and the entire period.

The rationale for this "perfect model" approach is that the statistical characteristics of the "observation" (one of the ensemble members) and the prediction (ensemble mean of remaining members) are the same, so the estimate of potential predictability is not affected by model biases (Koster et al., 2004; Kumar et al., 2014).

Generally, potential predictability is considered as the upper limit of forecasting skill, with an
 assumption that internal physics or at least the statistical characteristics of observation and model
 prediction are identical (Koster et al., 2004; Kumar et al., 2014).

In addition, the hit rate (HR) was also used to assess the seasonal predictability for extreme hydrologic 132 events (Ma et al., 2015), where the flooding condition was defined as the June-July mean precipitation 133 or moisture flux greater than 90th percentile of their climatology. Here, a forecast for flooding event can 134 be counted at a given grid or region when taking ensemble member 1 as observation and the average of 135 members 2–24 as the prediction: the HR was computed as $HR = \frac{a}{a+c} HR = \frac{a}{a+c}$, where *a* represents the 136 number of events that flooding is forecast and observed, b for flooding is forecast but not observed, and 137 c for observed flooding that is not forecast. Similar to the AC calculation, 24 HR values would be 138 obtained when each member was considered as the observation, and their average HR values was the 139 final potential predictability for extreme hydrologic events. 140

141 3. Results

142 3.1 Yangtze River 2016 pluvial flood and its associated atmospheric circulation

Figure 2a shows the spatial distribution of the 2016 June-July mean rainfall anomaly. Extreme pluvial flooding hit the middle and lower reaches of Yangtze River, where the area averaged precipitation increased by about 40% relatively to the climatology. In particular, continuous heavy rainfall <u>hit</u> pummelled-the Yangtze River basin, with rainfall anomalies locally exceeding 300_mm within 10 days (June 26-July 5; Yuan et al., 2018). Figure 2b shows that the June-July mean precipitation averaged over the Yangtze River basin ranks only second to the 1954 flood during the period 1948-2016, and is even heavier than the 1998 flood.

This Yangtze River extreme summer flood occurred in the context of the 2015/16 strong El Niño (Zhai 150 151 et al., 2016; Yuan et al., 2018). Generally, when the SST over the eastern tropical Pacific is warmer than normal in the preceding winter, the Yangtze region would experience a wetter summer, or even a 152 flood hazard. For instance, the catastrophic flooding of the Yangtze River in the summer of 1998 was 153 strongly influenced by the 1997/98 extreme El Niño (e.g., Lau and Weng, 2001). From November 2015 154 155 to January 2016, the seasonal mean SST anomaly in the NINOiño-3.4 region (NOAA's Oceanic NINONiño Index) peaked at 2.3 °C (L'Heureux et al., 2016), and returned to neutral condition until 156 May 2016. With the influence of the preceding El Niño signal, the western Pacific subtropical high 157 (WPSH) was stronger than climatology and located further west in the summer of 2016 through the 158 159 Pacific-East Asian teleconnection (e.g., Wang, 2000; Wu et al., 2003; Huang et al., 2007; Wang et al., 2014) and the Indo-western Pacific Ocean capacitor (Xie et al., 2016), so a large amount of moisture 160 was transported along its western flank, from the Indian ocean, South China Sea and Pacific oocean to 161 the middle and lower reaches of Yangtze River (Fig. 2c). As a result, there was a significantly 162 anomalous moisture band in the east-west direction characterized with the largest moisture transport 163

amount in the middle and lower reaches of Yangtze River, which was directly responsible for the 2016 summer flood (Fig. 2d).

166 3.2 Seasonal predictability of precipitation and moisture flux

Considering the association between intense moisture flux and heavy rainfall over the Yangtze River 167 basin, which is known within the canonical East Asian monsoon region (Ding and Chan, 2005), testing 168 whether atmospheric moisture flux is more predictable than precipitation at the seasonal time scale is 169 helpful for flood-control and disaster-relief. Figure 3 shows the predictions for June-July mean 170 anomalies of precipitation and corresponding moisture flux from the dynamical climate forecast model 171 172 CFSv2 for the 2016 summer flood at the first three-month leads. As compared with the observed precipitation, CFSv2 successfully captured the rainfall surplus across the middle and lower reaches of 173 the Yangtze River at 0.5-month lead (Fig. 3a), and predicted a visible moisture transport band along the 174 middle and lower reaches of the Yangtze River (Fig. 3b). The highest moisture flux anomaly occurred 175 over the southern bank of the Yangtze River, which corresponded exactly to the location of heavy 176 precipitation and flood. At 1.5-month lead, CFSv2 still performed well for the anomalous moisture flux, 177 but the predicted precipitation anomaly was much weaker than that at the 0.5-month lead (Figs. 3c-3d). 178 At the 2.5-month lead, the prediction skill of precipitation significantly weakened with almost no 179 180 anomaly (Fig. 3e), but the predicted moisture flux could reproduce the anomaly to some extent (Fig. 3f). In addition to the 2016 Yangtze flooding case, the potential predictability for June-July precipitation 181 and moisture flux at different lead times during 1982-2016 is also investigated. Figures 4a-4f depict the 182 spatial distribution of predictability for June-July mean precipitation and moisture flux at the 0.5-, 1.5-183 and 2.5-month leads respectively, where moisture flux has higher predictability than precipitation. The 184

highest AC values for moisture flux occur over the south of the Yangtze River where frequently suffers 185 from extreme summer pluvial flooding. At the 0.5-month lead, the AC values for precipitation are lower 186 than 0.3 over most areas (Fig. 4a), while they are higher than 0.3 and even close to 0.6 for moisture flux 187 predictability over the southern part of the Yangtze River basin (Fig. 4b). The AC values of 188 precipitation drops quickly with over forecast leads, and Fig. 4c shows that more than half of the 189 Yangtze region more than half of the AC values areis less than 0.2 over- the Yangtze region at the 190 when leading-1.5-month lead.; However, but the moisture flux still performs well with many AC values 191 higher than 0.3 at the 1.5-month lead, especially over and shows good predictability in the southeastern 192 193 mountain region (Fig. s. 4e-4d). The moisture flux at the 2.5-month lead has higher AC values even than precipitation at the 0.5-month lead (Fig. 4f). Meanwhile, it is evident that most areas of the 194 Yangtze River basin have significant predictability (at least at 90% confidence level) for the moisture 195 flux, but the predictability for precipitation is limited (Figs. 4a-4f). 196

197 Figure 4g indicates the corresponding spread for precipitation and moisture flux predictability throughout the middle and lower reaches of Yangtze River region (110°-123°E, 27°-34°N). The median 198 199 (mean) value for precipitation is 0.25 (0.23) at the 0.5-month lead, but reaches 0.37 (0.35) for the moisture flux. At the 2.5-month lead, the median (mean) value for moisture flux is 0.25 (0.24), which is 200 201 much higher than the value of 0.18 (0.16) for precipitation. The changes in potential predictability with 202 different forecast leads are also displayed in Figure 4h, based on both spatial and temporal samples for the Yangtze River basin. The difference between precipitation and moisture flux is statistically 203 significant (p < 0.05) with a two-tailed Student's t-test. It is evident that moisture flux has consistently 204 higher predictability than precipitation out to 8.5-month lead. Similar result is also found at the location 205

206 (30°N, 114°E) near Wuhan city (Fig. 4i), one of the big cities along the Yangtze River, which suffered
207 widespread inundation in the summer of 2016.

208 3.3 Varying predictability conditioned on different ENSO phases

As mentioned above, the Yangtze region in eastern China is one of the most strongly ENSO-affected 209 regions in the world, and the precipitation variability in this region is generally influenced by the 210 anomalous ENSO forcing (e.g., Wang, 2000; Wu et al., 2003; Ding and Chan, 2005). To explore their 211 covariability, here we performed a maximum covariance analysis (MCA, Bretherton et al., 1992) for the 212 preceding December-January-February mean SST (120°E-80°W, 10°S-60°N) and June-July mean 213 214 precipitation (100°E-150°E, 10°N-55°N) fields from 1948 to 2016. It is found that the second mode (MCA2) explains 23% of the variance, and its corresponding SST anomaly pattern is very similar to the 215 traditional ENSO-like pattern with a warm anomaly over the equatorial eastern Pacific and a horse-216 shoes pattern with cold anomalies over the western tropical and central nNorthern Pacific (Fig. 5a). 217 Meanwhile, its temporal evolution is strongly correlated with the NINO3.4 SST anomaly (r = 0.92, 218 black line in Fig. 5c). Correspondingly, the summer precipitation in the Yangtze region is above normal 219 significantly (Fig. 5b). Therefore, the Yangtze region is prone to experience a rainy or flooding summer 220 if the SST over the eastern tropical Pacific is warmer than normal in the preceding winter based on the 221 222 covariance analysis during the period 1948-2016, whether the predictability varies during different 223 ENSO phases should be investigated.

To explore the impacts of preceding ENSO El Niño signals on Yangtze precipitation and moisture <u>flux</u>
predictability, correlations and hit rates conditional on different ENSO phases (i.e., El Niño and La Niña)
at different leads are shown in Figure 6. It is found that the seasonal predictability of Yangtze summer

227	rainfall and moisture flux is much higher following El Niño years than La Niñas (Fig. 6a). The contrast
228	during different ENSO phases is more obvious for extreme events, and the potential detectability of
229	extreme moisture flux increases by 20% in post-El Niño summers as compared with the potential
230	detectability of extreme precipitation (Fig. 6b). This asymmetric performance during El Niño and La
231	Niña has drawn many attentions. One of the reasons is that the atmospheric response to tropical Pacific
232	SST anomaly is inherently nonlinear (Hoerling et al., 1997), where both the amplitude of SST anomaly
233	in the equatorial eastern Pacific and the associated atmospheric response are significantly larger during
234	El Niño than during La Niña episodes (Burgers and Stephenson 1999). Figure 6 also shows that the
235	predictability is high conditional on El Niños even out to 6.5-month lead, which is consistent with
236	previous studies. For instance, Sooraj et al. (2012) have mentioned that forecasting seasonal rainfall
237	anomalies over central tropical Pacific islands from El Niño winter into the following spring/summer is
238	skillful by using CFS, and Ma et al. (2015) have demonstrated high predictability for seasonal drought
239	over ENSO-affected regimes in southern China. The exception for 3.5-month lead forecast (started in
240	March) where the predictability conditioned on La Niña is slightly higher than El Niño (Fig. 6a) is
241	perhaps related to the 'spring predictability barrier', but such chaos disappears for extreme events (Fig.
242	6b),
243	Furthermore, CFSv2 predictions of atmospheric circulations associated with 500 hPa geopotential
244	height and 850 hPa wind and moisture flux are also investigated during different ENSO phases. As
245	shown in Figure 6c, there is an anomalously high pressure center over the subtropical western

247 impliesying that the WPSH is enhanced in post El Niño summers. Such circulation pattern would brings

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subtropical Pacific, which is a recurrent pattern in post-El Niño summers (Xie et al., 2016) and

the larger amounts of atmospheric moisture than normal from the southern oceans to the Yangtze River 248 basin, which corresponds well with extreme hydrologic events. The mechanism for this lag-impact of El 249 Niño on East Asia summer climate is the Indo-western Pacific ocean capacitor (IPOC), where the 250 coupled wind-evaporation-SST feedback over the Northwest Pacific in spring persists to trigger East 251 Asia-Pacific/Pacific-Japan (EAP/PJ) pattern that arises from the interaction of the anomalous anti-252 cyclone and North Indian Ocean warming in post-El Niño summers (Xie et al., 2016). On the contrary, 253 preceding La Niña winters are favorable to a low pressure anomaly in next summer, accompanied with 254 an abnormal cyclonic circulation, and thereby preventing the moisture from moving northwards to the 255 256 Yangtze region (Fig. 6d). It implies that the precipitation deficits or droughts are more likely to occur in this region in post-La Niña summers. The contrast is even obvious even for forecasts forat 6.5-month 257 lead forecasts (Figs. 6e-6f).- The differences in predicted circulation and associated moisture transport 258 largely result in Such predicted circulation discrepancy in different initial ocean conditions largely 259 260 determines higher predictability for extreme hydrologic events over the middle and lower reaches of the Yangtze River basin in post-El Niño summers (Hu et al., 2014). 261

262 4. Summary and Discussion

Previous studies have revealed that moisture flux has higher predictability than precipitation in weather forecasts over the northwestern Europe and the western U.S., which <u>areis</u> affected by westerlies and narrow bands of enhanced moisture transport known as atmospheric rivers (Lavers et al., 2014, 2016b). However, whether the atmospheric moisture <u>flux</u> is more predictable at seasonal time scales during a summer monsoon region is still unclear. Based on seasonal ensemble predictions from NCEP's operational CFSv2 model during 1982-2016, our results show that moisture flux has higher seasonal

predictability than precipitation over China's Yangtze River basin in summer. In addition, we also 269 investigated potential predictability of precipitation and moisture flux on weekly averaged fields in 270 June-July at subseasonal time scale. Results are similar to seasonal time scale, where the moisture flux 271 has a higher predictability than precipitation at different lead times (Fig. 7). Moreover, the potential 272 predictability may change under different climatic conditions. The seasonal predictability is much 273 higher when initialized in warm ENSO conditions not only for precipitation but also for moisture flux. 274 More importantly, the moisture flux shows higher detectability (hit rate) than precipitation for extreme 275 pluvial flooding events following El Niño winters. The results suggest that it may be possible to extend 276 277 the predictability of Yangtze River summer floods and to provide more reliable early warning by using atmospheric moisture flux predictions. However, to which degree that moisture flux is connected with 278 precipitation and floods might be model dependent. It is necessary to explore their connections in a 279 multi-model framework (e.g., NMME; Kirtman et al., 2014; Shukla et al., 2016). 280

This study extends previous findings on the predictability of precipitation and moisture flux at synoptic 281 scales (Lavers et al., 2014) to seasonal time scales, and from atmospheric river-affected regions to the 282 East Asian summer monsoon region. Given that the transport of atmospheric moisture from oceanic 283 source regions is important for extreme rainfall in monsoon regions (Gimeno et al., 2012), moisture flux 284 285 might also be useful for long-range forecasting over other areas affected by the monsoon and low-level jets. In fact, extreme precipitation and floods are found to be associated with large-scale moisture 286 transport over the North American monsoon (Schmitz and Mullen, 1996) and the South American 287 monsoon (Carvalho et al., 20110) regions. Extreme precipitation and floods usually occur accompanied 288 with intensive atmospheric moisture transport, especially over a large area such as the middle and lower 289

290 reaches of the Yangtze River. Given higher predictability of atmospheric moisture flux, it can be used 291 as a precursor for flooding forecasting, either directly linking moisture flux to streamflow prediction 292 through statistical techniques (e.g., conditional distribution or Bayesian methods), or adding moisture 293 flux information into precipitation prediction, and consequently improving floods prediction. Moreover, 294 it is suggested that assimilating moisture flux observations into numerical climate forecast models 295 would benefit the prediction of hydrological extremes.

The higher moisture <u>flux</u> predictability largely arises from more predictable large-scale circulation (Li et al., 2016), which strongly determines the <u>atmospheric</u> moisture transport. Although precipitation variability is affected by both large-scale moisture transport and localized process and features, such as condensation nuclei in the atmosphere and lifting movement, it is expected that moisture transport could still be used as a crucial source of predictability for flooding over monsoonal regimes, especially at long leads where meso-scale convection is still unpredictable at seasonal time scales.

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Acknowledgement. This work was supported by National Natural Science Foundation of China (91547103, 41605055), and the National Key R&D Program of China (2016YFA0600403). The authors thank Dr. Arun Kumar for helpful discussions. The authors acknowledge NCEP/EMC and IRI (http://iridl.ldeo.columbia.edu/SOURCES/.NOAA/.NCEP/.EMC/.CFSv2/) for making the CFSv2 hindcast and real-time forecast information available. We are also grateful for the constructive comments from three anonymous reviewers to improve the quality of this paper.

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Figure 1. An example of the potential predictability calculation, where the ensemble member 1 is the truth and the mean of the members 2-24 is the prediction. This is for 116°E and 28°N near to Wuhan city at (a-b) the 0.5-month lead and (c-d) the 1.5-month lead.



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423 Figure 2. The 2016 extreme summer flood. (a) Mean precipitation anomaly (shading, mm/day) during the June-July of 2016. (b) Time series of the June-July mean precipitation anomaly averaged over the 424 middle and lower reaches of Yangtze River basin (110-123°E, 27-34°N) in (a). (c) Anomaly of 500 hPa 425 geopotential height (shading, gpm) superimposed by absolute integrated horizontal moisture transport 426 between 1000 to 300 hPa layers 850 hPa vapor transports (vectors, kg•m⁻¹s⁻¹g/em•hPa•s). The thick 427 contour lines are 5880 gpm, implying the location of the West Pacific Subtropical High, where the 428 black denotes the June-July 2016 and the cyan is the climatology during 1982-2010. (d) Anomaly of 429 integrated horizontal moisture transport amount between 1000 to 300 hPa layers (shading, $\underline{k}\underline{K}$ g•m⁻¹s⁻¹). 430 431

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439 Figure 4. (a-f) Potential predictability (AC value, see Method) for June-July mean precipitation and atmospheric moisture flux at different lead times during 1982-2016 over the middle and lower reaches 440 of Yangtze River for the 0.5-, 1.5- and 2.5-month leads; the stippling indicates a 95% confidence level 441 according to a two-tailed Student's t-test. (g) Median, lower and upper quartiles, 1.5 times the 442 interquartile ranges for AC values for precipitation (black) and moisture (red) throughout the study 443 region (110-123°E, 27-34°N); outliers are displayed with + signs. (h-i) Potential predictability 444 throughout study region and Wuhan city (pink pentagram in (a)) at different lead times; the error bars 445 are standard deviations according to 24 members. 446



447

Figure 5. (a-b) Spatial and (c) temporal patterns of the second modes based on the maximum covariance analysis (MCA) for SST in preceding winter (December-January-February) and precipitation field in summer (June-July) for 1948-2016. Here the second MCA mode explains 23 % of the variance, as indicated in the square fraction of covariance (SFC).







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Figure 6. Potential predictability at different lead times in terms of (a) anomaly correlation (AC) for precipitation and moisture, and (b) hit rate (HR) for flood events (>90th percentiles) across the Yangtze River region conditioned on ENSO phases. (c-d) Composites of predicted anomalies of 500 hPa geopotential height (contour, gpm) superimposed by 850 hPa wind (vectors, m/s) and moisture flux (shading, g/cm•hPa•s) at the 0.5-month lead during different ENSO phases. (e-f) The same as (c-d), but for 6.5-month lead time.

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474		June 17, 2018	
475	Dr. Shraddhanand Shukla		
476	Editor		
477	Hydrology and Earth System Sciences		
478	RE: manuscript #hess-2018-112		
479			
480	Dear Dr. Shukla,		
481			
482	Thank you for your kind decision letter on our man	nuscript entitled "Extending seasonal predictability of	
483	Yangtze River summer floods" (hess-2018-112). We have carefully considered the reviewer's		
484	comments and incorporated them into the revised manuscript to the extent possible. We hope that you		
485	find the revised manuscript and the response to the reviews acceptable to HESS.		
486	The detailed responses to the comments are attached.		
487			
488	We appreciate the effort you spent to process the manuscript and look forward to hearing from you soon		
489			
490	Sincerely yours,		

Laga

- 492 Xing Yuan
- 493

494 Responses to the comments from Reviewer #1

We are very grateful to the Reviewer for the positive and careful review. The thoughtful comments have helped improve the manuscript. The reviewer's comments are italicized and our responses immediately follow.

498

This article by Wang and Yuan investigates seasonal predictability of water vapor flux and precipitation in the CFSv2 forecasts to see if water vapor flux could provide higher predictability of floods. I enjoyed reading the study and believe that this is an interesting study and one that is relevant to forecast users. I currently have two major concerns which I lay out below, and some further more minor comments. I hope the authors will find them of use. Major comments:

504 **Response:** We would like to thank the reviewer for the positive comments. Please see our responses 505 below.

506

507 I. While the language is generally very good, there are grammatical issues, so the paper would benefit 508 from being read by a native English speaker. Also, I think the authors need to be careful with their 509 terminology. For example, on many occasions they say "moisture" when in fact they mean "moisture 510 flux" or "water vapor flux". This needs to be corrected so that readers are not confused by the terms 511 (moisture could be viewed as total column water vapor).

512 **Response:** Thanks for the comments. We have corrected inappropriate terminology and proofread the 513 manuscript carefully. And the "moisture" has been replaced with "moisture flux" throughout the 514 manuscript.

515

2. The authors only analyse monthly fields and I believe that this is not sufficient. This is useful, but I think some testing should be done on at least weekly or 2-weekly averaged fields. The reason is that it will be easier (and more beneficial) to see when the predictability drops off in the model at a finer time resolution. For example, Luo and Wood (2006) have some good plots that may help you to consider other time averages.

521 **Response:** Thanks for the comments. According to the reviewer's suggestion, we have incorporated the 522 analysis for the potential predictability of precipitation and moisture flux for weekly averaged fields in 523 the revised manuscript.

524

525 We have used the S2S daily data, and have clarified them in the Section "2.2 CFSv2 seasonal hindcast 526 and real-time forecast data" as follows:

⁵²⁷ "In order to investigate the predictability at finer temporal resolution (e.g., weekly mean fields), the ⁵²⁸ CFSv2 daily reforecasts were obtained from the Subseasonal to Seasonal (S2S) prediction project for

the period of 1999-2010, with forecast lead times up to 45-days (Vitart et al. 2017). As for the June 1-7 weekly mean fields, the reforecasts started from May 18 were used as the first ensemble member, the

531 reforecasts started from May 19 were used as the second, and so on. This resulted in 14 ensemble

532 members, with forecast lead times from 1-day to 14-days. The above process was repeated for other

533 weekly averaged fields during June and July. This is called as the first group of ensemble subseasonal

forecasts, with lead times of 1-14days. The second group of ensemble reforecasts started from 17 May, May ..., and 30 May were formed similarly, with lead times of 2-15days, and so on." (L97-105 in the tracked version of the revised manuscript)

537

The AC values for both weekly mean precipitation and moisture flux during June-July at different lead times were obtained and shown in Figure R1 (Figure 7 in the revised manuscript). Results show that the moisture flux has a higher predictability than precipitation for weekly averaged fields at different lead times, which is consistent with the results on seasonal averaged fields. We have added the related discussion in Section 4 as follows:

⁵⁴³ "In addition, we also investigated potential predictability of precipitation and moisture flux on weekly ⁵⁴⁴ averaged fields in June-July at subseasonal scale. Results are similar to seasonal time scale, where the

545 moisture flux has a higher predictability than precipitation at different lead times (Fig. 7)." (L269-272)

546

547 Figure R1. (a-f) Potential predictability (AC value) for weekly mean precipitation and atmospheric 548 moisture flux at different lead times during June-July of 1999-2010 over the middle and lower reaches 549 of Yangtze River for the 1-14, 5-18 and 8-21 days leads; the stippling indicates a 95% confidence level

Lead time

- 550 according to a two-tailed Student's t-test. (g) Potential predictability throughout study region at
- 551 different lead times. Here, the daily CFSv2 reforecast were obtained from the S2S prediction project for
- 552 the period of 1999-2010.
- 553
- 554 Other comments:
- 555 Line 19: "atmospheric moisture" see major comment 1.
- Response: Thanks for the suggestion. We have corrected them throughout the manuscript. (L16, 18,19,
 L48, L84, etc.)
- 558

559 *Line 33: "ability to be predicted" is a strange phrase. Perhaps consider re-phrasing.*

- 560 Response: We have changed it to "the ability of the model to predict itself". (L33-34)
- 561

562 *Line 40: Precipitation is connected to mesoscale (or more local scale) circulation and orography.*

- 563 **Response:** Thanks for the comments. We have revised it as follows:
- ⁵⁶⁴ "The atmospheric moisture flux is supposed to be better predicted by large-scale climate models than ⁵⁶⁵ precipitation that is not only connected to mesoscale (or more local scale) circulation but also ⁵⁶⁶ influenced by the vertical convection and the localized orography". (L39-42)
- 567
- 568 Line 61: I think Lavers et al (2016a) investigated prediction skill, not predictability.
- 569 **Response:** Revised as suggested. (L63-64)
- 570
- 571 Line 75: What pressure levels were used? Please add some details here.
- **Response:** We have specified as "Monthly mean atmospheric fields including geopotential height, uwind, v-wind, and specific humidity at 300, 400, 500, 700, 850, 925 and 1000 hPa were derived from the ERA-Interim reanalysis". (L77-79)
- 575
- Lines 103-112: Why does the AC go from 0.33 to 0.44 (compare Figure 1b and 1d)? You would expect
 the predictability to drop with lead time.
- 578 **Response:** In general, the predictability drops over lead times, but not necessarily for any cases.
- 579 We plotted the results for all 24 ensemble members in Figure R2, and found that the AC for 0.5-month
- lead is not necessarily higher than 1.5-month lead. However, the average results for the 24 AC (Fig. R2c)
- 581 shows that AC decreases over leads on average.
- 582

Figure 2. The 2016 extreme summer flood. (a) Mean precipitation anomaly (shading, mm/day) during 618 the June-July of 2016. (b) Time series of the June-July mean precipitation anomaly averaged over the 619 middle and lower reaches of Yangtze River basin (110-123°E, 27-34°N) in (a). (c) Anomaly of 500 hPa 620 geopotential height (shading, gpm) superimposed by absolute integrated horizontal moisture transport 621 between 1000 to 300 hPa layers (vectors, kg•m⁻¹s⁻¹). The thick contour lines are 5880 gpm, implying the 622 location of the West Pacific Subtropical High, where the black denotes the June-July 2016 and the cyan 623 is the climatology during 1982-2010. (d) Anomaly of integrated horizontal moisture transport amount 624 (shading, $kg \cdot m^{-1}s^{-1}$). 625

617

Figure 3: What initialisation times are used in this figure (e.g. 1st May 2016)? Please consider adding
 to the caption.

Response: We have revised as "where the 0.5-month lead forecasts were initialized from mid-May to early June in 2016, 1.5-month lead forecasts were initialized from mid-Apr to early May in 2016, and so on." (L435-436)

632

Figure 6: Panels c-d. Perhaps a few extra contours should be added to more clearly show the 500 hPa geopotential height?

635 **Response:** Revised as suggested. (L452-458)

Figure 6. Potential predictability at different lead times in terms of (a) anomaly correlation (AC) for precipitation and moisture, and (b) hit rate (HR) for flood events (>90th percentiles) across the Yangtze River region conditioned on ENSO phases. (c-d) Composites of predicted anomalies of 500 hPa geopotential height (contour, gpm) superimposed by 850 hPa wind (vectors, m/s) and moisture flux (shading, g/cm•hPa•s) at the 0.5-month lead during different ENSO phases. (e-f) The same as (c-d), but for 6.5-month lead time.

644 Responses to the comments from Reviewer #2

645 We are very grateful to the Reviewer for the positive and careful review. The thoughtful comments have 646 helped improve the manuscript. The reviewer's comments are italicized and our responses immediately 647 follow.

648

649 This study found moisture flux has higher predictability than precipitation in summer in Yangtze River

650 basin, China. The predictability of precipitation and moisture are higher in post-El Niño summers than

651 those in post-La Niñas. The results extend the predictability of Yangtze River summer floods and to

652 provide more reliable early warning by using atmospheric moisture flux predictions. The research is

653 very interesting and significative. However, there are a few issues that the authors need to address

654 before the manuscript can be accepted. I recommend most of the issues I raise below just need

655 *clarification or justification*.

Response: We would like to thank the reviewer for the positive comments. Please see our responsesbelow.

658

659 We predict the precipitation in order to predict the flood. How to predict the flood using the moisture?

660 The authors maybe add some discussion.

661 **Response:** Thanks for the comments. We have added the discussion as follows:

"Extreme precipitation and floods usually occur accompanied with intensive atmospheric moisture 662 transport, especially over a large area such as the middle and lower reaches of the Yangtze River. Given 663 higher predictability of atmospheric moisture flux, it can be used as a precursor for flooding forecasting, 664 either directly linking moisture flux to streamflow prediction through statistical techniques (e.g., 665 conditional distribution or Bayesian methods), or adding moisture flux information into precipitation 666 prediction, and consequently improving floods prediction. Moreover, it is suggested that assimilating 667 moisture flux observations into numerical climate forecast models would benefit the prediction of 668 hydrological extremes." (288-295 in the tracked version of the revised manuscript) 669

670

671 *Line 133, 300m→ 300mm.*

672 Line 378, $Kg \cdot m \cdot ls \cdot l \rightarrow m \cdot ls^{-l}$

673 **Response:** Thanks for the comments. We have corrected them as suggested. (L146, L426)

Responses to the comments from Reviewer #3 675

We are very grateful to the Reviewer for the positive and careful review. The thoughtful comments have 676 helped improve the manuscript. The reviewer's comments are italicized and our responses immediately 677 follow. 678

679

The article entitled, "Extending seasonal predictability of Yangtze River summer floods" by Wang and 680

Yuan explores the seasonal predictability of both moisture flux and precipitation in the CFSv2 forecast 681

system. The study aims to determine whether moisture flux forecasts can be used to better predict for 682

summer flood prediction (compared to precipitation). I found the study interesting and potentially 683

684 useful to decision-makers and end-users in the region. However, I have several major concerns that I

hope the authors will address, as well as a number of minor comments. 685

Response: We would like to thank the reviewer for the positive comments. Please see our responses 686 below. 687

688

Major comments: 689

1. While much of the study is well written, there are numerous places in the text where there are 690

grammatical issues. These range from simple subject-verb agreement (as in the first sentence, "was" 691

should be replaced with "were"), to passages where the language is misleading and it is not clear what 692

the authors mean to say. The paper (and its corresponding conclusions) would benefit greatly from a 693

thorough proofread by a colleague who can help address and correct the language issues. 694

Response: Thanks for the comments. We have improved the clarification and carefully proofread the 695 manuscript, including the first sentence. 696

697

2. A major conclusion of the study is that the moisture flux can be better predicted than precipitation in 698

summers directly following ENSO events, and particularly El Niño. However, there is very limited 699

discussion of how and why El Niño impacts this area and therefore lends itself as a potential predictor 700

701 of moisture flux and hence, flooding in the region. Without providing some further discussion to the

paragraph that begins on line 220 that speaks directly to how ENSO is understood to impact the area 702

and how the plots shown in Figure 6 are consistent with this, I find that the major conclusions are not 703

fully supported by the study at present. For example, are the moisture flux vectors shown in Figure 6 704

related to the anomalous high, and is that known to be forced by El Niño? Some more explanation and 705 discussion is needed.

706

Response: Thanks for the comments. We have clarified as follows: 707

Section 3.3: "As mentioned above, the Yangtze region in eastern China is one of the most strongly 708 ENSO-affected regions in the world, and the precipitation variability in this region is generally 709

influenced by the anomalous ENSO forcing (e.g., Wang, 2000; Wu et al., 2003; Ding and Chan, 710

2005).....It is found that the second mode (MCA2) explains 23% of the variance, and its corresponding 711 712 SST anomaly pattern is very similar to the traditional ENSO-like pattern with a warm anomaly over the equatorial eastern Pacific and a horse-shoes cold anomalies over the western tropical and central 713 714 Northern Pacific (Fig. 5a). Meanwhile, its temporal evolution is strongly correlated with the NINO3.4 SST anomaly (r = 0.92, black line in Fig. 5c). Correspondingly, the summer precipitation in the Yangtze 715 region is above normal significantly (Fig. 5b)." Above all, there is no doubt that the El Niño signals 716 have an crucial role on the climate variability over the Yangtze region, especially on the precipitation 717 anomalies by impacting the large-scale circulation variation over the Northwestern Pacific Ocean and 718 the associated water vapor transport to the Yangtze region. When El Niño occurs in preceding winter, 719 there is always an enhanced western Pacific subtropical high (WPSH) accompanied with a weakened 720 East Asia summer monsoon (EASM) in the following summer, thereby resulting in an anomalously 721 722 anticyclonic circulation pattern over the northwestern Pacific that brings large amounts of atmospheric moisture from the oceans to the Yangtze River (Yuan et al. 2017). 723

724

725 In the revise version, we add some detailed discussion about the mechanism for the lag-impact of El 726 Niño on East Asia summer climate including how the El Niño forcings impact the atmospheric moisture 727 transport to the Yangtze region as follows:

"As shown in Figure 6c, there is an anomalously high pressure center over the subtropical western 728 Pacific, which is a recurrent pattern in post-El Niño summers (Xie et al., 2016) and implies that the 729 730 WPSH is enhanced. Such circulation pattern would bring larger amounts of atmospheric moisture than 731 normal from the southern oceans to the Yangtze River basin, which corresponds well with extreme hydrologic events. The mechanism for this lag-impact of El Niño on East Asia summer climate is the 732 733 Indo-western Pacific ocean capacitor (IPOC), where the coupled wind-evaporation-SST feedback over the Northwest Pacific in spring persists to trigger East Asia-Pacific/Pacific-Japan (EAP/PJ) pattern that 734 735 arises from the interaction of the anomalous anti-cyclone and North Indian Ocean warming in post-El Niño summers (Xie et al., 2016)." (L244-253 in the tracked version of the revised manuscript) 736

737

738 Minor comments:

739 1. Line 39-40, the sentence that mentions model precipitation being influenced by "meso-scale

- 740 convections" is unclear. Here, are the authors referring to mesoscale(local) circulation patterns that
- 741 impact precipitation? Also, it might be worth noting that convection schemes themselves (used to
- 742 parameterize finer scale processes) would also impact forecasted precipitation.
- 743 **Response:** Thanks for the comments. We have revised the manuscript as follows:
- "The atmospheric moisture flux is supposed to be better predicted by large-scale climate models than precipitation that is not only connected to mesoscale (or more local scale) circulation but also influenced by the vertical convection and localized orography (Lavers et al., 2014, 2016b)." (L39-42) 747

748 2. Line 75: The pressure levels of the variables studied should be identified.

Response: Thanks for the comments. We have specified as "Monthly mean atmospheric fields including geopotential height, u-wind, v-wind, and specific humidity at 300, 400, 500, 700, 850, 925 and 1000 hPa were derived from the ERA-Interim reanalysis". (L77-79)

752

3. In Figure 1, is there a reason why the AC is higher for the moisture flux at 1.5months lead-time
compared to 0.5 months? It would be good if the authors could provide some understanding of why this
is the case or if they believe it to be spurious because it is surprising.

- **Response:** Thanks for the comments. In general, the predictability drops over lead times, but not necessarily for any cases.
- 758 We plotted the results for all 24 ensemble members in Figure R1, and found that the AC for 0.5-month
- 759 lead is not necessarily higher than 1.5-month lead. However, the average results for the 24 AC (Fig. R1c)
- 760 shows that AC decreases over leads on average.

761

Figure R1. Potential predictability (AC value) when different ensemble member was taken as the truth
and the mean of the members was the prediction at Wuhan city for the (a) 0.5-and (b) 1.5-month leads.
(c) the final estimate of the potential predictability in Wuhan city.

- 765
- 766 *4. Line 124: There is no "b" in the equation on line 123.*
- 767 **Response:** We have removed it. (L137)
- 768

5. Lines 132-134: This sentence is awkward, particularly the use of the word "pummeled," please rewrite.

Response: Thanks for the comments. We have changed it as "In particular, continuous heavy rainfall hit
the Yangtze River basin, with rainfall anomalies locally exceeding 300 mm within 10 days (June 26July 5; Yuan et al., 2018)". (L145-147)

- 6. The sentence on Lines 174-177 is also awkward and does not clearly explain the results from Figure 775 4.
- 776

Response: Thanks for the comments. We have revised as "The AC values for precipitation drop quickly 777 778 with forecast leads, and Fig. 4c shows that more than half of the AC values are less than 0.2 over the Yangtze region at the 1.5-month lead. However, the moisture flux still performs well with many AC 779 values higher than 0.3 at the 1.5-month lead, especially over the southeastern mountain region (Fig. 4d)." 780 (L188-193) 781

782

7. Line 206: This sentence is a bit contradictory as it says "To explore the impacts of preceding El Nino 783

signals..." and then tells us that "hit rates conditional on different ENSO phases..." are shown in 784

Figure 6. Figure 6 shows both El Niño and La Niña hit rates, so really the authors are showing the 785

impacts of preceding ENSO events (not just El Niño as is written). Please switch "El Niño" in the 786

beginning of the sentence with "ENSO" and in the second mention of "ENSO" 787

phases, could add "(i.e. El Niño and La Niña)". 788

Response: Thanks for the comments. We have revised as suggested. 789

"To explore the impacts of preceding ENSO signals on Yangtze precipitation and moisture flux 790 791 predictability, correlations and hit rates conditional on different ENSO phases (i.e., El Niño and La Niña)

- at different leads are shown in Figure 6." (L224-226) 792
- 793

794 8. Lines 228-230 conclude that the different circulation patterns predicted for the two ENSO phases

795 determine a higher predictability for extreme hydrologic events in post-El Niño summers. However, why

is it necessarily higher predictability and not just a different signal that is predicted because of the 796

different ENSO events? This conclusion seems like a bit of a stretch to me without understanding of why 797

the El Niño signal would translate to higher predictability than La Niña based solely on the evidence 798

- 799 presented in the manuscript.
- **Response:** Thanks for the comments. We have added more explanations in the revised manuscript as 800 follows: 801
- "This asymmetric performance during El Niño and La Niña has drawn many attentions. One of the 802 reasons is that the atmospheric response to tropical Pacific SST anomaly is inherently nonlinear 803 (Hoerling et al., 1997), where both the amplitude of SST anomaly in the eastern equatorial Pacific and 804 the associated atmospheric response are significantly larger during El Niño than during La Niña 805 episodes (Burgers and Stephenson 1999)." (L230-234) 806
- "It implies that the precipitation deficits or droughts are more likely to occur in this region in post-807 LaNiña summers. The contrast is obvious even for forecasts at 6.5-month lead (Figs. 6e-6f). The 808 differences in predicted circulation and associated moisture transport largely result in higher 809 predictability for extreme hydrologic events over the middle and lower reaches of the Yangtze River 810
- basin in post-El Niño summers (Hu et al., 2014)." (256-261) 811

- 9. Line 373 references the "middle and lower reaches of Yangtze River basin." However, these areas
- are not previously defined in the text. I assume they may be the boxes outlined in Figure 2a, but this
- 815 *needs to be clarified.*
- 816 **Response:** We have now defined it in the Introduction section as follows:
- 817 "In present study, we aim to address the above questions by evaluating the seasonal predictability of
- 818 precipitation and moisture flux for the middle and lower reaches of Yangtze River (110-123°E, 27-34°N)
- 819 based on multisource observational data, and ensemble hindcasts and real-time forecasts from a
- dynamical seasonal forecast model Climate Forecast System version 2 (CFSv2; Saha et al., 2014) for the period of 1982-2016." (L68-72)
- 822

823 10. The legend for Figure 2c defines the 850 hPa moisture flux vectors in g/cm*hPa*s. I have never
824 seen this unit used before for moisture flux and would recommend it be converted to m/s kg*kg.

Response: Thanks for the comments. According to the suggestion from reviewer#1, we have used the total column-integrated moisture flux instead of that at the 850 hPa level in revised manuscript. The corresponding unit has also been converted to kg•m⁻¹s⁻¹. (L421-429)

Figure 2. The 2016 extreme summer flood. (a) Mean precipitation anomaly (shading, mm/day) during the June-July of 2016. (b) Time series of the June-July mean precipitation anomaly averaged over the

middle and lower reaches of Yangtze River basin (110-123°E, 27-34°N) in (a). (c) Anomaly of 500 hPa geopotential height (shading, gpm) superimposed by absolute integrated horizontal moisture transport between 1000 to 300 hPa layers(vectors, kg•m⁻¹s⁻¹). The thick contour lines are 5880 gpm, implying the location of the West Pacific Subtropical High, where the black denotes the June-July 2016 and the cyan is the climatology during 1982-2010. (d) Anomaly of integrated horizontal moisture transport amount (shading, kg•m⁻¹s⁻¹).

837

838 11. Figure 3: the different columns are plotted with a different longitudinal domain. It would be helpful

839 in comparing the precipitation to the moisture flux if all panels were plotted using the same longitude

840 *bounds*.

841 **Response:** Revised as suggested.

Figure 3. Spatial distributions of CFSv2 predicted anomalies of precipitation (shading, mm/day) and atmospheric moisture flux (shading, Kg•m⁻¹s⁻¹) in the June-July of 2016 at the 0.5-, 1.5- and 2.5-month leads, where the 0.5-month lead was initialized from mid-May to early June, 1.5-month lead was initialized from mid-Apr to early May, and so on.

848 12. Figure 4 seems to contradict what is shown in Figure 1 (see Minor Comment #3). The correlation
849 maps shown in Figure 4 indicate that Wuhan City has a lower AC value for lead-time 1.5 than lead time

850 0.5, but Figure 1d indicates that the AC is 0.44 for 1.5 month lead but only 0.33 for 0.5 month lead.

851 Why is there a discrepancy?

- 852 **Response:** Please see the response to minor Comment #3 above.
- 853

854 13. While the methods employed are interesting and the figures generally informative, I would 855 encourage some reorganization of Figures 2-6. Figures 2-3examine the anomalous 2016 event that the

text implies is related to the El Niño that occurs that year so when it is followed up by Figure 4 which

857 shows the potential predictability based on all years (1982-2016), it is a bit misleading. I would

- 858 recommend putting Figure 4 directly after Figure 1 and then continuing on
- 859 to the Figures detailing the 2015-2016 event.

Response: Our motivation of this study started from the prediction of the pluvial flood event over the 860 861 Yangtze region in the summer of 2016, as mentioned in the first paragraph of the Introduction section. Therefore, we first showed the observation and prediction for the 2016 summer in Figs. 2-3, and found 862 better prediction of moisture flux than the precipitation. Then, we analyzed the potential predictability 863 based on all hindcast and real-time forecasts during1982-2016, and found that moisture flux has a 864 higher predictability than precipitation. Finally, we explored the varying predictability conditioned on 865 866 different ENSO phases based on all observations, hindcast and real-time forecasts. We believe the logic is straightforward, so we would like to keep the original organization. 867

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