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Spatio-temporal patterns of the effects of precipitation variability
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      and land use/cover changes on long-term changes in sediment yield
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      in the Loess Plateau, China
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      Guangyao Gao<sup>1,2</sup>, Jianjun Zhang<sup>1</sup>, Yu Liu<sup>3</sup>, Zheng Ning<sup>1</sup>, Bojie Fu<sup>1,2</sup>, and Murugesu
 5
      Sivapalan<sup>4,5</sup>
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 7
      <sup>1</sup>State Key Laboratory of Urban and Regional Ecology, Research Center for
 8
      Eco-Environmental Sciences, Chinese Academy of Sciences, Beijing 100085, China
 9
      <sup>2</sup>Joint Center for Global Change Studies, Beijing 100875, China
10
      <sup>3</sup>Key Laboratory of Ecosystem Network Observation and Modeling, Institute of Geographical
11
      Sciences and Natural Resources Research, Chinese Academy of Sciences, Beijing 100101,
12
13
      China
      <sup>4</sup>Department of Geography and Geographic Information Science, University of Illinois at
14
      Urbana-Champaign, Champaign, Illinois, USA
15
      <sup>5</sup>Department of Civil and Environmental Engineering, University of Illinois at
16
17
      Urbana-Champaign, Urbana, Illinois, USA
18
19
      Correspondence to: Guangyao Gao (gygao@rcees.ac.cn)
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      Abstract
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      Within China's Loess Plateau there have been concerted revegetation efforts and
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      engineering measures since the 1950s aimed at reducing soil erosion and land degradation.
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      As a result, annual streamflow, sediment yield and sediment concentration have all
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      decreased considerably. Human induced land use/cover change (LUCC) was the dominant
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factor, contributing over 70% of the sediment load reduction, whereas the contribution of 26 precipitation was less than 30%. In this study, we use 50-year time series data (1961-2011), 27 showing decreasing trends in the annual sediment loads of fifteen catchments, to generate 28 spatio-temporal patterns in the effects of LUCC and precipitation variability on sediment 29 30 yield. The space-time variability of sediment yield was expressed notionally as a product of two factors representing: (i) effect of precipitation and (ii) fraction of treated land surface 31 area. Under minimal LUCC, the square root of annual sediment yield varied linearly with 32 precipitation, with the precipitation-sediment load relationship showing coherent spatial 33 34 patterns amongst the catchments. As the LUCC increased and took effect, the changes of sediment yield pattern depended more on engineering measures and vegetation restoration 35 campaign, and the within-year rainfall patterns (especially storm events) also played an 36 37 important role. The effect of LUCC is expressed in terms of a sediment coefficient, i.e., ratio of annual sediment yield to annual precipitation. Sediment coefficients showed a 38 steady decrease over the study period, following a linear decreasing function of the fraction 39 of treated land surface area. In this way, the study has brought out the separate roles of 40 precipitation variability and LUCC in controlling spatio-temporal patterns of sediment 41 42 yield at catchment scale.

43

44 **1 Introduction**

45 Streamflow and sediment transport are important controls on biogeochemical processes
46 that govern ecosystem health in river basins (Syvitski, 2003). Changes in soil erosion on
47 landscapes and the resulting changes in sediment transport rates in rivers have great

environmental and societal consequences, particularly since they can be brought about by 48 climatic changes and human induced land use/cover changes (LUCC) (Syvitski, 2003; 49 50 Beechie et al., 2010). Understanding the dominant mechanisms behind such changes at different time and space scales is crucial to the development of strategies for sustainable 51 52 land and water management in river basins (Wang et al., 2016). In recent decades, streamflows and sediment yields in large rivers throughout the world 53 have undergone substantial changes (Milly et al., 2005; Nilsson et al., 2005; Milliman et al., 54 2008; Cohen et al., 2014). Notable decreases in sediment yields have been observed in 55 56 approximately 50% of the world's rivers (Walling and Fang, 2003; Syvitski et al., 2005). Many studies have investigated the dynamics of streamflows and sediment yields at 57 different spatial and temporal scales (Mutema et al., 2015; Song et al., 2016; Gao et al., 58 59 2016; Tian et al., 2016). In addition to climate variability, LUCC, soil and water conservation measure (SWCM) and construction of reservoirs and dams have substantially 60 contributed to the sediment load reductions (Walling, 2006; Milliman et al., 2008; Wang et 61 al., 2011). While previous studies have certainly provided valuable insights into the 62 streamflow and sediment load changes, the distinctive roles of LUCC and precipitation 63 variability in changing sediment loads still need further investigation in large domains and 64 across gradients of climate and land surface conditions (Walling, 2006; Mutema et al., 65 2015). A particularly useful approach to the development of generalizable understanding of 66 the effects of precipitation variability and LUCC is a comparative analysis approach 67 focused on extracting spatio-temporal patterns of sediment yields based on observations in 68 multiple locations within the same region, or even across different regions. This is 69

especially valuable and crucial in areas with severe soil erosion and fragile ecosystems, e.g.,
the Loess Plateau (LP) in China, which is the motivation for the work presented in this
paper.

The LP lies in the middle reaches of the Yellow River (YR) Basin, and contributes 73 nearly 90% of the YR sediment (Wang et al., 2016). The historically severe soil erosion in 74 the LP is due to sparse vegetation, intensive rainstorms, erodible loessial soil, steep 75 topography and a long agricultural history (Rustomji et al., 2008). To control such severe 76 soil erosion, several SWCMs, including terrace and check-dam construction, afforestation 77 78 and pasture reestablishment, have been implemented since the 1950s (Yao et al., 2011; Zhao et al., 2017). A large ecological restoration campaign, the Grain-for-Green (GFG) 79 project converting farmland on slopes exceed 25° to forest and pasture lands, was launched 80 81 in 1999 (Chen et al., 2015). Furthermore, the climate in the LP region has been showing both warming and drying trends (i.e. increased potential evapotranspiration and reduced 82 precipitation) since the 1950s (Zhang et al., 2016). 83 These substantial LUCC have notably altered the hydrological regimes of the LP in 84 combination with the climate change. Consequently, the sediment yields within the LP have 85 showed a predictable declining trend over the past 60 years (Zhao et al., 2017), resulting in 86 approximately a 90% decrease of sediment yield in the YR (Miao et al., 2010, 2011; Wang 87 et al., 2016). Many other studies have detected the influences of LUCC and precipitation 88 variability on sediment load changes within the LP. Rustomji et al. (2008) estimated that 89 the contribution of catchment management practices to the decrease of annual sediment 90

91 yield ranged between 64 and 89% for eleven catchments in the LP during 1950s-2000.

92	Zhao et al. (2017) examined the spatio-temporal variation of sediment yield from 1957 to
93	2012 across the LP, and indicated that the adoption of large-scale SWCMs led to significant
94	reduction of sediment yield between Toudaoguai and Tongguan stations and large
95	reservoirs operation played a critical role in sediment yield reduction between Tongguan
96	and Huayuankou stations. Zhang et al. (2016) pointed that the combined effects of climate
97	aridity, engineering projects and vegetation cover change have induced significant
98	reductions of sediment yield between 1950 and 2008. Wang et al. (2016) found that
99	engineering measures for soil and water conservation were the main factors for the
100	sediment load decrease between the 1970s and 1990s, but large-scale vegetation restoration
101	campaigns also played an important role in reducing soil erosion since the 1990s.
102	On the basis of the outcomes of these studies, it is now generally accepted that the
103	largest reductions of sediment yield within the LP resulted from LUCC. However, this is
104	general knowledge covering the whole region, and given the significant variability of
105	climate and catchment characteristics across the LP (Sun et al., 2015a; Sun et al., 2015b), it
106	is important to go further and explore how these might affect spatio-temporal patterns of
107	sediment yield. Exploration of these patterns is important for sustainable ecosystem
108	restoration and water resources planning and management within the LP. They will also
109	serve as the basis for future research aimed at the development of more generalizable
110	understanding of landscape and climate controls on sediment yields at the catchment scale.
111	Most of the sediment yield of the LP was produced in the Coarse Sandy Hilly
112	Catchments (CSHC) region (Fig. 1) located in the central region of the LP. The CSHC
113	supplied over 70% of total sediment load in the YR, especially coarse sand (Rustomji et al.,

114 2008). This region was the focus of our efforts to investigate the variation of sediment load from 15 catchments within the region within the LP. The specific objectives of this study 115 were, therefore, to: (1) attribute the temporal changes in sediment yield to changes in both 116 precipitation variability and LUCC over the entire study period (1961-2011) within the 117 118 CSHC region, (2) extract spatio-temporal trends in sediment yields on the basis of annual 119 sediment yield data, (3) separate the contributions of precipitation variability and fractional area of LUCC to the observed spatio-temporal patterns of sediment yields, and pave the 120 way for more detailed process-based studies in the future. 121

122 **2** Materials and methods

123 **2.1 Study area**

The CSHC region covers the area between the Toudaoguai and Longmen hydrological 124 125 stations in the mainstream of the YR (Fig. 1). The main stream that flows through the CSHC region is 733 km long and its drainage catchment covers 12.97×10^4 km², which is 126 accounting for 14.8% of the entire YR Basin. The CSHC region is characterized by arid to 127 semi-arid climate conditions. The annual precipitation in the region during 1961-2011 was 128 437 mm on average, and varied from lower than 300 mm in the northwest to 580 mm in the 129 southeast (McVicar et al., 2007). The precipitation that occurs during the flood season 130 (June-September) is usually in the form of rainstorms with high intensity and accounts for 131 72% of the annual rainfall total. Correspondingly, about 45% of the annual runoff and 88% 132 of the annual sediment yield within the region are produced during the flood season. The 133 northwestern part of the CSHC is relatively flat while the southeastern part is more finely 134 dissected (Rustomji et al., 2008). 135

136	Fourteen main catchments along a north-south transect within the CSHC study area
137	were chosen for the study (Fig. 1). These catchments account for 57.4% of the CSHC area,
138	and contribute about 70% and 72% of streamflow and sediment load of the overall CSHC,
139	respectively, based on observed hydrological data during 1961-2011 (Rustomji et al., 2008;
140	Yao et al., 2011). Characteristics of these catchments are presented in Table 1 and Fig. 2,
141	showing that the catchments present strong climate and land surface gradients. The
142	catchments in the northwestern part (#1-6) had relatively lower mean annual precipitation
143	(380 mm < \overline{P} <445 mm, where \overline{P} is mean annual precipitation over 1961-2011) and low
144	growing season (April-October) LAI (0.41 <lai<0.48, area="" index),<="" is="" lai="" leaf="" td="" the="" where=""></lai<0.48,>
145	while the corresponding values for catchments in the southeastern part (#7-14) were
146	470-570 mm and 0.63 <lai<3.26, respectively.<="" td=""></lai<3.26,>
147	The entire CSHC region is considered as an additional "catchment" and it is also
148	examined independently. The streamflow and sediment load for the whole region were
149	taken to be equal to the differences of corresponding measurements between the
150	Toudaoguai and Longmen gauging station. The average annual precipitation, streamflow
151	and sediment load of the region during 1961-2011 was 437.27 mm, 33.30 mm and 5.17 Gt,
152	respectively. Both the annual river discharge and sediment load across the region showed
153	significant decreasing trends (-0.82 mm yr ⁻¹ , p <0.001 and -0.19 Gt yr ⁻¹ , p <0.001,
154	respectively) over the past five decades, whereas precipitation decreased only slightly
155	$(-0.93 \text{ mm yr}^{-1}, p=0.25)$ (Fig. 3).

2.2 Data collection

157 Monthly streamflow and sediment load data during 1961-2011 were provided by the

Yellow River Conservancy Commission of China. Daily rainfall data from 1961 to 2011 at 158 66 meteorological stations in and around the region (Fig. 1) were obtained from the 159 National Meteorological Information Center of China. The spatially average of rainfall data 160 was carried out using the co-kriging interpolation algorithm with the DEM as an additional 161 162 input. The hydro-meteorological data (including annual precipitation, P [mm], streamflow, Q [mm], and sediment load, S [t]), specific sediment yield defined as SSY=S/A [t km⁻²], 163 where A is the drainage area of the hydrological station $[km^2]$, sediment concentration 164 defined as SC=S/(Q.A) [kg m⁻³] and the sediment coefficient defined as $C_s=SSY/P$ [t km⁻²] 165 mm⁻¹] were estimated for each catchment. 166 The mean catchment slope gradient based on the ASTER GDEM data with a resolution of 167 30 m and soil data (scale 1:500,000) were provided by the National Earth System Science 168 169 Data Sharing Infrastructure (http://www.geodata.cn). The land use information as at 1975, 1990, 2000 and 2010 was determined with Landsat MSS and TM remote sensing images at a 170 spatial resolution of 30 m. Six land use types were classified, i.e., forestland, cropland, 171 grassland, construction land, water body, and barren land. The LAI data during 1982-2011 172 were obtained from the Global Land Surface Satellite (GLASS) NDVI Series with spatial 173 resolution of 1 km (www.landcover.org, Zhao et al., 2013). The total areas impacted by 174 various SWCMs (i.e., afforestation, grass plantation, terraces and check-dams) in each 175 catchment during1960s-2000s were obtained from Yao et al. (2011). 176 2.3 Trend test 177

178 The non-parametric Mann-Kendall (M-K) test method proposed by Mann (1945) and

179 Kendall (1975) was used to determine the significance of the trends in annual

meteorological and hydrological time series. A precondition for using the MK test is to remove the serial correlation of climatic and hydrological series. In this study, the trend-tree pre-whitening (TFPW) method of Yue and Wang (2002) was used to remove the auto-correlations before the trend test. There was no residual autocorrelation remaining after performing the TFPW. A *Z*-statistic was obtained from the M-K test on the whitened series. A negative value of *Z* indicates a decrease trend, and vice versa. The magnitude of the slope of the trend (β) was estimated by (Sen, 1968; Hirsch et al., 1982):

187
$$\beta = \operatorname{Median}\left[\frac{x_j - x_i}{j - i}\right] \quad \text{for all } i < j \tag{1}$$

188 where x_i and x_j are the sequential data values in periods *i* and *j*, respectively.

189 **2.4 Attribution analysis of changes in sediment yield**

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The time-trend analysis method was used to determine the quantitative contributions of 190 191 LUCC and precipitation variability to sediment yield changes. This method is primarily 192 designed to determine the differences in hydrological time series between different periods (reference and validation periods) with different LUCC conditions (Zhang et al., 2011). In 193 194 this method, a regression equation between precipitation and sediment yield is developed and evaluated during the reference period, and the established equation is then used to 195 estimate sediment yield during the validation period. The difference between measured and 196 197 predicted sediment yields during the validation period represents the effects of LUCC, and the residual changes are caused by precipitation variability. The governing equations of the 198 time-trend analysis method can be expressed as: 199

$$SSY_1 = f(P_1) \tag{2}$$

$$SSY_2' = f(P_2) \tag{3}$$

$$\Delta SSY^{\text{LUCC}} = \overline{SSY_2} - \overline{SSY_2'} \tag{4}$$

$$\Delta SSY^{Pre} = \left(\overline{SSY_2} - \overline{SSY_1}\right) - \Delta SSY^{LUCC}$$
(5)

where SSY' is the predicted sediment yield, subscripts 1 and 2 indicate the reference and 204 validation periods, respectively. $\overline{SSY_1}$ and $\overline{SSY_2}$ represent mean measured sediment yield 205 during the reference and validation periods, respectively, and $\overline{SSY'_2}$ represents mean 206 predicted sediment yield during the validation period. ΔSSY^{LUCC} and ΔSSY^{Pre} are sediment 207 yield changes during the validation period associated with LUCC and precipitation 208 variability, respectively. Rustomji et al. (2008) found that the square root of annual 209 210 sediment yield in the catchments of the Loess Plateau was linearly related to annual precipitation. This was, therefore, used in this study as the motivation to develop the 211 precipitation-sediment yield relationship during the reference period: 212

213
$$\sqrt{SSY} = aP + b \tag{6}$$

In this study, the full data period of 1961-2011 was divided into three phases (1961-1969, 1970-1999 and 2000-2011). The first period was considered the reference period as the effects of human activities were slight and could be ignored (Wang et al., 2016). During the second stage, numerous SWCMs were implemented. For the third stage, a large ecological restoration campaign (GFG project) was launched in 1999.

219 **3 Results and discussion**

220 **3.1 Changes of land use/cover**

221 The CSHC region has undergone extensive LUCC caused by the implementation of

- 222 SWCM and vegetation restoration projects (e.g., the GFG project). Fig. 4 shows the
- distribution of land use types of the region in 1975, 1990, 2000 and 2010. More than 90%

224	of the whole area was occupied by the cropland, forestland and grassland. The area of
225	cropland decreased by 26.72% and forestland increased by 53.15%, and there was no
226	significant change for the area of grassland (increase of 4.21%) in the CSHC region from
227	1975-2010. The majority of changes occurred during 2000-2010 due to the GFG
228	(reforestation) project (26.67% decrease and 36.21% increase for cropland and forestland,
229	respectively). The transition from cropland to forestland was greater in the catchments of
230	the southeastern part (especially in catchments #7-#9) than that in the northwestern part
231	(Fig. 4). In the period 1975 to 2000, the increase of forestland was 26.34% and 4.55% in
232	the southeastern and northwestern part, respectively, and the change of cropland was
233	negligible (only -0.39% and 0.22%, respectively). During 2000-2010, the forestland
234	increased by 47.79% and 18.30%, and the cropland decreased by 44.84% and 21.04% in
235	the southeastern and northwestern part, respectively.
236	The SWCMs implemented in the LP included both biotic treatments (e.g., afforestation
237	and grass-planting) and engineering measures (e.g., construction of terrace and check-dam
238	and gully control projects). Afforestation, grass-planting and construction of terraces were
239	seen as the slope measures, while building of check-dams and gully control projects were
240	the measures on the river channel. Although the area utilized for engineering measures was
241	much smaller than the biotic treatments, they immediately and substantially trap
242	streamflow and sediment load. The fraction of the treated area (area treated by erosion
243	control measures relative to total catchment area) increased from 3.95% in the 1960s to
244	28.61% in the 2000s (Fig. 5). The increase of the treated area was greatest during the 1980s
245	as a result of comprehensive management of small watersheds and the 2000s due to the

246	GFG project since 1999. Some decreases of SWCM areas (i.e. afforestation and
247	check-dams) occurred during the 1990s (Fig. 5) as some planted trees were died due to
248	drought and some small and medium check-dams were fully deposited by sediment and
249	then subsequently destroyed by floods.
250	The growing season LAI of the whole region changed from 0.74 during 1982-1999 to
251	0.81 during 2000-2011, an increase of 10.16% (Fig. 5). The LAI did not show significant
252	increase during 1982-1999 (0.003 yr ⁻¹ , $p=0.11$), and it increased significantly during
253	2000-2011 (0.024 yr ⁻¹ , p <0.01). The increase of growing season LAI during 1982-2011
254	was greater for the catchments in the southeastern part (0.009 yr^{-1}) compared to the
255	northwestern part (0.004 yr ⁻¹), especially after 2000 (Fig. 6). In the period from 1982-1999
256	to 2000-2011, the average increase of growing LAI of the fourteen sub-catchments was
257	0.088 yr^{-1} (0.010-0.183 yr}{-1}), with the increase of 0.114 yr $^{-1}$ and 0.053 yr $^{-1}$ in the
258	southeastern and northwestern part, respectively.
259	3.2 Trends of hydro-meteorological and sediment yield variables
260	Table 2 shows the trends in annual P , Q , SSY , SC and C_s of the fifteen catchments during the
261	period 1961-2011. The annual P showed a decline trend in all catchments except the Jialu
262	catchment, but the changing trend is only significant in the Xinshui and Zhouchuan
263	catchments ($p < 0.05$). The annual Q, SSY, SC and C _s showed significant decreasing trends in
264	all the catchments, and most of the decreases were at the 0.001 significance level. For the
265	fourteen sub-catchments, the average decrease rates of annual values of Q , SSY, SC and C_s
266	were 0.86 mm yr ⁻¹ (0.24-1.66 mm yr ⁻¹), 190.06 t km ⁻² yr ⁻¹ (26.47-398.82 t km ⁻² yr ⁻¹), 2.73
267	kg m ⁻³ yr ⁻¹ (0.69-4.70 kg m ⁻³ yr ⁻¹) and 0.38 t km ⁻² mm ⁻¹ yr ⁻¹ (0.04-0.87 t km ⁻² mm ⁻¹ yr ⁻¹),

268	respectively. The changing rates of Q , SSY, SC and C_s for the whole region were -0.85 mm
269	yr^{-1} , -131.52 t km ⁻² yr^{-1} , -2.06 kg m ⁻³ yr^{-1} and -0.27 t km ⁻² mm ⁻¹ yr^{-1} , respectively. The annual
270	average reductions in the whole region were equivalent to 2.56%, 3.30%, 2.01% and 3.07%
271	of the mean annual values of Q , SSY, SC and C_s , respectively.
272	The mean and the coefficient of variation, C_v , representing inter-annual variability of
273	annual values of P , Q , SSY , SC and C_s for the fifteen catchments during the three phases
274	(reference period-1, period-2 and period-3) are shown in Fig. 7. Compared to standard
275	deviation, the C_v value was better able to indicate the inter-annual variability of precipitation,
276	streamflow and sediment load among the catchments with distinctly different average values.
277	Compared to the reference period, the mean annual precipitation decreased by 11.73%
278	(6.36%-15.69%) and 10.64% (5.88%-16.7%) on average in period-2 and period-3,
279	respectively. From period-2 to period-3, the change of mean annual precipitation was slight
280	(increased by 1.32% on average) with a decrease of 2.45%-5.87% in four catchments and an
281	increase in the remaining catchments (0.35% - 8.29%). The variability of annual P also
282	decreased as indicated by the reductions of C_v values during period-2 and period-3 (Fig. 7a).
283	In contrast to annual P , the reductions of mean annual Q , SSY, SC and C_s were clearly more
284	evident. With respect to the reference period, the reduction was 34.41% (9.45%-54.72%),
285	48.02% (17.98%-67.61%), 24.20% (-9.93%-47.77%) and 39.31% (4.64%-63.5%) for <i>Q</i> , <i>SSY</i> ,
286	SC and C_s during period-2, and the decreasing rate was even more in period-3 with values of
287	64.82% (36.72%-84.19%), 88.23% (64.94%-97.64%), 67.81% (17.28%-91.12%) and 85.85%
288	(63.51%-96.97%), respectively. C_v of annual Q increased in eight catchments, with the
289	remaining ones showing decreasing trends (Fig. 7b), while C_v values for SSY, SC and C_s

increased in all catchments (Figs 7c-7e). The above results indicate substantially different
behaviors of the changes among precipitation, streamflow and sediment load.

292 **3.3 Quantitative attribution of sediment yield decline**

The effects of precipitation change and LUCC on sediment yield reductions in period-2 and 293 294 period-3 were quantified using Eqs. (2-6) and the results are shown in Fig. 8. The form of 295 Eq. (6) during the reference period is shown in Table 3. The analysis showed that both decreased precipitation and increased area treated with erosion control measures 296 contributed to the observed sediment load reduction, and that LUCC played the major role. 297 298 On average, LUCC and precipitation change contributed 74.39% and 25.61%, respectively, to sediment load reduction from the reference period to period-2, with their respective 299 contributions to sediment load reduction from the reference period to period-3 being 88.67 300 301 and 11.33%. The effect of LUCC in period-3 was greater than in period-2 as the land use/cover (see Figs. 4-5) and vegetation coverage (see Fig. 6) had undergone substantial 302 changes due to the ecological restoration campaigns launched during period-3. From 303 period-2 to period-3, the contribution of precipitation was negative for sediment yield 304 reduction in eleven catchments where the annual precipitation slightly increased and thus 305 the contribution of LUCC was larger than 100% (Fig. 8c). In the remaining four catchments, 306 the average contribution of LUCC increased to 83.96%. 307 In broad terms there are two factors that govern annual sediment yield of a catchment: 308 precipitation and landscape properties (soil, topography and vegetation). Precipitation is the 309 primary driver of runoff and, therefore, directly influences the sediment transport capacity 310 of streamflow and sediment yield at the catchment scale. Higher precipitation means higher 311

streamflow, which is the immediate driver of erosion and sediment transport. Landscape 312 properties not only have an impact on the volume or intensity of streamflow, but also 313 determine the erodibility of the soil. Correlations between the potential factors 314 (precipitation, percentage area of afforestation, pasture plantation, terracing, check-dams 315 316 and construction land, and LAI) and sediment yield change between different stages (see 317 Table 4) showed that check-dam construction was the dominant factor for sediment yield reduction from reference period to period-2. Pasture plantation and check-dam construction 318 acted as the dominant factors for sediment yield from reference period to period-3. The 319 320 increase of precipitation mitigated the reduction of sediment yield to some degree from period-2 to period-3. 321

Based on the above results, the variation of SSY mainly depended on precipitation in the 322 323 reference period before LUCC took effect and any spatial patterns of SSY in the catchments were controlled by differences in annual precipitation and land surface conditions. During 324 the validation period (period-2 and period-3) when increased LUCC had taken effect, SSY 325 decreased considerably. The decrease of precipitation was insignificant and LUCC 326 contributed over 70% of the sediment yield reduction. In this case, the temporal changes of 327 SSY depended more on the fraction of treated surface area and precipitation possibly played 328 a secondary role. The spatial pattern of the impacts of precipitation on sediment yield was 329 dependent on the landscape properties among catchments. Guided by this framework, data 330 were next analysed to generate separate spatial and temporal patterns constituting 331 332 respective components of the spatio-temporal patterns.

333 3.4 Spatial-temporal pattern of the impacts of precipitation on sediment yield

334	The regression equations of $\sqrt{SSY} = aP + b$ are shown in Table 3. The spatial distributions
335	of precipitation-sediment relationships during the three stages are shown in Fig. 9. During
336	the reference period, the correlation between precipitation and sediment yield was
337	significant in eleven catchments ($p < 0.05$) with the coefficient of determination (R^2) ranged
338	from 0.48 to 0.87 (Table 3). Furthermore, the precipitation-sediment yield relationship
339	varied from catchment to catchment and showed a spatial pattern. The correlation
340	coefficient between precipitation and sediment yield was greater for catchments in the
341	northwestern part with average R^2 value of 0.75 and p value of 0.007 compared to those in
342	the southeastern part where the average R^2 and p values were 0.48 and 0.059, respectively
343	(Table 3). Based on the slopes of the regression equations between annual precipitation and
344	sediment yield, the fourteen catchments were classified into four groups (Group-1: $a>0.3$,
345	Group-2: 0.2< <i>a</i> <0.3, Group-3: 0.1< <i>a</i> <0.2 and Group-4: 0< <i>a</i> <0.1), which indicate that the
346	sediment production capability of annual precipitation is different among the catchments
347	(Fig. 9a). The four catchments in the northwestern part (#1-3 and 5) had the greatest
348	regression slopes of $a>0.3$ (Group-1) and the Shiwang catchment had the lowest regression
349	slope of 0.07 (Group-4). Most of the catchments in the southeastern part were in the second
350	group of $0.2 < a < 0.3$. Overall, the regressed equations were significant for most of the
351	catchments, and were suitable for estimating the relative contributions of LUCC and
352	precipitation variability to sediment yield changes.
353	Compared to the reference period, the correlation between precipitation and sediment
354	yield during the period-2 decreased in the catchments, as indicated by lower R^2 values in

Table 3. The slopes of the regression lines in the period-2 decreased in most of the

356	catchments with respect to the reference period, except in Huangfu, Gushan and Kuye
357	catchments which increased slightly. Furthermore, the spatial patterns of the
358	precipitation-sediment yield relationship during these two periods were somewhat different
359	(Figs. 9a and 9b). From the reference period to period-2, Jialu catchment moved from
360	Group-1 to Group-2 and five catchments moved from Group-2 to Group-3.
361	During period-3, the correlation between precipitation and sediment yield was weaker
362	compared to the reference period and period-2 (Table 3). The relationships between
363	precipitation and sediment yield were not significant in all the catchments (Table 3). The
364	slopes of the regression lines during period-3 decreased sharply (Table 3). Six catchments
365	(five in the north-western part and one in the south-eastern part) had negative regression
366	slopes (Fig. 9c). This result indicates that the sediment production capability of annual
367	precipitation decreased greatly during period-3, and the increase of precipitation amount in
368	some catchments did not lead to increased sediment yield. Furthermore, the spatial patterns
369	of precipitation-sediment relationship during period-3 were clearly different from those
370	during the reference period and period-2 (compare Fig. 9c against Figs. 9a-9b). There were
371	only three groups with two catchments having regression slopes of $0.1 \le a \le 0.2$, six
372	catchments having regression slopes of $0.1 \le a \le 0.2$ and six catchments having negative
373	regression slopes.
374	The aforementioned analysis of the precipitation-sediment yield relationship in
375	different periods clearly indicates that the impacts of precipitation on sediment yield
376	declined with time. The impacts were different among catchments, with a clear spatial
377	pattern. The effects of precipitation on the sediment yield were greater in the north-western

part compared to those in the south-eastern part. The decreased effects of precipitation on
sediment yield with time were consistent with the significant reductions of sediment
coefficient (Table 2) and the decreased contribution of precipitation to sediment load
reduction (25.61% and 11.33% in period-2 and period-3, respectively). During period-2,
the LUCC were mainly induced by SWCM, especially engineering measures. During
period-3, the combined effects of substantial vegetation cover and conservation measures
further weakened the effects of precipitation on sediment load reduction.

385 **3.5 Spatial-temporal pattern of the impacts of land use/cover on sediment yield**

In order to quantify the effects of SWCM on sediment load reduction, the relationships between the decadal sediment coefficient and the fraction of area treated with erosion control measures in the 15 catchments were analysed and the results are presented in Table 5. The decadal sediment coefficient (\overline{SC}) decreased linearly with the fraction of treated land surface area (A_c) in all catchments:

$$SC = -mA_c + n \tag{7}$$

The correlations were significant in eleven catchments ($p \le 0.05$) with R^2 ranging from 392 0.78 to 0.99 (Table 5). The effects of SWCM on sediment load change show a spatial pattern. 393 The correlation between sediment coefficients and conservation measures were stronger in 394 catchments located in the north-western part compared to that in the south-eastern part (Table 395 5). Based on the slope of the regression equation between the sediment coefficient and 396 fraction of the treated area, the catchments were classified into three groups in Fig. 10 397 (Group-1: 0.8<*m*<1.2, Group-2: 0.4<*m*<0.8 and Group-3: 0<*m*<0.4), which indicated that the 398 degree of sediment load impacted by conservation measures was different among the 399

catchments. The average *m* value was 0.73 and 0.37 for the catchments in the north-western
and south-eastern part, respectively. Half of the catchments in the north-western part were in
Group-1 and the other half were in Group-2, whereas six of the eight catchments in the
south-eastern part were in Group-3 with lowest regression slope.

404 **3.6 Discussion**

Differences in catchment characteristics, including land use/cover, soil properties and 405 topography, as well as precipitation characteristics, are clearly the reason for the spatial 406 patterns in the precipitation-sediment yield relationship (Morera et al., 2013; Mutema et al., 407 408 2015). The lower vegetation cover was the main reason for the greater effects of precipitation on sediment yield in the northwestern part. In order to fully explore this, the 409 mapping of information of catchment characteristics into sediment yield models and 410 411 simulations under different climate scenarios would be needed (Ma et al., 2014; Achete et al., 2015). In this context, the inter-annual and intra-annual patterns of variability of 412 precipitation, including the distribution of storm events, may also contribute to the 413 observed spatial patterns of precipitation-sediment yield relationship. 414 As LUCC took effect during period-2 and period-3, and despite the much reduced role 415 of precipitation in driving changes in sediment yield, within-year temporal rainfall patterns 416 did play an important role in the observed changes of sediment yield, given that most of the 417 sediment yield was produced during a few key storm events. The correlation between 418 419 sediment yield and storm events with daily precipitation amount larger than 20 mm (including storm numbers, precipitation amount of storms) in the CSHC region during 420 different decades were investigated (see Table 6). The analysis showed that the sediment 421

yield was significantly correlated with storm numbers in the 1960s, 1970s and 1980s 422 (p < 0.05), and precipitation amount of storms in the 1960s and 1970s (p < 0.05). This result 423 indicated the critical role of storm events in sediment yield, especially during the periods 424 before substantial LUCC took effect. 425 Looking into this in more detail and taking the Yanhe catchment as an example, the 426 precipitation amount during the rainy season (May-October when sediment load was 427 measured) in 2003 and 2004 was 514.31 mm and 389.05 mm, respectively, whereas the 428 sediment load in 2004 (2427.37 \times 10⁴ t) was about over four times of that in 2004 429 $(590.04 \times 10^4 \text{ t})$. As shown in Fig. 11, there were six days with precipitation amounts over 430 20 mm and the maximum daily precipitation amount on 25th August was 27.85 mm in 2003, 431 and the values in 2004 were five days and 46.34 mm on 10th August. Furthermore, heavy 432 rainfall events were distributed in every month in 2003, whereas they were concentrated in 433 July and August in 2004. There were five evident peaks of sediment load with the sum of 434 1646.24×10^4 t (67.82% of annual total) in 2004, especially the one on 10^{th} August 435 produced 784.53×10^4 t sediment load (32.32% of annual total) (Fig. 11b). In contrast, there 436 were three peaks of sediment load in 2003, and the maximum value was only 139.97×10^4 t 437 (Fig. 11a). Therefore, apart from annual precipitation amounts, within-year rainfall patterns 438 should also be considered when investigating the effects of precipitation on 439 temporal-spatial changes of streamflow and sediment load. 440 The sediment load reductions in the CSHC region were primarily caused by the LUCC 441 and the implementation of SWCM. The cropland area decreased 9733.91 km² (8.73% of 442 region area) and the forestland area increased 7662.50 km² (6.87% of region area) in the 443

region from 1975 to 2010. Most of the increase in forestland area was converted from 444 cropland area induced by the GFG or reforestation project. As a result of the land use change, 445 vegetation cover increased greatly and it substantially contributed to the decreases of runoff 446 and sediment production. The SWCMs, such as afforestation and engineering measures were 447 448 the major interventions in the study area to reduce the runoff-sediment generation from precipitation and retain streamflow and sediment load within the catchment. Establishing 449 perennial vegetation cover was considered as one of the most effective measures to stabilize 450 soils and minimize erosion (Farley et al., 2005; Liu et al., 2014). It was reported that both 451 452 runoff coefficient and sediment concentration of catchments in the LP decreased significantly and linearly with the vegetation cover (Wang et al., 2016). The engineering structures mainly 453 included creation of terrace and building of check-dams and reservoirs, which reduced flood 454 455 peaks and stored water and sediment within the catchment. There were about 59, 874 check-dams in the region which trapped about 9842×10^4 t of sediment per year 456 (approximately 19% of annual sediment yield) during the past six decades (Yao et al., 2011). 457 Over time, the effectiveness of engineering measures decreased as they progressively filled 458 with sediments, and vegetation restoration played a greater role in controlling soil erosion. 459 Conclusions 460 4

Through analyses of hydrological and sediment transport data, this study has shown that long-term decreasing trends in sediment loads across fifteen large sub-catchments located in the CSHC region for the period 1961-2011. The study was particularly aimed at extracting spatio-temporal patterns of sediment yield and attributing these patterns to the broad hydro-climatic and landscape controls. The effects of precipitation variability and

466 land use/cover changes on sediment yield were investigated in detail.

467	Over the study period, the total area undergoing erosion control treatment went up
468	from only 4% to over 30%. This included to decrease of cropland by 27%, increase of
469	forestland by 53% and grassland by 4% from 1975-2010. Over the same period annual
470	precipitation decreased by not more than 10%. As a result of the erosion control measures,
471	there were major reductions in streamflow (65%), sediment yield (88%), sediment
472	concentration (68%) and sediment efficiency, i.e., annual sediment yield/annual
473	precipitation (86%) over the entire 50-year period.
474	The observed data in the 15 study catchments also exhibited interesting
475	spatio-temporal patterns in sediment yield. The study attempted to separate the relative
476	contributions of annual precipitation and LUCC to these spatio-temporal patterns. Before
477	LUCC took effect, the data indicates a linear relationship between square root of annual
478	sediment yield and annual precipitation in all 15 catchments, with highly variable slopes of
479	the relationship between the catchments, which exhibited systematic spatial patterns, in
480	spite of some scatter. As LUCC increased and took effect, the scatter increased and the
481	slopes of the sediment yield vs precipitation relationship became highly variable and lost
482	any predictive power. The study then looked at the controls on sediment coefficient instead
483	of sediment yield, thus eliminating the effect of precipitation and enabling a direct focus on
484	landscape controls. The results of this analysis found that sediment coefficient was heavily
485	dependent on the area under land use/cover treatment, exhibiting a linear decreasing
486	relationship. Even here, there was a considerable variation in the slope of the relationship
487	between the 15 catchments, which exhibited a systematic spatial pattern.

488	Preliminary analyses presented in this study suggest that much of the sediment yield in
489	the LP may be caused during only a few major storms. Therefore, the seasonality and
490	intra-annual variability of precipitation may play important roles in annual sediment yield,
491	which may also explain the spatial patterns of sediment yield and the effects of the various
492	LUCC. Also, the precipitation threshold for producing sediment yield would have increased
493	greatly as a result of SWCM and vegetation restoration in the LP. Exploration of these
494	questions in detail will require a more physically based model that can account for fine
495	scale rainfall variability and catchment characteristics. This is the next immediate step in
496	our investigations, and will be reported on in the near future.
497	
498	Data availability. All the data used in this study are available upon request.
499	
500	Competing interests. The authors declare that they have no conflict of interest.
501	
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607	Figure	captions
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- Figure 1. Location of the studied catchments in the Coarse Sandy Hilly Catchments(CSHC) region within the Loess Plateau.
- 610 Figure 2. Spatial distribution of (a) annual precipitation (1961-2011), (b) growing season
- 611 leaf area index (LAI, 1982-2011), (c) soil type and (d) slope in the study area.
- Figure 3. Annual precipitation, streamflow and sediment load for the whole CSHC regionduring 1961-2011.
- 614 **Figure 4.** Land use and cover of the study area in (a) 1975, (b) 1990, (c) 2000 and (d)

615 **2010**.

- Figure 5. The changes of soil and water conservation measures area and growing season
 LAI in the study area.
- 618 Figure 6. Long-term trends in growing season LAI changes over (a) 1982-2011, (b)
- 619 1982-1999 and (c) 2000-2011 in the study area. Inset in each figure shows the
- 620 frequency distribution of the LAI trends.
- 621 Figure 7. The changes of (a) precipitation, (b) streamflow, (c) sediment yield, (d) sediment
- concentration and (e) sediment coefficient during different stages (1961-1969,

623 1970-1999 and 2000-2011).

- 624 Figure 8. Contributions of precipitation and land use/cover to reductions of sediment load
- from (a) reference period (P1) to period-2 (P2), (b) reference period (P1) to period-3 (P3)
- and (c) period-2 (P2) to period-3 (P3).
- 627 **Figure 9.** Spatial distribution of slope *a* in the regression equation $\sqrt{SSY} = aP + b$ during
- 628 (a) reference period (1961-1969), (b) period-2 (1970-1999) and (c) period-3

- 629 (2000-2011). *SSY* is specific sediment yield, and *P* is precipitation.
- 630 **Figure 10.** Spatial distribution of slope *m* in the regression equation $\overline{SC} = -mA_c + n$. \overline{SC} is
- 631 the decadal average sediment coefficient, and A_c is the percentage of the area affected by
- soil and water conservation measures in the catchments.
- **Figure 11.** Daily precipitation and sediment load of the Yanhe catchment during rainy
- 634 season (May-October) in (a) 2003 and (b) 2004.

 Table 1. Long-term hydrometeorological characteristics (1961-2011) and growing season leaf area index (LAI) (1982-2011) of the studied catchments in the Loess Plateau.

ID Catchment			Slope	Area	Annual average						
ID	Catchment	Gauging station	(°)	(km ²)	P (mm)	Q (mm)	$\frac{SSY}{(t \text{ km}^{-2})}$	SC (kg m ⁻³)	$\frac{C_s}{(t \text{ km}^{-2} \text{ mm}^{-1})}$	LAI	
1	Huangfu	Huangfu	7.8	3175	388.95	36.34	11608.86	275.90	27.35	0.412	
2	Gushan	Gaoshiya	9.8	1263	422.49	49.55	12398.68	189.57	25.98	0.440	
3	Kuye	Wenjiachuan	6.3	8515	394.63	59.25	9099.60	114.99	21.17	0.427	
4	Tuwei	Gaojiachuan	5.8	3253	402.82	97.53	4454.47	38.44	10.16	0.406	
5	Jialu	Shenjiawan	10.4	1121	445.51	49.22	9645.19	142.19	20.03	0.480	
6	Wuding	Baijiachuan	6.8	29662	384.32	36.39	3089.61	74.09	7.67	0.460	
7	Qingjian	Yanchuan	15.9	3468	485.58	38.93	8747.17	190.57	17.35	0.626	
8	Yanhe	Ganguyi	16.5	5891	516.09	34.08	6604.90	166.31	12.45	0.920	
9	Shiwang	Dacun	15.2	2141	572.16	32.99	798.89	20.32	1.31	3.261	
10	Qiushui	Linjiaping	13.0	1873	469.02	34.83	7818.21	185.79	15.75	0.938	
11	Sanchuan	Houdacheng	14.6	4102	486.23	50.37	3444.56	53.39	6.63	1.887	
12	Quchan	Peigou	14.6	1023	539.73	30.24	7492.57	192.01	13.68	0.934	
13	Xinshui	Daning	14.0	3992	529.96	29.22	3004.96	86.81	5.23	1.752	
14	Zhouchuan	Jixian	15.3	436	530.06	30.13	4951.15	107.99	8.55	1.165	
15	CSHC	Toudaoguai and Longmen	10.5	129654	437.27	33.30	3988.04	102.42	8.73	0.765	

ID	Catchment -	Р		Q		SSY		SC		C_s		
ID		Ζ	β (mm yr ⁻¹)	Ζ	β (mm yr ⁻¹)	Ζ	$\beta (t \text{ km}^{-2} \text{ yr}^{-1})$	Ζ	$\beta (\mathrm{kg}\mathrm{m}^{-3}\mathrm{yr}^{-1})$	Ζ	β (t km ⁻² mm ⁻¹ yr ⁻¹)	
1	Huangfu	-0.57 ^{ns}	-0.52	-4.82***	-0.99	-4.50***	-323.24	-1.97*	-2.58	-4.71***	-0.80	
2	Gushan	-0.78 ^{ns}	-1.16	-5.02***	-1.47	-4.90***	-398.82	-3.75***	-3.92	-5.15***	-0.87	
3	Kuye	-0.49 ^{ns}	-0.37	-5.98***	-1.66	-5.41***	-288.83	-4.61***	-3.22	-5.60***	-0.63	
4	Tuwei	-0.24 ^{ns}	-0.27	-7.88***	-1.57	-5.20***	-130.34	-4.37***	-0.98	-5.59***	-0.30	
5	Jialu	0.19 ^{ns}	0.26	-7.55***	-1.42	-5.36***	-298.10	-3.80***	-3.89	-5.60***	-0.69	
6	Wuding	-0.39 ^{ns}	-0.37	-6.60***	-0.54	-4.55***	-79.19	-3.33***	-1.35	-4.94***	-0.20	
7	Qingjian	-0.73 ^{ns}	-0.56	-2.06*	-0.24	-3.01**	-138.54	-3.09**	-3.53	-2.73**	-0.30	
8	Yanhe	-1.19 ^{ns}	-1.17	-3.22**	-0.34	-3.36***	-115.18	-3.30***	-3.07	-3.10**	-0.22	
9	Shiwang	-1.20 ^{ns}	-1.50	-4.01***	-0.61	-6.26***	-26.47	-5.43***	-0.69	-6.12***	-0.04	
10	Qiushui	-0.28 ^{ns}	-0.35	-5.80***	-0.97	-6.98***	-290.44	-5.00***	-4.00	-5.98***	-0.55	
11	Sanchuan	-1.43 ^{ns}	-1.71	-6.09***	-0.96	-5.35***	-108.69	-5.13***	-1.60	-5.99***	-0.21	
12	Quchan	-0.94 ^{ns}	-1.14	-3.23**	-0.42	-3.65***	-173.16	-3.72***	-4.12	-3.46***	-0.29	
13	Xinshui	-2.37*	-2.71	-5.57***	-0.70	-5.92***	-106.30	-3.77***	-1.92	-5.60***	-0.19	
14	Zhouchuan	-2.21*	-2.48	-7.20***	-0.79	-5.86***	-183.49	-6.73***	-4.70	-7.12***	-0.35	
15	CSHC	-0.67 ^{ns}	-0.55	-5.91***	-0.85	-5.70***	-131.52	-4.26***	-2.06	-5.67***	-0.27	

Table 2. Mann-Kendall trend analysis results for the annual precipitation (*P*), streamflow (*Q*), specific sediment yield (*SSY*), sediment concentration (*SC*), sediment coefficient (C_s) during 1961-2011.

^a ***, ** and * indicate the significance levels of 0.001, 0.01 and 0.05, respectively. ns indicates the significance levels exceeds 0.05.

Table 3. The linear regression equations between square root of specific sediment yield and annual precipitation ($\sqrt{SSY} = aP + b$) during three stages (1961-1969, 1970-1999 and 2000-2011).

ID	Catalument	Reference period (1961-1969)			Period-2 (1970-1999)			Period-3 (2000-2011)			
ID Catchi	Catchment	Regression equation	R^2	р	Regression equation	R^2	р	Regression equation	R^2	р	
1	Huangfu	y = 0.341x + 12.041	0.78	0.002	y = 0.397x - 11.454	0.40	0.000	y = 0.135x + 5.842	0.12	0.277	
2	Gushan	<i>y</i> =0.349 <i>x</i> +8.237	0.84	0.001	y = 0.354x - 5.627	0.37	0.000	y = 0.076x + 10.415	0.09	0.344	
3	Kuye	y = 0.323x + 9.939	0.67	0.007	y = 0.325x - 3.904	0.35	0.001	y = 0.037x + 8.208	0.03	0.564	
4	Tuwei	y = 0.218x + 12.635	0.87	0.000	y = 0.188x + 1.648	0.22	0.008	y = -0.030x + 27.644	0.03	0.613	
5	Jialu	y = 0.382x + 6.976	0.78	0.004	y = 0.222x + 11.867	0.13	0.049	y = 0.072x + 7.131	0.03	0.616	
6	Wuding	y = 0.174x + 20.544	0.53	0.027	y = 0.151x + 7.546	0.26	0.004	y = 0.107x - 1.511	0.17	0.182	
7	Qingjian	y = 0.232x + 20.923	0.48	0.040	y = 0.173x + 29.319	0.16	0.027	y = 0.096x + 8.344	0.05	0.522	
8	Yanhe	y = 0.243x + 0.741	0.39	0.070	y = 0.126x + 32.699	0.16	0.031	y = 0.006x + 39.338	0.00	0.973	
9	Shiwang	y = 0.070x + 10.935	0.27	0.150	y = 0.079x-7.837	0.24	0.006	y = -0.007x + 9.426	0.01	0.769	
10	Qiushui	y = 0.257x + 30.738	0.60	0.014	y = 0.239x-2.814	0.29	0.002	y = -0.111x + 72.39	0.06	0.448	
11	Sanchuan	y = 0.191x + 15.053	0.36	0.089	y = 0.174x - 9.652	0.42	0.000	y = -0.056x + 37.680	0.06	0.432	
12	Quchan	y = 0.202x + 34.590	0.72	0.016	y = 0.132x + 29.685	0.09	0.104	y = -0.199x + 119.247	0.11	0.300	
13	Xinshui	y = 0.202x - 6.593	0.71	0.004	y = 0.184x - 17.464	0.53	0.000	y = 0.015x + 16.822	0.01	0.823	
14	Zhouchuan	y = 0.207x + 20.226	0.33	0.090	y = 0.245x - 31.399	0.32	0.001	y = -0.035x + 26.145	0.06	0.460	
15	CSHC	y = 0.218x + 5.689	0.70	0.005	y = 0.174x + 2.912	0.35	0.001	y = 0.001x + 24.996	0.00	0.994	

Table 4. The regression models for sediment yield change (ΔSSY) in d	lifferent stages.
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Period	Regression model	R^2	р
Reference period vs. Period-2	<i>ΔSSY</i> =-0.135-0.850 ×ΔDam	0.886	0.000
Reference period vs. Period-3	$\Delta SSY = -0.067 - 0.659 \times \Delta Dam - 0.081 \times \Delta Pasture$	0.928	0.023
Period-2 vs. Period-3	$\Delta SSY=-0.105-0.488 \times \Delta Dam+0.058 \times \Delta P-0.129 \times \Delta Pasture$	0.905	0.003

 ΔDam and $\Delta Pasture$ are changes in percentage area of check-dams and pasture plantation, respectively. ΔP is changes of annual precipitation over the two compared periods.

ID	Catchment	Regression equation	R^2	р
1	Huangfu	y = -0.67x + 45.88	0.85	0.025
2	Gushan	y = -0.90x + 46.66	0.82	0.034
3	Kuye	y = -0.83x + 38.32	0.89	0.017
4	Tuwei	y = -0.48x + 19.94	0.98	0.002
5	Jialu	y = -1.20x + 53.20	0.97	0.002
6	Wuding	y = -0.31x + 16.92	0.97	0.003
7	Qingjian	y = -0.31x + 24.70	0.48	0.193
8	Yanhe	y = -0.26x + 18.54	0.79	0.045
9	Shiwang	y = -0.15x + 3.01	0.87	0.020
10	Qiushui	y = -0.87x + 35.69	0.80	0.040
11	Sanchuan	y = -0.28x + 13.32	0.78	0.046
12	Quchan	y = -0.29x + 21.02	0.52	0.169
13	Xinshui	y = -0.20x + 8.63	0.72	0.069
14	Zhouchuan	y = -0.61x + 17.89	0.61	0.118
15	CSHC	y = -0.54x + 17.74	0.99	0.000

Table 5. Regression equations between the decadal sediment coefficient and percentage of the area affected by soil and water conservation measures ($\overline{SC} = -mA_c + n$) in the catchments.

Daardaa		Р	λ	storm	P_{ss}	P _{storm}		
Decades	r	р	r	р	r	р		
1960s	0.772	0.015*	0.808	0.008**	0.718	0.029*		
1970s	0.266	0.458	0.714	0.020*	0.695	0.026*		
1980s	0.775	0.009**	0.633	0.050*	0.527	0.117		
1990s	0.865	0.001***	0.591	0.072	0.572	0.084		
2000s	0.118	0.715	0.006	0.986	0.138	0.669		

Table 6. Pearson correlation coefficients (r) and two-tailed significance test values (p) between sediment yield and annual precipitation (P), number of storms (N_{storm}) and precipitation amount of storms (P_{storm}) during different decades of the CSHC region.

***, ** and * indicate the significance levels of 0.001, 0.01 and 0.05, respectively.



Figure 1. Location of the studied catchments in the Coarse Sandy Hilly Catchments (CSHC) region within the Loess Plateau.



Figure 2. Spatial distribution of (a) annual mean precipitation (1961-2011), (b) growing season leaf area index (LAI, 1982-2011), (c) soil type and (d) slope in the study area.



Figure 3. Annual precipitation, streamflow and sediment load for the whole CSHC region during 1961-2011.



Figure 4. Land use and cover of the study area in (a) 1975, (b) 1990, (c) 2000 and (d) 2010.



Figure 5. The changes of soil and water conservation measures area and growing season LAI in the study area.



Figure 6. Long-term trends in growing season LAI changes over (a) 1982-2011, (b) 1982-1999 and (c) 2000-2011 in the study area. Inset in each figure shows the frequency distribution of the LAI trends.





Figure 7. The changes of (a) precipitation, (b) streamflow, (c) sediment yield, (d) sediment concentration and (e) sediment coefficient during different stages (1961-1969, 1970-1999 and 2000-2011).



Figure 8. Contributions of precipitation and land use/cover to reductions of sediment load from (a) reference period (P1) to period-2 (P2), (b) reference period (P1) to period-3 (P3) and (c) period-2 (P2) to period-3 (P3).

Figure 9. Spatial distribution of slope *a* in the regression equation $\sqrt{SSY} = aP + b$ during (a) reference period (1961-1969), (b) period-2 (1970-1999) and (c) period-3 (2000-2011). *SSY* is specific sediment yield, and *P* is precipitation.

Figure 10. Spatial distribution of slope *m* in the regression equation $\overline{SC} = -mA_c + n$. \overline{SC} is the decadal average sediment coefficient, and A_c is the percentage of the area affected by soil and water conservation measures in the catchments.

Figure 11. Daily precipitation and sediment load of the Yanhe catchment during rainy season (May-October) in (a) 2003 and (b) 2004.