

1 **A discrete wavelet spectrum approach for identifying non-monotonic trend**
2 **patterns of hydroclimate data**

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23 **Abstract:** The hydroclimatic process is changing non-monotonically and identifying its trend
24 pattern is a great challenge. Building on the discrete wavelet transform theory, we developed a
25 discrete wavelet spectrum (DWS) approach for identifying non-monotonic trend patterns in
26 hydroclimate time series and evaluating their statistical significance. After validating the DWS
27 approach using two typical synthetic time series, we examined annual temperature and potential
28 evaporation over China from 1961-2013, and found that the DWS approach detected both the
29 “warming” and the “warming hiatus” in temperature, and the reversed changes in potential
30 evaporation. Further, the identified non-monotonic trend patterns showed stable significance
31 when the time series was longer than 30 years or so (i.e., the widely defined “climate” timescale).
32 The significance of trends in potential evaporation measured at 150 stations in China, with an
33 obvious non-monotonic pattern, was underestimated and was not detected by the Mann-Kendall
34 test. Comparatively, the DWS approach overcame the problem and detected those significant
35 non-monotonic trends at 380 stations, which helped understand and interpret the spatiotemporal
36 variability of the hydroclimatic process. Our results suggest that non-monotonic trend patterns
37 of hydroclimate time series and their significance should be carefully identified, and the DWS
38 approach proposed has the potential for wide use in hydrological and climate sciences.

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40 **Key words:** trend identification; discrete wavelet spectrum; decadal variability; statistical
41 significance; Mann-Kendall test

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45 **1. Introduction**

46 Climate and hydrological processes are exhibiting great variability (Allen and Ingram,
47 2002; Trenberth et al., 2014). Quantitatively identifying changing signals in the hydroclimate
48 process is of great socioeconomic significance (Diffenbaugh et al., 2008; IPCC, 2013) as an
49 important basis for hydrological modelling, understanding the future hydroclimatic regimes,
50 and water resources planning and management. However, it remains a challenge to both
51 scientific and social communities. The simplest and the most straightforward way to identify
52 changes in the hydroclimate process would be to fit a monotonic (e.g., linear) trend at a certain
53 time period at which a significance level would be assigned by a statistical test. Among the
54 methods used for the detection of trends, the Mann-Kendall non-parametric test is most widely
55 used and has been successfully applied in studies on climate change and its impact, when the
56 time series is almost monotonic as required and a statistical threshold of ± 1.96 is set to judge
57 the significance of trends at 95% confidence level (Burn and Hag Elnur, 2002; Yue et al., 2002).
58 However, due to its nonlinear and nonstationary nature, the hydroclimate process is changing
59 and developing in a more complicated way rather than a monotonic trend way at large time
60 scales (Cohn and McMahon, 2005; Milly et al., 2008). For example, a debate on the recent
61 change of global air temperature has been receiving enormous public and scientific attention
62 that the global air temperature increased during 1980-1998 passing most statistical significance
63 tests and has since stabilized till now, widely called “global warming hiatus” (Kosaka and Xie,
64 2013; Roberts et al., 2015; Medhaug et al., 2017). Another known example is “evaporation
65 paradox” (Brutsaert and Parlange, 1998; Roderick and Farquhar, 2002) that potential
66 evaporation has worldwide declined from the 1960s, again passing most statistical significance

67 tests, but then reversed after the 1990s. In practice, for the hydroclimate time series, non-
68 monotonicity is more the rule rather than the exception (Dixon et al., 2006; Adam and
69 Lettenmaier, 2008; Gong et al., 2010). Therefore, identifying the non-monotonic trend pattern
70 hidden in those hydroclimate time series and assessing its statistical significance present a
71 significant research task for understanding hydroclimatic variability and changes at large time
72 scales.

73 Among those methods presently used in time series analysis, the wavelet method,
74 including both continuous and discrete wavelet transforms, has the superior capability of
75 handling nonstationary characteristics of the time series at multi-time scales (Percival and
76 Walden, 2000; Labat, 2005), so it may be more suitable for identifying non-monotonic trend
77 patterns in hydroclimate time series at large time scales. In a seminal work, Torrence and
78 Compo (1998) placed the continuous wavelet transform in the framework of statistical analysis
79 by formulating a significance test. Since then, the continuous wavelet method has become more
80 applicable and rapidly developed to estimate the significance of variability in climate and
81 hydrological studies. Especially, the continuous wavelet spectrum (i.e., continuous wavelet
82 variance) was established to detect those significant variabilities in the hydroclimate process
83 (Labat et al., 2000). However, in the continuous wavelet results of a time series, a known
84 technical issue is the “data redundancy” (Gaucherel, 2002; Nourani et al., 2014), which is the
85 redundant information across timescales leading to more uncertainty.

86 On the contrary, the other type of wavelet transform, i.e., the discrete wavelet transform,
87 has the potential to overcome that problem of data redundancy, in that those wavelets used for
88 discrete wavelet transform must meet the orthogonal properties. Therefore, the discrete wavelet

89 method can be more effective to identify and describe the non-monotonic trend pattern in a time
90 series (Almasri et al., 2008; de Artigas et al., 2006; Kallache et al., 2005; Partal and Kucuk,
91 2006; Nalley et al., 2012). However, there lacked an effective discrete wavelet spectrum in the
92 wavelet methodology without which uncertainty in the discrete wavelet-aided identification of
93 a trend could not be accurately estimated, and the significance level of the identified trend could
94 not be quantitatively evaluated either. For overcoming the problem, Sang et al. (2013) discussed
95 the definition of trend, and proposed a discrete wavelet energy function-based method for the
96 identification of trends, with the basic idea of comparing the difference of discrete wavelet
97 results between hydrological data and noise. The method used a proper confidence interval to
98 assess the statistical significance of the identified trend, in which the key equation for
99 quantifying trend's significance was based on the concept of quadratic sum. However,
100 computation of the quadratic sum disobeys the customary practice of computing variance in
101 spectral analysis. By using the quadratic sum, the significance of a non-monotonic trend cannot
102 be reasonably assessed, because it neglects the big influence of trend's mean value. For instance,
103 for those trends with small variations but big mean values, the quadratic sums are big values,
104 based on which the statistical significance of trends would inevitably be over-assessed.
105 Therefore, evaluation of the statistical significance of a non-monotonic trend in a time series
106 should be based on its own variability, and the influence of other factors should also be
107 eliminated.

108 By combining the advantages of discrete wavelet transform and successful practice in
109 spectral analysis methods, this study aimed at developing a practical but reliable discrete
110 wavelet spectrum approach for identifying non-monotonic trend patterns in hydroclimate time

111 series and quantifying their statistical significance, and further improving the understanding of
 112 non-monotonic trends by investigating their variation with data length increase. To do that,
 113 Section 2 presents details of the newly developed approach building on the wavelet theory and
 114 spectrum analysis. In Section 3, we use both synthetic time series and annual time series of air
 115 temperature and potential evaporation over China as examples to investigate the applicability
 116 of the approach, which is followed by discussion and conclusion in the final section.

117 **2. A discrete wavelet spectrum approach**

118 Here we develop an approach, termed as “discrete wavelet spectrum approach,” for
 119 identifying non-monotonic trend patterns in hydroclimate time series, in which the discrete
 120 wavelet transform (DWT) is used first to separate the trend pattern at large time scales, and its
 121 statistical significance is then evaluated by using the discrete wavelet spectrum, whose
 122 confidence interval is quantified and described through Monte-Carlo test.

123 Following the wavelet analysis theory (Percival and Walden, 2000), the discrete wavelet
 124 transform of a time series $f(t)$ with a time order t can be expressed as:

$$125 \quad W_f(j, k) = \int_{-\infty}^{+\infty} f(t) \psi_{j,k}^*(t) dt \quad \text{with} \quad \psi_{j,k}(t) = a_0^{-j/2} \psi(a_0^{-j}t - b_0k) \quad (1)$$

126 where $\psi^*(t)$ is the complex conjugate of the mother wavelet $\psi(t)$; a_0 and b_0 are constants, and
 127 integer k is a time translation factor; and $W_f(j, k)$ is the discrete wavelet coefficient under the
 128 decomposition level j (i.e., time scale a_0^j). In practice, the dyadic DWT is used widely by
 129 assigning $a_0=2$ and $b_0=1$:

$$130 \quad W_f(j, k) = \int_{-\infty}^{+\infty} f(t) \psi_{j,k}^*(t) dt \quad \text{with} \quad \psi_{j,k}(t) = 2^{-j/2} \psi(2^{-j}t - k) \quad (2)$$

131 The highest decomposition level M is determined by the length L of series $f(t)$, and can be
132 calculated as $\log_2(L)$ (Foufoula-Georgiou and Kumar, 2014). The sub-signal $f_j(t)$ in the original
133 series $f(t)$ under each level j ($j = 1, 2, \dots, M$) can be reconstructed as:

$$134 \quad f_j(t) = \sum_k W_f(j, k) \psi^*(2^{-j}t - k) \quad (3)$$

135 where the sub-signal $f_j(t)$ at the highest decomposition level (when $j=M$) defines and describes
136 the non-monotonic trend pattern of the series $f(t)$, as generally understood. However, it should
137 be noted that a meaningful trend closely depends on the time scale concerned. If the variability
138 of series $f(t)$ on a certain smaller time scale K ($K < L$) is concerned, the proper decomposition
139 level can be determined as $\log_2(K)$, then the sum of all those sub-signals at the time scale equal
140 to and bigger than K can be the non-monotonic trend pattern identified.

141 Sang (2012) discussed the influence of the choice of the mother wavelet and
142 decomposition level, as well as noise types on the discrete wavelet decomposition of time series,
143 and further proposed some methods to solve for them. By conducting Monte-Carlo experiments,
144 he found that the seven wavelet families (126 mother wavelets) used for DWT can be divided
145 into three types, and recommended the first type, by which wavelet energy functions of diverse
146 types of noise data keep stable and thus have little influence on the wavelet decomposition of
147 time series. Specifically, one chooses an appropriate wavelet, according to the relationship of
148 statistical characteristics among the original series, de-noised series and removed noise, chooses
149 a proper decomposition level by analyzing the difference between energy function of the
150 analyzed series and that of noise, and then identifies the deterministic components (including
151 trend) by conducting significance testing of DWT. These methods are closely built on the
152 composition and variability of hydroclimate time series at different time scales. They were used

153 here to accurately identify and describe the non-monotonic trend pattern in a time series, and
154 assess its statistical significance.

155 Further, to establish a reliable discrete wavelet spectrum (DWS) of time series, we need
156 to specify a spectrum value $E(j)$ for each sub-signal $f_j(t)$ (in Eq. 3), based on which we can
157 quantitatively evaluate its importance and statistical significance. Following the general
158 practice in conventional spectral analysis methods (Fourier transform, maximum entropy
159 spectral analysis, *etc.*), here we define $E(j)$ at the j th level by taking the variance of $f_j(t)$:

$$160 \quad E(j) = \text{var}(f_j(t)) \quad (4)$$

161 It can accurately quantify the intensity of variation of sub-signals (including trend) by
162 eliminating the influence of their mean values, which is different from the quadratic sum-based
163 method proposed by Sang et al. (2013). For hydroclimate time series, both stochastic and
164 deterministic components generally have distinctive characteristics from purely noise
165 components (Sang et al., 2012; Rajaram et al., 2015). Due to the grid of dyadic DWT (Partal
166 and Cigizoglu, 2008), discrete wavelet spectra $E_r(j)$ of various noise types strictly follow an
167 exponentially decreasing rule with a base 2 (Sang 2012):

$$168 \quad E_r(j) = 2^{-j} \quad (5)$$

169 The discrete wavelet spectra of deterministic components and that of noise are obviously
170 different. Hence, we define the DWS of noise data as the “reference discrete wavelet spectrum
171 (RDWS)”, based on which we evaluate the statistical significance of the non-monotonic trend
172 pattern of a time series.

173 To be specific, we design a technical flowchart to show how we develop the DWS
174 approach for identifying the non-monotonic trend pattern of time series, and also for evaluating
175 the statistical significance of that trend pattern (see details in Figure 1):

176 (1) For the series $f(t)$ with length L to be analyzed, we normalize it, and decompose it using
177 the DWT method in Eq. (2) and (3);

178 (2) We calculate the discrete wavelet spectrum of the series $f(t)$ by Eq. (4);

179 (3) For comparison, we then use the Monte-Carlo method to generate normalized noise data
180 N with the same length as the series $f(t)$, and compute its RDWS by Eq. (4). Considering
181 that discrete wavelet spectra of diverse types of noise data consistently follow Eq. (5),
182 here we generate noise data following the standard normal probability distribution;

183 (4) We repeat the above step 5000 times, and calculate the mean value and variance of the
184 spectrum values (in Eq. 4) of the normalized noise data N at each decomposition level j .
185 Based on it, we estimate an appropriate confidence interval of RDWS at the concerned
186 confidence level. In this study, we considered 95% confidence level;

187 (5) In comparing DWS of the series $f(t)$ and the confidence interval generated by that of
188 noise (i.e., RDWS), we identified the deterministic components under the highest
189 decomposition level as the non-monotonic trend pattern of the series, and judged
190 whether it was significant. Specifically, if the spectrum value of the analyzed series' sub-
191 signal under the highest level was above the confidence interval of RDWS, it was
192 considered that the non-monotonic trend pattern was statistically significant; otherwise,
193 if the spectrum value of the sub-signal under the highest level fell into the confidence
194 interval of RDWS, it was not statistically significant;

195 (6) If a smaller time scale K is concerned, we can use the decomposition level $\log_2(K)$,
196 instead of M , and then repeat the steps (1-5) to identify the non-monotonic trend pattern
197 at that time scale.

198 In the following section, we mainly investigate the applicability and reliability of the DWS
199 approach for identifying the non-monotonic trend and assessing its significance, and further
200 investigate the variation of non-monotonic trend with data length increase to improve our
201 understanding of trend at large time scales.

202 < **Figure 1** >

203 **3. Results**

204 **3.1 Synthetic series analysis**

205 To test and verify the reliability of the developed discrete wavelet spectrum (DWS)
206 approach for identifying the non-monotonic trend pattern of a time series, we considered the
207 general hydrological situations and generated two synthetic series data, with known signals and
208 noise a priori. For investigating the variation of non-monotonic trend with data length increase,
209 we set the length of the two synthetic series as 200, and the noise in them followed a standard
210 normal probability distribution. The first synthetic series S1 consisted of an exponentially
211 increasing line and a periodic curve (with a periodicity of 200) with some noise content (Figure
212 2, left panel); and the second synthetic series S2 was generated by including a hemi-sine curve,
213 a periodic curve (with a periodicity of 50) and some noise content (Figure 2, right panel). Using
214 the MK test and considering monotonic trends, series S1 showed a significant increase but the
215 trend of series S2 was not significant.

216 When using the DWS approach (Figure 1), we considered the time scale as data length,
217 and used the Daubechies (db8) wavelet to decompose series S1 into seven (i.e., $\lfloor \log_2 200 \rfloor$) sub-
218 signals using Eq. (2) and Eq. (3). Then, we took the sub-signals under the seventh level as the
219 defined non-monotonic trend pattern. As shown in Figure 2 (left panel), the identified non-
220 monotonic trend pattern in series S1 was similar to the true trend pattern. However, the linear
221 fitting curve (a monotonic curve) could not capture the detail of the non-monotonic trend pattern.
222 The same approach applied to series S2 in Figure 2 (right panel) and the conclusion did not
223 change. Moreover, for series S2 with large variability at large time scales, the linear fitting
224 curve or other monotonic curves may not be physically meaningful.

225 < **Figure 2** >

226 We computed the discrete wavelet spectra of the two synthetic series using Eq. (4), and
227 used the reference discrete wavelet spectrum with 95% confidence interval to evaluate the
228 statistical significance of their non-monotonic trend patterns. That is, if the red point at a certain
229 data length was above the 95% confidence bar, described by the blue line in Figure 3, it was
230 considered that the trend pattern was significant at 95% confidence level. Using the DWS
231 approach, the trend pattern of series S1, which was quasi-monotonic, was found significant
232 (Figure 3a) as in the MK test (Figure 3c), but the non-monotonic series S2 showed a significant
233 trend pattern (Figure 3b), which was greatly different from the MK test (Figure 3d).

234 In Figure 3, we also presented the significance of the identified trend patterns of the two
235 series using both our DWS approach and the MK test, and we changed the data length to
236 investigate the stability of the statistical significance of the non-monotonic trend pattern.
237 Generally, it would have more uncertainty when evaluating the statistical significance of trend

238 pattern with a shorter length, corresponding to a bigger 95% confidence interval. Using our
239 DWS approach, the 95% confidence interval (i.e., the height of blue bars in Figure 3) for
240 evaluating the statistical significance of trend pattern generally decreased with the increase of
241 data length, as expected. However, in the MK test, the significance was always determined by
242 the constant thresholds of ± 1.96 , regardless of the data length.

243 In the DWS results in Figure 3, the significance levels of non-monotonic trend patterns
244 did not consistently decrease with data length, but showed some fluctuation, as the proportions
245 of different components (including trend) in the original series varied with data length.
246 Furthermore, one would expect that if the trend pattern of a series at a certain length was
247 identified statistically significant, the trend pattern would extend with the increase of data length,
248 thus its significance may be more stable with a larger length of data considered. Using our DWS
249 approach, the trend pattern of series S1 was significant when the data length was larger than 55
250 (Figure 3a), being similar to the result of the MK test (Figure 3c). The trend pattern of the series
251 S2 was statistically significant when the data length was larger than 75 (Figure 3b). However,
252 using the MK test, the monotonic trend of series S2 was significant only when the data length
253 was between 40 and 185 (Figure 3d). In summary, the significance of trend pattern identified
254 by our DWS approach was more stable than that detected by the MK test, demonstrating the
255 advantage of the DWS approach in dealing with non-monotonic variation of hydroclimate time
256 series.

257 < **Figure 3** >

258 **3.2 Observed data analysis**

259 We used the annual time series of air temperature (denoted as TEM) and potential
260 evaporation (denoted as PET) over China to further verify the applicability of our developed
261 DWS approach for identifying non-monotonic trend patterns of a time series. These time series
262 were obtained from the hydroclimate data measured at 520 meteorological stations over China,
263 with the same measurement years from 1961 to 2013. The data have been quality-checked to
264 ensure their reliability for scientific research. The PET data were calculated from the Penman-
265 Monteith approach (Chen et al., 2005).

266 The average time series of TEM and PET measured at 520 stations were first considered.
267 Given the general nonstationary nature of observed hydroclimate time series, linear trends or
268 more generally monotonic curves could not capture the trend pattern with large interdecadal
269 variation and therefore were not particularly physically meaningful. In Figure 4 (left panel), we
270 presented the average annual TEM time series visually showing nonstationary characteristics
271 and non-monotonic variation. The TEM series decreased till the 1980s with fluctuations and
272 then sharply rose till the 2000s, followed by a decreasing tendency. The large fluctuation of the
273 average air temperature after the late 1990s is the well-known phenomenon of the “global
274 warming hiatus” (Roberts et al., 2015). The linear fitting curve obviously missed out the more
275 complicated trend pattern of the observed temperature time series. Using our DWS approach,
276 we decomposed the TEM series into five (i.e., $\langle \log_2 53 \rangle$) sub-signals using Eq. (2) and Eq. (3),
277 and took the sub-signals under the fifth level as the trend pattern, which realistically presented
278 the nonstationary variability of temperature at large time scales (Figure 4, left panel).

279 We also applied the DWS approach to the average annual PET time series. In the time
280 series of PET (Figure 4, right panel), there was a decreasing trend for the period from 1961 to

281 the 1990s, which is the well-known “evaporation paradox” leading to controversial
282 interpretations continuing over the last decade of hydrological cycles (Brutsaert and Parlange,
283 1998; Roderick and Farquhar, 2002). That decreasing trend was then followed by an abrupt
284 increase in around the 1990s, almost the same time when solar radiation was observed to be
285 reversing its trend, widely termed as “global dimming to brightening” (Wild, 2009).
286 Surprisingly, after the mid-2000s, PET started to decrease again (Figure 4, right panel).
287 Sometimes, one would propose to fit linear curves for separate time periods. Again, linear
288 curves could not capture the overall non-monotonic trend pattern of the PET series. Using the
289 same DWS approach, we identified the non-monotonic trend pattern of the PET time series
290 (Figure 4, right panel), which captured the two turning points of the changing trends in the
291 1990s and the 2000s.

292 < **Figure 4** >

293 The changes of trends in terms of magnitudes and signs for different periods led to the
294 difficulty in assessing and interpreting the significance of trends. For example, the PET time
295 series showed a significant decrease using the MK test ($-3.76 < -1.96$) during 1961-1992 (Figure
296 5d). At that moment before the reversed trend reported, the significant decrease could be
297 literally interpreted, as that PET had significantly declined and might be declining in the future.
298 However, the PET time series reversed after the 1990s and again in the 2000s, coming with an
299 insignificant overall trend for the whole period of 1961-2013. For the more or less monotonic
300 time series of the TEM series (1961-2013), the MK test detected a significant increase ($6.00 >$
301 1.96) (Figure 5c), which led to the surprise when air temperature was reported to have stopped
302 increasing after the late 1990s. In summary, it becomes vital to develop an approach for testing

303 the significance of trend pattern, which is suitable for non-monotonic time series, as it is an
304 important basis and a prerequisite for hydrological simulation and prediction at decadal scales.

305 In this study, building on the discrete wavelet transform, we proposed an operational
306 approach, i.e., DWS, for evaluating the significance of non-monotonic trend pattern in TEM
307 (Figure 5a) and PET (Figure 5b) series. For comparison purposes, we also conducted the
308 significance test for the two time series using the MK test (Figure 5c and 5d). Similar to Figure
309 3, we changed the data length to investigate the stability of statistical significance (Figure 5).
310 Again, results indicated that the 95% confidence interval for evaluating the statistical
311 significance of non-monotonic trend pattern generally decreased with data length, which was
312 different from the constant thresholds ± 1.96 adopted in the MK test. The significance test
313 using our DWS approach appeared to be more stable with data length than the MK test (Figure
314 5). Using our DWS approach, the trend pattern in the TEM series became significant when the
315 data length increased to 30, and the significance was more stable when it was greater than 35
316 (Figure 5a). For the case of the PET series, the trend pattern became statistically significant
317 when the data length was larger than 25 (Figure 5b). The findings here have important
318 implications for non-monotonic hydroclimate time series analysis, in that the timescale of
319 defining *climate* and *climate change* by the World Meteorological Organization is usually 30
320 years (Arguez and Vose, 2011) and in hydrological practice it is between 25-30 years.

321 For the whole time series investigated here, whose length was larger than 30 years, we
322 were able to examine the significance using the developed DWS approach. Combining the trend
323 pattern in Figure 4 (left panel) and the significance test in Figure 5a, we confirmed that the
324 trend pattern of the TEM time series from 1961-2013 identified in this study was significant at

325 95% confidence interval. Similarly, the trend pattern in PET was also significant (Figure 4 right
326 panel and Figure 5b). The significance test results suggested that the three main stages of the
327 series (red lines, Figure 4) were detectable as the overall trend pattern from the variability of
328 the series and were vital to understanding how the temperature and the PET series were
329 changing at interdecadal scales. In particular, the reversed change in PET and its significance
330 can be revealed by our DWS approach, which can provide more useful and physically
331 meaningful information.

332 < **Figure 5** >

333 We further detected and evaluated the significance of non-monotonic trends of the PET
334 time series measured at 520 stations for investigating their spatial difference. Because the trends
335 in the annual TEM time series were quasi-monotonic, and they were statistically significant at
336 most of the stations, no matter using our DWS approach or the MK test, more details of TEM
337 data were not repeated here. As for the trend patterns in the PET data, the results gotten from
338 our DWS approach (Figure 6, left panel) and those in the MK test presented substantial
339 differences. When conducting the statistical significance test using the MK test, the monotonic
340 trends were detected as significant in those the annual PET time series measured at 230 stations.
341 Significant downward monotonic trends were mainly found in the southern part of the Songliao
342 River basin, the Haihe River basin, the Huaihe River basin, some regions in South China, and
343 Northwest China. Significant upward monotonic trends were mainly found in the northern part
344 of the Songliao River basin, the upper reach of the Yellow River basin, the southwest corner of
345 China, and some regions in the Yangtze River Delta.

346 Comparatively, significant non-monotonic trends in the PET time series were detected at
347 380 stations throughout China. That means that those annual PET time series measured at 150
348 stations (28.8% of the total stations and mainly in the south part of China) mainly indicated
349 non-monotonic variations rather than monotonic trends at interdecadal scales, with similar
350 phenomena as shown in Figure 4 (right panel), and their significance was underestimated by
351 the MK test, which can only handle monotonic trends. Previous studies (Zhang et al., 2016;
352 Jiang et al., 2007) indicated that potential evaporation was influenced by more physical factors
353 (precipitation, air temperature, wind speed, relative humidity, etc.) in the southern part of China
354 rather than the northern part; thus, the potential evaporation process in South China presented
355 a more complex variability and was more difficult to detect and attribute its physical causes. As
356 a result, it is known here that the annual potential evaporation process in most parts of China
357 indicated significance variability at interdecadal scales, but it was underestimated by the
358 conventional MK test; moreover, only considering monotonic trends would cause a great
359 difficulty in accurately understanding the temporal and spatial variability of potential
360 evaporation and hydroclimate process in China, and also would be unfavorable for hydrological
361 predictions at interdecadal scales. Our results suggest that the non-monotonic trend pattern of
362 hydroclimate time series and its significance should be carefully identified and evaluated.

363 < **Figure 6** >

364 **4. Summary and Conclusion**

365 Climate and hydrological processes are changing non-monotonically. Identification of
366 linear (or monotonic) trends in hydroclimate time series, as a common practice, cannot capture
367 the detail of the non-monotonic trend pattern in the time series at large time scales, and then

368 can lead to misinterpreting climatic and hydrological changes. Therefore, revealing the trend
369 pattern of the time series and assessing its significance from the usually varying hydroclimate
370 process remains a challenge. To that end, we develop the discrete wavelet spectrum (DWS)
371 approach for identifying the non-monotonic trend in hydroclimate time series, in which the
372 discrete wavelet transform is used first to separate the trend pattern, and its statistical
373 significance is then evaluated by using the discrete wavelet spectrum (Figure 1). Using two
374 typical synthetic time series, we examine the developed DWS approach, and find that it can
375 precisely identify non-monotonic trend pattern in the synthetic time series (Figure 2) and has
376 an advantage in significance testing (Figure 3).

377 Using our DWS approach, we identify the trend pattern in the annual time series of average
378 temperature and potential evaporation over China from 1961-2013 (Figure 4). The identified
379 non-monotonic trend patterns precisely describe how temperature and PET are changing at
380 interdecadal scales. Of particularly interest here is that the DWS approach can help detect both
381 the “warming” and the “warming hiatus” in the temperature time series, and reveal the reversed
382 changes and the latest decrease in the PET time series. The DWS approach can provide other
383 aspects of information on the trend pattern in the time series, i.e., the significance test. Results
384 show that the trend pattern becomes more significant and the significance test becomes more
385 stable when the time series is longer than a certain period like 30 years or so, the widely defined
386 “climate” time scale (Figure 5). Using the DWS approach, in both time series of mean air
387 temperature and potential evaporation, the identified trend patterns are found significant (Figure
388 5). Moreover, significance of trend patterns in the PET time series obtained from the DWS
389 approach and the MK test has obviously different spatial distributions (Figure 6). The variability

390 of hydroclimate process at large time scales, especially for non-monotonic trend patterns, would
391 be underestimated by the MK test, which causes a great difficulty in understanding and
392 interpreting the spatiotemporal variability of hydroclimate process. Comparatively, the
393 developed DWS approach can quantitatively assess the statistical significance of non-
394 monotonic trend pattern in the hydroclimate process, and so can meet practical needs much
395 better.

396 In summary, our results suggest that the non-monotonic trend pattern of hydroclimate time
397 series and its statistical significance should be carefully identified and evaluated, and the DWS
398 approach developed in this study has the potential for wider use in hydrological and climate
399 sciences.

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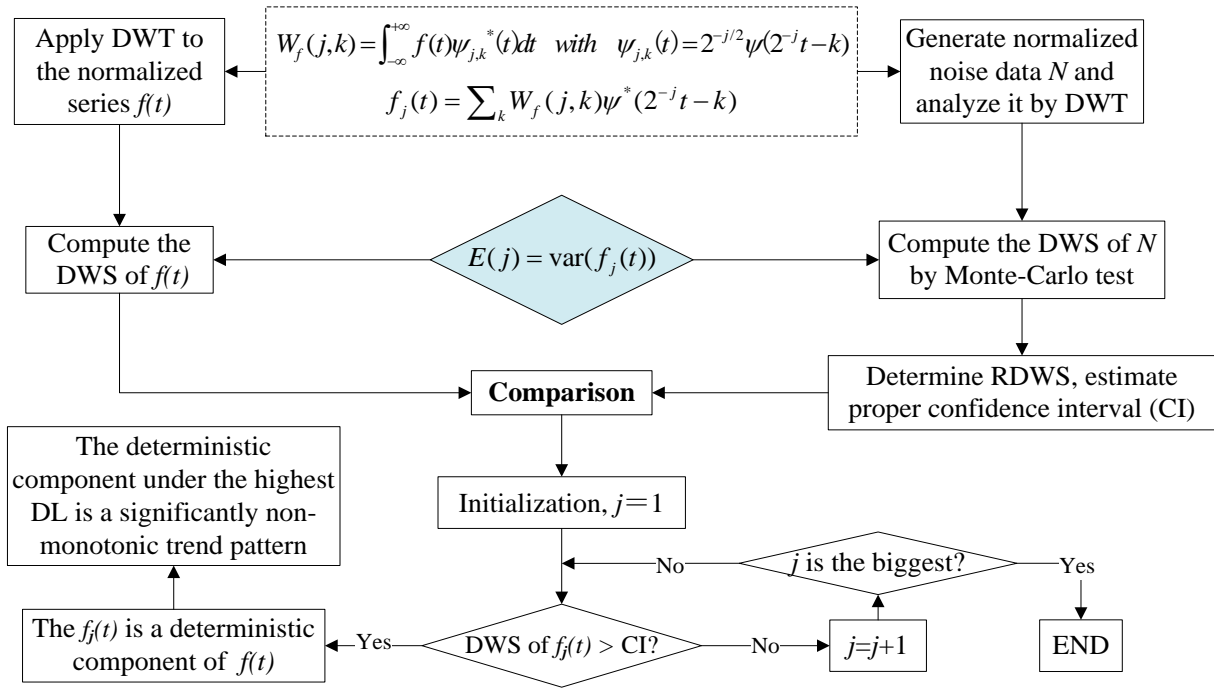
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520 **Figures**

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525 **Figure 1.** Technical flowchart for identification of the non-monotonic trend pattern in a time

526 series using the discrete wavelet spectrum (DWS) approach developed, where the discrete

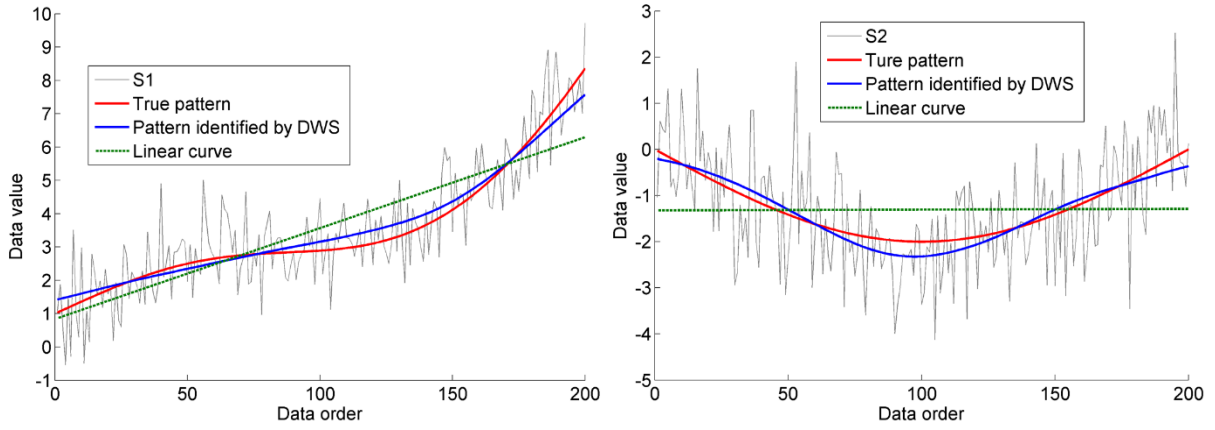
527 wavelet transform (DWT) method was used to decompose the time series, and the reference

528 discrete wavelet spectrum (RDWS) with certain confidence interval (CI) was used for the

529 evaluation of significance.

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533 **Figure 2.** Non-monotonic trend patterns in the synthetic series S1 and S2 identified by the

534 discrete wavelet spectrum (DWS) approach, and the linear trends in the two series.

535 Synthetic series S1 is generated as: $S1=1.112^{0.1t}+0.8\times\sin(0.01\pi t)+\alpha$; and synthetic series S2

536 is generated as: $S2=\sin(0.04\pi t)+2\times\sin(\pi+0.005\pi t)+\alpha$, where α is a random process

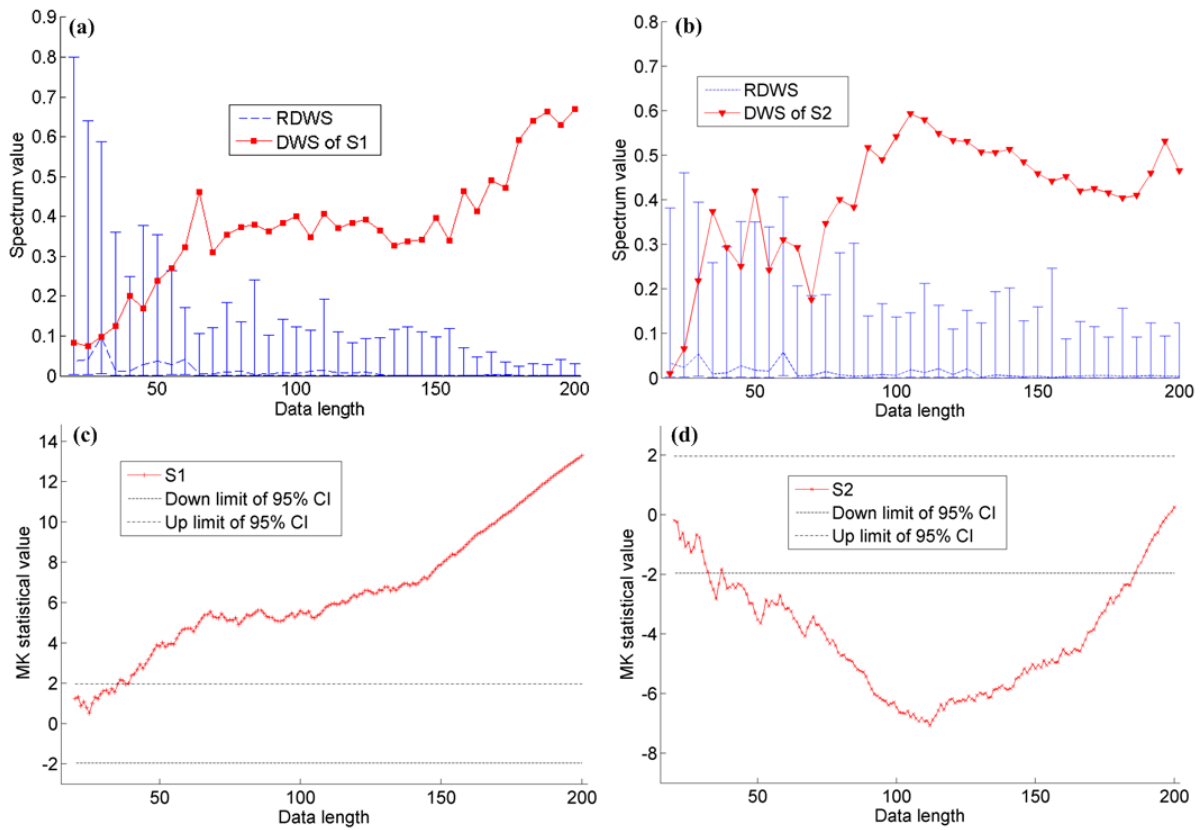
537 following the standard normal distribution.

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543 **Figure 3.** Evaluation of statistical significance of non-monotonic trend patterns in the synthetic

544 series S1 (a) and S2 (b) with different data length by the discrete wavelet spectrum (DWS)

545 approach, and the results by the Mann-Kendall (MK) test (c and d). In figure a and b, the

546 blue line is the reference discrete wavelet spectrum (RDWS) with 95% confidence interval

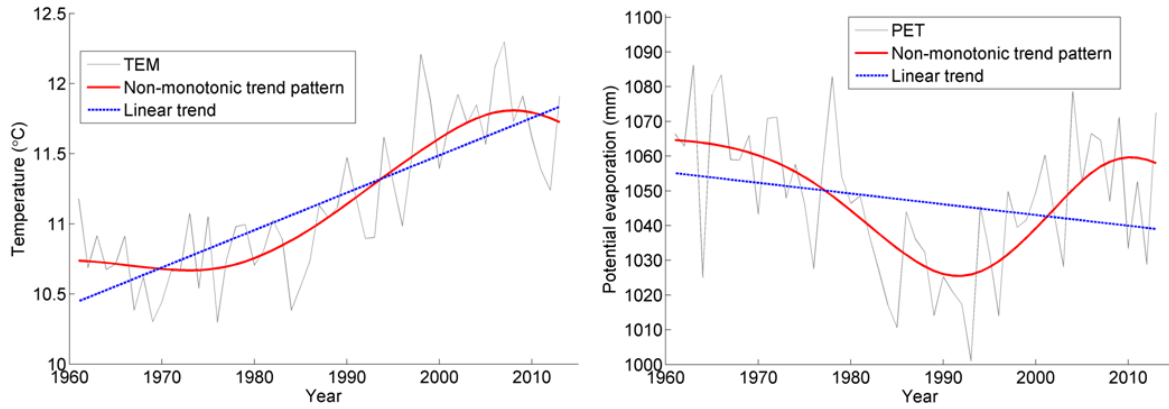
547 under each data length; if the red point at certain data length is above the blue bar, it is

548 thought that the trend pattern is significant at 95% confidence level. In figure c and d, the

549 two black dash lines indicate 95% confidence interval (CI) with the thresholds of +/- 1.96

550 in the MK test.

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553 **Figure 4.** Non-monotonic trend patterns in the annual time series of the mean air temperature

554 (TEM) and the potential evaporation (PET) over China from 1961-2013 identified by the

555 discrete wavelet spectrum (DWS) approach, and the linear trends in the two series.

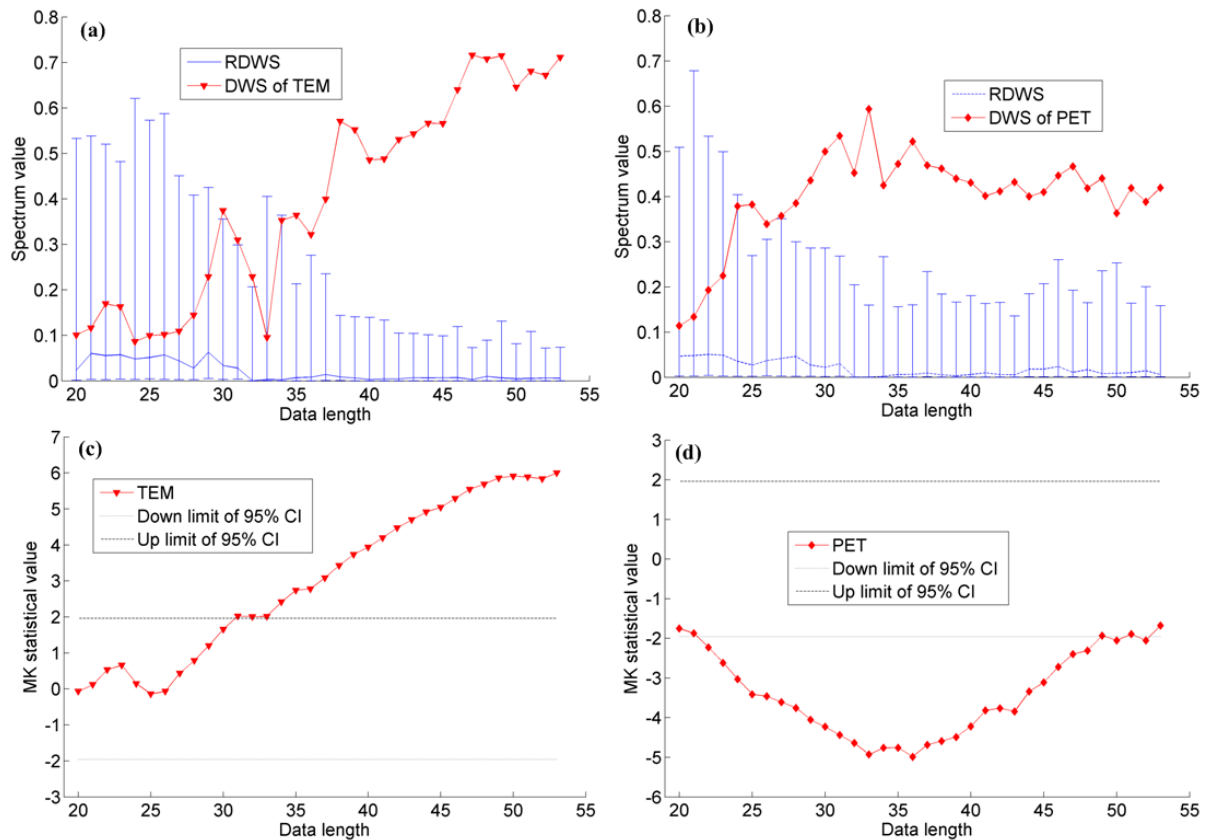
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Figure 5. Evaluation of statistical significance of non-monotonic trend patterns in the annual

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time series of the mean air temperature (TEM, a) and the potential evaporation (PET, b)

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over China with different data length by the discrete wavelet spectrum (DWS) approach,

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and the results by the Mann-Kendall (MK) test (c and d). In figure a and b, The blue line is

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the reference discrete wavelet spectrum (RDWS) with 95% confidence interval under each

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data length; and in figure c and d, the two black dash lines indicate 95% confidence interval

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(CI) with the thresholds of +/- 1.96 in the MK test.

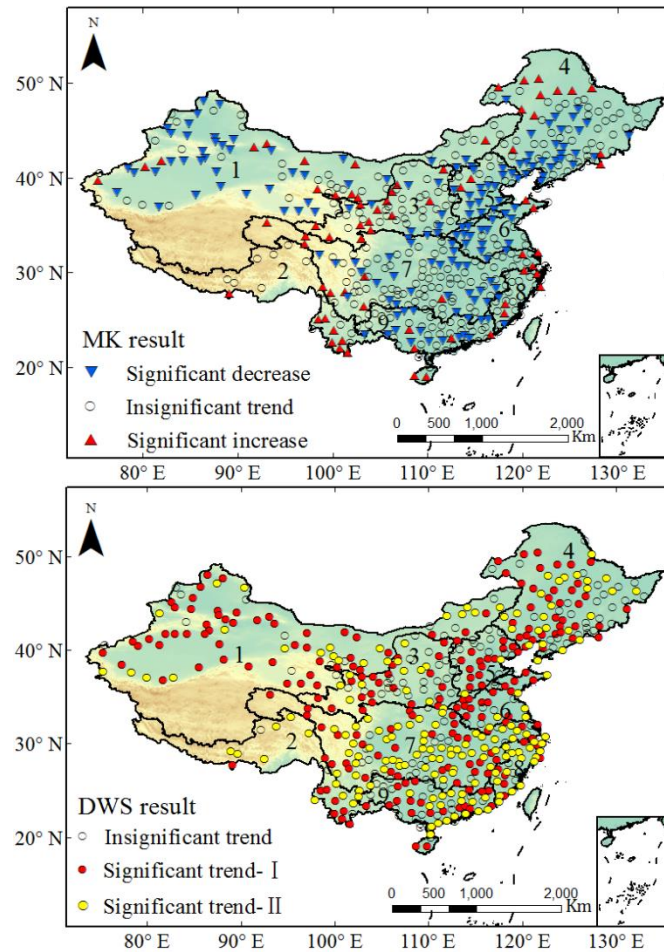
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575 **Figure 6.** Spatial distribution of the significance of trends in the annual potential evaporation

576 data during 1961-2013 and measured at 520 weather stations over China. The result above

577 was gotten from the Mann-Kendall (MK) test. The result below was gotten from the

578 discrete wavelet spectrum (DWS) approach developed, in which significant trend-I means

579 those significant trends (at 230 stations) can be identified by both the DWS approach and

580 the MK test, but significant trend-II means those significant trends (at 150 stations) can

581 only be identified by the DWS approach but not the MK test. 1, the Northwest Inland River

582 basin; 2, the Southwest River basin; 3, the Yellow River basin; 4, the Songliao River basin;

583 5, the Haihe River basin; 6, the Huaihe River basin; 7, the Yangtze River basin; 8, the

584 Southeast River basin; and 9, the Pearl River basin.