## **Reply to Referee #1**

It seems Gong and his/her colleagues have substantially improved the manuscript in the resubmitted version. Generally, the manuscript is written clearly and understandable, but some grammars are still need to be checked and confirmed, probably by a native speaker. I like the discussion about the human impact on evaporation, i.e. vegetation degradation and sand dunes bulldozing. The impact of vegetation degradation did not only change the vegetation cover but also modify the soil conditions. I agree with the authors that the processes are complex, and still needs to be further investigated. Summarily, these relative long-term and intensive land surface water and energy observations are important for us to understand the interaction between land surface and atmosphere and even groundwater, especially in this semiarid region where the ecosystems are vulnerable. But there is still space to improve the quality of this manuscript before publication.

Answer: Thank you for your positive comments. The manuscript has been checked and revised by the English-speaking editors of AJE (the supplement is the certificate).

## Other comments:

L24: I think it might be okay to generalize the results a little bit to "improve our understanding . . . in the fragile ecosystems of semiarid regions."

Answer: We changed the significance of this study in the abstract with the following sentences, please see lines 18-20 in the revised manuscript.

"This study could improve our understanding of the effects of land use/cover change on ET in the fragile ecosystem of semiarid regions and provide a scientific reference for the sustainable management of regional land and water resources"

## L34: Not clear info. Please rephrase.

Answer: We rephrased this sentence with the following sentences.

"In terms of physical processes, ET is affected by net radiation (Valipour et al., 2015), water vapor pressure deficit (Zhang et al., 2014), wind speed (Falamarzi et al., 2014), and soil water stress (Allen et al., 1998). Moreover, vegetation condition is also a crucial factor influencing ET (Tian et al., 2015; Wang et al., 2011; Piao et al., 2006; Mackay et al., 2007)", please see lines 29-33 in the revised manuscript.

## L 65: limiting factor for. . .

Answer: We replaced the expression of "<u>limiting factor on vegetation</u>" with "<u>limiting factor for vegetation</u>", please see line 115 in the revised manuscript.

## L66: what do you mean by "common droughts"?

Answer: In this sentence, the "common droughts" referred to droughts. In order to avoid misunderstanding, we deleted "common" in this sentence, please see line 116.

## L 62-73: Some detailed information might better go to study site section.

Answer: We used the information in this paragraph to describe the typical characteristics of Mu Us Sandy Land, including the sand dunes, biological soil crusts

(BSCs) and dry sand layer, which result in complex ET process. Therefore, we thought it might be better to leave some information about these typical land surface properties in this paragraph. And following your suggestion, we moved the following sentences describing the vegetation and water condition of Mu Us Sandy Land to the section 2.1 (Site description), please see lines 115-117 in the revised manuscript.

"Shortage of water is the critical limiting factor for vegetation in this regions, and drought-enduring vegetation are prevailed as a result of droughts (Wang et al., 2002; Wu, 2006). There are at least 117 shrub and semi-shrub species within the Mu Us Sandland (Dong and Zhang, 2001)."

L 81: in situ field. . . ?

Answer: Yes, we corrected it.

L88: "doubtful" is a strong word. You'd better change it. Answer: Following your suggestion, we revised it by replacing "doubtful" with "may induce uncertainty", please see line 65 in the revised manuscript.

L97: "... is little learned..." reads awkward. Please rephrase. Again, "field observations" Answer: We revised the sentence from "To our knowledge, there is little learned of ET under native sparse shrubland and continuous field observations under land degradation and vegetation rehabilitation conditions." to "Continuous field observations under both land degradation and vegetation rehabilitation processes have rarely been documented, especially in the semiarid shrubland", please see lines 74-76 in the revised manuscript.

Yes, we corrected the word "field".

## L102 probably change measurements to measurement

Answer: We revised the word "measurements" to "measurement", please see line 93 in the revised manuscript.

L123: is it better to say "water demand"? Answer: Yes, we revised it.

L141: "as time went on. . .". Please keep the same tense in one sentence. Answer: We revised it.

L187-189: It might be better to briefly describe how you calculated latent heat flux. Answer: We added a brief description of latent heat flux calculations by eddypro in the section 2.3 (flux data processing) with the following sentences, please see lines 179-185 in the revised manuscript.

"10 Hz 3-dimensional wind speed and water vapor concentrations that collected by EC technique were processed to half-hourly latent heat flux ( $\lambda E_{T}$ ) using Eddypro processing software (v5.2.0, LI-COR, Lincoln, NE USA). The main principle is that

 $\lambda E_{\rm T}$  can be expressed as  $\rho_a w' q'$  (where w' is the fluctuation of vertical wind

speed, q' is the fluctuation of specific humidity and  $\rho_a$  is the air density). The software also applies the quality control of data, including spike removal, tilt correction, time lag compensation, turbulent fluctuation blocking and spectral corrections."

## L198: what do you mean by "immediately"?

Answer: We used "immediately" to emphasize that we used the values before and after the data gap. In order to avoid the confusion, we deleted the word "immediately" here, please see line 201 in the revised manuscript.

## L266: How did you determine the factors in this equation?

Answer: In our study,  $\theta_r$  was calculated by  $\theta$  at different depths ( $\theta_i$ , i = 5, 10, 20, 40, 60, 80, 120, 160 cm). A schematic diagram is showed in the following.





In the layer I (0-5 cm), the soil water profile was assumed triangular, while in the layers II, III, IV, V, VI, VII and VIII, the soil water profiles were assumed trapezoidal. Therefore,  $\theta_r$  was calculated by the following equation,

$$\theta_{\rm r} = \frac{0.5 \begin{bmatrix} 5\theta_5 + (\theta_5 + \theta_{10}) * (10 - 5) + (\theta_{10} + \theta_{20}) * (20 - 10) \\ + (\theta_{20} + \theta_{40}) * (40 - 20) + (\theta_{40} + \theta_{60}) * (60 - 40) \\ + (\theta_{60} + \theta_{80}) * (80 - 60) + (\theta_{80} + \theta_{120}) * (120 - 80) \\ + (\theta_{120} + \theta_{160}) * (160 - 120) \end{bmatrix}}{160}$$

We added the above information in the text (please see lines 308-312 in the revised manuscript). Besides, we revised the Eq.13 in the text as well.

L291-294: Is this the commonly used method to calculate NDVI? If so, you do not need to mention these details. And I have no idea why you describe the NDVI\_Terra and NDVI\_Aqua. Can you clarify?

Answer: Yes, this method is commonly used to calculate NDVI. As we found that there were slight differences ( $|NDVI_{Terra} - NDVI_{Aqua}| = 0.01 \pm 0.0075$ ) between the calculated daily  $NDVI_{Terra}$  and  $NDVI_{Aqua}$ , we calculated NDVI by averaging  $NDVI_{Terra}$  and  $NDVI_{Aqua}$  in our study in order to eliminate the impacts caused by such differences.

We added the above information in the section 3.2.2 (vegetation parameter), please see lines 272-277 in the revised manuscript. In addition, we followed your suggestion to simplify the descriptions of the method to calculate NDVI by firstly deleting the superfluous sentences to describe  $NDVI_{Terra}$  and  $NDVI_{Aqua}$  (e.g., L449-451, L453-461, L472-473 and L474-475 in the marked-up manuscript). Then we rephrased the sentences to state the method to calculate NDVI by averaging  $NDVI_{Terra}$  and  $NDVI_{Aqua}$  (please see lines 267-278 in the revised manuscript).

L398-399: Since NDVI is a normalized factor and you derived the NDVI\_w based NDVI, I do not think it is meaningful to quantify the impact of NDVI on evaporation. This relationship might be changed in different cases and even with different time series datasets. You can describe this relation, but it is probably not suitable as a highlight and mention it in Abstract.

Answer: We agree with you that the relationship between NDVI and ET differs in different cases or with different time series. We also discussed this point in the manuscript (section 5.1). The main purpose for describing the relationship is to compare our result with other studies in different cases to show how strong NDVI affects ET. By our survey, the relationship between NDVI and ET were reported mostly in forests and grassland. Thus, it is meaningful to fill the gap to quantify the effect of NDVI on ET in shrubland.

In summary, we described this relationship in the main body for comparison, and following your suggestion, we deleted this description in the abstract.

L412: do you mean "compared Period I with. . . "?

Answer: Yes, we revised the sentence from "<u>compared to period I with natural land</u> <u>use/cover condition, …</u>" to "<u>compared with Period I, …</u>", please see line 426 in the revised manuscript.

L431-434: The first-order control of evaporation is a long time debate. I agree with the conclusion, but this research might be not directly related to this conclusion. I suggest the authors weaken the tone, to use "probably" or "very likely" etc.

Answer: We followed your suggestion and revised the word by replacing "mainly" in this sentence with "likely", please see line 445 in the revised manuscript.

L545: "tolerant to" is probably followed by some "vices", not survive. Please rephrase. Answer: We revised the sentence from "<u>...because shrubs are more tolerant to survive</u> in water-starved ecosystems" to "<u>... because shrubs are easier to survive in waterlimited ecosystems</u>", please see lines 545-546 in the revised manuscript.

## L 550: more water than "what"?

Answer: The missing "what" in this sentence was the word "grass". However, we deleted this sentence "<u>As potato consumes much more water than grass</u>". Because we have emphasized the fact that potato consumes more water than grass in this paragraph, please see line 842-843 in the marked-up manuscript.

#### **Reply to Referee #2**

This study assessed the relationships between evapotranspiration (ET) and change of land by analyzing the eddy-covariance measurements of actual ET together with data of a number of its potentially influencing factors including normalized vegetation index, soil water content as well as climate variables to estimate potential ET. Data are collected from a case study with different periods reflecting changes of land-use conditions, which provides further evidence to support the statistical analyses. The manuscript is well written and the knowledge promoted is a clear contribution to the understanding of how ET processes can potentially change with land-use in semiarid regions. I think this study is suitable for publication after moderate revision, with improved clarity and better flow especially for the Introduction and Methodology please see my major comments below.

Answer: Thank you for your positive comments. We revised our manuscript by carefully following your comments and suggestions.

## Major comments:

1. The Introduction launched quite well with highlighting the importance of the assessing relationship between ET to vegetation conditions in arid/semiarid regions (L28-37), followed by a comprehensive literature review explaining relevant physical mechanisms (L38-61). However, the third paragraph (L62-79) seems to be a bit disjointed as the flow of ET/vegetation stops and shifts to the case study, whereas paragraph 4 (L80-98) returns to the ET/vegetation flow and paragraph 5 again introduces the study site. I think the easiest way to improve the flow is by swapping paragraph 3 and 4 (I found this can in fact fit better with your current connecting sentences between paragraphs i.e. L60-61, L96-98). So you would have: Paragraph 1: importance of the assessing relationship between ET to vegetation conditions in arid/semiarid regions Paragraph 2: physical mechanisms on how vegetation can influence ET (finish with L60-61 which then leads to the method of assessing these impacts) Paragraph 3: method to assess the vegetation impact on ET (finish with L96-98 which then leads to the case study in a sparse shrubland) Paragraph 4: introducing the case study and how it can contribute to the above-mentioned knowledge gap. I'd also recommend combining this use some discussion from the current paragraph 5 (L99-101) to help justifying the choice of the case study. Paragraph 5: I'd recommend to leave this paragraph purely as a summary of the study (as current L102-104), and maybe elaborate a little bit with highlighting the significance of the study. I think the above structure can allow the storyline about ET/vegetation relationship to complete before introducing the study site, which provides a smoother transition and also better justification on the use of Mu US sandland as the case study.

Answer: Thank you for your constructive suggestions. According to your comments, we swapped paragraph 3 and 4 to make these paragraphs jointed, please see lines 58-92 in the revised manuscript.

In addition, we also combined the following sentences to paragraph 4 to help justify the choice of the study site.

"Coincidentally, two processes of land use/cover changes (land degradation and

vegetation rehabilitation) have occurred at the edge of the Mu Us Sandy Land, providing us a unique opportunity to study the effects of land use/cover change on ET."

Furthermore, in paragraph 5, we highlighted the significance of our study by adding the sentence "<u>Our results were expected to provide a scientific reference for the sustainable management of regional land and water resources in the context of intensive agricultural reclamation</u>", please see lines 95-97 in the revised manuscript.

2. I appreciate the comprehensiveness of Section 2 which covers the details of data collections methods and models used to analyze different data variables. However, I found that Section 2.3.2 become a bit confusing with introducing models related to a number of variables. As this section describes the methods employed for the core analyses of the study, I think the clarity can be further improved by using further subsections for individual variables. In addition, I think the methods used for data analyses should be introduced as well. Currently the statistical methods used for data analyses are mainly described in the Results section (e.g. L325-327, L380-384, L386-L389). I think it can be clearer to summarize them in the Section 2.3.2 instead (probably as an overview in the start of this section). In this way you can better justify why these analyses are conducted and how they help to answer the research questions, while purely focusing on the results and interpretation in the Results section. And then the readers can get an overall understanding on the data analyses to be conducted and knowing what to expect in the Results section. So I'd suggest the following structure for Section 2.3.2: Sub-section 1: overview – introducing the variables which are needed for analyzing the impact of ET and vegetation conditions (these will be detailed in the following sub-sections), and what analyses will be conducted with these variables (e.g. as those introduced in L325-327, L380-384 and L386-L389 etc.) Sub-section 2: estimating potential ET Sub-section 3: estimating soil water content Sub-section 4: estimating NDVI ...

Answer: Together based on your comment 2 and 3, we separated the pervious section 2 into two sections (section 2 and section 3). In the new section 2, we mainly introduced the case study information (including site information, the measurements in our study site) and data (we thought the data here included raw data and processed data). Thus, we thought it might be better to move the previous section 2.3.1 (flux data processing) into this new section 2 as subsection 2.3. We re-arranged and revised section 2 based on the following structure:

2 Case study and data

- ...2.1 Site description
- ....2.2 Field measurements
- .....2.2.1 Eddy covariance system
- .....2.2.2 Other measurements
- ...2.3 Flux data processing (lines 178-212)

While in section 3, we mainly introduced the methods to calculate the footprint and the variables that have controls on ET. Following your suggestion, we added a subsection 3.3 (statistical analysis), including the statistical methods that described in

the Results section previously (as you referred in this comment 2, e.g., L380-384, L386-L389 in previous manuscript). In addition, we also added the information about the reason why we chose linear function to simulate the correlations between ET and its three controlling factors (please see our reply to the minor comment 12).

For the sentence you mentioned in this comment 2 (e.g., L325-327 in the previous manuscript), which described the purpose to calculate energy balance closure, we prefer to leave them in the Result Section (lines 343-346). Because energy balance closure is a common concept and there is no need to describe it in the method part. In addition, if we move the sentence to the method part, the continuity of section 4.1 will be broken.

Therefore, we re-arranged and revised the section 3 according to the following structure:

- 3 Methodology
- ...3.1 Footprint model
- ...3.2 Method of analyzing controlling factors of ET
- .....3.2.1 Potential evapotranspiration
- .....3.2.2 Vegetation parameters
- .....3.2.3 Soil water stress
- ...3.3 Statistical analysis (new subsection, lines 320-329).

**3.** I think the Section 2 (Material and Methods) is a bit too long trying to cover different aspects including case study, measurements of raw data, data processing and analyzing. In my opinion a better way to organize these is to break Section 2 into two sections, for example as: Section 2. Case study and data (note:I'd use 'data' to refer to the raw measurements here rather than in the next section, where you introduce data-processing and analyzing.) 2.1 site description 2.2 measurements ... Section 3. Methodology 3.1 flux data processing 3.2 footprint model 3.3 method of analyzing controlling factors of ET (and if you agree with my last comment, the sub-sections can go below:) 3.3.1 ... 3.3.2 ... ...

Answer: Please see our reply to the comment 2.

#### **Minor comments:**

**1.** L30: 'ET' - please define acronym when it first appears in the text, and please also check if all other acronyms are properly defined.

Answer: We added the definition of ET in the text, please see line 26 in the revised manuscript.

**2.** L101: '4' - please spell out numbers less than 10 i.e. as 'four-year'. Answer: We revised it.

**3.** L111: please delete the repeated 'temperate'. Also, is there a better way to introduce the climate zone, as currently it seems like a 'noun train' ('temperate semiarid continental monsoon climate'). You can find some examples on improving 'noun train' from http://www.webwritingthatworks.com/DGuideCOG5b.htm.

Answer: We have studied the guidelines from the link you provided. Together based on

the guidelines and other scholars' studies (Yang et al., 2015; Wu and Ci, 2002). We thought the sentence might be better by changing it from "the study site is in a temperate semiarid continental temperate monsoon climate" to "This site is a semiarid area with temperate continental monsoon climate", please see line 104 in the revised manuscript.

#### **References:**

Yang, Y., Bu, C., Mu, X., Zhang, K.: Effects of differing coverage of moss - dominated soil crusts on hydrological processes and implications for disturbance in the Mu Us Sandland, China. Hydrological Processes, 29(14), 3112-3123, 2015.

Wu, B., Ci, L. J..: Landscape change and desertification development in the Mu Us Sandland, Northern China. Journal of Arid Environments, 50(3), 429-444, 2002.

**4.** L194-195: Would there be any impact on the results from this data removal, and would this be a limitation of the study? This should be briefly discussed (Maybe in the Discussion or Conclusion section?).

Answer: We thought there might be little impact of this data removal on our results due to the following reasons.

Firstly, in our study, the missing and rejected  $\lambda E_{\rm T}$  values almost occurred during nighttime (89.1% in Period I, 91.3% in Period II, 92.6% in Period III and 88.7% in Period IV), which were mainly caused by insufficient electric power supply in low air temperature environment and the low turbulence during the nighttime.

Secondly, in the nighttime, the change in  $\lambda E_{\rm T}$  is small, and ET values are close to zero. Therefore, after removal of the nighttime data, the errors of the gap-filled nighttime values based on the neighboring good data are small. Besides,  $\lambda E_{\rm T}$  values of nighttime accounted a very small proportion to the daily ET.

Thirdly, the ratios of the missing and rejected data points are not so high. For example, Falge et al. (2001) have reported that during quality control procedure of 28 flux sites, there was an average of 31% missing or rejected values of  $\lambda E_{\rm T}$  values. Wever et al. (2002) reported that there was 15% missing or rejected values of  $\lambda E_{\rm T}$  values during the quality control procedure. Mauder et al. (2006) have reported that there was an average of 20% missing or rejected values of  $\lambda E_{\rm T}$  values by 20 flux sites. Therefore, the ratio of rejected and missing half-hourly data in each period was reasonable and the dataset of  $\lambda E_{\rm T}$  after quality control procedure is reliable.

We added the above reasons in the text, please see lines 188-197 in the revised manuscript.

#### References:

Falge, E., Baldocchi, D., Olson, R., Anthoni, P., Aubinet, M., Bernhofer, C., Burba, G., Ceulemans, R., Clement, R., Dolman, H., Granier, A., Gross, P., Grunwald, T., Hollinger, D., Jensen, N. O., Katul, G., Keronen, P., Kowalski, A., Lai, C. T., Law, B. E., Meyers, T., Moncrieff, H., Moors, E., Munger, J. W., Pilegaard, K., Rannik, U., Rebmann, C., Suyker, A., Tenhunen, J., Tu, K., Verma, S., Vesala, T., Wilson, K. and Wofsy, S.: Gap filling strategies for long term energy flux data sets. Agricultural and Forest Meteorology, 107, 71-77, 2001.

Mauder, M., Liebethal, C., Göckede, M., Leps, J. P., Beyrich, F., Foken, T.. Processing and quality

control of flux data during LITFASS-2003.Boundary-Layer Meteorology, 121(1), 67-88, 2006. Wever, L. A., Flanagan, L. B., & Carlson, P. J. (2002). Seasonal and interannual variation in evapotranspiration, energy balance and surface conductance in a northern temperate grassland. Agricultural and Forest Meteorology, 112(1), 31-49.

**5.** L208-211: It would be clearer if these lines can be presented as individual formulae (i.e. in the format of L219). Also, according to L205, the 'n' in 'Rn' should be subscripted - please also check that the use of other symbols is consistent throughout the text.

Answer: After considering your comment, we thought it might be better to change these lines from several formulae to the following form: " $f = a * (R_n - G)^2 + b * (R_n - G) + c$  (Period I: a = 0.0014, b = 0.075, c = 10.69, R = 0.77; Period II: a = 0.0012, b = 0.056, c = 17.69, R = 0.67; Period III: a = 0.0014, b = 0.16, c = 13.24, R = 0.75; Period IV: a = 0.0015, b = -0.083, c = 25.87, R = 0.69)", please see lines 206-209 in the revised manuscript.

In addition, we revised the  $R_{\rm n}$  and checked the use of other symbols throughout the text.

## 6. L246: 'psychrometric constant' - what is the value of the constant?

Answer: We added the equation of psychrometric constant (Eq.6 in the revised manuscript), please see lines 252-255 for detailed information.

**7.** L248: 'U2' - where is it in Equation (5)?

Answer:  $U_2$  is used to calculate the aerodynamic resistance ( $r_a$ , Eq.7 in the manuscript). We moved the equation of calculating  $U_2$  from Eq.6 to Eq.8 for better understand, please see lines 259-262 for detailed information.

**8.** L337: 'Ds' - not defined as in Minor comment #1. Also, how are the data of Ds obtained? I couldn't seem to find it in Section 2.2.2 (other measurements).

Answer: Thank you for your kind remind. We added the information of measurement that obtained  $D_s$  in the section of "other measurements" with the following sentence. "Sunshine duration ( $D_s$ ) is measured by a sunshine recorder (CSD3; KIPP&ZONEN, Delft, the Netherlands)."

**9.** L337: 'normal' - I think 'average monthly' would be a better description here. Answer: We revised it.

**10.** L347: Figure 4 has not been introduced in the text yet, should it be mentioned somewhere between L336-337?

Answer: Yes, we added the sentences of "Four-year and long-term (1954-2014) average monthly values of  $D_s$ ,  $T_a$ ,  $R_H$ , and P are shown in Fig.4." in the section 4.2, please see lines 356-358 in the revised manuscript.

**11.** L380: 'relationships' - 'correlations' would be a more accurate description. Answer: Thank you for your suggestion. We revised it, please see line 403 in the revised manuscript.

**12.** L389-390: the r2 only investigates linear relationships - are you expecting any non-linear relationships which are not covered here and would this be a limitation? This can be briefly discussed.

Answer: We already used several common functions (e.g., exponential function, linear function, logarithmic function and quadratic function) to fit the correlations between ET and its controlling factors ( $E_{TP}$ , NDVI and  $f_s$ ). The values of determination coefficient ( $R^2$ ) are listed in the following Tab. 1.

According to the results that showed in the following Tab.1,  $R^2$  of the linear function is generally the highest. Therefore, we chose the linear function to fit the correlations between ET and its three influencing factors in our study.

We added the above information in the section 3.3 (Statistical analysis), please see lines 321-326 in the revised manuscript.

Table 1. The determination coefficient  $(R^2)$  of the correlations between ET and the three controlling factors.

	ET and E <sub>TP</sub>	ET and NDVI	ET and f <sub>s</sub>
Exponential function	0.46	0.52	0.27
Linear function	0.46	0.52	0.28
Logarithmic function	0.43	0.51	0.19
Quadratic function	0.45	0.51	0.28

**13.** L464: The term 'BSC' has already been defined in L68. Answer: We deleted the definition here.

**14.** L565: It should be worth highlighting some significance and contributions of this study towards the end of conclusion.

Answer: We highlighted the significance of our study at the end of conclusion with the following sentences:

"Furthermore, our results suggest that when we simulate the impact of land use/cover change on hydrological processes, vegetation factor might not be the unique factor to parameterize, instead, the integrated effects of land surface and vegetation conditions should be considered. Our study also provides a scientific reference to the regional sustainable management of water resources in the context of intensive agricultural reclamation.", please see lines 566-571 in the revised manuscript.

# **15.** Fig. 6: I don't think the use of different shapes is necessary given that you are using multiple panels?

Answer: We revised Fig.6.

**16.** L884 (title of Fig. 6): 'r: Pearson's correlation significance' should r be 'Pearson's correlation coefficient' instead?

Answer: Yes, we revised it in the title of Fig.6, please see line 944 in the revised manuscript.

1	Manuscript Title
2	Monitoring the variations of evapotranspiration due to-the land use/cover changes in a
3	semiarid shrubland
4	Tingting Gong, Huimin Lei, Dawen Yang, Yang Jiao, Hanbo Yang
5	State Key Laboratory of Hydroscience and Engineering, Department of Hydraulic
6	Engineering, Tsinghua University, Beijing, 100084, China
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8	
9	Abstract
10	Evapotranspiration ( $E_T$ ) is an important process in the hydrological cycle, and
11	vegetation change is a primary factor that affects $E_{\rm T}$ . In this study, an attempt is made
12	towe analyzed the annual and inter-annual characteristics of $E_{\rm T}$ using continuous
13	observation data from eddycovariance (EC) measurements over four periods-years (1 <sup>st</sup>
14	<u>1</u> July 2011 to <del>30<sup>th</sup> 30</del> June 2015) at in a study site located in a semiarid shrubland (of
15	the-Mu Us Sandy landLand,) of China. The Normalized normalized difference
16	vegetation index (NDVI) was demonstrated as as the predominant factor that influences
17	the seasonal variations in $E_{\rm T}$ . Normalization method was adopted to exclude the effects
18	of potential evapotranspiration ( $E_{TP}$ ) and soil water stress ( $f_s$ ) on $E_T$ . Vegetation
19	phenological process was validated to have a remarkable positive effect on
20	nNormalized $E_{\rm T}$ in a rate of 1.86 (the slope of normalized $E_{\rm T}$ per NDVI) increased
21	remarkably along with vegetation greening. Additionally, At the annual scale, along
22	withBoth on during the the land degradation process and vegetation rehabilitation

processes, both  $E_T$  and normalized  $E_T$  both increased as a result of due to an the 23 integrated effects of the changes in vegetation types, topography, and soil surface 24 characteristics. We discussed several possibilities that might lead to the increase. Our 25 study suggested that when we simulated the impact of land use/cover change on 26 hydrological processes, the integrated effects of land surface and vegetation conditions 27 should be considered. This study could improve our understanding of the effects of land 28 use/cover changes on  $E_{\rm T}$  in the fragile ecosystems of semiarid regions and provide a 29 scientific reference for the sustainable management of regional land and water 30 31 resources<del>Our work may promote our knowledgeimprove our understanding about the</del> characteristics of ET of the mix land use/cover condition (sparse shrubland and 32 grassland) in the fragile ecosystems of Mu Us Sandlandsemiarid regions. -33 34 Key words: evapotranspiration; normalized difference vegetation indexvegetation phenology; land use/cover change; eddy covariance; Mu Us Sandlandsemiarid 35 *climate*region 36

38	Arid and semiarid biomes cover approximately about 40% of the Earth's terrestrial
39	surface (Fernández, 2002). Previous studies have shown that more than 50% of
40	precipitation (P) is consumed by <u>evapotranspiration (E<sub>T</sub>)</u> (Yang et al., 2007; Liu et al.,
41	2002), ). Moreover, a slight change in $E_{\rm T}$ could have significant influences on water
42	<u>cycle a</u> and that the ratio of $E_{\rm T}/P$ could increase to even 90% or more in semiarid and
43	arid areasin these regions (Mo et al., 2004; Glenn et al., 2007). Therefore, a slight
44	change in $E_{\rm T}$ would have significant influences on water cycle in arid and semiarid
45	regions. In terms of physical processes, $E_{\rm T}$ is affected by not only netelimatological
46	factors (e.g., direct solar radiation, air temperature, vapor pressure deficit and wind
47	speed) but also vegetation conditions. For example, direct solar radiation and the
48	ambient air temperature provide the energy to make E <sub>T</sub> happen (Valipour et al., 2015),-
49	wWater vapor pressure deficit-is the driving force to remove water vapor from the
50	evaporating surface to the surrounding atmosphere (Zhang et al., 2014), - Besides, the
51	replacement of the saturated air with drier air depends greatly on wind speed (Falamarzi
52	et al., 2014), and soil moisture stress (Allen et al., 1998). Besides Moreover, For another
53	side, vegetation condition parameter is also a crucial factor influencing $E_{T}$ .
54	dynamics differ under different vegetation conditions (Tian et al., 2015; Wang et al.,
55	2011; Piao et al., 2006; Mackay et al., 2007). $E_{\rm T}$ is not only affected by climatic factors
56	(e.g., radiation, temperature, and relative humidity), but also affected by vegetation
57	conditions (Tian et al., 2015; Wang et al., 2011; Piao et al., 2006; Mackay et al., 2007).
58	As such, there has been an important push to understand how $E_{T}$ responds to vegetation

## 59 conditions in these regions.

60	Vegetation change mainly integrates include the phenological change (temporal)
61	and land use/cover change (spatial). The pPhenological change reflectss the response
62	of plants to climate change (vegetation greening and browning processes) (Ge et al.,
63	2015), which actively controls $E_{\rm T}$ process through internal physiologiesy such as
64	stomatal conductance (Pearcy et al., 1989)-and, as well as stomatal-the numbers and
65	sizes of stomata (Turrell, 1947). In general, transpiration is in-directly proportional to
66	stomatal conductance at the the leaf-level scale (Leuning et al., 1995). Meanwhile, aAt
67	<u>the</u> canopy scale, $E_{\rm T}$ is positively proportional to surface conductance, that which is an
68	integration of stomatal conductance and leaf area (Ding et al., 2014). Thus, as a good
69	indicator of vegetation phenological change, many studies have found that $E_{\rm T}$ was-is
70	positively related to vegetation indexes such as Normalized difference vegetation index
71	(NDVI) (Gu et al., 2007). Land use/cover change influences $E_T$ by means of modifying
72	vegetation species with with different transpiration rates, radiation transfers within
73	canopy (Martens et al., 2000; Panferov et al., 2001), topography (Lv et al., 2006),
74	albedos (Zeng et al., 2009), soil texture (Maayar and Chen, 2006), litter coverage (Wang,
75	1992), and biological soil crusts (BSCs) (Yang et al., 2015, Fu et al., 2010; Liu, 2012;
76	Eldridge and Greene, 1994). These complex processes result in no consensus about the
77	effects of land use/cover changes on $E_{\rm T}$ . For example, during the land degradation
78	process, some researchers found that warming air temperature increase was the
79	dominant-main cause to make E <sub>T</sub> increase (Zeng and Yang, 2008; Li et al., 2000;
80	Deffema and Freire, 2001). In contraryBy contrast, a decline in $E_T$ wasfound to

decrease-along with deforestation process because of less-less transpiration (Snyman, 2001; Souza and Oyama, 2011) or higher albedos (Zeng et al., 2002). Moreover, no changes indifferences of  $E_{\rm T}$  during land degradation was-were also reported (Hoshino et al., 2009). Therefore, the impacts of land use/cover changes on  $E_{\rm T}$  still deserve further investigations. As such Thus, there has been an important push to better understand how  $E_{\rm T}$  responds to vegetation conditionchange, especially to the land use/cover changes in these regions.

Three methods were usually employed to assess the impacts of land 88 89 <u>use/covervegetation change on  $E_{\rm T}$ : numerical models, paired comparative approaches</u> and the in situ filed field observations. In these methods, numerical models are widely 90 used (Twine et al., 2003; Feddema et al., 2005; Kim et al., 2005; Li et al., 2009; 91 92 Cornelissen et al., 2013; Mo et al., 2004). HHowever, model parameterization of vegetation condition is a big challenge, as the aforementioned complex underlying 93 mechanisms mentioned above may cannot be completely considered in the models. 94 95 Therefore, the simulated impact effects of land use/covervegetation change on E<sub>T</sub> is are highly dependent on the model parameterizations, which may lead to and there might 96 be largeinducer uncertainty with the resulting conclusions may be doubtful 97 (Cornelissen et al., 2013; Li et al., 2009). Paired comparative approach is -often 98 considered as the best method, nonetheless, but it is difficult to find two similar medium 99 and large-sized sites with similar meteorological conditions but different land 100 use/covervegetation conditions (Li et al., 2009; Lorup et al., 1998). Moreover, the 101 method of <u>Inin situ field observations is also a widely used method for</u>to investigate 102

long-term land-atmosphere exchange-measurementss. However, the land use/cover 103 conditions at the sites are usually generally stable, and only the responses of  $E_{\rm T}$  to 104 vegetation phenologyical changes can be studied observed. For example, such as the 105 characteristics of E<sub>T</sub> variations under in grassland (Zhang et al., 2005), mixed plantation 106 (cork oak, black locust and arborvitae) (Tong et al., 2014), vineyard (Li et al., 2015) 107 and grazed steppe (Chen et al., 2009; Vetter et al., 2012). CHowever, To our knowledge, 108 there is little learned of E<sub>T</sub>-dynamics under nativesemiarid sparse shrubland are less 109 well understood and continuous filed field observations under both land degradation and 110 111 vegetation rehabilitation condition processes have rarely been documented, especially in the semiarid shrubland, are littlenot documented. 112

The Mu Us Sandy land-Land is a semiarid shrubland ecosystem at-on the northern 113 114 margin of the Loess Plateau in China, The area covering covers an area of only 40,000 km<sup>2</sup> (Dong and Zhang, 2001). The region and is ecologically fragile (Yang et al., 2007). 115 Shortage of water is the critical limiting factor on vegetation, and drought enduring 116 117 vegetation are prevailed as a result of common droughts (Wang et al., 2002; Wu, 2006). There are at least 117 shrub and semi-shrub species have been found within the Mu Us 118 Sandland (Dong and Zhang, 2001). In such arid and semiarid an ecosystem, sand dunes 119 and biological soil crusts (BSCs) are commonly observed (Gao et al., 2014; Yang et al., 120 2015; Li and Li, 2000; Liu, 2012). Due to the exists existence of BSCs (Yang et al., 121 2015; Fu et al., 2010; Liu, 2012) and dry sand layers (Wang et al., 2006; Feng, 1994; 122 Liu et al., 2006; Yuan et al., 2007), soil evaporation have has been effectively retained, 123 therefore, the Mu Us Sandy Land Sandland holdscontains abundant groundwater (Li 124

125	and Li, 2000). During the past decades, rapid land use/cover changes has have taken
126	placeoccurred in this region due to agricultural reclamation (Wu et al., 1997; Wu and
127	Ci, 2002; Wang et al., 2009; Ostwald and Chen, 2006; Zhang et al., 2007), leading to a
128	dramatic changes in vegetation conditions. With respect to the specific question of
129	whether land use/cover changes will lead to increases in $E_{\rm T}$ or not, a continuous
130	measurement of $E_{\rm T}$ under different land use/cover conditions is <u>needed-required</u> in this
131	region. <u>Our study site is at the edge of the Mu US Sandland.</u> Coincidentally, two
132	processes of land use/cover changes (land degradation and vegetation rehabilitation)
133	hasve occurred at our study site located at this site the edge of the Mu Us Sandy Land,
134	with long term observations, which providesing us a unique opportunity to study the
135	effects of land use/cover change on $E_{\underline{T}\underline{z}}$ .
136	Three methods were usually employed to assess the impacts of land use/cover
137	change on $E_{I}$ : numerical models, paired comparative approaches and the in situ filed
138	observations. In these methods, numerical models are widely used (Twine et al., 2003;
139	Feddema et al., 2005; Kim et al., 2005; Li et al., 2009; Cornelissen et al., 2013; Mo et
140	al., 2004). However, model parameterization of vegetation condition is a big challenge
141	as the complex underlying mechanisms mentioned above cannot be completely
142	considered in the models. Therefore, the simulated impacts of land use/cover change
143	on $E_{\rm T}$ is highly dependent on the model parameterizations, and the resulting conclusions
144	may be doubtful (Cornelissen et al., 2013; Li et al., 2009). Paired comparative approach
145	is often considered as the best method, but it is difficult to find two similar medium and
146	large-sized sites with different land use/cover conditions (Li et al., 2009; Lorup et al.,
-	

147	1998). In situ observation is also a widely used method for long-term land-atmosphere
148	exchange measurements. However, the land use/cover conditions at the sites are usually
149	stable, and only the responses of $E_{\rm P}$ to vegetation phenology change can be studied. For
150	example, the characteristics of $E_{\rm T}$ under grassland (Zhang et al., 2005), mixed
151	plantation (cork oak, black locust and arborvitae) (Tong et al., 2014), vineyard (Li et
152	al., 2015) and grazed steppe (Chen et al., 2009; Vetter et al., 2012). To our knowledge,
153	there is little learned of $E_{\rm T}$ under native sparse shrubland and continuous filed
154	observations under land degradation and vegetation rehabilitation conditions are not
155	documented.
156	Hence, Our study site is at the edge of the Mu US Sandland. Coincidentally, land
157	degradation and vegetation rehabilitation has occurred at this site, which provides us a
158	unique opportunity to study the effects of land use/cover change on <i>E</i> <sub>T</sub> Based Our study
159	site is at the edge of the Mu US Sandland. Coincidentally, land degradation and
160	vegetation rehabilitation has occurred at this site, which provides us a unique
161	opportunity to study the effects of land use/cover change on E <sub>T</sub> -based on the four-4-
162	year measurements of $E_{\rm T}$ by eddy covariance techniques, this study analyzed the
163	seasonal and inter-annual variations of in $E_T$ , and discussed the possible reasons for the
164	responses of $E_{\rm T}$ to land use/cover changes. Our analysis results were expected to
165	provide a scientific reference tofor athe sustainable management of regional land and
166	water resources management in the context of intensive agricultural reclamation.
167	-On the basis of understanding the impacts of land use/cover change by human
168	activity on $E_{\text{T}}$ , more effective approaches should be implemented to maintain the water
1	

## 171 2 Materials and methods Case study and data

172 <u>2.1</u>Site description

The study was carried outconducted at the Yulin flux site (N 38 26 ; E 109 47 ; 173 1233 m), which was established in June 2011. and This site is located in a landform 174 transition zone that changes from the Mu Us Sandy LandSandy land to the north 175 Shaanxi Loess Plateau (Fig. -1). The study site is in a temperate continental semiarid 176 continental temperate monsoon climate. According to the long-term climate data (1951-177 178 2012) from a meteorological station in Yulin (Fig.\_-1), the annual precipitation varies varied from 235 mm to 685 mm, with a mean of 402 mm, and more than 50% of annual 179 precipitation is fallingfalled in the monsoon season (July-September). The mean annual 180 181 air temperature is-was 8.4 °C during over the past 61 years. The dominant soil type is sand (98% sand) (saturated soil water content: 0.43 m<sup>3</sup>m<sup>-3</sup>, field capacity: 0.16 m<sup>3</sup>m<sup>-3</sup>, 182 residual moisture content: 0.045 m<sup>3</sup>m<sup>-3</sup>). There are widely distributed fixed sand dunes 183 and semi-fixed sand dunes around the site, and the depth of the dry sand layer is 10 cm 184 (Wang et al., 2006). The mean groundwater depth of-at our study site from 1<sup>st</sup>-1 July 185 2011 to <del>30<sup>th</sup>-30</del> June 2015 was 3.5 m. 186

187

### [Figure 1 is to be inserted here]

Shortage of water is the critical limiting factor for vegetation in arid and
 semiaridthis regionsites, and drought-enduring vegetation (e.g., shrubs) areis prevailed
 as a result of droughts (Wang et al., 2002; Wu, 2006). There are at least 117 shrub and
 semi-shrub species have been found within the Mu Us Sandland (Dong and Zhang,

192 2001). The experimental The study site is mainly covered with mixed vegetation, one kind of vegetation is: the native drought-enduring shrubs with low water demands 193 194 demand such as (e.g., Artemisia ordosica and Salix psammophila) (Fig. 2a); the other kind is and the sparse grassland (that mainly distributed at the bottom of sand dunes 195 196 because of the better soil moisture condition) (Lv et al., 2006). The maximum root depth of the shrubs was approximately 160 cmThey constitute the dominant vegetation in Mu 197 Us Sandland (An et al., 2011) and are adapted well to semiarid and arid sites. According 198 to our observations around the flux tower on 14<sup>th</sup> June 2011, the maximum root depth 199 of the shrubs was approximately 160 cm. Xiao et al. (2005) reported that the growing 200 season of Artemisia ordosica and Salix psammophila spanned from late April to late 201 September. Therefore, we defined the period from 1<sup>st</sup>-1 May to 30<sup>th</sup>-30 September as 202 the vegetation growing season for data analysis in this study. On 15<sup>th</sup>-15 August 2011 203 and  $7^{\text{th}}$ -7 September 2011, we did surveys about of the vegetation coverage with by 204 randomly selected selecting 7-seven samples around the flux tower (5  $\times$  500 cm  $\times$  500 205 cm and 2  $\times$  1000 cm  $\times$  1000 cm). and We found that the vegetation coverage was 206 28.2% in August and 27.9%, respectively. in September. 207

208

#### [Figure 2 is to be inserted here]

At the end of June 2012, the land use/cover condition around the east<u>ern</u> area <u>portion</u> of <u>the</u> flux tower began to be changed by farmers (the natural vegetation <u>including the</u> leaves and branches <u>was-were</u> cut-off, and the sand dunes were bulldozed) (Fig.\_-2c), converting part of the natural vegetated land to bare <u>soilland</u>, with the planning of planting potatoes in the future. As time <u>goes-went</u> on, natural grass 214 gradually grew out gradually in the area of bare land before planting potatoes were planted. Thus Thus, our study period (1 July 2011 to 30 June 2015) can be was divided 215 into four periods according to the land use/cover conditions: (a) Period I (1<sup>st</sup>-1 July 2011 216 to 30<sup>th</sup> 30 June 2012), the period with was the natural land use/cover condition (i.e., 217 mixed sparsely distributed shrubs and grass) (Fig. 2a and Fig. 2b); (b) Period II (1st-1) 218 July 2012 to <del>30<sup>th</sup>-30</del> June 2013), the was the transitional period when with the land 219 use/cover condition starting started to change with (partial some natural vegetation 220 being cut-offremoved and sand dunes being bulldozed); (c) Period III (1<sup>st</sup> July 2013 to 221 222 30<sup>th</sup> June 2014), was the period when the land use/cover condition constituted two parts, one was: the natural vegetation zone and the other was the bare soil zone (Fig. 2c); and 223 (d) Period IV (1<sup>st</sup> July 2014 to 30<sup>th</sup> June 2015), was the period when the bare soil zone 224 225 was-gradually covered by re-growinging grass (Fig. 2d).

226

227 2.2 <u>Field Measurements measurements</u>

228 2.2.1 Eddy covariance system <u>measurements</u>

229 Net exchange of water vapor between atmosphere and canopy at this site is measured by the eddy-covariance (EC) flux measurements, which assess the fluxes of 230 land-atmosphere (such as water and energy) (Baldocchi et al., 2001). The data are 231 essential for the estimation of the water and energy balance (Franssen et al., 2010). At 232 our site, the EC system is installed at a height of 7.53 m above the ground surface, using 233 234 CSAT3 three-dimensional sonic anemometers (Campbell Scientific Inc., Logan, UT, 235 USA) for wind and temperature fluctuations measurements and a LI-7500A open-path infrared gas analyzer (LI-COR, Inc., Lincoln, NE, USA) for water vapor content 236

237 measurement.

#### 238 2.2.2 Other measurements

239 Net radiation  $(R_n)$  is measured by a net radiometer (CNR-4; KIPP&ZONEN, Delft, the Netherlands), including four radiometers measuring the incoming and reflected 240 241 short-wave radiation  $(R_{\rm S})$ , and incoming and outgoing long-wave radiation  $(R_{\rm L})$ . Sunshine dDuration  $(D_S)$  is measured by a sunshine recorder (CSD3; KIPP&ZONEN, 242 243 Delft, the Netherlands). Wind speed and direction (05103, Young Co. Traverse City, MI, USA) are measured at 10 m above the ground surface. Precipitation (P, mm) is 244 245 recorded with a tipping bucket rain gauge (TE525MM; Campbell Scientific Inc., Logan, UT, USA) installed at a height of 0.7 m above the ground surface. Air temperature ( $T_a T_a$ ) 246 and relative humidity  $(R_{\rm H}RH)$  are measured by a temperature and relative humidity 247 248 probe (HMP45C; Campbell Scientific Inc., Logan, UT, USA) at a height of 2.6 m above the ground surface. Soil water content ( $\theta$ ) is measured by Time Domain Reflectometry 249 (TDR) sensors (CS616; Campbell Scientific Inc., Logan, UT, USA), soil temperature 250 251  $(T_s)$  is measured by thermocouples (109; Campbell Scientific Inc., Logan, UT, USA), 252 and soil heat flux (G) is measured by heat flux plates (HFP01SC; Campbell Scientific Inc., Logan, UT, USA) at a depth of 0.03 m below the ground surface. These ground 253 254 variables (G,  $\theta$ ,  $T_s T_s$ ) are measured beneath the surface at two profiles (1): a plant canopy <u>patch\_profile</u> and (2) a bare soil <u>patchprofile</u>.  $\theta$  and  $T_sT_s$  are measured at 255 depths of 5, 10, 20, 40, 60, 80, 120 and 160 cm below the ground surface. Groundwater 256 257 table is measured by an automatic sensor (CS450-L; Campbell Scientific Inc., Logan, UT, USA), which is installed in a groundwater well close to the tower.\_ 258

#### 2.3 2.3 Flux data processing

261	10 Hz 3-dimensional wind speed and water vapor concentrations that collected by
262	EC technique were processed to half-hourly average-latent heat flux ( $\lambda E_T$ ) using
263	Eddypro processing software (v5.2.0, LI-COR, Lincoln, NE USA). The main principle
264	is that $\lambda E_{\rm T}$ the flux can be expressed as [] $\rho_a w' q'$ (where w' is the fluctuation
265	of vertical wind speed, $q'$ is the fluctuation of specific humidity and $\rho_a$ is the air
266	density)., The software alsowhich applied the the most recent methods for flux
267	corrections, conversions and quality control of data, including, dspike removal, tilt
268	correction, time lag compensation, turbulent fluctuation blocking and spectral
269	corrections were contained in this software. The quality control was performed on the
270	half-hourly output files, and calculated $\lambda E_{T}$ was flagged as 0 (excellent quality), 1
271	(good quality) and 2 (bad quality, removed and need to be gap-filled), respectively. The
272	basic principle of the technique is that flux is calculated ?????. The software provides
273	almost all the essential correction procedures including (1) detection and elimination of
274	spikes; (2) tilt correction; (3) sensor separation correction; (4) density fluctuation
275	correction (Webb et al., 1980). The calculated half hourly flux datasets were further
276	filtered for the remaining spikes, instrument malfunctions, and poor quality, according
277	to the following criteria (Papale et al., 2006): (1) incomplete half-hourly measurement,
278	mainly caused by power failure or instrument malfunction; (2) rainy events; and (3)
279	outliers caused by occasional spikes for unknown reasons. The ratiospercentages of
280	half-hourly $\lambda E_{\rm T}$ values rejected removed (including missing and rejected) through the
281	quality control procedure arewere 17.3% in Period I, 20.2% in Period II, 16.5% in
1	12

282	Period III, and 18.6% in Period IV., and Aalmost all missing and the rejected removed
283	$\lambda E_{\rm T}$ values occurred during the nighttime (89.1% in Period I, 91.3% in Period II, 92.6%)
284	in Period III, and 88.7% in Period IV). During the nighttime, the change in $\lambda E_{T}$ was
285	small, and $E_{T}$ values were close to zero., and besides, $\lambda$ Therefore, after removal of
286	the nighttime $\lambda E_{\rm T}$ values, the errors of the gap-filled nighttime values based on the
287	neighboring good data were small. Besides, nighttime $\lambda E_{T}$ values accounted for only
288	a small proportion of the daily $E_{\rm T}$ . Furthermore, T the ratio percentages of rejected and
289	missing data in our study are close to other scholars' results. The reported percentage
290	was summarized in a range of 15%~31% (For example, Falge et al(, 2001; Wever et
291	al., 2002; Mauder et al., 2006) have reported that during quality control procedure, there
292	was an average of 31% missing or rejected values of $\lambda E_{\rm T}$ values by 28 flux sites.
293	Wever et al. (2002) reported that there was 15% missing or rejected values of $\lambda E_{T}$
294	values during the quality control procedure. Mauder et al. (2006) have reported that
295	there was an average of 20% missing or rejected values of $\lambda E_{T}$ values by 20 flux sites.
296	In addition, $\lambda E_{T}$ values during nighttime changed steady and close to zero, coupling
297	with the fact that they accounted a very small proportion throughout whole day.
298	Therefore, the dataset of $\lambda E_{\rm T}$ after quality control procedure was isconsidered
299	reliable-to-use.
300	After quality control, missing and rejected data were gap-filled in order to create
301	continuous datasets. Three methods were applied in the gap-filling procedure: (1) Daily
302	averaged flux data were calculated by firstly gap-filled half-hourly data. Llinear
303	interpolation was used to fill gaps of less than 1-h by calculating an average of the

304	values before and after the data gap .; (2) for Larger gaps (gaps that more larger than 1-
305	h but lesssmaller than 7 -days) in flux data, the were replaced by average values using
306	mean diurnal variation (MDV) methods (Falge et al. 2001) was used. This method is
307	adopted by FLUXNET for standardized gap-filling; (3). We found that the daily $\lambda E_{T}$
308	<u>had the best correlation with daily available energy <math>(R_{n} = G)</math> rather than other</u>
309	environmental variables such as vapor pressure deficit (VPD) and NDVI. Therefore,
310	for some large gaps that largermore than 7 -days andbut lesssmaller than 15 days in
311	<u>daily <math>\lambda E_{\rm T}</math>, we fitted the relationship between daily <math>\lambda E_{\rm T}</math> and the daily available energy</u>
312	<u>flux (<math>R_n</math> – G) in each period. We chose the function <math>f</math> with the highest coefficient of</u>
313	correlation (R) in each period (Yan et al., 2013), and the function was expressed as
314	$f = a * (R_n - G)^2 + b * (R_n - G) + c$ (Period I: $a = 0.0014, b = 0.075, c = 10.69, R$
315	= 0.77; Period II: a = 0.0012, b = 0.056, c = 17.69, R = 0.67; Period III: a = 0.0014, b
316	= 0.16, $c = 13.24$ , $R = 0.75$ ; and Period IV: $a = 0.0015$ , $b = -0.083$ , $c = 25.87$ , $R = 0.69$ ).
317	Then, we used the fitted function $f$ in each period to estimate the daily $\lambda E_{\rm T}$ values
318	of large gaps. In addition, We chose the function $f_{-}$ with the highest coefficient of
319	correlation (R) in each period (Yan et al., 2013). The function $f$ of each period was
320	showed in Tab.1. Large gaps of more than 7-days but less than 15 days did mostly
321	<u>occurappeared in the winter, which accounted for a small proportion of annual <math>\lambda E_{T.\tau}</math></u>
322	[Table 1 is to be inserted here]

324 2.3 Data and methodology <u>3 Methodology</u>

325 2.3.1 Flux data processing

326 The half-hourly latent heat flux ( $\lambda$ ET) data were calculated by EddyPro software

327	(www.licor.com/eddypro) based on the raw data collected from the EC technique, and
328	it is widely used because it is comprehensive, freely available and use-friendly (Fratini
329	et al., 2014). The calculated half hourly flux datasets were filtered for spikes,
330	instrument malfunctions, and poor quality, according to the following criteria (Papale
331	et al., 2006): (1) incomplete half-hourly measurement, mainly caused by power failure
332	or instrument malfunction; (2) rainy events; and (3) outliers caused by occasional spikes
333	for unknown reasons. The ratios of data removed through this procedure are 17.3% in
334	Period I, 20.2% in Period II, 16.5% in Period III and 18.6% in Period IV.
335	Daily averaged flux data were calculated by firstly gap-filled half-hourly data.
336	Linear interpolation was used to fill gaps less than 1-h by calculating an average of the
337	values immediately before and after the data gap. Larger gaps (gaps more than 1-h but
338	less than 7-days) in flux data were replaced by average values using mean diurnal
339	variation (MDV) methods (Falge et al. 2001). This method is adopted by FLUXNET
340	for standardized gap filling. We found that the daily $\lambda E_{\rm T}$ had the best correlation with
341	daily available energy (Rn-G) rather than other environmental variables such as vapor
342	pressure deficit (VPD) and NDVI. Therefore, for some large gaps more than 7-days and
343	less than 15 days in daily $\lambda E_{\rm T}$ , we fitted the relationship between daily $\lambda E_{\rm T}$ and daily
344	available energy flux $(R_n - G)$ in each period. Then we used the fitted function $f$ to
345	estimate the daily $\lambda E_{\rm T}$ of gaps. We chose the function $f$ with the highest coefficient
346	of correlation ( <i>R</i> ) in each period (Yan et al., 2013). The function $f$ of each period was
347	Large gaps of more than 7-days did occur in the winter.
348	$\lambda = 0.0014 (Rn - G)^2 + 0.0746 (Rn - G) + 10.69$ (Period I, $R = 0.77$ ), $\lambda E_T = 0.77$

349	$0.0012(Rn-G)^2 + 0.0559(Rn-G) + 17.69$ (Period II, $R = 0.67$ ) $\lambda E_T =$
350	$0.0014(Rn-G)^2 + 0.16(Rn-G) + 13.244$ (Period III, $R = 0.75$ ) $\lambda E_T = 0.75$
351	$0.0015(Rn-G)^2 - 0.0834(Rn-G) + 25.868$ (Period IV, $R = 0.69$ ), respectively.

252	Largo	one of 1	nore ther	7_dou	e did oo	our in th	winter
552	Luigo g	ups or i	nore than	i / duy			- willter.

Period	Formula	R
Ŧ	$\lambda E_T = 0.0014  (Rn - G)^2 + 0.0746  (Rn - G) + 10.69$	<del>0.77</del>
Ħ	$\lambda E_{\tau} = 0.0012 (Rn - G)^2 + 0.0559 (Rn - G) + 17.69$	<del>0.67</del>
Ħ	$\lambda E_T = 0.0014(Rn - G)^2 + 0.16(Rn - G) + 13.244$	<del>0.75</del>
HH	$\lambda E_T = 0.0015(Rn - G)^2 - 0.0834(Rn - G) + 25.868$	<del>0.69</del>

354

## 355 2.3.2 Footprint model<u>3.1 Footprint model</u>

In order to determine the contributing source area of flux at our site, scalar flux footprint model proposed by Heisieh et al. (2000) was used. The analytic model accurately described describes the relationship between the footprint, observation height, surface roughness, and atmospheric stability. The footprint fetch  $F_f F_f$  is was calculated by as follows:

361 
$$F_f/Z_m = D/(0.105 \times k^2) Z_m^{-1} |L|^{1-Q} Z_u^Q$$
 (1)

where k is the von Karman constant (=0.40), D and Q are the similarity constants (for stable conditions: D = 0.28, Q = 0.59; for near neutral and neutral conditions: D = 0.97, Q = 1; for unstable conditions: D = 2.44, Q = 1.33), L is the Obukhov Length,  $Z_m$ is the height of wind instrument (=7.5310.0 m),  $Z_u$  is defined as (Heish-Hsieh et al., 2000);

367 
$$Z_u = Z_m (\ln(Z_m/Z_0) - 1 + Z_m/Z_0)$$
 (2)

368 where  $Z_0$  is the height of momentum roughness (0.05 m).

369

370 2.3.3 Methods of analyzing controlling factors on  $E_{\rm T}3.2$  Method of analyzing

371 <u>controlling factors on  $E_{\rm T}$  ET</u>

It is generally recognized that potential evapotranspiration ( $E_{TP}$ ), vegetation condition and soil water content are the three main factors controlling-that control  $E_{T}$ (Lettenmaier and Famiglietti, 2006; Chen et al., 2014). In order to decouple the effect of vegetation change from the integrated effects of these three factors on  $E_{T}$ , we used a simple equation which is-was similar with-to the FAO single crop coefficient method (Irrigation and Drainage Paper No. 56 (FAO-56)), and This equation is can be expressed as follows<sub>7</sub>:

379 
$$E_{\rm T} = E_{\rm TP} \times f_{\nu}$$
 (vegetation)  $\times f_s$  (soil water) (3)

where  $f_{\nu}$  (vegetation) represents the effect of vegetation change on  $E_{T_{\tau}}$  and  $f_s$  (soil water) represents the effect of soil water content on  $E_T$ .

382 <u>Moreover, By transforming the Eq.3,</u>  $f_{\nu}$  (vegetation) can be <u>regarded as the</u> 383 <u>normalized  $E_{\rm T}$ , which eliminates the effects of atmospheric and soil water content on</u> 384  $E_{\rm T}$  and can be expressed asby rearranging Eq. 3,

385 
$$f_v(\text{vegetation}) = E_T / [E_{TP} \times f_s(\text{soil water})]$$
 (4)

386 where  $f_{\psi}$  (vegetation) can also be regarded as the normalized  $E_{T}$  which eliminates the 387 effects of atmospheric and soil water content.

388 <u>3.2.1 Potential evapotranspiration</u>

389  $E_{\text{TP}} \text{ (mm day}^{-1)}$  was estimated by the following equation (Maidment, 1992) which 390 is a modification of the Penman equation,

391 
$$E_{\rm TP} = \frac{\Delta}{\Delta + \gamma} (R_{\rm n} - G) + \frac{\rho_a c_p / r_a}{\Delta + \gamma} \frac{\rm VPD}{\lambda}$$
(5)

where the units of  $R_n$  and G are mm d<sup>-1</sup>;  $\rho_a$  is the air density (= 3.486  $\frac{P_a}{275 + T_a T}$ , kg m<sup>-</sup> 392 <sup>3</sup>, where  $P_a P$  is the atmospheric pressure in kPa and  $T_a T$ \_\_is air temperature in degrees 393 Celsius);  $c_p$  is the specific heat of moist air (=1.013 kJ kg<sup>-1</sup>  $\mathbb{C}^{-1}$ );  $\Delta$  is the slope of 394 the saturation vapor-pressure-temperature curve (kPa  $\mathbb{C}^{-1}$ );  $\gamma$  is the psychrometric 395 constant (kPa  $\mathbb{C}^{-1}$ ); VPD is the difference <u>of between</u> the mean saturation vapor 396 pressure ( $e_s$ , kPa) and actual vapor pressure ( $e_a$ , kPa); and  $\lambda$  is the latent heat of 397 vaporization of water (=2.51 MJ kg<sup>-1</sup>);  $\gamma$  is the psychrometric constant (kPa  $\mathbb{C}^{-1}$ ), 398 which is calculated by the following equation,; 399  $\gamma = \frac{c_p P_a}{s \lambda}$ 400 (6)where  $\varepsilon_{\rm is}$  is the ratio of the molecular weight of water vapor to that for dry air 401 <u>(=0.622).</u> 402  $-U_2$  - is the daily wind speed at a height of 2.0 m (m s<sup>-1</sup>), which was simulated by the 403 wind speed at the height of 10.0 m (m s<sup>-1</sup>), 404

405 
$$U_{z} = U_{10} \frac{4.87}{\ln(67.8+10-5.42)}$$
(6)  
406  $r_{a}$  is the aerodynamic resistance, which was can be calculated as follows (Penman,  
407 1948; 1963);:  
408  $r_{a} = \frac{4.72[ln(\frac{Z_{h}}{Z_{a0}})][ln(\frac{Z_{h}}{Z_{0a0}})]}{1+0.536U_{2}}$ 
(6)  
409 (7)

410 where  $Z_h$  is the height at which meteorological variables are measured (2 m), and

411  $Z_{a00}$  is the aerodynamic roughness of surface (0.00137 m) (Penman, <u>1948;</u> 1963);  $U_2$ 

412 is the daily wind speed at a height of 2.0 m (m s<sup>-1</sup>), which and it was simulated calculated

413 by the wind speed at the height of  $10.0 \text{ m} (U_{10}, \text{m s}^{-1})$ ,

$$U_2 = U_{10} \frac{4.87}{\ln(67.8*10-5.42)}$$
(8)

415 The effects of soil water content on  $E_{\rm T}$  can be described in three stages (Idso et al., 416 1974), stage 1: the soil water is enough to satisfy the potential evaporation rate ( $f_{\rm g}$ =1); 417 stage 2: the soil is drying and water availability limits  $E_{\rm T}$  ( $0 < f_{\rm g} < 1$ ); and stage 3: the soil 418 is dry and evaporation can be considered negligible ( $f_{\rm g}$ =0). We used daily soil water 419 content of the root depth ( $\theta_{\rm F}$ ) to estimate  $f_{\rm g}$  by the following expression (Hu et al., 420  $\frac{2006}{5}$ ,

$$421 \qquad f_{s} = \begin{pmatrix} =1 & \theta_{r} > \theta_{k} \\ =0 & \theta_{r} < \theta_{w} \\ = \frac{\theta_{r} - \theta_{w}}{\theta_{k} - \theta_{w}} & \theta_{w} \le \theta_{r} \le \theta_{k} \end{pmatrix} \qquad ()$$

422 where  $\theta_{w}$  is the wilting value,  $\theta_{k}$  is the stable field capacity which is considered to 423 be equivalent to 60% of the field capacity (Lei et al., 1988; Wang et al., 2008).  $\theta_{r}$  (m<sup>3</sup> 424 m<sup>-3</sup>) is the mean soil water content from surface to the depth of 160 cm (root zone) and 425 was calculated by measured soil water contents at different depths,

426 
$$\theta_{\rm F} = \frac{0.5 \left[ 10^{\theta} + 15^{\theta} + 15^{\theta} + 30^{\theta} + 30^{\theta} + 20^{\theta} + 40^{\theta} + \frac{10^{\theta}}{60^{\theta}} + \frac{10^{\theta}}{60^{\theta}} + \frac{10^{\theta}}{60^{\theta}} + \frac{10^{\theta}}{10^{\theta}} + \frac{10^{\theta}}{10^{\theta}}$$

427

 $\Theta$ 

428 Site-averaged soil water content of each depth ( $\theta_i$ ; i = 5, 10, 20, 40, 60, 80, 120, and 429 160 cm) was calculated by taking a weighted average of the measured values in the 430 canopy and bare surface patches,

431  $\theta_{i} = M \times \theta_{i,e} + (1 - M) \times \theta_{i,b}$  () 432 where  $\theta_{i,e}$  and  $\theta_{i,b}$  refer to the measured soil water content of canopy patch and bare 433 soil patch at the depth of *i* cm, respectively; M is the monthly vegetation coverage of 434 undisturbed zone, and it was calculated by monthly Normalized Difference Vegetation 435 Index (NDVI) values (Gutman and Ignatov, 1998), 436  $M = (NDVI - NDVI_{min})/(NDVI_{max} - NDVI_{min})$  ()

where NDVI<sub>max</sub> is the maximum value (0.8 in this study); NDVI<sub>min</sub> is the minimum
value (0.05 in this study) (Gutman and Ignatov, 1998). The calculated monthly M (27.6%
and 24.2%) was consistent with the measured vegetation coverage in August 2011
(28.2%) and September 2011 (27.9%) at our study site.

- 441
- 442 <u>3.2.2 Vegetation parameters</u>

In this study, vegetation phenology <u>is-was</u> represented by Moderate Resolution
Imaging Spectroradiometer (MODIS)-NDVI data when the land use/cover condition<u>s</u>

445	is were fixed. NDVI is sufficiently stable to reflect the seasonal changes of any
446	vegetation (Huete et.al, 2002). Higher NDVI usually generally represent reflect the
447	greater photosynthetic capacity (greenness) of vegetation canopy (Gu et al., 2007;
448	Tucker, 1979). The daily NDVI was calculated by daily surface reflectance (at 250 m)
449	data-from_MODIS/Terra (MOD09GQ, http://reverb.echo.nasa.gov) and MODIS/Aqua
450	(MYD09GQ, http://reverb.echo.nasa.gov) Surface Reflectance (at 250m) data within
451	the footprint and the equation was, source area were chosen to calculate NDVI.,
452	$NDVI = \frac{NIR - VIS}{NIR + VIS} $ (9)
453	The Surface Reflectance data of MODIS/Terra (MOD09GQ) and MODIS/Aqua
454	(MYD09GQ) were downloaded from reverb (http://reverb.echo.nasa.gov). MODIS
455	Reprojection Tool (MRT) (Kalvelage and Willems, 2005) was used to reject the daily
456	Surface Reflectance data to the Universal Transverse Mercator (UTM) The calculation
457	of NDVI is based on its definition,
458	$NDVI_{\overline{\text{Terra or Aqua}}} = \frac{NIR - VIS}{NIR + VIS}$
459	(12)
460	where NDVI <sub>Terra</sub> and NDVI <sub>Aqua</sub> are the NDVI values calculated from MODIS/Terra
461	and MODIS/Aqua reflectance data, respectively; NIR is the spectral response in the
462	near-infrared band (857 nm); ) and VIS is the visible red radiation band (645 nm).
463	MODIS Reprojection Tool (MRT) (Kalvelage and Willems, 2005) was used to
464	rejectconvert the daily Surface Reflectance data to the Universal Transverse Mercator
465	(UTM). In this study, NDVI was calculated by using MODIS/Terra data (MOD09GQ)
466	(_NDVI <sub>Terra_</sub> ) and MODIS/Aqua data (MYD09GQ) (_NDVI <sub>Aqua_</sub> )

1	
467	(http://reverb.echo.nasa.gov), respectively. As we found that there were slight
468	<u>differences</u> ( $ NDVI_{Terra} - NDVI_{Aqua}  = 0.01 \pm 0.0075$ ) between $NDVI_{Terra}$ and
469	NDVI <sub>Aqua</sub> , we calculated NDVI by averaging NDVI <sub>Terra</sub> and NDVI <sub>Aqua</sub> in order to
470	eliminate the impacts of such differences. In order to eliminate the poor quality data
471	valuesT, the calculated NDVI data series stackvalues needs to be were then firstly
472	filtered to remove anomalous hikes hikes and drops (Lunetta et al., 2006). Hikes and
473	drops were eliminated by removing the values that suddenly decreased or increased,
474	and the then smoothing spline method was used to produce a smoother profile. In this
475	study, daily NDVI value was averaged from NDVI <sub>Terra</sub> and NDVI <sub>Aqua</sub> .
476	Theoretically, land use/cover changes can be evaluated by comparing the land
477	use/cover maps in two different periods. However, the transient land use/cover maps
478	are were unavailable at our site. Therefore, we separated the study area within the the
479	footprint area-into two zones: we assigned the undisturbed zone without any land
480	use/cover changes was deemed _as zone A, and assigned the disturbed zone with land
481	use/cover changes was deemed as zone B. In zone A, vegetation condition changes
482	included only vegetation phenological changes; however, in zone B, there were not only
483	vegetation phenological changes but also land use/cover changes. By assumingBased
484	on the assumption that the phenological changes caused by climate in the two zones are
485	were the same, we defined an indicator $(D_{lu})$ to be the <u>as a</u> measure of land use/cover
486	change <u>s</u> :

$$487 \qquad D_{lu} = M_{\rm A} - M_{\rm B} \tag{10}$$

488 where,  $M_A$  and  $M_B$  are the <u>monthly</u> vegetation coverages of zone A and zone B,

489	respectively. The monthly vegetation coverage was calculated by monthly NDVI values
490	(Gutman and Ignatov, 1998):
491	$M = (NDVI - NDVI_{min}) / (NDVI_{max} - NDVI_{min}) $ (11)
492	where NDVI <sub>max</sub> is the maximum value (0.8 in this study) and NDVI <sub>min</sub> is the
493	minimum value (0.05 in this study) (Gutman and Ignatov, 1998). The calculated
494	monthly M values (27.6% and 24.2%) were consistent with the measured vegetation
495	coverages in August 2011 (28.2%) and September 2011 (27.9%) at our study site. The
496	calculation of vegetation coverage will be described in section 3.2.3.
497	3.2.3 Soil water stress
498	The effects of the soil water content on $E_{\rm T}$ can be described in three stages (Idso et
499	al., 1974), stage 1: the soil water is enough to satisfy the potential evaporation rate
500	( $f_s=1$ ); stage 2: the soil is drying and water availability limits $E_T$ ( $0 < f_s < 1$ ); and stage 3:
501	the soil is dry and evaporation can be considered negligible ( $f_s=0$ ). We used daily soil
502	water content inof the root depth $(\theta_r)$ to estimate $f_s$ by the following expression (Hu
503	<u>et al., 2006)<del>,</del></u> :
504	$f_{s} = \begin{cases} = 1 & \theta_{r} > \theta_{k} \\ = 0 & \theta_{r} < \theta_{w} \\ = \frac{\theta_{r} - \theta_{w}}{\theta_{k} - \theta_{w}} & \theta_{w} \le \theta_{r} \le \theta_{k} \end{cases} $ (142)
505	where $\theta_{w}$ is the wilting value, and $\theta_{k}$ is the stable field capacity which is considered
506	to be equivalent to 60% of the field capacity (Lei et al., 1988; Wang et al., 2008). $\theta_r$
507	(m <sup>3</sup> m <sup>-3</sup> ) is the mean soil water content from surface to the depth of 160 cm (root zone)
508	and was calculated by measured soil water contents at different depths ( $\theta_i$ ; $i = 5, 10, 20,$
509	40, 60, 80, 120, and 160 cm). From land surface to the depth of 5 cm, the soil water

profile was regarded assumed as a triangletriangular, while inat other depths, the soil
water profiles were treated assumed as trapezoidstrapezoidal. Therefore, the root-zone
soil moisture of root zone was calculated equation was as;
$0.5 \begin{bmatrix} 5\theta_{5} + (\theta_{5} + \theta_{10}) * (10 - 5) + (\theta_{10} + \theta_{20}) * (20 - 10) \\ + (\theta_{20} + \theta_{40}) * (40 - 20) + (\theta_{40} + \theta_{60}) * (60 - 40) \\ + (\theta_{60} + \theta_{80}) * (80 - 60) + (\theta_{80} + \theta_{120}) * (120 - 80) \\ + (\theta_{120} + \theta_{160}) * (160 - 120) \end{bmatrix} $ $(122)$
$\theta_{\rm r} = \frac{1}{160} $ (123)
where, the sSite averaged soil water content of each depth ( $\theta_i$ ; ( $i = 5, 10, 20, 40, 60, $
80, 120, and 160 cm) was calculated by taking a weighted average of the measured
values in the canopy and bare surface patches,
$\theta_{i} = M \times \theta_{i,c} + (1 - M) \times \theta_{i,b}$
<u>(134)</u>
where $\theta_{i,c}$ and $\theta_{i,b}$ refer to the measured soil water contents of canopy patch and bare
soil patch at the depth of $i$ cm, respectively; <u>M</u> M is the monthly vegetation coverage
of undisturbed zone, and it was calculated by monthly Normalized Difference
Vegetation Index (NDVI) values (Gutman and Ignatov, 1998),
$M = (NDVI - NDVI_{min}) / (NDVI_{max} - NDVI_{min}) $ (14)
where NDVI <sub>max</sub> is the maximum value (0.8 in this study); NDVI <sub>min</sub> is the minimum
value (0.05 in this study) (Gutman and Ignatov, 1998). The calculated monthly M (27.6%
and 24.2%) was consistent with the measured vegetation coverage in August 2011
(28.2%) and September 2011 (27.9%) at our study site.
3.3 Statistical analysis
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552 <u>3-4</u> Results

553 34.1 Footprint and energy balance closure

Based on the footprint model, we got the half-hourly scatter data-of footprint fetch (Eq.\_-(2)), and according to the wind rose <u>diagram</u> (Fig.\_-3a), the prevailing wind direction<u>s</u> in-<u>at</u> this site were northwest and southeast,-<u>. Therefore</u>, <u>so</u>-we chose an ellipse to enclose the scatters and simulated the footprint (Fig.\_-3b). The long axis of the simulated ellipse is 1682 m, and the short axis is 1263 m. There were was 93% halfhourly flux data within the ellipse under unstable conditions.

Additionally, We-we measured the boundary of zone B in October 2013 when the land use/cover condition in zone B had stopped to-changeing (Fig. 3b). There were 11 pixels (250 m × 250 m) in zone A, while there are \_ and 19 pixels (250 m × 250 m) in zone B, and thus, in the following part of calculating the weight-averaged NDVI (NDVI<sub>w</sub>) within the footprint fetch, we chose the weighted coefficient as  $\beta =$ 11/(11 + 19).

566

### [Figure 3 is to be inserted here]

567 EC system performance was assessed by the energy balance closure which was calculated by conducting the linear regression between available energy  $(R_n - G)$  and 568 the sum of surface fluxes ( $\lambda E_{T} + H$ ), which is also used to examine the quality of flux 569 570 data (Wilson et al., 2002). In order to validate EC measurements and examine the quality of flux data, we used daily data in period I to conduct the linear regression 571 between available energy  $(R_{\rm H} - G)$  and the sum of surface fluxes  $(\lambda E_{\rm T} + H)$ . The linear 572 regression yielded a slope of 0.87, an intercept of -1.42 W m<sup>-2</sup>, and an  $R^2$  of 0.82. These 573 574 indicators indicated suggested that the measurements at our experimental site provided reliable flux data, and that the EC measurements underestimated the sum of <u>the</u> surface
fluxes to the extent of 13%. <u>A lot ofMany</u> researchers have investigated <u>the</u> energy
imbalance (Barr et al., 2006; Wilson et al., 2002; Franssen et al., 2010), and there is a
consensus that it is difficult to examine the exact reasons <u>leading to the for the</u> imbalance.

580 34.2 Characteristics of environmental variables

A brief summary of the key environmental variables will beisis presented in this 581 section. Four-year averaged monthly sunshine duration (D<sub>s</sub>), T<sub>a</sub>, R<sub>H</sub>, P and long-term 582 (1954-2014) averaged monthly values of D<sub>s</sub>, T<sub>a</sub>, R<sub>H</sub>, and P wereare showed in Fig. 4. 583 Monthly  $D_s D_s$  was much higher than the <u>long-term normal average monthly</u> values, of 584 <u>1954-2014</u> except in July and September. The highest value of monthly  $D_s D_s$  was 585 586 observed in May (299.5 h) and the lowest was observed in February (206.6 h). The <u>Seasonal seasonal characteristics of  $T_a$  that showed a highly similar trend with that of</u> 587 the long-term average monthly values of 1954-2014 normal, and the differences were 588 less than 1  $^{\circ}C$ ,  $^{\circ}C$  except in July, January and March. The highest value of monthly 589  $T_a T_a$  was <u>observed</u> in July (22.1 <u>C</u>  $\sim$ ) and the lowest was <u>observed</u> in December (-590 591 8.1<u>°C</u>°C). The values of  $R_{\rm H}RH$  showed were almost lower than the long-term average monthly values of 1954-2014 normal, especially in March and April. The highest  $R_{\rm H}R_{\rm H}$ 592 593 was <u>observed</u> in September (65.4%) and the lowest was <u>observed</u> in March (35.1%). The seasonal distributions of P were consistent with the <u>long-term average monthly</u> 594 595 values of 1954-2014 normal, and 89.7% of P happened occurred in the growing season. The value of P was highest in July was the highest (120.5 mm) and in January was the 596

597 lowest<u>in January</u> (0.3 mm).

598

#### [Figure 4 is to be inserted here]

The inter-annual characteristics of daily\_ $T_{a_1}$   $D_{s_1}$   $R_{H_1}$   $T_{a_1}$   $R_{H_2}$   $D_{s_1}$   $P_{r_2}$ , groundwater level (GWL), and total *P* in the growing season of each period were are listed in Tab.

602

#### [Table 1 is to be inserted here]

The values of  $T_a T_a$ ,  $R_H R_H$ ,  $P_2$  and  $\theta_r$  in the growing season of Period IV were the lowest compared with-to those in other three periods. Periods I\_--III are-were all wet years, while Period IV was the <u>a</u> dry year. The values of  $\theta_r$  in Periods I\_--III were basically the same similar, however,  $\theta_r$  decreased by 0.0113 m<sup>3</sup> m<sup>-3</sup> in Period IV. The mean GWL in Period III was the shallowest.

608

609 34.3 Seasonal variations in  $E_{\rm T}$  due to climate variability

The Seasonal seasonal curve of  $E_{\rm T}$  in each year had a single peak value (Fig. 5a), 610 with the higher  $E_{\rm T}$  appearing mostly in the growing season while the lower appeared in 611 the non-growing season. The daily  $E_{\rm T}$  was in a ranged from 0.0 mm day<sup>-1</sup> to 6.8 mm 612 day<sup>-1</sup> during the four periods, the highest  $E_{\rm T}$  appeared was observed on 22<sup>th</sup> June 2013 613 after when. The highest  $E_{T}$  appeared at, which was the day after a continuousal rainfall 614 event that started from 19<sup>th</sup> June 2013 to 21<sup>th</sup> June 2013 (90.3 mm), E<sub>T</sub> rates normally 615 increase rapidly after rainfall events. The lowest  $E_{\rm T}$  appeared was on 28<sup>th</sup> November 616 2012, which was in the frozen period (late November to early March in-at our study 617 site). In-On rainy days, E<sub>TP</sub> (Fig. 5b) was much-lower due to lower net radiation and air 618 temperature.  $E_{\text{TP}}$  was in the ranged of from 0.2 mm day<sup>-1</sup> that appeared in December 619

620 2011 to 17.9 mm day<sup>-1</sup> that appeared in September 2013.

621

## [Figure 5 is to be inserted here]

622 The Seasonal seasonal NDVI curve with for natural land use/cover condition (in zone A during Periods I---IV and in zone B during Period I) represented the process of 623 natural vegetation phenology and it had one a single peak value in each year (Fig. -5c). 624 In early May, the seasonal NDVI curve began to increase accompanied by thatas and 625 the native vegetation began to entered the growing season, and reached to thea 626 maximum value (0.27±0.01) was reached in July or August. In the winter, the daily 627 628 NDVI-basically stayed at a remained relatively constant value ( $0.13\pm0.01$ ).  $f_s$  (Fig. -5d) increased rapidly in response to rainfall events of more than 5 mm a day, \_\_and also 629 decreased rapidly one or two days later after rainfall events. During From late 630 631 November to early March, there was a frozen period atim this site, and soil water content was below the wilting point. The groundwater level fluctuated changed obviously in 632 the monsoon season (July to September) and mildly in the winter (December to 633 634 February).

635 \_\_The relationships between  $E_{\rm T}$  and the three factors ( $E_{\rm TP}$ , NDVI<sub>w</sub> (NDVI<sub>w</sub> = 636 NDVI<sub>A</sub> ×  $\beta$  + NDVI<sub>B</sub> × (1 -  $\beta$ )), fs) were analyzed and were shown in Fig. 6 (a, b, c) 637 by daily data in Period I. Because in Period I, the land use/cover condition within 638 footprint was undisturbed. Data in rainy days was removed, because in rainy days,  $E_{\rm T}$ 639 was gap-filled instead of actual measured.



### [Figure 6 is to be inserted here]

641

In order to figure out the major seasonal factor that control ET at our study site,

642	significant T-test was calculated to evaluate the degree of correlation. The linear
643	correlations between ET and the three factors both passed the 95%-t-test confidence
644	<u>level.</u> The linear correlations between $E_{\rm T}$ and the three controlling factors both all
645	passed the 95% <i>t</i> -test at a 95% confidence level. The <u>The determination coefficient</u>
646	$(R^2=0.52)$ value of the correlation between $E_T$ and $NDVI_w$ (NDVI <sub>w</sub> =
647	$NDVI_A \times \beta + NDVI_B \times (1 - \beta))$ was the largest, indicating that NDVI was a
648	dominant factor that controlling highly correlated with the daily variations of in $E_{\rm T}$ . To
649	better quantify the effects of <u>the</u> phenological process on $E_{\rm T}$ , <u>the correlation between</u>
650	daily normalized $E_{\rm T}$ ( $f_{\rm v}$ ) and NDVI <sub>w</sub> in Period I were was analyzed (Fig. 7a).
651	[Figure 7 is to be inserted here]
652	<u>A Positive positive</u> linear regression was found between $f_v (f_{\psi} = E_{\psi}/(E_{\psi} \times f_{\psi}))$
653	and $NDVI_w$ (Fig7a). The slope of the linear regression was used to evaluate the
654	controlling degree between $f_{v}$ normalized $E_{T}$ and vegetation phenological process.
655	which. The positive regression stated the direct positive relationship between NDVI <sub>w</sub>
656	and normalized $E_{\rm T}$ , indicated ing that when NDVI <sub>w</sub> increases increased one unit, it will

657 <u>would</u> contribute  $f_v$  normalized  $E_T$  to increase about 1.86 units.

658

659 <u>34</u>.4 Inter-annual variations in  $E_{\rm T}$  due to land use/cover changes

During the four periods, in zone A, the NDVI values of each period were basically the samesimilar because the land use/cover condition was not changed. While in zone B, the peak values of NDVI first\_ly-declined from 0.28 to 0.15 (Period I to Period III) due to the change of the land use/cover condition changed condition changed from 664 <u>mixed vegetation</u> to bare soil, and then it The peak NDVI value then increased to 0.22 665 (Period IV) due to the grass recovery (Figure 5(c). 5c). An interesting phenomenon was 666 found accompanied by the changing process of land use/cover conditions:  $E_T$  in the 667 growing season of each period was gradually observed to be increaseding from Period 668 <u>I to III (Tab.2), while it -  $E_T$  in Period IV increased strongly greatly in Period IV even</u> 669 with less precipitation, because a mass of soil water and ground water was consumed 660 to satisfy the  $E_T$  demand (Fig. 5e).

671

#### [Table 2 is inserted to be here]

Compared-to\_with Period I-with natural land use/cover condition,  $D_{lu}D_{lu}$  values of Period II and Period III gradually increased, and while  $D_{lu}D_{lu}$  of Period IV decreased. Taking August in each period as an example, in August of Period I,  $D_{lu}$  was 0.2%, while in August of Periods II, Period III and Period \_IV,  $D_{lu}$  were 2.9%, 12.6%, and 8.6%, respectively. In order to eliminate the influence of vegetation phenological change on  $E_{\rm T}$ , we chose the growing season of each period to analyze the correlations between  $f_{\rm v}$  normalized  $E_{\rm T}$  and  $D_{\rm lu}$ .

Quantitative results of the <u>correlation</u>relationship between  $D_{lu}D_{ht}$  and normalized  $E_{T}$ -( $f_v$ ) are are shown in Fig. 7b. According to the dynamic path showed in Fig.7b, wFrom Period I to Period III, When-with the changed land surface characteristics (the natural vegetation in Zone zone B was cut offcleared (Period I-III),; the fixed and semifixed sand dunes were bulldozed, the BSCs and dry sand layers were were disappeared) (Period I-III), normalized  $E_T$  (i.e.,  $f_v$ ) increased and the increment was more evident in Period III (from 78.5 to 88.1). When the land use/cover condition of in zone B gradually changed <u>from bare soil</u> to sparse grassland due to the self-restoring capacity of nature, normalization  $E_{\rm T}$ -( $f_{\rm v}$ )-increased more significantly (from 88.1 to 111.3).

688

689 4-<u>5</u>Discussion

690 4<u>5</u>.1 Implications of the <u>impacts effects</u> of phenological change on  $E_{\rm T}$ 

The correlations between  $E_{\rm T}$  and its controlling factors infer<u>red</u> that at our experimental site, NDVI was the predominant factor that influences the seasonal variations on <u>in</u>  $E_{\rm T}$ . The <u>strong</u> positive linear relationship between  $f_{\rm v}$  normalized  $E_{\rm T}$ and NDVI <u>suggests suggested</u> that transpiration is <u>was probably</u> <u>mainly</u> controlled by the stomatal conductance and the numbers of stomata, which is are proportional to leaf area (Pearcy et al., 1989; Turrell, 1947), rather than the atmospheric water demand represented by  $E_{\rm TP}$ .—

Various studies have tested assessed the relationships correlation between 698 vegetation phenological changes and  $E_{\rm T}$ , and these results generally showed consistent 699 and positive linear relationships (Nouri et al., 2014; Rossato et al., 2005; Duchemin et 700 701 al., 2006; Glenn et al., 2008). However, with for different vegetation species, phenological changes have effects on  $E_{\rm T}$  in-to\_different degrees. Relative strong 702 regressions between NDVI and  $E_{\rm T}$  have been reported at forested sites (Loukas et al. -(, 703 2005; Nouri et al., 2014; Chong et al., 2007) and grass-covered sites (Kondoh and 704 Higuchi, 2001; Nouri et al., 2014), with have analyzed the relationships between NDVI 705 and  $E_{\rm T}$  in Greece, and relative strong regressions were found in forested sites ( $\mathbb{R}^2=0.78$ ). 706 707 Kondoh and Higuchi (2001) investigated the correlation between NDVI and  $E_{\rm T}$  in a 708 grass-covered site at the university of Tsukuba, and a very high determination

coefficients ( $\mathbb{R}^2 = 0.92$ ) higher than 0.7, was showed to revealing the strong control of 709 phenological changes on  $E_{\rm T}$ . Nouri et al. (2014) have analyzed the relationships 710 between NDVI and  $E_{T}$  in forests and grasses, and they found that determination 711 coefficient of forests ( $R^2=0.94$ ) was higher than the grassland ( $R^2=0.88$ ). Chong et al. 712 713 (2007) have found a strong relationship between NDVI and  $E_{\rm T}$  in forests and moist savanna of Africa. Thus, we speculate that, for high dense vegetated ecosystems, 714 phenological changes might may have a strong and significant control on  $E_{\rm T}$ . However, 715 in low vegetation vegetated cover conditionecosystems such as sparse shrubland in this 716 717 study, the relationship between  $E_{\rm T}$  and seasonal vegetation phenological changechange is is thus positive but relatively weak. 718

719

720 45.2 Possible reasons for the effects of land use/cover changes

During Periods I—IV, the land use/cover conditions of at\_our experimental site has undergonewent two processes, : one was the land degradation process (Periods II—III), while the other was the and vegetation rehabilitation process (Period IV). Interesting phenomenon phenomena\_was-were found that-during these two processes: (1)  $E_{\rm T}$  and normalized  $E_{\rm T}$  values were both increased, and (2) and normalized  $E_{\rm T}$  increased much faster from Period III to IV/during the vegetation rehabilitation process than that from Period I to III/during the land degradation process.

The impact of phenological change<u>s</u> on  $E_T$  demonstrated that  $E_T$  <u>willwould</u> decrease along with the leaf browning. Thus, we expected that  $E_T$  <u>will-would</u> also decrease if only leaves were cleared by human activities. However, during Periods II---III, not only 731 leaves were cleared, but also other land surface properties (all branches were cut-off, sand dunes (fixed and semi-fixed) were bulldozed, and the dry sand layers and -the 732 733 biological soil crusts (BSCs) were destroyed) were changed, making-resulting in the complex land use/cover condition-complexs. All these changed land surface properties 734 735 might contribute to the increase of  $E_{\rm T}$ . The exists of dry sand layers and BSCs were demonstrated to effectively restrained the soil evaporation rates (Wang et al., 2006; Lv 736 et al., 2006; Liu et al., 2006; Chen and Dong, 2001; Yang et al., 2015; Fu et al., 2010; 737 Liu, 2012). However, the bulldozing of sand dunes at our experimental site made the 738 739 elevation of the flat soil surface be lower than the average elevation of the undisturbed soil surface (about approximately - 1.5 m lower, Figure 2(d). 2d), which resulted 740 741 that making the groundwater depth was much shallower than before the pre-disturbance 742 depth. Thus, it is was hard for the formation of dry sand layers with shallower groundwater depthlevel. In this situation with the destroyed BSCs and the disappeared 743 dry sand layers, the sufficient groundwater supply (Li and Li, 2000) accelerated the loss 744 745 of water that stored in shallow soil through evaporation. The enhancement effect of soil 746 evaporation offsetoffset the inhibition inhibiting effect of transpiration by due to leaves 747 clearing, which made  $E_{\rm T}$  increase.

A secondary reason for the <u>enhancement increase</u> of soil evaporation was that <u>more</u> solar radiation was absorbed by soil layer<u>the soil layer absorbed more solar radiation</u> during <u>the</u> land degradation process. In Period I<u>, with natural vegetation</u>, the radiation absorbed by <u>the</u> shadowed soil was the solar radiation transmitted into the canopy of shrub<u>s</u> and grass. However, <u>with when</u> the natural vegetation <u>being was cut offcleared</u>,

753	the leaves and the branches were also removed, which made the shadowed soil exposed
754	and enhanced the radiation absorbed by the soil, thus contribute to the thereby increase
755	increasing of soil evaporation (Martens et al., 2000; Panferov et al., 2001).
756	BesidesMoreover, the removal of leaves and branches and the disappearance of sand
757	dunes both altered the land surface albedos, which . Various scholars have demonstrated
758	that changes of land surface albedos could directly alter the solar radiation absorbed by
759	the land surface (Dirmeyer and Shukla, 1994; Greene et al., 1999), subsequently leading
760	to the change in $E_{\rm T}$ .
761	<u>We found sS</u> ome inconsistent results regarding the response of $E_T$ dynamics
762	toduring and the possible reason contributed to the change of $E_{\rm T}$ were found from the
763	previous analyses that aim at studying the characteristics of $E_{\rm T}$ with during land
764	degradation process were reported in previous studies were reported. Although aA
765	portion of For instance, studies reported that $-\underline{E_{T}}$ decreased during the land degradation
766	process withbut due to different reasonsthe causes of this decrease were demonstrated
767	different. For example, Li et al. (2013) have analyzed the features of $E_{\rm T}$ during land
768	degradation process in Qinghai-Tibet Plateau, and they found that warming air
769	temperature was the main cause to enhance $E_{T}$ . Souza and Oyama (2011) and However,
770	some other scholars have opposite conclusions. For example, Snyman (2001) have
771	compared $E_{\rm T}$ of natural grassland and degraded grassland resulted from overgrazing in
772	a semi-arid are of South Africa, and he found demonstrated that $E_{\rm T}$ was smaller of the
773	degraded grasslanddecreased during the land degradation process due to less decreased
774	transpiration in semiarid regions Souza and Oyama (2011) have demonstrated that

775 desertification in a semi-arid area of Northeast Brazil contributed to the decrease of  $E_{\rm T}$ due to the loss of transpiration from natural vegetation. Lu et al. (2011) considered that 776 777 have found that  $E_{\rm T}$  was lower in disturbed grazed grassland compared to the undisturbed grassland, and the lower soil water content was the thought to be main reason 778 779 to the explanation to result infor the decrease of  $E_{\rm T}$  in the land degradation process. Mao and Cherkauer (2009) <u>also have demonstrated reported a decrease of that  $E_{\rm T}$ </u> 780 decreased when land use/cover condition was changed converted from forests to grass 781 782 cropland in the Great Lakes region. HoweverHowever, some or 783 contrarydifferent contrasting phenomenonresponse results werewasere also reported regarding the effects of land degradation on  $E_{\rm T}$  by other scholars. Furthermore, Hoshino 784 et al. (2009) have demonstrated found that there was no difference in  $E_{\rm T}$  during the land 785 786 degradation by process associated with overgrazing in a semi-arid Mongolian grassland, and they thought hypothesized that the reason for this lack of change might be the short 787 time of grazing time (2 years). Furthermore, Li et al. (2013) demonstrated that the 788 789 warming air temperature was the main cause of the enhanced  $E_{\rm T}$  during the land 790 degradation process in Qinghai-Tibet Plateau. Throughout the above researches of  $E_{\rm T}$ under-during land degradation process, we found it was harddifficult to accurately 791 describe the features of  $E_{T_{a}}$  even when the land degradation was only 792 manifestperformed by less vegetation coverage. Therefore, in-at our study site with 793 complex land surface properties (sand dunes, dry sand layers and BSCs), the impact 794 795 <u>effect</u> of land degradation on E<sub>T</sub> was much more complicated.

During the vegetation rehabilitation process (Period IV),  $f_{\rm V}$ -normalized  $E_{\rm T}$ 

797	increased significantly due to the rehabilitation of grass in zone B, even though with
798	less precipitation was observed compared with other three periods (Period I, II and III).
799	The rehabilitation of grass, rather than shrub <u>s</u> , was due to the sufficient groundwater
800	supply resulted from bulldozing the sand dunes. Previous researchers reported that
801	sparse shrubs more commonly grew at the top of sand dunes and grass grew at the
802	bottom of sand dunes, because the differences between groundwater depth-level and the
803	top of sand dunes was was larger than that the that between groundwater depth-level
804	and the bottom of the sand dunes (Lv et al., 2006; Chen and Dong, 2001). Because
805	transpiration increases with the greening of vegetation greening (this wasas
806	demonstrated in section $34.3$ , the regrowing grass will would enhance plant
807	transpiration supplied by the sufficient groundwater. More importantly, the
808	transpiration rate of grass is higher than that of shrubs because shrubs are more tolerant
809	to droughteasier to survive in water-limited conditions (Yang et al., 2014; Wang et al.,
810	2002; Wu, 2006). Therefore, in the vegetation rehabilitation process, the increasing
811	rateenhancement of transpiration rate in Period IV was much higher than that in Periods
812	IIII. Consistent conclusions of $E_{\rm T}$ increase <u>due to the enhancement of transpiration</u>
813	during vegetation rehabilitation process were reported. For example, Qiu et al. (2011)
814	Yang et al. (2014) and Sun et al (2006) have demonstrated that in the vegetation
815	rehabilitation process, $E_{\rm T}$ increased and more water was consumed and less rainfall
816	would infiltrate deeper soil layer (Qiu et al., 2011; Yang et al., 2014; Sun et al., 2006;
817	Li et al., 2009)Yang et al. (2014) and Sun et al (2006)-also considered suggest ET
818	would increase with vegetation rehabilitation due to the increase of transpiration.

819	Furthermore, Li et al. (2009) have reported that $E_{\rm T}$ increased when land use/cover
820	condition converted from shrubland to grassland. Meanwhile, the regrowing grass can
821	<u>could</u> reduce the radiation absorbed by <u>the</u> soil and hence reduce <del>the</del> soil evaporation.
822	However, the interception of radiation by the grass canopy was expected to be smaller
823	than that by the mixed shrub and grass canopy in Periods IIII because the leaf area
824	index of grass_ <u>iswas</u> smaller than the sum of leaf area and stem area index <u>es</u> of <u>the</u>
825	mix of shrubs and grass. Therefore, the reduction of soil evaporation in Period IV may
826	<u>might</u> be small compared with the increment of soil evaporation in Periods IIII.
827	We noticed that the groundwater level <u>GWL</u> decreased continuously from after
828	Period III due to the enhancement of $E_{\rm T}$ by the re-growth of grass and relative lower
829	precipitation, and the regrowing grass has a high <u>er</u> transpiration rate compared with <u>than</u>
830	that of the native mixed shrub and grass-ecosystem. Therefore, we hypothesized that if
831	the land use/cover condition of zone B continues continues to be grassland in over the
832	<u>next</u> several years, the groundwater <u>level_level</u> will decrease due to the larger
833	consumption, making the soil water condition gradually become poor <u>er</u> for the growing
834	growth of grass. Then, in this situation, the grassland is expected to degrade to
835	shrubland in zone B because shrubs are more tolerant easier to survive in water-starved
836	limited ecosystems. Furthermore, iOn the other hand, potato n the next few years, potato
837	will be planted in zone B. However, the water requirement of potato-is studied to
838	consume wasis more than 320 mm in the growing season (Qin et al., 2013; Liu et al.,
839	2010) and the <u>water</u> consumption is more than that of natural grass (Qin et al., 2013,
840	2014; Hou et al., 2010) <u>BesidesThus</u> , <u>irrigation is necessary for planting potato needs</u>

to irrigate several times during the growing season in water-limited environmentlimited
ecosystems (Fulton et al., 1970; Liu et al., 2010; Fabeiro et al., 2001). As potato
consumes much more water than grass, oOur results implied imply that the groundwater
level level may might continue to decrease faster with the growth of potato in the future,
which may lead to a more fragile ecosystem.

846

847 <u>5-6</u>Conclusion

In this study, seasonal and inter-annual features of  $E_{\rm T}$  were analyzed. The dDaily 848  $E_{\rm T}$  was in a range from 0.0 mm day<sup>-1</sup> to 6.8 mm day<sup>-1</sup> during the four periods. NDVI 849 was the predominant factor that influences the seasonal variations in  $E_{\rm T}$ , and -850 v $\forall$ egetation greening had a positive effect on  $E_{\rm T}$ . During the land degradation process 851 (Periods II---Period-III), when natural vegetation (including leaves and branches), sand 852 dunes, dry sand layers and BSCs were all bulldozed by human activities,  $E_{\rm T}$  was 853 observed to increased with at a mild rate. In During the vegetation rehabilitation process 854 855 (Period IV) with sufficient groundwaterless precipitation,  $E_{\rm T}$  also increased with at a faster rate than that in the degradation process. Our study demonstrated that, when land 856 use/cover condition changed by human activities, the underlying mechanisms that leads 857 to the <u>changes increase</u> of  $E_{\rm T}$  were complex, and vegetation types, topography and soil 858 surface characteristics <u>may allcould</u> contribute to the changes in  $E_{T}$ . Furthermore, 859 oour results suggest that when we simulate the impact of land use/cover change on 860 hydrological processes, vegetation factor might not be the unique factor to parameterize, 861 instead, the integrated effects of land surface and vegetation conditions should be 862

863	considered. Our study also provides a scientific reference to the regional sustainable
864	management of water resources in the context of intensive agricultural reclamation
865	-On the basis of understanding the impacts of destroying virginal ecology by
866	human activity on water cycle, calling for paying more attention to water cycle and
867	conserve the water environment.

868

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### 1207 Figure and table captions

Fig. 1. Location of the Loess Plateau and the-map of study site (LP: the Loess Plateau;
black triangle: flux tower; white triangle: Yulin meteorological station; ①: Tu River;
2: Yuxi River; ③: Yellow River).

1211

Fig. 2. Land use/cover conditions of <u>at</u> the study site <u>over the Loess Plateau</u>: (a) the natural land use/cover condition of shrubland (photo was taken <u>at on</u> 6<sup>th</sup> August 2011); (b) the natural land use/cover condition of grassland (photo was taken <u>at on</u> 7<sup>th</sup> September 2011); (c) the undisturbed zone (natural vegetation) and the disturbed zone (bare soil) in the land degradation process (photo was taken <u>at on</u> 26<sup>th</sup> April 2013); (d) the undisturbed zone (natural vegetation) and the disturbed zone (grassland) in-<u>during</u> the vegetation rehabilitation process (photo was taken <u>at on</u> 16<sup>th</sup> August 2014).

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1220 Fig. 3. Diagrams of wind rose and footprint: (a) wind rose of the study site by using half-hourly wind speed and wind direction data; and (b) simulated footprint by ellipse 1221 1222 (the long axis is 1682\_m, and the short axis is 1263\_m; zone A is is the source area that in which have not encountered any land use/cover changecondition did not change, 1223 while zone B is is the source area that in which have experienced land use/cover 1224 condition did change by due to human activities; the white triangle is the flux tower). 1225 1226 Fig. 4. Seasonal characteristics of four-year and long-term (1954-2014, from Yulin 1227 <u>meteorological station</u>) average monthly values of: (a) sunshine duration  $(D_S DS)$ ; (b) 1228

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<u>air</u> temperature  $(T_a)$ ; (c) relative humidity  $(R_H R_H)$ ; and (d) total precipitation (P) of four

1230 1231

periods at the study site and climatological\_\_normalmonthly\_average (<u>of 1954-2014</u>
 <u>from climate ological normal in Yulin meteorological station).</u>

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Fig. 5. Seasonal and inter-annual characteristics of daily (a) evapotranspiration ( $E_{\rm T}$ , mm); (b) potential evapotranspiration ( $E_{\rm TP}$ , mm); (c) NDVI in zone A and zone B within the the footprint; (d) precipitation (P, mm); (ed) the soil water stress of the root zone ( $f_{\rm s}$ ) and (e) the groundwater level (GWL, m) – duringfrom 1<sup>st</sup>–1 July 2011 to 30<sup>th</sup>–30 June 2015<del>.</del>

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Fig. 6. The correlations between daily evapotranspiration ( $E_T$ , mm) and its controlling factors: (a) daily potential evapotranspiration ( $E_{TP}$ , mm); (b) daily weight-averaged NDVI (NDVI<sub>w</sub>) within the footprint (NDVI<sub>w</sub>); (c) daily soil water stress of the root zone ( $f_s$ ) in Period I by excluding the data in-on rainy days (r: Pearson's correlation significance coefficient; T: Tt-test significance).

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Fig. 7. Quantitative analysis between of (a) the correlations between (a) vegetation phenological change (NDVI<sub>w</sub>) and daily normalized  $E_T (f_v = E_T / (E_{TP} \times f_s))$  in Period I (excluded the data in on rainy days and frozen days); ) and (b) the indicator of land use/cover change ( $D_{lu}$ ) and total normalized  $E_T (f_v = E_T / (E_{TP} \times f_s))$  of in the growing season in of each period.

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Table <u>121</u>. Daily air temperature ( $T_a T_{a}$ , <u>C</u>-C), relatively humidity ( $R_H R_H$ , %), total sunshine duration ( $\mathcal{D}_s \mathcal{D}_S \mathcal{D}_S$ , h), soil water content of the root zone ( $\theta_r$ , m<sup>3</sup> m<sup>-3</sup>), the groundwater level (GWL, m), and total precipitation (P, mm) in 1954-2014 and in the growing season of each period (<u>Because-because</u> there were some missing data in Period IV (from 12<sup>th</sup> September 2014 to 23<sup>th</sup> November 2014 and from 13<sup>th</sup> March 2015 to 22<sup>th</sup> April 2015), we <del>got rid of excluded</del> data in these two time ranges of Periods I---III and 1954-2014)

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ItemsVariable	1954-2014	Ι	II	III	IV
<u></u> <i>T</i> a <sup></sup> (℃)	19.8	19.6	20.4	19.9	19.3
$R_{\rm H}R_{\rm H}$ (%)	57.7	57.3	54.9	53.4	52
D <sub>S</sub> <u>Ds Ds</u> (h)	213.3	220.7	215.8	218.2	220.7
P (mm)	329.8	357.1	384.1	330.2	199.8
$\theta_r \ (\mathrm{m^3  m^{-3}})$	_	0.077	0.077	0.076	0.064
GWL (m)	_	-3.8	-3.6	-3.0	-3.5

Table 232. Typical values of total evapotranspiration ( $E_{\rm T}$ , mm), total potential evapotranspiration ( $E_{\rm TP}$ , mm), the indicator of land use/cover change ( $D_{\rm lu}$ ), the soil water stress of the root zone ( $f_{\rm s}$ ), and normalized  $E_{\rm T}$  (i.e.,  $f_{\rm v}$  (=  $E_{\rm T}/(E_{\rm TP} \times f_{\rm s})$ )) in the growing season of each period. (Because because there were some missing data in Period IV (from 12<sup>th</sup> September 2014 to 23<sup>th</sup> November 2014 and from 13<sup>th</sup> March 2015 to 22<sup>th</sup> April 2015), we removed the values of  $E_{\rm T}$ ,  $E_{\rm TP}$  and  $f_{\rm s}$  of in these two time

1267 ranges <u>in of</u> Period<u>s</u> I---III).

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	Items	$E_{\mathrm{T}}$	$E_{\mathrm{TP}}$	$D_{\mathrm{lu}} \overline{D_{\mathrm{tu}}}$	$f_s$	$f_{ m v}$
	Period <sub>s</sub>	(mm)	(mm)	(%)	(dimensionless)	(dimensionless)
	Ι	238.4	876.1	-0.2	0.62	78.1
Growing	II	236.5	870.7	4.6	0.63	79.9
season	III	292.1	956	10.4	0.59	86.3
	IV	332.2	937	6	0.37	111.9



1271 Fig.\_1














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