

1 Extending flood forecasting lead time in large watershed by coupling WRF QPF with
2 distributed hydrological model

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14 **Abstract.** Long lead time flood forecasting is very important for large watershed flood
15 mitigation as it provides more time for flood warning and emergency responses. Latest
16 numerical weather forecast model could provide 1-15 days quantitative precipitation
17 forecasting products at grid format, by coupling this product with distributed
18 hydrological model could produce long lead time watershed flood forecasting products.
19 This paper studied the feasibility of coupling the Liuxihe Model with the WRF QPF for
20 a large watershed flood forecasting in southern China. The QPF of WRF products has
21 three lead time, including 24 hour, 48 hour and 72 hour, with the grid resolution being
22 20kmx20km. The Liuxihe Model is set up with freely downloaded terrain property; the
23 model parameters were previously optimized with rain gauge observed precipitation,
24 and re-optimized with WRF QPF. Results show that the WRF QPF has bias with the
25 rain gauge precipitation, and a post-processing method is proposed to post process the
26 WRF QPF products, which improves the flood forecasting capability. With model
27 parameter re-optimization, the model's performance improves also. This suggests that
28 the model parameters be optimized with QPF, not the rain gauge precipitation. With the
29 increasing of lead time, the accuracy of WRF QPF decreases, so does the flood

30 forecasting capability. Flood forecasting products produced by coupling Liuxihe Model
31 with WRF QPF provides good reference for large watershed flood warning due to its
32 long lead time and rational results.

33

34 **Key words :** WRF, Liuxihe Model, Flood forecasting, lead time, parameter
35 optimization

36

37 **1 Introduction**

38 Watershed flood forecasting is one of the most important non-engineering measures for
39 flood mitigation(Tingsanchali, 2012, Li et al., 2002), significant progresses in
40 watershed flood forecasting have been made in the past decades(Borga et al., 2011,
41 Moreno et al., 2013). Lead time is a key index for watershed flood forecasting,
42 especially for large watershed (Toth et al., 2000, Han et al., 2007). Only flood
43 forecasting products with long lead time are useful as they could provide enough time
44 for flood warning and flood emergency responses. In the long practice of flood
45 forecasting, ground based rain gauge measured precipitation is the main input for flood
46 forecasting model, but as this kind of precipitation is the rainfall falling to the ground
47 already, so it has no lead time. This makes the watershed flood forecasting with very
48 short lead time (Jasper et al., 2002), and could not satisfy the requirement of flood
49 warning (Shim et al., 2002) in lead time, particularly in large watershed, thus reducing
50 the value of the flood forecasting products in watershed flood mitigation.

51

52 The developed numerical weather prediction models in the past decades could provide
53 longer lead time quantitative precipitation forecast(QPF) product at grid format. The

54 lead time for the latest weather prediction model could be as long as to 1~15 days
55 (Buizza, 1999, Ahlgrimm et al., 2016). By coupling the weather prediction model QPF
56 with flood forecasting model, the flood forecasting lead time thus could be extended.
57 This provides a new way for large watershed flood forecasting (Jasper et al., 2002,
58 Zappa et al., 2010, Giard and Bazile, 2000). Many numerical weather prediction models
59 have been proposed and put into operational use, such as the European Centre Medium-
60 Range Weather Forecasts (ECMWF) Ensemble Prediction System (EPS) (Molteni et
61 al., 1996, Barnier et al., 1995), the weather research and forecasting (WRF) model
62 (Skamarock, 2005, 2008, Maussion, 2011), the numerical weather forecast model of
63 Japan Meteorological Agency (Takenaka et al., 2011, Gao and Lian, 2006), the
64 numerical forecast model of China Meteorological Agency (Li and Chen, 2002), and
65 others.

66

67 Watershed flood forecasting relies on a hydrological model for computation tool, while
68 the precipitation is the model's driving force. The earliest hydrological model is
69 regarded as the Sherman unit-graph (Sherman, 1932), which belongs to the category of
70 lumped hydrological model. Many lumped hydrological models have been proposed,
71 such as the Sacramento model (Burnash, 1995), the NAM model (DHI, 2004), the
72 Xinanjiang model (Zhao, 1977), among others. The lumped hydrological model regards
73 the watershed as a whole hydrological unit, thus the model parameter is the same over
74 the watershed, but this is not true, particularly for a large watershed. The precipitation
75 the lumped hydrological model uses is averaged over the watershed also. This further
76 increases the model's uncertainty in large watershed flood forecasting as it is well
77 known that the precipitation distribution over the watershed is highly uneven. The QPF
78 produced by numerical weather prediction model forecasts precipitation at grid format,

79 which provides detailed precipitation distribution information over watershed. This is
80 another advantage of QPF. The lumped hydrological model could not take the
81 advantage of gridded WPF products.

82

83 The latest development of watershed hydrological model is the distributed hydrological
84 model (Refsgaard et al., 1996), which divides the watershed into grids, and different
85 grids could have their own precipitation, terrain property and model parameter. Hence
86 a distributed hydrological model is the ideal model for coupling WRF QPF for
87 watershed flood forecasting. The first proposed distributed hydrological model is SHE
88 model (Abbott et al.1986a, 1986b), and now many distributed hydrological models
89 have been proposed, and a few have been used for watershed flood forecasting, such as
90 the SHE model (Abbott et al.1986a, 1986b), the WATERFLOOD model (Kouwen,
91 1988), the VIC model (Liang et al., 1994), the WetSpa model (Wang et al., 1997), the
92 Vflo model (Vieux et al., 2002), the WEHY model(Kavvas et al., 2004), the Liuxihe
93 model (Chen et al., 2009, 2011), among others.

94

95 As distributed hydrological model calculates the hydrological process at grid scale, so
96 the computation time needed for running the distributed hydrological model is huge even
97 for a small watershed. This limits the model's application in watershed flood forecasting,
98 particularly in large watershed. Model parameter uncertainty related to distributed
99 hydrological model also impacted its application. But with the development of parallel
100 computation algorithm for distributed hydrological model and its deployment on
101 supercomputer (Chen et al., 2013), the computation burden is not a great challenge of
102 distributed hydrological modeling anymore. Also with the development of automatical

103 parameter optimization of distributed hydrological model in flood forecasting (Madsen
104 et al., 2003, Shafii et al., 2009, Xu et al., 2012, Chen et al., 2016), the model parameters
105 could be optimized, and the model's performance could be improved largely. With these
106 advances, now distributed hydrological model could be used for large watershed flood
107 forecasting.

108

109 In this paper, the WRF QPF is coupled with a distributed hydrological model-the
110 Liuxihe model for a large watershed flood forecasting in southern China. The spatial
111 and temporal resolution of WRF QPF is at 20km*20km and 1 hour respectively with
112 three lead times, including 24 hour, 48 hour and 72 hour. The WRF QPF has a similar
113 precipitation pattern with that estimated by rain gauges, but overestimates the averaged
114 watershed precipitation, and the longer the WRF QPF lead time, the higher the
115 precipitation overestimation. Since WRF QPF has systematic bias compared with rain
116 gauge precipitation, a post-processing method is proposed to post process the WRF
117 QPF products, which improves the flood forecasting capability. The Liuxihe Model is
118 set up with freely downloaded terrain property. The model parameters were previously
119 optimized with rain gauge observed precipitation, and re-optimized with WRF QPF.
120 With model parameter re-optimization, the model's performance improved. Model
121 parameters should be optimized with QPF, not the rain gauge precipitation. Flood
122 forecasting products produced by coupling Liuxihe Model with WRF QPF provide
123 good reference for large watershed flood warning due to their long lead time and
124 rational results.

125 **2 Study area and data**

126 2.1 Study area

127 Liujiang River Basin(LRB) is selected as the studied area, which is the largest first
128 order tributary of the Pearl River with a drainage area of 58270 km²(Chen et al., 2017).
129 LRB is in the monsoon area with heavy storms that induced severe flooding in the
130 watershed, and caused huge flood damages in the past centuries. Fig. 1 is a sketch map
131 of LRB.

132

133 Fig. 1 is here

134

135 2.2 Rain gauge precipitation and river flow discharge

136 Precipitation of 68 rain gauges within the watershed in 2011, 2012 and 2013 was
137 collected and used in this study to compare with the WRF QPF. Precipitation data are
138 at one hour interval. River discharge near the watershed outlet is collected also for this
139 same period. As this study focuses on watershed flood forecasting, so only the
140 precipitation and river discharge during the flood events are prepared. There is one
141 flood event in each year. The flood events are numbered as flood event 2011, flood
142 event 2012 and flood event 2013 respectively.

143 **3 WRF QPFs and their post-processing**

144 3.1 WRF model

145 All simulations for this study were conducted with the Advanced Research WRF
146 (WRF-ARW) model version 3.4 (Skamarock et al. 2008). WRF-ARW model is 3-D,

147 non-hydrostatic, fully compressible, and has the terrain-following sigma coordinate
148 system. The model is considered as the next generation's medium range weather
149 forecasting model, and can simulate different weather processes from cloud scale to
150 synoptic scale, especially in horizontal resolution of 1 ~ 10 km. The model also
151 integrates the advanced numerical methods and data assimilation techniques, a variety
152 of physically process schemes, and multiple nested methods and the capability of being
153 used in different geographical locations. WRF-ARW model satisfies the needs of
154 scientific research and practical applications for this study.

155

156 Prior studies have been shown in quantitative precipitation forecasting by using WRF-
157 ARW model. For instance, Pennelly et al. (2014) employed the WRF model to predict
158 three precipitation events of Alberta, Canada, and compared the precipitation with 48
159 hour leading time predicted by the model with rain gauges. The results showed that
160 Kain-Fritsch cumulus parameterization overestimated the value of precipitation
161 invariably. Eiserloh and Chiao (2015) used WRF-ARW with data assimilation to
162 investigate an Atmospheric River event over Northern California. Maussion et al. (2011)
163 compared the capability of WRF model in retrieving monthly precipitation and
164 snowfall at three different spatial resolution including 30, 10 and 2 km domains over
165 Tibet. Their results showed that the model was able to recapture monthly precipitation
166 and snowfall. Pan et al. (2012) used two WRF simulation groups between pre-process
167 and post-process in Heihe river basin, and compared and analyzed the mean bias error,
168 root mean square error and correlation coefficient of the two WRF groups. Huang et al.
169 (2011) found that variations in the microphysical process parameterization schemes
170 had much more influence on precipitation than that of cumulus parameterization
171 schemes, especially for a torrential rain attributed to large-scale forcing that mainly

172 resulted from stratus clouds. Kumar et al. (2008) used WRF model to study a heavy
173 rain in 2005, their results showed that WRF model could reproduce the storm event and
174 its dynamical and thermo-dynamical characteristics. Hong and Lee (2009) conducted a
175 triply nested WRF simulation for convective initiation of a thunderstorm. Givati et al.
176 (2012) predicted the hiemal precipitation event of 2008 and 2009 based on WRF model
177 in upstream of the Jordan River, and coupled WRF model with hydrological model-
178 HYMKE to simulate the velocity and discharge of Jordan River. Sensitivity experiment
179 of WRF microphysical schemes by Niu et al. (2007) have shown the adequate
180 performance of precipitation predicted associated with region, center location and
181 rainfall intensity. Xu et al.(2007) compared the hiemal continuous precipitation process
182 predicted with the estival results by WRF model, the results showed that the KF scheme
183 was better than BM scheme in summer. Hu et al. (2008) found that the parameterization
184 scheme of WRF model was related to the model resolution, and the parameterization
185 scheme should be selected by the resolution of WRF model.

186 3.2 Configuration of WRF for LRB B

187 The WRF-ARW was applied to LRB following the configurations by Li et al. (2015).
188 More information about LBR can be found in Li et al. (2015) and Chen et al. (2017).
189 The model domain is centered at 23.8N, 109.2W with the Lambert conformal projection.
190 The vertical structure includes 28 levels with the focus on the lower-levels of
191 troposphere. The initial and time-dependent lateral boundary conditions are supplied
192 from NCEP Global Forecast System (GFS) 3-hourly global analysis at 0.5 °horizontal
193 resolution. The model domain has a 20 km grid resolution. The single-moment 3-class
194 microphysics (WSM3) parameterization (Hong and Lim, 2006) is adopted for this study.
195 Kain-Fritsch cumulus parameterization (Kain, 2004) as well as the YSU boundary layer
196 microphysics scheme (Hong et al., 2006) are used. Other physics schemes used

197 include the NOAH scheme for the land surface physics (Ek et al., 2003), the Goddard
198 scheme for the shortwave radiation physics (based on Chou and Suarez,1994), and
199 Rapid Radiative Transfer Model (RRTM) scheme for the longwave radiation physics
200 (Mlawer et al., 1997).

201

202 The spatial and temporal resolution of WRF is at 20km x 20km and 1 hour, respectively.
203 The entire Liujiang River Basin is covered by total 156 grid points of the WRF model.
204 The simulated QPF for flood events in years 2011 to 2013 were produced with three
205 different lead time (i.e., 24 hours, 48 hours and 72 hours), respectively. As shown in
206 Figs. 2-4, the WRF QPF products in in three different years, while (a) is the rain gauge
207 precipitation, (b) is the WRF QPF with 24 hour lead time, (c) is the WRF QPF with 48
208 hour lead time, and (d) is the WRF QPF with 72 hour lead time.

209 Fig. 2 is here

210 Fig. 3 is here

211 Fig. 4 is here

212 3.3 Evaluation of WRF QPF and rain gauges precipitation

213 Comparisons of WRF QPF and rain gauge precipitation are performed. From the
214 simulated results, as shown in Figs. 2, 3 and 4, it appears that the temporal
215 precipitation pattern of both products is similar, although there are some insignificant
216 differences. To make further comparison, the accumulated precipitation of the three
217 flood events averaged over the watershed are calculated and listed in Table 1.

218

219 Table 1 is here

220

221 As summarized in Table 1, it could be found that the WRF QPF accumulated
222 precipitation has obvious bias with rain gauge accumulated precipitation. For all the
223 three flood events, the WRF QPF accumulated precipitation are higher than those
224 measured by rain gauges. In other words, the WRF QPF overestimates the
225 precipitation. For flood event 2011, the overestimated watershed averaged
226 precipitation of WRF QPF with lead time of 24 hour, 48 hour and 72 hour are 23%,
227 32% and 55% respectively. For the flood event in 2012, they are 16%, 37% and 71%
228 respectively. They are 50%, 73% and 95% respectively from the event in 2013. The
229 results suggest that longer the WRF QPF lead time, the higher chance of
230 overestimation.

231 3.4 WRF QPF statistical calibrations

232 From the simulated results (c.f., Fig. 2, 3 and 4, and Table 1), the WRF QPF has
233 significant bias compared to rain gauge precipitation. Assuming the rain gauge
234 precipitation is correct, the WRF QPF needs to be further calibrated. In order to do so,
235 the WRF QPF is further post-processed based on the rain gauge precipitation to
236 correct the systematic error of WRF QPF. The principle of WRF QPF statistical
237 calibrations proposed in this study is to keep the areal averaged event accumulated
238 precipitation from both model and rain gauge products to be equivalent. In other
239 words, the statistical approach is to nudge the WRF QPF precipitation to rain gauge
240 results.

241

242 Based on this principle, the WRF QPF post-processing procedure is summarized as
243 follows:

244

245 1) Calculate the areal average precipitation of the WRF QPF for each flood events
246 over the watershed as following equation.

$$247 \quad \bar{P}_{WRF} = \frac{\sum_{i=1}^N P_i F_i}{N} \quad (1)$$

248 Where, \bar{P}_{WRF} is the areal average precipitation of WRF QPF of one flood event, P_i is
249 the precipitation on WRF grid i, F_i is the surface area of WRF grid i divided by the
250 whole watershed drainage area, N is the total number of WRF grids.

251

252 2) Calculate the areal average precipitation of the rain gauges with the following
253 equation.

$$254 \quad \bar{P}_2 = \frac{\sum_{j=1}^M P_j}{M} \quad (2)$$

255 Where, \bar{P}_2 is the areal average precipitation of the rain gauges network, P_j is the
256 precipitation observed by jth rain gauge, M is the total number of rain gauges.

257

258 3) The precipitation of every WRF QPF grids then could be revised with the
259 following equation.

260
$$P'_i = P_i \frac{\bar{P}_2}{\bar{P}_{WRF}} \quad (3)$$

261 Where, P'_i is the revised precipitation of ith WRF grid.

262 With the above WRF QPF statistical calibration methods, the WRF QPF of flood
263 event 2011, 2012 and 2013 are post-processed, and will be used to couple with the
264 Liuxihe Model for flood simulations.

265 **4 Hydrological model**

266 4.1 Liuxihe Model

267 Liuxihe model is a physically based fully distributed hydrological model proposed
268 mainly for watershed flood forecasting (Chen, 2009, Chen et al., 2011), and has been
269 used in a few watersheds for flood forecasting(Chen, 2009, Chen et al., 2011, 2013,
270 2016, Liao et al., 2012 a, b, Xu et al., 2012 a, b). In Liuxihe Model, runoff components
271 are calculated at grid scale, runoff routes at both grid and watershed scale. Runoff
272 routing is divided into hill slope routing and river channel routing by using different
273 computation algorithm. Liuxihe Model proposed an automatic parameter optimization
274 method using PSO algorithm (Chen et al., 2016), which largely improves the model's
275 performance in watershed flood forecasting. Now Liuxihe Model is deployed on a
276 supercomputer system with parallel computation techniques (Chen et al., 2013) that
277 largely facilitates the model parameter optimization of Liuxihe Model.

278

279 Chen et al. (2017) set up Liuxihe Model in LRB with freely downloaded terrain
280 property data from the website at a spatial resolution of 200m*200m, and optimized
281 model parameters with observed hydrological data. The model was validated by
282 observed flood events data, and the model performance was found rational and could

283 be used for real-time flood forecasting. The model only uses rain gauge precipitation,
284 so its flood forecasting lead time is limited. In this study, the Liuxihe Model was set up
285 in LRB and the optimized model parameters were be used in this study as the first
286 attempt. Fig. 5 shows the model structure.

287

288

Fig.5 is here

289

290 4.2 Liuxihe Model parameter optimization

291 While the model parameters optimization by Chen et al. (2017) is done by using the
292 rain gauge precipitation, this study uses the WRF QPF as the precipitation input. So the
293 parameters of Liuxihe Model that were set up in LRB may not be appropriate for
294 coupling the WRF QPF. For this reason, considering Liuxihe Model is a physically
295 based distributed hydrological model, the parameters were optimized again by using
296 the WRF QPF flood event in 2011. Hence, the WRF QPF is the post-processed one, not
297 the original one. Results of parameter optimization are shown in Fig. 6. Among them,
298 (a) is the objective function evolution result, (b) is the parameters evolution result, and
299 (c) is the simulated flood process by using the optimized model parameters. To compare,
300 the simulated flood process of flood event 2011 was also drawn in Fig. 6(c).

301

302

Fig. 6 is here

303

304 From the result of Fig. 6(c), it may be seen that the optimized model parameters with
305 WRF QPF improved the flood simulation when compared to the corresponding flood
306 simulation based on gauge precipitation. This means parameter optimization with

307 WRF QPF is necessary.

308 4.3 Coupling WRF QPF with Liuxihe Model for LRB flood forecasting

309 When the Liuxihe Model set up for LRB flood forecasting (Chen et al., 2017) was
310 employed to couple with the WRF QPF, the model spatial resolution remained to be
311 200m*200m. As the spatial resolution of WRF QPF is at 20km*20km, the WRF QPF
312 was downscaled to the resolution of 200m*200m by using the nearest downscaling
313 method, the same spatial resolution of the flood forecasting model.

314 **5 Results and discussions**

315 5.1 Effects of WRF post-processing

316 The original WRF QPF and the post-processed QPF were used to couple with the
317 Liuxihe Model. In this simulation, the original model parameters that were optimized
318 with the rain gauge precipitation were employed, not the re-optimized model
319 parameters. The simulated results are shown in Fig. 7, 8 and 9.

320

321 Fig. 7 is here

322 Fig. 8 is here

323 Fig. 9 is here

324

325 From the above results, it could be seen that the simulated flood discharges with the
326 original WRF QPF are much lower than the observed ones. But with post-processed
327 WRF QPF used, the simulated flood discharge increased and became much more close
328 to the observation. This implies that the flood forecasting capability has been improved

329 by post-processing of WRF QPF. To further compare the three results, 5 evaluation
330 indices, including Nash-Sutcliffe coefficient(C), correlation coefficient(R), process
331 relative error(P), peak flow relative error(E) and water balance coefficient(W) were
332 calculated and listed in Table 2.

333

334

Table 2 is here

335

336 From the results of Table 2, it has been found that all the 5 evaluation indices have been
337 improved by coupling the post-processed WRF QPF. For example, for flood event 2011
338 with 24 hour lead time, the Nash-Sutcliffe coefficient/C, correlation coefficient/R,
339 process relative error/P, peak flow relative error/E and coefficient of water balance/W
340 with original WRF QPF are 0.65, 0.88, 35%, 14% and 1.44 respectively, but those with
341 the post-processed WRF QPF are 0.75, 0.93, 23%, 8% and 1.15 respectively. For flood
342 event 2012 with 48 hour lead time, the above 5 evaluation indices with original WRF
343 QPF are 0.63, 0.75, 48%, 12% and 1.43 respectively, and are 0.75, 0.84, 26%, 8% and
344 1.32 respectively with the post-processed WRF QPF. For flood event 2013 with 72 hour
345 lead time, the above 5 evaluation indices with original WRF QPF are 0.44, 0.75, 129%,
346 45% and 1.66 respectively, and are 0.55, 0.82, 98%, 23%, 1.25 respectively with the
347 post-processed WRF QPF. It is obvious that with the post-processed WRF QPF, the
348 evaluation indices are improved substantially. These results show that WRF QPF post
349 processing could improve the flood forecasting capability because the WRF QPF is
350 more close to the observed precipitation after post-processing. So it should be practiced
351 for real-time flood forecasting.

352

353 5.2 Results comparison for different model parameters

354 The model parameters optimized with rain gauge precipitation and WRF QPF are
355 different; so different parameter values will result in different model performance. To
356 analyze this effect, the flood events of 2012 and 2013 with two different sets of model
357 parameters values are simulated, and are shown in Fig. 10 and Fig. 11 respectively.
358 Only the post-processed WRF QPF are coupled in this simulation.

359

360 Fig. 10 is here

361 Fig. 11 is here

362

363 From the above figures it may be that the simulated flood results with re-optimized
364 model parameters are better than those simulated with the original model parameters.
365 The simulated flood discharge with the re-optimized model parameters matches. To
366 further compare the two results, 5 evaluation indices, including Nash-Sutcliffe
367 coefficient(C), correlation coefficient(R), process relative error(P), peak flow relative
368 error(E) and water balance coefficient(W) are calculated and listed in Table 3.

369

370 Table 3 is here

371

372 From the results of Table 3, it is found that the results of flood simulation based on the
373 re-optimized model parameters have better evaluation indices. All evaluation indices
374 for those based on re-optimized model parameters are improved. For example, for flood
375 event 2012 with 24 hour lead time, the Nash-Sutcliffe coefficient/C, correlation
376 coefficient/R, process relative error/P, peak flow relative error/E and coefficient of

377 water balance/W with original model parameters are 0.58, 0.82, 35%, 12% and 1.08
378 respectively, but those with the re-optimized model parameters are 0.74, 0.86, 28%, 8%
379 and 0.95 respectively. For flood event 2013 with 48 hour lead time, the 5 indices with
380 the original model parameters are 0.62, 0.86, 22%, 13% and 1.24 respectively, and are
381 0.68, 0.89, 18%, 9% and 1.06 respectively for those with re-optimized model
382 parameters. So it could be said that in coupling the WRF QPF with distributed
383 hydrological model, the model parameters need to be re-optimized with the WRF QPF.
384 This finding implies that the precipitation pattern has obvious impact on model
385 parameters. It should be considered, and model parameter optimization is a rational way
386 for considering this effect.

387 5.3 Flood simulation accuracy with different lead time

388 To compare the model performance with different lead time, the flood events with 3
389 different lead times are simulated and shown in Fig. 12. The model parameters are the
390 re-optimized ones, and the QPF is the post-processed QPF.

391

392 Fig. 12 is here

393

394 From the results of Fig. 12, it could be seen that the flood simulation result get worse
395 as the lead time increases, i.e., the model performance with 24 hour lead time is better
396 than that with 48 hour lead time, and the model performance with 48 hour lead time is
397 better than that with 72 hour lead time. The simulated hydrological process with 24
398 hour lead time is very similar to that simulated with rain gauge precipitation. To further
399 compare the results, 5 evaluation indices, including Nash-Sutcliffe coefficient(C),
400 correlation coefficient(R), process relative error(P), peak flow relative error(E) and
401 water balance coefficient(W) were calculated and listed in Table 4.

402

403

Table 4 is here

404

405 From the results of Table 4, it is found that the simulated flood events with 24 hour lead
406 time have best evaluation indices, and are very close to those simulated with rain gauge
407 precipitation. For flood event 2012, the Nash-Sutcliffe coefficient/C, correlation
408 coefficient/R, process relative error/P, peak flow relative error/E and coefficient of
409 water balance/W with rain gauge are 0.82, 0.89, 20%, 5% and 0.8 respectively, while
410 those with 24 hour lead time are 0.74, 0.86, 28%, 8% and 0.95 respectively, those with
411 48 hour lead time are 0.63, 0.84, 48%, 12% and 1.32 respectively, and are 0.56, 0.56,
412 56%, 18% and 1.54 respectively for 72 hour lead time. For flood event 2013, the Nash-
413 Sutcliffe coefficient/C, correlation coefficient/R, process relative error/P, peak flow
414 relative error/E and coefficient of water balance/W with rain gauge are 0.95, 0.92, 8%,
415 6% and 1.08 respectively, while those with 24 hour lead time are 0.87, 0.87, 9%, 12%
416 and 1.02 respectively, those with 48 hour lead time are 0.62, 0.86, 22%, 13% and 1.24
417 respectively, and are 0.61, 0.87, 75%, 17% and 1.66 respectively for 72 hour lead time.
418 This finding means that the current WRF QPF capability is lead-time dependent, and
419 with the increasing lead time, the practical value of WRF QPF gets lower.

420 **6 Conclusion**

421 In this study, the WRF QPF was coupled with a distributed hydrological model-the
422 Liuxihe model, for large watershed flood forecasting, and three lead times of WRF QPF
423 products, including 24 hours, 48 hours and 72 hours are tested. WRF QPF post
424 processing method is proposed and tested, model parameters are re-optimized by using
425 the post-processed WRF QPF, model performances are compared among various

426 conditions. Based on the results of this study, the following conclusions could be drawn:

427

428 1) The quantitative precipitation forecasting produced by WRF model has a similar
429 pattern with that estimated by rain gauges temporally, but overestimated the averaged
430 watershed precipitation for the event accumulated total precipitation. The longer the
431 WRF QPF lead time, the higher the precipitation overestimation. For flood event 2011,
432 the overestimated watershed averaged precipitation of WRF QPF with lead times of 24
433 hour, 48 hour and 72 hour are 23%, 32% and 55% respectively. For flood event 2012,
434 these are 16%, 37% and 71% respectively, while for flood event 2013, these are 50%,
435 73% and 95% respectively.

436

437 2. WRF QPF has systematic bias compared with rain gauge precipitation, and this
438 bias could be reduced via post-processing. Principle used in this study for WRF QPF
439 post processing is effective and could improve the flood forecasting capability. For
440 flood event 2011 with 24 hour lead time, the Nash-Sutcliffe coefficient/C, correlation
441 coefficient/R, process relative error/P, peak flow relative error/E and coefficient of
442 water balance/W with original WRF QPF are 0.65, 0.88, 35%, 14% and 1.44
443 respectively, but those with the post-processed WRF QPF are 0.75, 0.93, 23%, 8%
444 and 1.15 respectively. For flood event 2012 with 48 hour lead time, the above 5
445 evaluation indices with original WRF QPF are 0.63, 0.75, 48%, 12% and 1.43
446 respectively, and are 0.75, 0.84, 26%, 8% and 1.32 respectively with the post-
447 processed WRF QPF. For flood event 2013 with 72 hour lead time, the above 5
448 evaluation indices with original WRF QPF are 0.44, 0.75, 129%, 45% and 1.66
449 respectively, and are 0.55, 0.82, 98%, 23%, 1.25 respectively with the post-processed
450 WRF QPF.

451

452 3. Hydrological model parameters optimized with the rain gauge precipitation need to
453 be re-optimized using the post-processed WRF QPF, this improves the model
454 performance significantly. That is, in coupling the distributed hydrological model with
455 QPF for flood forecasting, the model parameters should be optimized with the QPF
456 produced by WRF. For flood event 2012 with 24 hour lead time, the Nash-Sutcliffe
457 coefficient/C, correlation coefficient/R, process relative error/P, peak flow relative
458 error/E and coefficient of water balance/W with original model parameters are 0.58,
459 0.82, 35%, 12% and 1.08 respectively, but those with the re-optimized model
460 parameters are 0.74, 0.86, 28%, 8% and 0.95 respectively. For flood event 2013 with
461 48 hour lead time, the 5 indices with the original model parameters are 0.62, 0.86, 22%,
462 13% and 1.24 respectively, and are 0.68, 0.89, 18%, 9% and 1.06 respectively for those
463 with re-optimized model parameters.

464

465 4. The simulated floods by coupling WRF QPF with distributed hydrological model are
466 rational and could benefit the flood management communities due to their longer lead
467 times for flood warning. They provide a good reference for large watershed flood
468 warning. But with the lead time getting longer, the flood forecasting accuracy is getting
469 lower. For flood event 2012, the Nash-Sutcliffe coefficient/C, correlation coefficient/R,
470 process relative error/P, peak flow relative error/E and coefficient of water balance/W
471 with rain gauge are 0.82, 0.89, 20%, 5% and 0.8 respectively, while those with 24 hour
472 lead time are 0.74, 0.86, 28%, 8% and 0.95 respectively, those with 48 hour lead time
473 are 0.63, 0.84, 48%, 12% and 1.32 respectively, and are 0.56, 0.56, 56%, 18% and 1.54
474 respectively for 72 hour lead time. For flood event 2013, the Nash-Sutcliffe
475 coefficient/C, correlation coefficient/R, process relative error/P, peak flow relative

476 error/E and coefficient of water balance/W with rain gauge are 0.95, 0.92, 8%, 6% and
477 1.08 respectively, while those with 24 hour lead time are 0.87, 0.87, 9%, 12% and 1.02
478 respectively, those with 48 hour lead time are 0.62, 0.86, 22%, 13% and 1.24
479 respectively, and are 0.61, 0.87, 75%, 17% and 1.66 respectively for 72 hour lead time.

480 **7 Data availability**

481 The Rain gauge precipitation and river flow discharge data were provided by the
482 Bureau of Hydrology, Pearl River Water Resources Commission, China exclusively
483 used for this study. The WRF QPF results were provided by Yuan Li, and has been
484 published and cited in this paper (Li et al. 2015). The Liuxihe Model used in this
485 study are provided by Yangbo Chen, and has been published and cited in this paper
486 (Chen et al. 2017).

487

488 **Competing interests.** The authors declare that they have no conflict of interest.

489

490 **Acknowledgements**

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492 (funding no. 201301070), the National Science Foundation of China (funding no.
493 50479033), and the Basic Research Grant for Universities of the Ministry of Education
494 of China (fundingno.13lgjc01).

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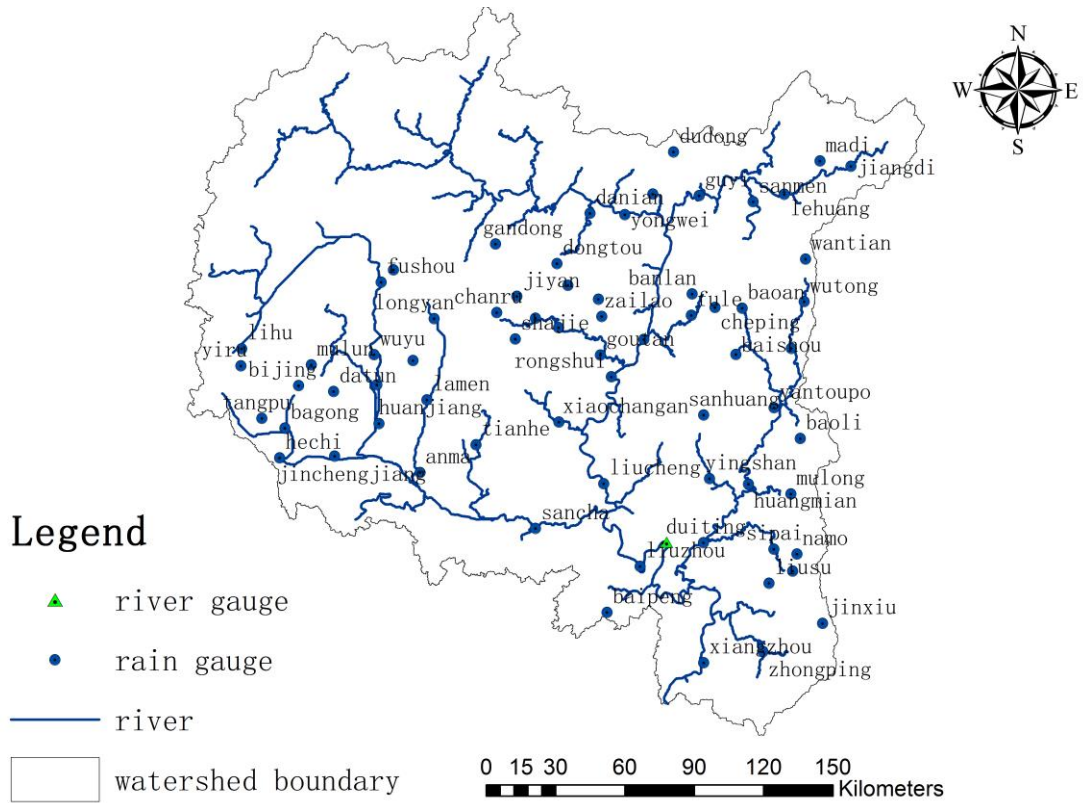
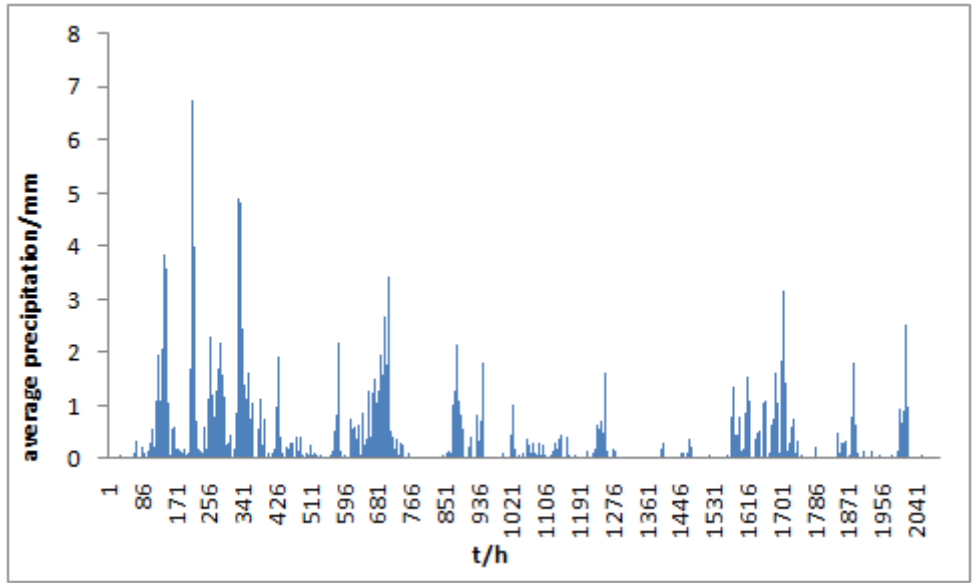


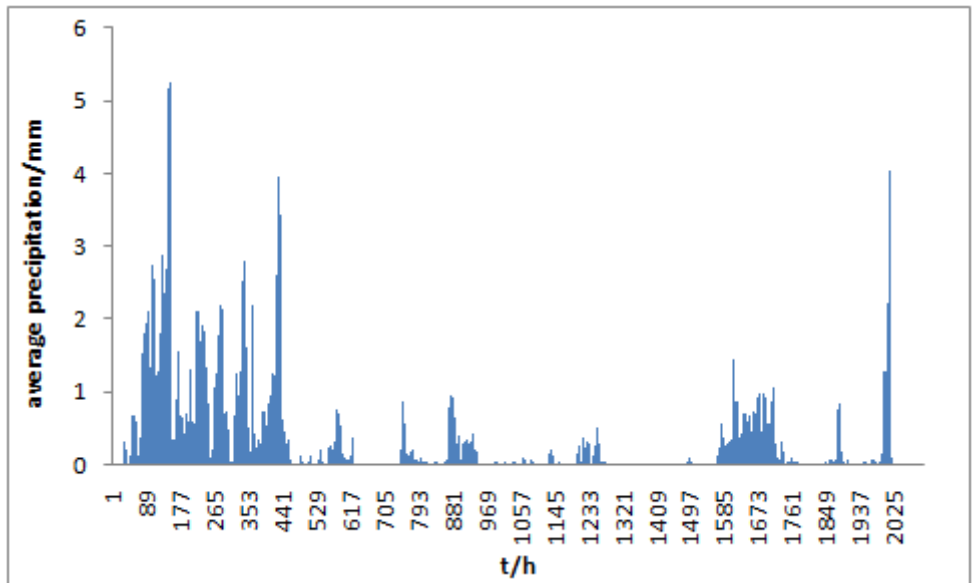
Fig. 1 Sketch map of Liujiang River Basin(Chen et al., 2017)



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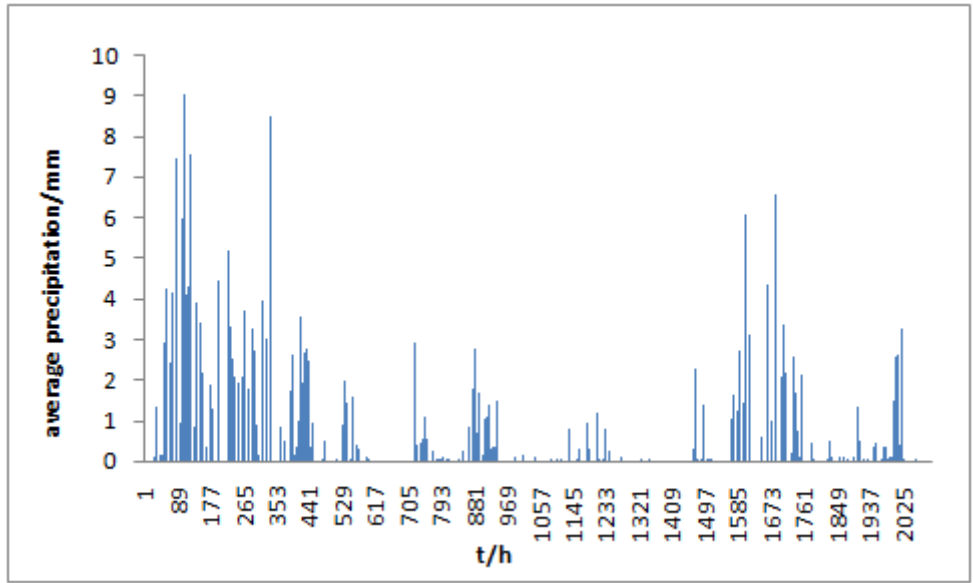
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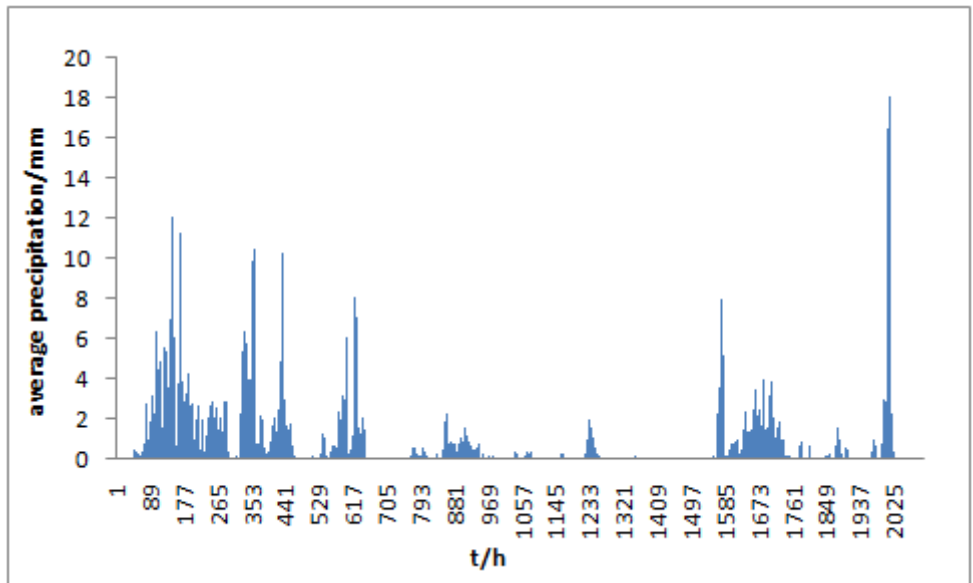
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(b)



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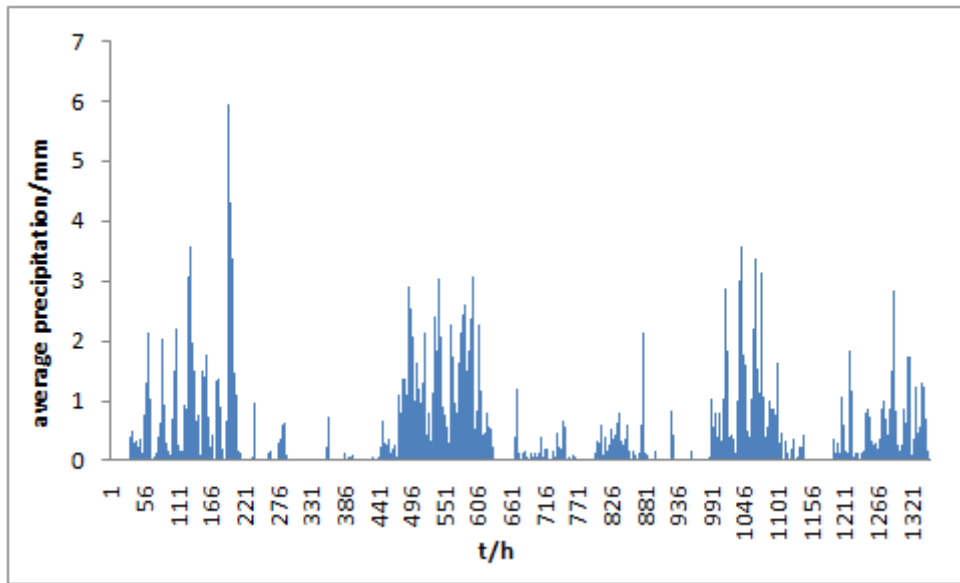


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(d)

510 Fig. 2 Precipitation pattern comparison of two precipitation products(2011), (a) is the
511 average precipitation of rain gauges, (b) is the average precipitation of WRF with 24
512 hour lead time, (c) is the average precipitation of WRF with 48 hour lead time, (d) is
513 the average precipitation of WRF with 72 hour lead time.

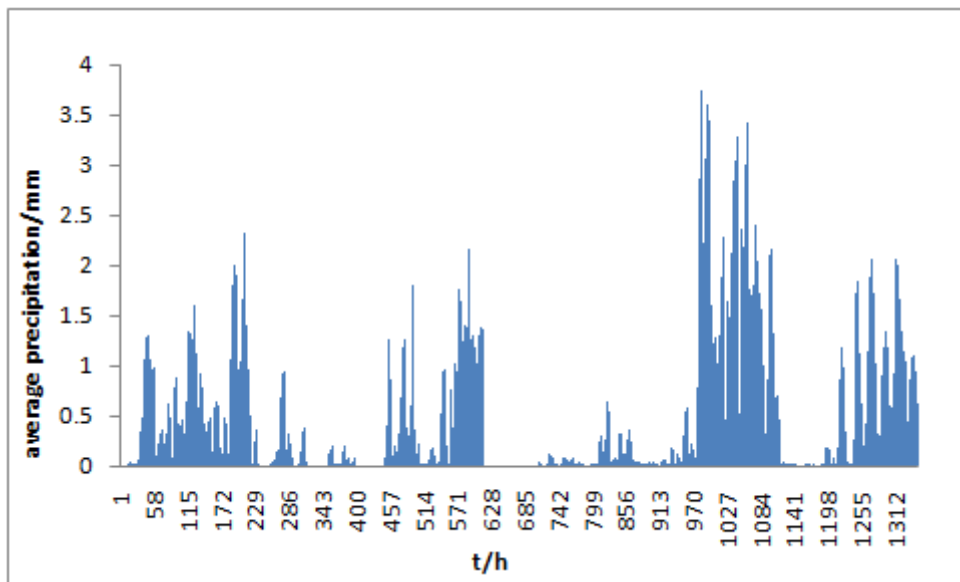
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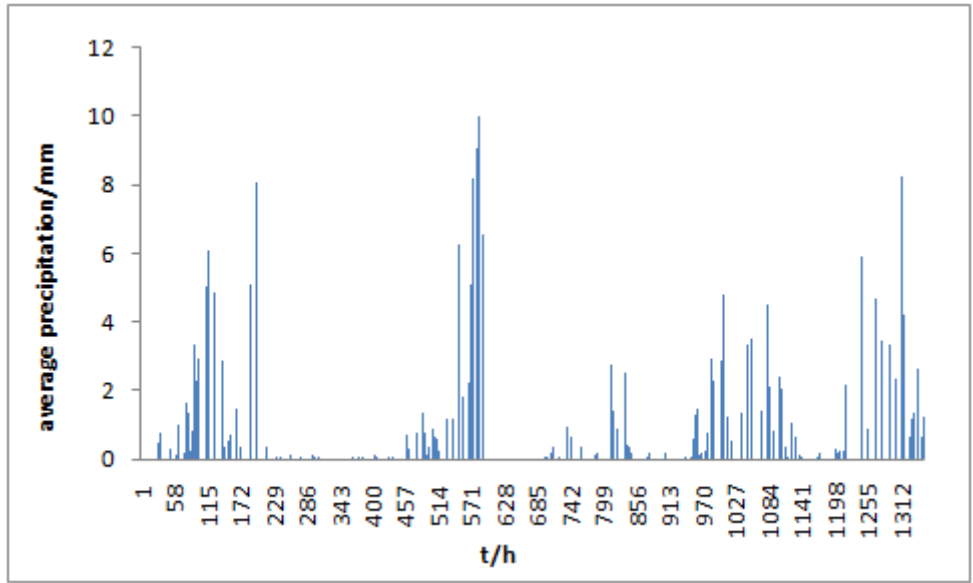
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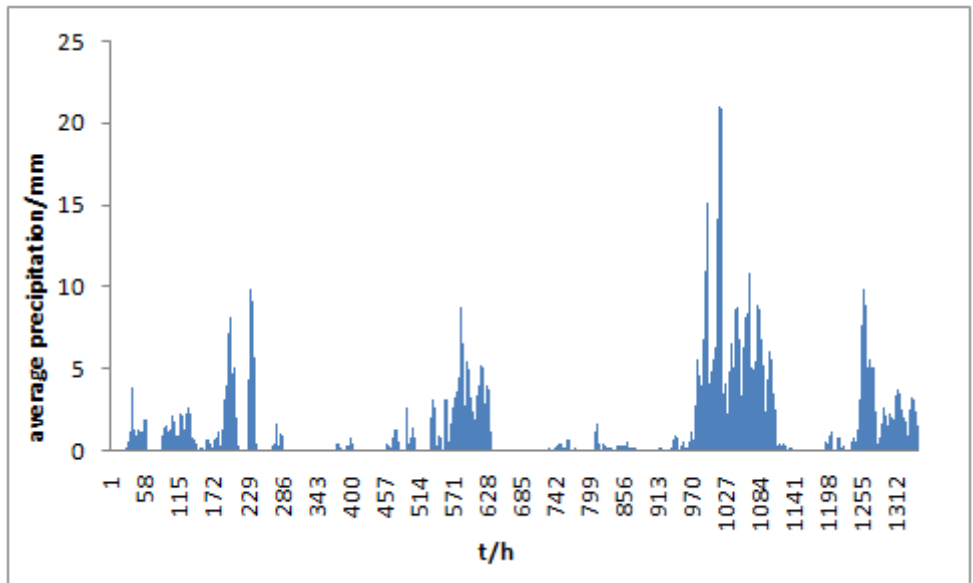
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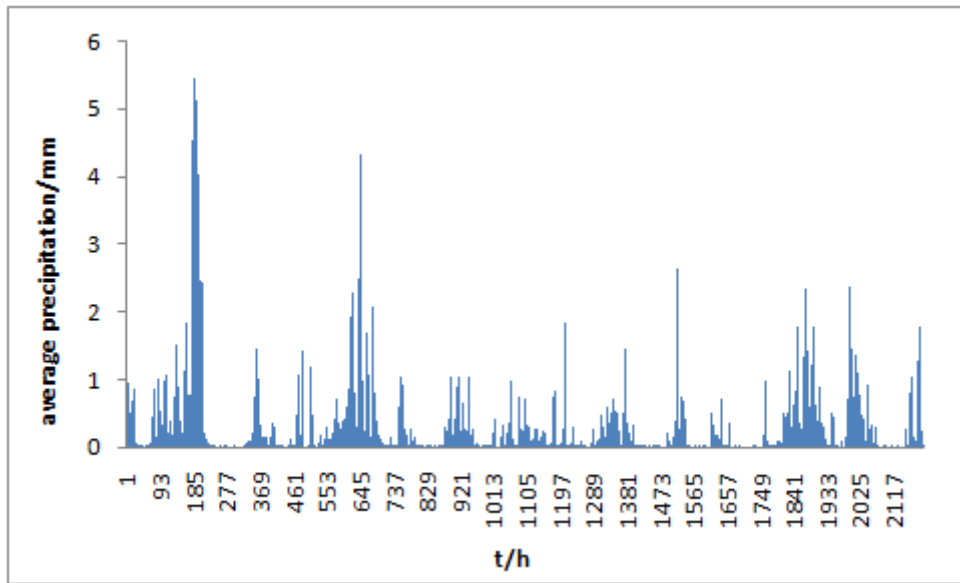


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523 Fig. 3 Precipitation pattern comparison of two precipitation products(2012) , (a) is the
524 average precipitation of rain gauges, (b) is the average precipitation of WRF with 24
525 hour lead time, (c) is the average precipitation of WRF with 48 hour lead time, (d) is
526 the average precipitation of WRF with 72 hour lead time.

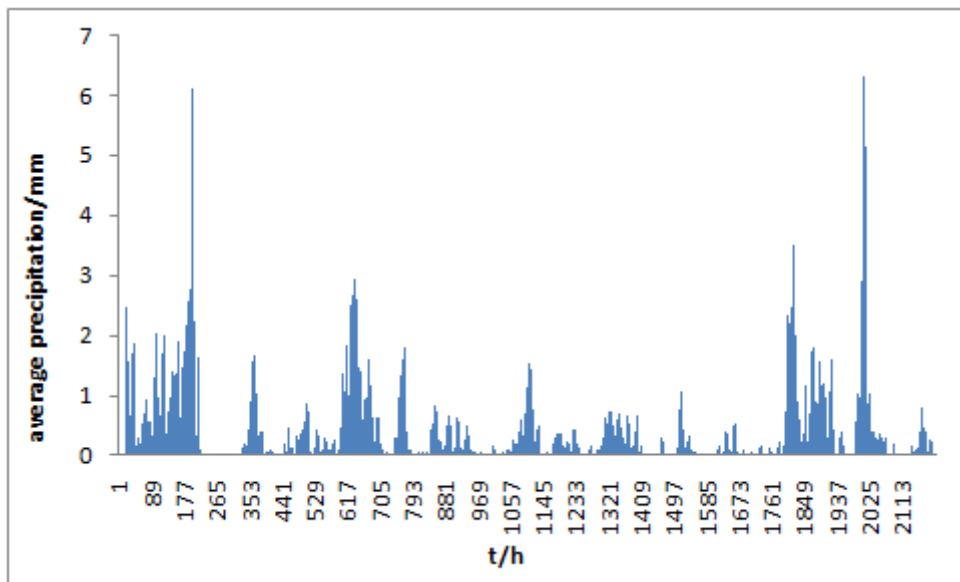
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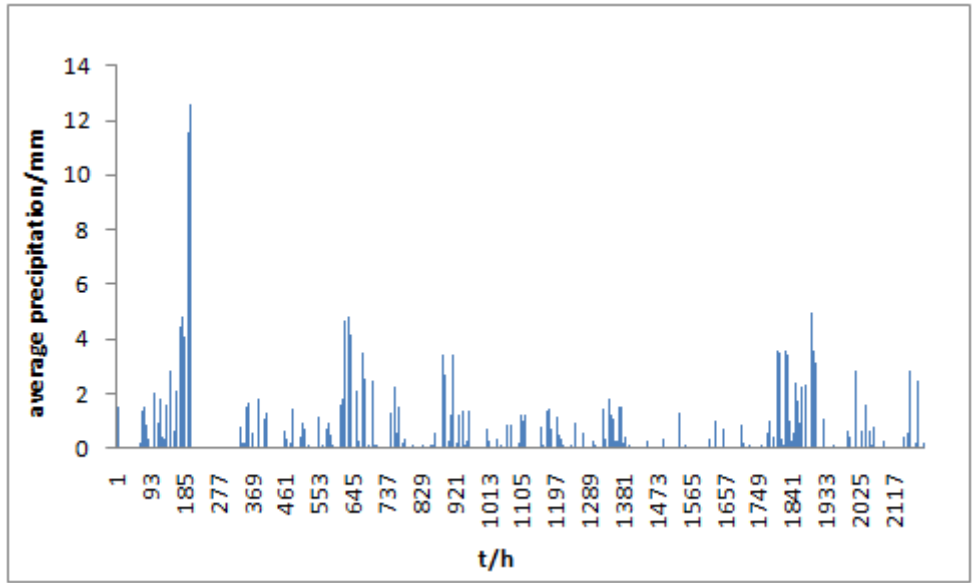
(a)



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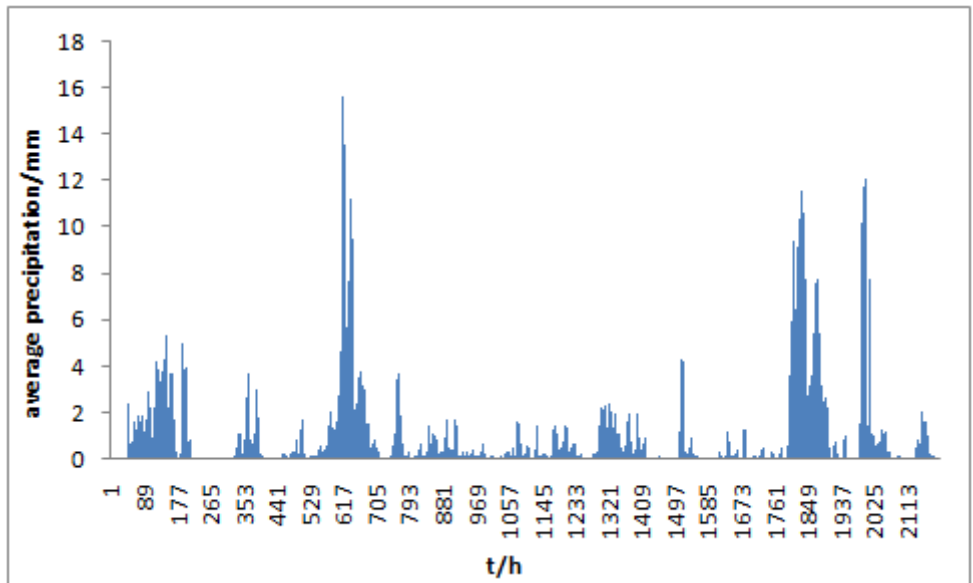
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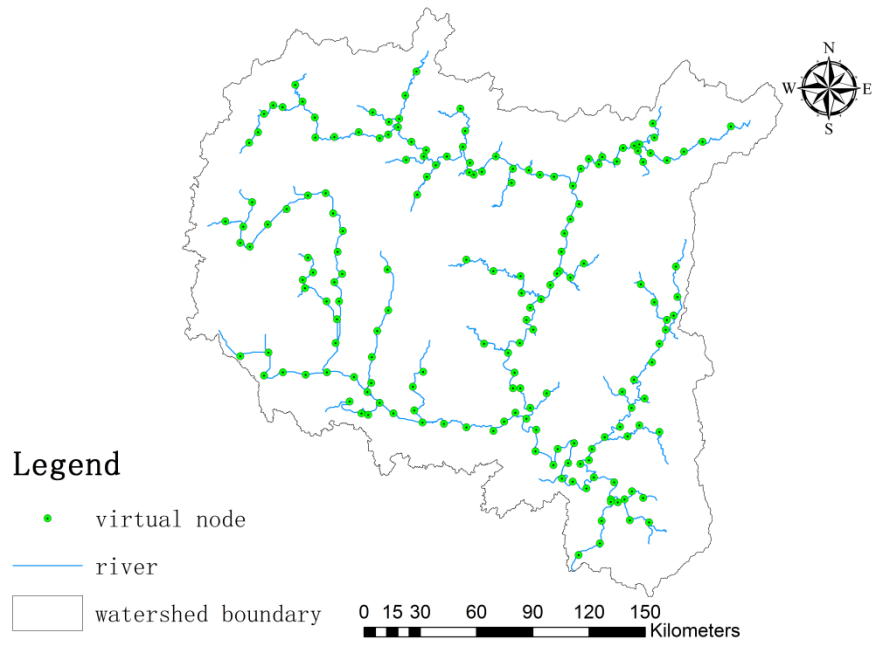


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(d)

536 Fig. 4 Precipitation pattern comparison of two precipitation products(2013), (a) is the
 537 average precipitation of rain gauges, (b) is the average precipitation of WRF with 24
 538 hour lead time, (c) is the average precipitation of WRF with 48 hour lead time, (d) is
 539 the average precipitation of WRF with 72 hour lead time.



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541 Fig.5 Liuxihe Model structure of LRB (200m×200m resolution, Chen et. al., 2017)

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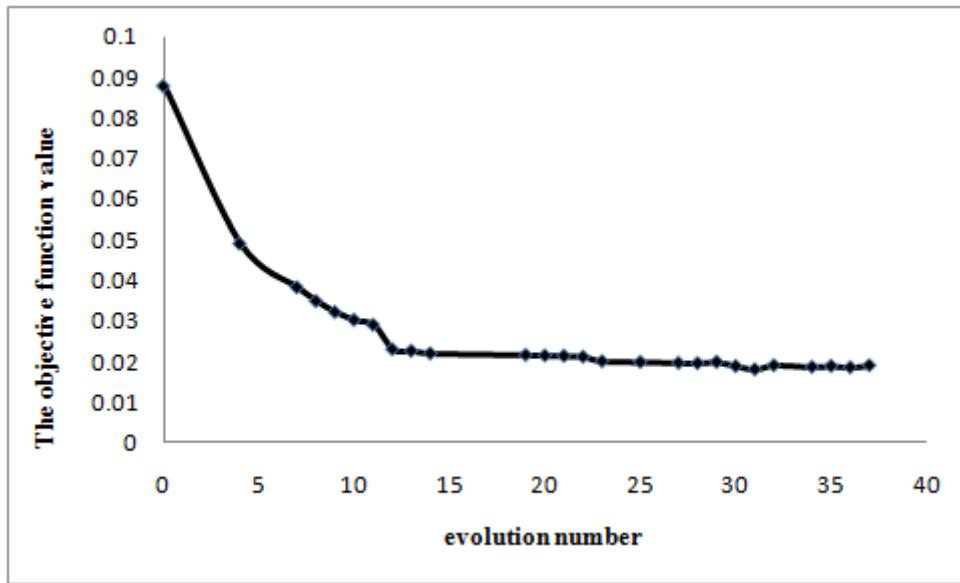
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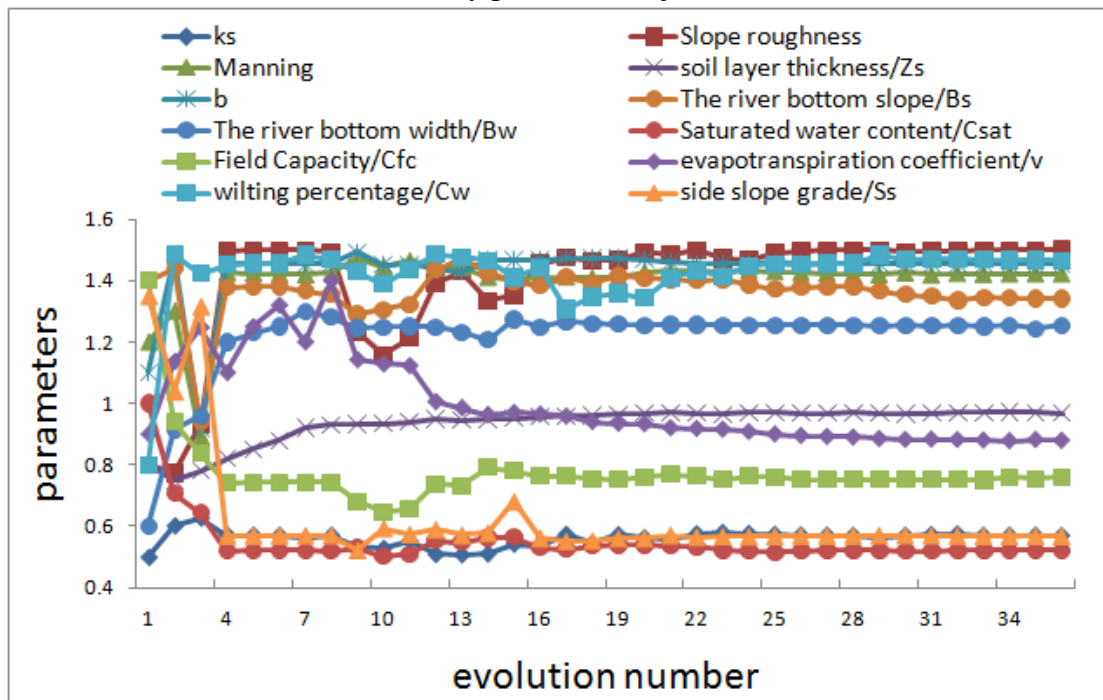
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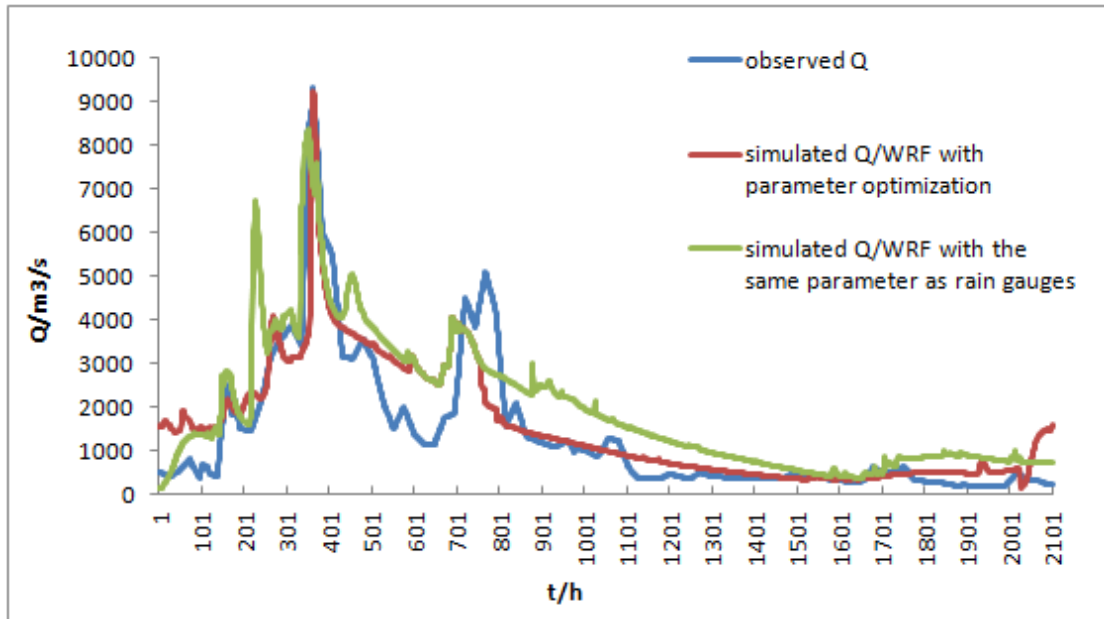
(a) Evolutionary process of objective function



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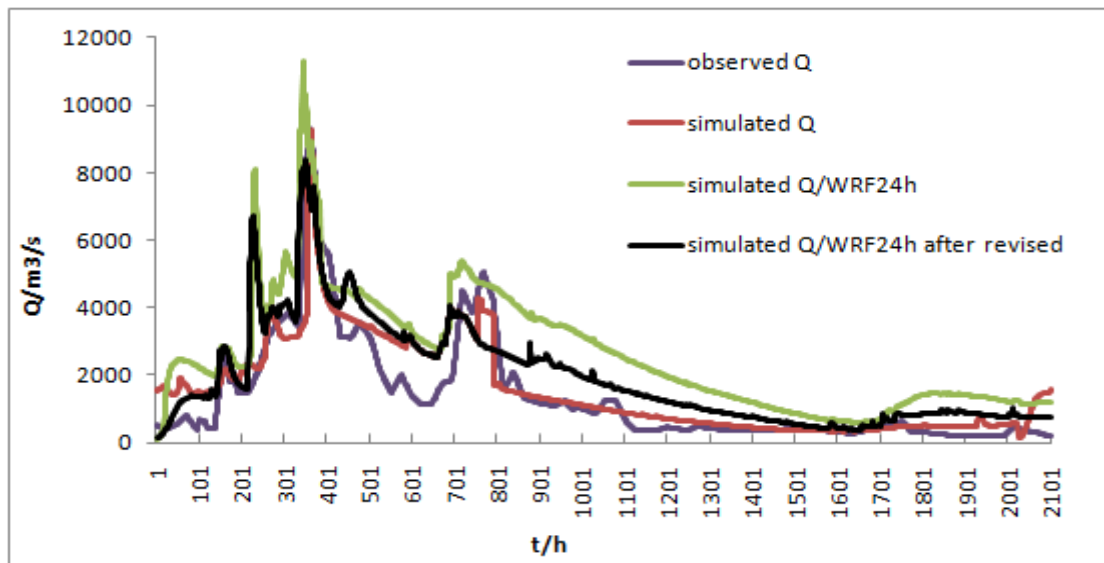
(b) Parameter evolution process



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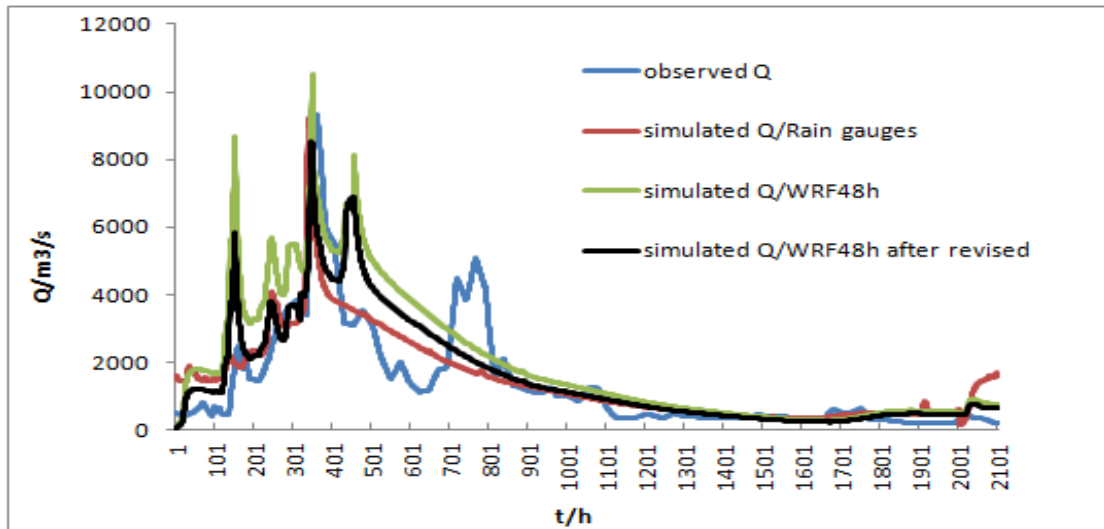
(c) Simulated flood process with optimized model parameters

Fig. 6 Parameter optimization results of Liuxihe Model for LRB with WRF QPF



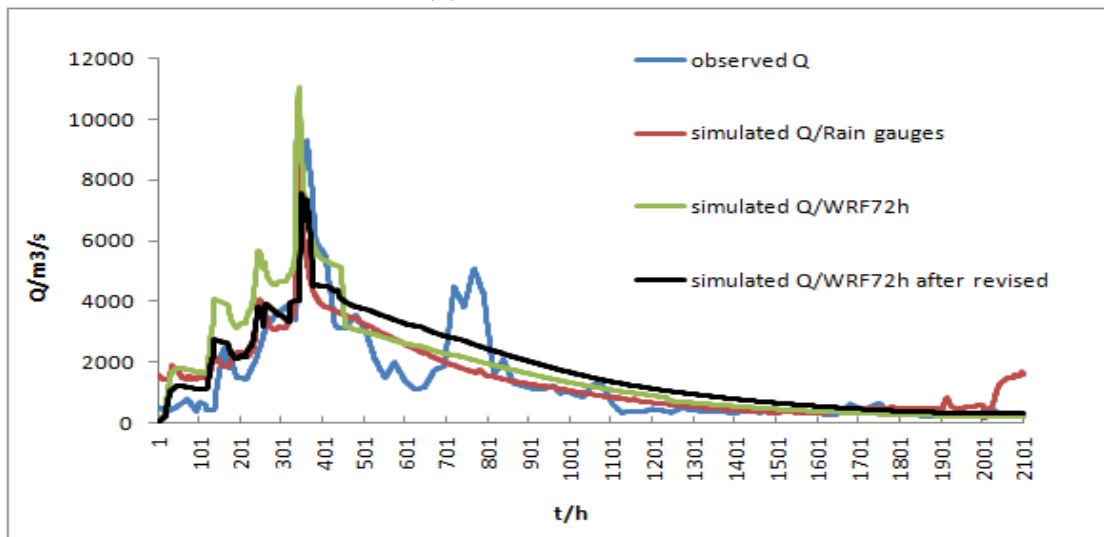
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(a) 24 hour lead time



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(b) 48 hour lead time

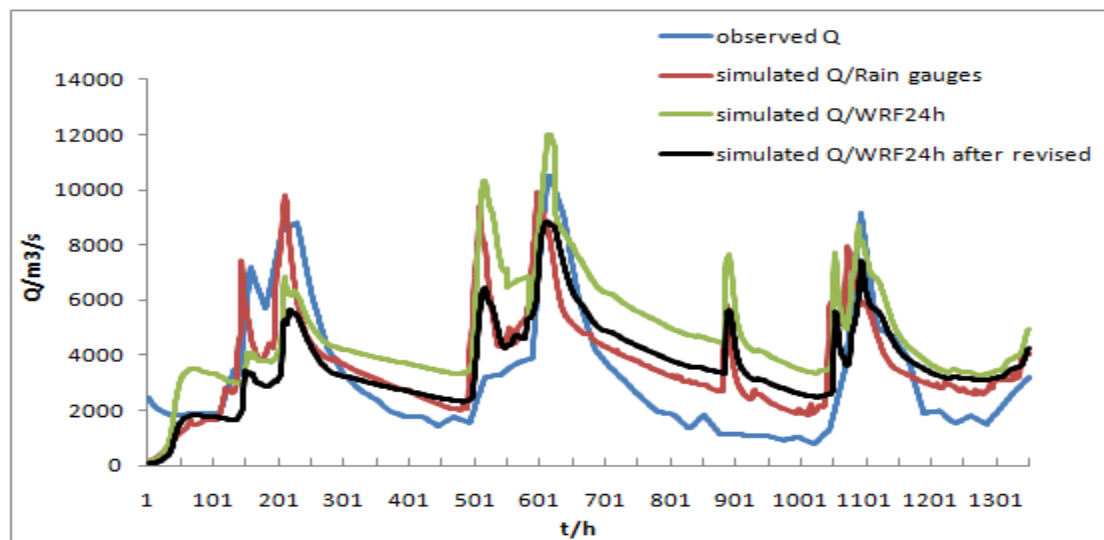


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(c) 72 hour lead time

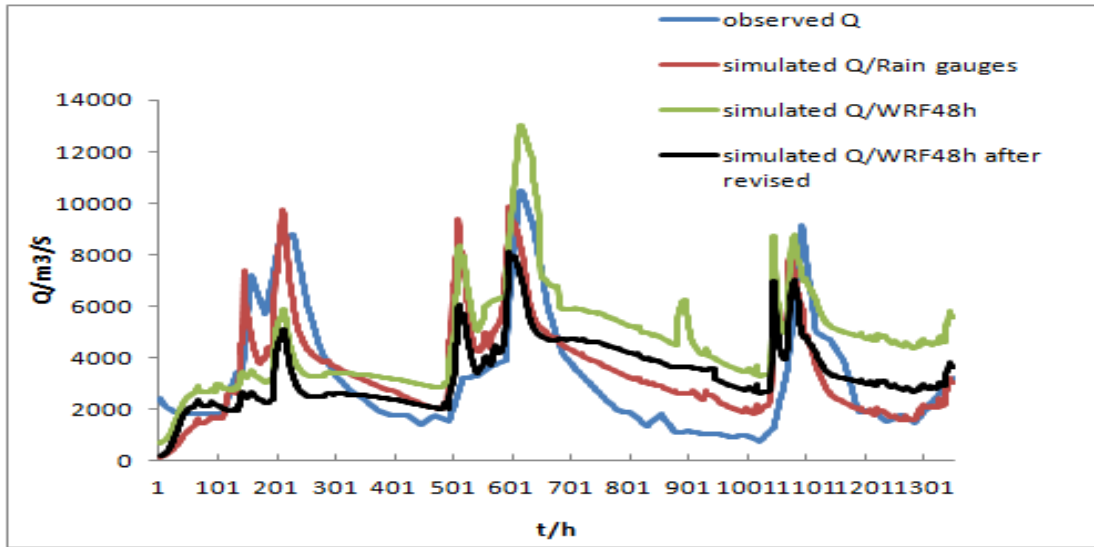
Fig. 7 Coupled flood simulation results with original model parameters (2011)

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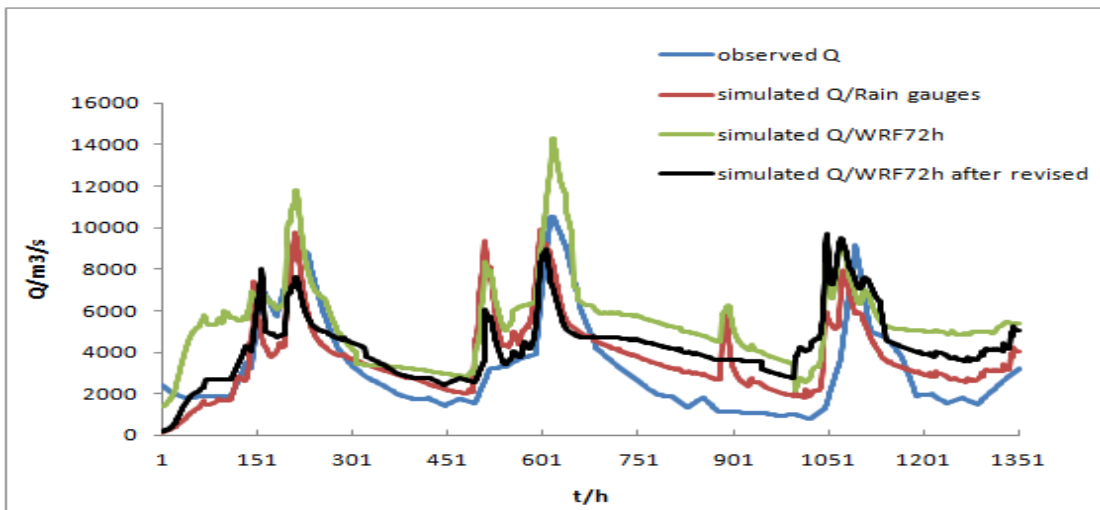
(a) 24 hour lead time



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(b) 48 hour lead time



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(c) 72 hour lead time

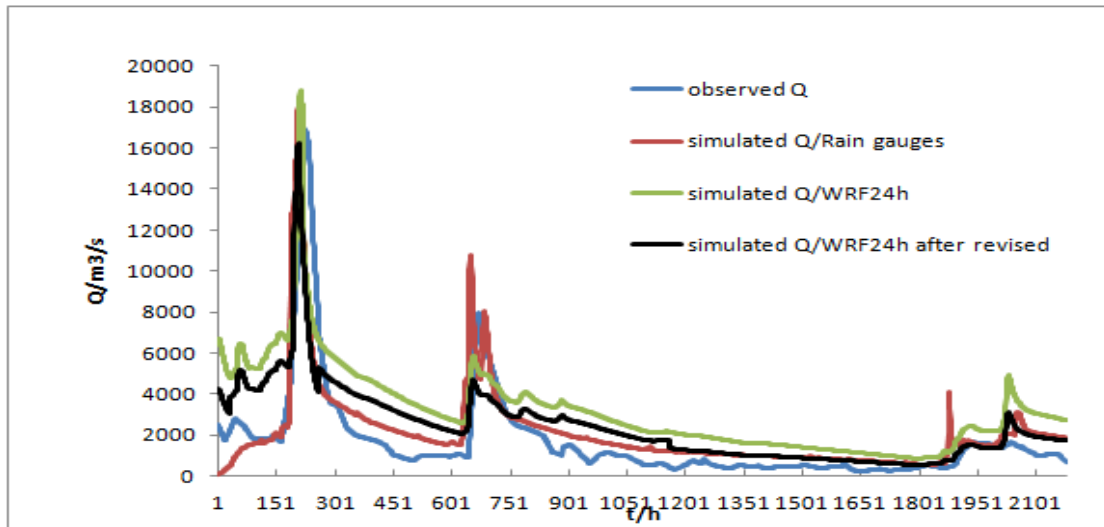
Fig. 8 Coupled flood simulation results with original model parameters(2012)

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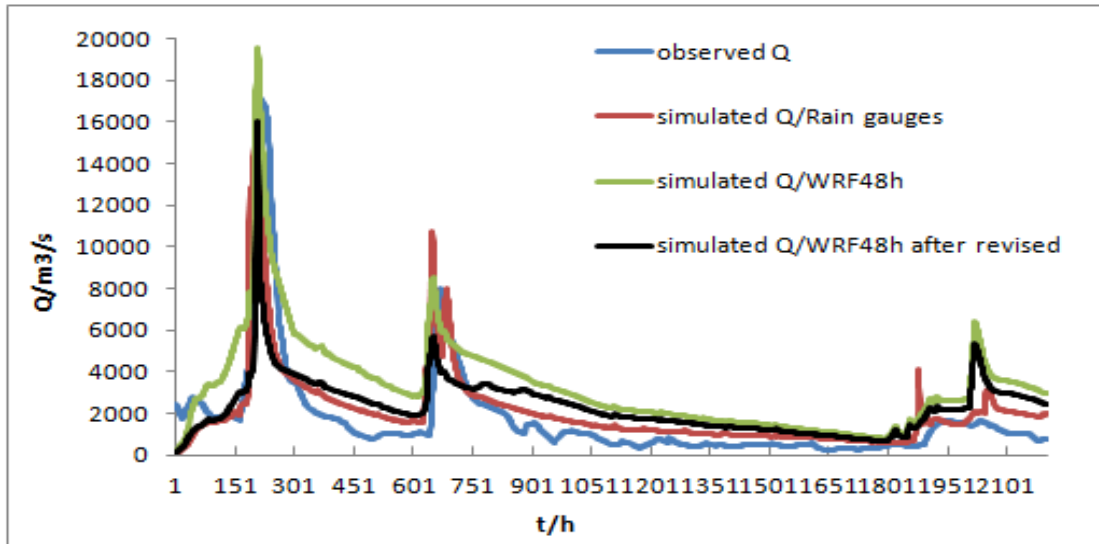
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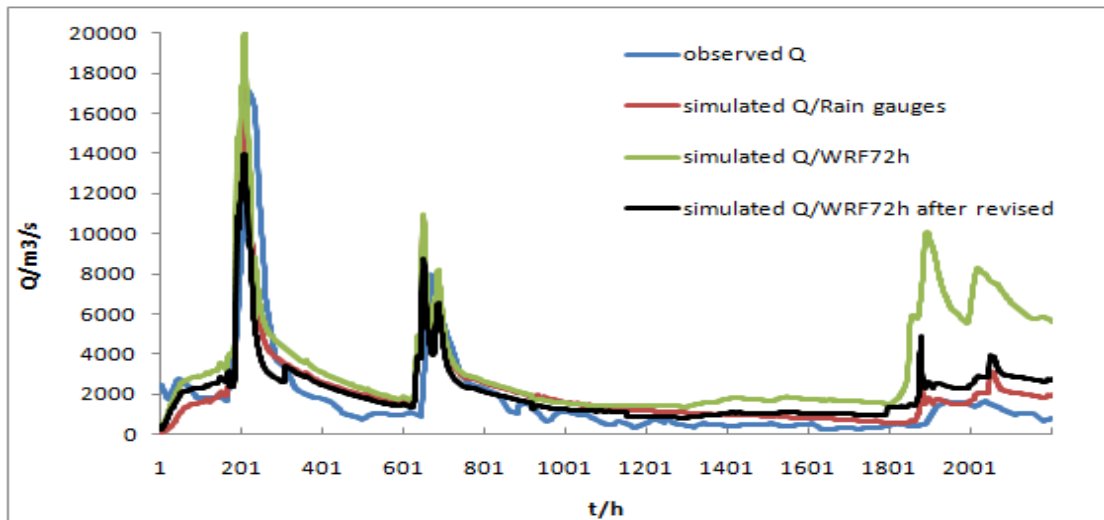
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(a) 24 hour lead time



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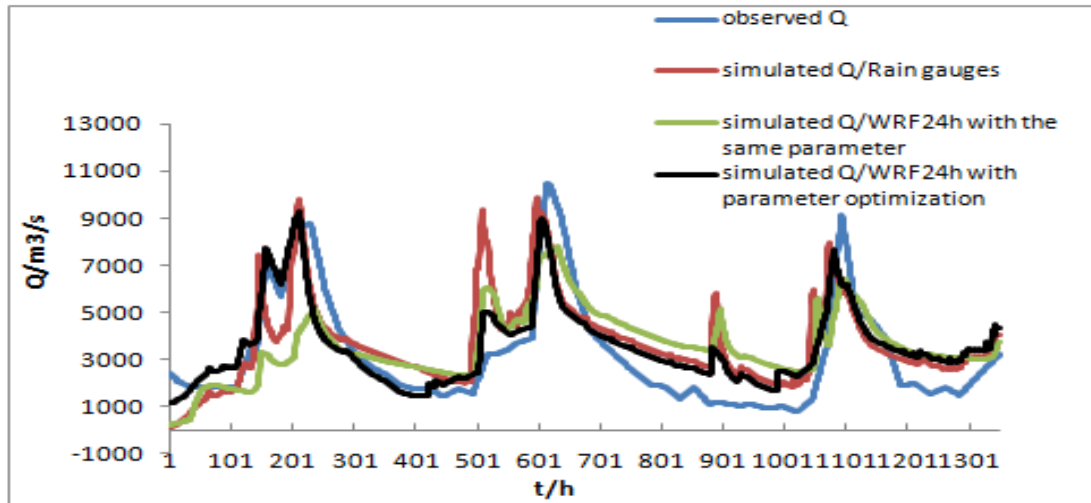
(b) 48 hour lead time



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(c) 72 hour lead time

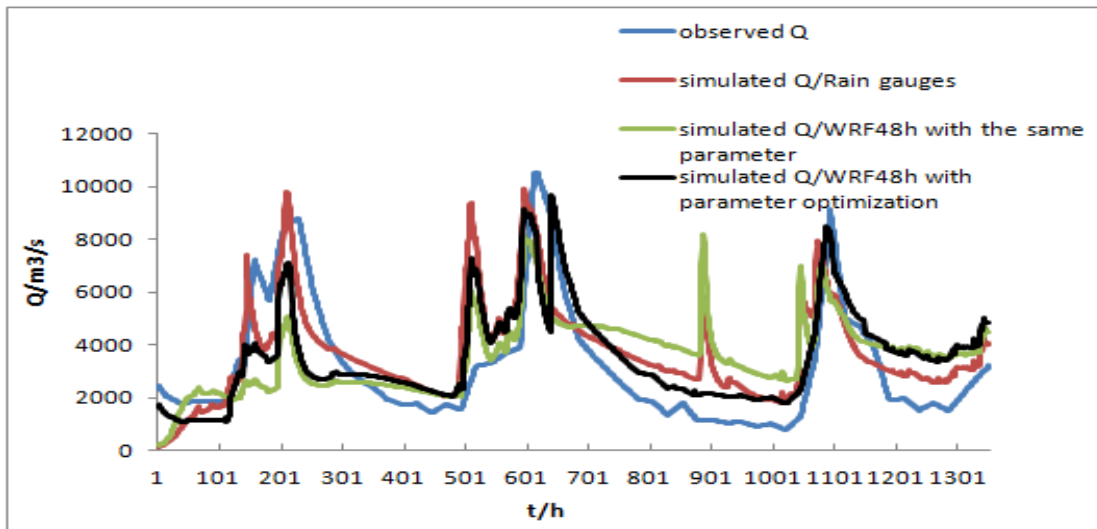
Fig. 9 Coupled flood simulation results with original model parameters (2013)



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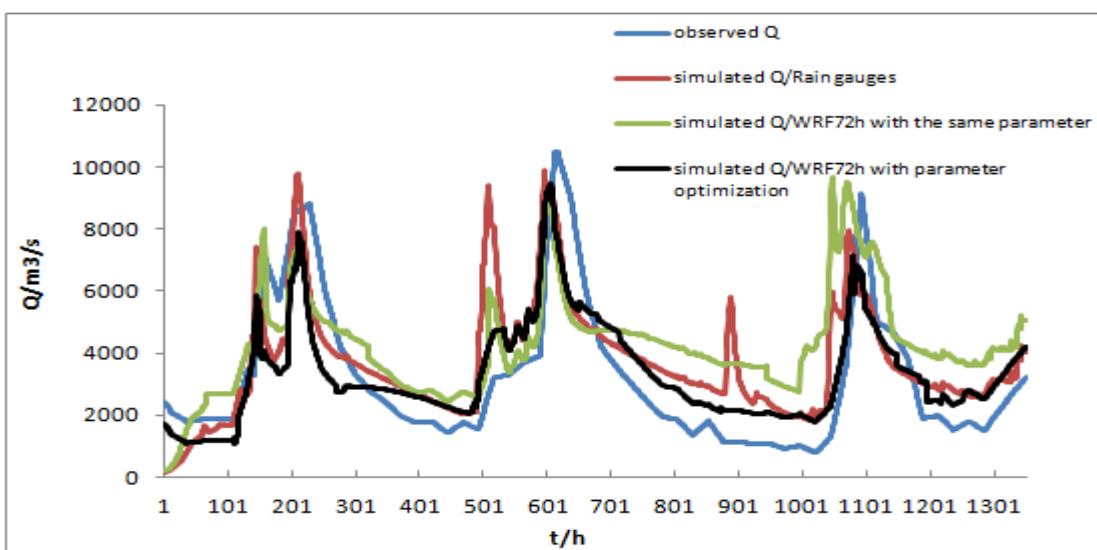
(a) 24 hour lead time



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(b) 48 hour lead time



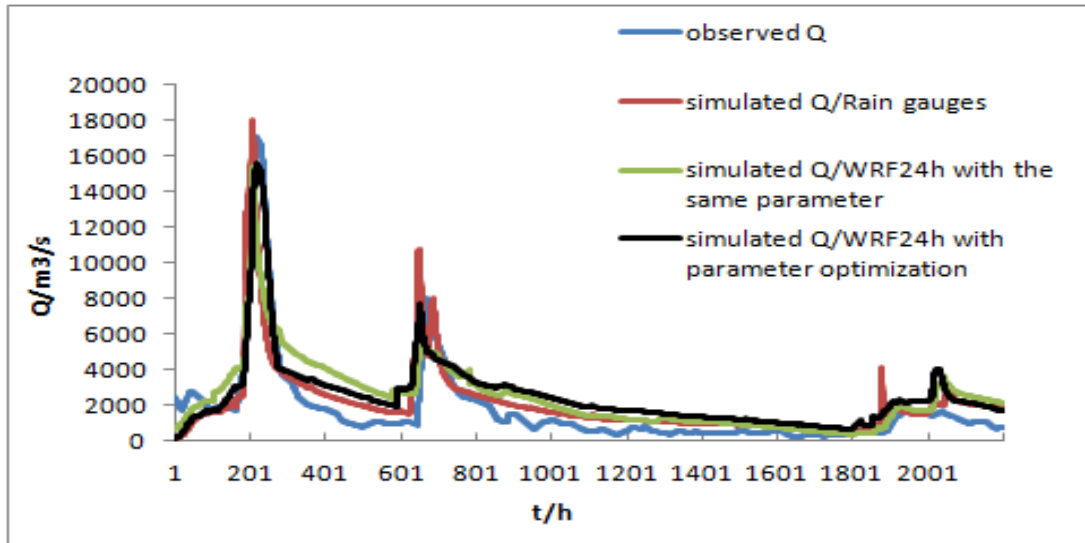
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(c) 72 hour lead time

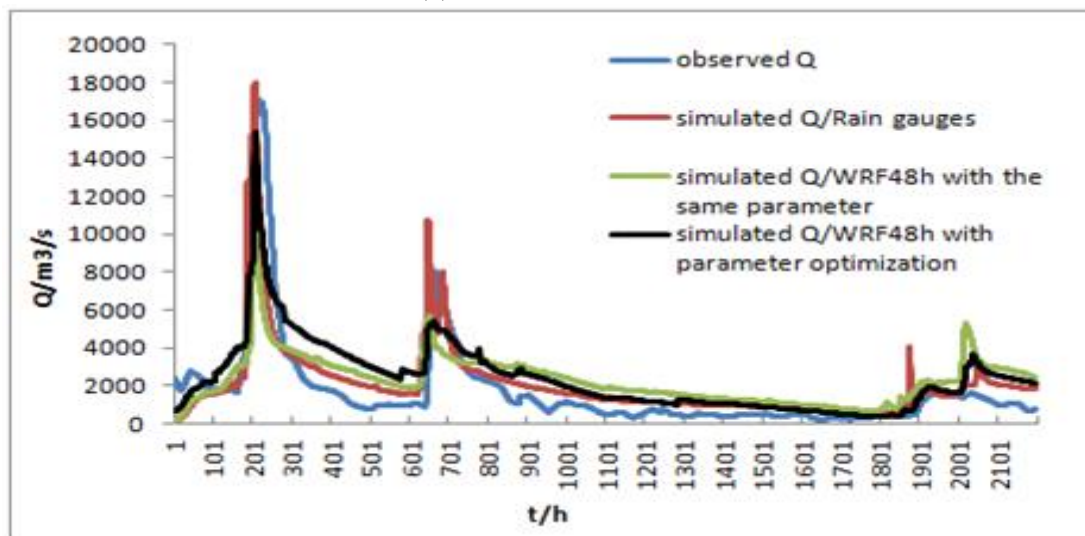
Fig. 10 Coupled flood simulation results with re-optimized model parameters (2012)



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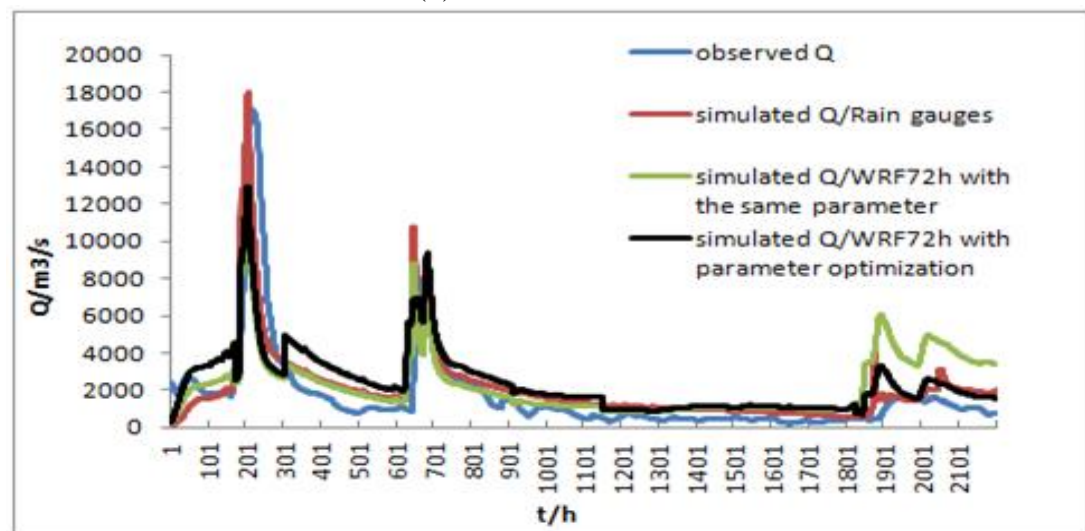
(a) 24 hour lead time



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(b) 48 hour lead time



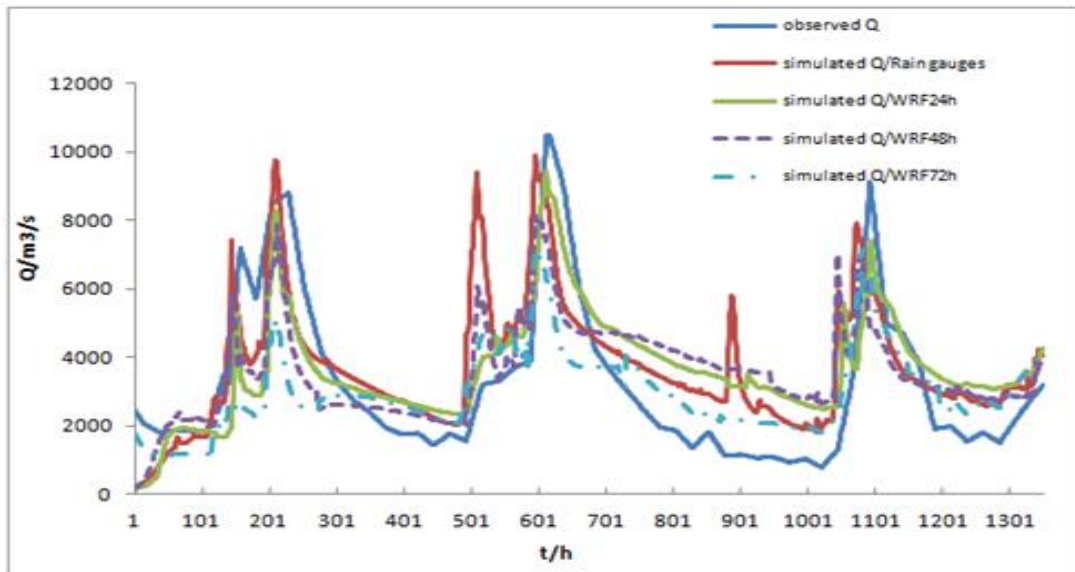
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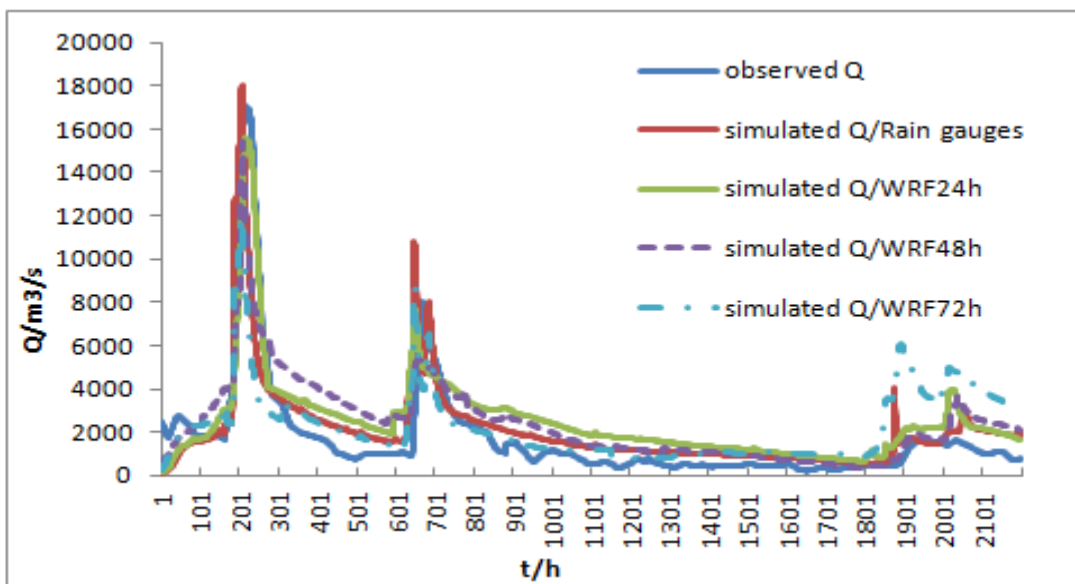
(c) 72 hour lead time

Fig. 11 Coupled flood simulation results with re-optimized model parameters (2013)



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(a) Flood event 2012



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(b) Flood event 2013

Fig. 12 Simulated results with different lead time

608 **Tables**

609 Table 1 Precipitation comparison of two products

Flood event no.	Precipitation products	average precipitation(mm)	relative bias %
2011	rain gauges	0.22	
	WRF/24h	0.27	23
	WRF/48h	0.29	32
	WRF/72h	0.34	55
2012	rain gauges	0.38	
	WRF/24h	0.44	16
	WRF/48h	0.52	37
	WRF/72h	0.65	71
2013	rain gauges	0.22	
	WRF/24h	0.33	50
	WRF/48h	0.38	73
	WRF/72h	0.43	95

610

611

612 Table 2 Evaluation indices of simulated flood events with post-processed WRF QPF

Rain type	statistical index	201101010	20120101	20130101
WRF/24h	Nash-Sutcliffe coefficient/C	0.65	0.66	0.65
	Correlation coefficient/R	0.88	0.73	0.83
	Process relative error/P	0.35	0.57	0.19
	Peak flow relative error/E	0.14	0.18	0.25
	The coefficient of water balance/W	1.44	1.35	1.38

WRF/24h after revised	Nash-Sutcliffe coefficient/C	0.75	0.75	0.75
	Correlation coefficient/R	0.93	0.82	0.85
	Process relative error/P	0.23	0.35	0.11
	Peak flow relative error/E	0.08	0.12	0.16
	The coefficient of water balance/W	1.15	1.08	1.12
WRF/48h	Nash-Sutcliffe coefficient/C	0.58	0.63	0.5
	Correlation coefficient/R	0.78	0.75	0.8
	Process relative error/P	0.52	0.48	0.34
	Peak flow relative error/E	0.41	0.12	0.24
	The coefficient of water balance/W	1.52	1.43	1.51
WRF/48h after revised	Nash-Sutcliffe coefficient/C	0.64	0.75	0.62
	Correlation coefficient/R	0.82	0.84	0.86
	Process relative error/P	0.45	0.26	0.22
	Peak flow relative error/E	0.34	0.08	0.13
	The coefficient of water balance/W	1.22	1.32	1.24
WRF/72h	Nash-Sutcliffe coefficient/C	0.45	0.48	0.44
	Correlation coefficient/R	0.68	0.36	0.75
	Process relative error/P	0.64	0.62	1.29

	Peak flow relative error/E	0.31	0.35	0.45
	The coefficient of water balance/W	1.67	1.54	1.66
WRF/72h after revised	Nash-Sutcliffe coefficient/C	0.52	0.58	0.55
	Correlation coefficient/R	0.75	0.45	0.82
	Process relative error/P	0.53	0.52	0.98
	Peak flow relative error/E	0.11	0.22	0.23
	The coefficient of water balance/W	1.15	1.14	1.25

613

614 Table 3 Evaluation indices of simulated flood event with different model parameters

parameter type	statistical index	201101010	20120101	20130101
Coupling model 24h/originally optimized model parameters	Nash-Sutcliffe coefficient/C	0.75	0.58	0.75
	Correlation coefficient/R	0.93	0.82	0.85
	Process relative error/P	0.23	0.35	0.11
	Peak flow relative error/E	0.08	0.12	0.16
	The coefficient of water balance/W	1.15	1.08	1.12
Coupling model24h /re- optimized model parameters	Nash-Sutcliffe coefficient/C	0.78	0.74	0.87
	Correlation coefficient/R	0.95	0.86	0.87
	Process relative error/P	0.19	0.28	0.09
	Peak flow relative error/E	0.06	0.08	0.12
	The coefficient of water balance/W	1.03	0.95	1.02

Coupling model 48h/originally optimized model parameters	Nash-Sutcliffe coefficient/C	0.64	0.75	0.62
	Correlation coefficient/R	0.82	0.84	0.86
	Process relative error/P	0.45	0.26	0.22
	Peak flow relative error/E	0.34	0.08	0.13
	The coefficient of water balance/W	1.22	1.32	1.24
Coupling model 48h /re-optimized model parameters	Nash-Sutcliffe coefficient/C	0.72	0.75	0.68
	Correlation coefficient/R	0.86	0.87	0.89
	Process relative error/P	0.32	0.22	0.18
	Peak flow relative error/E	0.21	0.06	0.09
	The coefficient of water balance/W	1.05	1.12	1.06
Coupling model 72h/originally optimized model parameters	Nash-Sutcliffe coefficient/C	0.52	0.75	0.55
	Correlation coefficient/R	0.75	0.45	0.82
	Process relative error/P	0.53	0.52	0.98
	Peak flow relative error/E	0.11	0.22	0.23
	The coefficient of water balance/W	1.15	1.14	1.25
Coupling model 72h /re-optimized model parameters	Nash-Sutcliffe coefficient/C	0.62	0.72	0.61
	Correlation coefficient/R	0.78	0.56	0.87
	Process relative error/P	0.38	0.32	0.75
	Peak flow relative error/E	0.09	0.18	0.17
	The coefficient of water balance/W	1.08	1.02	1.05

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616 Table 4 Evaluation indices of simulated flood event with different lead time

Rain type	statistical index	20120101	20130101
Rain gages	Nash-Sutcliffe coefficient/C	0.82	0.95
	Correlation coefficient/R	0.89	0.92
	Process relative error/P	0.2	0.08
	Peak flow relative error/E	0.05	0.06
	The coefficient of water balance/W	0.8	1.08
WRF/24h	Nash-Sutcliffe coefficient/C	0.74	0.87
	Correlation coefficient/R	0.86	0.87

	Process relative error/P	0.28	0.09
	Peak flow relative error/E	0.08	0.12
	The coefficient of water balance/W	0.95	1.02
WRF/48h	Nash-Sutcliffe coefficient/C	0.63	0.62
	Correlation coefficient/R	0.84	0.86
	Process relative error/P	0.48	0.22
	Peak flow relative error/E	0.12	0.13
	The coefficient of water balance/W	1.32	1.24
WRF/72h	Nash-Sutcliffe coefficient/C	0.56	0.61
	Correlation coefficient/R	0.56	0.87
	Process relative error/P	0.56	0.75
	Peak flow relative error/E	0.18	0.17
	The coefficient of water balance/W	1.54	1.66

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