



1 **ENSO-triggered floods in South America:**
2 **correlation between maximum monthly discharges during strong events**

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7

8 **Abstract**

9 ENSO-triggered floods altered completely the annual discharge of many watersheds of South America. Anomalous
10 years as 1941, 1982-83, 1997-98 and 2015-16 signified enormous fluvial discharges draining towards the Pacific
11 Ocean, but also to the Atlantic. These floods affected large cities built on medium-latitudinal Andes (Lima, Quito,
12 Salta), but also those located at floodplains, as Porto Alegre, Blumenau, Curitiba, Asunción, Santa Fe and Buenos
13 Aires. Maximum discharge months are particular and easily distinguished along time series from watersheds located
14 at the South American Arid Diagonal. At watersheds conditioned by precipitations delivered from the Atlantic or
15 Pacific anti-cyclonic centers, the ENSO-triggered floods are more difficult to discern. The floods of 1941 affected
16 70,000 inhabitants in Porto Alegre. In 1983, Blumenau city was flooded during several days; and the Paraná River
17 multiplied 15 times the width of its middle floodplain. That year, the Colorado River in Northern Patagonia
18 connected for the last time to the Desagüadero – Chadileuvú - Curacó system and its delta received saline water for
19 the last time. During strong ENSO years the water balances of certain piedmont lakes of Southern Patagonia are
20 modified as the increases in snow accumulations cause high water levels, with a lag of 13 months. The correlation
21 between the maximum monthly discharges of 1982-83 and 1997-98 at different regions and watersheds indicates
22 they can be forecasted for future floods triggered by same phenomena. South American rivers can be classified
23 therefore into ENSO-affected and ENSO-dominated for those within the Arid Diagonal that are exclusively subject
24 to high discharges during those years.

25 Keywords: floods, El Niño-Southern Oscillation, South America, maximum monthly discharges

26



1. Introduction

El Niño events were known before the Spanish colonized the Peru region because of their consequences on the anchovy fishery. They were also known by seasonal heavy rainfalls and rapid floods in tropical South America. However, Jules Verne exaggerated these flash floods as occurring on the Pampas plains in his book *Les enfants du Capitaine Grant* (1868, reproduced in 1962 in the Disney's movie *In search of the Castaways*). These floods do not occur as rapid; they are the response of several weeks or months with rains over the average.

Their origin is well known: immense volumes of water transported across the Pacific Ocean during certain anomalous years, the so-called “*El Niño*” or ENSO years (Vargas et al. 2000; Andreoli and Kayano 2005). Although this interannual anomalous years are known by their climatic and oceanographic consequences, their hydrological responses in South American rivers have not been carefully reported. One to the main reason is the lack of information about rains records (Sun et al., 2015) and also of long and continuous hydrological records (Ward et al. 2016).

Rapid floods at the Andes watersheds occur during strong ENSO years, impacting in Peru, Ecuador and Northern Chile. They are significantly recorded when they affect arid watersheds comprised within the South American Arid Diagonal. However, these interannual floods also affect extended Atlantic watersheds of rivers as the Paraguay, Bermejo, Pilcomayo and Salado. This manuscript reported the available records of these floods -in their monthly periodicity-, compiled at national agencies of different countries (Ecuador, Peru, Chile, Paraguay, Brazil, Argentina and Uruguay). Environmental and social impacts of these floods in South America were reported considering specially that the floods triggered by El Niño-La Niña are significantly longer (Ward et al. 2016).

2. Climate

Central South America has a subtropical to temperate climate. Humidity is provided from the east by trade winds from the anti-cyclonic center of the South Atlantic. Further south, humidity is also provided by westerly winds from the South-Pacific anti-cyclonic center. Between both humid areas, the Arid South American Diagonal (ASAD) extends from N to S, connecting the Atacama and Patagonian deserts (Fig. 1). Climate was considered as the main variable governing the suspended sediment yields from catchments basins located to the E of the Andes between Ecuador and Bolivia, either in its variability or indirectly conditioning the vegetation cover (Pepin et al. 2013). Along the coast of Chile, rains increases from north to south (Valdés-Pineda et al. 2014; Araya Ojeda and Isla 2016). On the other hand, along the Eastern Patagonia coast, rains increase northwards (Coronato et al. 2008).





Fig. 1. A. Anti-cyclonic centers ejecting winds from the east and west. B. South America is characterized by the ASAD connecting Atacama and Patagonia deserts. C. Major rivers of South America.

South America has significant temporal changes in its interannual precipitation. Precipitation has a linear response to El Niño occurrences (Andreoli and Kayano 2005). In regard to long-term precipitation, south of 15°S, there were positive jumps east of the Andes, with a negative trend toward the west (Minetti and Vargas 1997). Historical positive jumps occurred between 1946 and 1960 while the negative trend diminished from north (Antofagasta station) to south (Islote Evangelista meteorological station).

In Southern Chile, and according to records measured at Valdivia, there was a significant decrease in precipitations between 1901 and 2005 (González-Reyes and Muñoz 2013). On the other coast, at the Argentine Pampas, there was an increase in 50-200 mm in the annual rains comparing two intervals: 1947-1976, and 1977-2006 (Forte Lay et al. 2008). The Pampa Region increases its Precipitation rates during the last decades of 20th century (Scarpatti and Capriolo 2013). Several authors point to the early 70's as the epoch of significant increases in runoffs of the rivers Paraguay and Paraná (Pasquini and Depetris 2007). Notwithstanding this natural climatic scheme, significant variations in South America should be assigned to changes in the land use and land covers (Clark et al. 2012).

71

72 3. Methods

Monthly hydrological records were compiled and analyzed from the databases of different South American countries (Table 1).

74

Peru	http://www.senamhi.gob.pe/
Ecuador	http://www.serviciometeorologico.gob.ec/caudales-datos-historicos/
Chile	http://snia.dga.cl/BNAConsultas/reportes
Brazil	http://hidroweb.ana.gov.br/HidroWeb.asp?ToItem=4100
Argentina	http://www.mininterior.gov.ar/obras-publicas/rh-base.php

Table 1. Web pages of the hydrological records of different countries of South America.

Historical maps and TM images of the Landsat satellites were compared in order to discriminate the extension of flooding episodes from normal conditions.

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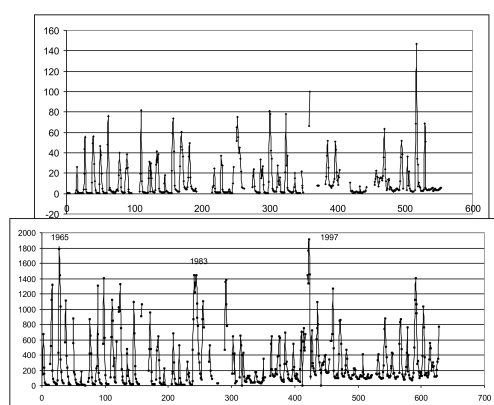
79 4. Results

80 4.1. Ecuador and Peru

According to the Ecuadorian INAMHI institution, the largest floods connected to strong El Niño phenomena occurred in 1977-78, 1982-83 and 1997-98. The floods of 1983 were triggered by enormous amounts of rainfall at Western Ecuador (Rossel et al. 1996). The impacts caused by the strong ENSO of 1997-98 were estimated according to different economic sectors: agriculture (43.6 MU\$), infrastructure of the sanitary sector (27.5 MU\$), housing (3.2 MU\$) and industry (9.5 MU\$; Vaca 2010).



86 The Daule River (Fig. 1) was flooded in 1965, 1983, 1997-98 and 2012 (Fig. 2). Normally these maximum
87 discharges occurred during the first months of the year (January to May). The worst floods were during the first
88 months of 1998 with discharges over 1300 m³/s (Fig. 2). At the boundary between Ecuador and Peru, the Zarumilla
89 River increased its discharge during the years 1965, 1973, 1983 and 1998.



90
91 *Fig. 2. Monthly discharges of the Carrizal River (top) and Daule River in La Capilla showing peaks in 1965, 1983,*
92 *1997 and 2012.*

93 Perú hydrological statistics are published every year by the SENAMHI (2016). The Rimac River (Fig. 1) flooded in
94 1925 with a maximum daily discharge of 600 m³/s; an event considered the first “meganiño” of the 20th century
95 (Rocha Felices 2011). It was also flooded in 1941 (385 m³/s) and 1955 (380 m³/s). Historical data from Perú
96 indicates that there is a patchy distribution between different basins (Waylen and Caviedes 1986).

97

98 4.2. Chile

99 Although Chilean floods may occur by different origins, 71% are associated to rainfalls. However, rainfalls are
100 assumed to be diminishing in a long-term scenario (González-Reyes and Muñoz, 2013). Those floods associated
101 exclusively to strong ENSO events occur northwards of 36°S (Rojas et al. 2014). However, significant discharges
102 also occur at the south, but masked to other floods triggered by local rains (Araya Ojeda and Isla 2016).

103 There is not a definite effect of ENSO anomalies along the whole Chile. Those rivers of Northern Chile comprised
104 within the South American Arid Diagonal are specifically subject to anomalous precipitations. The two debris flow
105 recorded in Antofagasta in 1940 (Vargas et al 2000) could have been also connected to the strong ENSO of
106 1941. The 1982-83 and 1997-98 ENSO rainfalls affected significantly Northern Chile (Meza 2013; Vargas et al.
107 2006).

108 In Central Chile, the higher discharges of the Aconcagua River (Fig. 1) were related to ENSO events but with a
109 certain delay (Waylen and Caviedes 1990). For the interval 1901-2005 there was a significant reduction of annual
110 precipitation for the Valdivia region, southern Chile (González-Reyes and Muñoz, 2013). It has been proposed that
111 the reduction in water yields in South-Central Chile is caused by land-use changes derived from the replacement of
112 native forest by exotics (Little et al. 2009); afforestation significantly affect runoff at the Biobio Region (Iroumé and
113 Palacios 2013).



114

115 4.3. Brazil

116 Anomalous years affected some cities of Brazil. The floods of 1941 affected 70,000 inhabitants at the riverine area
 117 of Porto Alegre (Fig. 3). City authorities constructed a dike in order to prevent another flood of the Guaíba fluvial
 118 complex (Loitzenbauer et al. 2012).



119

120 *Fig. 3. A) Fluvial area of Porto Alegre flooded in 1941. B) Present area.*

121

122 The floods of 1983 of the Itajaí-Açu River caused the destruction of 30,000 houses at Blumenau. The level of the
 123 river raised 16 m over normal level and stayed high for several days. 80% of the Itajaí County was affected. The
 124 Iguaçu River, an affluent to the Paraná River (Fig. 1), flooded Curitiba in 1982 and 1983 signifying losses of 10,000
 125 and 78,000 MU\$S respectively at some neighborhoods (Tucci and Petry 2006).

126 4.4. Bolivia

127 The Pilcomayo River (Fig. 1) has a maximum discharge of 3500 m³/s, about 45 times its minimum discharge (80
 128 m³/s; Rabicaluc 1986). This river has an alluvial fan of 210,000 km² with several abandoned channels (Iriondo et al.
 129 2000). The floods of the Upper Pilcomayo River of 1983 and 1984 (Fig. 4) increased 2-3 times the amount of
 130 sediment transported in suspension (Malbrunot 2006). During normal years the river transports less than 1 x 10⁶ tons
 131 of sediments; in 1984 it transported 2 and 3 millions of tons. The city of Villamontes (Tarija) is usually flooded by
 132 the Pilcomayo River. Although a hydrologic gauge was installed in 1941, it operated randomly. A maximum level of
 133 7.98 m was measured in March, 1984 with a maximum discharge of 7000 m³/s (Ribstein and Peña 1993). The last
 134 flood was recorded during the beginning of 2018.

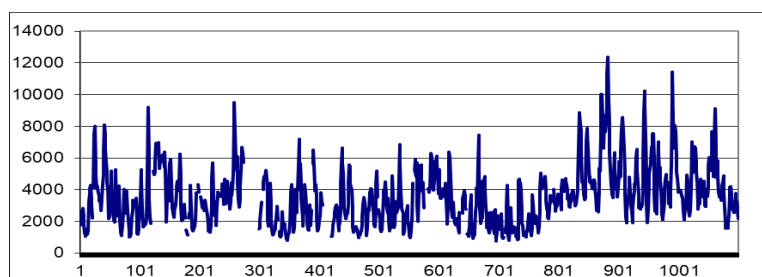


Fig. 4. Pilcomayo River floods at Puerto Pilcomayo with the peaks in 1983, 1997 and 2015.

The Bermejo River (Fig. 1) also flows from Bolivia to Argentina, to the Paraná River. Based on historical archives, its hydrological cycles have been reconstructed (Prieto and Rojas 2015). Floods progressively increased since 1800. The deforestation has increased the climatic effects. Floods frequency diminished during the first half of the 20th century but increased significantly to the end of that century. These rivers that flow from the Andes to the Parana River (Paraguay, Pilcomayo, Bermejo) carried significant amount of particulate and dissolved substances. The plume of the Paraguay River persists isolated from the Upper Paraná water during approximately 225 km (Campodónico et al. 2015).

4.5. Paraguay

The Republic of Paraguay is located between three rivers (Paraguay, Paraná and Pilcomayo), all belonging to the Río de la Plata watershed (Baez et al. 2014). The Paraguay River (Fig. 1) is about 2800 km long draining an area of about 1,095,000 km² (Collischonn et al. 2001). This large basin should be analyzed according to two regions: the northern related to the Amazonas River system, and the southern, subject to ENSO-triggered floods (Drago et al. 2008). This watershed is in close relation to the Patiño Aquifer (Monte Domecq and Baez Benítez 2007) and the Pantanal wetlands (Collischonn et al. 2001). During the winters of 1982 and 1983 the river had discharges of 9712 and 10663 m³/s, respectively (Monte Domecq et al. 2003; Barros et al. 2004). The two largest floods of the riverine areas of Asunción occurred in 1983 (63 m over MSL) and 1992 (62.3 m). During the floods of 1997/98 24,975 inhabitants were evacuated from Asunción, and 54,000 inhabitants from other departments (Neembucú, Concepción, Cordillera and Chaco). One of the major risks of the Paraguay River floods is that concerning to the Cateura waste disposal of the Asunción city (Fig. 5). This dumping site is on the floodplain and very close to the international boundary with Argentina.



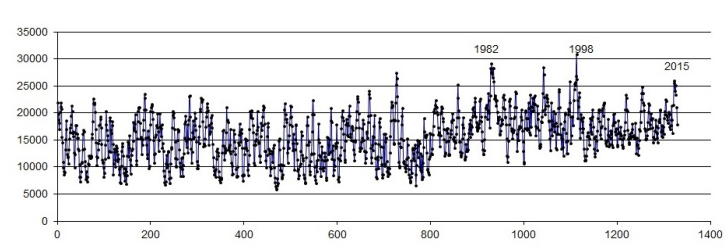
Fig. 5. During the ENSO 2015-2016 the waste disposal site of Asunción was flooded very close to the international boundary with Argentina.



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162 **4.6. Argentina and Uruguay**

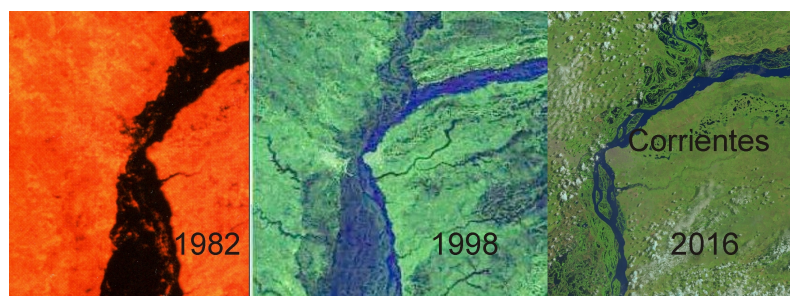
163 The Rio de la Plata received significant amount of water from the Upper Paraná (83%), 20% from the Paraguay
 164 River, and about 7% from the rivers flowing from the west (Bermejo, Pilcaomayo and Salado; Pasquini and
 165 Depertris 2010). The Paraná River (Fig. 1) flooded systematically during the last strong ENSOs (1982-83, 1997-98
 166 and 2015; fig. 6) It multiplied 15 times the widths of its floodplain during the floods of 1982-83 (Drago 1989). This
 167 extraordinary event signified high monthly streamflows in Corrientes during a year and half (Camilloni and Barros
 168 2003).



169

170 *Fig. 6. Hydrological record of the Paraná River at Timbúes Station (1901-2016).*

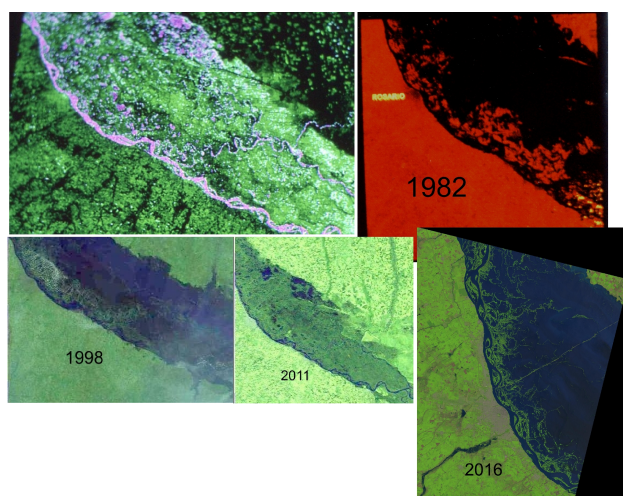
171 In Corrientes city, the discharge surpassed 10,000 m³/s from July 1982 to December 1983 (Camilloni and Barros
 172 2003), also affecting the localities of Resistencia, Barranqueras, Puerto Vilelas and Fontana (Fig. 7). The Paraná
 173 River at Barranqueras reached the maximum 8.6 m level. Below the General Belgrano Bridge the discharge was
 174 58,000 m³/s.



175

176 *Fig. 7. The floods of 1982, 1998 and 2016 affected the cities of Corrientes and Resistencia (Argentina).*

177 These floods signified the transport of subtropical floating plants (*Eichhornia crassipes*, also known as “water
 178 hyacinth”) to temperate areas, and carrying dangerous fauna with them (snakes, spiders and lizards). Several fluvial
 179 harbors as Rosario, Campana, Zárate and Buenos Aires were restricted in their operability during these events. At
 180 the floodplain close to Rosario, the peak flows of 1982-83, 1992, and 1997-98 exceeded 30,000 m³/s (Fig. 8).
 181 During these extraordinary floods, the floodplain stores between 23 and 123 x 10⁶ tons/year (García et al. 2015).



182

183 *Fig. 8. The Paraná River flooded several times restricting the operation of the harbor of Rosario.*

184 The Uruguay River (**Fig. 1**) flooded in 1941, 1983 and 1997-98 (Isla and Toldo 2013). It has a mean discharge of
 185 4315 m³/s (Evarsa 2006). Harmonic analysis shows a dry period during the 1950-1960 decade, recorded also at the
 186 Paraná watershed (Krepper et al. 2003).

187 The Colorado River (Northern Patagonia) is assumed to deactivate from the northern portion of the watershed
 188 during the Holocene. During the floods of 1982-83, the whole watershed connected for the last time and saline water
 189 arrived to the delta plain (Isla and Toldo 2013). Proglacial lakes of eastern Patagonia were also affected during
 190 ENSO years: increments in the amount of snow during ENSO years produce high water levels of these lakes with a
 191 lag of 13 months (Pasquini et al. 2008).

192 **5. Maximum floods**

193 Comparing the best recorded strong ENSOs (1982-83 and 1997-98) they produced similar maximum discharges in
 194 Chile (Araya Ojeda and Isla 2016). The strong ENSO of 1997-98 was stronger in Ecuador (Daule and Carrizal
 195 rivers). However, and comparing their maximum monthly discharges, these floods are correlated (Fig. 9). This
 196 correlation is useful to forecast the maximum discharges expected for future strong ENSOs. During the last two
 197 centuries, the three strong ENSOs occurred in less than 40 years (1982-82, 1997-98 and 2015-2016) and are
 198 therefore indicating a higher frequency in regard to previous years.

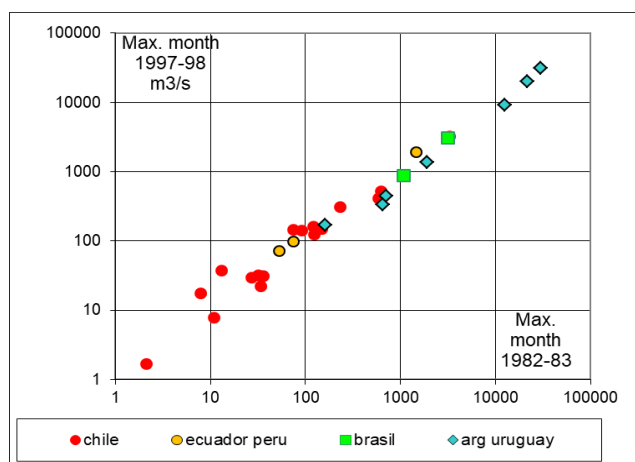


Fig. 9. Comparison of maximum discharges (m³/s) of the floods of 1982-83 and 1997-98.

6. Discussion

The ENSO events recorded during the Holocene are highly controversial (Clement et al. 2000). Although this review is based on hydrological measurements from countries with short series, there are some records that could be applied as historical and tree-ring records. Paleoclimatic studies indicate that in Northern Chile (north of 30° S) there was an absence of heavy rainfalls between 8400 and 5300 years BP in conjunction to a decrease in ENSO activities at the Eastern Pacific Ocean (Ortega et al. 2012). Based on tree rings from the Bermejo River region, it was stated that for the last three centuries there was significant increments in the frequency, intensity and duration of floods and droughts since the second half of the 20th century (Ferrero et al. 2015). The last five extreme wet events occurred since 1814, the last three in the last 40 years. However, there were significant droughts in Western Pampas: the “Pampas Dust Bowl” occurred between 1930 and 1940 (Viglizzo and Clark 2006). Summarizing, for Northeastern Argentina 1901-1960 was a dry period while 1970-2003 was characterized by wet conditions (Lovino et al. 2014). ENSO floods occur with a different delay between the high-relief Andes watersheds draining towards the Pacific Ocean and those meandering towards the Atlantic Ocean. In Patagonia, the delay between the snow recharge and the raise in the piedmont lakes levels is about 13 months (Pasquini et al. 2008).

ENSO cycles do not only affect the hydrological records of South America. They also affect rivers of China causing variations in their sediment discharge (Liu et al. 2017), and can therefore considered a good predictor for flood-affected and flood-destroyed crop areas (Zhang et al. 2016).

ENSO cycles, either Niños or Niñas, have significant effects on the global price of wheat. Niños cause reductions of 1.4% in its production while Niñas cause reductions of 4% (Ubilava 2017). Niños have positive effects regarding crop yields at the Argentine Pampa; maize and wheat yields increase during ENSOs, while the increase in soybean only occurred along some areas (Magrin et al. 1998). On the other hand, sunflower yields diminish during ENSO years. These increments are more significant in the north of the Pampas region (Fernández Long et al. 2011).

7. Conclusions



- 226 1. Strong ENSO floods affected South America in 1941, 1982-83, 1997-98 and 2015-16.
- 227 2. Rivers from the South American Arid Diagonal are only affected by ENSO floods. Those outside the
- 228 diagonal can be also affected by anomalous precipitations derived from the Atlantic or Pacific oceans.
- 229 3. Comparing the monthly discharges of several rivers, the 1982-82 and 1997-98 floods were of similar
- 230 magnitudes and should be considered to forecast future strong events and to organize mitigating plans.

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234 Pages.

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