Extending flood forecasting lead time in large watershed by coupling WRF QPF with

2

distributed hydrological model

3

Ji Li¹, Yangbo Chen¹, Huanyu Wang¹, Jianming Qin¹, Jie Li²

4	¹ Department	of	Water	Resources	and	Environment,	Sun	Yat-sen	University,
5	Guangzhou 51	1027	5, Chin	a					
6	² Hydrology B	urea	u, Pear	l River Wat	er Re	sources Commi	ission,	Guangzh	nou 510370,
7	China								

8

9 Correspondence to: Yangbo Chen (eescyb@mail.sysu.edu.cn)10

11 Abstract. Long lead time flood forecasting is very important for large watershed 12 flood mitigation as it provides more time for flood warning and emergency responses. Latest numerical weather forecast model could provide 1-15 days quantitative 13 14 precipitation forecasting products at grid format, by coupling this product with distributed hydrological model could produce long lead time watershed flood 15 16 forecasting products. This paper studied the feasibility of coupling the Liuxihe Model with the WRF QPF for a large watershed flood forecasting in southern China. The 17 QPF of WRF products has three lead time, including 24 hour, 48 hour and 72 hour, 18 f_{1} grid resolution f_{1} 20kmx20km. The Liuxihe Model is set up with freely 19 downloaded terrain properting model parameters were previously optimized with 20 rain gauge observed precipitation, and re-optimized with WRF QPF. Results show 21 22 that the WRF QPF has bias with the rain gauge precipitation, and a post-processing 23 method is proposed to post process the WRF QPF products, which improves the flood 24 forecasting capability. With model parameter re-optimization, the model's performance improves als $\boxed{12}$ suggests that the model parameters be optimized with 25 QPF, not the rain gauge precipitation. With the increasing of lead time, the accuracy 26 27 of WRF QPF decreases, so does the flood forecasting capability. Flood forecasting products produced by coupling Liuxihe Model with WRF QPF provides good 28

reference for large watershed flood warning due to its long lead time and rationalresults.

31

32 Key words : WRF, Liuxihe Model, Flood forecasting, lead time, parameter
 33 optimization

34

35 **1 Introduction**

36 Watershed flood forecasting is one of the most important non-engineering measures 37 for flood mitigation(Tingsanchali, 2012, Li et al., 2002), significant progresses in watershed flood forecasting $l_{\Pi=}^{2}$ been made in the past decades (Borga et al., 2011, 38 39 Moreno et al., 2013). Lead time is a key index for watershed flood forecasting, 40 especially for large watershed (Toth et al., 2000, Han et al., 2007). Only flood forecasting products with long lead time = seful as = puld provide enough time for 41 42 flood warning and flood emergency responses. In the long practice of flood forecasting, ground based rain gauge measured precipitation is the main input for 43 44 flood forecasting model, but as this kind of precipitation is the rainfall falling to the 45 ground already, so it has no lead time. This makes the watershed flood forecasting 46 with very short lead time (Jasper et al., 2002), and could not satisfy the requirement of 47 flood warning (Shim et al., 2002) in lead time, particularly in large watershed, thus reducing the value of the flood forecasting products in watershed flood mitigation. 48

49

50 The developed numerical weather prediction mo_{\square} in the past decades could provide 51 longer lead time quantitative precipitation forecast(QPF) product at grid form \square he 52 lead time for the latest weather prediction model could be as long as to 1~15 days 53 (Buizza, 1999, Ahlgrimm et al., 2016). By coupling the weather prediction model QPF with flood forecasting model, the flood forecasting lead time thus could be extended \square 54 this provides way for large watershed flood forecasting (Jasper et al., 2002, 55 Zappa et al., 2010, Giard and Bazile, 2000). Many numerical weather prediction 56 models have been proposed and put into operational use, such as the European Centre 57 58 Medium-Range Weather Forecasts (ECMWF) Ensemble Prediction System (EPS) 59 (Molteni et. al., 1996, Barnier et. al., 1995), the weather research and forecasting (WRF) model (Skamarock, 2005, 2008, Maussion, 2011), the numerical weather 60 61 forecast model of Japan Meteorological Agency (Takenaka et al., 2011, Gao and Lian, 2006), the numerical forecast model of China Meteorological Agency (Li and Chen, 62 2002), and others. 63

64

Watershed flood forecasting relies on $\frac{1}{12}$ drological model for computation tool, while 65 66 the precipitation is the model's driving force. The earliest hydrological model is regarded as the Sherman unit-graph (Sherman, 1932), which belongs to the category 67 of lumped hydrological model. Many lumped hydrological models have been 68 69 proposed, such as the Sacramento model (Burnash, 1995), the NAM model (DHI, 70 2004), the Xinanjiang model (Zhao, 1977), among others. The lumped hydrological model regards the watershed as a whole hydrological unit, thus the model parameter is 71 72 the same over the watershed, but this is not true, particularly for a large watershed. The precipitation the lumped hydrological model $\sqrt{12}$ is averaged over the watershed 73 also nis further increases the model's uncertainty in large watershed flood 74 forecasting as it is well known that the precipitation distribution over the watershed is 75 highly uneven. The QPF produced by numerical weather prediction model forecasts 76 77 precipitation at grid format, which provides detailed precipitation distribution information over watershe
is another advantage of QPF. The lumped hydrological
model could not take the advantage of gridded WPF products.

80

81 The latest development of watershed hydrological model is the distributed 82 hydrological model (Refsgaard et. al., 1996), which divides the watershed into grids, and different grids could have their own precipitation, terrain property and model 83 parameter, distributed hydrological model is the ideal model for coupling WRF 84 QPF for watershed flood forecasting. The first proposed distributed hydrological 85 model is SHE model (Abbott et. al.1986a, 1986b), and now many distributed 86 hydrological models have been proposed, and a few have been used for watershed 87 88 flood forecasting, such as the SHE model (Abbott et. al.1986a, 1986b), the 89 WATERFLOOD model (Kouwen, 1988), the VIC model (Liang et. al., 1994), the WetSpa model (Wang et. al., 1997), the Vflo model (Vieux et. al., 2002), the WEHY 90 91 model(Kavvas et al., 2004), the Liuxihe model (Chen et. al., 2009, 2011), among 92 others.

93

94 As distributed hydrological model calculates the hydrological process at grid scale, so 95 the computation time needed for runing the distributed hydrological model is huge even for a small watershed, ich limits the model's application in watershed flood 96 97 forecating, particularly in large watershed. Model parameter uncertainty related to 98 distributed hydrological model also impacted its application. But with the development of parallel computation algorithm for distributed hydr $\frac{2}{\Pi =}$ gical model 99 100 and its deployment on supercomputer (Chen et. al., 2013), the computation burden is 101 not a challenge of distributed hydrological modeling anymore. Also with the development of automatical parameter optimization of distributed hydrological model in flood for find (Madsen et. al., 2003, Shafii et. al., 2009, Xu et. al., 2012, Chen et. al., 2016), the model parameters could be optimized, and the model's performance could be improved largely. With these advances, now distributed hydrological model could be used for large watershed flood forecasting.

107

In this paper, the WRF QPF is coupled with t 108 Liuxihe model for a large watershed flood forecasting in southern China. The spatial 109 and temporal resolution of WRF QPF is at 20km*20km and 1 hour respectively with 110 three lead t_{\square} , including 24 hour, 48 hour and 72 hour. The WRF QPF has a similar 111 ttern with that estimated by rain gauges, but overestimates the averaged watershed 112 precipitation, and the longer the WRF QPF lead time, the higher the precipitation 113 overestimation \square RF QPF has systematic bias compared with rain gauge precipitation, 114 115 a post-processing method is proposed to post process the WRF QPF products, which improves the flood forecasting capability. The Liuxihe Model is set up with freely 116 downloaded terrain property e model parameters were previously optimized with 117 rain gauge observed precipitation, and re-optimized with WRF QPF. With model 118 parameter re-optimization, the model's performance improved $\frac{1}{1-1}$ odel parameters 119 should be optimized with QPF, not the rain gauge precipitation. Flood forecasting 120 121 reference for large watershed flood warning due to 122 123 results.

124 **2 Studied area and data**

125 2.1 Studied area

Liujiang River Basin(LRB) is selected as the studied area, which is the largest first order tributary of the Pearl River with a drainage area of 58270 km²(Chen et. al., 2016). LRB is in the monsoon area with heavy storms that ind f s severe flooding in the watershed, and caused huge flood damages in the past centuries. Fig. 1 is a sketch map of LRB.

131

132

Fig. 1 is here

133

134 2.2 Rain gauge precipitation and river flow discharge

Precipitation of 68 rain gauges within the watershed in 2011, 2012 and 2013 was collected and used in this study to compare with the WRF QPF. Precipitation data one hour interval. River discharge near the watershed outlet is collected also for this same period. As this study for on watershed flood forecasting, so only the precipitation and river discharge during the flood events are prepared. There is one flood event in each year e flood events are numbered as flood event 2011, flood event 2012 and flood event 2013 respectively.

142 **3 WRF Q** and st-processing

143 3.1 WRF model

144 The WRF model (Skamarock et. al., 2005, 2008) is considered as the next 145 generation's medium term weather forecasting model, and can simulate different weather processes from cloud scale to synoptic scale, especially in horizontal resolution of $1 \sim 10$ km. Also, it integrates the advanced numerical methods and data assimilation techniques, a variety of phy ly process schemes, and multiple nested methods and the capability of being used in different geographical locations. The development of WRF model satisfies the needs of scientific research and practical application, and could be further improved and strengthened. Now WRF model has replaced the previously used MM5 model.

153

Many studies have been carried out in quantitative precipitation forecasting by using 154 WRF mode pr example, Kumar et al. (2008) used WRF model to study a heavy rain 155 in 200 is result showed that WRF system could reproduce the storm event and its 156 dynamical and thermo-dynamical characteristics. Hong and Lee (2009) set up a triply 157 nested WRF model to simulate the initiation of a thunderstorn 158 sensitivity test. Maussion et. al. (2011) compared the capability of WRF model in 159 retrieving monthly precipitation and snowfall at three different spatial resolution 160 including 30 = 0 and 2 km 1 = 0 result showed that WRF model had a good 161 162 performance in simulating monthly precipitation and snowfall in Tibet. Givati et al. (2012) predicted the hiemal precipitation event of 2008 and 2009 based on WRF 163 164 model in upstream of the Jordan River, and coupled WRF model with hydrological model-HYMKE to simulate the velocity and discharge of Jordan River. Pennelly et. al. 165 (2014) employed WRF model to predict three precipitation events of Alberta, 166 Canada, and compared the precipitation with 48 hour $lea \begin{bmatrix} 1 \\ 1 \end{bmatrix}$ g time predicted by WRF 167 model in the precipitation observed by rain gauges in the result showed that 168 Kain-Fritsch scheme overestimated the value of precipitation inv $\frac{1}{12}$ bly. Zhang (2004) 169

introduced the WRF version 2 and grapes 3d variation assimilation 170 and real-time forecasting results of weather conditions showed that WRF model had a 171 172 good performance in forecasting all kinds of weather conditions and had the ability to predict the air quality. Niu et. al. (2007) tested the sensitivity of microphysical sch 173 to a typical heavy rain based on WRF model, and analyzed the performance of 174 recipitation predicted from the precipitation region, $\frac{1}{12}$ nter position and rainfall 175 intensity. Xu et. al. (2007) compared the hiemal continuous precipitation process 176 predicted with the estival results by WRF mode $\begin{bmatrix} 2 \\ 1 \end{bmatrix}$ he results showed that the KF 177 scheme was better than BM scheme in summer. Hu et. al. (2008) found that the 178 179 parameterization scheme of WRF model was related to the model resolution, and the 180 parameterization scheme should be selected by the resolution of WRF model. Huang et. al. (2011) found that variations in the microphysical process 181 parameterization schemes had much more influence on precipitation than that of 182 cumulus parameterization schemes, especially for a torrential rain attributed to 183 184 large-scale forcing that mainly resulted from stratus clouds. Wang and Ma (2011) introduced the application of WRF model fr_{Π}^2 the physics parameterization scheme, 185 real-time simulation study and the comparison with MM5 model in China in recent 186 decade. Pan et. al. (2012) used two WRF simulation groups between pre-process and 187 post-process in Heihe river basin, and compared and analyzed the mean bias error, 188 root mean square error and correlation coefficient of the two WRF groups. 189

190 $3.2 \text{ WRF } Q_{\square} = \text{ of } LRB$

The WRF model (version 3) was set up in LRB by Li et. al. (2014) e model domain is centered at 23.8N, 109.2W, and the projection is Lambert conformal projection. The vertical structure includes 28 layers covering the whole troposphere. The WRF

single-moment 3-class microphysics parameterization, i.e., Kain-Fritsch (Kain, 194 2004) and cumulus parameterization (Hong and Lim, 2006) were adopted for 195 precipitation simulation. The parameterization $\operatorname{sch}_{\square=}^{\square=}$ of WRF $\operatorname{the}_{\square=}^{\square=}$ hore than that of 196 other mesoscale numerical weather prediction (NWP 197 physical parameterization schemes: microphysical process, cumulus, land surface 198 processes, atmospheric radiation and planetary boundary layer. There are 13 199 microphysical process parameterization scheme $\frac{1}{10}$ urdue Lin scheme was used in this 200 study as microphysical process. The parameterization scheme of precipitation was 201 improved based on the scheme n = n et al. (1983) as well as Rutledge and Hobbs (1983), 202 which is more mature than other schemes and is suited to simulate the high resolution 203 real time data. 204

205

The spatial and temporal resolution of WRF is at 20km*20km and 1 hour respectively so there are 156 WRF grids in LRB. QPF products in 2011, 2012 and 2013 were produced at 3 different lead time respectively 24 hours, 48 hours and 72 hours. Fig. 2, 3 and 4 are WRF QLE n three different years, while (a) is the rain gauge precipitation, (b) is the WRF QPF with 24 hour lead time, (c) is the WRF QPF with 48 hour lead time, and (d) is the WRF QPF with 72 hour lead time.

- 213 Fig. 2 is here
- Fig. 3 is here
- 215 Fig. 4 is here
- 216

217	3.3	Compa	rison c	of WRF	OPF	and rain	gauges	precip	oitation
					•		G		

218	WRF QPF and rain gauge precipitation are compared in this study. From the results of
219	Fig. 2, 3 and 4, it could found that the temporal precipitation pattern of both
220	products are similar $\begin{bmatrix} 1 \\ 1 \end{bmatrix}$ ere are some kinds of differences, but the difference $\begin{bmatrix} 1 \\ 1 \end{bmatrix}$ not
221	significant. To make further comparison, the accumulated precipitation of the three
222	flood events averaged over the watershed are calculated and listed in Table 1.
223	
224	Table 1 is here
225	
226	From the results of Table 1, it could be t and that the WRF QPF accumulated
227	precipitation has obvious bias with rain gauge accumulated precipitation. For all the
228	three flood events, the WRF QPF accumulated precipitation are higher than those
229	estimated by rain gauge, i. e., the WRF QPF overestimates the precipitation. For flood
230	event 2011, the overestimated watershed averaged precipitation of WRF QPF with
231	lead time of 24 hour, 48 hour and 72 hour are 23%, 32% and 55% respectively, for
232	flood event 2012, they are 16%, 37% and 71% respectively, for flood event 2013
233	are 50%, 73% and 95% respectively
234	lead time, the higher the overestimation.

235 3.4 WRF QPF post-processing

From the results of Fig. 2, 3 and 4, and Table 1, the WRF QPF has significant bias

237 with rain gauge precipitation. If the rain gauge precipitation is assumed correct, then

238 WRF QPF has error. So in this study the WRF QPF is post-processed based on the

rain gauge precipitation to correct the systematic error of WRF QPF. The principle of

WRF QPF post-processing proposed in this study is to keep the areal averaged event
accumulated precipitation from both products are similar, i.e., to adjust the WRF QPF
precipitation to make its event accumulated precipitation equal to that of rain gauge.
Based on this principle, the WRF QPF post-processing procedure is summarized as
follows:

245

248
$$\overline{P}_{WRF} = \frac{\sum_{i=1}^{N} P_i F_i}{N} \quad (1)$$

249 Where, \overline{P}_{WRF} is the areal average precipitation of WRF QPF of one flood event, P_i is 250 the precipitation on WRF grid i, F_i is the drainage area of WRF grid i, *N* is the total 251 number of WRF grids.

252 2) Calculate the areal average precipitation of the rain gauges with the following253 equation.

254
$$\overline{P}_2 = \frac{\sum_{j=1}^M P_j}{M}$$
 (2)

255 Where, \overline{P}_2 is the areal average precipitation of the rain gauges network, P_j is the 256 precipitation observed by jth rain gauge, *M* is the total number of rain gauges. 257 3) The precipitation of every WRF QPF groups then could be revised with the

- 11 -

259
$$P_i' = P_i \frac{\overline{P_2}}{\overline{P_{WRF}}}$$
 (3)

260 Where, P_i is the revised precipitation of ith WRF grid.

With the above WRF QPF post-processing method, the WRF QPF of flood event 2011, 262 2012 and 2013 were post-processed, and will used to couple with the Liuxihe 263 Model for flood simulation.

264 3 Hydrological model

265 3.1 Liuxihe Model

266 Liuxihe model is a physically based fully distributed hydrological model proposed mainly for watershed flood forecasting (Chen, 2009, Chen et. al., 2011), and has been 267 used in a few watershed bod forecasting (Chen, 2009, Chen et. al., 2011, 2013, 2016, 268 Liao et. al., 2012 a, b, Xu et. al., 2012 a, b). In Liuxihe Model, runoff components are 269 calculated at grid scale, runoff routes at both grid and watershed scale. Runoff routing 270 is divided into hill \square ope routing and river channel routing by using different 271 computation algorithm. Liuxihe Model proposed an automatic parameter optimization 272 method using PSO algorithm (Chen et. al., 2016), which largely improves the model's 273 274 performance in watershed flood forecasting. Now Liuxihe Model is deployed on a supercomputer system with parallel computation techniques (Chen et. al., 2013) that 275 largely facilitates the model parameter optimization of Liuxihe Model. 276

277

278 Chen et. al. (2016) set up Liuxihe Model in LRB with freely downloaded terrain 279 property data from the website at a spatial resolution of 200m*200m, and optimized 280 model parameters with observed hydrological data. The model was validated by



- 13 -

305 with WRF QPF improved $\mathbf{\overline{1}}$ ich than that simulated with $\mathbf{\overline{1}}$ n gauge precipitation $\mathbf{\overline{1}}$ is 306 means parameter optimization with WRF QPF is necessary.

307 3.3 Coupling WRF QPF with Liuxihe Model for LRB flood forecasting

- Liuxihe Model set up for LRB flood forecasting (Chen et. al., 2016) mployed
 to couple with the WRF QPF, the model spatial resolution reference is to be 200m*200m.
 As the spatial resolution of WRF QPF is at 20km*20km, the WRF QPF was
 downscaled to the resolution of 200m*200m by using the nearest downscaling
- 312 method, the same spatial resolution of the flood forecasting model.

313 4 Results and discussions

314 4.1 Effects of WRF post-processing

- The original WRF QPF and the post-processed QPF and the couple with the Liuxihe Mode to the simulation, the original model parameters that the ptimized with the rain gauge precipitation to the re-optimized model parameters be simulated results are shown in Fig. 7, 8 and 9. Fig. 7 is here Fig. 8 is here Fig. 9 is here
- 323

From the above results, it could be seen that the simulated flood discharges with the original WRF QPF \square nuch lower than the observed one \square ut with post-processed WRF QPF used, the simulated flood discharge increased and \square uch more close to the observation is implies that the flood forecasting capability has been improved by post-processing of WRF QPF. To further compare the three results, 5 evaluation indices, including Nash-Sutcliffe coefficient(C), correlation coefficient(R), process relative error(P), peak flow relative error(E) and water balance coefficient(W) calculated and listed in Table 2.

- 332
- 333

Table 2 is here

334

335 From the results of Table 2, it has been found that all the 5 evaluation indices have been improved by coupling the post-processed WRF QPF. For example, \int_{1}^{2} lood 336 event 2011 with 24 hour lead time, the Nash-Sutcliffe coefficient/C, correlation 337 coefficient/R, process relative error/P, peak flow relative error/E and coefficient of 338 water balance/W with original WRF QPF are 0.65, 0.88, 35%, 14% and 1.44 339 respectively, but those with the post-processed WRF QPF are 0.75, 0.93, 23%, 8% 340 and 1.15 respectively. $\int_{1}^{2} =$ flood event 2012 with 48 hour lead time, the above 5 341 evaluation indices with original WRF QPF are 0.63, 0.75, 48%, 12% and 1.43 342 respectively, and are 0.75, 0.84, 26%, 8% and 1.32 respectively with the 343 post-processed WRF QPF. post-processed WRF QPF. 344 evaluation indices with original WRF QPF are 0.44, 0.75, 129%, 45% and 1.66 345 respectively, and are 0.55, 0.82, 98%, 23%, 1.25 respectively with the post-processed 346 WRF OPF. It is obvious that with the post-processed WRF OPF, the evaluation 347 indices are improved much pre. These results show that WRF QPF post processing 348 could improve the flood forecasting capability because the WRF $\begin{bmatrix} 2\\ 1 \end{bmatrix}$ PF is more close 349 to the observed precipitation after post-processing $\begin{bmatrix} 1 \\ 1 \end{bmatrix}$ it should be d $\begin{bmatrix} 1 \\ 1 \end{bmatrix}$ for real-time 350 flood forecasting. 351

4.2 Results comparison for different model parameters

The model parameters optimized with rain gauge precipitation and WRF QPF are different different parameter will here different model performance. To analyze this effect, the flood events of 2012 and 2013 with two different sets of model parameter e simulated, and are shown in Fig. 10 and Fig. 11 respectively.

- 358
- 359

Fig. 10 is here

360

Fig. 11 is here

361

From the above sults it has been found that the simulated flood results with re-optimized model parameters f etter than the simulated with the original model parameter f is simulated flood discharge with the re-optimized model parameter more fitting the observation. To further compare the two results, 5 evaluation indices, including Nash-Sutcliffe coefficient(C), correlation coefficient(R), process relative error(P), peak flow relative error(E) and water balance coefficient(W) are calculated and listed in Table 3.

- 369
- 370

Table 3 is here

371

From the results of Table 3, it hat the results of flood simulation based on the re-optimized model parameters have better evaluation indices. All evaluation indices for the based on re-optimized model parameters seample, for flood event 2012 with 24 hour lead time, the Nash-Sutcliffe

coefficient/C, correlation coefficient/R, process relative error/P, peak flow relative 376 error/E and coefficient of water balance/W with original model parameters are 0.58, 377 0.82, 35%, 12% and 1.08 respectively, but those with the re-optimized model 378 parameters are 0.74, 0.86, 28%, 8% and 0.95 respectively. For flood event 2013 with 379 48 hour lead time, the 5 indices with the original model parameters are 0.62, 0.86, 380 22%, 13% and 1.24 respectively, and are 0.68, 0.89, 18%, 9% and 1.06 respectively 381 for those with re-optimized model parameters. So it could be said that in coupling the 382 WRF QPF with distributed hydrological model, the model parameters n_{\square} to be 383 re-optimized with the WRF QPF. This finding implies that the precipitation pattern 384 has obvious impact nodel parameters, should be considered, and model 385 parameter optimization is a rational way for considering this effect. 386

387

388 4.3 Flood simulation accuracy with different lead time

To compare the model performance with different lead time, the flood events with 3 different lead time simulated and shown in Fig. 12 e model parameters are the re-optimized ones, and the QPF is the post-processed QPF.

392

393

Fig. 12 is here

394

From the results of Fig. 12, it could be seen that the flood simulation results is worse as the lead time increases, i.e., the model performance with 24 hour lead time is better than that with 48 hour lead time, and the model performance with 48 hour lead time is better than that with 72 hour lead time. The simulated hydrological process with 24 hour lead time is very similar where the simulated with rain gauge precipitation. To further compare the results, 5 evaluation indices, including Nash-Sutcliffe

401	coefficient(C), correlation coefficient(R), process relative error(P), peak flow relative
402	error(E) and water balance coefficient(W) $figure$ calculated and listed in Table 4.
403	

Table 4 is here

405

From the results of Table 4, it has not found that the simulated flood events with 24 406 hour lead time h_{\square}^{\square} pest evaluation indices, and h_{\square}^{\square} ery close to h_{\square}^{\square} simulated with 407 rain gauge precipitation. For flood event 2012, the Nash-Sutcliffe coefficient/C, 408 409 correlation coefficient/R, process relative error/P, peak flow relative error/E and coefficient of water balance/W with rain gauge are 0.82, 0.89, 20%, 5% and 0.8 410 respectively, $\frac{1}{12}$ pse with 24 hour lead time are 0.74, 0.86, 28%, 8% and 0.95 411 respectively, those with 48 hour lead time are 0.63, 0.84, 48%, 12% and 1.32 412 respectively, and are 0.56, 0.56, 56%, 18% and 1.54 respectively for 72 hour lead time. 413 For flood event 2013, the Nash-Sutcliffe coefficient/C, correlation coefficient/R, 414 process relative error/P, peak flow relative error/E and coefficient of water balance/W 415 with rain gauge are 0.95, 0.92, 8%, 6% and 1.08 respectively 416 time are 0.87, 0.87, 9%, 12% and 1.02 respectively, those with 48 hour lead time are 417 0.62, 0.86, 22%, 13% and 1.24 respectively, and are 0.61, 0.87, 75%, 17% and 1.66 418 respectively for 72 hour lead time. This finding means that the current WRF QPF 419 420 capability is lead-time dependent, and with the increasing of lead time, the practical value of WRF QPF gets lower. 421

422 **5 Conclusion**

423 In this study, the WRF QPF \square oupled with a distributed hydrological model-the 424 Liuxihe mode large watershed flood forecasting, and three lead tip of WRF QPF products, including 24 hours, 48 hours and 72 hours are tested. WRF QPF post processing method is proposed and tested, model parameters are re-optimized by using the post-processed WRF QPF, model performed are compared among v conditions. Based on the results of this study, the following conclusions could be drawn:

430

431 1) The quantitative precipitation forecasting produced by WRF model has a similar pattern with that estimated by rain gauges temporally, but overestimated the averaged 432 watershed precipitation $\begin{bmatrix} 1 \\ 1 \\ 1 \end{bmatrix}$ he event accumulated total precipitation $\begin{bmatrix} 1 \\ 1 \\ 2 \end{bmatrix}$ d the longer 433 the WRF QPF lead time, the higher the precipitation overestimation. For flood event 434 2011, the overestimated watershed averaged precipitation of WRF QPF with lead t_{Π} 435 of 24 hour, 48 hour and 72 hour are 23%, 32% and 55% respectively 436 2012, these are 16%, 37% and 71% respectively, while for flood event 2013, these are 437 50%, 73% and 95% respectively. 438

439

440 2. WRF QPF has systematic bias compared with rain gauge precipitation, and this 441 bias could be reduced via post-processing. Principle used in this study for WRF QPF 442 post processing is effective and could improve the flood forecasting capability. For 443 flood event 2011 with 24 hour lead time, the Nash-Sutcliffe coefficient/C, correlation coefficient/R, process relative error/P, peak flow relative error/E and coefficient of 444 445 water balance/W with original WRF QPF are 0.65, 0.88, 35%, 14% and 1.44 respectively, but those with the post-processed WRF QPF are 0.75, 0.93, 23%, 8% 446 and 1.15 respectively. For flood event 2012 with 48 hour lead time, the above 5 447 448 evaluation indices with original WRF QPF are 0.63, 0.75, 48%, 12% and 1.43 449 respectively, and are 0.75, 0.84, 26%, 8% and 1.32 respectively with the

- 19 -

post-processed WRF QPF. For flood event 2013 with 72 hour lead time, the above 5
evaluation indices with original WRF QPF are 0.44, 0.75, 129%, 45% and 1.66
respectively, and are 0.55, 0.82, 98%, 23%, 1.25 respectively with the post-processed
WRF QPF.

454

3. Hydrological model parameters optimized with the rain gauge precipitation $n_{\square=1}^{\square=1}$ 455 to be re-optimized using the post-processed WRF QPF, this improves the model 456 performance $\arg e_{\Pi}$, in coupling stributed hydrological model with QPF for 457 flood forecasting, the model parameters should be optimized with the QPF produced 458 by WRF. For flood event 2012 with 24 hour lead time, the Nash-Sutcliffe 459 460 coefficient/C, correlation coefficient/R, process relative error/P, peak flow relative 461 error/E and coefficient of water balance/W with original model parameters are 0.58, 0.82, 35%, 12% and 1.08 respectively, but those with the re-optimized model 462 parameters are 0.74, 0.86, 28%, 8% and 0.95 respectively. For flood event 2013 with 463 464 48 hour lead time, the 5 indices with the original model parameters are 0.62, 0.86, 22%, 13% and 1.24 respectively, and are 0.68, 0.89, 18%, 9% and 1.06 respectively 465 466 for those with re-optimized model parameters.

467

4. The simulated floods by coupling WRF QPF with distributed hydrological model rational and could benefit the flood management communities due to find onger lead
time for flood warning provides a good reference for large watershed flood warning.
But with the lead time getting longer, the flood forecasting accuracy is getting lower.
For flood event 2012, the Nash-Sutcliffe coefficient/C, correlation coefficient/R,
process relative error/P, peak flow relative error/E and coefficient of water balance/W
with rain gauge are 0.82, 0.89, 20%, 5% and 0.8 respectively, find se with 24 hour lead

475 time are 0.74, 0.86, 28%, 8% and 0.95 respectively, those with 48 hour lead time are 0.63, 0.84, 48%, 12% and 1.32 respectively, and are 0.56, 0.56, 56%, 18% and 1.54 476 respectively for 72 hour lead time. For flood event 2013, the Nash-Sutcliffe 477 478 coefficient/C, correlation coefficient/R, process relative error/P, peak flow relative error/E and coefficient of water balance/W with rain gauge are 0.95, 0.92, 8%, 6% 479 and 1.08 respectively, $\begin{bmatrix} -2 \\ -2 \end{bmatrix}$ se with 24 hour lead time are 0.87, 0.87, 9%, 12% and 1.02 480 respectively, those with 48 hour lead time are 0.62, 0.86, 22%, 13% and 1.24 481 482 respectively, and are 0.61, 0.87, 75%, 17% and 1.66 respectively for 72 hour lead 483 time.

484 Acknowledgements

This study is supported by the Special Research Grant for the Water Resources Industry (funding no. 201301070), the National Science Foundation of China (funding no. 50479033), and the Basic Research Grant for Universities of the Ministry of Education of China (fundingno.13lgjc01).

Figures







(b)







Fig. 2 Precipitation pattern comparison of two precipitation products(2011), (a) is the average precipitation of rain gauges, (b) is the average precipitation of WRF with 24 hour lead time, (c) is the average precipitation of WRF with 48 hour lead time, (d) is the average precipitation of WRF with 72 hour lead time.

502









H

Fig. 3 Precipitation pattern comparison of two precipitation products(2012), (a) is the
average precipitation of rain gauges, (b) is the average precipitation of WRF with 24
hour lead time, (c) is the average precipitation of WRF with 48 hour lead time, (d) is
the average precipitation of WRF with 72 hour lead time.

t/h

(d)











Fig. 4 Precipitation pattern comparison of two precipitation products(2013), (a) is the average precipitation of rain gauges, (b) is the average precipitation of WRF with 24 hour lead time, (c) is the average precipitation of WRF with 48 hour lead time, (d) is the average precipitation of WRF with 72 hour lead time.





535 Fig.5 Liuxihe Model structure of LRB (200m ×200m resolution, Chen et. al., 2016)



(a) Evolutionary process of objective function











585 Tables

Flood event no.	Precipitation products	average precipitation(mm)	relative bias %
	rain gauges	0.22	
2011	WRF/24h	0.27	23
2011	WRF/48h	0.29	32
	WRF/72h	0.34	55
	rain gauges	0.38	
2012	WRF/24h	0.44	16
2012	WRF/48h	0.52	37
	WRF/72h	0.65	71
	rain gauges	0.22	
2012	WRF/24h	0.33	50
2015	WRF/48h	0.38	73
	WRF/72h	0.43	95

586 Table 1 Precipitation comparison of two products

587

588

589 Table 2 Evaluation indices of simulated flood events with post-processed WRF QPF

Rain type	statistical index	201101010	20120101	20130101
	Nash-Sutcliffe coefficient/C	0.65	0.48	0.65
	Correlation coefficient/R	0.88	0.73	0.83
WRF/24h	Process relative error/P	0.35	0.57	0.19
	Peak flow relative error/E	0.14	0.18	0.25
	The coefficient of water balance/W	1.44	1.35	1.38

	Nash-Sutcliffe coefficient/C	0.75	0.58	0.75
	Correlation coefficient/R	0.93	0.82	0.85
WRF/24h after	Process relative error/P	0.23	0.35	0.11
icviscu	Peak flow relative error/E	0.08	0.12	0.16
	The coefficient of water balance/W	1.15	1.08	1.12
	Nash-Sutcliffe coefficient/C	0.58	0.63	0.5
	Correlation coefficient/R	0.78	0.75	0.8
WRF/48h	Process relative error/P	0.52	0.48	0.34
	Peak flow relative error/E	0.41	0.12	0.24
	The coefficient of water balance/W	1.52	1.43	1.51
	Nash-Sutcliffe coefficient/C	0.64	0.75	0.62
	Correlation coefficient/R	0.82	0.84	0.86
WRF/48h after revised	Process relative error/P	0.45	0.26	0.22
	Peak flow relative error/E	0.34	0.08	0.13
	The coefficient of water balance/W	1.22	1.32	1.24
	Nash-Sutcliffe coefficient/C	0.45	0.66	0.44
WRF/72h	Correlation coefficient/R	0.68	0.36	0.75
	Process relative error/P	0.64	0.62	1.29

	Peak flow relative error/E	0.31	0.35	0.45
	The coefficient of water balance/W	1.67	1.54	1.66
	Nash-Sutcliffe coefficient/C	0.52	0.75	0.55
	Correlation coefficient/R	0.75	0.45	0.82
WRF/72h after	Process relative error/P	0.53	0.52	0.98
levised	Peak flow relative error/E	0.11	0.22	0.23
	The coefficient of water balance/W	1.15	1.14	1.25

591 Table 3 Evaluation indices of simulated flood event with different model parameters

parameter type	parameter type statistical index		20120101	20130101
	Nash-Sutcliffe coefficient/C	0.75	0.58	0.75
	Correlation coefficient/R	0.93	0.82	0.85
Coupling model 24h/originally optimized model	Process relative error/P	0.23	0.35	0.11
parameters	Peak flow relative error/E	0.08	0.12	0.16
	The coefficient of water balance/W	1.15	1.08	1.12
	Nash-Sutcliffe coefficient/C	0.78	0.74	0.87
Coupling	Correlation coefficient/R	0.95	0.86	0.87
model24h	Process relative error/P	0.19	0.28	0.09
/re-optimized	Peak flow relative error/E	0.06	0.08	0.12
model parameters	The coefficient of water balance/W	1.03	0.95	1.02

	Nash-Sutcliffe coefficient/C	0.64	0.75	0.62
Coupling model	Correlation coefficient/R	0.82	0.84	0.86
48h/originally	Process relative error/P	0.45	0.26	0.22
optimized model	Peak flow relative error/E	0.34	0.08	0.13
parameters	The coefficient of water balance/W	1.22	1.32	1.24
	Nash-Sutcliffe coefficient/C	0.72	0.75	0.68
	Correlation coefficient/R	0.86	0.87	0.89
Coupling model	Process relative error/P	0.32	0.22	0.18
4011/Ie-optimized	Peak flow relative error/E	0.21	0.06	0.09
model parameters	The coefficient of water balance/W	1.05	1.12	1.06
	Nash-Sutcliffe coefficient/C	0.52	0.75	0.55
Coupling model	Correlation coefficient/R	0.75	0.45	0.82
72h/originally	Process relative error/P	0.53	0.52	0.98
optimized model	Peak flow relative error/E	0.11	0.22	0.23
parameters	The coefficient of water balance/W	1.15	1.14	1.25
	Nash-Sutcliffe coefficient/C	0.62	0.72	0.61
Counting model	Correlation coefficient/R	0.78	0.56	0.87
72h /ra optimized	Process relative error/P	0.38	0.32	0.75
model parameters	Peak flow relative error/E	0.09	0.18	0.17
model parameters				

593

Table 4 Evaluation indices of simulated flood event with different lead time

Rain type	Rain type statistical index		20130101
	Nash-Sutcliffe coefficient/C	0.82	0.95
	Correlation coefficient/R	0.89	0.92
Rain gages	Process relative error/P	0.2	0.08
	Peak flow relative error/E	0.05	0.06
	The coefficient of water balance/W	0.8	1.08
	Nash-Sutcliffe coefficient/C	0.74	0.87
WRF/24h	Correlation coefficient/R	0.86	0.87
	Process relative error/P	0.28	0.09

	Peak flow relative error/E	0.08	0.12
	The coefficient of water balance/W		1.02
	Nash-Sutcliffe coefficient/C	0.63	0.62
	Correlation coefficient/R	0.84	0.86
WRF/48h	Process relative error/P	0.48	0.22
	Peak flow relative error/E	0.12	0.13
	The coefficient of water balance/W	1.32	1.24
	Nash-Sutcliffe coefficient/C	0.56	0.61
	Correlation coefficient/R	0.56	0.87
WRF/72h	Process relative error/P	0.56	0.75
	Peak flow relative error/E	0.18	0.17
	The coefficient of water balance/W	1.54	1.66

597 **References**

598	[1]	Abbott, M.B. et al.: An Introduction to the European Hydrologic System-System
599		HydrologueEuropeen, 'SHE', a: History and Philosophy of a Physically-based,
600		Distributed Modelling System, Journal of Hydrology, 87, 45-59, 1986.
601	[2]	Abbott, M.B. et al.: An Introduction to the European Hydrologic System-System
602		HydrologueEuropeen, 'SHE', b: Structure of a Physically based, distributed
603		modeling System, Journal of Hydrology, 87, 61-77, 1986.
604	[3]	Ahlgrimm, Maike, Richard M, Forbes, Jean-Jacques Morcrette, and Roel A. J.
605	r. 1	Neggers ARM's Impact on Numerical Weather Prediction at ECMWF[J]. 2016.
606		(57):1-12.
607	[4]	Barnier, B., L. Siefridt, P. Marchesiello. Thermal forcing for a global ocean
608		circulation model using a three-year climatology of ECMWF analyses. Journal of
609		Marine Systems ,6:363-380,1995.
610	[5]	Borga, M., Borga, E.N. Anagnostou, G. Bloschl d, J.D. Creutine. Flash flood
611		forecasting, warning and risk management: the HYDRATE project[J].
612		Environmental science&policy, 2011,(14) :834-844.
613	[6]	Buizza, R., M. Miller and T. N. Palmer. Stochastic representation of model
614		uncertainties in the ECMWF Ensemble Prediction System[J]. Q. J. R. Meteorol.
615		Soc. 1999,(125): 2887-2908.
616	[7]	Burnash, R. J. C "The NWS river forecast system-catchment modeling."
617		Computer models of watershed hydrology, V. P. Singh, ed., Water Resource
618		Publications, Littleton, Colo., 311–366, 1995.
619	[8]	Chen, Yangbo. Liuxihe Model, China Science and Technology Press, September
620		2009.
621	[9]	Chen, Yangbo, Ren, Q.W., Huang, F.H., Xu, H.J., and Cluckie, I.: Liuxihe Model
622		and its modeling to river basin flood, Journal of Hydrologic
623		Engineering, 16, 33-50, 2011.
624	[10	Chen, Yangbo, Yi Dong, Pengcheng Zhang.Study on the method of flood
625		forecasting of small and medium sized catchment, proceeding of the 2013
626		meeting of the Chinese Society of Hydraulic Engineering, 1001-1008, 2013.
627	[11	Chen, Yangbo, Ji Li, Huijun Xu. Improving flood forecasting capability of
628		physically based distributed hydrological model by parameter optimization.
629	F 4 9	Hydrology & Earth System Sciences, 20,375-392, 2016.
630	[12	Chen, Y., Li, J., Wang, H., Qin, J., and Dong, L.: Large watershed flood
631		forecasting with high resolution distributed hydrological model, Hydrol. Earth
632	F 1 0	Syst. Sci. Discuss., doi:10.5194/hess-2016-489, in review, 2016.
633	[13	Danish Hydraulic Institute(DHI).MIKETI: A Modeling System for Rivers and
634	F1 4	Channels User-guide Manual[R].DHI, 2004.
635	[14	Gao, Songying, Lian qiangsun. Inspection and evaluation numerical forecast
636		product of Japan in precipitation forecasting in Dandong[J].Meteorological, 2006,
637	F 1 7	(0): /9-83.
638	[15	JGiard, D. and E. Bazile. Implementation of a New Assimilation Scheme for Soil
639		and Surface Variables in a Global NWP Model[J]. Monthly weather

640 review.2000,(128): 997-1015.

641	[16] Givati, A., Barry L., Yubao Liu, and Alon Rimmer. Using the WRF Model in an
642	Operational Stream flow Forecast System for the Jordan River[J].,
643	2012,(51):285-299. DOI: 10.1175/JAMC-D-11-082.1.
644	[17] Han, Dawei, Terence Kwong, and Simon Li. Uncertainties in real-time flood
645	forecasting with neural networks[J]. Hydrological. Process, 2007, (21): 223–
646	228.
647	[18] Hong, S.Y. and Lim, J The WRF Single-Moment 6-Class Microphysics Scheme
648	(WSM6)[J]. Journal of the Korean Meteorological Society, 2006, 42(2):129–51.
649	[19] Hong, Song-You, Ji-Woo Lee, Assessment of the WRF model in reproducing a
650	flash-flood heavy rainfall event over Korea[J]. Atmospheric Research, 2009.
651	(93):818–831.
652	[20] Hu, Xiangiun, Jianhong Tao, Fei Zheng, Na Wang, Tiejun Zhang, Shixiang, Liu,
653	and Dacheng Shang. Synopsis the parameterized scheme of physical process of
654	WRF Gansu Science and Technology 24, 73-75 2008
655	[21] Huang, Haibo, Chunyan Chen, and Wenna Zhu. Impacts of Different Cloud
656	Microphysical Processes and Horizontal Resolutions of WRF Model on
657	Precipitation on Forecast Effect METEOROLOGICAL SCIENCE AND
658	TECHNOLOGY 39 529-536 2011
659	[22] Jasper Karsten Joachim Gurtz and Herbert Lang Advanced flood forecasting in
660	Alpine watersheds by coupling meteorological observations and forecasts with a
661	distributed hydrological model[1] Journal of Hydrology 2002 (267) 40–52
662	[23] Kain J.S. The Kain-Fritsch convective parameterization: An update, Journal of
663	Applied Meteorology and Climatology 43, 170–181, 2004
664	[24] Kayyas M Chen Z Dogrul C Yoon I Ohara N Liang L Aksov H
665	Anderson M Yoshitani I Fukami K and Matsuura T (2004) "Watershed
666	Environmental Hydrology (WEHY) Model Based on Upscaled Conservation
667	Equations: Hydrologic Module." I. Hydrol. Eng. 2004. 6(450), 450-464
668	[25] Kouwen N WATELOOD: A Micro-Computer based Flood Forecasting System
669	based on Real-Time Weather Radar Canadian Water Resources Journal
670	13 62-77 1988
671	[26] Kumar Anil I Dudhia R Rotunno Dev Nivogi and U C Mohanty Analysis of
672	the 26 July 2005 heavy rain event over Mumbai India using the Weather
673	Research and Forecasting (WRF)model[1] Ouarterly Journal of the royal
674	meteorological society 200(134):1897-1910
675	[27] Li Hongyan Hanbing Liu Ximin Yuan Shukun Liu The recognition theory of
676	ANN and its application in flood forecasting[1]. Shui Li Xue Bao, 2002, 06:
677	15-19.
678	[28]Li, Yuan, G.H. Lu, Z.Y. Wu, and Jun Shi, Study of a dynamic downscaling
679	scheme for quantitative precipitation forecasting. Remote Sensing and GIS for
680	Hydrology and Water Resources IAHS Pub.
681	doi:10.5194/piahs-368-108-2015.108-113.2015
682	[29]Li, Zechun and Dehui Chen The development and application of the operational
683	ensemble prediction system at national meteorological center[J]. Journal of
684	Applied Meteorological Science, 2002, (13):1-15
685	[30] Liang X. Lettenmaier, D.P. Wood, E.F. and Burges, S.J. A simple
686	hydrologically based model of land surface water and energy fluxes for general
687	circulation models. J. Geophys. Res. 99 14415-14428 1994
688	[31] Liao, Zhenghong, Yangbo Chen, Xu Huijun, Yan Wanling Ren Oiwei, Parameter
689	Sensitivity Analysis of the Liuxihe Model Based on E-FAST Algorithm. Tropical
690	Geography, 2012, 32(6):606-612.

691	[32] Liao, Zhenghong, Yangbo Chen, Xu Huijun, He Jinxiang, Study of Liuxihe Model
692	for flood forecast of Tiantoushui Watershed, Yangtze River, 2012, 43(20): 12-16.
693	[33] Lin, Y L, Farley R D, and Orville H D. Bulk parameterization of the snow field in
694	a cloud model. Journal of Climate and Applied Meteorology, 22, 1 065-1 092,
695	1983.
696	[34], HParameter estimation in distributed hydrological catchment modelling using
697	automatic calibration with multiple objectives, Advances in Water Resources,
698	26,205-216, 2003.
699	[35] Maussion, F., D. Scherer, R. Finkelnburg, J. Richters, W. Yang, and T. Yao.WRF
700	simulation of a precipitation event over the Tibetan Plateau, China – an
701	assessment using remote sensing and ground Observations[J].Hydrol. Earth Syst.
702	Sci., 2011,(15): 1795–1817.doi:10.5194/hess-15-1795-2011.
703	[36] Molteni, F., R. Buizza, T.N. Palmer and T. Petroliagi. The ECMWF Ensemble
704	Prediction System: Methodology and validation. Meteorol. Soc., 1996, 122:
705	73-119.
706	[37] Moreno, H. A., Enrique R. Vivoni, David J. Gochis. Limits to Flood Forecasting
707	in the Colorado Front Range for Two Summer Convection Periods Using Radar
708	Nowcasting and a Distributed Hydrologic Model[J]. Journal of
709	Hydrometeorology, 2013, (14) :1075-1097.
710	[38] Niu, Junli and Zhihui Yan. The impact on the heavy rain forecast based on
711	physical process of WRF.SCIENCE & TECHNOLOGY INFORMATION.23,
712	42-45, 2007.DOI:10.3969/J.188n.1001-9960.2007.23.011.
/13	[39] Pan, Alaoduo, Aln Li, Younua Kan, and Chao Liu. Impact of Underlying Surface
/14	INFORMATION ON WRF MODELIN HEINE KIVER BASIN.PLATEA
/15	UMETEOROLOGI, 51,057-007,2012.
710	simulating heavy precipitation in Alberta[I] Atmospheric Research 135
718	136 172–192 2014
710	[41]Refsgaard I C 1997 "Parameterisation calibration and validation of distributed
720	hydrological models " I Hydrol 198 69–97
721	[42] Rutledge Steven A and Peter V. Hobbs TheMesoscale and Microscale Structure
722	and Organization of Clouds and Precipitation in Midlatitude Cyclones WII: A
723	Model for the "Seeder-Feeder" Process in Warm-Frontal Rainbands IOURANI.
724	OF THE ATMOSPHERIC SCIENCES 40 1185-1206 1983
725	[43]Shafii, M. and Smedt, F. De: Multi-objective calibration of a distributed
726	hydrological model (WetSpa) using a genetic algorithm. Hydrol. Earth Syst. Sci.
727	13. 2137–2149. 2009.
728	[44] Sherman, L. K., "Streamflow from rainfall by the unit-graph method," Eng.
729	News-Rec., 1982, 108, 501–505.
730	[45] Shim, Kyu-Cheoul, Darrell G. Fontane, M.ASCE, John W. Labadie, and
731	M.ASCE. Spatial Decision Support System for Integrated River Basin Flood
732	Control. [J].Journal of Water Resources Planning and Management,2002, 128(3):
733	190-201.DOI: 10.1061/(ASCE)0733-9496(2002)128:3(190)
734	[46] Skamarock, William C., Joseph B. Klemp, Jimy Dudhia,, David O. Gill, Dale M.
735	Barker, Wei Wang, and Jordan G. Powers. A Description of the Advanced
736	Research WRF Version 2[M].NCAR TECHNICAL NOTE,NCAR/TN-468,STR,
737	2005.
738	[47] Skamarock, William C., Joseph B., Klemp Jimy, Dudhia David, O. Gill, Dale M.
739	Barker Michael, G. Duda, Xiangyu, Huang Wei Wang, Jordan G. Powers. A

- 44 -

740	Description of the Advanced Research WRF Version 3 [M].NCAR TECHNICAL
741	NOTE, NCAR/IN-468,SIR, 2008.
742	[48] Lakenaka, Hideaki, Lakashi Y. Nakajima, Akiko Higurashi, Atsushi Ujayahi Tamia Takamura, Dashial T. Dinkan and Tamunki Nakajima Estimation
743	Higuchi, Lamio Lakamura, Rachiel T. Pinker, and Teruyuki Nakajima. Estimation
744	of solar radiation using a neural network based on radiative transfer[J]. Journal of Complexity Descendent, 116, D09215, 1 26 doi:10.1020/2000ID012227.2011
745	Geophysical Research, 116, D08215: 1-26,doi:10.1029/2009JD013337, 2011.
746 747	[49] Ingsanchall, 1. Orban flood disaster management [J].Procedia Engineering, 2012. (32):25 -37.
748	[50]Toth, E., A. Brath, A. Montanari, Comparison of short-term rainfall prediction
749	models for real-time flood forecasting [J]. Journal of Hydrology 2000, (239):
750	132–147.
751	[51] Vieux, B. E., and Vieux, J. E.: VfloTM: A Real-time Distributed Hydrologic
752	Model[A]. In:Proceedings of the 2nd Federal Interagency Hydrologic Modeling
753	Conference, July 28-August 1, Las Vegas, Nevada, Abstract and paper on
754	CD-ROM, 2002.
755	[52] Wang, Xiaojun, Hao Ma. Progress of Application of the Weather Research and
756	Forecast (WRF) Model in China. ADVANCES IN EARTH
757	SCIENCE,26,1191-1199,2011.
758	[53] Wang, Z., Batelaan, O., De Smedt, F.: A distributed model for water and energy
759	transfer between soil, plants and atmosphere (WetSpa). Journal of Physics
760	andChemistry of the Earth 21, 189-193, 1997.
761	[54]Xu, Guoqiang, Xudong Liang, Hui Yu, Liping Huang, and JishanXue.
762	Precipitation Simulation Using Different Cloud-Precipitation Schemes for a
763	Landfall Typhoon.PLATEA UMETEOROLOGY, 26, 891-900, 2007.
764	[55]Xu, Huijun, Yangbo Chen, Zeng Biqiu, He Jinxiang, Liao Zhenghong,
765	Application of SCE-UA Algorithm to Parameter Optimization of Liuxihe
766	Model, Tropical Geography, 2012.1, 32(1): 32-37.
767	[56]Xu, Huijun, Yangbo Chen, Li Zhouyang, He Jinxiang. Analysis on parameter
768	sensitivity of distributed hydrological model based on LH-OAT
769	Method[J].Yangtze River,2012, 43(7): 19-23.
770	[57]Zappa, Massimiliano, Keith J. Beven, Michael Bruen, Antonio S. Cofino, Kok,
771	Eric Martin, Pertti Nurmi, Bartlome Orfila, Emmanuel Roulin, Kai Schroter,
772	Alan Seed, Jan Szturc, Bertel Vehvilainen, Urs Germann, and Andrea Rossa.
773	Propagation of uncertainty from observing systems and NWP into hydrological
774	models: COST-731 Working Group 2[J].Atmospheric Science Letters.2010, (11):
775	83–91.
776	[58]Zhang, Guocai. Progress of Weather Research and Forecast (WRF) Model and
777	Application in the United States.Meteorological,12, 27-31,2004.
778	[59] Zhao, R. J Flood forecasting method for humid regions of China, East China

College of Hydraulic Engineering, Nanjing, China, 1977.