Dear Editor,

Thank you very much for your constructive comments. We took them into consideration when revising the manuscript. Regarding your comments 'the clarifications which are needed to make it easier to understand exactly what you have done', we added more information in the section of method (2.4, 2.5, and 2.6), which will better explain the purpose and research procedures of our work. Additionally, we have revised discussion section and added more explanation to clarify the mechanism of hydrological effect of land use change. As to the English word selection and expression, we got help from a native speaker to check the manuscript, and all writing error have been corrected. We believe that all of these revisions improved the quality substantially. Thanks again for your time.

Responses to reviewers

We thank two reviewers very much for their insightful comments and suggestions. All comments are valuable and helpful for improving the quality of the manuscript and have been carefully checked and addressed. Please see our point-by-point responses below. A kindly note: reviewers' comments are in blue color and our responses are in black color; page and line numbers cited in our responses refer to the revised manuscript with track changes for easy evaluation. A track-changed version of the revised manuscript can be found after the response to reviewers.

Reviewer #1:

Qiu et al.- Spatiotemporal response of the water cycle to land use conversions in a typical hilly-gully basin on the Loess Plateau, China- studies the hydrological response of land cover and land use change in a hilly-gully basin of the Loess Plateau in China. This region is very lack of water resources and land use and management play very important role in increasing its resilience and sustainability. This study based on modeling approach adds new and important information for effective water resource conservation and management. Meanwhile it would serve as a useful reference with broad impact for similar arid region worldwide. However, the present form is lack of some details on model construction and validation, especially for the dry period simulation. I would also suggest the author check and improve the input data accuracy. English word selection and expression accuracy need improvement too. There are too many usage of "exhibit" in the text. The following specific comments provide more suggestion for author to consider in their manuscript revision. **Response:** We thank for the positive evaluation and instructive comments and suggestions.

(1) According to the reviewer's suggestion, we added more details on model construction, and this procedure is as follows: A geographic information system (GIS) interface, ArcSWAT (Version 2012.10_2.18), was used to set up the model. The DEM was used to delineate the watershed, and this process produced 88 subbasins. The 30 m land use map of 1990 and soil map were from the Chinese Academy of Sciences, and the digitalized soil data were from the Data Center of West China (http://westdc.westgis.ac.cn/). We used the multiple Hydrological Response Unit (HRU) option in ArcSWAT to represent each unique field (combination of land cover, soil, and slope) as a separate unit, leading to a total of 1,136

HRUs using a threshold value of 10% for the three categories. To meet the needs of the model, the SWAT codes for the main land use types in the study area were defined as cropland (AGRR), woodland (FRST), grassland (RNGE), residential land (URBN), and barren land (BARR) in the attribute data (*Please see P6 L22 ~ P7 L7*).

(2) In terms of model verification, we firstly calibrated the SWAT model with 10-year (1983–1992) monthly streamflow data of Ganguvi hydrological station using SWAT-CUP (SUFI2 algorithm). Then the model was validated using data from the subsequent eight years (1993–2000). To evaluate the model performance numerically, we used three statistical measures, including percent bias (PBIAS), Nash-Sutcliffe efficiency (NSE), and the coefficient of determination (squared correlation coefficient, r^2). Typically, a model simulation is considered satisfactory with NSE>0.5, -25% \leq PBIAS \leq 25%, and r² \geq 0.5 (Wu and Chen, 2013; Neupane and Kumar, 2015) (*Please see P7 L7 ~ P7 L17*). In the calibration and validation periods, the statistics of NSE, PBIAS and r² ranged from 0.51~0.70, 15.7%~16.9% and 0.71~0.82, respectively, indicating a satisfactory performance. However, PBIAS value showed that SWAT underestimated the streamflow in both calibration and validation periods. Visual inspection as shown in Fig. 3 indicated that the underestimation mainly occurred in the dry period. The potential reasons for the underestimation can be categorized as "measurement quality" and "model behavior" reasons, "Measurement quality" refers to missing precipitation data or streamflow measuring error. The study area lies in the hilly-gullied loess region, and the complex and rugged terrain results in highly variable precipitation and difficulties in spatially estimating the precipitation (Cao et al., 2006). Additionally, this region has a limited number of meteorological and hydrological stations. In this study, five meteorological stations were involved; however, only three stations are located inside the basin, with the remaining two occurring outside the basin (Fig. 1). It has been reported that the accuracy of streamflow prediction mainly depends upon the precipitation gage numbers and their locations (Cao et al., 2006; Mul et al., 2009). Thus, the insufficient precipitation records and the distances among the meteorological stations resulted in poor streamflow simulation. "Model behavior" refers to model limitations such as the inadequate representation of the physical mechanisms of hydrological processes. For example, SWAT uses total daily precipitation and does not consider rainfall intensity within a day; thus, it can underestimate streamflow for some precipitation events (Qiu et al., 2012). Another example is the use of runoff curve number to simulate the surface runoff behavior. This approach does not account for saturation excess runoff or contributions from variable source areas (Garen and Moore, 2005; Easton et al., 2008). Although the simulation results in our research met the calibration criteria, the model performance can be improved if precipitation data from additional gages become available in the basin. Above information was added in the paragraph one of the discussion section (Please see P13 L10 ~ P14 L5).

(3) We thank the reviewer for pointing out our language problems (e.g., too many use of a word). Now we have checked the manuscript carefully and revised it further to improve its quality and accuracy.

Line 22-23, pp1: "These results suggest that the expansive revegetation of sloping land could reduce runoff generation, particularly in woodland areas, but these effects could reduce the soil water volume in the region." What are the percentages of runoff and soil water volume reductions?

Response: Following the reviewer's suggestion, we calculated the decrease percentage. Compared with the land use

condition in 2010, the negative effects were most evident where cropland with a slope $\geq 15^{\circ}$ was converted to woodland, with decreases in surface runoff and soil water of 17.1% and 6.4%, respectively. We added this information in the abstract (*please see P1 L22 ~ P1 L24*).

Line 6, pp2: change to "considered as?" **Response:** Done (*please see P2 L19*).

Line 23, pp2: "between runoff and increased LULC" what kind of land cover and land use is that?

Response: In the literature, it was reported that mean runoff in 2005–2010 declined significantly compared with the time period of 1997–2005, and this can be attributed mainly to changes in land use/land cover, i.e., increases in forests, shrubs, and grasses, and decreases in sloping farmlands. Thus, we rephrased the sentence to make the description more accurate (*please see P3 L13-15*).

Line 27-29, pp3: need valid data sources for the climate information.

Response: The climate information of study area was based on the collected observation data, and the data was provided by the Meteorological Institute of Shaanxi Province, China. We added the data source in the manuscript accordingly (*please see P5 L4*).

Line 4, p4: figure 2 shows that the streamflow is about 80% of precipitation. This is against common sense. Pls check your data again.

Response: The units of the streamflow (10^6 m^3) and precipitation (mm) are different, and the calculated ratio of streamflow to precipitation is around 6.6%.

Line 10-20, pp5: lulc of 2010 was used as baseline. How about the lulc of 1990 and 2000? Did you use these two maps as model calibration and validation? It seems only 2010 map was used for the slope related scenario creation. It is a little bit confusing here. You may need to clarify more on the model and scenario configurations.

Response: We have added more details on model set up and scenario configurations based on the reviewer's comments. We firstly calibrated the SWAT model using land use data of 1990. Then we test different land use data of 1990, 2000 and 2010 under the same climate data from 1986-2015 to separate the unique effect of land use change. Finally, the hypothetical land use conversion scenarios were developed from land use condition of 2010 and its hydrological effects were assessed using same procedures with evaluating historic land use change effects (*please see P6 L21 ~ P8 L18*).

Line 29-30, pp5: do you mean the precipitation input miss the streamflow simulations? If so, you may need to prove/show that in a new figure.

Response: Indeed, here we tried to explain the reason for the discrepancy between measured precipitation and streamflow during the dry period. We deleted this sentence and gave the explanations in the section of Discussion (*please see P13 L11 ~ P14 L5*).

Table 3: change "Ratio" to "Percentage" for consistence.**Response:** Done (*please see P26 Table 3*).

Line 23, pp6: it should be "in the same period" and "resulting the small net change" **Response:** Done (*please see P10 L9*).

Line 28, pp8: change "classes" to "types" **Response:** Done (*please see P14 L6*).

Line 11, pp9: what the difference between surface runoff and streamflow?

Response: In SWAT, surface runoff is flow that occurs along the sloping surface (i.e., quick-response overland flow), which contributed to streamflow. It occurs whenever the net rainfall exceeds the rate of infiltration or the it was saturated. Streamflow is the water flowing downhill through creeks, streams, and rivers toward the outlet of the basin, and it includes surface runoff, lateral flow within soil profile, and base flow.

Line 23, pp9: why cropland has the highest soil water storage?

Response: The reason can be explained as follows: Firstly, in the Yanhe basin, the cropland was mainly distributed in the bottom region of the valley and on land with gentle slope. The valley bottom generally receives more water flow from the hillslope and has lower temperature than higher regions, resulting in more soil moisture (Wang et al., 2012). Furthermore, land with gentle slope has higher soil water retention ability than does land with steep slope (Famiglietti et al., 1998). Secondly, our results showed that the cropland had lower evapotranspiration per unit area than did the other land use types. Thus, these interrelated above two reasons led to the highest soil water storage. This explanation has been added to the revised manuscript (*please see P14 L9 ~ P14 L15*).

Line 4, pp10: sloped cropland or sloping cropland? Pls make this term consistent throughout the manuscript.

Response: Thanks for the suggestion, and we used "sloping cropland" in the whole manuscript (please see P16 L7-8).

Line 13, pp10: "transition analysis"? Pretty new and never heard. Pls consider changing it. Otherwise more definition should be given.

Response: "Land us transition analysis" summarizes the amount of certain land that changes from each category at the initial time to each category at the subsequent time, and it has been widely used in landscape ecology and GIS studies of land use to quantitatively estimate the rate of change. To avoid confusion, we took "land use change analysis" in the revision (*please see P9 L20 & P26 L1*).

Line 24, pp10: change planting wood to afforestation? **Response:** Done (*please see P17 L3*).

The study (Qiao et al., 2017, Woody plant encroachment reduce annual runoff and shifts runoff mechanisms in the tallgrass prairie, USA. Water Resource Research) provides some evidence of water budget difference between woodland and grassland and would be a useful reference for this study.

Response: Thanks for recommending a valuable reference to our study, and we have cited this article in our revision (*please see P14 L21-24*).

Reviewer #2:

General Comment: This manuscript investigated the response of main hydrological components to land use change and basin management policy in the loess region of China. It is interesting to see the spatial distribution of hydrological responses at the unit (HRU) level and the quantitative assessment of the roles of different land use types in basin hydrology. In particular, it is attractive to interpret the mechanism of the streamflow decrease associated with vegetation restoration on the Loess Plateau and thus may have a broad readership. Overall, the article was well organized, well written, and easy to follow. All figures and tables are of high quality and informative. What I would suggest is give more discussion on the reason why the different land use types show different hydrological characteristics (e.g., soil water and ET) in this region. That would help reader better understand the findings from the article.

Response: We thank the reviewer for his/her positive remarks and constructive comments. We considered the comments of reviewer and added more discussion on hydrological characteristics. For soil water, explanations for differences among different land use types need to combine the topography and land cover on the land surface. In our study, woodland had lowest surface runoff and water yield, because woodland areas capture more rainfall and uptake more water than other land use types (e.g., cropland and grassland) (Wang et al., 2012; Jian et al., 2015), resulting in a lower infiltration, runoff, and water discharge (Wang et al., 2013; Duan et al., 2016). Qiao et al. (2017) reported that the reduced runoff in areas with woody plant relative to grassland areas is associated with shift in runoff generation mechanisms: thus, the shift from saturation excess overland flow to infiltration excess overland flow might also have contributed to the reduced runoff in woodland areas observed in the present study. Additionally, compared with other vegetation types, woodland areas might lose more water through ET. This pattern was evident from the change in soil water, with less soil water in woodland than in other areas under the same precipitation amount. Less soil water was observed in woodland because forests in the Yanhe basin generally grow on landform with high slopes; our analysis indicated that more than 62% of woodland was located on slopes $\geq 15^{\circ}$ (Table 5). Steeper slopes generally retain less soil water due to low infiltration and rapid surface drainage (Famiglietti et al., 1998). Thus, a large amount of precipitation was associated with forest growth and ET rather than discharge out of the basin as surface runoff and streamflow. (Please see P14 L16 $\sim P15 L7$). Cropland had the highest soil water which can be partly attributed to the topography of cropland. The cropland area was mainly distributed in the bottom region of the valley and on land with gentle slope. The valley bottom generally receives more water flow from the hillslopes and has lower temperature than higher regions, resulting in more soil water (Wang et al., 2012). Furthermore, land with gentle slope has higher soil water retention ability than does land with steep slope (Panagopoulos et al., 2011). In addition, cropland had lower evapotranspiration per unit area than did the other land use types. In contrast, grassland had lower soil water, which can be attributed to its higher soil water loss through evapotranspiration. Wang et al. (2012) reported that grass cover types cannot protect the soil surface from solar radiation, leading to greater water loss via direct evaporation (*please see P14 L9 ~ P14 L16*). As for ET on the Loess Plateau region, it has been widely demonstrated that the woodland had highest ET, followed by grassland, and cropland (Xiao et al., 2013; Wang et al., 2012).

Specific comments:

P2 L6: Use "sloping cropland" here to guarantee the consistence, because you use "sloping land" in the whole manuscript.

Response: Corrected (please see P16 L7-8).

P7 L26 Add "of" between "trend" and "surface runoff".

Response: We rehearsed this sentence, and revised version is 'the decrease in surface runoff' (please see P11 L22).

P7 L29 Replace "because of the decreased area of cropland" by "because of its decreased area". **Response:** Corrected (*please see P12 L2*).

P8 L1 Replace "deceasing" by "decrease". **Response:** Corrected (*please see P12 L5*).

P9 L1-12 the response of soil water in woodland are well discussed, how about in cropland and grassland and try to explain more.

Response: As responded to the reviewer #1, highest soil water in cropland may be due to two related reasons: Firstly, in the Yanhe basin, the cropland area was mainly distributed in the bottom region of the valley and on the land with gentle slope, and valley bottom generally receives more water flow from the hillslope and has lower temperature than higher regions, resulting in more soil moisture (Wang et al., 2012); gentle sloping area had higher soil water retention ability compared with steep sloping area (Panagopoulos et al., 2011). Secondly, our results showed that the cropland had lower evapotranspiration per unit area than did the other land use types. Thus, these interrelated two main reasons led to the highest soil water storage in cropland. In contrast, grassland had lower soil water, which can be attributed to its higher soil water loss through evapotranspiration. Wang et al. (2012) reported that grass cover types cannot protect the soil surface from solar radiation, leading to greater water loss via direct evaporation (*please see P14 L9 ~ P14 L16*).

P9 L14 Rephrase the sentence and clearly state what are the main forms of land use change in the Yanhe basin.

Response: It was rephrased, and the revised sentence is "After the implementation of the GFGP, the area of cropland decreased continuously because it was transformed into grassland and woodland from 1990 to 2010, and conversions among cropland, grassland and woodland were the main forms of land use change in the Yanhe basin" (*please see P15 L8-10*).

P9 L20 Where are the locations of 26 watersheds?

Response: This study reported in the literature was conducted globally. The researchers selected 26 catchments around the world and evaluated the effect of afforestation on water yield. To avoid confusion, we added "globally" in this sentence (*please see P15 L15-16*).

P18: "Ratio" should be changed to "Percentage" in Table 3. **Response:** Corrected (*please see P26 Table3*).

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Spatiotemporal response of the water cycle to land use conversions in a typical hilly-gully basin on the Loess Plateau, China

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10 Abstract. The hydrological effects of the 'Grain for Green' project (GFGP) on the Loess Plateau have been largely extensively debated due to the complexity of the water system and its multiple driving factors. The aim of this study was to investigate the response of the hydrological cycle to the GFGP measures based on in a case study of the Yanhe basin, a typical hilly-gully area on the Loess Plateau of China. First, we analyzed the land use and land cover (LULC) changes from 1990 to 2010. Then, we evaluated the effects of LULC changes and sloping land conversion on the main hydrological components in the basin 15 considering the land surface characteristics and climate impacts. The using the Soil and Water Assessment Tool (SWAT) was used for this analysis. The results indicated that farmland cropland exhibited a decreasing trend declining from 40.2% of the basin area in 1990 to 17.6% in 2010, and that the woodland and grassland areas correspondingly increased due to the implementation of the GFGP in the basin. Due to With the land use changes from 1990 to 2010, surface runoff and the water yield exhibited showed a decreasing trends which was mainly due to decrease in surface runoff, whereas In contrast, 20 evapotranspiration (ET) showed an increased increasing trend over the same period, resulting in a persistent decrease in soil water, Additionally, c The conversion of onverting sloping cropland areas with slopes $>15^{\circ}$ or $>25^{\circ}$ to grassland and or woodland hadexerted negative effects on surface runoff, the water yield and soil water and a positive effect on ET. Compared with the land use condition in 2010, the negative effects were most evident where cropland with a slope $\geq 15^{\circ}$ was converted to woodland, with decrease in surface runoff and soil water of 17.1% and 6.4%, respectively. converting cropland areas with

slopes $\geq 15^{\circ}$ or $>25^{\circ}$ to grassland and woodland had negative effects on surface runoff, the water yield and soil water and a positive effect on ET. The magnitudes of the hydrological effects generated by sloping cropland to woodland conversion were greater than those for sloping cropland to grassland conversion. These results suggest that the expansive revegetation reforestation of on sloping land in the loess hilly-gully region could reduce decreased runoff generationwater yield and increased ET, particularly in woodland areas, but these effects could reduce resulting in reduced the soil water volume in the region. Overall, th The results of this study can be used to improve support sustainable land use planning and water resource management on the Loess Plateau in China.

1 Introduction

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Land surface change is one of the most important drivers of eco-hydrological changes (Li et al., 2009; Bloeschl et al., 2007).
The impacts of land use and land cover (LULC) changes on water resources and hydrological processes in a river basin are mainly reflected in the overland surface runoff, streamflow, evapotranspiration (ET), and soil moisture (Bari and Smettem, 2004; Chawla and Mujumdar, 2015; Liu et al., 2008b; Zucco et al., 2014). Many studies have investigated the interactive mechanismsrelationships between land use patterns and basin hydrology, and they have found that the characteristics of basin hydrology vary among different land use patterns _______ due not only to different variation in land use types (Wang et al., 2012; Jian et al., 2015; Duan et al., 2016), but also to the spatial heterogeneity of LULC (Chu et al., 2010; Liu et al., 2013). However, debate exists among ecohydrologists regarding the effects of past and ongoing land use changes because of the spatial and temporal complexity of hydrological processes (Lørup et al., 1998; Lopez-Moreno et al., 2011; Alkama et al., 2013; Liu et al., 2016).

Land use planning in China is considered <u>as a crucial strategy</u> for the sustainable management of a river basin system<u>s</u> and has been widely adopted for ecological restoration and water resource protection, especially in the Loess Plateau region, which is <u>famous-well known</u> for its fragile ecology (Liu et al., 2008a; Zhang et al., 2009; Zhen et al., 2014). The river basins on the Loess Plateau are important because of the dense population, intensive cultivation, and the high demand for water in the area. Unfortunately, this region is characterized by insufficient water resources and severe soil erosion, and it has historically experienced vegetation degradation and desertification (Zhao et al., 2013; Guo et al., 2002). Thus, land use change and basin hydrology have attracted a considerable deal of attention, and <u>since the 1970s</u>, soil and water conservation practices have been implemented in the area to <u>overcome-mitigate</u> the <u>ever</u>-increasing environmental problems <u>since the 1970s</u>. In 1999, the 'Grain for Green' project (GFGP) was launched by the Chinese government in the Loess Plateau region, and with the primary goal was to of retiringe and converting steep croplands (slope $\geq 15^{\circ}$) to green lands (Zhou et al., 2012; Liu et al., 2008a). It was reported that the vegetation coverage on the Loess Plateau increased from 6.5% in the 1970s to 51% in 2010 (Wang et al., 2012), and approximately 16,000 km² of rain-fed cropland was converted to planted vegetation during the past decade (Feng et al., 2016). Consequently, the hydrological processes on the sloping land and in the river systems have changed, but the extent of these changes and the<u>ir relationships-ties</u>-to LULC change remain topies of scientific research<u>unclear</u>.

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In the past decades, the hydrological effects of land use change have been widely explored across different spatiotemporal 10 scales on the Loess Plateau. Huang et al. (2003) investigated the runoff response to afforestation using a paired watershed analysis and revealed that forest revegetation reduced annual runoff, with and the reduction amount increased increasing with the tree age of trees. McVicar et al. (2007, 2010) developed a decision support tool for China's revegetation program and simulated annual streamflow based on revegetation planning. Wei et al. (2015) observed-found a strong inverse relationship between runoff and conversion to forests, shrubs, and grasses during the period 2005-2010 in a typical watershed on the Loess 15 Plateau. strong inverse relationship between runoff and increased LULC using statistical analysis from 1997 through 2000. Feng et al. (2016) reported that revegetation increased ET and resulted in a significant (P<0.001) decrease in the ratio of streamflow to precipitation on the Chinese Loess Plateau. Zuo et al. (2016) combined statistical tests and hydrological modeling to assess the effects of land use on runoff and found that the water resources in the upstream region decreased more than did those in the downstream region. Liang et al. (2015) used an elasticity and decomposition model based on the Budyko 20 framework to simulate and forecast the hydrological effects of ecological restoration and demonstrated that ecological restoration played a dominant role in the reduction of streamflow in 14 main subbasins on the Chinese Loess Plateau. Another study showed that the GFGP can potentially increase the soil water content and water yield and decrease the runoff and ET (Tian et al., 2016). Moreover, some researchers stated have reported that woody species consume more water by evapotranspiration ET than do other vegetation types (Wang et al., 2012; Yang et al., 2014), and some studies have documented that large-scale reforestation has greatly-decreased the water yield and exacerbated water scarcity (Sun et al., 2006), gradually 25

leading to soil desiccation (Chen et al., 2008; Chen et al., 2007; Wang et al., 2008). In general, these studies indicated that the GFGP on the Loess Plateau has had an evident influence on basin hydrology; however, because these studies were performed from different perspectives or concentrated on single hydrological elements, the effects of ecosystem restoration on the water balance have not been clarified. In addition, most of theprevious studies have been based on statistical methods with and short time scales. Although numerical models are useful tools for quantitatively assessing the hydrological responses to environmental changes, the existing previous modeling studies mainly focused on the water discharge in a river channel, and with few studies focused on the have analyzed the spatial features of hydrological responses. Spatially studying such responses could improve watershed management and the development of strategies for water resource optimization. Therefore, a comprehensive understanding of how LULC change and <u>its</u> spatial heterogeneity in LULC affect the water balance is essential for long-term land use planning and water resource management.

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In this study, the Soil and Water Assessment tool-Tool (SWAT) was applied to investigate the hydrological impacts of land use changes, including the those resulting from sloping land conversion (SLC) program, on the Loess Plateau. The Yanhe basin, a main tributary of the middle reach of the Yellow River that has undergone large-scale revegetation and SLC, was selected to for analyzinge the hydrological response mechanisms using the via modeling technique. The specific objectives were as follows: 1) to investigate the spatiotemporal variations in key water balance components as a result of LULC change, 2) to evaluate the potential effects of land conversion on water availability under the SLC program, and 3) to examine the changes in soil water storage under different land use condition. The results provide a useful reference for sustainable land use planning and water resources management on the Loess Plateau in China.

2 Materials and methods

20 **2.1 Study area**

The Yanhe basin, which is located in northern Shaanxi Province, China $(36^{\circ}21'-37^{\circ}19' \text{ N} \text{ and } 108^{\circ}38'-110^{\circ}29' \text{ E})$, is a typical hilly loess area on the Loess Plateau. The drainage area of the basin is 7591 km², and its elevation ranges from 560 m to 1760 m (Fig. 1). The slope of the basin varies from 0° to 85.3°, with a mean value of 17.7°. The main channel of the Yanhe River is 284.3 km long and originates in Jingbian County. It flows from northwest to southeast through Zhidan County, Ansai County,

Yanan City, and Yanchang County before entering the Yellow River. The soil in the basin developed from loess deposits, and the dominant soil type is loessial soil which is classified as Calcaric <u>eambisols</u> (FAO 2014). The Yanhe basin characterized by a semi-arid continental climate with warm and concentrated precipitation in summer and cold, dry winters with occasional snowfall. <u>According to the collected meteorological data from the Meteorological Institute of Shaanxi Province</u>, <u>China</u>, <u>1</u>The precipitation from 1952–2015 ranged from 300 mm to 803 mm, with a mean annual value of 495 mm. Additionally, the mean annual maximum and the minimum air temperatures <u>over this period</u> were 17.4°C and 4.2°C, respectively. Grassland, <u>farmlandcropland</u>, and woodland (mainly artificial woods) are the dominant land use types in this region. Most crops are cultivated on sloping lands, and woodlands are generally located on the steeper parts of the landform. The mean annual streamflow at the most downstream station of Ganguyi was 205×10^6 m³ from 1952 through 2008, and the streamflow from June to September accounted for 64.1% of the total annual discharge at this station. The GFGP was implemented in the basin in 1999, and the observed streamflow <u>exhibited showed</u> a decreasing trend during the 2010s (Fig. 2).

2.2 Data

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The SWAT model setup requires <u>data inputs including topography</u>, soil, <u>land use</u>, <u>climate and streamflow discharge</u>. <u>A digital</u> elevation model (DEM) of the basin with a 30-m resolution was obtained from the National Geomatics Center of China
(<u>http://www.ngcc.cn/)</u>. <u>TheA soil survey datamap, including with</u> 20-m resolution <u>a raster map with soil property database at a 20 m resolution, wereas obtained from the Institute of Soil and Water Conservation (ISWC). Chinese Academy of Sciences (CAS), and we developed a user's soil database required for the model that linked to the digital soil map using ArcGIS (version 10.2). LULC data from 1990, 2000 and 2010 with a-30-m resolution were supplied by the Institute of Remote Sensing and Digital Earth, CAS, and <u>-</u><u>T</u>the land used area was divided intoclassified into six main land use types according to the ecosystem classification system of China, including: grassland, woodland, cropland, water, residential land and barren land, according to the ecosystem classification system of China. Required <u>___daily</u> meteorological data for the SWAT model comprise daily precipitation, maximum and minimum temperature, relative humidity, wind speed and solar radiation, which were available which were collected from 1980 to 2015 from five county level meteorological stations: Jingbian, Zhidan, Aansai, Yan'an and Yanchang. These data were provided byfrom the Meteorological Institute of Shaanxi Province, China. The datasets consist of *Yanchang*. These data were provided byfrom the Meteorological Institute of Shaanxi Province, China. The datasets consist of *Yanchang*. These data were provided byfrom the Meteorological Institute of Shaanxi Province, China. The datasets consist of *Yanchang*. These data were provided byfrom the Meteorological Institute of Shaanxi Province, China. The datasets consist of *Yanchang*. These data were provided byfrom the Meteorological Institute of Shaanxi Province, China. The datasets consist of *Yanchang*. Thes</u>

daily precipitation, maximum and minimum temperature, relative humidity, wind speed and solar radiation, for five county level meteorological stations including Jingbian, Zhidan, Aansai, Yan'an and Yanchang stations. Observed monthly streamflow data at the Ganguyi hydrological station were obtained from the Yellow River Conservancy Commission (YRCC) from 1980-2010. The data were used for model calibration and validation, although some data gaps (missing values) were present. Ganguyi station is the most downstream station on the Yanhe River and controls an area of 5891 km². A digital elevation model (DEM) of the basin with a 30-m resolution was obtained from the National Geometrics Center of China (http://www.ngee.en/)-The soil data including a raster map with soil property database at a 20-m resolution, were obtained from the Institute of Soil and Water Conservation (ISWC), Chinese Academy of Sciences (CAS). LULC data from 1990, 2000 and 2010 with a 30 m resolution were supplied by the Institute of Remote Sensing and Digital Earth, CAS. The land area was divided into six main land use types, including grassland, woodland, cropland, water, residential land and barren land, according to the ecosystem

elassification system of China.

2.3 Model description

The SWAT model (version 2012), which was developed by the U.S. Department of Agriculture (USDA)—Agricultural Research Services (ARS), is a physically based, temporally continuous, and distributed watershed-scale hydrological model

15 (Neitsch et al., 2011: Douglas-Mankin et al., 2010). The SWAT-modeled hydrological cycle is based on the water balance, as documented by Arnold et al. (1998). SWAT has been widely tested and successfully used to explore the effects of climate and land use/management changes on watershed hydrology and water quality (Nyeko, 2015; Panagopoulos et al., 2011; Gassman et al., 2014; Douglas-Mankin et al., 2010; Zhang et al., 2013; Bosch et al., 2010; Wu and Chen, 2013; Xu et al., 2012; Zhang et al., 2011; Zhang et al., 2008; Qiu et al., 2012). Detailed descriptions of the mechanisms and structure of the SWAT model can be found in several literatureselsewhere (Neitsch et al., 2011; Arnold et al., 2012).

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2.4 Model setup and calibration/validation

A geographic information system (GIS) interface, ArcSWAT (Version 2012.10 2.18), was used to set up the model. A-The 30m-DEM was used to delineate the watershed with a threshold sub-basin area of 50 km², resulting in 88 sub-basins. Then, the land use map of 1990 and the soil map were imported and overlaid to create the hydrological response units (HRUs). HRUs are portions of a sub-basin that represents unique combinations of land use, soil and slope. To accurately reflect the spatial variability in the basin, multiple hydrological response units (HRUs) were selected, using a minimum threshold value of 5% for each land use, soil and slope categories, resulting in 1136 HRUs. To meet the needs of the model was used to delineate the basin. With the threshold area of subbasins set to 50 km² this process produced 88 subbasins. The land use map of 1990 and soil map were used to parameterize the SWAT model... the SWAT codes for the main land use types in the study area were defined as cropland (AGRR), woodland (FRST), grassland (RNGE), residential land (URBN), and barren land (BARR) in the attribute data. The SWAT model was calibrated using monthly streamflow data from Ganguvi hydrological station, and we accurately reflect the spatial variability in the basin, multiple hydrological response units (HRUs) were selected. A single HRU represents a unique combination of land cover, soil type and slope, and 1136 HRUs were established in the basin. We usadopted the SWAT-CUP (SUFI-2) -programmodule to perform a sensitivity analysis of-model parameters, and the SUFI 2 algorithm was used for optimization, SWAT was calibrated Calibration was conducted -using the first ten-year (1983–1992) record of of monthly-streamflow, and the -and-model was then validated using data from the subsequent eight years (1993–2000). Additionally, a three-year warm-up period (1980-1982) was used to minimize the effects of uncertain initial conditions (e.g., soil water storage) in the model simulation. To evaluate the model performance numerically, we used three statistical measuresincluding -: percent bias (PBIAS), Nash-Sutcliffe efficiency (NSE) (Nash and Sutcliffe 1970), and the coefficient of determination (squared correlation coefficient, r²). Typically, a model simulation is considered satisfactory with NSE>0.5, - $25\% \leq PBIAS \leq 25\%$, and $r^2 > 0.5$ (Neupane and Kumar, 2015; Wu and Chen, 2013).

2.5 Historical land use change and hydrological simulation Modeling scenarios

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In order to understand the historical land use change of the Yanhe basin, we conducted the spatiotemporal analysis of land use change using three land use conditions (land use maps from 1990, 2000 and 2010, corresponding to LU1990, LU2000 and LU2010, respectively), and the amount of specific land use types that changes from the initial time to the subsequent time was identified by overlaying the land use maps for different years. Then, the different land use data was taken as input of the calibrated SWAT model to isolate the effect of land use change under the historical climate from 1986-2015. Finally, the main hydrological components of water cycle (surface runoff, subsurface flow, ET, water yield and soil water) under different land use condition was assessed quantitatively at the HRU level. In this study, we used three land use conditions (land use maps in 1990, 2000 and 2010, corresponding to LU1990, LU2000 and LU2010, respectively) to evaluate the hydrological impacts of past land use changes from 1986 through 2015. To assess the long term effects of SLC on the hydrological cycle in the Yanhe basin, the land use condition of 2010 was set as the baseline scenario (BS), and four land conversion scenarios were reestablished based on the land use of 2010 and the SLC policy of the GFGP. During the baseline period, the model was driven by the 30 years (1986 2015) of climate data. Then, we established four hypothetical SLC scenarios. Scenario 1 (S1) and scenario 2 (S2) refer to the conversion of farmland on slopes steeper than 25° to grassland and woodland, respectively, and scenario 4 (S4) assumes that farmland on slopes greater than or equal to 15° is converted to grassland and woodland, respectively. The above four scenarios were implemented using the same climate forcing data to isolate the effects of SLC change.

2.6 Hypothetical SLC scenarios

SWAT was applied for assessing the long-term effects of SLC on the hydrological cycle in the Yanhe basin. The land use condition of 2010 was set as the baseline scenario (BS), and four hypothetical land conversion scenarios depicting the potential land use pattern were reestablished based on the land use of 2010 and the SLC policy of the GFGP. Scenario S1 and scenario S2 involve the conversion of cropland on slopes steeper than 25° to grassland and woodland, respectively, and scenario S3 and

15 S2 involve the conversion of cropland on slopes steeper than 25° to grassland and woodland, respectively, and scenario S3 and scenario S4 involve the conversion of cropland on slopes greater than or equal to 15° to grassland and woodland, respectively. To isolate the hydrological effects only caused by different land management, SWAT model with BS and four hypothetical scenarios were set up and run for 30 years using the historical climate forcing data from 1986-2015.

3 Results

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20 **3.1 Model performance evaluation**

We performed a global sensitivity analysis to examine 18 parameters that are potentially related to streamflow and identified the most sensitive 8 parameters for subsequent model calibration (Table 1). The selected 8 parameters represent key hydrological components such as surface runoff, soil water capacity and conductivity, ET, and groundwater recharge in this

region. Although the calibrated model slightly underestimated streamflow, the model performance was satisfactory in the calibration period (1983-1992), with NSE, r² and PBIAS values of 0.51, 0.71 and 15.7%, respectively (Fig. 3a). In the 8-year (1993-2000) validation period, the <u>streamflow characteristics were well captured by the calibrated model exhibited good</u> performance, with an NSE value of 0.82, although PBIAS <u>still</u>-indicated a 16.9% underestimation (Fig. 3b). In both the calibration and validation periods, the underestimation mainly occurred in <u>the spring</u> and winter, <u>when during which time</u> little discharge occurs in the watershed. Thus, the model performed poorly during the dry period. Furthermore, a discrepancy can be observed between measured precipitation and streamflow during the dry period because some small peaks in observed streamflow could not have been caused by the observed precipitation amount. The potential reasons for this issue include missing precipitation events due to a limited number of stations or errors in streamflow measurements. Nonetheless, both visual comparison and numerical evaluation indicated that the overall model performance was acceptable for simulating the hydrological processes in the Yanhe basin.

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3.2 Land use changes in Yanhe basin from 1990 to 2010

FarmlandCropland, woodland and grassland were the primary land use types in the Yanhe basin, and rapid land use change occurred from 1990 to 2010 (Table 2). <u>A downward trend was observed in Farmlandcropland exhibited a decreasing trend</u>,
which declineding from 40.2% of the basin area in 1990 to 17.6% in 2010-, whereas upward trends were observed in <u>Bb</u>oth grassland and woodland <u>exhibited increasing trends</u>, although these <u>upward</u> trends differed from each other. Grassland increased slowly from 44.7% to 45.3% between 1990 and 2000, followed by a rapid increase from 45.3% to 55.1% between 2000 and 2010. Woodland exhibited an increase of 73.9% from 1990 to 2000 followed by a smaller increase of 7.8% from 2000 to 2010. To identify the details of the land use conversion patterns, we explored the mutual transition conversion between among three major land use types <u>farmlandcropland</u>, grassland, and woodland (Table 3). The <u>conversion transition</u> of farmlandcropland to other land use types was observed in the periods of 1990-2000 and 2000-2010, while whereas the conversions of grassland and or woodland to other land use types were only evident from 1990-2000. The farmlandcropland area decreased dramatically from 1990 to 2010 because it was largeprimarily due to conversionted to woodland and or grassland, with conversion percentages of 16.7% and 39.5%, respectively, from 1990 to 2000 and of 6.4% and 33.2%.

respectively, from 2000 to 2010, respectively. Spatially, As shown in Fig. 4-shows that, the conversions mostly occurred in the central and northwestern parts of the basin from 1990–2000 and were scattered from 2000 to 2010. We also examined the land use conversions over the 20-year periods (1990 to 2010) and found that 19.6% and 54.5% of farmlandcropland was transformed to woodland and grassland, respectively, indicating evident LULC change due to the implementation of the GFGP policy in the basin (Table 3). Although the percentages of grassland and woodland increased, 20.1% of grassland and 19.5% of woodland were converted into farmlandcropland from 1990 to 2000. Additionally, mutual conversions were observed between grassland and other land use types, but the total area of grassland changed only slightly from 1990 to 2000. For example, 1204.6 km² of farmlandcropland and 281.5 km² woodland were converted into grassland, respectively, in the same period, resulting in_-a small net_change in the net area (Table 3). Additionally, despite the increase in woodland from 1990 to 2010, 12.6% and 31.1% of the woodland from 1990 were was converted into farmlandcropland and grassland, respectively, particular with much of this conversion occurring in the southern part of the basin.

3.3 Water balance components under different land use types

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To further analyze the quantitative impacts of LULC changes on the hydrological cycle, the main hydrological components of individual land use types were assessed via model simulation. The analysis results indicated that the water balance varied among different land use types in the basin (Fig. 5). The overland surface runoff per unit area from residential land was the highest, with an average value of 58.9 mm, followed by th<u>ose of at from</u> cropland and grassland, at 19.2 mm and 9.3 mm, respectively. Woodland <u>exhibited had</u> the lowest surface runoff of 3.7 mm, which was approximately <u>about</u>-80.7% lower than that of cropland and 62.2% lower than that of grassland (Fig. 5a). The water yield from individual land use types <u>exhibited had</u> a trends similar to <u>that those</u> of surface runoff. These results implied that the conversions of cropland and grassland to woodland could reduced runoff and streamflow at the regional scale. <u>By In</u> contrast, the ET of woodland areas was the highest, followed by those that of grassland, cropland, and residential land. With the exception of residential land, tThe soil water from otherof all land use types <u>except residential land</u> exhibited a <u>trend n inverse trend compared withopposite</u> that of ET (cropland > grassland > woodland); <u>thus</u>, woodland areas used more soil water for ET and generated less runoff and streamflow. Fig. 5b

shows the total water volumes on <u>for</u> the different land use types in the basin. The highest volumes of surface runoff and water yield were associated with cropland because of its large water volume per unit area and large total area in the basin. Grassland had the largest soil water storage and highest ET in the basin. The surface runoff, water yield, soil water, and ET of woodland were the lowest among different land use types, <u>with the except_ion offor</u> residential land, due to the small total area of woodland in 1990.

3.4 Hydrological impacts of historical land use changes

The impacts of land use changes were simulated through by applying the model to three land use scenarios (LU1990, LU2000, and LU2010) with the same climate forcing data (from 1986 to 2015). The results showed that simulated streamflow at the outlet of the Yanhe basin exhibited a negative responses to land usedownward trend with land use changes frombetween 1990 and 2010, and the magnitude of the decrease-reduction under LU2010 was larger than that under LU2000 (Fig. 6a). Further analysis showed that relatively large responses mainly occurred from May to September (Fig. 6b). Table 4 presents the simulated average annual components of the water balance in the Yanhe basin. A decreasing trend was observed in S₂urface runoff generated at the basin scale exhibited a decreasing trend as land use changes occurred from 1990 to 2010, and it-with decreased by of 16.6% in under LU2000 and 29% in under LU2010. Land use changes had minor effects on subsurface flow, with a slight increase of 1.8% in under LU2000 and of 2.7% underin LU2010 eompared with relative to flow under that in LU1990. Similar to the subsurface flow, ET displayed a weak increasing trend in the basin due to land use changes. The water yield in the basin under LU2000 and LU2010 changed in a similar manner as surface runoff, with decreases of 5.6% and 9.9%, respectively, compared with that relative to water yield of LU1990. The soil water decreased by 11% from 1990 to 2010, with a rapid decrease from 1990-2000 and a slow decrease afterwardthereafter.

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The spatial distributions of the main hydrological components at the HRU scale are shown in Fig. 7. The central part of the basin <u>exhibited produced high</u> surface runoff, <u>while whereas the most</u> western portion and <u>the southern edge of the basin</u> <u>exhibited had</u> little runoff (Fig. 7a and 7e). Additionally, the decreasing trend <u>e in surface runoff from 1990 to 2010 mainly</u> <u>occurred on was concentrated in</u> the middle part of the basin from 1990 to 2010. To further understand the hydrological changes associated with land use change, we analyzed the long-term (30-year) average total water volumes of the main hydrological

components (Fig. 8). The results showed that decreases in total surface runoff mainly occurred in cropland areas because of the its decreased area-of-eropland. The associated decreased percentages were 26.8% under LU2000 and 57.3% under LU2010 relative to that the surface runoff of LU1990 (Fig. 8a). Although the total surface runoff generated in grassland and woodland areas increased from LU1990 to LU2010, the total magnitude of these increases could not compensate for the decreasing decrease associated with cropland. Thus, the total surface runoff of the basin decreased. The water yield exhibited displayed a trend similar to that of surface runoff from LU1990 to LU2010 (Fig. 7b and f)—the total water yield decreased by 60.9% in cropland areas and increased by 18.8% and 136.4% on-in grassland and woodland, respectively. These changes were due to the dramatic conversion of cropland to grassland and woodland between LU1990 and LU2010 (Fig. 8b). Similar to the spatiotemporal trends in the water yield, ET increased under LU2010 compared with that officiative to LU1990 (Fig. 7c and g), whereas it decreased by 58.7% in cropland and increased by 21.4% and 124.0% for in grassland and woodland, respectively, from LU1990 to LU2010, respectively (Fig. 8c). In terms of the spatial distribution of soil water, an evident decrease occurred under LU2010, and with the highest decrease occurred_occurring_in the north-central part of the region (Fig. 7d and h). Although the total soil water increased in woodland exhibited an increasing trend with land use change, it was the total soil water was lowest in woodland when compared with the soil water in cropland and grassland areas under the all three land use conditions scenarios (Fig. 8d).

3.5 Potential impacts of hypothetical SLC

The potential hydrological effects of SLC were projected through by inputting four land conversion scenarios that were reconstructed based on the LU2010 baseline (Table 5). The changes in land use in S1 led to the conversion of 131.6 km² cropland with slope >25° to grassland, which resulted in slight decreases in surface runoff, water yield and soil water, whereasile minor increase were found in subsurface flow and ET exhibited minor increase. With the same acreage of cropland to woodland conversion in S2, the induced hydrological effects were similar to those in S1, whereas the differences between S2 and the baseline were larger in magnitude than those between S1 and the baseline. When the land conversion was extended to cropland with slope \geq 15°, the trends of hydrological change were further strengthened. For example, in S3, grassland increased by 23.9% and 15.3% on land surfaces with slopes of 15°-25° and >25°, respectively, resulting in 12.4% and 2.7%

decreases in surface runoff and the water yield, respectively, compared relative to the baseline. Conversely, the subsurface flow and ET-minimally increased slightly by 2.2% and 0.2%, respectively. Although soil water consistently exhibited a negative response, the decrement decreases wereas small. When the same area of cropland with slope $\geq 15^{\circ}$ was converted to woodland in S4, the surface runoff and water yield decreased by 17.1% and 6.4% compared relative to the baseline; while subsurface flow and ET exhibited increased by 3.9% and 0.3% increases, respectively. Soil water notably changed in S4, decreasing by 6.4% compared with relative to the baseline soil water. Fig. 9 illustrates the spatial response of soil water to land use conversion scenarios compared with relative to the baseline of 2010. Soil water decreases in S4 mainly occurred along the southern edge of the basin and the north-central part of the basin, which are areas of have a high elevation and high slope gradient (Fig. 9d).

4 Discussion

- In this study, the performance of the SWAT model was evaluated using monthly streamflow data from the Ganguyi hydrological station. Although the calibrated model slightly underestimated streamflow in the dry period in both the calibration and validation periods, the three statistical indexes (NSE, PBIAS, r²) indicated that the modelling accuracy was acceptable. The potential reasons for the underestimation can be categorized as "measurement quality" and "model behavior" reasons. "Measurement quality" refers to missing precipitation data or streamflow measuring error. The study area lies in the hilly-gullied loess region, and the complex and rugged terrain results in highly variable precipitation and difficulties in spatially estimating the precipitation (Cao et al., 2006). Additionally, this region has a limited number of meteorological and hydrological stations. In this study, five meteorological stations were involved; however, only three stations are located inside the basin, with the remaining two occurring outside the basin (Fig. 1). It has been reported that the accuracy of streamflow prediction mainly depends upon the precipitation gage numbers and their locations (Cao et al., 2006; Mul et al., 2009). Thus, the insufficient precipitation records and the distances among the meteorological stations such as the inadequate
 - representation of the physical mechanism of hydrological processes. For example, SWAT uses total daily precipitation and does not consider rainfall intensity within a day; thus, it can underestimate streamflow for some precipitation events (Oiu et al., 2012). Another example is the use of runoff curve number to simulate the surface runoff behavior. This

approach does not account for saturation excess runoff or contributions from variable source areas (Garen and Moore, 2005: Easton et al., 2008). The potential reasons for this issue include missing precipitation events due to a limited number of stations or errors in streamflow measurements. Although the simulation results in our research met the calibration criteria, the model performance can be improved if precipitation data from additional gages become available in the basin.

FarmlandCropland, grassland and woodland were the primary land use elasses types in the Yanhe basin, and the sum of their area accounted for more than 95% of the total area during the our study period under study. To investigate the effects of land use change and management practices on the water balance, three land use scenarios were assessed using the SWAT model. The contributions of individual land use types to basin hydrology differed. Cropland had the highest soil water which can

- 10 be partly attributed to the topography of cropland. The cropland area was mainly distributed in the bottom region of the valley and on land with gentle slope. The valley bottom generally receives more water flow from the hillslope and has lower temperatures than higher regions, resulting in more soil water (Wang et al., 2012). Furthermore, land with gentle slope has higher soil water retention ability than does land with steep slope (Panagopoulos et al., 2011). In addition, cropland had lower ET per unit area than did the other land use types. In contrast, grassland had lower soil water, which
- 15 can be attributed to its higher soil water loss through ET. Wang et al. (2012) reported that grass cover types cannot protect the soil surface from solar radiation, leading to greater water loss via direct evaporation. Woodland and, to a lesser extent, Compared to cropland, ggrassland and woodland areas were generally associated with reduced lower surface runoff and decreases in the water yield of in the basin, particularly woodland. This result is consistent with those of previous studies. which showed demonstrating that woodland areas captures more rainfall and uptake more water than do other land use types 20 (e.g., cropland and grassland) (Jian et al., 2015; Wang et al., 2012), resulting in a reduced lower runoff - and water discharge (Wang et al., 2013; Duan et al., 2016). Additionally, Qiao et al. (2017) reported that the reduced runoff in areas with woody plant relative to grassland areas is associated with shift in runoff generation mechanisms; thus, the shift from saturation excess overland flow to infiltration excess overland flow might also have contributed to the reduced runoff in woodland areas observed in the present study. Additionally, compared with other vegetation types, wwoodland areas might lost lose more water through ET-compared to other vegetation types. This trend-pattern was demonstrated evident by from the change in soil water.
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exhibiting with less soil water on their woodland than in other areas at under the same precipitation compared to the trends in other areas amount. Less In addition, another possible explanation for the low soil water volumes in was observed in woodland areas was that because forests in the Yanhe basin generally grow on landform with high slopes. O; our analysis indicated that more than 62% of woodland was located on slopes \geq 15° (Table 5). Steeper slopes generally retain less soil water due to low infiltration-rates and rapid surface drainage (Famiglietti et al., 1998). Thus, a large amount of precipitation was associated with forest growth and ET rather than discharge out of the basin as surface runoff and streamflow. Such water patterns prevent water loss but at the expense of reduced soil water in the region.

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After the implementation of the GFGP, the area of cropland decreased continuously because it was transformed into grassland and woodland from 1990 to 2010, and the conversions among land use types were the main forms of land use change 10 in the Yanhe basin. These results are consistent with previous findings regarding trends in land use change on the Loess Plateau (Li et al., 2016). We compared monthly and annual average streamflow under three land use scenarios, and the results showed that the decreased average monthly streamflow in the rainy season is the primary mechanism for the decrease in annual average streamflow in the basin. In addition, we found that the decrease in surface runoff was the main reason for the streamflow decrease in the basin, and our quantitative evaluation suggested that surface runoff decreased by 29.1% from 1990 to 2010 due to land use change-from 1990 to 2010. A similar conclusion was drawn by Farley et al. (2005), who studied 26 watersheds globally to assess the hydrological effect of afforestation and found that annual runoff was reduced by 31-44% on average. Notably, the average annual volume of soil water exhibited presented an evident decrease under the 2010 land use scenario. This trend may might be due to two interrelated reasons. First, the area of cropland decreased dramatically in 2010, while and cropland exhibited was characterized by the highest soil water volume per unit area relative to other land use types (Fig. 5a). leading to the large decrease in soil water. Second, a certain proportion of cropland was converted to grassland and woodland, which had the lowest soil water per unit area; thus, the soil water decreased due to the cropland area reduction, and this decrease could not be offset by increases in total soil water in grassland and woodland areas (Fig. 8). Although the area of woodland peaked in 2010, the total soil water in woodland areas increased only slightly-increased. This trend could have occurred main reason for this phenomenon might be that because more water was consumed by woodland than cropland and grassland. Additionally, cropland areas with high slopes were converted to woodland after the GFGP was implemented (Zhou et al., 2012),

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whereas and areas with steep slopes have lower soil moisture retention potentials than <u>do</u> areas with gentle slopes (Pachepsky et al., 2001). Thus, topographical factors also played important roles in the spatial heterogeneity of the water balance (Qiu et al., 2001; Bi et al., 2008).

Four future SLC scenarios were proposed to demonstrate the hydrological effects of land use changes.-Increases in grassland or-and_woodland cover due to the land conversion of cropland with slopes ≥15° or >25° had negative effects on surface runoff, the-water yield and soil water, whereas and positive effects on subsurface flow and ET-increased. We found that the magnitudes of the hydrological effects of the conversion of slopedsloping cropland to woodland were greater than those associated with the conversion of slopedsloping cropland to grassland. This result suggests that the expansion of woodland could reduce runoff generation and drainage because of overland flow retention and intensification of ET in this region. However, the revegetated sloping land was prone to reduced soil moisture due to its steep slope. Some studies have reported that revegetation can cause soil water shortages in both the near-surface soil and deep soil layers (Farley et al., 2005; Jian et al., 2015), likely resulting in soil desiccation (Wang et al., 2011; Fu et al., 2012). Therefore, watershed management should consider all water balance components (Duan et al., 2016), and vegetative structure and management measures should be optimized to improve-the ecohydrological functions and promote-the watershed sustainability.

15 **5** Conclusions

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In this study, we investigated the land use changes from 1990 to 2010 based on a transition analysis. Cfound that cropland, grassland and woodland were the dominant land use types in the Yanhe basin, and land use conversion occurred among these types occurred since 1990 due to the implementation of the GFGP. The decrease in cropland led to increase in grassland and woodland. The impacts of LULC changes on the water balance components were assessed quantitatively using the SWAT model___and_three periods of land use maps and four hypothetical SLC scenarios based on the GFGP policy on the Loess Plateau. Our analysis showed that cropland was associated with the highest surface runoff and water discharge per unit area, followed by grassland and woodland. These differences can partly explain the relationships between hydrological characteristics and land use change at the basin scale. Surface runoff and water yield exhibited decreaseding trends due to land use changes from 1990 to 2010; while subsurface flow and ET increased. Consequently, soil water decreased between 2000

and 2010. By adopting four cropland land conversion scenarios, we found that the function of reducing surface runoff was more effective when croplands with slopes $\geq 15^{\circ}$ were converted into grassland or woodland compared to than when converting areas with slopes $\geq 25^{\circ}$ were converted. Furthermore, planting woodafforestation on sloping land had greater hydrological effects than did planting grass. Notably, surface runoff and soil water decreased and ET increased. Overall, this study provides useful information for land use planning and soil and water conservation on the Loess Plateau, and Ffurther studies are required to investigate the optimization of the vegetative structure and the avoidance of undesired hydrological effects.

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Parameters	Description	Range	Optimized value/ percent change
CN2	Soil conservation service (SCS) runoff curve number	-20%~20%	-4.03% ^r
SOL_AWC	Soil available water capacity	-20%~20%	10.71% ^r
SOL_K	Soil saturated hydraulic conductivity (mm/h)	-20%~20%	7.36% ^r
ESCO	Soil evaporation compensation factor	0~1	0.51 ^v
EPCO	Plant uptake compensation factor	0~1	0.65 ^v
ALPHA_BF	Baseflow alpha factor (day)	0~0.5	0.36 ^v
GWQMN	Threshold depth of water in the shallow aquifer required for return flow to occur (mm)	100~1200	906.20 ^v
SURLAG	Surface runoff lag coefficient	0.05~6	2.93 v
REVAPMN	Threshold depth of water in the shallow aquifer for "revap" to occur (mm)	100~1200	499.40 ^v

Table 1 Sensitive parameters for streamflow simulation and calibrated values

Note: the superscripts "r" and "v" in the column of optimized value indicate the percent changes based on initial values and replaced parameter values, respectively.

	Farmland	<u>Cropland</u>	Gras	sland	Woo	dland	Resider	ntial land	W	ater	Barre	en land
	Area (km ²)	Ratio <u>Pc</u> <u>t.</u> (%)										
1990	3051.2	40.2	3390.4	44.7	1059.6	13.9	21.4	0.28	65.4	0.86	3.1	0.040
2000	2246.4	29.6	3436.3	45.3	1842.4	24.3	43.9	0.58	17.2	0.23	0.7	0.009
2010	1338.3	17.6	4178.7	55.1	1986.3	26.2	66.4	0.87	16.6	0.22	0.5	0.006

Table 2 Land use changes in the Yanhe basin from 1990 to 2010

	1990-2000		2000-2010		1990-2010	
	Area (km ²)	Ratio <u>Pct.</u> (%)	Area (km ²)	Ratio <u>Pct.</u> (%)	Area (km ²)	RatioPct (%)
FarmlandCropland to woodland	509.2	16.7	143.6	6.4	598.5	19.6
FarmlandCropland to grassland	1204.6	39.5	745.4	33.2	1661.8	54.5
FarmlandCropland to residential land	15.9	0.5	20.7	0.9	28.5	0.9
FarmlandCropland to barren land	0.3	0.0	0.0	0.0	0.1	0.0
Grassland to woodland	762.3	22.5	0.6	0.0	796.7	23.5
Grassland to farmlandcropland	683.1	20.1	1.5	0.0	410.4	12.1
Grassland to residential land	7.1	0.2	0.6	0.0	10.4	0.3
Grassland to barren land	0.4	0.0	0.0	0.0	0.3	0.0
Woodland to farmland <u>cropland</u>	206.5	19.5	0.0	0.0	133.8	12.6
Woodland to grassland	281.5	26.6	0.0	0.0	330.1	31.1
Woodland to residential land	5.2	0.5	0.5	0.0	9.4	0.9
Woodland to barren land	0.0	0.0	0.0	0.0	0.0	0.0
Barren land to woodland	0.2	7.4	0.0	0.0	0.2	7.4
Barren land to grassland	0.8	25.9	0.0	0	1.0	31.7
Barren land to farmlandcropland	2.0	65.1	0.0	0.0	1.7	54.6
Barren land to residential land	0.0	0.0	0.0	0.0	0.2	6.4

Table 3 Primary patterns of land use transition change in the Yanhe basin from 1990 to 2010

Table 4 Simulated average annual values of hydrological components in the Yanhe basin under different land use conditions

	Surface runoff (mm)	Subsurface flow (mm)	ET (mm)	Water yield (mm)	Soil water (mm)
LU1990	15.1	22.2	461.8	37.3	123.0
LU2000	12.6	22.6	464.5	35.2	113.1
LU2010	10.7	22.8	466.5	33.6	109.5

		FarmlandCropland	Woodland	Grassland	Residential land	Barren land
	<15°	814.3	761.2	1677.7	52.5	0.3
LU2010	15°~25°	392.7	808.4	1640.8	10.8	0.1
	>25°	131.6	417.3	860.9	3.2	0.0
	<15°	_	_	_		_
S 1	15°~25°	_	—	—		_
	>25°	0	417.3	992.5	3.2	0.0
<	<15°	_		_		_
S2	15°~25°	_	_	_	_	_
	>25°	0	548.9	860.9	3.2	0.0
<15	<15°	_		_		_
S3	15°~25°	0	808.4	2033.5	10.8	0.1
	>25°	0	417.3	992.5	3.2	0.0
S4	<15°	_			_	_
	15°~25°	0	1201.0	1640.8	10.8	0.1
	>25°	0	548.9	860.9	3.2	0.0

Note: '—' indicates the same value as that of LU2010.

Table 6 Simulated average annual values	of hydrological components	s under different scenario	s of slope land conversion

	Surface runoff (mm)	Subsurface flow (mm)	ET (mm)	Water yield (mm)	Soil water (mm)
S 1	10.4	22.9	466.8	33.3	109.0
S2	10.3	23.0	466.9	33.2	107.7
S3	9.37	23.3	467.6	32.7	107.5
S4	8.87	23.7	467.8	32.6	102.5

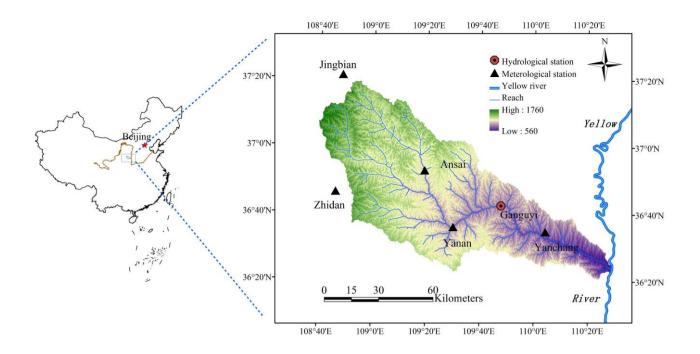


Fig. 1 Map of the location and elevation of the Yanhe basin showing the elevation and the locations with of the distribution of meteorological and hydrological stations

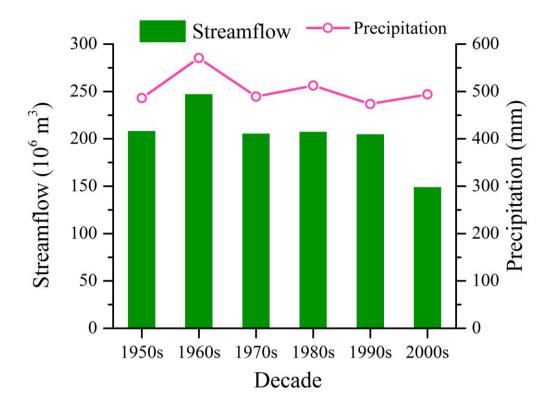


Fig. 2 Precipitation and streamflow changes in different decades in the Yanhe basin

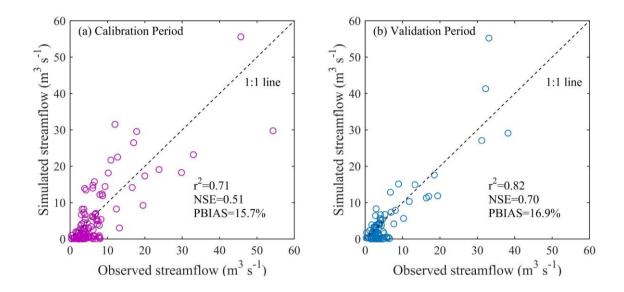


Fig. 3 Model performance for the (a) 10-year (1983-1992) calibration period and the (b) 8-year (1993-2000) validation period

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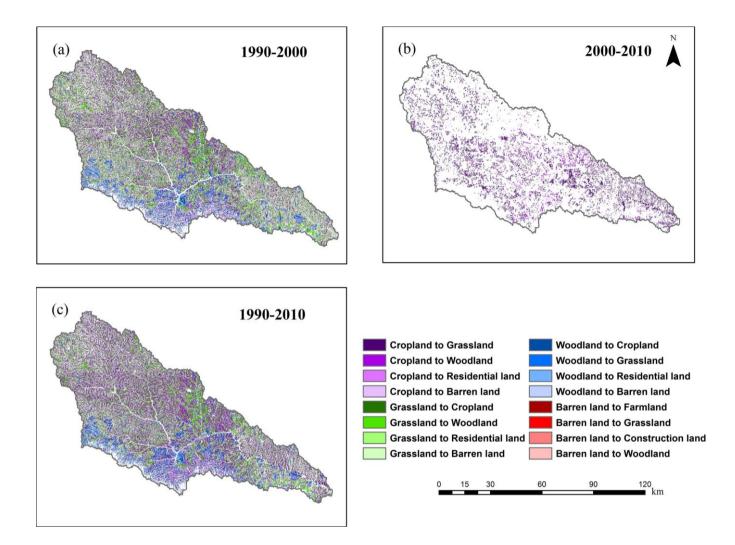


Fig. 4 Land use transitions changes from 1990 to 2010 in the Yanhe basin from 1990 to 2010

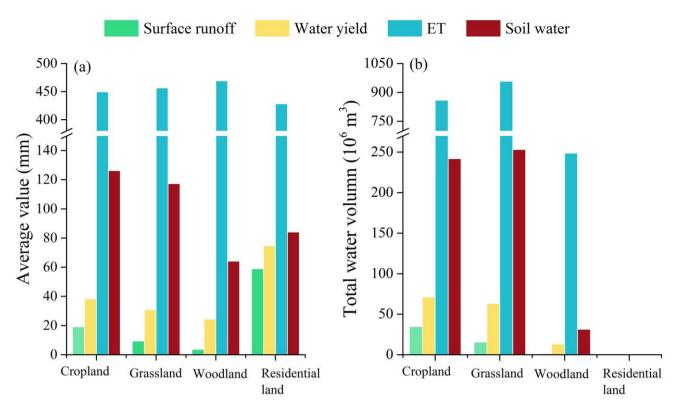


Fig. 5 Calculated (a) unit area and (b) total annual hydrological components <u>in-for</u> the 18-year simulation period (1983-2000) for different land use types in the Yanhe basin

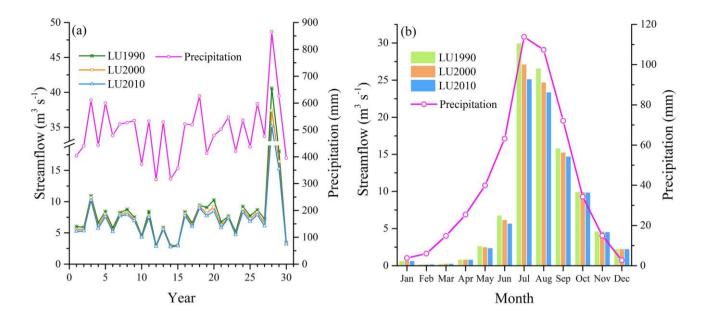


Fig. 6 (a) Simulated annual streamflow and (b) simulated multiyear average monthly streamflow (the data were averaged from 1986 to 2015) at the outlet of the Yanhe basin under different land use conditions

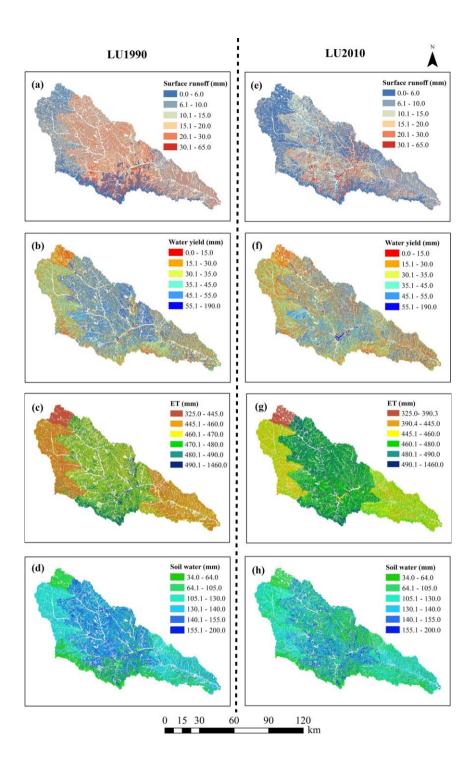


Fig. 7 Simulated spatial distribution of (a) surface runoff, (b) water yield, (c) ET and (d) soil water under land use change between 1990 and 2010

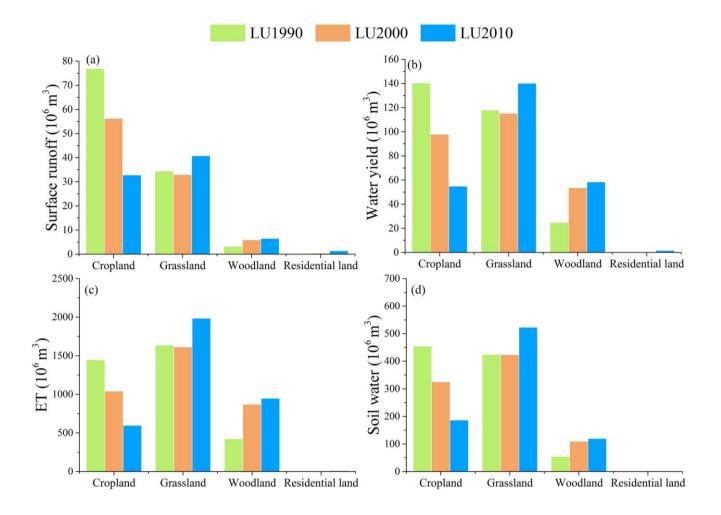


Fig. 8 Calculated average annual total water volume under different land use scenarios

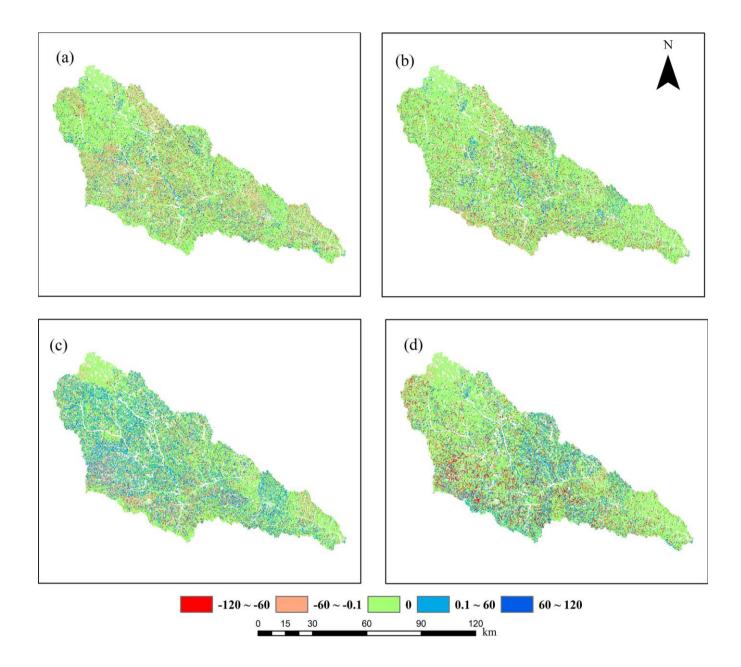


Fig. 9 Calculated soil water difference between the baseline scenario and (a) S1, (b) S2, (c) S3 and (d) S4

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