Topography significantly influencing low flows in snow-dominated watersheds

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Abstract. Watershed topography plays an important role in determining the spatial heterogeneity of ecological, geomorphological, and hydrological processes. Few studies have quantified the role of topography in various flow variables. In this study, 28 watersheds with snow-dominated hydrological regimes were selected with daily flow records from 1989 to 1996. These watersheds are located in the Southern Interior of British Columbia, Canada and range in size from 2.6 to 1,780 km^2. For each watershed, 22 topographic indices (TIs) were derived, including those commonly-used in hydrology and other environmental fields. Flow variables include annual mean flow (Q_{mean}), $Q_{10\%}$, $Q_{25\%}$, $Q_{50\%}$, $Q_{75\%}$, $Q_{90\%}$, and annual minimum flow (Q_{min}), where $Q_{x\%}$ is defined as the daily flow that occurred each year at a given percentage (x). Factor analysis (FA) was first adopted to exclude some redundant or repetitive TIs. Then, multiple linear regression models were employed to quantify the relative contributions of TIs to each flow variable in each year. Our results show that topography plays a more important role in low flows (flow magnitudes $\leq Q_{75\%}$) than high flows. However, the effects of TIs on different flow magnitudes are not consistent. Our analysis also determined five significant TIs: perimeter, slope length factor, surface area, openness, and terrain characterization index. These can be used to compare watersheds when low flow assessments are conducted, specifically in snow-dominated regions with the watershed size less than several thousand square kilometres.

Keywords: Topographic indices; Flow magnitudes; Snow-dominated; Relative contribution; Low flow.

1. Introduction

Topography plays a critical role in geomorphological, biological, and hydrological processes (Moore et al., 1991; Quinn et al., 1995). Many topographic indices (TIs) have been derived to describe the spatial patterns of a landscape (Yokoyama et al., 2002), locate spatial patterns of species (Jenness, 2004), and simulate spatial soil moisture (Park et al., 2001). In hydrology, hydrological responses are forced by climatic inputs (e.g. precipitation) but are controlled by topography and other factors such as land use and land cover (Beven and Kirkby, 1979; Hewlett and Hibbert, 1967). In describing the role of topography in hydrology, numerous TIs have been developed and applied to help understand hydrological processes and to explain the variation between watersheds (Moore et al., 1991). Although the importance of topography in controlling various flow magnitudes has been widely recognized (Price, 2011), the quantitative relationship between specific TIs and various flow variables is not well understood.

TIs can be categorized into two groups, namely, primary and secondary (or compounded) indices (Moore et al., 1988). Primary indices (e.g., slope, elevation, and aspect) are normally directly calculated from a digital elevation model (DEM), while secondary indices are the combination of primary indices that are used to explain the role of topography in geomorphology, biology, and hydrological processes. For instance, the topographic wetness index (TWI) is defined as: $\ln(\alpha/\tan\beta)$, where α is the upslope contributing area per unit contour length and β is the slope. TWI is a required

primary input for TOPMODEL and other hydrological applications (Beven, 1995; Beven and Kirkby, 1979; Quinn et al., 1995). Hydrological studies mainly focus on primary TIs and a few secondary TIs, however, these have limited explanatory powers as they largely fail to explain the variation in hydrological processes. Several TIs (e.g., terrain characterization index, topographic openness) have been widely adopted in geomorphology and biology, but are seldom used in hydrological studies. A thorough examination of existing TIs is needed to identify those that best account for hydrological variations between watersheds.

Studies of watershed topography on hydrological processes often include topics such as specific discharge (Karlsen et al., 2016), spatial baseflow distribution (Shope, 2016), transit time (McGuire et al., 2005; McGuire and McDonnell, 2006), and hydrological connectivity (Jencso and McGlynn, 2011). These studies were often based on a short period of data (< 5 years), limiting our ability to draw general conclusions on how topography affects hydrological processes. Moreover, hydrological responses are compounded by the spatially diverse effects of climate, vegetation, soil, and topography (Li et al., 2017; Wei et al., 2018; Zhang et al., 2017). For example, several hydrological models have been applied to test the effects of spatial distribution of a hydrological variable (e.g., specific discharge, soil moisture, or groundwater recharge) (Erickson et al., 2005; Gómez-Plaza et al., 2001; Li et al., 2014). However, the effects of topography alone on hydrology are not usually addressed in those studies. Finally, understanding how topography influences hydrology has significant implications for sustainable

management of aquatic ecosystems (Zhang et al., 2016). Therefore, the major objectives of this study were: 1) to examine the role of topography in various flow magnitudes in 28 selected watersheds with snow-dominated hydrological regimes in the Southern Interior of British Columbia, Canada; and 2) to identify the most important topographic indices that can be used to compare variations in flow magnitudes between watersheds under similar climatic conditions.

2. Data

2.1. Study watersheds

In this study, 28 watersheds were selected ranging in size from 2.6 to 1,780 km^2 (Fig. 1 and Table S1 in the Supplementary Information (SI)). The watersheds are located between 51ºN 122ºW and 49ºN 118ºW in the Southern Interior of British Columbia, Canada where hydrological regimes are snow-dominated. In this region, the Pacific Decadal Oscillation shifted from a cool to a warm phase around 1977 (Fleming et al., 2007; Wei and Zhang, 2010), resulting in more precipitation and lower temperatures, and consequently affecting hydrological regimes. In addition, an extensive mountain pine beetle infestation caused large-scale forest cover change from 2003 onwards. To avoid the uncertainties associated with these perturbations and maximize the sample size, the period of 1989-1996 was selected, during which daily flow records of selected stations are complete. In addition, we further confirm that vegetation changes (using LAI as a proxy) did not significantly alter annual mean flow during this period (see the SI, section 2).

Annual mean temperature (T) and precipitation (P) of the study watersheds were calculated from the ClimateBC dataset (Wang et al., 2006). ClimateBC is a standalone program that extracts and downscales PRISM (Daly et al., 2008) monthly climate normal data and calculates seasonal and annual climate variables for specific locations based on latitude, longitude, and elevation. Annual P, T, and potential evapotranspiration (PET) were determined at a spatial resolution of 500 meters and averaged for each watershed. PET was calculated using the Hargreaves method (Liu et al., 2016; Wang et al., 2016; Zhang and Wei, 2014). The average mean annual P and PET of all 28 watersheds were 813 ± 205 mm and 586 ± 58 mm for 1989-1996, respectively (Fig. 1 and Figs. S1-S4).

2.2. Topographic indices

Based on availability and representation of TIs in literature, 22 topographic indices (TIs) were derived using a gridded DEM at a spatial resolution of 25 m (Table 1). The DEM, geospatial streamflow networks, lakes and wetland coverage were obtained from GeoBC (Government of British Columbia, http://www2.gov.bc.ca/gov/content/data/about-data-management/geobc/geobc-products). All data were transformed to the same projected coordinate system prior to the calculations of TIs. Calculation of the TIs was made in ArcGIS 10.4.1 (ERSI®) and SAGA GIS 2.1.2. Detailed information on the calculation and interpretation of TIs can be found in the references listed in Table 1.

Figure 1. Locations and elevations of the 28 study watersheds, and their hydrometric stations (red star sign).

Table 1. 22 topographic indices and descriptions.

3. Methods

3.1. Definitions of the selected flow variables

Annual mean flow (Q_{mean}) and other flow variables generated by the annual flow duration curve (FDC) were used in this study. The selected flow variables include: $Q_{10\%}$, $Q_{25\%}$, $Q_{50\%}$, $Q_{75\%}$, $Q_{90\%}$, and annual minimum flow (Q_{min}) (mm day^{-1}). They are defined as the daily flow that occurred each year at the given percentage. For example, $Q_{90\%}$ is the flow magnitude at 90% of the time in a year (Cheng et al., 2012). To account for the confounding effects of climate, each annual flow variable was standardized with annual (P) and expressed as Q_X/P.

3.2. Factor analysis

Because some initially selected TIs may be highly related, the first step was to introduce factor analysis (FA) to reduce the number of TIs while still retaining important topographic information. FA can be interpreted in a similar manner as principal component analysis (PCA). The major difference between the two approaches is that FA not only considers the total variance but also makes the distinction between common and unique variances (Lyon et al., 2012). As TIs were calculated in a region with

similar topography, the average TIs at the watershed level are not varied dramatically between watersheds (McGuire and McDonnell, 2006; Price, 2011). Therefore, to ensure better differentiation, the standard deviations of TIs of a watershed were used for the FA test. It should be noted that the flow variables were not included in the FA test.

Three criteria were used in the FA procedure to exclude redundant TIs: Kaiser-Meyer-Olkin (KMO), Bartlett's test, and anti-image correlation. KMO is a measure of sampling adequacy, which tests that partial correlations among variables are small enough to ensure the validity of the FA test. Bartlett's test of sphericity assesses the level of correlation between the variables in the FA to determine if the combination of variables is suitable for such analysis (Lyon et al., 2012). The diagonals of anti-image correlation matrix are a measure of sampling adequacy of specific TIs, which ensures that TIs are adequate for the FA. If a TI makes the FA indefinite, namely KMO < 0.7, Bartlett' test P > 0.05, and the diagonals of anti-image correlation < 0.7, then this TI is excluded from further consideration. With this iterative approach, the group of TIs with the largest KMO, Barlett's test P <0.05, and the diagonals of anti-image correlation > 0.7 are determined as the final group of TIs. In this study, FA tests were conducted in the IBM® SPSS® Statistics Version 22.

3.3. Relative contributions of each TI to flow variables

The nonparametric Kendall's tau correlation examined the statistical correlations between flow variables and the FA-selected TIs in the 28 study watersheds. If a significant correlation is detected, it indicates high topographic control on that flow variable. Multiple linear regression (MLR) models were then built for each year between 1989 and 1996 for each flow variable (see section 5 in the SI for details). The purposes of the MLR models were: 1) to further exclude those TIs that were insignificantly related to flow variables, and 2) to quantify the relative contributions of the selected TIs to each flow variable in each year. In the MLR models, each flow variable was treated as a dependent variable, while all FA selected TIs were regarded as independent variables. To exclude insignificant TIs to flow variables, all the selected 11 TIs by the FA were initially included in the MLR model. The ANOVA test was then adopted to identify the statistical significance between TIs and each flow variable in each year. If one TI was insignificant ($P>0.05$), then it was removed from the model. The ANOVA test was then re-run for the rest TIs to ensure that all significant TIs were selected. By this trial and error process, the final models with only significant TIs to flow variables were determined. The detailed procedure can also be found in Li et al. (2014). Then, the R package "relaimpo" was used to quantify the relative contributions of the selected TIs to each flow variable (Gromping, 2006). In particular, the relative contributions of each dependent variable to R^2 were calculated for each model. In this way, the contributions of the significant TIs to flow variables were derived for each MLR model.

To quantify the role of each TI in regulating each flow variable, we defined a contribution index (*CI*), which can be expressed as: $CI = O \times C$, where, *O* is the number of the selected TIs appearing in the final multiple linear regression models (the maximum number is eight because each flow variable is studied for eight years), and *C* is the average relative contribution of each TI to the specific flow variable . Therefore, a higher *CI* indicates a higher influence of that TI on a flow variable. The lumped *CI* of each TI to each flow variable is treated as the total contribution of the TI to all flow variables. Finally, the TIs with the CI values that are higher than the average of the lumped CI of all TIs were determined as the final set of TIs. As such, TIs with higher contributions were selected for the flow variables.

4. Results

4.1. Factor analysis

A subset of 11 TIs was selected from the initial 22 calculated TIs using the FA procedure. The KMO test (0.853), Barlett's test (P<0.001), and the diagonals of anti-image correlation (> 0.7) on the 11 TIs further confirm that our selected TIs are adequate to represent topographical characteristics of a watershed in our study region. In the FA analysis, the first and second factors explained 80.9% and 11.7% of total variance of TIs in the selected watersheds, respectively (Fig. 2). TIs included in the first factor are: DDG, LS, openness, relief, slope, SA, TCI, and TRI, while those in the second factor include SCA, perimeter, and UCA. Based on the definition of each TI, we further conclude that the first factor

represents watershed roughness or complexity, while the second factor describes the watershed size. Therefore, the selected 11 TIs were subsequently classified into two groups representing complexity (Group A) and area (Group B) (Fig. 2).

Figure 1. Factor analysis of topographic indices (TIs) among 28 watersheds. The first and second factors explained 80.9% and 11.7% of the total variance, respectively.

4.2. Relative contributions of TIs to flow variables

The nonparametric Kendall's tau test revealed significant correlations between the TIs and each flow variable from 1989 to 1996 (Tables S2 to S8 in SI). The number of significant TIs in each year increased from 1 to 11 with decreasing in flow magnitudes. A larger number of the TIs were correlated to low flow variables ($Q_{75\%}$, $Q_{90\%}$, and Q_{min}), indicating that topography plays a more pronounced role in regulating lower flows. Here, flows lower than or equal to $Q_{75\%}$ are defined as the low flows. Thus, MLR models were only carried out for $Q_{75\%}$, $Q_{90\%}$, and Q_{min}.

The regression models between flow variables and the selected TIs were all significant ($P<0.01$) with R^2 values ~ 0.5, indicating that selected TIs can be used to explain the variations of flow variables (Table S9). The details of MLR models are listed in the SI. The TIs that were included in each model are shown in Fig. 3 and Tables S10-S13. The relative contributions of each TI to $Q_{75\%}$, $Q_{90\%}$, and Q_{min}

showed large variations between years (Fig. 3 A to C, and Tables S10-S12). Fig. 3 revealed that each TI influences flow variables differently (e.g. SA played a prominent role in $Q_{90\%}$ but did not significantly contribute to the variation of $Q_{min,}$ while openness had the opposite role). This means that each TI cannot be used to explain variation in the same way for each flow variable. Based on the lumped CI values, the relative contribution of the perimeter TI was the highest in Group B as well as among all TIs (Fig. 3). In Group A, the TIs above the average of the lumped CI were LS, SA, openness, and TCI receiving high contributions to low flows. Therefore, we conclude that the above-mentioned five TIs are significant topographic indices influencing low flow variables, which can be used to assess and compare low flows in any watershed studies where there are similarities in watershed sizes and climate.

Figure 3. Relative contributions of each topographic index (TI) to $Q_{75\%}$ (Panel A), $Q_{90\%}$ (B), and Q_{min} (Panel C) from 1989 to 1996. Note that the numbers above the bars indicate the number of years when the given TIs were included in the regression models. Panel D: Contribution index (CI) of the 11 topographic indices (TIs) selected by factor analysis (FA) to $Q_{75\%}$, $Q_{90\%}$, and $Q_{min,}$ respectively. The dashed line is the average of the lumped CI.

5. Discussion

In this study, our results show that a limited number of TIs are significantly related to the Q_{mean}, $Q_{10\%}$, and $Q_{25\%}$, suggesting that topography plays a limited role in the variations of annual mean flow and high flows. This study area is characterized by snow-dominated hydrological regimes, with high flows

(e.g., $Q_{25\%}$ or greater) coming predominantly from the snow-melt process in early March to late May (e.g., Fig. S5). Snow-melting processes are significantly related to elevation and climate in our study region (Winkler et al., 2005). As the study watersheds in the Southern Interior of BC, Canada have similar elevation ranges and climate variability, it is not surprising that only a limited TIs were significantly related to high or mean flows. In contrast, more TIs were significantly associated with low flows, suggesting that topography plays a more important role in low flows than high flows in the study region. Low flows often occur in the later summer (late August) and winter (October to February) (e.g., Fig. S5), and are mainly driven by groundwater discharges and small amounts of precipitation. A watershed with more complexities of topography would likely have higher water retention ability due to longer flow paths and residence time and consequently promote more groundwater recharge and higher low flows (Price, 2011). Our Kendall's tau correlation tests uncover a positive relationship between the selected TIs and low flow variables (Tables S6-S8), indicating that the rougher or more complex that a watershed is, the higher the low flow yields are. Therefore, we conclude that topography plays a more important role in low flows than in high flows in the study region.

Five TIs including perimeter, LS, SA, openness, and TCI were identified as the major contributors to flow variables in this study. As far as we know, no studies have quantified topographic controls on various flow magnitudes. Nevertheless, the relationship between topography and the mean transit time (McGuire and McDonnell, 2006), temporal specific discharge (Karlsen et al., 2016), and hydrological

connectivity (Jencso and McGlynn, 2011) have been investigated. It is no doubt that topography is one of the major contributors to hydrological variations (Price, 2011; Smakhtin, 2001). Although these studies pinpointed specific TIs and their interactions with hydrological responses, only a limited number of TIs were quantitatively assessed. In contrast, a total number of 22 TIs were calculated for 28 watersheds in this study. The much higher number of TIs that were initially included, along with the filtering methods applied allowed us to select more suitable and significant TIs. Through this study design, we expect that the selected five TIs can effectively be used to support assessment or comparisons of low flows between watersheds in the study region. It should also be noted that we only selected the first five TIs that had substantially higher contributions than the other calculated TIs. The rest of the hydrological-related TIs showed only a minor ability to explain flow variations.

Among the selected five TIs, perimeter, a primary TI, is commonly used in scientific studies to describe the characteristics of watershed topography. Our study further proves that it has a large influence on low flow variables. However, four secondary TIs (LS, SA, openness, and TCI) are mainly used in geomorphology to characterize ruggedness or roughness of landscapes and to identify topographic functioning of ecosystems. For examples, TCI has been used to map soil organic matter concentration (Zeng et al., 2016). Park et al. (2001) revealed that TCI is a better TI to predict soil depth than TWI, plan curvature, and profile curvature (see definitions in Table 1). LS is one of the key inputs to the universal soil loss equation (USLE) being used to quantify soil erosion hazards (Desmet and Govers,

1996). SA was used to estimate animal species and habitat (Jenness, 2004), and map the spatial patterns of a floodplain (Scown et al., 2015). Openness was initially adopted to identify the boundary of different geological units and can be used to identify surface convexities and concavities, which is better than the commonly used profile and plan curvature (Yokoyama et al., 2002). In this study, the selected five TIs were initially filtered by the FA test, indicating that each selected TI has uniqueness in describing watershed topographic characteristics, and is more outperformed than the other tested TIs in describing variation in flow variables in our study region. Therefore, we expect that selected five TIs can be applied to support hydrological analysis and modelling.

To our surprise, some commonly-used TIs in hydrology, such as slope, median elevation, upslope contribution area, wetland areas, etc. are not included in the FA list as the topographic information contained in those primary TIs also exists in some secondary TIs. Secondary TIs have the advantage of describing the hydrology-related landscapes in fuller detail. For example, slope is directly included in calculations of the TWI (secondary) and DDG (secondary). UCA (primary TI) is included in TWI and TCI (secondary). It is also worth mentioning that some secondary TIs played critical roles in determining the spatial heterogeneity of ecological and geomorphological processes, but their roles in hydrological processes in our study region were not demonstrated. These TIs were, therefore, not selected in this study. For example, the TWI has long been used as a key input variable for TOPMODEL, and is an indicator of soil moisture (Beven, 1995). Our study identified that TWI and

wetland area were not significantly related to flow variables, indicating that these factors played a limited role in the selected flow variables in our region. This may be because the watersheds in this area need to overcome a soil moisture storage threshold prior to releasing water (Karlsen et al., 2016). In summary, our selected five TIs significantly represent low flow characteristics of the watersheds in the Southern Interior of British Columbia, Canada, which is characterized by a snow-dominated hydrological regime.

In our study, FA was first introduced to exclude reductant TIs. The MLR models were further employed to exclude the insignificant TIs to various flow magnitudes. There is a concern whether the TIs excluded by FA were significantly related to flow magnitudes, but not included in the MLR models. To address this concern, we re-ran a sub-set of analyses by conducting the MLR models for Q_{min} in 1989, 1990, and 1991 as examples (section 6 in the SI). The results showed that the MLR models would include several redundant TIs (e.g., wetland coverage, roundness, stream length) if the FA test was not conducted (Table S14). Kendall's tau correlation tests uncovered that they were not significantly correlated with low flows (Table S8). This further demonstrates that some redundant and nonsignificant TIs would be introduced if the FA test was not initially conducted. In addition, our research framework was developed based on a large number of watersheds with long-term hydrological data all in a similar climatic region that is snow-dominated. As far as we know, there are no such studies conducted in either snow- or rain-dominated regions. Our research methodology or

framework can be applied in rain-dominated regions to assess how topography controls flow in any regions where there are sufficient long-term monitoring data.

There are several uncertainties in our study. Firstly, hydrological responses are the resultant effects of climate, soil, vegetation, topography, and geology (Price, 2011; Smakhtin, 2001). In this study, LAI represented the variations of vegetation cover in different watersheds was included in our analysis in order to minimize the effects of different levels of vegetation coverage in the study watersheds. However, it was excluded by the FA test, confirming that the differences in vegetation cover and their effects were minor so that our selected watersheds are comparable in terms of forest coverage. In addition, the climate could be a confounding factor affecting our comparisons. In this study, annual flow variables were standardized by the annual precipitation to minimize the effects of climate on flow variables. In this way, the effects of climate variability were considered to some extent but not in full detail. For examples, in extreme dry or wet years (e.g., 1994 and 1996 are the driest and wettest years in our study period, respectively) (Fig. S2-S5), hardly any of the TIs were significantly correlated to $Q_{90\%}$ (Table S7), but almost all were correlated to Q_{min} (Table S8). This is because that the flows at $Q_{90\%}$ in 1994 and 1996 in the selected watersheds occurred in the winter (October-February) (e.g., Fig. S5), while Q_{min} in the majority of watersheds were in the late summer (August and September). The Kendall's tau correlation tests (Table S7) also confirm that topography played a minor role in $Q_{90\%}$, but a significant role in Q_{min} (Table S8). Thus, hydrological responses are mainly controlled by the

combined effects of climate and topography. Secondly, LS and TCI are commonly-used indicators of soil erodibility, and were included in this study to capture the influence of soil conditions. For example, Park et al. (2001) showed that the areas with high values of TCI have high rates of soil erosion. Therefore, the variation of soil properties was considered to some degrees by the inclusion of these TIs, but not in full detail. In fact, it is impossible to derive accurate soil properties over a large area. Thirdly, potential impacts of the DEM resolution on the calculation of TIs were not considered. However, Panagos et al. (2015) indicated that the resolution of 25 meters is adequate for calculation of slope factor at the European scale. Similarly, DEM resolutions were not found to significantly affect the calculation of DDG (Hjerdt et al., 2004). Thus, considering the same DEM data with the resolution of 25-meter, we assume that the error caused by the DEM resolution would also be minor.

6. Conclusion

This study concludes that topography plays a significant role in low flows, while its role in high flows is limited. A total number of five topographic indices, including perimeter, LS (slope length factor), SA (surface area), openness, and TCI (topographic characteristic index), were identified with significant contributions to low flow variables. It is recommended that those above-mentioned five TIs can be used to assess the magnitude of low flows in a study region which is characterized by a snow-dominated hydrological regime with watershed sizes up to several thousand square kilometres. Our

research methodology can be applied to other regions for investigating how topography controls flow magnitudes.

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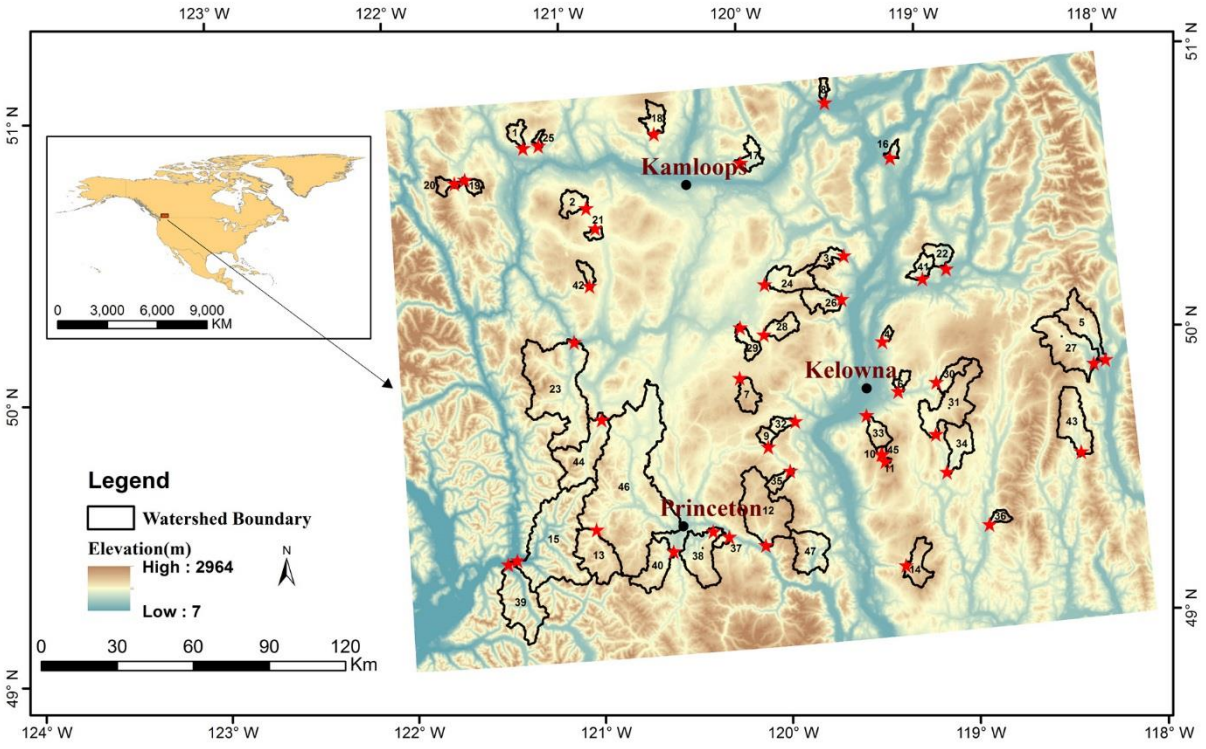
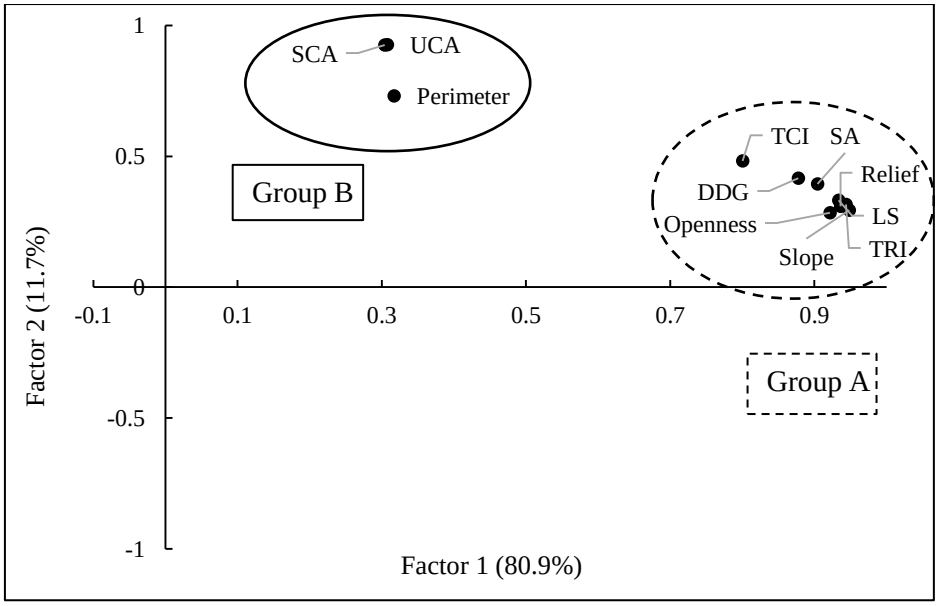


Figure 1. Locations and elevations of the 28 study watersheds, and their hydrometric stations (red star sign).



490 **Figure 2.** Factor analysis of topographic indices (TIs) among 28 watersheds. The first and second factors explained 80.9% and 11.7% of the total variance, respectively.

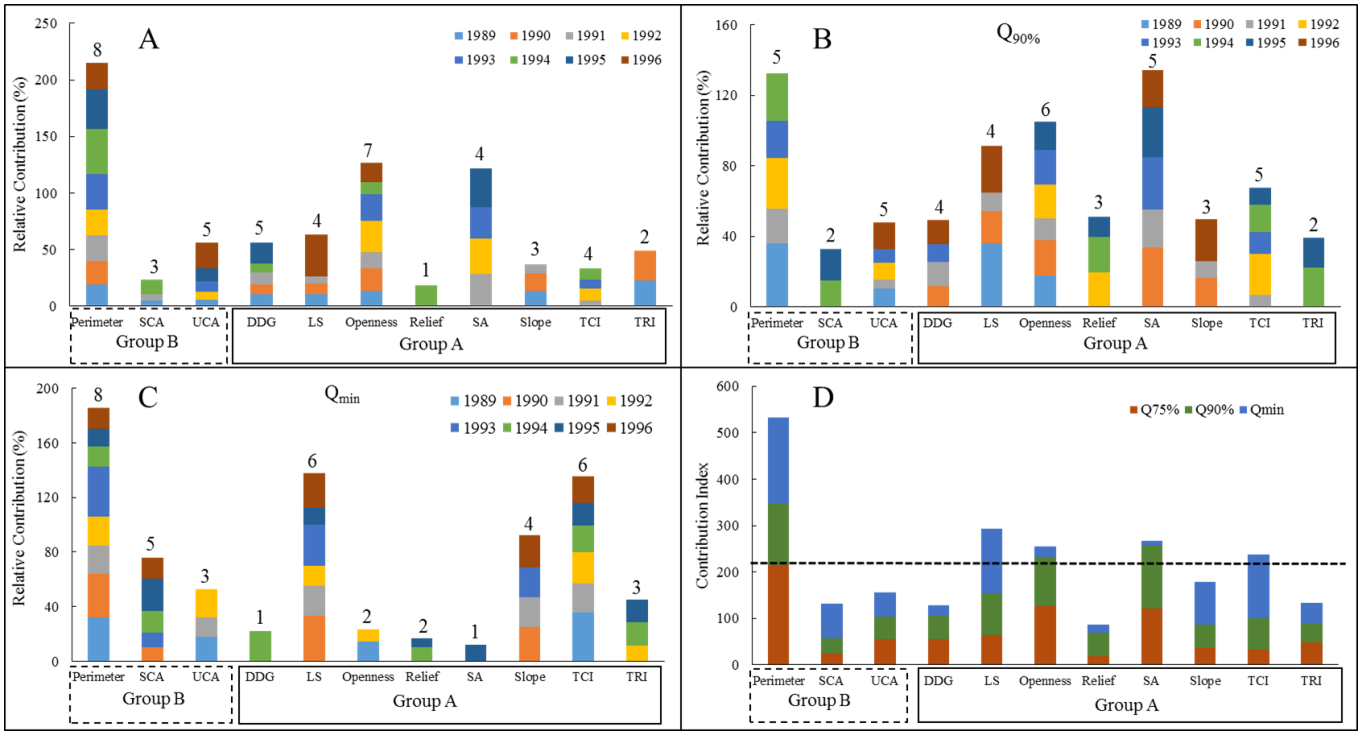


Figure 3. Relative contributions of each topographic index (TI) to $Q_{75\%}$ (Panel A), $Q_{90\%}$ (B), and Q_{min} (Panel C) from 1989 to 1996. Note that the numbers above the bars indicate the number of years when the given TIs were included in the regression models. Panel D: Contribution index (CI) of the 11 topographic indices (TIs) selected by factor analysis (FA) to $Q_{75\%}$, $Q_{90\%}$, and Q_{min}, respectively. The dashed line is the average of the lumped CI.

Table 1. Topographic indices and descriptions

No.	Abbreviation	TI	Description and references
1	**UCA**	Upslope Contributing Area	UCA is the area that can potentially produce runoff to a given location (Erskine et al., 2006).
2	**DDG**	Downslope Distance Gradient	DDG is a hydrologic measure of the impact of the local slope characteristics on a hydraulic gradient. Values are low on concave slope profiles and high on convex slope profiles (Hjerdt et al., 2004).
3	DDGD	Downslope Distance Gradient Difference	The difference between DDG and local or neighbour gradients (Hjerdt et al., 2004).
4	FLD	Downstream Flow Length	The downslope distance of a pixel along the flow path to the outlet of a watershed (Greenlee, 1987).
5	ME	Median Elevation	Median elevation among all DEM pixels in a watershed.
6	**Relief**	Relief	The difference between the highest and lowest elevations within a local analysis window. 11×11 grid cell window is used in this paper.
7	Roughness	Roughness	Roughness is calculated as 1/cos(slope) of each DEM pixel.
8	**Slope**	Slope Degree	Slope degree of each DEM pixel (Burrough et al., 2015).
9	**LS**	Slope Length Factor	LS is a combined factor of slope length and slope gradient. It represents the ratio of soil loss per unit area on a site to the corresponding loss from a 22.1 m long experimental plot with a 9% slope (Desmet and Govers, 1996).
10	**SCA, also known as A_s**	Specific Contributing Area	Upslope contributing area per unit length of contour (Quinn et al., 1991).
11	STD	Stream Density	Ratio of the sum of all stream length to watershed area.
12	**TCI**	Terrain Characterization Index	TCI = Cs *$\log_{10}$ (SCA), where Cs is the surface curvature index; The higher positive TCI values reflect higher aggradation of soil materials at a certain point along the hillslope (Park and van de Giesen, 2004).
13	**TRI**	Terrain Ruggedness Index	TRI expresses the degrees of difference in elevation among adjacent cells (Riley, 1999). It calculates the sum changes between a grid cell and its eight neighbour grid cells. Higher values indicate more ruggedness of a watershed.
14	TPI	Topographic Position Index	TPI ≈ 0 indicates flat area. TPI>0 tends towards ridge tops and hilltops. TPI< 0 tends towards the valley and canyon bottoms (Jenness, 2006). A 9×9 grid cell window is used in this paper.
15	TWI	Topographic Wetness Index	TWI = ln (SCA/Tan(slope)), it shows the spatial distribution of zones of surface saturation and soil water content (Ambroise et al., 1996; Quinn et al., 1995).

16	Wetland	Wetland Coverage	Percentage wetland area to total watershed area.
17	Length	Length of Main River	The total length of mainstream.
18	Roundness	Roundness Coefficient	The ratio of watershed area to the area of a circle with the same perimeter. A lower value indicates a longer and narrow watershed.
19	**Openness**	Positive Topographic Openness	Describe the degree of dominance or enclosure of a location on an irregular surface. Values are high for convex and low for concave forms, respectively (Yokoyama et al., 2002).
20	**SA**	Surface Area	The land area of each DEM, which may provide a better estimation of the surface roughness than the planimetric area (Jenness, 2004). A lower value indicates a more gentle topography.
21	**Perimeter**	Perimeter of a Watershed	The perimeter of a watershed.
22	Total	Total Curvature	The standard curvature combines profile and planform curvatures (Moore et al., 1991).

505 Note: Bold and underlined TIs are selected by the Factor Analysis test.