

We appreciate the editor and both referees for their valuable and insightful comments, which have greatly improved our manuscript. Below we describe the modifications made according to their comments. For clarity, comments are given in italics and blue, and our responses are given in plain text. The line numbers within brackets indicate the location of the modifications in the revised manuscript. The revised manuscript with all revisions tracked is appended at the end of this document.

-Referee#1-

There are many researches focusing on reasons, causes and modelling of nonstationarity of hydrological extremes such as: Xihui Gu, Qiang Zhang, Vijay P. Singh, Peijun Shi, 2017. Nonstationarities in the occurrence rate of heavy precipitation across China and its relationship to climate teleconnection patterns. International Journal of Climatology, DOI: 10.1002/joc.5058. Xihui Gu, Qiang Zhang, Vijay P. Singh, Peijun Shi, 2017. Changes in magnitude, frequency and timing of heavy precipitation across China and its potential links to summer temperature. Journal of Hydrology, 547, 718-731. Xihui Gu, Qiang Zhang, Vijay P. Singh, Peijun Shi, 2017. Nonstationarity in timing of extreme precipitation across China and impact of tropical cyclones. Global and Planetary Change, 149, 153-165. Xihui Gu, Qiang Zhang, Vijay P. Singh, Lin Liu, 2016. Nonstationarity in the occurrence rate of floods in the Tarim River basin, China, and related impacts of climate indices. Global and Planetary Change, 142, 1-13. Qiang Zhang, Xihui Gu, Vijay P. Singh, Mingzhong Xiao, Xiaohong Chen, 2015. Evaluation of flood frequency under non-stationarity resulting from climate change and human activities in the East River basin, China. Journal of Hydrology, 527, 565-575. Qiang Zhang, Xihui Gu, Vijay P. Singh, Mingzhong Xiao, Chong-Yu Xu, 2014. Stationarity of annual flood peaks during 1951-2010 in the Pearl River basin, China. Journal of Hydrology, 519, 3263-3274. What are the motivations, research objectives and novel points of this current study when compared to standing researches? My strong suggestion is that thorough literature review is pretty necessary. New findings, new ideas, new methods, if any, should be pointed out with enough citations to justify authors' statements.

AUTHORS' REPONSE: Thank you for introducing the overlooked references and for the good suggestion of enhancing the literature review. Following the advice of the reviewer, we have made a more comprehensive literature survey by citing and discussing important recent publications in the field, including those introduced by the reviewer. We have also improved the description of motivations, research objectives and novel points of this current study. The related paragraphs have been changed into the following:

“In hydrological analysis and design, conventional frequency analysis estimates the statistics of a hydrological time series based on recorded data with the stationary hypothesis which means that this series is ‘free of trends, shifts, or periodicity (cyclicality)’ (Salas, 1993). However, global warming and human forces have changed

climate and catchment conditions in some regions. Time-varying climate and catchment conditions can affect all aspects of the flow regime, i.e. changing the frequency and magnitude of floods, altering flow seasonality, and modifying the characteristics of low flows, etc. The hypothesis of stationarity has been suspected (Milly et al., 2008). If this problematic method is still used, the frequency analysis may lead to high estimation error in hydrological design. Therefore, considerable literatures have introduced the concept of hydrologic nonstationarity into analysis of various hydrological variables, such as annual runoff (Arora, 2002; Jiang et al., 2017; Jiang et al., 2015; Liu et al., 2017; Xiong et al., 2014; Yang and Yang, 2013), flood (Chen et al., 2013; Gilroy and Mccuen, 2012; Gu et al., 2016; Kwon et al., 2008; López and Francés, 2013; Tang et al., 2015; Xiong et al., 2015b; Yan et al., 2016; Zhang et al., 2014; Zhang et al., 2015), low flow (Du et al., 2015; Jiang et al., 2014; Liu et al., 2015), precipitation (Cheng and AghaKouchak, 2014; Gu et al., 2017a, b, c; Mondal and Mujumdar, 2015; Shahabul Alam et al., 2014; Villarini et al., 2010) and so on. Compared with the literatures on annual runoff, floods and precipitation, the literatures on the nonstationary analysis of low flow are relatively limited.” [Lines 62-78]

“To our knowledge, compared with the nonstationary flood frequency analysis, the studies on the nonstationary frequency analysis of low-flow series is not very extensive because of incomplete knowledge of low flow generation (Smakhtin, 2001). Most of these studies explain nonstationarity of low-flow series only by using climatic indicators or a single indicator of human activity. However, the indicators of catchment conditions (e.g. recession rate) related to physical hydrological processes have seldom been attached in nonstationary modeling of low flow series. This lack of linking with hydrological process makes it impossible to accurately quantify the contributions of influencing factors for the nonstationarity of low flow series, and such a scientific demand for tracing the sources of nonstationarity of low-flow series and qualifying their contributions motivated the present study...” [Lines 100-109]

“The goal of this study is to trace origins of nonstationarity in low flows through developing a nonstationary low-flow frequency analysis framework with the consideration of the time-varying climate and catchment conditions (TCCCs) and human activity (HA). In this framework, the climate and catchment conditions are quantified using the eight indices, i.e., meteorological variables (total precipitation P , mean frequency of precipitation events λ , temperature T and potential evapotranspiration ET), basin storage characteristics (base-flow index BFI , recession constant K) and aridity indexes (climate aridity index AI_{ET} , the recession-related aridity index AI_K). The specific objectives of this study are: (1) to find the most important index to explain the nonstationarity of low-flow series; (2) to determine the best subset of TCCCs indices and/or human activity indices (i.e., population POP , irrigation area IAR , and gross domestic product GDP) for final model through stepwise selection method to identify nonstationary mode of low-flow series; and (3) to quantify the contribution of selected explanatory variables to the nonstationarity.” [Lines 131-142]

There are no exact and/or results included in the Abstract section. Or only limited words describing results. More details and particularly in a quantitative way should be provided for description of results and conclusions

AUTHORS' REPONSE: The reviewer is correct. In the modified abstract, we have provided more quantitative results and conclusions. In the revision of the second part of the Abstract, the description of results and findings has been modified as following:

“The results from stepwise regression for the optimal explanatory variables show that the variables related to irrigation, recession, temperature and precipitation play an important role in modeling. Specifically, analysis of annual minimum 30-day flow in Huaxian shows that AI_K is of the highest relative importance among the optimal variables, followed by IAR (note to reviewer: *Irrigated area – a newly added index in the revised version*), BFI and AI_{ET} ; and nonstationary GA distribution model with these optimal variables has an AIC value of 207.0, while the AIC values of other models just with AI_K or time as explanatory variables or without any variable are 217.4, 225.5, 232.3, respectively. We conclude that the incorporation of multiple indices related to low-flow generation permits tracing various driving forces.” [Lines 29-36]

In Introduction section, it was noticed that there are numerous researches focused on nonstationary low flow frequency analysis. However, no novel points were listed and hence research motivations were not well justified. Besides, as a tributary of the Yellow River, evaporation or evapotranspiration, irrigation, population, GDP and so on should be included as factors influencing low flow changes. Related works have been done using Budyko framework by Prof. Dawen Yang from Tsinghua University and Prof. Qiang Zhang from Beijing Normal University and other colleagues from China. Besides, I still have no idea about how the authors developed the framework to evaluate low flow frequency from a nonstationary perspective.

AUTHORS' REPONSE: Thanks to the reviewer for pointing out this. Firstly, in the introduction section, the related part has been reorganized and modified in order to clarify our research motivations more clearly. Related part for study motivation, refer to the response of the first comment above.

Secondly, following the reviewer's advice, besides the following indices (K , AI_K and BFI) that are related to human activities and the indices (K , AI_K and BFI) that are linked to physical hydrological processes, in the revised version, we have included the irrigation area (IAR), the Gross Domestic Product (GDP), and population (POP) indices. The process of their change with time has been presented (see Fig. 5 in the revised manuscript). The Pearson correlation coefficients between low-flow series and these indices have been presented (see Fig. 6 in the revised manuscript). The models (M2b, M5, and M6, as described in Table 2 in the revised manuscript) are added. The summary of their results have been presented in Table 5 in the revised manuscript. Analysis of all new results has been shown in Figs. 7, 8, 9, 10 and 11 in the revised manuscript. Besides, we have added the following statements in the discussion section

in the revised manuscript.

“The related researches (Jiang et al., 2015; Yang and Yang, 2011; Yang and Yang, 2013; Zhang et al., 2015) have applied the Budyko framework to analyze the impacts of climate change and/or human activity on annual runoff. Indeed, for annual runoff, the Budyko framework is a good method because it used the mean annual water-energy balance equation to consider generation process of total runoff. Unfortunately, to our knowledge, there is a lack of equation derived from basic physics laws for generation process of low flows. Therefore, we emphasize the importance of TCCCs variables to modeling of low-flow nonstationarity.” [Lines 507-513]

Thirdly, the framework is composed of the time varying and GLM methods, and the method of stepwise selection for TCCCs indices and human activity indices. And to address this comment, we have added a flow chat of methodology (Fig. 1 in the revised manuscript) to explain how the framework is organized.

In Method section, a working framework should be formulated besides some descriptions.

AUTHORS’ REPONSE: Thank you for your good suggestion. Following the reviewer’s suggestion we have added a flow chart of methodology, as show in Fig. 1 in the revised manuscript.

Why the authors choose Weihe River basin as a case study? Are there any unique features of the study region when compared to other alternative rivers?

AUTHORS’ REPONSE: The nonstationarity of annual runoff in Weihe River basin has been shown to be very significant (Lin et al., 2012; Xiong et al., 2014). The previous studies have demonstrated that the climate change and human activities play an important role in annual runoff changes. When compared to other alternative rivers, the nonstationarity mode of low flows in the study region is so complex that it is difficult to be identified due to the influence of various factors. This feature aroused our interest in choosing the study area. We try to demonstrate that the nonstationarity of low flows in this basin is caused by multiple factors and more effective analysis model should incorporate not only a single climate index or human activity indices but also the other climate indices and catchment condition indices. We have clarified this point.

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Zhang, S., Yang, H., Yang, D., and Jayawardena, A. W.: Quantifying the effect of

vegetation change on the regional water balance within the Budyko framework, Geophysical Research Letters, 43, n/a-n/a, 2015.

-Referee#2-

General Comment *This work covered an interesting topic. It is qualified for HESS after a minor revision. Authors incorporated multiple variables into time-varying model by GLM, and called this a nonstationary mode considering TCCCs. They calculated and compared AIC of this mode with that of the stationary mode and the nonstationary mode with a single covariate in two stations in Weihe. Then they concluded this TCCCs nonstationary mode was the optimal one for nonstationary low flow frequency analysis in Weihe.*

AUTHORS' REPONSE: Thank you for your positive evaluation and a good summary of the paper.

It's a pity that they did a lot of work without clearly stating their motivation. Authors first raised an issue in review that the previous studies in low flow failed to provide a link between hydrological process and frequency analysis, and this made it difficult for tracing the origins of low flow change. While readers might think they intend to trace these origins (which was also hinted by the title), they defended that "the goal of this study is to develop a nonstationary low-flow frequency analysis framework". It is better for them to keep consistent in the whole introduction section.

AUTHORS' REPONSE: Thank you for your comments and the good suggestion. This is also pointed out by reviewer 1. Following the reviewer's advice, we have also better stated our study motivation as following.

"To our knowledge, compared with the nonstationary flood frequency analysis, the studies on the nonstationary frequency analysis of low-flow series are not very extensive because of incomplete knowledge of low flow generation (Smakhtin, 2001). Most of these studies explain nonstationarity of low-flow series only by using climatic indicators or a single indicator of human activity. However, the indicators of catchment conditions (e.g. recession rate) related to physical hydrological processes have seldom been attached in nonstationary modeling of low flow series. This lack of linking with hydrological processes makes it impossible to accurately quantify the contributions of influencing factors for the nonstationarity of low flow series, and such a scientific demand for tracing the sources of nonstationarity of low-flow series and qualifying their contributions motivated the present study..." [Lines 100-109]

We have also explicitly defined and stated the study objectives in the 5th paragraph of the Introduction Section, as follows:

"The goal of this study is to trace origins of nonstationarity in low flows through developing a nonstationary low-flow frequency analysis framework with the consideration of the time-varying climate and catchment conditions (TCCCs) and

human activity (HA). In this framework, the climate and catchment conditions are quantified using the eight indices, i.e., meteorological variables (total precipitation P , mean frequency of precipitation events λ , temperature T and potential evapotranspiration ET), basin storage characteristics (base-flow index BFI , recession constant K) and aridity indexes (climate aridity index AI_{ET} , the recession-related aridity index AI_K). The specific objectives of this study are: (1) to find the most important index to explain the nonstationarity of low-flow series; (2) to determine the best subset of TCCCs indices and/or human activity indices (i.e., population POP , irrigation area IAR , and gross domestic product GDP) for final model through stepwise selection method to identify nonstationary mode of low-flow series; and (3) to quantify the contribution of selected explanatory variables to the nonstationarity.” [Lines 131-142]

Besides, to better show the advantage of this framework, which was composed of the time-varying and GLM method, they should compare it with other models using only climatic indicators or a single indicator of human activity, just as they mentioned in the review, not just the mode with either AIK or BFI as the explanatory variable.

AUTHORS’ REPONSE: Thank you for the comment. Our study had included the model with climate indicators. But, indeed, the model with a single indicator of human activity (e.g. irrigation, population, GDP as mentioned by the first reviewer) was not involved in the original submission. Thus to address this comment, the main and supplementary texts have been revised to compare the nonstationary mode considering TCCCs with the nonstationary mode considering human activity (irrigation, population, GDP), as also stated in the reply to reviewer 1. The process of their change with time has been presented (see Fig. 5 in the revised manuscript). The Pearson correlation coefficients between low-flow series and these indices have been presented (see Fig. 6 in the revised manuscript). The models (M2b, M5, and M6, as described in Table 2 in the revised manuscript) are added. The summary of their results has been presented in Table 5 in the revised manuscript. All new results have been shown in Figs. 7, 8, 9, 10 and 11 in the revised manuscript. Statements in the Results Section have been added and revised, as shown in the revised manuscript with tracked changes.

In addition, there are some mistakes and improper statements in this paper; outlines of methods and results are unclear, and the discussion is weak. It is better for authors to put together contents of results and discussion, and further discuss their results and compared with other related works.

AUTHORS’ REPONSE: Thank you for your comment. The mistakes and improper statements have been carefully checked and corrected in the revised version; to clarify methods, a flow chart of methodology and a table which summarizes the explanatory variables have been added; and we have revised the contents of results and discussion, following the reviewer’s good suggestion. The revised manuscript has included further discussion of results and comparison with other related works as following:

“Overall, the causes of nonstationarity in category for two gauging stations have no clear difference, but have some differences in the relative importance. As shown in Table 5, when modeling the low-flow series of Huaxian using TCCCs variables, the optimal model (M4) preferred the variables are related to recession process; however, for Xianyang, the preferred variables are related to temperature. The reason for this may be that as a downstream station, Huaxian station suffers more intensive human activity, so that the importance of temperature change to the low-flow change is reduced, and meanwhile the importance of streamflow recession (related to the capability of water storage) change is enhanced.” [Lines 493-501]

“The related researches (Jiang et al., 2015; Yang and Yang, 2011; Yang and Yang, 2013; Zhang et al., 2015) have applied the Budyko framework to analyze the impacts of climate change and/or human activity on annual runoff. Indeed, for annual runoff, the Budyko framework is a good method because it used the mean annual water-energy balance equation to consider generation process of total runoff. Unfortunately, to our knowledge, there is a lack of equation derived from basic physics laws for generation process of low flows. Therefore, we emphasize the importance of TCCCs variables to modeling of low-flow nonstationarity.” [Lines 507-513]

Specific Comment The logic of review in the introduction is not smooth. Some references mentioned in the paragraph starting from Line 52, such as Lars Gottschalk's work, were badly concluded and they'd better be put in the next paragraph.

AUTHORS' REPONSE: Thank you for pointing out this and for your good suggestion. To address your comment, we have revised the introduction as mentioned above.

A flow chart of methodology is needed.

AUTHORS' REPONSE: This is a good point. To address your comment, we have added it (Fig. 1 in the revised manuscript).

Line 127 Meaning of this sentence is obscure.

AUTHORS' REPONSE: The sentence has been revised as following: “The distribution type used to build the nonstationary model is outlined”

Further explanation for the selection of 8 candidate variables is needed.

AUTHORS' REPONSE: Thank you for the comment. To address this comment, the 1st paragraph of ‘Section 2.3 Candidate explanatory variables’ has been revised. And the reason for the selection of 8 candiadte variables has been listed in Table 3 in the revised manuscript.

Indices more related to irrigation, like irrigation area, need to be considered, since (Line278) In the Weihe basin, the impacts of agricultural irrigation on runoff have been found to be significant.

AUTHORS' REPONSE: Thank you for the comment. Following reviewer's suggestions, we have included this index (irrigation area) as mentioned above.

Both those 8 explanatory variables and data resources can be summarized in two tables.

AUTHORS' REPONSE: This is a good point. To address this comment, we have revised the text and added Table 3 in the revised manuscript.

I don't see much use in Figure 2.

AUTHORS' REPONSE: Thank you for the comment. Following your suggestion, the Figure has been deleted in the revised manuscript.

Why do you need to study all the series from AM1, 7, 15 to 30?

AUTHORS' REPONSE: The main reason for including four series is to investigate whether the time scale of the series affects the nonstationary mode. As shown in Figure 7 (the revised manuscript), the effect of time scale is existed but limited. In response to this good comment, we have revised the part of the Multiple Covariate Analysis Section to focus on the AM_{30} series.

In some subplans in Figure 8, AIC of either M2 or M3 is worse than M1. What is the probable cause? The conclusion in Line 391 cannot be directly generated from Figure8.

AUTHORS' REPONSE: We have explained that this phenomenon mainly appears in the AM_1 and AM_7 series. AM_1 and AM_7 series are more vulnerable, which means that multiple causes can affect them. The nonstationary mode with one or two physical explanatory variables (M2 or M3) cannot work well for AM_1 and AM_7 . However, the overall decreased trend caused by multiple factors is consistent with the nonstationary mode with time (M1).

What is the impact of location difference on the different AIC results in two stations? Needs to add discussion.

AUTHORS' REPONSE: This is a good point. Following the reviewer's suggestion, we have added a supplemental part to the discussion section (see lines 482-489 in the revised manuscript).

The standard of selecting M4 variables with stepwise selection method needs to be further clarified.

AUTHORS' REPONSE: Following the reviewer's suggestion, we have clarified the standard of the models' variables using Fig. 1 in the revised manuscript.

Table5 and 6 can be merged into one table.

AUTHORS' REPONSE: Agree, corrected.

Formula 2, no need to put "i=" on the top

AUTHORS' REPONSE: Corrected.

Table2, add explanation for parameters down below the table

AUTHORS' REPONSE: Corrected.

The definition, reason of selection, and formula of 8 indices should be listed in a table.

AUTHORS' REPONSE: Corrected.

Line228, 234, 242 add blank space before the paragraph (need to check in the whole paper)

AUTHORS' REPONSE: Corrected.

Line298 slash tag between “n” and “day” is missing (check the whole paper)

AUTHORS' REPONSE: Corrected.

Line304 mistake in time tense

AUTHORS' REPONSE: Corrected.

Line388 incomplete sentence

AUTHORS' REPONSE: Corrected.

Figure1 mark the location of Weihe in the map of China with a rectangular frame

AUTHORS' REPONSE: Corrected.

Figure3 adding R

AUTHORS' REPONSE: Corrected.

Figure 3 &4 lines are too thick

AUTHORS' REPONSE: Corrected.

Figure5&6 differences among colors are too delicate to be seen

AUTHORS' REPONSE: Corrected.

Table 3 &4 add division lines among rows of different stations

AUTHORS' REPONSE: Corrected.

Mistake in references, year of “Bivariate frequency analysis of nonstationary low-flow series based on the time-varying copula” was 2015

AUTHORS' REPONSE: Thank you for pointing out this. We have corrected this.

References

Jiang, C., Xiong, L., Wang, D., Liu, P., Guo, S., and Xu, C.-Y.: Separating the impacts of climate change and human activities on runoff using the Budyko-type equations with time-varying parameters, *Journal of Hydrology*, 522, 326-338, 2015.

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Thanks again to the editor and the two reviewers for providing professional and insightful comments and advices which have significantly improved the revised version of the manuscript.

Sincerely,

Bin Xiong, Lihua Xiong, Jie Chen, Chong-Yu Xu, Lingqi Li

Multiple Causes of Nonstationarity in the Weihe Annual Low Flow Series

Bin Xiong¹, Lihua Xiong^{1*}, Jie Chen¹, Chong-Yu Xu^{1,2}, Lingqi Li¹

1 State Key Laboratory of Water Resources and Hydropower Engineering Science, Wuhan University, Wuhan 430072, P.R. China

2 Department of Geosciences, University of Oslo, P.O. Box 1022 Blindern, N-0315 Oslo, Norway

** Corresponding author:*

Lihua Xiong, PhD, Professor

State Key Laboratory of Water Resources and Hydropower Engineering Science

Wuhan University, Wuhan 430072, P.R. China

E-mail: xionglh@whu.edu.cn

Telephone: +86-13871078660

Fax: +86-27-68773568

16 Abstract:

17 Under the background of global climate change and local anthropogenic activities, multiple
18 driving forces have introduced ~~a variety of~~ various non-stationary components into low-flow series.
19 This has led to a high demand on low-flow frequency analysis that considers nonstationary
20 conditions for modeling. In this study, through a nonstationary frequency analysis framework ~~of~~
21 ~~low-flow frequency analysis has been developed on basis of~~ with the Generalized Linear Model
22 (GLM) to consider time-varying distribution parameters. ~~In GLMs, the candidate multiple~~
23 explanatory variables were ~~as~~ incorporated to explain ~~the time-varying the variation in low-flow~~
24 distribution parameters. These variables are comprised of the three indices of human activities (i.e.,
25 population POP, irrigation area IAR, and gross domestic product GDP) and the eight measuring
26 indices of the climate and catchment conditions ~~in low-flow generation~~, (i.e., total precipitation P ,
27 mean frequency of precipitation events λ , temperature T , potential evapotranspiration ET , climate
28 aridity index AI_{ET} , base-flow index BFI , recession constant K and the recession-related aridity
29 index AI_K). This framework was applied to model the annual minimum flow series of both
30 Huaxian and Xianyang gauging stations in the Weihe River, China. ~~Stepwise regression analysis~~
31 ~~was performed to obtain the best subset of those candidate explanatory variables for the final~~
32 ~~optimum model.~~ The results from stepwise regression for the optimal explanatory variables show
33 that ~~the inter-annual variability in the variables of those selected best subsets~~ the variables

related to irrigation, recession, temperature and precipitation plays an important role in modeling ~~annual low flow series~~. Specifically, analysis of annual minimum 30-day flow in Huaxian shows that AI_K is of the highest relative importance among the optimal variables, followed by IAR , BFI and AI_{ET} , and nonstationary GA distribution model with these optimal variables ~~has~~ an AIC value of 207.0, while the AIC values of other models just with AI_K or time as explanatory variables or without any variable are 217.4, 225.5, 232.3, respectively. ~~AI_K is of the highest relative importance among the best subset of eight candidates, followed by BFI and AI_{ET} .~~ We conclude that ~~The~~ the incorporation of multiple indices related to low-flow generation permits tracing various driving forces. The established link in nonstationary analysis will be beneficial to ~~predict~~ analyze future occurrences of low-flow extremes in similar areas.

Keywords: Climate Change; Streamflow Recession; Multiple Factors; Nonstationarity; Low-flow Frequency Analysis;

1. Introduction

Low flow is defined as the ‘flow of water in a stream during prolonged dry weather’ (WMO, 1974). Yu et al. (2014) quantitatively described a low flow event as a segment of hydrograph during a period of dry weather with discharge values below a preset (relatively small) threshold. According to WMO (2009), annual minimum flows averaged over several days can be used to

measure low flows. During low-flow periods, the magnitude of river flow will greatly restrict its various functions (e.g. providing water supply for production and living, diluting waste water, ensuring navigation, meeting ecological water requirement). Therefore, ~~the~~ investigation of the magnitude and frequency of low flows is of primary importance for engineering design and water resources management (Smakhtin, 2001). ~~For~~ In recent years, low flows, as an important part of river flow regime, have been attracting ~~the an~~ increasing attentions of hydrologists and ecologists, ~~due to in the context of~~ the significant impacts of climate change and human activities ~~on most functions (e.g. providing water supply for production and living, diluting waste water, ensuring navigation, meeting ecological water requirement) of river flow during low flow periods.~~ (Bradford and Heinonen, 2008; Du et al., 2015; Kam and Sheffield, 2015; Kormos et al., 2016; Liu et al., 2015; Sadri et al., 2015; Smakhtin, 2001; WMO, 2009). In general, under the impact of a changing environment, combinations of multiple factors, such as precipitation change, temperature change, irrigation area change and construction of reservoirs, can drive various patterns of streamflow changes (Liu et al., 2017; Tang et al., 2015). Unfortunately, when subjected to a variety of influencing forces, low flow is more vulnerable than high flow or mean flow. Therefore, it is a pretty important issue in hydrology to identify low-flow changes, track multiple driving factors and quantify their contributions from the perspective of hydrological frequency analysis.

In hydrological analysis and design, conventional frequency analysis estimates the statistics of a hydrological time series based on recorded data with the stationary hypothesis which means that this series is “free of trends, shifts, or periodicity (cyclicality)” (Salas, 1993). However, global warming and human forces have changed climate and catchment conditions in some regions. Time-varying climate and catchment conditions can affect all aspects of the flow regime, i.e. changing the frequency and magnitude of floods, altering flow seasonality, and modifying the characteristics of low flows, etc. The hypothesis of stationarity has been suspected (Milly et al., 2008). If this problematic method is still used, the frequency analysis may lead to high estimation error and costly in hydrological design. Therefore, considerable literatures have introduced the concept of hydrologic nonstationarity into analysis of various hydrological variables, such as annual runoff (Arora, 2002; Jiang et al., 2017; Jiang et al., 2015; Liu et al., 2017; Xiong et al., 2014; Yang and Yang, 2013), flood (Chen et al., 2013; Gilroy and Mccuen, 2012; Gu et al., 2016; Kwon et al., 2008; López and Francés, 2013; Tang et al., 2015; Xiong et al., 2015b; Yan et al., 2016; Zhang et al., 2014; Zhang et al., 2015), low flow (Du et al., 2015; Jiang et al., 2014; Liu et al., 2015), precipitation (Cheng and AghaKouchak, 2014; Gu et al., 2017a, b, c; Mondal and Mujumdar, 2015; Villarini et al., 2010) and so on. Compared with the literatures on annual runoff, floods and precipitation, the literatures on the nonstationary analysis of low flow are very relatively limited.

88 Previous hydrological literatures on frequency analysis of nonstationary ~~low-flow~~
89 hydrological series mainly focus on two aspects: development of nonstationary method and
90 exploration of covariates reflecting changing environments. Strupczewski et al. (2001) presented
91 the method of time-varying moment which assumes that the hydrological variable of interest obeys
92 a certain distribution type, but its moments change over time. The method of time-varying moment
93 was modified to be the method of time-varying parameter values for the distribution representative
94 of hydrologic data (Richard et al., 2002). Villarini et al. (2009) presented this method using the
95 Generalized Additive Models for Location, Scale, and Shape Parameters (GAMLSS) (Rigby and
96 Stasinopoulos, 2005), a flexible framework to assess nonstationary time series. The time-varying
97 parameter method can be extended to the physical covariate analysis by replacing time with any
98 others physical covariates (Du et al., 2015; Jiang et al., 2014; Kwon et al., 2008; López and
99 Francés, 2013; Liu et al., 2015; Villarini et al., 2010; Villarini and Strong, 2014). For example,
100 Jiang et al. (2014) used reservoir index as an explanatory variables based on the time-varying
101 copula method for bivariate frequency analysis of nonstationary low-flow series in Hanjiang River,
102 China. Du et al. (2015) took precipitation and air temperature as the explanatory variables to
103 explain the inter-annual variability in low flows of Weihe River, China. Liu et al. (2015) took Sea
104 Surface Temperature in Nino3 region, the Pacific Decadal Oscillation, the sunspot number (3 years
105 ahead), the winter areal temperature and precipitation as the candidate explanatory variables to

106 explain the inter-annual variability in low flows of Yichang station, China. Kam and Sheffield
107 (2015) ascribed the increasing inter-annual variability of low flows over the eastern United States
108 to North Atlantic Oscillation and Pacific North America.

109 ~~Low flows are more vulnerable to influences of climate change and human activities than~~
110 ~~high flows. However, To our knowledge,~~ compared with the nonstationary flood frequency
111 analysis, the studies on the nonstationary frequency analysis of low-flow series is not very
112 extensive because of incomplete knowledge of low flow generation (Smakhtin, 2001). Most of
113 these studies explain nonstationarity of low-flow series only by using climatic indicators or a
114 single indicator of human activity. However, the indicators of catchment conditions (e.g. recession
115 rate) related to physical hydrological processes have seldom been attached in nonstationary
116 modeling of low flow series. ~~This leads to lack of linking with hydrological process, which in~~
117 ~~turn would exclude further analysis, such as accurately tracing origins of change in low flow~~
118 ~~series~~ This lack of linking with hydrological processes makes it impossible to accurately quantify
119 the contributions of influencing factors for the nonstationarity of low flow series, and such a
120 scientific demand for tracing the sources of nonstationarity of low-flow series and qualifying their
121 contributions motivated the present study. The knowledge of low-flow generation has been
122 increased by efforts of hydrologists, which can help develop physical covariates to address
123 nonstationarity. Low flows generally originate from groundwater or other delayed outflows

124 (Smakhtin, 2001; Tallaksen, 1995). Their generation relates to both an extended dry weather
125 period (leading to a climatic water deficit) and complex hydrological processes which determine
126 how these deficits propagate through the vegetation, soil and groundwater system to streamflow
127 (WMO, 2009). Thus, not only climate conditions drivers (e.g. potential evaporation exceeds
128 precipitation), but also catchment conditions drivers (e.g. the faster hydrologic response rate to
129 precipitation) can cause low flows.

130 The significant factors such as precipitation, temperature, evapotranspiration, streamflow
131 recession, large-scale teleconnections and human forces may play important roles in influencing
132 low-flow generation (Botter et al., 2013; Giuntoli et al., 2013; Gottschalk et al., 2013; Jones et al.,
133 2006; Kormos et al., 2016; Roderick et al., 2013; Sadri et al., 2015). Gottschalk et al. (2013)
134 presented a derived low flow probability distribution function with climate and catchment
135 characteristics parameters (i.e., the mean length of dry spells λ^{-1} and recession constant of
136 streamflow K) as its distribution parameters. Botter et al. (2013) derived “a measurable index”
137 (λ^{-1}/K) which can be used for discriminating erratic river flow regimes from persistent river flow
138 regimes. Recently, in- Van Loon and Laaha (2015) used climate and catchment characteristics (e.g.
139 the duration of dry spells in precipitation and the base flow index) to explain the duration and
140 deficit of hydrological drought event and offered a further understanding of low-flow generation.
141 These studies indicated that climate and catchment conditions play an important role in producing

low flows.

The goal of this study is to trace origins of nonstationarity in low flows through developing a nonstationary low-flow frequency analysis framework with the consideration of the time-varying climate and catchment conditions (TCCCs) and human activity (HA). ~~The goal of this study is to develop a nonstationary low flow frequency analysis framework with the consideration of the time varying climate and catchment conditions (TCCCs).~~ In this framework, the climate and catchment conditions are quantified using the eight indices, i.e., meteorological variables (total precipitation P , mean frequency of precipitation events λ , temperature T and potential evapotranspiration ET), basin storage characteristics (base-flow index BFI , recession constant K) and aridity indexes (climate aridity index AI_{ET} , the recession-related aridity index AI_K).

The specific objectives of this study are: (1) to find the most important index to explain the nonstationarity of low-flow series; (2) to determine the best subset of TCCCs indices and/or human activity indices (i.e., population POP , irrigation area IAR , and gross domestic product GDP) for final model through stepwise selection method to identify nonstationary mode of low-flow series; and (3) to quantify the contribution of selected explanatory variables to the nonstationarity.

~~The non-stationary frequency analysis with TCCCs developed in this study is able to give the trace of nonstationary low flow drivers and to estimate the contribution of each driver to the change in low flow series.~~

This paper is organized as follows. Section 2 describes the methods. ~~We describe t~~The Weihe River basin and available data sets used in this study are described in Section 3, followed by a presentation of the results and discussion in Section 4. Section 5 summarizes the main conclusions.

2 Methodology

The flowchart of how to organize the nonstationary low-flow frequency analysis framework is shown in Fig. 1. The whole process is divided into three steps. The first step is preliminary analysis, including the graphical presentation of both explanatory variables and low-flow series, the statistical test for nonstationarity and the correlations between each explanatory variable and each low-flow series. The second step is single covariate analysis for the most important explanatory variable. The third step is multiple covariate analysis for the optimal combination. We use a low-flow frequency analysis model and stepwise regression method to accomplish the last two steps. In ~~this~~the following sub-sections, first, the low-flow frequency analysis model is constructed based on the nonstationary probability distributions method, in which distribution parameters serving as response variables can vary as functions of explanatory variables. Second, ~~the candidate distributions are described to determine the different types of nonstationary frequency curves.~~the distribution types used to build the nonstationary model are outlined. Then, the ~~eight~~ candidate explanatory variables related to the time-varying climate and catchment

~~conditions (TCCCs) and human activity (HA) are clarified~~
~~presented to incorporate time-varying~~
~~climate and catchment conditions (TCCCs) into distribution models for the nonstationary~~
~~frequency analysis.~~ Finally, estimation of model parameters and selection of models are illustrated.

<Figure 1>

2.1 Construction of the low-flow nonstationary frequency analysis model

Generally, a nonstationary frequency analysis model can be established based on the time-varying distribution parameters method (Du et al., 2015; López and Francés, 2013; Liu et al., 2015; Richard et al., 2002; Villarini and Strong, 2014). For the nonstationary probability distribution $f_Y(Y_t|\theta^t)$, let Y_t be a random variable at time t ($t=1,2,...,N$) and vector $\theta^t=[\theta_1^t,\theta_2^t,...,\theta_m^t]$ be the time-varying parameters. The number of parameters m in hydrological frequency analysis is generally limited to three or less. The function relationship between the k^{th} parameter θ_k^t and the multiple explanatory variables is expressed as follows:

$$g_k(\theta_k^t)=h_k(x_1^t,x_2^t,...,x_n^t) \quad (1)$$

where $x_1^t,x_2^t,...,x_n^t$ are explanatory variables; n is the number of explanatory variables; $g_k(\cdot)$ is the link function which ensures the compliance with restrictions on the sample space and is usually set to natural logarithm for the given negative predictions; $h_k(\cdot)$ is the function for nonstationary modeling. The theory of Generalized Linear Model (Dobson and Barnett, 2012) is used to build function relationships between distribution parameters and their explanatory

195 variables. In GLMs, the response relationship can be generally expressed as

$$196 \quad g_k(\theta'_k) = \alpha_{0k} + \sum_{i=1} \alpha_{ik} x'_i \quad (2)$$

197 where α_{ik} ($i = 0, 1, 2, \dots, n, k = 1, \dots, m$) are the GLM parameters.

198 In order to ~~give a further nonstationary analysis~~ compare the nonstationary models constructed
 199 by various combinations of explanatory variables, Eq. (2) is modified in this study using
 200 dimensionless method for the standard GLM parameters. The value of θ'_k could be assumed to be
 201 equal to its mean ($\bar{\theta}_k$) when all explanatory variables are equal to their mean (\bar{x}_i), i.e.,

$$202 \quad \theta'_k(x'_1 = \bar{x}_1, x'_2 = \bar{x}_2, \dots, x'_n = \bar{x}_n) = \bar{\theta}_k \quad (3)$$

203 Eq. (2) is then modified as

$$204 \quad g_k\left(\frac{\theta'_k}{\bar{\theta}_k}\right) = \beta_{0k} + \sum_{i=1}^{i=n} \beta_{ik} z'_i$$

$$z'_i = \frac{x'_i - \bar{x}_i}{s_i}, \quad i = 1, 2, \dots, n \quad (4)$$

$$\beta_{0k} = g_k\left(\frac{\theta'_k}{\bar{\theta}_k} \middle| \theta'_k = \bar{\theta}_k\right) = g_k(1)$$

205 where z'_i is normalized explanatory variables; s_i is the standard deviation of x'_i ;
 206 β_{ik} ($i = 1, 2, \dots, n, k = 1, \dots, m$) are the standard GLM parameters. Let the link function $g_k(\cdot)$ be the
 207 natural logarithmic function $\ln(\cdot)$ and θ'_l be the distribution parameter in $[\theta'_1, \theta'_2, \dots, \theta'_m]$ with
 208 most significant change, the degree of nonstationarity in low flow series can be defined as
 209 $\ln(\theta'_l) - \ln(\bar{\theta}_l)$. Then, the contribution c'_i of each explanatory variable x'_i to $\ln(\theta'_l) - \ln(\bar{\theta}_l)$

210 could be defined as

211
$$c_i^t = \beta_{it} \frac{x_i^t - \bar{x}_i}{s_i} \quad (5)$$

212 **2.2 Candidate distribution functions**

213 We need to select the form of probability distribution $f_Y(\cdot)$ to determine what type of
214 nonstationary frequency curves will be produced. Various probability distributions have been
215 compared or suggested in modeling of low-flow series (Du et al., 2015; Hewa et al., 2007; Liu et
216 al., 2015; Matalas, 1963; Smakhtin, 2001). An extensive overview of distribution functions for low
217 flow is given in Tallaksen et al. (2004). Following these recommendations, we consider five
218 distributions, i.e. Pearson-III (PIII), Gamma (GA), Weibull (WEI), Lognormal (LOGNO) and
219 Generalized Extremes Value (GEV) as candidates in this study (Table 1). In the case of Pearson-III
220 distribution, considering that the parameter θ_3 of Pearson-III as lower bound should approach
221 zero and the parameter θ_3 of GEV is quite sensitive and difficult to be estimated, we assume
222 them to be constant in this study.

223 **2.3 Candidate explanatory variables**

224 We look for variables $x_1^t, x_2^t, \dots, x_n^t$ that can explain parts of the variations in distribution
225 parameters θ^t . From the perspective of low-flow generation, the dependency between low-flow
226 regime and both climate and catchment conditions has been presented by previous studies (Botter

et al., 2013; Gottschalk et al., 2013; Van Loon and Laaha, 2015). We focus on eight measuring indices: total precipitation, mean frequency of precipitation events, temperature, potential evapotranspiration, climate aridity index, base-flow index, recession constant and recession-related aridity index. These indices were chosen to incorporate time-varying climate and catchment conditions (TCCCs) in nonstationary modeling, of low-flow frequency and serving as candidate explanatory variables. The values of them at each year could be estimated from hydro-meteorological data. Annual precipitation (P) and temperature (T) are calculated directly by meteorological data. The remaining TCCCs indices need to be estimated indirectly. Detailed estimation procedures are shown ~~as follows~~ in following subsections. In addition to TCCCs indices, the three indices of human activity (irrigation area, population and gross domestic product) are included, and the reasons for selecting all indices are summarized in Table 2.

2.3.1. Annual mean frequency of precipitation events (λ)

Annual mean frequency of precipitation events is defined as an index to represent the intensity of precipitation recharge to the streamflow:

$$\lambda = \frac{1}{W} \sum_{w=1}^{w=W} \frac{N_w(A)}{t_r} \quad (6)$$

where $N_w(A)$ is the number of daily rainfall events A (with values more than the threshold 0.5 mm) in w^{th} windows with a length t_r ; W is the number of windows.

244 **2.3.2. Annual climate aridity index (AI_{ET})**

245 The ratio of annual potential evaporation to precipitation, commonly known as the climate
246 aridity index, has been used to assess the impacts of climate change on annual runoff (Arora, 2002;
247 Jiang et al., 2015). The climate aridity index largely reflects the climatic regimes in a region and
248 determines runoff rates (Arora, 2002). Therefore, we choose the annual climate aridity index as a
249 measure of time-varying climate and catchment conditions and estimate its value in a whole region
250 using

$$251 \quad AI_{ET} = \frac{ET}{P} \quad (7)$$

252 where P is annual areal precipitation (mm); ET is annual areal potential evapotranspiration.
253 The Hargreaves equation (Hargreaves and Samani, 1985) is applied to calculate ET using the
254 R-package ‘Evapotranspiration’ (Guo, 2014).

255 **2.3.3. Annual base-flow index (BFI)**

256 The base flow index (BFI) is defined as the ratio of base flow to total flow. This index has
257 been applied to quantify catchment conditions (e.g. soil, geology and storage-related descriptors)
258 to explain hydrological drought severity (Van Loon and Laaha, 2015). We also choose annual base
259 flow index (BFI) as a measure of TCCCs. BFI is estimated using a hydrograph separation
260 procedure in R-package ‘lfstat’ (Koffler and Laaha, 2013).

261 2.3.4. Annual streamflow recession constant (K)

262 Recession constant is an important catchment characteristic index measuring the time scale of
263 the hydrological response and reflecting water retention ability in the upstream catchment (Botter
264 et al., 2013). Various estimation methods have been developed to extract recession segments and to
265 parameterize characteristic recession behavior of a catchment (Hall, 1968; Sawaske and Freyberg,
266 2014; Tallaksen, 1995).

267 In this study, annual recession analysis (ARA) is performed to obtain annual streamflow
268 recession constant (K). In ARA, the linearized Duperoy-Boussinesq equation is used to parameterize
269 characteristic recession behavior of a catchment and is written as

$$270 -\frac{dQ_t}{dt} = \frac{1}{K} Q_t \quad (8)$$

271 where Q_t is the value at time t . Eq. (8) is investigated by plotting data points $\frac{dQ_t}{dt}$ against Q_t
272 of all extracted recession segments from hydrographs at each year. The criteria of recession
273 segments extraction ~~is~~are based on the Manual on Low-flow Estimation and Prediction (WMO,
274 2009). Then, the annual recession rate (K^{-1}) is estimated as the slope of fitted straight line of these
275 data points with least square method. We calculated K using R-package 'lfstat' (Koffler and
276 Laaha, 2013).

277 2.3.5. Annual recession-related aridity index (AI_K)

278 In this study, recession-related aridity index is defined as the ratio of recession rate (K^{-1}) to

279 mean precipitation frequency (λ), denoted as

280
$$AI_K = \frac{K^{-1}}{\lambda} \quad (9)$$

281 This ratio plays an important role in controlling on-river flow regime (Botter et al., 2013;
282 Gottschalk et al., 2013) and serves as an indicator measuring the recession-related aridity degree of
283 the streamflow in river channel. For example, faster recession process or lower precipitation
284 frequency may lead to increased runoff loss or decreased precipitation supply. Consequently, the
285 higher the value AI_K is, the more likely low flow events occur, and vice versa.

286 2.4 Parameter estimation

287 The model parameters including $\bar{\theta}_k (k=1,2,...,m)$ and $\beta_{ik} (i=1,2,...,n,k=1,...,m)$ are
288 estimated. $\bar{\theta}_k (k=1,2,...,m)$ are estimated from outputs of stationary frequency analysis through
289 maximum likelihood method. We have

290
$$L(\bar{\theta}_1, \bar{\theta}_2, ..., \bar{\theta}_m) = \sum_{t=1}^{t=N} \ln \left[f_Y(y_t | \bar{\theta}_1, \bar{\theta}_2, ..., \bar{\theta}_m) \right] \quad (10)$$

291 where y_t is observed low flow at time t ; N is the number of samples. The parameters
292 $\beta_{ik} (i=1,2,...,n,k=1,...,m)$ are estimated through maximum likelihood method to produce
293 nonstationary low-flow frequency curves:

$$L \begin{pmatrix} \beta_{11}, \dots, \beta_{n1} \\ \dots \\ \beta_{1m}, \dots, \beta_{nm} \end{pmatrix} = \sum_{t=1}^{t=N} \ln \left\{ f_Y \left(y_t \mid \theta_1' \left(z_1^t, \dots, z_n^t \mid \beta_{11}, \dots, \beta_{n1} \right), \dots, \theta_m' \left(z_1^t, \dots, z_n^t \mid \beta_{1m}, \dots, \beta_{nm} \right) \right) \right\} \quad (11)$$

The residuals (normalized randomized quintile residuals) are used to test the goodness-of-fit of fitted model objects (Dunn and Symth, 1996):

$$\hat{r}_t = \Phi^{-1} \left(F_Y \left(y_t \mid \hat{\theta}' \right) \right) \quad (12)$$

where $F_Y(\cdot)$ is the cumulative distribution of y_t ; $\Phi^{-1}(\cdot)$ is the inverse function of the standard normal distribution. The distribution of the true residuals \hat{r}_t converges to standard normal if the fitted model is correct. Worm plot (Buuren and Fredriks, 2001) is used to check whether \hat{r}_t have a standard normal distribution.

2.5 Model selection

Model selection contains the selection of the type of probability distribution and the selection of the explanatory variables to explain the response variables (i.e., distribution parameters θ_1 and θ_2). In order to obtain the final optimal model, the selection of the explanatory variables for θ_1 and θ_2 is conducted by a stepwise selection strategies (Stasinopoulos and Rigby, 2007; Venables, 2002): [i.e.](#) select a best subset of candidate explanatory variables for θ_1 using a forward approach (which starts with no explanatory variable in the model and tests the addition of each explanatory variable using a chosen model fit criterion); given this subset for θ_1 select another subset for θ_2 (forward). The stepwise selection strategies can get a series of stepwise models with different

311 numbers of explanatory variables, as shown in Fig1. In order to detect how the number of
 312 explanatory variables influences the performance of the model for describing non-stationarity, we
 313 investigate the ~~five-eight~~ types of stepwise models as shown in Table 3: the zero-covariate model
 314 or stationary model (M0), the time covariate model (M1), single physical covariate model M2
 315 (single TCCCs covariate model M2a or single HA covariate model M2b), ~~the double physical~~two
 316 TCCCs covariates model (M3) ~~and,~~ the optimal ~~number physical~~TCCCs covariates model (M4),
 317 ~~as shown in Table 2~~ the optimal HA covariates model (M5) and the final model (M6). The model
 318 fit criterion is based on the Akaike's information criterion (Akaike, 1974) as shown by the
 319 following

$$320 \quad AIC = -2ML + 2df \quad (13)$$

321 where ML is the log-likelihood in Eq. (11) and df is the number of degrees of freedom. The
 322 model with the lower AIC value was considered better.

323 3. Study Area and Data

324 3.1. The study area

325 The Weihe River, located in the southeast of the Northwest Loess Plateau, is the largest
 326 tributary of the Yellow River, China. The Weihe River has a drainage area of 134 766 km²,
 327 covering the coordinates of 33°42'-37°20'N 104°18'-110°37'E (Fig. 42). This catchment
 328 generally has a semi-arid climate, with extensive ~~sub-humid~~ continental monsoonal influence.

329 Average annual precipitation of the whole area over the period 1954-2009 is about 540 mm, and
330 has a wide range (400-1000 mm) in various regions. Under the significant impacts of climate
331 change and human activities in the Weihe River basin in recent decades, the hydrological regime
332 of the river has changed over time (Du et al., 2015; Jiang et al., 2015; Xiong et al., 2015a).

333 <Figure ~~1~~2>

334 In the Weihe basin, the impacts of agricultural irrigation on runoff have been found to be
335 significant (Jiang et al., 2015; Lin et al., 2012). Lin et al. (2012) mentioned that the annual runoff
336 of the Weihe River was significantly affected by irrigation diversion of the Baoji Gorge irrigation
337 area. The irrigated area of Baoji Gorge Irrigation Area increased over time since the founding of
338 P.R. China in 1949, and due to one influential irrigation system project in that area, it became more
339 than twice of the original ~~one~~one-irrigation area since 1971. Jiang et al. (2015) demonstrated that in
340 the Weihe basin, irrigated area, as compared with the other indices e.g. population, gross domestic
341 product and cultivated land area, was a more suitable human explanatory variable for explaining
342 the time-varying behavior of annual runoff. Within the above background, it is important to
343 considering the effects of human activities that mainly originate from irrigation diversion, and
344 especially for studying low flow series in this basin. In this study, we use the available data
345 (1980-2005) of the irrigation diversion system on plateau in Baoji Gorge Irrigation Area in Zhang
346 (2008) to provide some information for the knowledge of low flow generation. The estimations of

347 annual recession rate (K^{-1}) by the daily streamflow data are expected to incorporate the
348 information of impacts of water diversions on the low flows in the river channel.

349 3.2. ~~Streamflow d~~Data

350 We used daily streamflow records (1954-2009) provided by the Hydrology Bureau of the
351 Yellow River Conservancy Commission from both Huaxian station (with a drainage area of 106
352 500 km²) and Xianyang station (with a drainage area of 46 480 km²). Low-flow extreme events
353 were selected from the daily streamflow series using the widely-used annual minimum series
354 method (WMO, 2009). AM_n is the annual minimum ~~n~~n-day flow during hydrological year
355 ~~defined to start beginningt~~ on 1 March. Consequently, AM_1 , AM_7 , AM_{15} and AM_{30} are selected as
356 low-flow extreme events in this study. The original measure unit of streamflow data ($\text{m}^3 \cdot \text{s}^{-1}$) is
357 converted to $10^{-4} \text{ m}^3 \cdot \text{s}^{-1} \cdot \text{km}^{-2}$ ~~by dividing by the corresponding drainage area (km^2)~~ for
358 convenience of comparison of results between the Huaxian and Xianyang gauging stations

359 ~~3.3. Precipitation and temperature data~~

360 We downloaded~~ed~~ daily total precipitation and daily mean air temperature records for 19
361 meteorological stations over the basin from the National Climate Center of the China
362 Meteorological Administration (source: <http://cdc.cma.gov.cn>). The areal average daily series of
363 both variables above Huaxian and Xianyang stations are calculated using the Thiessen polygon
364 method (Szolgayova et al., 2014; Thiessen, 1911). The annual average temperature (T) and annual

total precipitation (P) over the period 1954-2009 are calculated for each catchment.

Human activity data (i.e. gross domestic product, population and irrigation area) were taken from annals of statistics provided by the Shaanxi Provincial Bureau of Statistics (<http://www.shaanxitj.gov.cn/>) and Gansu Provincial Bureau of Statistics (source: <http://www.gstj.gov.cn/>).

4. Results and discussion

4.1. Identification of nonstationarity

~~Figure 2 shows that the Weihe River basin is characterized by a warm and humid summer (June, July, and August) with low ratio of irrigated diversion, and by a cold and dry winter (December, January, and February) with high ratio of irrigated diversion. The majority of the low flow events in this basin occur in these two seasons and show a bimodal frequency distributions of occurrence with two peaks in February and June, respectively (Fig. 2a). This result implies that the generation of low flows may be influenced by more than one factor such as high ratio of irrigated diversion, high air temperature or lack of precipitation.~~

◁Figure 2▷

Graphical representation and statistical test provide a preliminary analysis for low-flow nonstationarity. The graphical representations of time-series data help visualize the trends of

related variables (i.e. low-flow, TCCCs and HA variables), the density distributions of TCCCs
variables and the correlations between low-flow variables and these explanatory variables. In Fig.
3. Overall~~overall~~, four annual minimum streamflow series (AM_1 , AM_7 , AM_{15} and AM_{30}) in
both Huaxian and Xianyang gauging stations show overall decreasing trends, as indicated by the
fitted (dashed) trend lines ~~in Fig. 3~~. Compared with Huaxian, Xianyang has a larger runoff
modulus (the flow per square kilometer) and a larger decrease in annual minimum streamflow
series. For example, the decline slope of AM_{30} is $-0.0725 (10^{-4} \text{ m}^3 \cdot \text{s}^{-1} \cdot \text{km}^2/\text{yr})$ in Huaxian
station ~~which while Xianyang station it is larger than~~ $-0.1338 (10^{-4} \text{ m}^3 \cdot \text{s}^{-1} \cdot \text{km}^2/\text{yr})$ ~~in Xianyang~~
~~station.~~

<Figure 3>

Figure 4 shows the kernel density estimations and time processes of ~~the eight candidate~~
~~explanatory TCCCs~~ variables ~~(Sect. 2.3) reflecting the TCCCs~~ for both Huaxian (H) and Xianyang
(X) stations. The results show that these variables have different variation patterns. For example,
the mean frequency of precipitation events (λ) has a decreasing trend, while temperature (T) has
an increasing trend. As presented by Fig. 5, three HA variables have a significant upward trend,
especially the irrigation area IAR which is increased greatly after about 1970, suggesting that the
impact of human activities in this basin has increased over time.

<Figure 4>

<Figure 5>

The significance of trends in the four annual minimum streamflow series and ~~eight~~ explanatory TCCCs variables is tested by the Mann-Kendall trend test (Kendall, 1975; Mann, 1945; Yue et al., 2002), and the change -points in these series are detected by the Pettitt's test (Pettitt, 1979). The results in Table 3-4 show that in both Huaxian and Xianyang stations, the decreasing trends in all the four low-flow series (AM_1 , AM_7 , AM_{15} and AM_{30}) and two explanatory variables (λ and P), and the increasing trends in T , ET , and AI_{ET} are significant at the 0.05 level (Table 34), but BFI shows no significant trends. However, K and AI_K had significantly decreasing trends only in Huaxian station ($p\text{-value} < 0.05$). The results of change-point detection show that all low-flow series are located at 1968-1971 ($p\text{-value} < 0.05$) except AM_{30} at Xianyang station whose change point is located at 1993 ($p\text{-value} < 0.05$); for the eight candidate explanatory variables, the change points of the variables related to temperature (T , ET , AI_{ET}) in both stations are located at 1990-1993 ($p\text{-value} < 0.05$), the change points of the variables related to precipitation (λ , P) in both stations are close at 1984-1990 ($p\text{-value} \leq 0.186$) and the change points of the variables related to streamflow recession (K , AI_K) in Huaxian station are located at 1968-1971 ($p\text{-value} < 0.05$). However, BFI in both stations and K and, AI_K in Xianyang

417 station show no significant change points.

418 A preliminary attribution analysis is performed using the Pearson correlation matrix to
419 investigate the relations between the annual minimum series and eight candidate explanatory
420 variables. Figure ~~65~~ indicates that there are significant linear correlations between the four
421 minimum low-flow series (AM_1 , AM_7 , AM_{15} and AM_{30}) and all the explanatory variables
422 ~~except~~ *GDP*, ~~with~~have the absolute values of Pearson correlation coefficients larger than 0.27
423 ($p\text{-value} < 0.05$). These potential physical causes of nonstationarity in low flows are further
424 considered by establishing low-flow nonstationary model with TCCCs and HA variables in the
425 following section.

426 <Figure ~~56~~>

427 4.2. Nonstationary frequency analysis models

428 4.2.1 Single covariate models

429 Figure ~~76~~ presents the AIC values of the ~~three-four~~ types of models (M0, M1, M2a and
430 M2b~~M2, M1 and M0~~) fitted for the low flow series (AM_1 , AM_7 , AM_{15} and AM_{30}). Some
431 interesting results are shown as follows. First, nonstationary models (M1, M2a and M2b~~M2 and~~
432 ~~M1~~) have lower AIC values than stationary model (M0), which suggests that nonstationary models
433 are worth considering. Second, for Huaxian station, irrespective of the chosen explanatory

434 variables, the distribution type plays an important role in modeling nonstationary low flow series.
 435 For example, PIII, GA and WEI distributions in AM_{15} and AM_{30} ~~most~~ cases have lower AIC
 436 values than LOGNO and GEV distributions. However, for Xianyang, choosing a suitable
 437 explanatory variable may be more important than choosing a distribution type. For example,
 438 variables t , P , T , ~~and~~ AI_{ET} , POP and IAR in most cases have lower AIC values than the
 439 other explanatory variables. Finally, in Huaxian, the lowest AIC values ~~the best M2 models~~ for
 440 modeling AM_1 , AM_7 , AM_{15} and AM_{30} are found in GEV M2b IAR, LOGNO M2b IAR,
 441 PIII M2a AI_K and GA M2a AI_K , respectively ~~are all found in the M2- AI_K model (using AI_K~~
 442 ~~as an explanatory variable)~~; while in Xianyang, the lowest AIC values ~~the best M2 models~~ for
 443 modeling AM_1 , AM_7 , AM_{15} and AM_{30} are ~~all found in the M2- K , M2- AI_{ET} , M2- AI_{ET}~~
 444 ~~and M2- T model~~ GEV M2b IAR, GEV M2b IAR, PIII M2b IAR and GEV M2b IAR,
 445 respectively. These results indicated that ~~in for explaining nonstationarity of low flow in~~ Huaxian
 446 station, IAR is the most dominant HA variables, and AI_K is the most dominant TCCCs variable
 447 ~~causing nonstationarity in AM_1 , AM_7 , AM_{15} and AM_{30}~~ ; while in Xianyang, the most
 448 dominant HA variables is IAR , the most dominant TCCCs variables ~~causing nonstationarity in~~
 449 AM_1 , AM_7 , AM_{15} and AM_{30} are K , AI_{ET} , AI_{ET} and T , respectively. ~~Table 4~~
 450 ~~summarizes the above analysis.~~

451 <Figure 67>

Figure 87 shows the diagnostic assessment of ~~the best M2 model~~ the GA_M2 model (GA_M2 with the optimal explanatory variable) for AM_{30} in both Huaxian and Xianyang stations. The centile curves plots of GA_M2 (Figs. 7a-8a and 7b8b) show the observed values of AM_{30} , the estimated median and the areas between the 5th and 95th centiles. Figure 7a-8a shows the response relationship between AM_{30} and AI_K in Huaxian: the increase of AI_K means the smaller magnitude of low-flow events because a high value of AI_K (faster stream recession or fewer rainy days) may lead to faster water loss or less supply. In Fig. 7b8b, the higher values of IAR_T means the smaller magnitude of low flow events, which suggests that IAR_T plays an important role in driving low-flow generation in Xianyang. Figs 7e-8c and 7d-8d show that the worm points are within the 95% confidence intervals, thereby indicating a good model fit and a reasonable model construction.

<Figure 78>

4.2.2 Multiple covariate models

Figure 8-9 shows ~~that~~ the AIC values of stationary model (M0), time covariate model (M1), physical covariate models (M2a, M2b, M3, M4, M5 and M6) for AM_{30} ~~(M2, M3 and M4 with the corresponding optimal explanatory variables) for AM_1 , AM_7 , AM_{15} and AM_{30} in both Huaxian and Xianyang stations.~~ As shown in Fig. 9, M4 (nonstationary GA distribution with the

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optimal TCCCs variables) has a good performance; after adding the HA variables, M6 with the lowest AIC values is attained; it can be found that the combination of multiple TCCCs variables displays a major role in changing the low flows of Weihe River, but the influence of HA variables shouldn't be ignored. For all low flow series, the lowest AIC values are always found in the M4 models, suggesting that it is necessary to consider multiple explanatory variables for nonstationary modeling.

<Figure 89>

A summary of frequency analysis based on five types of models (M0, M1, M2, M3 and M4) for both Huaxian and Xianyang gauging stations nonstationary GA distribution AM_{30} is presented in Table 5 and Table 6, respectively. We choose to focus on M4, M5 and M6. When only using TCCCs variables to model nonstationary low-flow frequency distribution, the results of M4 show the optimal combination of explanatory variables for all low-flow series contains more than three variables. For example, for AM_{30} of Huaxian, the optimal combination of TCCCs variables includes AI_K , BFI and AI_{ET} . When only using HA variables, the results of M5 show IAR is important to the low flows in this area. And M4 has a better performance than M5. When using both TCCCs variables and HA variables, the results of M6 show the optimal combination contains multiple TCCCs variables and the irrigation area IAR . For Huaxian, the optimal combination of all explanatory variables is AI_K , IAR , BFI and P , while for Xianyang, the optimal

487 combination is IAR , AI_{ET} and BFI . For M4 and M3 models, the relative importance of
 488 selected explanatory variables is identified through the stepwise selection method. For instance,
 489 for AM_{30} in Xianyang (Table 5), temperature (T) with highest relative importance, followed
 490 orderly by P , BFI and K . We can also find that if ~~the candidate~~ two TCCCs variables are
 491 highly correlated, they do not seem to be selected as the explanatory variables at the same time.
 492 For example, ~~one of those variables~~ in terms of ~~only~~ air temperature (T), evapotranspiration (ET)
 493 and the climate aridity index (AI_{ET}), only one of them will appear in the optimal combination-a
 494 ~~best subset of eight candidates in the final optimum model~~. This suggests that multicollinearity
 495 problem in multiple variables analysis can be reduced, which will help obtain more reliable GLMs
 496 parameters for contribution analysis.

497 The diagnostic assessment of the ~~best M4 model (GA_M4)~~ GA_M6 model for AM_{30} at
 498 two stations is presented by Fig. 910. The centile curves plots of GA_M46 (Figs. 910a and 910b)
 499 show the more sophisticated nonstationary modeling than GA_M2 (Fig 78). When using GA_M46
 500 to model AM_{30} in Huaxian (Fig. 9a), similar to GA_M2, the lower low flows are found to also
 501 correspond to higher value of AI_K , but GA_M46 ~~are~~ is able to identify the more complex
 502 variation patterns of low flows through the incorporation of IAR , BFI and P . Figures 910c
 503 and 910d show that the data points of worm plots of GA_M46 are almost within the 95%
 504 confidence intervals, thereby indicating an acceptable model fit and a reasonable model

construction.

<Figure 910>

Figure 10—11 presents the contribution of each selected explanatory variable to $\ln(\theta'_1) - \ln(\bar{\theta}_1)$ in observation year based on GA_M46 for AM_{30} in Huaxian and Xianyang. We can find that for Huaxian, the simulation value of $\ln(\theta'_1)$ frequently occur below $\ln(\bar{\theta}_1)$ during the two periods of about 1970-1982 and 1993-2003, which is in accordance with the observed decrease in AM_{30} of Huaxian station during these periods. In the former period 1970-1982, ~~the largest negative contribution is found in both~~ AI_K ~~and~~ BFI ~~contribute a lot of negative amount to~~ $\ln(\theta'_1) - \ln(\bar{\theta}_1)$, ~~whereas during 1993-2003, the contribution of both~~ AI_K ~~and~~ BFI ~~becomes much less. However,~~ IAR ~~has almost equal negative contribution to~~ $\ln(\theta'_1) - \ln(\bar{\theta}_1)$ ~~in both periods. Unlike the former~~ ~~therethree~~ ~~variables, the significant negative contribution of~~ AI_{ET} ~~is only found in 1993-2003. For~~ AM_{30} ~~of Xianyang, the contribution of~~ IAR , AI_{ET} ~~and~~ BFI ~~is similar to that at~~ Huaxian station in two periods, however AI_K ~~is not included in the final model. In the latter period 1993-2003, the largest negative contribution was found in~~ AI_{ET} . These results suggest that the significant change of AI_K (mainly because of faster streamflow recession after nearly 1971) dominates the decrease in AM_{30} of Huaxian during 1970-1982, while after 1993, the significant change of AI_{ET} (due to decreasing precipitation and increasing evapotranspiration) has a main effect on the decrease in AM_{30} of Huaxian.

<Figure 11>

4.3. Discussion

The impacts of both human activities and climate change on low flows of the study area ~~of the Weihe basin~~ led to time-varying climate and catchment conditions (TCCCs). Nonstationary modeling for annual low flow series ~~considering using~~ TCCCs variables and/or HA variables as explanatory variables is clearly different from either the stationary model (M0) or the time covariate model (M1). The result demonstrates that considering multiple drivers (e.g. the variability in catchment conditions), especially in such an artificially influenced river, is necessary for nonstationary modeling of annual low flow series.

In this study area, nonstationary modeling considering TCCCs is supported by the following facts and findings. For human activities, an important milestone representative is the completion and operation of the irrigation system on plateau in Baoji Gorge Irrigation Area since 1971 (Sect. 3.1). Figure 5c shows the change of irrigation area in this basin. And T~~he~~ change-point detection test in Sect. 4.1 shows that significant change points of both annual recession constant (K) and low flow series occur exactly ~~in~~ at around 1971. This result demonstrates that changes in both K and AM_{30} may involve a consequence of this project. In addition to human activities, climate change also makes a considerable contribution to nonstationarity of low flows, as suggested by nonstationary modeling using TCCCs variables with stepwise analysis. Actually, climate driving

541 pattern may strengthen after nearly 1990, which is indicated by change-point detection test of both
542 annual mean temperature (T) and annual precipitation (P) as well as the behavior of annual low
543 flow series after nearly 1990. Therefore, the temporal variability in irrigation area, streamflow
544 recession, air temperature and precipitation (the frequency and volume of rain events) should be
545 the main driving factors of generating low flow regimes in this basin. Overall, the causes of
546 nonstationarity in category for two gauging stations have no clear difference, but have some
547 differences in the relative importance. As shown in Table 5, when modeling the low-flow series of
548 Huaxian using TCCCs variables, the optimal model (M4) preferred the variables are related to
549 recession process; however, for Xianyang, the preferred variables isare related to temperature. The
550 reason for this may be that as a downstream station, Huaxian station suffers more intensive human
551 activity, so that the importance of temperature change to the low-flow change is reduced
552 meanwhile the importance of streamflow recession (related to the capability of water storage)
553 change is improvenhanced.

554 Ignoring the negative impacts of the errors in estimating annual recession constant (K)
555 which are caused by insufficient data points of extracted stream segments at some wet years may
556 lead to the propagation of high errors in annual recession analysis, and accordingly affect the
557 quality of nonstationary frequency analysis when using K as an explanatory variable. Further
558 study will give more reliable estimation of K through improving annual recession analysis.

The related researches (Jiang et al., 2015; Yang and Yang, 2011; Yang and Yang, 2013; Zhang et al., 2015) have applied the Budyko framework to analyze the impacts of climate change and/or human activity on annual runoff. Indeed, for annual runoff, the Budyko framework is a better good method than the regression modeling method using in this study, because it used the mean annual water-energy balance equation to consider generation process of total runoff. Unfortunately, to our knowledge, there is a lack of the controls equation derived from basic physics laws for generation process of low flows. Therefore, we emphasize the importance of TCCCs variables to modeling of low-flow nonstationarity.

5. Conclusion

There is an increasing need to develop an effective nonstationary low-flow frequency model to deal with nonstationarities caused by climate change and time-varying anthropogenic activities. In this study, time-varying climate and catchment conditions (TCCCs) in the Weihe River basin were measured by annual time series of the eight indices, i.e., total precipitation (P), mean frequency of precipitation events (λ), temperature (T), potential evapotranspiration (ET), climate aridity index (AI_{ET}), base-flow index (BFI), recession constant (K), and the recession-related aridity index (AI_K).

The nonstationary distribution model was developed using both these eight TCCCs indices and/or there HA indices as candidate explanatory variables for frequency analysis of time-varying annual low flow series caused by multiple drivers. The main driving forces of the decrease in low flows in

577 | the Weihe River include reduced precipitation, warming climate, increasing irrigation area and
578 faster streamflow recession. Therefore, a complex deterioration mechanism resulting from these
579 factors demonstrates that in this arid and semi-arid area, the water resources could be vulnerable to
580 adverse environmental changes, thus portending increasing water shortages. The nonstationary
581 low-flow model considering TCCCs can provide the knowledge of low-flow generation
582 mechanism and give more reliable design of low flows for infrastructure and water supply.
583

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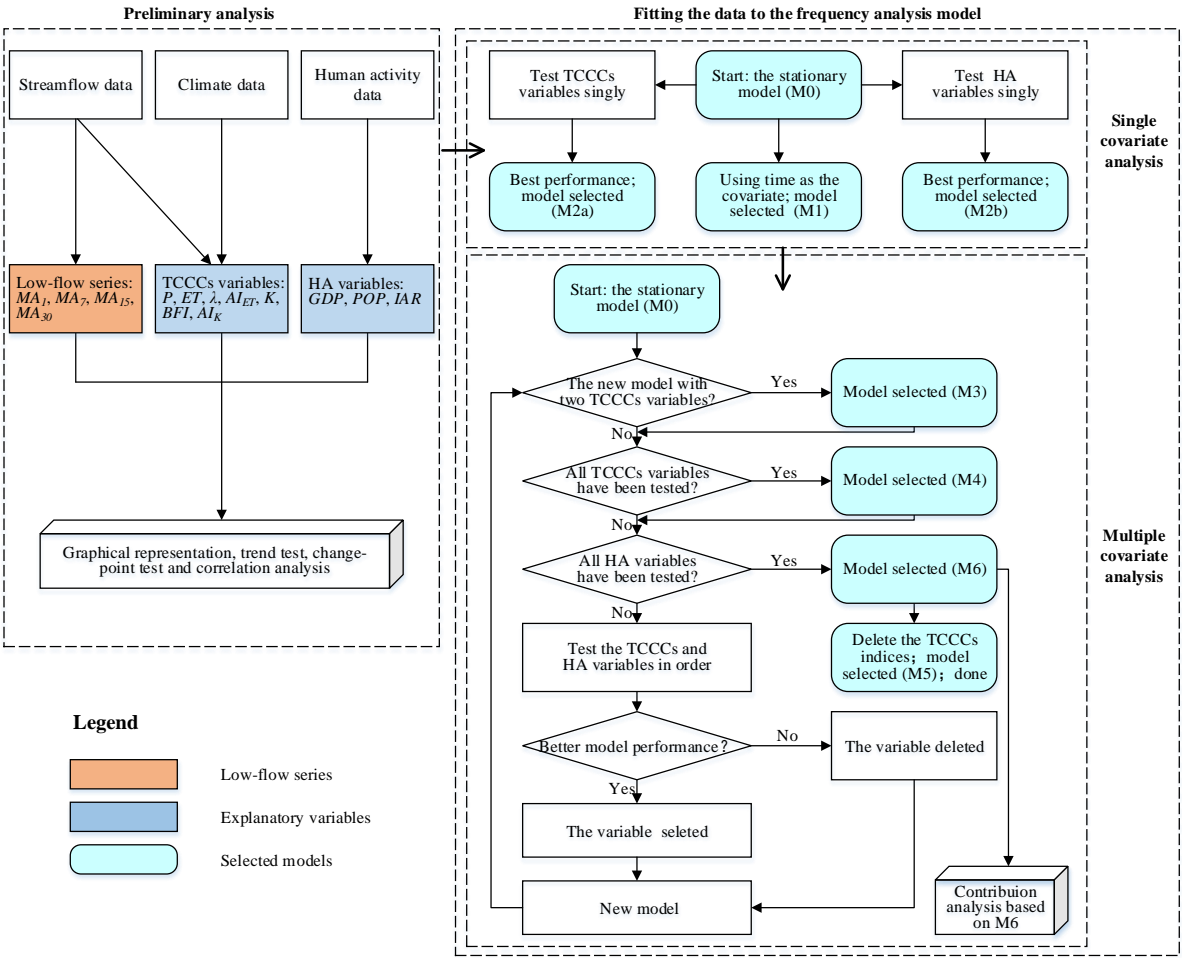
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753

754 **Figure**

755

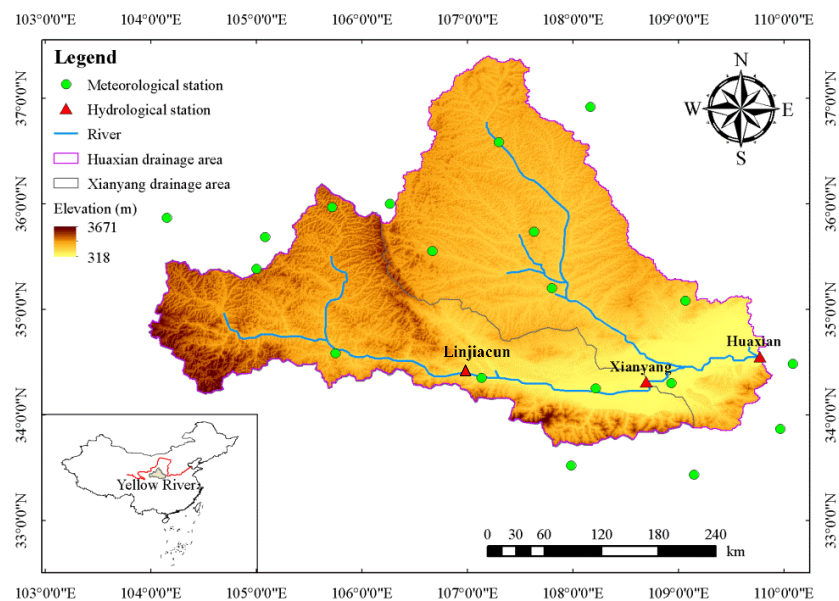


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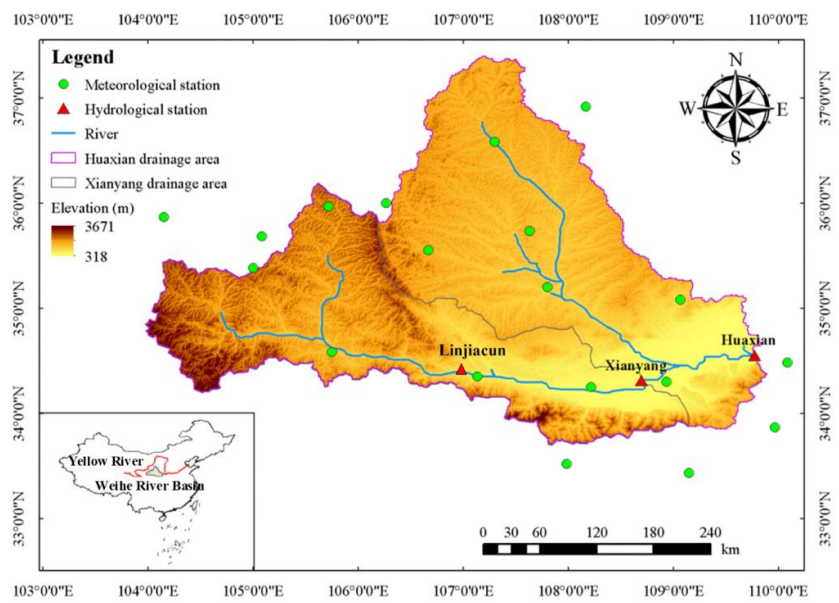
757 **Figure 1. The framework of nonstationary low-flow frequency analysis.**

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761 | Figure 42. Location, topography, hydro-meteorological stations and river systems of the Weihe
762 River basin.
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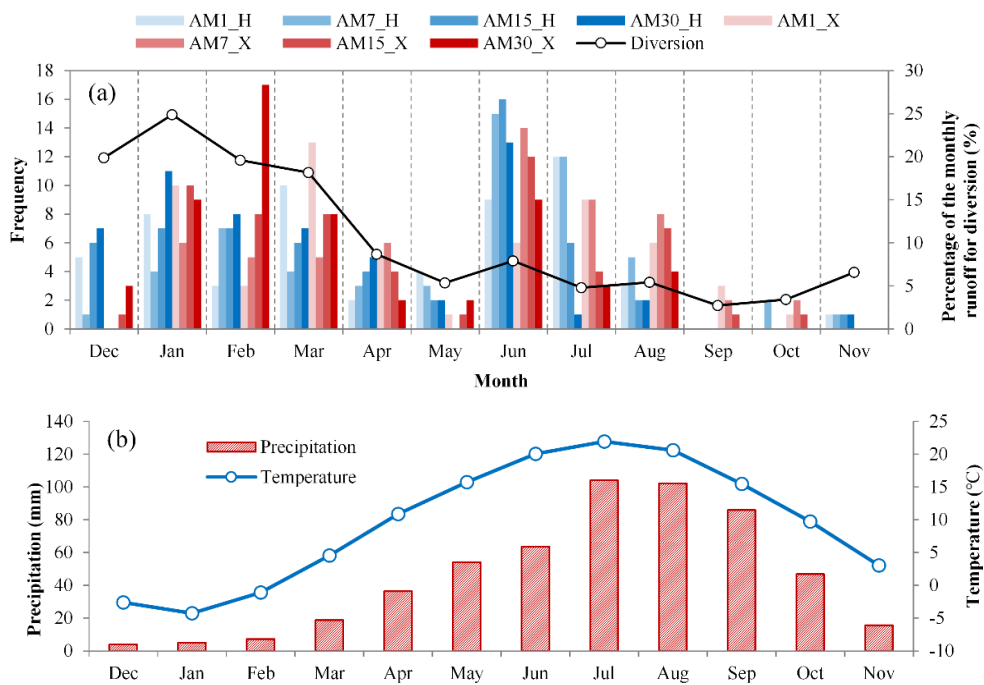


Figure 2. Overview of annual low flows and important environment factors using mean monthly data. (a) is frequency distributions of the occurrence time of the annual minimum flows with four durations at Huaxian (H) and Xianyang (X); the black line is mean monthly diversion (1980 to 2005) in Baoji Gorge area. (b) Mean monthly precipitation and temperature from 1954 to 2009.

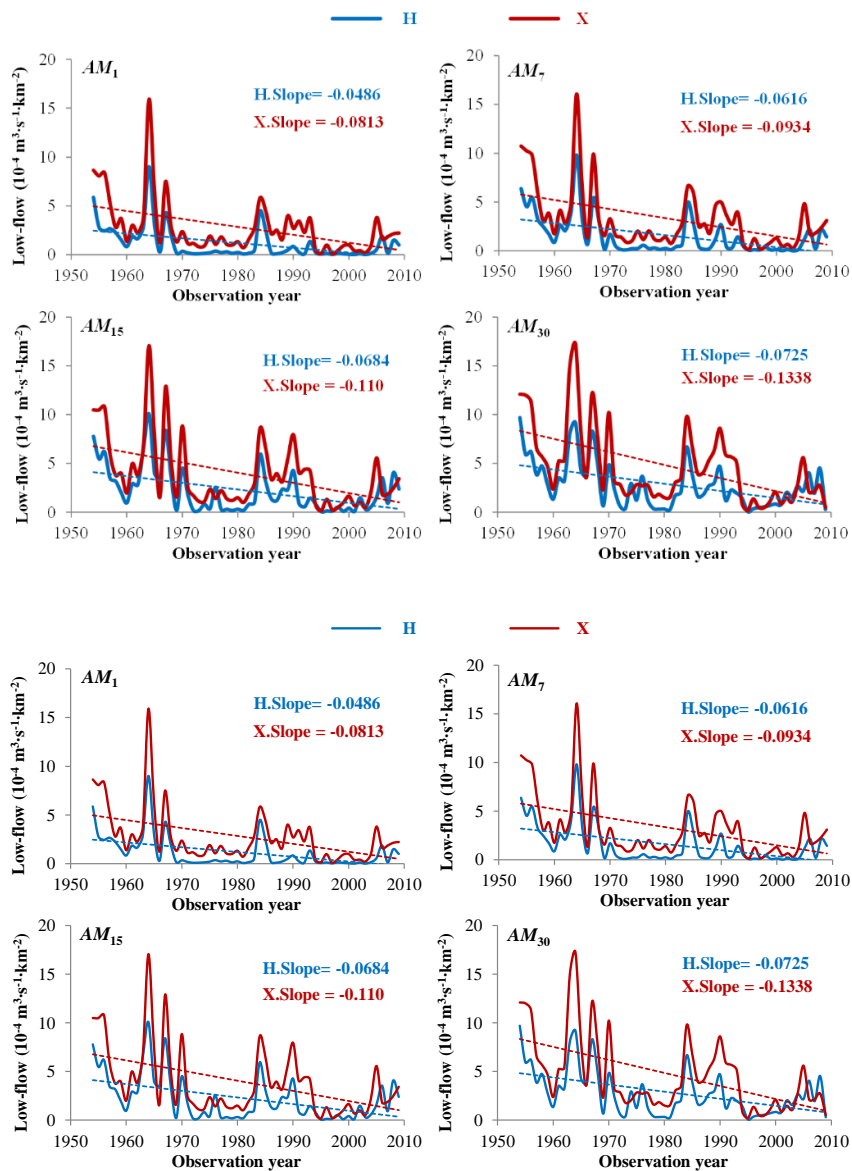
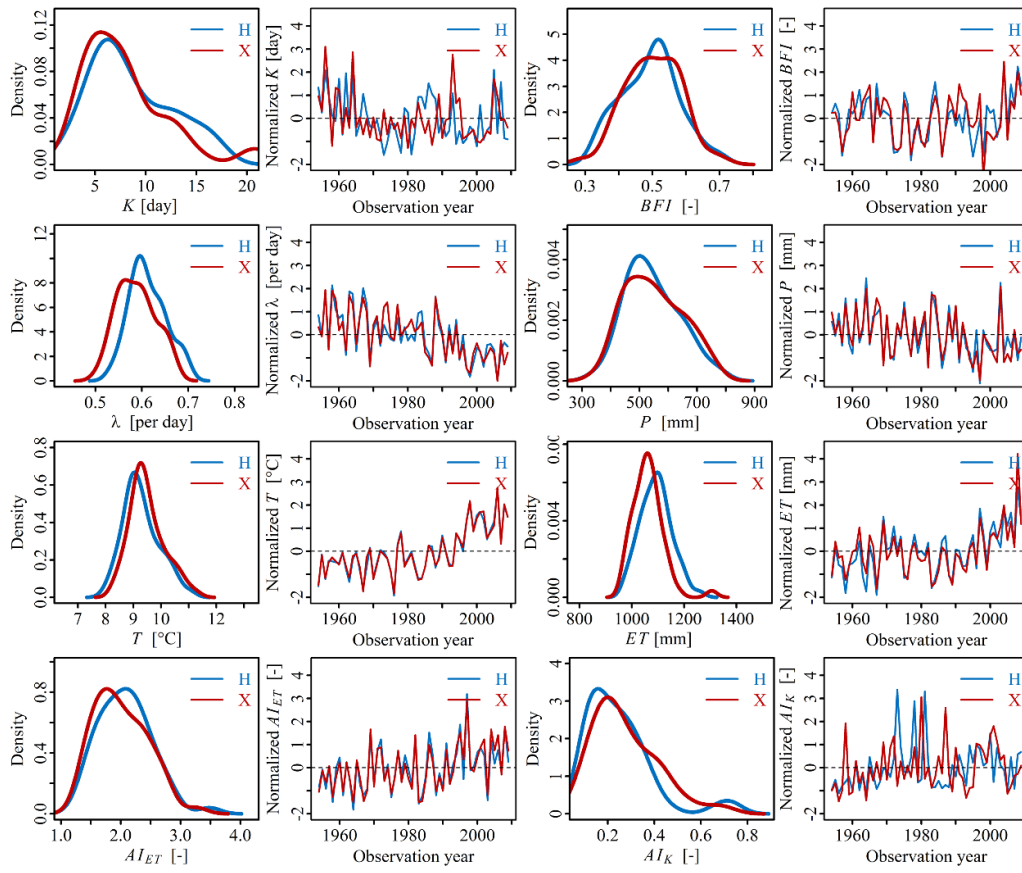


Figure 3. The annual minimum low flows and fitted trend lines in both Huaxian (H) and Xianyang (X) gauging stations.



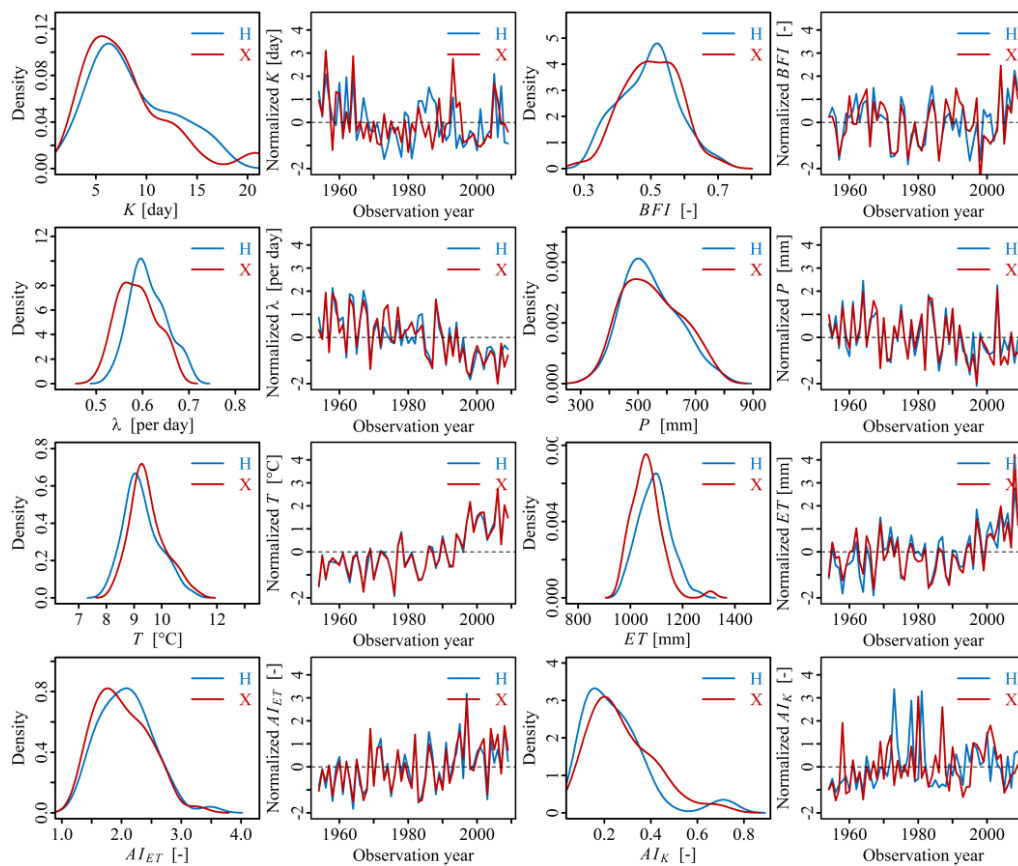


Figure 4. Frequency distributions (using the kernel density estimations) and ~~annual series of eight~~
~~candidate explanatory variables~~ time series processes of TCCCs variables in both Huaxian (H) and
Xianyang (X) stations.

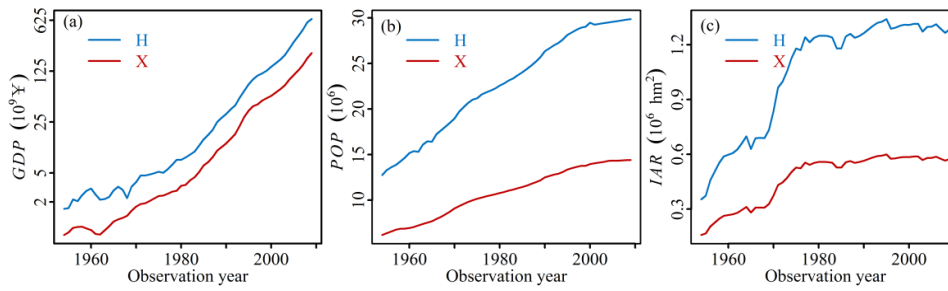
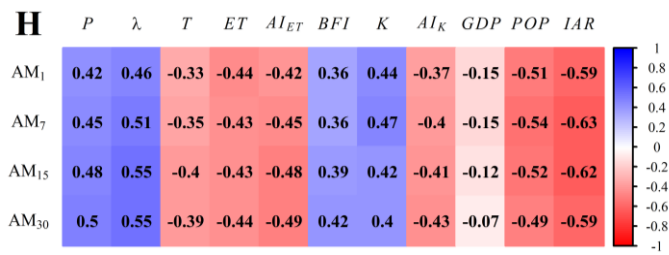
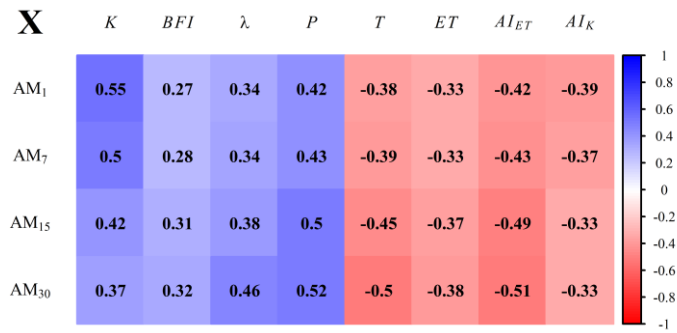
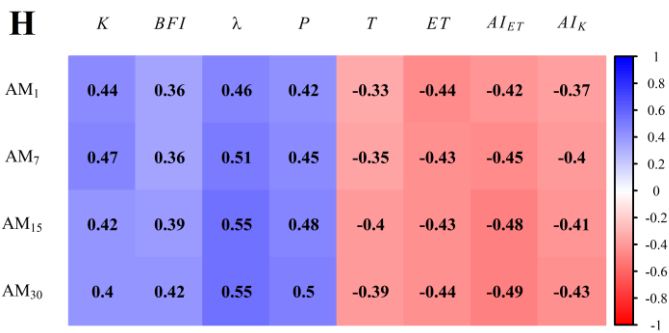


Figure 5. HA indices in both Huaxian (H) and Xianyang (X). (a), (b) and (c) are for population (*POP*), gross domestic production (*GDP*) and irrigated area (*IAR*), respectively.



788 Figure ~~56~~. The Pearson correlation coefficients matrix between the annual minimum flow series
789 and ~~eight~~ candidate explanatory variables in Huaxian (H) and Xianyang (X) stations; the darker
790 color intensity represents a higher level of correlation (blue indicates positive correlation, and red
791 indicates negative correlations).

792

H							X						
	LOGNO	PIII	GEV	WEI	GA	AIC		LOGNO	PIII	GEV	WEI	GA	AIC
AM_1	M0_none	112.4	112.9	114.3	108.6	110.9	114.3	227.2	226.5	228.0	227.2	226.3	228.0
	M1_I	94.8	97.0	90.3	94.4	97.0		211.1	210.6	209.5	211.5	209.9	
	M2_K	102.5	96.7	101.5	95.4	96.7		215.4	211.4	212.9	212.2	210.7	
	M2_BFI	103.4	99.2	97.6	96.7	99.2		220.8	217.0	218.8	218.9	217.7	
	M2_λ	99.9	99.1	97.8	96.9	99.1		217.2	216.1	218.6	217.7	216.5	
	M2_P	106.9	97.9	104.5	97.2	97.9		215.6	211.9	215.6	212.9	211.5	
	M2_T	101.6	102.8	99.9	100.5	102.8		211.9	212.8	211.5	214.4	212.7	
	M2_ET	101.5	99.9	100.5	98.3	99.9		217.0	216.6	214.9	218.4	217.1	
	M2_AI _{ET}	104.8	96.8	103.0	95.8	96.8		213.6	211.3	214.1	212.3	210.8	
M2_AI _K	103.6	95.8	101.8	95.0	95.8	90.3	218.2	214.0	217.4	215.2	214.0	209.5	
OPTIPAL MODEL: AIC(GEV, t)= 90.3							OPTIPAL MODEL: AIC(GEV, t)= 209.5						
AM_7	M0_none	151.8	155.0	158.2	153.5	155.2	158.2	243.5	243.8	245.8	245.1	244.1	245.8
	M1_I	134.7	138.5	137.9	139.1	140.7		227.9	228.2	228.0	229.7	227.9	
	M2_K	139.9	138.3	146.3	139.7	140.5		233.1	230.6	232.5	232.3	230.7	
	M2_BFI	141.8	141.5	146.1	142.1	143.7		235.9	234.2	236.2	236.8	235.4	
	M2_λ	137.7	139.4	142.4	140.3	141.6		233.8	233.3	236.8	235.5	234.2	
	M2_P	144.6	138.5	147.4	140.4	140.7		232.3	229.0	233.6	230.4	229.0	
	M2_T	141.6	144.6	144.4	145.2	146.7		229.1	230.1	230.5	232.2	230.3	
	M2_ET	141.6	142.6	147.0	143.7	144.8		234.4	234.0	234.7	236.4	235.0	
	M2_AI _{ET}	142.7	137.9	146.2	139.5	140.0		230.6	228.5	232.6	230.0	228.4	
M2_AI _K	139.9	135.7	140.7	137.7	138.0	134.7	235.6	232.2	236.3	234.1	232.8	227.9	
OPTIPAL MODEL: AIC(PIII, t)= 134.7							OPTIPAL MODEL: AIC(GA, t)= 227.9						
AM_{15}	M0_none	211.9	203.6	222.2	209.0	209.0	222.2	269.9	269.3	273.3	270.0	269.3	273.3
	M1_I	200.6	190.2	210.1	197.5	197.4		254.5	253.5	256.1	254.6	253.4	
	M2_K	199.9	190.7	208.0	198.0	198.0		262.0	259.0	262.9	260.8	259.8	
	M2_BFI	200.7	190.4	209.3	197.7	197.7		261.6	258.6	262.7	260.5	259.5	
	M2_λ	197.6	188.3	204.7	195.4	195.4		259.3	257.4	263.4	259.0	258.1	
	M2_P	200.6	187.6	210.3	195.0	194.9		256.2	251.6	258.9	252.4	251.4	
	M2_T	201.8	193.5	209.0	200.7	200.8		254.3	253.5	256.7	255.0	253.6	
	M2_ET	202.3	192.7	210.6	199.9	199.9		260.6	258.5	262.7	260.2	259.3	
	M2_AI _{ET}	199.4	187.3	209.0	194.6	194.5		254.7	251.3	258.2	252.1	251.0	
M2_AI _K	197.5	184.2	205.7	192.3	192.2	184.2	263.1	259.2	264.8	261.0	260.2	251.0	
OPTIPAL MODEL: AIC(PIII, AI _K)= 184.2							OPTIPAL MODEL: AIC(GA, AI _{ET})= 251.0						
AM_{30}	M0_none	241.4	237.3	244.1	236.5	236.3	244.1	290.7	290.6	293.4	290.8	289.8	293.4
	M1_I	233.0	225.4	233.5	225.4	225.5		271.6	270.4	271.5	271.8	270.1	
	M2_K	230.1	224.9	232.6	225.1	225.0		284.0	281.1	284.1	282.9	281.7	
	M2_BFI	226.2	223.5	228.5	224.0	223.5		282.3	278.8	284.1	280.3	279.3	
	M2_λ	228.5	221.4	227.1	221.1	221.1		276.9	274.7	279.6	276.0	274.8	
	M2_P	229.0	220.9	228.4	220.6	220.6		275.1	271.1	277.0	272.1	270.9	
	M2_T	233.4	227.2	232.2	227.7	227.5		270.8	270.3	271.1	272.1	270.1	
	M2_ET	232.9	226.2	233.3	226.4	226.4		280.9	278.3	282.1	279.9	278.7	
	M2_AI _{ET}	228.1	220.5	227.5	220.1	220.1		273.7	270.5	276.0	271.6	270.2	
M2_AI _K	224.5	217.5	220.8	217.7	217.4	217.4	283.9	279.9	284.0	281.5	280.5	270.1	
OPTIPAL MODEL: AIC(GA, AI _K)=217.4							OPTIPAL MODEL: AIC(GA, T)= 270.1						

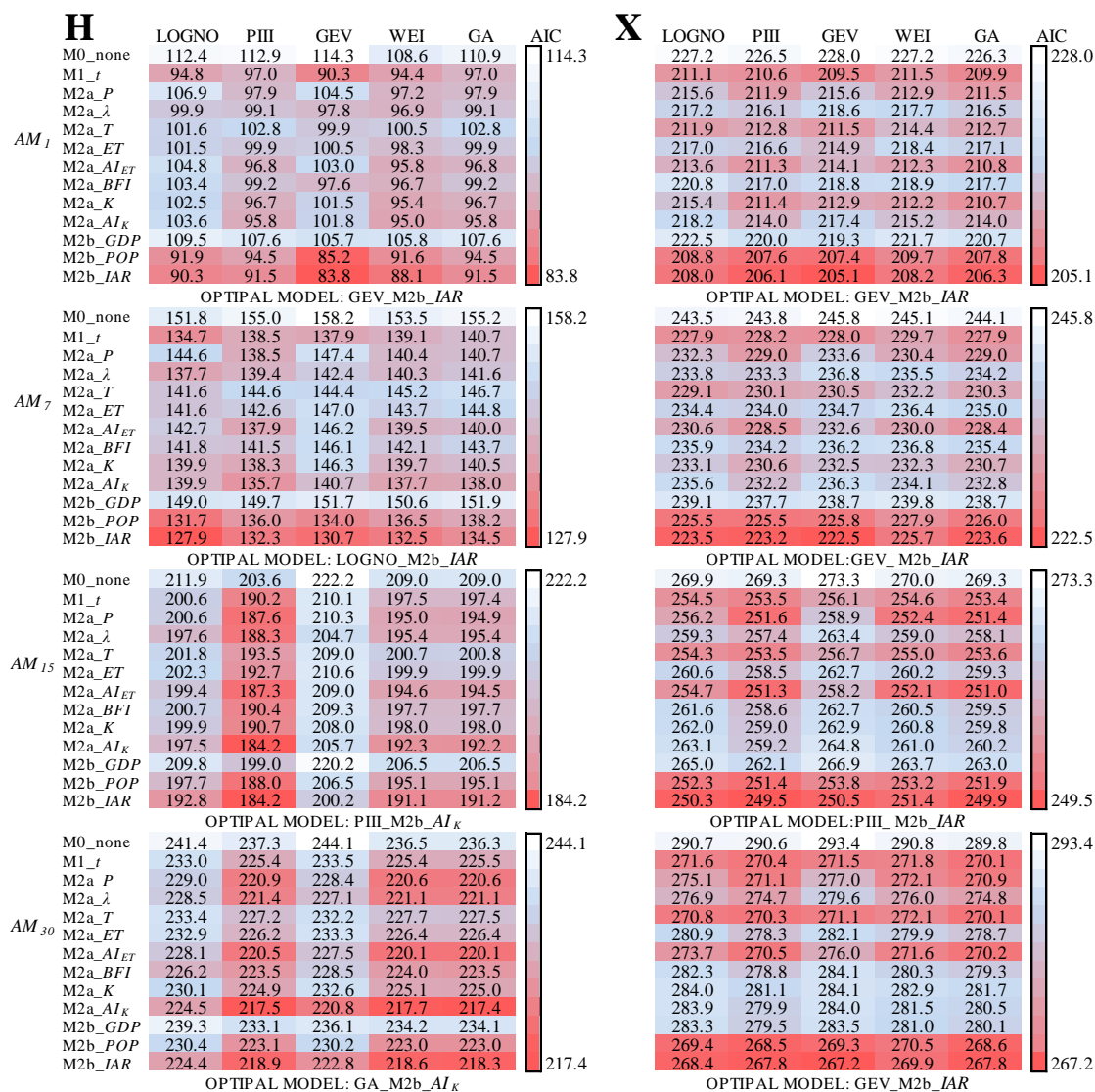
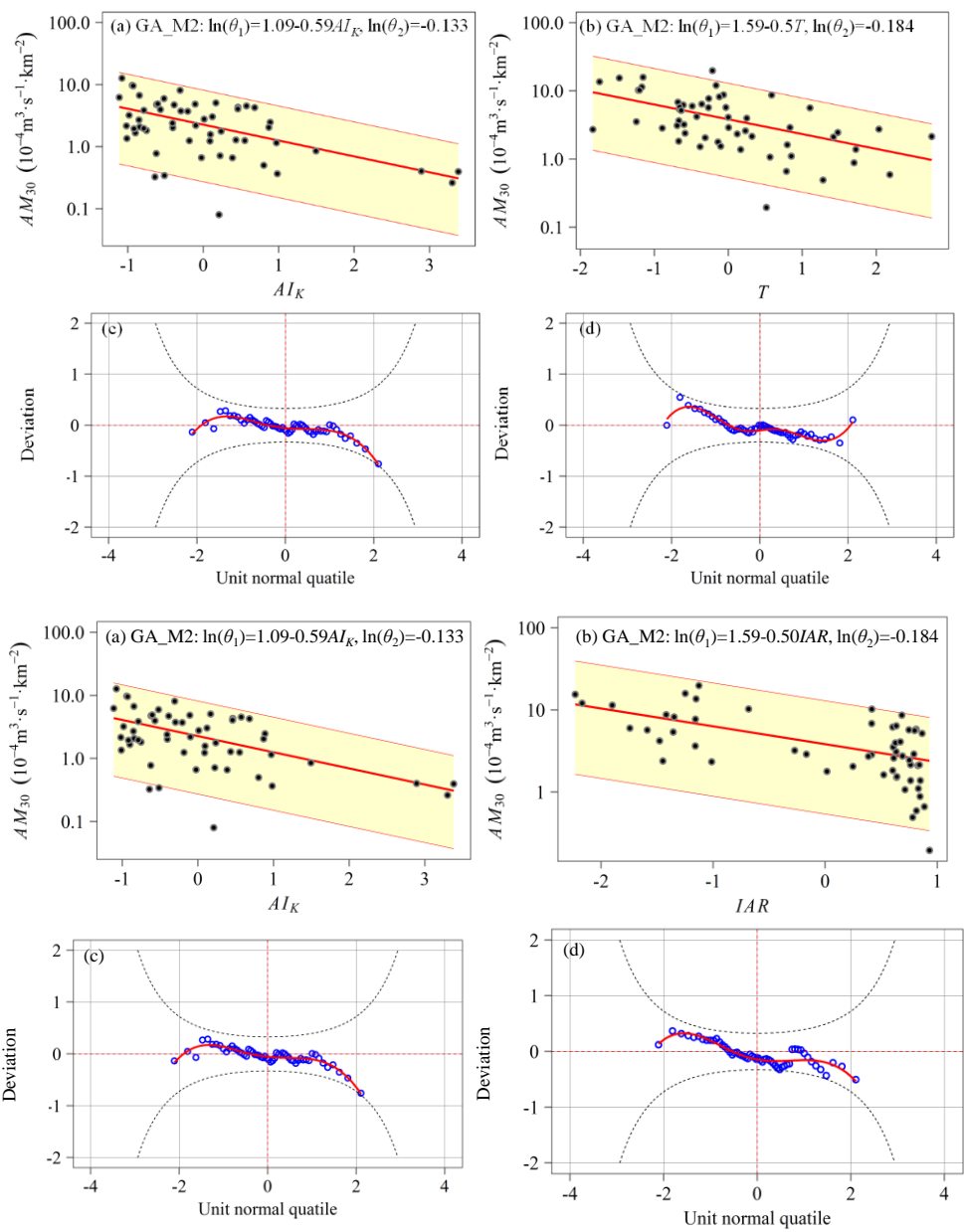


Figure 67. Comparisons among M0, M1 and M2 based on the AIC values for the four observed low-flow series in Huaxian (H) at left panel and Xianyang (X) at right panel; darker red color represents a higher goodness of fit.

H

X

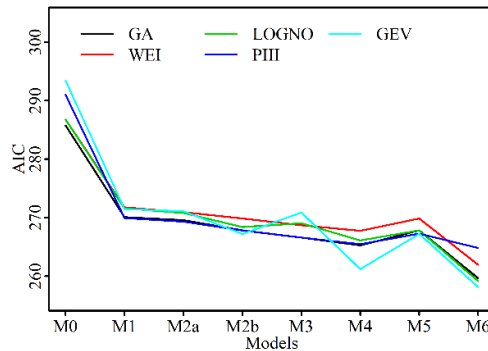
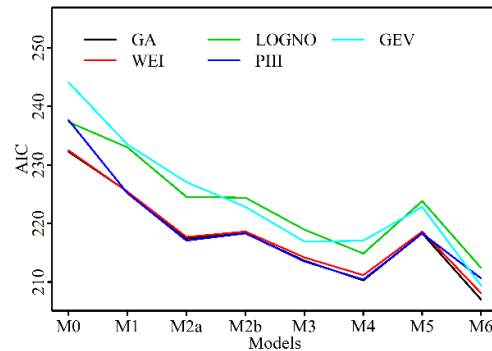
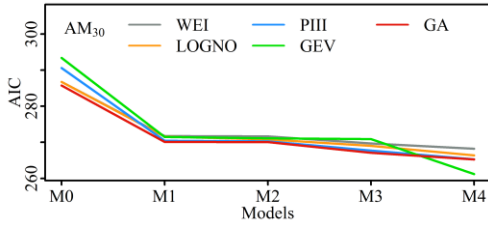
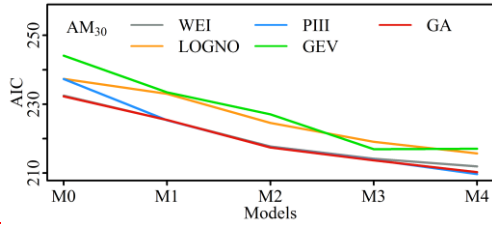
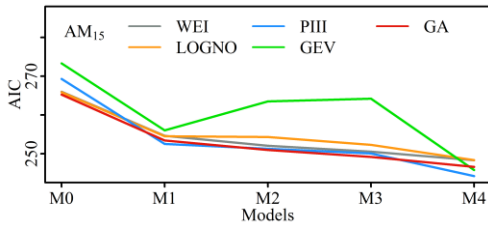
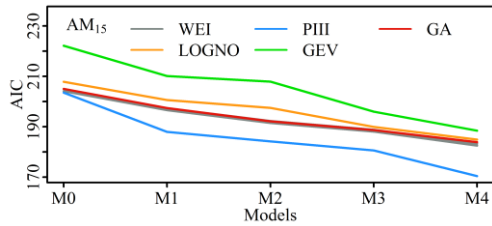
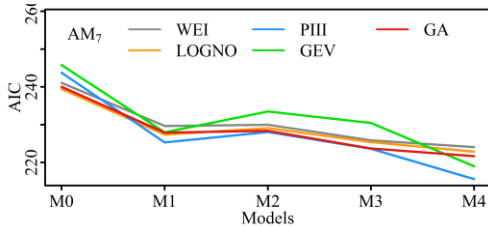
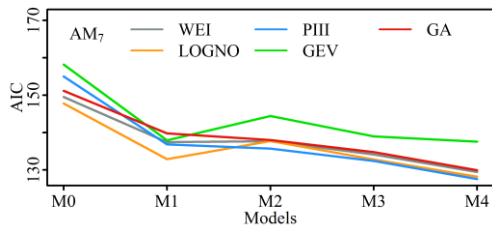
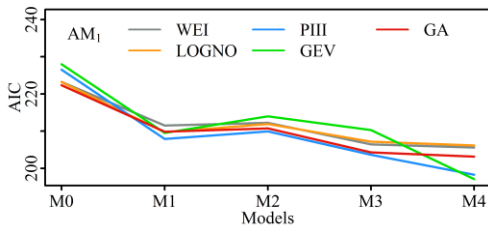
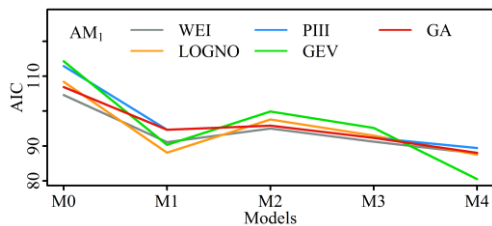


803 | Figure 78. Performance assessments of ~~the best M2 model (GA_M2)~~ GA_M2 for AM_{30} in
804 Huaxian (H) at left panel and Xianyang (X) at right panel. (a) and (b) are the centile curves plots
805 of GA_M2 (red lines represent the centile curves estimated by GA_M2; the 50th centile curves are
806 indicated by thick red; the yellow-filled areas are between the 5th and 95th centile curves; the
807 black points indicate the observed series); (c) and (d) are the worm plots of GA_M2 for the
808 goodness-of-fit test; a reasonable model fit should have the data points fall within the 95%
809 confidence intervals (between the two red dashed curves).

810

H

X



811

812

813 Figure 89. Comparisons ~~among of performance of~~ stationary model (M0), time covariate model
814 (M1) and physical covariate models (M2~~a~~, M2~~b~~, M3, M4, M5 and M6 with ~~the~~ their
815 corresponding optimal explanatory variables) for AM_{30} in Huaxian (H) at left panel and
816 Xianyang (X) at right panel.

817

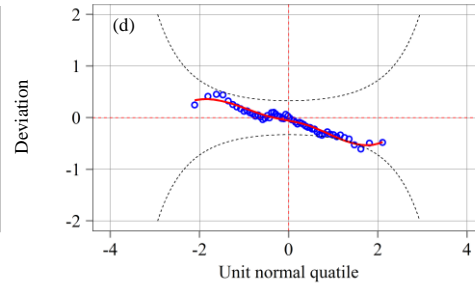
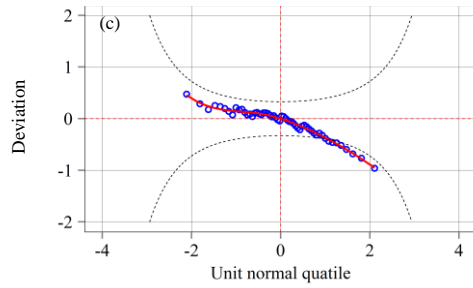
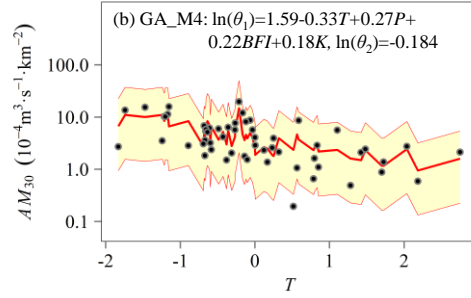
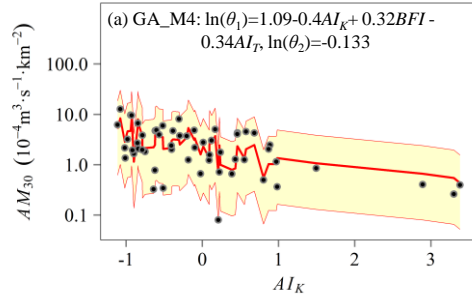
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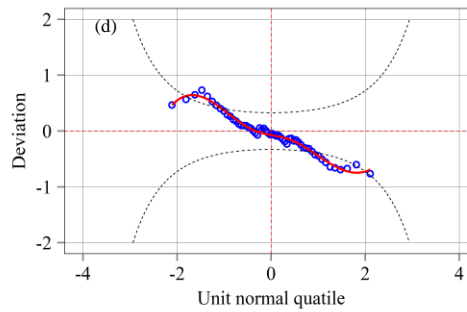
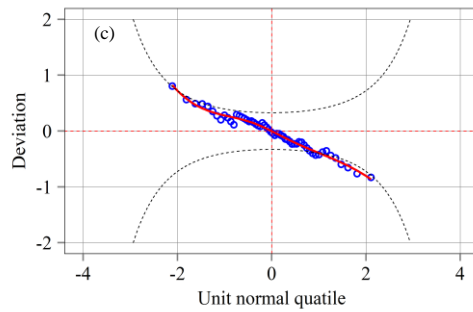
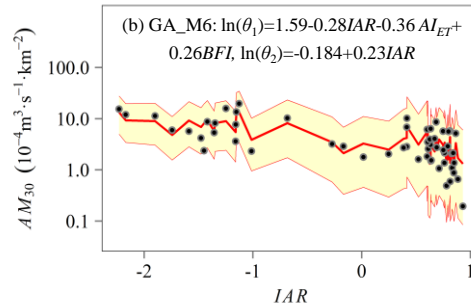
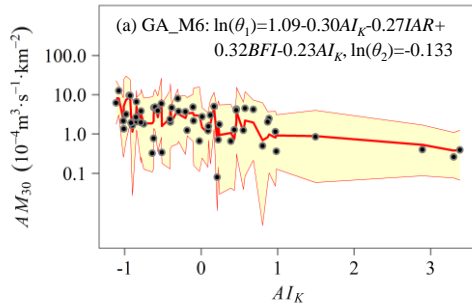
820

H

X



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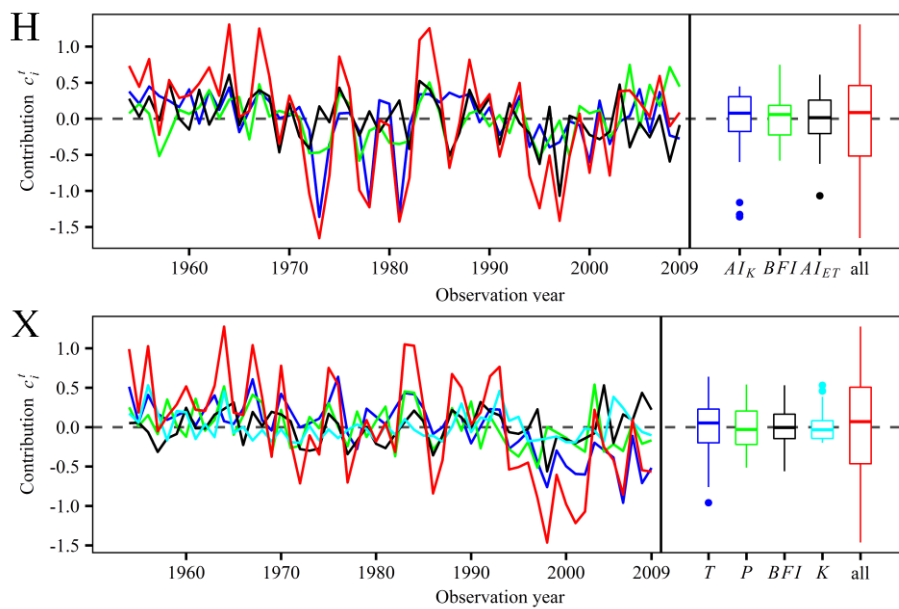


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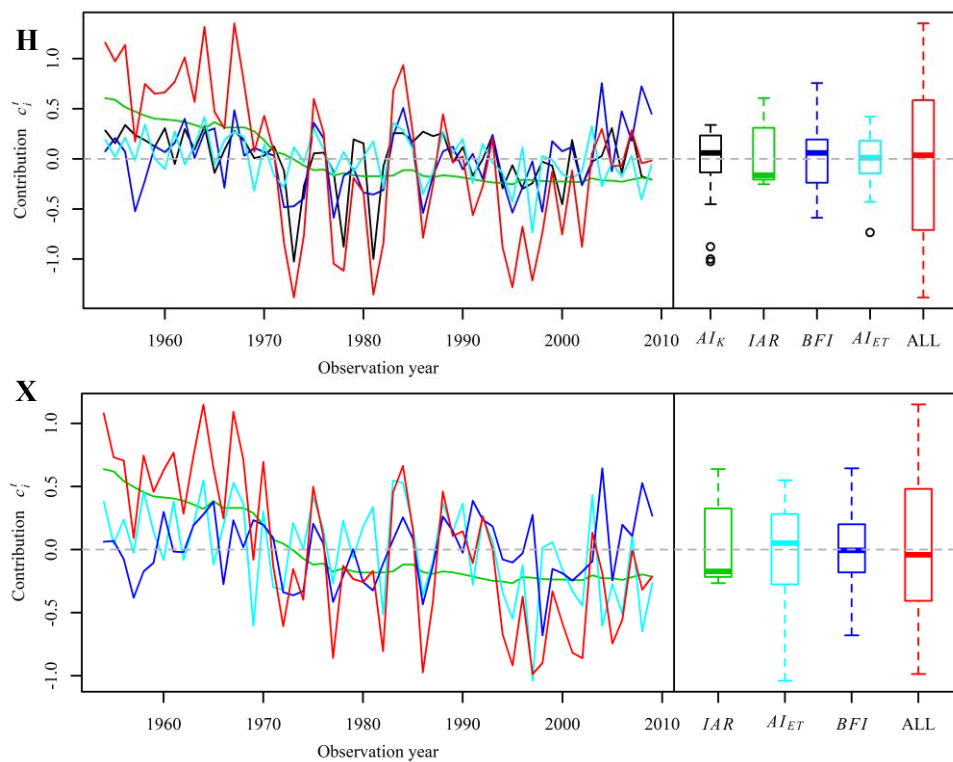
823 | Figure 910. Performance assessments of ~~the best M4 model (GA_M4)~~ GA_M6 for AM_{30} in
824 | Huaxian (H) at left panel and Xianyang (X) at right panel. (a) and (b) are the centile curves plots
825 | of GA_M4 GA_M6 (red lines represent the centile curves estimated by GA_M4 GA_M6; the 50th
826 | centile curves are indicated by thick red; the yellow-filled areas are between the 5th and 95th
827 | centile curves; the filled black points indicate the observed series); (c) and (d) are the worm plots
828 | of GA_M4 GA_M6 for the goodness-of-fit test; A reasonable model fit should have the data points
829 | fall within the 95% confidence intervals (between the two red dashed curves).

830

831



832



833 Figure ~~40~~11. Contribution of selected explanatory variables to $c'_i = \ln(\theta'_i) - \ln(\bar{\theta}_i)$ in different
834 periods based on GA_~~M4~~M6.

835

836

837

838

839 **Table**

840 Table 1. The probability density functions and moments (the mean and variance) for the candidate
841 distributions in this study.

Distributions	Probability density function	Distribution moments
Pearson-III	$f_Y(y \theta_1, \theta_2, \theta_3) = \frac{(y - \theta_3)^{1/\theta_2^2 - 1}}{\Gamma(1/\theta_2^2)(\theta_1 \theta_2^2)^{1/\theta_2^2}} \exp\left(-\frac{y - \theta_3}{\theta_1 \theta_2^2}\right)$ $y > \theta_3, \theta_3 > 0, \theta_1 > 0, \theta_2 > 0$	$E[Y] = \theta_1 + \theta_3$ $Var[Y] = \theta_1^2 \theta_2^2$
Gamma	$f_Y(y \theta_1, \theta_2) = \frac{(y)^{1/\theta_2^2 - 1}}{\Gamma(1/\theta_2^2)(\theta_1 \theta_2^2)^{1/\theta_2^2}} \exp\left(-\frac{y}{\theta_1 \theta_2^2}\right)$ $y > 0, \theta_1 > 0, \theta_2 > 0$	$E[Y] = \theta_1$ $Var[Y] = \theta_1^2 \theta_2^2$
Weibull	$f_Y(y \theta_1, \theta_2) = \left(\frac{\theta_2}{\theta_1}\right) \left(\frac{y}{\theta_1}\right)^{\theta_2 - 1} \exp\left(-\left(\frac{y}{\theta_1}\right)^{\theta_2}\right)$ $y > 0, \theta_1 > 0, \theta_2 > 0$	$E[Y] = \theta_1 \Gamma(1 + 1/\theta_2)$ $Var[Y] = \theta_1^2 \left[\Gamma\left(1 + \frac{2}{\theta_2}\right) - \Gamma^2\left(1 + \frac{1}{\theta_2}\right) \right]$
Lognormal	$f_Y(y \theta_1, \theta_2) = \frac{1}{y \theta_2 \sqrt{2\pi}} \exp\left\{-\frac{[\log(y) - \theta_1]^2}{2\theta_2^2}\right\}$ $y > 0, \theta_2 > 0$	$E[Y] = w^{\theta_2} e^{\theta_1}$ $Var[Y] = w(w - 1) e^{2\theta_1}$ $w = \exp(\theta_2^2)$
GEV	$f_Y(y \theta_1, \theta_2, \theta_3) = \frac{1}{\theta_2} \left[1 + \theta_3 \left(\frac{y - \theta_1}{\theta_2} \right) \right]^{-1/\theta_3 - 1} \exp\left\{-\left[1 + \theta_3 \left(\frac{y - \theta_1}{\theta_2} \right) \right]^{-1/\theta_3}\right\}$ $-\infty < \theta_1 < \infty, \theta_2 > 0, -\infty < \theta_3 < \infty$	$E[Y] = \theta_1 - \frac{\theta_2}{\theta_3} + \frac{\theta_2}{\theta_3} \eta_1$ $Var[Y] = \theta_2^2 (\eta_2 - \eta_1^2) / \theta_3^2$ $\eta_m = \Gamma(1 - m\theta_3)$

842

843 Table 2. Description of the developed nonstationary models using time, ~~or the indices of~~ TCCCs
844 ~~indices and/or HA indices~~ as explanatory variables.

Model- category	Codes	Distribution					Description	
		GA	WEI	LOGNO	PIII	GEV	Variable category	The numbers of variables
Stationary	M0	GA_M0	WEI_M0	LOGNO_M0	PIII_M0	GEV_M0	-	Zero
	M1	GA_M1	WEI_M1	LOGNO_M1	PIII_M1	GEV_M1	Time	One
Nonstationary	M2	GA_M2	WEI_M2	LOGNO_M2	PIII_M2	GEV_M2	TCCCs	One
	M3	GA_M3	WEI_M3	LOGNO_M3	PIII_M3	GEV_M3	TCCCs	Two
	M4	GA_M4	WEI_M4	LOGNO_M4	PIII_M4	GEV_M4	TCCCs	Identified by the stepwise selection

Model	Distribution					Description	
codes	GA	WEI	LOGNO	PIII	GEV	Variable category	The numbers of variables
<u>M0</u>	<u>GA_M0</u>	<u>WEI_M0</u>	<u>LOGNO_M0</u>	<u>PIII_M0</u>	<u>GEV_M0</u>	<u>-</u>	<u>Zero</u>
<u>M1</u>	<u>GA_M1</u>	<u>WEI_M1</u>	<u>LOGNO_M1</u>	<u>PIII_M1</u>	<u>GEV_M1</u>	<u>Time</u>	<u>One</u>
<u>M2a</u>	<u>GA_M2a</u>	<u>WEI_M2a</u>	<u>LOGNO_M2a</u>	<u>PIII_M2a</u>	<u>GEV_M2a</u>	<u>TCCCs</u>	<u>One</u>
<u>M2b</u>	<u>GA_M2b</u>	<u>WEI_M2b</u>	<u>LOGNO_M2b</u>	<u>PIII_M2b</u>	<u>GEV_M2b</u>	<u>HA</u>	<u>One</u>
<u>M3</u>	<u>GA_M3</u>	<u>WEI_M3</u>	<u>LOGNO_M3</u>	<u>PIII_M3</u>	<u>GEV_M3</u>	<u>TCCCs</u>	<u>Two</u>
<u>M4</u>	<u>GA_M4</u>	<u>WEI_M4</u>	<u>LOGNO_M4</u>	<u>PIII_M4</u>	<u>GEV_M4</u>	<u>TCCCs</u>	<u>Identified by the stepwise selection</u>
<u>M5</u>	<u>GA_M5</u>	<u>WEI_M5</u>	<u>LOGNO_M5</u>	<u>PIII_M5</u>	<u>GEV_M5</u>	<u>HA</u>	<u>Identified by the stepwise selection</u>
<u>M6</u>	<u>GA_M6</u>	<u>WEI_M6</u>	<u>LOGNO_M6</u>	<u>PIII_M6</u>	<u>GEV_M6</u>	<u>TCCCs+HA</u>	<u>Identified by the stepwise selection</u>

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Table 3. The summary of candidate explanatory variables and reason of selection.

<u>Category</u>	<u>Name</u>	<u>Indices</u>	<u>Reason of selection (related to)</u>	<u>Unit</u>
<u>TCCCs</u>	<u>P</u>	<u>Precipitation</u>	<u>Main supply source</u>	<u>mm</u>
	<u>λ</u>	<u>Mean frequency of precipitation events</u>	<u>Water supply intensity</u>	<u>per day</u>
	<u>T</u>	<u>Temperature</u>	<u>Evaporation loss</u>	<u>°C</u>
	<u>ET</u>	<u>Potential evapotranspiration</u>	<u>Evaporation loss</u>	<u>mm</u>
	<u>AI_{ET}</u>	<u>Climate aridity index</u>	<u>Degree of meteorological drought</u>	<u>-</u>
	<u>BFI</u>	<u>Base-flow index</u>	<u>Water storage capability</u>	<u>-</u>
	<u>K</u>	<u>Recession constant</u>	<u>Water storage capability</u>	<u>day</u>
	<u>AI_K</u>	<u>Recession-related aridity index</u>	<u>Both the water storage and supply capability</u>	<u>-</u>
<u>HA</u>	<u>IAR</u>	<u>Irrigation area</u>	<u>Both irrigation diversion and evaporation loss</u>	<u>10⁶ hm²</u>
	<u>POP</u>	<u>Population</u>	<u>Water withdrawal loss for agricultural, domestic and industrial purposes</u>	<u>10⁶</u>
	<u>GDP</u>	<u>Gross domestic product</u>	<u>Water withdrawal loss for agricultural, domestic and industrial purposes</u>	<u>10⁹ ¥</u>

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852 Table 34. The results of trend test and change-point detection for both the four low flow series and
853 eight candidate explanatory TCCCs variables in Huaxian and Xianyang stations.

Station	Variable	Mann-Kendall test		Pettitt's test	
		S	<i>p-value</i>	Change point	<i>p-value</i>
Huaxian	AM_1	-564	6.91E-05(***)	1968	1.34E-03(**)
	AM_7	-560	7.79E-05(***)	1968	1.44E-03(**)
	AM_{15}	-438	2.01E-03(**)	1971	4.85E-03(**)
	AM_{30}	-378	7.71E-03(**)	1971	9.96E-03(**)
	P	-292	3.97E-02(*)	1985	1.86E-01()
	λ	-632	8.20E-06(***)	1984	3.02E-04(***)
	T	752	1.11E-07(***)	1993	8.17E-06(***)
	ET	548	1.11E-04(***)	1993	1.98E-03(**)
	AI_{ET}	384	6.79E-03(**)	1990	6.03E-02()
	BFI	52	7.19E-01()	1998	3.88E-01()
	K	-312	2.79E-02(*)	1968	8.11E-02()
	AI_K	376	8.04E-03(**)	1971	3.60E-02(*)
Xianyang	AM_1	-517	2.65E-04(***)	1968	2.2E-03(**)
	AM_7	-483	6.58E-04(***)	1970	2.5E-03(**)
	AM_{15}	-474	8.29E-04(***)	1971	2.2E-03(**)
	AM_{30}	-570	5.78E-05(***)	1993	4.5E-04(***)
	P	-414	3.51E-03(**)	1990	1.45E-02(*)
	λ	-652	4.21E-06(***)	1984	6.00E-05(***)
	T	724	3.22E-07(***)	1993	5.41E-06(***)
	ET	372	8.74E-03(**)	1993	3.01E-03(**)
	AI_{ET}	454	1.37E-03(**)	1993	8.82E-03(**)
	BFI	64	6.56E-01()	2003	8.65E-01()
	K	-210	1.39E-01()	1966	2.03E-01()
	AI_K	290	4.11E-02(*)	1968	1.63E-01()

Signif. codes: 0 '***' 0.001 '**' 0.01 '*' 0.05 '.' 0.1 ' ' 1

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Table 4. The results of M2 models for modeling low flow series in Huaxian and Xianyang stations.

Station	Series	Optimal-variable	Optimal-distribution	AIC	Distribution parameters		
					$-\ln(\theta_1)$	$-\ln(\theta_2)$	θ_3
Huaxian	AM_t	AI_K	WEI	95.0	$-0.19 - 0.72 AI_K$	-0.418	-
	AM_7	AI_K	PIII	135.7	$-0.43 - 0.76 AI_K$	0.219	0.007
	AM_{15}	AI_K	PIII	184.2	$-0.83 - 0.75 AI_K$	0.105	0.069
	AM_{30}	AI_K	GA	217.4	$-1.09 - 0.59 AI_K$	-0.133	-
Xianyang	AM_t	K	GA	210.7	$-1.00 + 0.40 K$	-0.118	-
	AM_7	AI_{EF}	GA	228.4	$-1.17 - 0.45 AI_{EF}$	-0.139	-
	AM_{15}	AI_{EF}	GA	251.0	$-1.39 - 0.49 AI_{EF}$	-0.139	-
	AM_{30}	T	GA	270.1	$-1.59 - 0.50 T$	-0.184	-

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860 | Table 5. The summary of frequency analysis using GA distribution for AM_{30} in Huaxian and Xianyang.

Station	Model codes	Optimal variable	AIC	Distribution parameters		
				$\ln(\theta_1)$	$\ln(\theta_2)$	θ_3
Huaxian						
	GA_M0	z	232.3	1.09	-0.133	-
	GA_M1	t	225.5	$1.09-0.32t$	-0.133	-
	GA_M2	AI_K	217.4	$1.09-0.59AI_K$	-0.133	-
	GA_M2b	IAR	218.3	$1.09-0.47IAR$	-0.133	-
	GA_M3	AI_K, BFI	213.7	$1.09-0.50AI_K+0.32BFI$	-0.133	-
	GA_M4	AI_K, BFI, AI_{ET}	211.1	$1.09-0.40AI_K+0.32BFI-0.34AI_{ET}$	-0.133	-
	GA_M5	IAR	218.3	$1.09-0.47IAR$	-0.133	-
	GA_M6	AI_K, IAR, BFI, AI_{ET}	207.0	$1.09-0.30AI_K-0.27IAR+0.32BFI-0.23AI_{ET}$	-0.133	-
Xianyang						
	GA_M0	z	285.8	1.59	-0.184	-
	GA_M1	t	270.1	$1.59-0.48t$	-0.184	-
	GA_M2a	T	270.1	$1.59-0.50T$	-0.184	-
	GA_M2b	IAR	267.8	$1.59-0.50IAR$	-0.184	-
	GA_M3	T, P	267.1	$1.59-0.34T+0.32P$	-0.184	-
	GA_M4	T, P, BFI, K	265.4	$1.59-0.33T+0.27P+0.22BFI+0.18K$	-0.184	-
	GA_M5	IAR	267.8	$1.59-0.50IAR$	-0.184	-
	GA_M6	IAR, AI_{ET}, BFI	259.7	$1.59-0.28IAR-0.36 AI_{ET}+0.26BFI$	$-0.184+0.23IAR$	-

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Table 5. The summary of frequency analysis for four annual low flow series of Huaxian.

Series	Model-codes	Optimal-variable	AIC	Distribution parameters		
				$\ln(\theta_1)$	$\ln(\theta_2)$	θ_3
AM_t	WEI_M0	-	104.6	-0.19	-0.418	-
	WEI_M1	t	91.1	$-0.19-0.84t$	$-0.418-0.30t$	-
	WEI_M2	AI_K	95.0	$-0.19-0.72AI_K$	-0.418	-
	WEI_M3	$AI_{K1}-BFI$	91.3	$-0.19-0.58AI_K+0.55BFI$	-0.418	-
	WEI_M4	$AI_{K1}-BFI, ET, \lambda$	87.9	$-0.19-0.39AI_K+0.61BFI-0.54ET$	$-0.418+0.27\lambda$	-
AM_{15}	PIH_M0	-	155.0	0.43	0.219	0.007
	PIH_M1	t	136.8	$0.43-0.59t$	$0.219+0.19t$	0.007
	PIH_M2	AI_K	135.7	$0.43-0.76AI_K$	0.219	0.007
	PIH_M3	$AI_{K1}-BFI$	132.4	$0.43-0.65AI_K+0.48BFI$	0.219	0.007
	PIH_M4	$AI_{K1}-BFI, AI_{ET}, \lambda, P$	127.5	$0.43-0.62AI_K+0.57BFI-0.60AI_{ET}$	$0.219-0.32\lambda-0.30AI_K+0.21P$	0.007
AM_{15}	PIH_M0	-	203.5	0.83	0.105	0.069
	PIH_M1	t	188.0	$0.83-0.46t$	$0.105+0.208t$	0.069
	PIH_M2	AI_K	184.2	$0.83-0.75AI_K$	0.105	0.069
	PIH_M3	$AI_{K1}-BFI$	180.6	$0.83-0.65AI_K+0.43BFI$	0.105	0.069
	PIH_M4	$AI_{K1}-BFI, \lambda, K$	170.4	$0.83-0.70AI_K+0.42BFI$	$0.105-0.36\lambda-0.71AI_K-0.43K$	0.069
AM_{30}	GA_M0	-	232.3	1.09	-0.133	-
	GA_M1	t	225.5	$1.09-0.32t$	-0.133	-
	GA_M2	AI_K	217.4	$1.09-0.59AI_K$	-0.133	-
	GA_M3	$AI_{K1}-BFI$	213.7	$1.09-0.5AI_K+0.32BFI$	-0.133	-
	GA_M4	$AI_{K1}-BFI, AI_F$	211.1	$1.09-0.4AI_K+0.32BFI-0.34AI_F$	-0.133	-

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868 Table 6. The summary of frequency analysis for four annual low flow series of Xianyang.

Series	Model-codes	Optimal-variable	AIC	Distribution-parameters	
				$-\ln(\theta_1)$	$-\ln(\theta_2)$
AM ₁	GA_M0	-	222.3	1.0	-0.118
	GA_M1	t	209.9	1.0-0.44 t	-0.118
	GA_M2	K	210.7	1.0+0.4 K	-0.118
	GA_M3	K, T	204.3	1.0+0.37 K -0.38 T	-0.118
	GA_M4	K, T, BFI, λ	203.2	1.0+0.33 K -0.32 T +0.27 BFI	-0.118-0.17 λ
AM ₂	GA_M0	-	240.1	1.17	-0.139
	GA_M1	t	227.9	1.17-0.42 t	-0.139
	GA_M2	AI_{EF}	228.4	1.17-0.45 AI_{EF}	-0.139
	GA_M3	$AI_{EF}-K$	223.7	1.17-0.38 AI_{EF} +0.31 K	-0.139
	GA_M4	AI_{EF}, K, BFI, λ	221.7	1.17-0.31 AI_{EF} +0.3 K +0.28 BFI	-0.139-0.2 λ
AM ₁₅	GA_M0	-	265.3	1.39	-0.139
	GA_M1	t	253.4	1.39-0.43 t	-0.139
	GA_M2	AI_{EF}	251.0	1.39-0.49 AI_{EF}	-0.139
	GA_M3	$AI_{EF}-K$	249.2	1.39-0.45 AI_{EF} +0.24 K	-0.139
	GA_M4	AI_{EF}, K, BFI, λ	246.6	1.39-0.36 AI_{EF} +0.23 K +0.32 BFI	-0.139-0.21 λ
AM ₃₀	GA_M0	-	285.8	1.59	-0.184
	GA_M1	t	270.1	1.59-0.48 t	-0.184
	GA_M2	T	270.1	1.59-0.5 T	-0.184
	GA_M3	T, P	267.1	1.59-0.34 T +0.32 P	-0.184
	GA_M4	T, P, BFI, K	265.4	1.59-0.33 T +0.27 P +0.22 BFI +0.18 K	-0.184

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