Author response to reviewer's comments

Reviewer #1

On behalf of myself and the co-authors, I take the opportunity to thank the anonymous reviewers for their constructive comments, questions and editions. We have responded to all questions and comments, as discussed below. Most of the given comments and suggestions by the reviewers were relevant, and accordingly we have updated the manuscript significantly. We feel the quality and readability of the paper have been improved significantly.

General Comments

Comment 1: The study is relevant to the broader community in terms of illustrating the need to consider global change (not only climate change), including land use and land management as well as changes thereof in water resources planning. The study should be able to provide needed information for the basin water managers as well. Although the methods used are not novel, a study of this nature in that area provides new results. The paper claims to assess the changes temporally and spatially. Although a spatial description is given, there is no map or link to figures displaying the results spatially thus it is not possible to determine spatial patterns in the changes and the text was difficult to follow. The only map of the catchment provided does not illustrate the catchments referred to in the text. The spatial aspect of the analysis needs to be strengthened. Linked to this are the weak conclusions drawn.

Response: The authors absolutely agree with the reviewer that the text was difficult to follow as there were no maps that could show the result clearly. We would like to thank the reviewer as this comment has helped us to improve the readability of the paper by including important results on different maps to show the spatial variability. We have included two maps (1 for streamflow and 1 for rainfall) to show the spatial variability across the basin. The rainfall and streamflow trends (increasing, decreasing and no change) for each respective station can now be easily recognized from the maps. Following the reviewers comment, we have also modified Fig. 1 to include the sub-basins, level of the stations on the map. This could help readers to easily identify where the sub-catchments and monitoring stations are located in the basin The decreasing, increasing and no change trends designated by different icons on the maps are also clearly indicating the spatial variability of the parameters. Accordingly, the discussion part on the spatial variability of rainfall and streamflow has also been improved in the manuscript.

Comment 2: The paper states the relation of the trends observed in streamflow to the hydropower dam and the land management changes. Spatially these relationships are not illustrated and there is no quantitative assessment undertaken. It would strengthen the paper to have a stronger relationship between the changes in streamflow and these changes demonstrated.

Response: The authors would like to thank the reviewer for his/her comment. The aim of this study was understanding the existing relationships between the hydrological and rainfall trends which is a pre-requisite for investigating the effect of land management dynamics on the hydrological processes and streamflow of the basin. The result confirmed that the observed streamflow variabilities were not caused by rainfall but rather by land management. Quantifying the spatiotemporal relationships of observed streamflow changes with the physical

mechanisms that have caused the variabilities are recommended for further study. Consequently, the authors consider including quantified relationship between streamflow variability and land management practices to be beyond the scope of the paper. Yet, by taking the comment into account we have improved the paper by including such information from some literature existing in the basin.

Comment 3: The reference to the monthly results in the paper when they are not actually included needs revision. If they are important/relevant enough to discussed in the paper, then they should be provided.

Response: The authors totally agree with the reviewer and summary results from Mann-kendall and Pettitt tests for monthly and seasonal rainfall have now been provided as supplementary file in Table S1 and Table S2, respectively. Results of monthly streamflow analyses from Pettitt test is also included in Fig.S3 for further references. The manuscript has been also improved accordingly.

Comment 4: Beyond this, there were several grammatical and spellings errors in the manuscript. A thorough proof read is required. Other comments are included in the attached pdf.

Response: we would like to thank the reviewer for his/her significant contribution to improve grammatical and spelling errors. Based on the comments and suggestions from the reviewer, the following changes has been made in the document

Specific comments

Page 3 Line 24: delete "the" from the sentences Response: thank you, corrected in the document

P3L26: Delete "from" from the sentences

Response: Grammar corrected

P3L28: unnecessary phrase "topography of the"

Response: Repetition avoided

P3L31-32: repeated word "climate", "himud?" and change" ranging" to ranges.

Response: Thank you for the comment and Grammar corrected in the file

P4L15: Delete, "in the" from the sentences

Response: Deleted and the sentences is rephrased

P5L14: Correct sentences to "rainfall stations with less than 10% missing data have been used"

Response: Sentences improved accordingly

P5L20: Grammar, "were" changed to are.

Response: corrected

P6L8: Improve sentences to "performed the best and were recommended for the basin"

Response: Improved according to the suggestion

P6L17: change "were" these to these were

Response: corrected

P6L25-27: Need to show on map to be able to place in context of the catchment

Response: We have improved Fig 1 to include the sub-basins and all rainfall and streamflow stations. Moreover, we have developed and included two maps that shows the distribution of the observed trends across the sub-basins. Discussion on spatial variability of streamflow based on the map has added

P6L28: remove "out" from sentences

Response: removed

P7L5: Change "where" to when

Response: the whole sentences is improved

P7L9: add "s" after tend and change "of" to in

Response: Grammar corrected, 'trend' changed to trends, of deleted

P7L14: unnecessary word "approach"

Response: deleted

P8L14: change "seen" to found

Response: changed

P8L15: use streamflow is some instances and stream flow in others

Response: Corrected to streamflow in the whole document

P9L17-19: Meaning unclear

Response: sentences rephrased to make it clear

P9L19: change "that" to if or why

Response: replaced by why

P9LL27: In the text you described the spatial variation but not shown in figures. I think maps showing this would improve the paper.

Response: Thank you for the advice and the main of rainfall (annual) and streamflow (dry season) are now presented on a map. The decreasing, increasing and no change trends are shown in the map (Fig. 3 for annual rainfall and Fig. 5 for dry season streamflow) using different icons and the spatial variability can now easily identified for readers.

P10L3: Both of which have the highest percentage of missing data

Response: Thank you for the remark and this information is now added to this sentences. Their discrepancy from the other stations could be because of unreliable data.

P10L9: Either the result should be provided or this should be omitted

Response: Summary results on monthly rainfall trends is added as supplement document (Table S1).

P10L13: again-not sure that this is necessary. Was the change point test simply not done? In the next sentences though you refer to a figures as an example of the annual and seasonal time steps.

Response: Thank you for the comment. We did not include all results because of space and based on the reviewers recommendation monthly result (Table S1) and seasonal results (Table S2) are provided as supplementary file for further references.

P10L19-20: But no monthly results are shown so how is this possible to deduce? Response: Monthly rainfall result is now included as supplementary file (Table S1)

P11L5: correct spelling "gagging"

Response: corrected

P11L13: rephrase sentences to "did not significantly changed"

Response: rephrased

P11L15: change "decreasing of" to decrease in and "in" to at

Response: all grammar mistakes corrected

P11L29: correct spelling "ungagged"

Response: corrected

P12L6 and L7: put "an" before abrupt, replace "of" by in and "in" to at

Response: all corrected

P12L11-17. Why are the monthly results not included?

Response: Results on monthly streamflow from Pettitt test included as a supplement document (Fig. S3) for further references. Spatial variability of monthly flow is discussed based on the graphs.

P12L31: replace "in" before four station and three stations by at

Response: replaced

P13L7: What is meant by this?

Response: modified to short period fluctuations of hydrology

P13L9: IHA parameters or variables? Response: Parameters included after IHA

P15L8: change "compare to compared

Response: Corrected

P15L17: remove significant

Response: removed

P15L24: change "week" to weak

Response: corrected

Reviewer # 2

General comments

Comment 1: The study presented historical trend analysis of rainfall and streamflow data in the Upper Tekeze-Atbara River Basin, Ethiopia. The statistical methods (Mann-Kendall and Pettitt tests) have been well applied in many other studies, hence the methodology is not unique. However attributing hydrological changes, which were shown to be not statistically related to rainfall in the Tekeze-Atbara basin but rather to land and water resource management changes in the catchment will be of considerable use to managers and policy makers. This manuscript will therefore represent an important scientific knowledge addition to the region in question. I have provided annotated comments in the attached supplementary file, but my general comments are as follows: Greater emphasis could be placed on discussion of spatial variability of hydrological change, some of the sub-headings suggested this would be discussed but it never really materialised in the manuscript. I have suggested to amend fig 1 accordingly and thereby develop this discussion from there. Some of the charts to need to be reorganised to provide this spatio-temporal context to the data – see annotations.

Response: We agree with the reviewer that the spatial analyses is not discussed in-depth in the paper. Therefore, we have modified Fig 1 to include the sub-basins. Moreover, we have also included two maps (Fig. 3 for rainfall, and Fig. 5 for streamflow) to show the spatial variability across the basin. The two figures allow easy detection of trends in the basin.

Charts (in Fig 2 to Fig. 3) are also modified following the given suggestions. Accordingly, the discussion part on the spatial variability of rainfall and streamflow has been improved in the manuscript (track version)

Comment 2: At some points in the manuscript there is some discussion on monthly data but this is not presented, only the aggregated seasonal/annual data. This data should be presented, perhaps as a supplementary file. I have pointed out also that other pre-statistical data is discussed at length but not given – again I think there is place for this at least as a supplementary file as the authors entrust the reader to believe some of the conclusions drawn from this pre-analysis.

Response: We have added summary results of the Mann-Kendall and Pettitt tests for monthly and seasonal rainfall as supplementary file in Table S1 and Table S2, respectively. While, results of monthly streamflow analyses from Pettitt test is given in Fig. S3. Additionally, we prepared supplementary files for TFPW of rainfall (Fig. S1) and streamflow (Fig. S2). Then it is possible to check time series data before and after removal of serial correlation. In this file, samples of all time scales (monthly, seasonal and annul) are included. Therefore, discussions in the manuscript has been also improved accordingly.

Comment 3: Whilst conclusions are drawn on why hydrological changes have occurred in the basin as a result of land use/water management changes rather than being linked to rainfall, the authors point to subsequent work to unpack the causes (which will assumingly be published somewhere at a later date), one expects that grey literature/policy documents etc could be cited in this manuscript, if available, to substantiate the arguments in the interim – at editors discretion of course.

Response: The result of this study showed that streamflow variabilities are not caused by changes of rainfall. More likely attributed to land use management. However, as given in the paper, linking streamflow changes to landuse land cover changes requires a lot of analysis of satellite imageries during last four decades, which is beyond the scope of this paper. In fact this will be our next exercise to quantify LULC changes, and assess relation with stream flow changes. However, we attempted to support our argument by referring to some literature within the basin but at smaller scale as well as from neighbouring basin.

Detailed response on the specific issues

P1L24: delete"s" from alterations, avoid comma between seasonal, annual flows Changed

P1L25-26: Rephrase sentence" a significant increasing/decreasing patterns" Rephrased to show both significant increasing and decreasing trends

P2L16-17: change "reported that an increasing of trend of rainfall" to reported that a trend of increasing rainfall

Changed

P2L21: Analysed trend of hydro-climatic variables on what parameters?

Types of parameters used for analyses are now included in the sentences, which are streamflow and rainfall

P2L29: change "did not show statistically" to did not show a statistically, not clear what the annual flow statistics refer to - changes in mean annual runoff perhaps?

Corrected, modified into mean annual runoff

P3L3: change "trend" to trends and 50 and 60 years into 50 to 60 Corrected

P3L4: 30-40 what

Corrected to 30-40 years

P3L4-5: improve sentences into "Meanwhile, record lengths less than 25 years tended to show statistically significant increasing trends"

Improved in the text to the above sentences.

P3LL6: This paragraph could benefit from a discussion on 'hydro-climatic zones' i.e. how does apparent hydrological change, if anthropogenically driven differ between humid and semi-arid regions?

Discussion on climatic zone included in the paragraph. The spatio-temporal runoff generation in semi-arid areas is strongly non uniform as runoff generation controlling factors are different from that of a humid environment.

P3L10: Remained constant in what respect - rainfall depth (volumes) or intensity, or both? Corrected into amount of rainfall

P3L10-13: suggest amend first part of sentence to: 'Despite the importance of streamflow to assure sustainable water resource utilisation and food security.....'

Sentences improved.

P3L22: change river to River

Corrected

P3L23: change "near Ras Deshen" to near the Ras Deshen

Corrected

P3L29: Put "the" before elevation

Corrected

P3L30-32: change "himud" to humid and rangaing to ranges

Grammar corrected

P4L4: change "The river flow pattern follows that of rainfall in to "The river flow pattern typically follows that of rainfall"

Changed

P4L10: at what scale and for how long have these SWC activities been taking place? Scale included in the text

P4L14-15: outlet of what?, avoid "s" after alters and re-write last part of sentence Corrected to outlet of the basin, s avoided and last sentences has been improved in the text

P4L24: change "temporal and spatial" to spatio-temporal Changed

P4L19-21: remove "s" from depends and efforts, delete scrutinizing of from the sentences Changed

P4L26-27: change "varied" to varying, "remove despite that" and add whilst before gauging station

Corrected

P4L28: change "the" before 30 years to A 30 years

P5L6-7: what was the criteria for these rainfall datasets being deemed reliable? Criteria are described in the next paragraph. The text of the manuscript has been updated accordingly

P5L10: add "an" before acceptable

Included

P5L16: apply full length names for CHIRPS and TRMM etc in the first paragraph of section 3.1.1. and use only the acronyms here Changed

P5L17: why was 25km radius selected?

A 25 km radius area average of rainfall was taken considering the satellite data resolution and to avoid the effect of topography complex on the rainfall estimation. This sentences is also updated in the text.

P6L3: change "on the" to for the

Corrected

P6L17: Change "were these" to these were

Corrected

P6L16-17: explain why longer records are better, and why shorter period enhances spatial coverage - the point of this sentence is not clear.

Longer period is to increase identification of dominant trends while shorter period allows to include more representative stations even with less recording period. More explanation is included in the text.

P6L20: why is 20 years important?

To account spatial variability and obtaining reliable result from trend test tool. Also explained in the paper.

P6L28: "change explicitly point to" to which suggests

Corrected

P6L30: perhaps use better word than scanty - what do you mean, patchy and unverified? Modified to "very limited"

P6L33: change "its" to for and "checked" to check

Corrected

P6L33-34: OK - but how did you use the other stations for flow verification - whilst you explain double mass and residual mass plots below, it's not clear from your explanation what the other gauges are used for exactly - just needs some clarification.

Comparing between upstream and downstream stations can give preliminary evidence on the reliability of data. We compared each stations with nearby stations just to get first-hand information on the quality of data. Text has also been updated

P7L4: why was two years selected as criteria for omission? why two weeks avoided

Response: More than two years data were excluded from the analysis in order to avoid unreliable time series data whilst below two years were included to enhance spatial coverage by including stations with some missing data. Similarly, more than two weeks of missing values during the rainy season were avoided as 80% of the annual rainfall is coming from this season. Clarification on this issue is also included to the text document.

P7L9: change "Man-kendall" to Mann-Kendall

Corrected

P8L16: For the reader, it will be worthwhile to see the difference before and after the TFPW analysis - as supplementary file to accompany the manuscript.

Result on TFPW is now included as supplementary file. Comparison of rainfall (Fig. S1) and streamflow (Fig. S2) after and before TFPW from all time scales (monthly, seasonal and annul) are included.

P8L19: change "trend-free-prewhitened" change to TFPW, "the" added before Pettitt Changed

P8L25: why was a user defined 5% chosen - any reason for this - does this literature commonly use this value?

It is commonly used in the literatures for hydro-climatic trend analyses. Text is improved with more explanations.

P8L33: add "the" before Indicators

Corrected

P9L2: 20 IHA parameters are not listed here - and only 8 referred to in Table 6. Perhaps add table describing all the parameters used.

Corrected to 8 parameters.

P9L21: what is meant by non-normal data?

It was to mean unreliable data, now changed to unreliable data

P9L24: Do you mean natural storage heterogeneity e.g. variable vadose zone properties - or man-made storage (dams etc)?

Man-made storage, text modified accordingly

P9L23: change "flow is very small" to discharges are very low Changed accordingly

P9L23-24: not sure I follow your point here - why will low flows cause temporal dependency? - do you mean that hydrographs are relatively homogenous between gauges during low flow conditions?

It is to mean similarity of flows between consecutive days and to make it explicit the text is now rephrased.

P9L27: not much discussion in the manuscript on the spatial variability element and this needs to be developed further - one way of achieving this will be spatial plots of increasing/decreasing trends - see my comment on fig 1.

The discussion of spatial variability has been improved by including different maps showing the distribution of sub-basins and the main results with different icons, Fig. 3, Fig. 5. To make it more sound the caption is also changed into "rainfall variability over the basin".

P10L6-7: repeated sentences "The result of rainfall analyses shows no trend" Removed

P10L6-7: Grammar mistakes

Corrected

P10L9: refer to monthly data in this section, but none given in the tables - may be worth including in supplementary file to accompany paper.

Summary of monthly result is now included as supplementary file (Table S1)

P10L17-10: why not? it would be useful to see the statistical indices you refer to. Again - maybe to add in supplementary file.

Supplementary file included (Table S2)

P10L15-19: refer to monthly data in this section, but none given in the tables - may be worth including in supplementary file to accompany paper

Supplementary file for all monthly result is now included (Table S1 and Table S2)

P10L25: meaning what exactly - are you referring to the ITCZ?

ITCZ included in bracket

P10L27-30: did the authors referred to in this paragraph use trend detection methods or alternatively infer trends?

They have applied same methodology of trend analyses and text is also modified to reflect that.

P11L8: avoid repetitive sentences

Removed

P11L14: sorry - I may have missed previously was the flow data disaggregated to seasonal data sets?

Seasonal data sets were obtained by aggregating the monthly data of each season

P11L15: change decreasing to decrease

Changed

P11L29: change "ungagged" to ungauged

Corrected throughout the document

P12L14: change have to has

Corrected

P13L1: land degradation through what mechanism(s)?

Deforestation and over cultivation on the rugged topography, document also modified

P13L13: what is a water condition? do mean hydrological condition?

Changed to hydrological condition

P13L26: due to homogenisation of the low flow and peak flow hydro graph?

Modified

P14L10: is there any literature whether peer reviewed or grey literature that can be used to cite for land-use changes in the study region?

Discussions on land use change and its effect on hydrological processes is included from previous literatures which was studied at watershed levels.

P14L27: as above - but it will be useful to infer on some causative mechanisms from any literature that is available, as quick reference - i.e you have spoken of water abstractions, but not land-use changes explicitly

Discussions on effect of land use added from few existing studies at small watershed level. To make it sound this discussion has merged with the above (P12L10).

P20L7: remove "age" from Table 1

Removed

P22L6: change "annual flow" in Table 3 to annual average flow

Corrected

P22L5: level the station names in Table 3 on fig 1.

Included in the figure

P26.Fig.1: may be worth replotting map in grayscale, with rainfall and flow gauges - added value may be to use icons that show increasing/decreasing or static trends. Include stations names and -sub-catchments referred to in manuscript. Map will need to be reproduced at greater dpi resolution. Lat/Long in Tables are given in Degrees Decimal this map should also be

All suggestions are included in the figure and the map is now significantly improved.

P27.Fig. 2: If having a combined plot then use the legend only once, consider plotting all four against one x-axis with four charts plotted one of top of each other with individual y-axis. The individual chart titles NSE, BIAS, RMAE, r, will be the y-axis label

We found this comment is also very important and now modified to have one x-axis and contains only one legend.

P28.Fig. 3: It would be useful to plot the data above each other - with all sharing the same x-axis - for context plot 1950-2020. That way if these gauges are labelled in fig 1, we can see the spatial variability.

Similarly, modified all charts under this fid to shared same x-axis for the context of 1950-2020

P29Fig. 4: again, these plots should have same x-axis scale 1970-2020 remove clutter by moving the labels 'annual', 'rainy season' etc to top of the chart column. Improved to the recommended style.

Temporal and spatial changes of rainfall and streamflow in the Upper Tekeze–Atbara River Basin, Ethiopia

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Abstract. The Upper Tekeze–Atbara river sub—basin—partbasin, part of the Nile basin, is characterized by high temporal and spatial variability of rainfall and streamflow. In spite of its importance for sustainable water use and food security, the changing patterns of streamflow and its association with climate change is not well understood—in the basin. This study aims at improving the understanding of the linkages between rainfall and streamflow trends and identifying the possible drivers of streamflow variabilities in the basin. Trend analyses and change point detections of rainfall and streamflow were analysed using Mann-Kendall and Pettitt tests, respectively, using data records for 21 rainfall and 9 streamflow stations. The nature of changes and linkages between rainfall and streamflow were carefully examined for monthly, seasonal and, annual flows, as well as Indicators of Hydrological Alterations (IHA).

The trend and change point analyses found that 20 of the tested 21 rainfall stations did not show statistically significant changes. In contrast, trend analyses on the streamflow showed a-both significant increasing and /decreasing patterns. A decreasing trend in the dry (October to February), short (March to May), main rainy seasons (June to September) and annual totals is dominant in 6 out of the 9 stations. Only one out of nine gauging stations experienced increasing flow significantly in the dry and short rainy seasons., This increasing trend is attributed to the construction of Tekeze hydropower dam above upstream thise station in 2009. Overall, streamflow trends and change point timings were found to be inconsistent among the stations. Changes in streamflow without significant change in rainfall suggests other factors than rainfall to drive the change. Most likely Weak linkages between rainfall and streamflow trends indicate that the observed changes in streamflow regimes could be due to changes in catchment characteristics of the basin. Further studies are needed to verify and quantify the hydrological changes shown in statistical tests by identifying the physical mechanisms behind those changes. The findings from this study are useful as a pre-requisite for studying the effects of catchment management dynamics on the hydrological variabilities in the basin.

40 **Keywords:** Streamflow variability, Trend analyses, Tekeze River Basin, Statistical test

1. Introduction

Recent changes in climatic conditions combined with other anthropogenic factors have increased the concern of the international community on water resources management in basins

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(Jones et al., 2015; Zhang et al., 2008). Understanding climate change and its impact on hydrological variability is important for water management-, and thus has received attention from researchers in different parts of the world (e.g. Kim et al., 2008; Ma et al., 2008; Pano et al., 2010; Tekleab et al., 2013; Wang et al., 2013; Zhan et al., 2014; Zhao et al., 2015). These studies investigate how climate change reflected in changing rainfall patterns affects the hydrological regimes of river basins.

Identifying the trends and linkages between rainfall and streamflow is fundamental to understand the influence of climate change on the hydrological variability of a basin. Many studies (e.g., IPCC, 2013; Shi et al., 2013; Tekleab et al., 2013; Tesemma et al., 2010; Zhao et al., 2015) have shown that rainfall is the primary atmospheric factor that directly affects the streamflow patterns. The impact of climate change on hydrology varies from place to place. For example, Ma et al. (2008) for the arid region of northwest China, Zhang et al. (2011) for China, Zhao et al. (2015) for the Wei river basin of China, Love et al. (2010) for the Limpopo river of Southern Africa and Abeysingha et al. (2015) for the Gomti river basin in north India, found that a decreasing trend of rainfall resulted in an significant reduction in the streamflow. In contrast, Masih et al. (2011) in the Zagros mountains of Iran, Wilk and Hughes (2002) in South India and Abdul Aziz and Burn (2006) in the Mavkenzie river basin of Canada, reported that an a trend of increasing trend of rainfall have has significantly increased river flows. There are also a number of studies (e.g., Tekleab et al., 2013; Gebremicael et al., 2013; Wang et al., 2015; Hannaford, 2015; Saraiva et al., 2015) who found that changes in rainfall are not sufficient to explain the trends in the streamflow.

In Ethiopia, few studies have analysed the trend of hydro-climatic variables including, streamflow and rainfall. Conway and Hulme (1993) reported declining annual rainfall over the Blue Nile and Atbara basins resulting in a reduction of river flows between 1945 and 1984. In contrast, recent investigations by Tesemma et al. (2010), Tekleab et al. (2013) and Gebremicael et al. (2013) agreed that rainfall over the Upper Blue Nile basin did not show a statistically significant trend for the last 40 years (1964-2005). Despite that the pattern of rainfall remained constant, hydrological flows in the basin showed a heterogeneous trend. Rainy (June to September) and dry (October to February) season flows at the Upper Blue Nile basin outlet (EI Diem) have significantly increased and decreased, respectively, while the mean annual flow runoff did not show a statistically significant trend. This implies that trends observed in the river flows may not be attributed to climate change but rather to changes in catchment characteristics. The lack of consensus in the literature may also show that there is still considerable uncertaintyies about the impact of climate change on the hydrological regimes of

the region. The length of the statistical record has a direct implication on the results of the trend analyses and some of the observed discrepancies could be because of applying different periods of time series data. For example, Dixon et al. (2006) investigated the impact of record length on the trend pattern of stream-flow in Wales and central England and their results indicated that trends over 50 and to 60 years showed a statistically significant increasing trend, while for a medium record length (30–40 years) remained constantno such trend was detected. Meanwhile, Record lengths less than 25 years tended to showed statistically significant increasing trends. This shows that trend analyses are sensitive to the time domain and careful attention should be given during analyses. Moreover, the variability in climatic zone within a basin may also influence the hydrological regimes. Many studies (e.g. Castillo et al., 2003; Yair & Kossovysky, 2002; Li & Sivapalan, 2011) reported that the spatio-temporal runoff generation in semi-arid areas is strongly non uniform as runoff generation controlling factors are different from that of a humid environment.

With regard to the Tekeze-Atbara river basin, no-it lacks comprehensive study of the hydroclimatic trends-exist. Seleshi and Zanke (2004) attempted to investigate the pattern of rainfall over the upper part of Tekeze River basin by considering only one climatic station. Their output demonstrated that the amount of rainfall remained constant for the past 40 years (1962–2002). Despite the importance of streamflow for theto ensure sustainable use of water resource utilization and ensuring food security in the semi-arid regions of the country, long-term trends and change point of flow regimes and the association with climate change are not yet well understood. Therefore, it is important to understand the connections between rainfall and streamflow trends of the basin and establish whether hydrological variability is driven by changes in climate or by changes in catchment characteristics or both. This paper is intended to (i) investigate the spatiotemporal variability of rainfall and streamflow in the headwaters of Tekeze basin, (ii) identify any abrupt changes if significant trends exists, and (iii) explore the impact of climate change reflected in change in rainfall patterns on the hydrological variability of the basin.

2. Study area descriptions

The study area is the upper part of Tekeze River basin, located in Northern Ethiopia between longitude $37.5^{\circ} - 39.8^{\circ}$ E and latitude $11.5^{\circ} - 14.3^{\circ}$ N (Fig .1). The Tekeze Rriver originates in the southern part of the basin near the RasDeshen Mountains and flows in the northern direction and then turns towards west flowing into north-eastern Sudan, where the river joins the Atbara River (Zenebe, 2009; Belete, 2007). This basin is one of the major tributaries of the Nile River

which drains from an area of 45,694 km² at the Embamadre gauging station (Fig. 1). The mean annual flow at this point is 5.4 10⁹ m³ yr⁻¹, which is about 66 % of the total annual flow whhere the Atbara joins the main Nile. The topography of the basin is characteriszed by ruagged topography consisting of mountains, highlands and terrains of gentle slopes. EThe elevation of the basin varies from 834 m.a.s.l at the basin outlet in Embamadre to more than 4528 m.a.s.l in the Ras Dashen mountains.

The general climate of the basin is semi-arid in the east and north and partly humid in the south, where rainfall ranges from below 400 mm yr⁻¹ in the east to more than 1200 mm yr⁻¹ in the south (Belete, 2007; Zenebe, 2009). More than 70 % of the total annual rainfall falls in two months (July and August). The variations of rainfall over the basin are mainly associated with the seasonal migration of the inter-tropical convergence zone (ITCZ) and complex topography (Nyssen et al., 2005). The river flow pattern typically follows that of rainfall. Maximum stream flowdischarge occurs in August, while it ceases completely during the dry season from October to February.

Dominant land use in the basin includes cultivable land (>7070 .5%), open grass-land, sparsely grown woodland, bushes and shrubs and exposed rocks (Tefera, 2003). This basin is characterized by severe land degradation through deforestation, over-grazing and cultivation on the rugged topography. However, it is also known for its more recent experiences with soil and water conservation (SWC) activities (Alemayouh et al., 2009; Nyssen et al., 2010). Physical SWC structures (Alemayouh et al., 2009; Negusse et al., 2013) and biological SWC measures through plantation and exclosures (Descheemaeker et al., 2006, 2008; Belay et al., 2014) have been practiseed in the semi-arid parts of the basin. These interventions have been implemented at watershed level for the last three decades. —In addition, one large hydropower dam inaugurated in 2009 is found approximately 83 km upstream of the basin outlet (Fig. 1) which may also alters the downstream flow regimes. in the downstream.

3. Data and Methods

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Temporal and spatialSpatio-temporal datasets of rainfall and streamflow are required for the trend and change point analyses. These statistical analysis analyses directly depends on the quality and length of the time series data. Therefore, muchmore efforts were was given to verify the accuracy of the rainfall and streamflow data. Scrutinizing of the time series data are summarized in section 3.1 and 3.2.

3.1. Rainfall data

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For this study, daily rainfall data since 1953 were used from 21 stations located within and surrounding the basin (Fig. 1 and Table 1). These data were provided by the Ethiopian National Meteorological Service Agency. After scrutiny of all stations, only 21 out of more than 75 stations in the basin were considered for further analyses. Despite that tThe length of the records varied varying from station to station, whilst all gauging stations with at least 30 years of continuous and relatively good quality of observed data were taken into account. The A 30 years record period is a reasonable minimum length for applying statistical trend analyses of rainfall (Love et al., 2010; Longobardi and Villani, 2009). The location and general information of all rainfall stations are shown in Fig. 1 and Table 1.

3.1.1. Rainfall data analyses and validation

Visual inspection, linear and multiple regression analyses between neighbouring stations and other global datasets, including New_LocClime software package (Grieser et al., 2010), Climate Hazards Group Infrared Precipitations (CHIRPS) (Funk et al., 2014) and Tropical Rainfall Measuring Mission's (TRMM) (Simpson et al., 1988), were applied for data analyses and validation, detecting outliers, filling missing values and reliability checking for all gauging stations. The rainfall datasets were found to be reliable to be used for statistical analyseis for the following screening criteria-

The coefficient of variation in annual rainfall of the basin rangeds from 18 % in the southern to 33 % in the eastern and northern parts of the basin. As presented shown in Table 1, all but two stations have a coefficient of variation below 30 % which is an acceptable limit for data validation (Medvigy and Beauliew, 2011; Sushant et al., 2015). To ensure data continuity and integrity, missing rainfall data of less than 1 year were estimated from global and neighbouring stations and data gaps larger than 1 year were excluded from the analyses. Based on these data screening and analyses methods, rainfall stations with less than 10 % missing data have been used in the analyses.

Satellite data including Climate Hazards Group Infrared Precipitations (CHIRPS) and Tropical Rainfall Measuring Mission's (TRMM) were used to validate and filling the missing values. Among many rainfall estimates (RFE) in the area, e.g., (RFE, CMORPH, ERA40, CHIRPS, TRMM), the last two were used to validate and fill in missing data. Both satellite data sources have a relatively high resolution (TRMM 0.25° and CHRPS 0.05°) and, are commonly used in Africa (Shukala et al., 2014; Katsanos et al., 2015). Detailed descriptions of these rainfall

products are documented in many publications (e.g. Dinku et al., 2007; Funk et al., 2014; Katsanos et al., 2015; Simpson et al., 1988). Before using them for validation and reliability checking, the rainfall products were first compared directly with observed rainfall of selected stations with good quality data. As the observed rainfall data in the region are sparse and unevenly distributed, the point data were not interpolated into gridded time series. Instead, area averaged time series of satellite rainfall products around each gauging station (~ 25 km radius) were taken for the comparison. A 25 km radius of area average rainfall was considered to account the satellite data resolutions as well as to avoid the effect of topography complex on the rainfall estimation. A common time period (1998-2015) of the satellite (y) and ground rainfall (x) data were considered for the comparison.

The performance of the satellite products in estimating the amount of rainfall around the gauging stations were evaluated using statistical measures shown in Table 2. Their full descriptions can be found in Toté et al. (2015), Thiemig et al. (2012), Derin and Yilmaz (2014). Both CHIRPS and TRMM (3B42v7) satellite products were evaluated against monthly rainfall values from seven observed stations. The performance results show that CHIRPS outperformed TRMM on for the majority of stations (Fig. 2). CHIRPS rainfall data showed a good correlation with observed rainfall in the stations. Accordingly, the CHIRPS rainfall data were used for reliability checking and filling missing values of all stations before trend analyses. It has been reported that this product has a better performance in Africa (Funk et al., 2014; Toté et al., 2015). Hessel (2015) compared 10 satellite rainfall products over the Nile basin and CHIRPS products performed the best and were recommended for the basin. The likely reason for its better performance could be because of its availability at a high resolution of 0.05 degrees.

3.2. Streamflow data

Streamflow data from all gauging stations in the basin were obtained from the Ethiopian Ministry of Water Resources and Energy. Although the recording of flow data over the basin started in the late 1960s, it was discontinued for most of the gauging stations during the civil war in the 1980s. To this effect, only nine out of the total 39 stations have an extended period of more than 20 years data and these were these used in the analyses. Spatio-temporal trends can be affected by the chosen length of records. A longer period of historical data increases the visibility of dominant trends and increases the reliability of results from trend analyses, while shorter period enhances spatial coverage of streamflow by including more representing stations at different parts of the basin. To better account the spatial variability of streamflow, A-a length

of more than 20 years data is can be desirable for trend analyses of streamflow (Abdul Aziz and Burn, 2006; Saraiva et al., 2015; Abeysingha et al., 2015). The sensitivity of trend to the length of flow record is discussed in the introduction (section 1). The location and general information of all flow stations used are summarized in Fig. 1 and Table 3. The average annual flows of each station indicate that hydrological responses are spatially uneven over the basin. For example, Illala and Werie have higher streamflow per unit area as compared to Genfel and Geba 1 tributaries. Despite there is a large difference in the drainage area (Table 3 and Fig.1), approximately, the same volume of runoff is contributed to the Geba 1 from Genfel and Illala tributaries. Moreover, looking into the drainage area of Geba 2 (4590 km²) and Werie (1770 km²), more water is discharged from Werie (5 %) than from Geba 2 (6 %) to the basin outlet (Tekeze 2) which explicitly point tosuggest a high variability in hydrological response to catchment characteristics.

As hydro-meteorological data in the basin, if not in all basins in Ethiopia, is scanty-very limited and with many gaps, it is critical to carefully screen and checks their quality before using them for analyses. Hence, the raw data were visually inspected and screened for mistyped typos and outliers. Each station was carefully checked its for data consistency in by comparing to the nearby, upstream and downstream stations. Relationships between neighbouring stations can give a preliminary evidence on the reliability of time series data, provided that there is no manmade water storage above the station (Hong et al., 2009). Identified unreliable data were fixed after comparing its upper and lower boundary limits. Furthermore, heterogeneity of the time series data was also detected using the double mass curve and residual mass plot methods. The monthly hydrological flow data were aggregated from the daily data and the seasonal and annual data was calculated from the monthly data. In order to remove unreliable data whilst including more stations to increase spatial coverage, Finally, Missing missing data for more than two years were excluded in-from the analyseis. However, Deduring the peak rainy season missing data for more than two weeks were excluded from the analyses. The reason to exclude only two weeks was to minimize untrustworthy data as where more than 80 % of the river flow is generating during only two months (July and August).

3.3. Trend analyses method

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To identify the trends of in rainfall and streamflow, a non-parametric Mann-Kendall (Kendall, 1975) statistical test is applied. The Mann-Kendall test (MK), is a rank based method that has been widely used to detect the trend of hydro-climatic time series data in different parts of the

world (e.g Abdul Aziz and Burn, 2006; Gebremicael et al., 2013; Jones et al., 2015; Návar, 2015; Tekleab et al., 2013; Mohamed & Savenije, 2014; Wang et al., 2015). The procedure of MK testing approach starts by calculating the MK statistic s using Eq. (1) (Yue et al., 2002).

$$S = \sum_{i=1}^{n-1} \sum_{j=i+1}^{n} \operatorname{sgn}(x_j - x_i), \text{ where } \operatorname{sgn}(\theta) = \begin{cases} +1 \\ 0 \\ -1 \end{cases} \quad \begin{bmatrix} \theta > 0 \\ \theta = 0 \\ \theta < 0 \end{bmatrix}$$
 (1)

Where xj and xk are the data values in time j and k and j>k, respectively, and n is the length of data set. The normalized test statistics Z of MK test and the variance VAR(S) were calculated as shown in Eq. (2) and Eq. (3).

$$V(S) = \frac{1}{18} [n(n-1)(2n+5)] \tag{3}$$

where s and v(S) are the Kendall's statistics and Variance, respectively. The MK test calculates Kendall's statistics s, which is the sum of the difference between data points and a measure of associations between two samples (Kendall's tau). The MK test, accepts the null hypothesis if $-Z \le Z$ cr $\le Z$, where Zcr is critical value of the normalized statistics Z at 5 % confidence level (1.96). Positive and negative values of those parameters (z, s, and tau) indicate an 'upward trend' and 'downward trend', respectively. In order to evaluate the trend results, the Z value combined with the computed two-tailed probability (P) were compared with the user defined confidence level (5 %) of the standard normal distribution curve. The MK test is commonly used and suitable to identify trends in water resources as it is not affected by the distribution, outliers and missing values of time series data (Yue et al., 2002, 2003; Zhang et al, 2008, 2011).

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The existence of serial correlation in the time series data may affect trend detection in the non-parametric trend test methods (Masih et al., 2011; Zhang et al., 2011). The Trend-Free Pre-Whitening (TFPW) method (Yue et al., 2003) was employed to avoid serial correlations in the data. This method is found to be the most powerful tool to remove a serial correlation time series if it exists (Yue et al., 2003; Burn et al., 2004; Tekleab et al., 2010; Mohamed & Savenije, 2014). The data series was pre-whitened using the formula shown in Eq. (4).

$$Y_1 = Y_t - r1.Y_{t-1} (4)$$

Where r_1 is the estimated serial correlation coefficient, Y_t is trended series for time interval t, and Y_1 is data series without auto-regressive, and Y_{t-1} is the original time series value. Detailed descriptions of TFPW can be seen found in literatures (e.g., (Yue et al., 2003; Burn et al., 2004; Tekleab et al., 2010, Gao et al., 2011). Finally, the MK test was applied to the trend free prewhitened TFPW data series for analysing the gradual change in the rainfall and stream-flows.

3.4. Change point detection method

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To estimate the occurrence of an abrupt change, a non-parametric Pettitt test (Pettitt, 1979) was applied to the trend-free pre-whitened-TFPW data series. The Pettitt test is a rank-based and distribution-free test for identifying if there is a significant change between cumulative functions before and after a time instant. The Pettitt test considers a sequence of random variables $X_1, X_2, ..., X_T$ that can have a change point at τ if X_t for $t, ..., \tau$ have common distribution function $f_1(x)$ and X_t for $t = \tau + 1, ..., T$ have a common distribution function $f_2(x)$, and $f_1(x) \neq f_2(x)$. The test statistics $K_T = Max \mid U_t, T \mid , 1 \leq \tau < T$ and associated probability (P) used to test were computed using equations found in Gao et al. (2011). The test was evaluated against a user-defined significance level (5 %) and P values less than 5% were considered as a statistically significant change in the data series. The 5% significance level was chosen as it is commonly used in the hydro-climatic trend analyses (e.g. Tekleab et al., 2013; Gebremicael et al., 2013; Saraiva Okello et al., 2015). This Pettitt technique has been widely used to detect time change points in the hydro-climatic data (e.g., Ma et al., 2008; Love et al., 2010; Gao et al., 2011; Zhang et al., 2011; Gebremicael et al., 2013; Tekleab et al., 2013).

3.5. Hydrological alteration indicators

Temporal and spatial streamflow variability can also be characterized and quantified using hydrologic alteration indicators. The Indicators of Hydrologic Alteration (IHA) software developed by the US Nature Conservancy (Mathews and Richter, 2007) were applied to assess the degree of hydrological alteration. Twenty out of Eight parameters out of the total 33 IHA parameters were selected for this study. The selected parameters are magnitude and duration of annual extreme water conditions (e.g. 1–day, 3–day, 7–day minimum and maximum flows), rate and frequency of water condition changes (e.g. rising rate and falling rate) and magnitude of monthly water conditions (e.g. monthly flows). Such IHA parameters are common in

characterizing of hydrological regimes influenced by climate and anthropogenic factors (Tayler et al., 2003; Mathews and Richter, 2007; Masih et al., 2011; Saraiva et al., 2015). The consistency of those parameters were was analysed and compared with the user-defined P values (5 %).

4. Results and discussion

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According to NMSA (1996), major seasons in the study area are the rainy (June-September), dry (October-February) and short rainy seasons (March-May). Before detecting trends in precipitation and hydrological flows, serial correlation existence in all datasets were tested at monthly, seasonal and annual scales. Accordingly, $9.1\frac{\%}{}$ – 18 % of the monthly, seasonal and annual data of the rainfall stations were found to have a statistically significant auto-correlation at 95% confidence level. This indicates that a false significant trend might have appeared in some of the stations when actually there is no trend because of auto-correlation the stations when actually there is no trend in the data series. Similarly, $3.2\frac{\%}{}$ – 6 % of the monthly and seasonal flows showed statistically significant auto-correlation. -It is unclear that why the monthly and seasonal fluxes showed stronger autocorrelation than annual. The probable reason could be because of storage properties in the catchments, non-normalunreliable data and missing values (Hirsch and Slake, 1984; Abeysingha et al., 2015). Furthermore, continuous constant observations in the dry months, where the river flow discharges are very small low may have increased the degree of temporal dependencysimilarity among consecutive observations. To avoid such spurious trend detection, serial correlation problems in all time periods were eliminated using Trend-Free-Pre-Whitening techniques before trend analyses. Comparison of rainfall and streamflow before and after TFPW at different scales (monthly, seasonal and annual) is presented as a supplementary file in Fig. S1 and Fig. S2, respectively.

4.1. Temporal and spatial Rainfall variability of over the basin rainfall

The presence of monotonic increasing/decreasing trends in monthly, seasonal and annual rainfall of 21 gauging stations were was tested using the MK test. The results for seasonal and annual rainfall are summarized in Table 4. Positive and negative values of Z statistics are showingshow increasing and decreasing trends, respectively. Bolded—Z statistics in bold illustrates statistically significant trends of rainfall. The spatial distribution of observed significant and non-significant trends of annual rainfall over the basin is also given in Fig. 3.-

Results of the trend analyses were used to identify if the time series of annual and seasonal rainfall had a statistically significant trend in the last 30–60 years (Table 4). The results of the rainfall analyses shows no trends. Fig. 3 shows the spatial variability of rainfall on annual scale throughout the basin. Except for two stations (Axum and Shire) (annual), and one station (rainy season), there is no significant in the rainfall trends in the Upper Tekezofe the basin. Both Axum and Shire stations which are located in the North West part of the basin showed an increasing and decreasing trend, respectively (Fig. 3). The possible reason for obtaining different result in these stations could be because of unreliable data as both stations have the highest percentage of missing data comparing to the remaining stations. However, although statistically not significant, statistical indices of the test revealed a tendency of decreasing rainfall patterns in the eastern and northern part of the basin during the main rainy season and the entire year (Table 4). Meanwhile, there anis an increasing tendency in the southern and western parts of the basin for the same time scales. With regard to monthly rainfall, despite there is was some temporal and spatial variability of monthly rainfall, no dominant trends are found for in the majority of the months (results not provided here Table S1)).

The Pettitt test was used to identify if there is a breakpoint in the data series. Similar to the MK test, the majority of rainfall stations did not show statistically significant change points at 5 % significance level (see Table S2). For example, change point for annual and seasonal rainfall of some stations can be seen in Fig. 43. Fig. 43a and 43de illustrate that annual rainfall in Mekelle (AP) and Adigrat stations do not show an upward or downward shift in the given time domain. An increasing and decreasing change point of annual and short rainy season (March–May) rainfall in Axum and AbiAdi stations are shown in Fig. 43b and 43cd, respectively.

In general, most of the rainfall stations across the basin did not experiences a statistically significant trend at 5 % significance level. The result also reveals that even though there is no dominant trend, monthly rainfall over the basin is observed to be is more variable than compared to the seasonal and annual rainfall over the basin (Table S1). The possible reason could be the amount of rainfall in a given month is shifted to the next or previous month and the station might be receiving the same amount of seasonal rainfall but varying in distribution among the months. Another possible reason could be topographically induced climate. For example, Bizuneh (2013) noticed that monthly rainfall variability in Siluh catchment of Geba basin is dependent on altitude. Furthermore, observed monthly rainfall variability might be due to large-scale oscillation (e.g. ITCZ) variability rather than long—term climate variability in the basin.

The results of this study are consistent with Seleshi and Zanke (2004) who found no significant trends of rainfall at Mekelle station. Results from neighbouring catchments of similar climate characteristics (e.g., Tekleab et al., 2013; Gebremicael et al., 2013) and applying the same methods of trend analyses have also shown that the pattern of rainfall remained constant for the last 40 years which is in agreement with our finding (e.g., Tekleab et al., 2013; Gebremicael et al., 2013).

4.2. Streamflow variability

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4.2.1. Long term trends of streamflow

Streamflow of nine gauging stations (Table 3) was analysed for long-term trend detection using MK and Pettitt test. Table 5 summarizes the results from the MK test, where and positive/and negative values of Z statistics associated with the computed probability (P-value) shows indicate an increasing/or decreasing trends. Statistically significant trends are shown in bold. Furthermore, the observed trends are also presented in Fig. 5 to show the spatial variability of streamflow over the basin. The positive, negative and zero icons in the map indicate an increasing, decreasing and no change of dry season streamflow trends across the basin, respectively.

Annual flow patterns exhibited a declining trend in the majority of stations although this time is more pronounced in the eastern part of the basin (Table 5). The change is found to be statistically significant at Siluh, Genfel, Geba stations. Interestingly, although there is a dominant decreasing pattern in the majority of the tributaries, the annual flow at Embamadre station did has not significantly changed. Seasonal streamflow of the stations were also analysed to further scrutinize temporal and spatial variability.

The analyses found that a significant decreasing of decrease of dry season flow has occurred in for most stations (Table 5 and Fig. 5). During the short rainy season, a decreasing trend has occurred in for five stations, some of these trends being statistically significant. Nevertheless, in the dry and short rainy seasons flow has significantly increased only at the basin outlet (Embamadre). Streamflow showed a significant increasing trend at Embamadre and a non-significant increasing trend at Kulmesk during the main rainy season. In contrast, a gradual decreasing pattern of flow was recorded in the remaining gauging stations with this change being significant in four stations (Table 5)... Most of the decreasing trends were observed in the

eastern and northern parts of the basin where land degradation is believed to be very high compared to the southern and western parts of the basin (Fig. 5).

The Mmajority of the gauging stations did not show a consistent trend in monthly streamflow. For example, discharge in of Siluh and Genfel stations is characterized by a decreasing trend in most months. In Agulae and Illala catchments, Aa significant decreasing trend is found in April and May flows of Agulae and Illala watersheds while the remaining months observed a decreasing trend that was not statistically significant (Table 5). The combination of Siluh, Genfel and Agulae tributaries at Geba station near Mekelle showed a decreasing trend in all months. Monthly flow patterns of the Upper Tekeze River Basin, the sum of all gauged (Table 3) and ungagged ungauged tributaries at Embamadre station, revealed a significantly increasing trend in April and May while all other months remained unchanged (Table 5).

The Pettitt test was also applied to identify an abrupt change of streamflow in the catchments. The change points of annual and seasonal flow for selected stations are shown in Fig.- 46. For the annual flow, stations did not show consistent shifts across all stations. For example, annual flows in Siluh and Geba catchments shifted downward after 1992 and 2002, respectively, while no significant abrupt change was observed in Genfel and Tekeze at Embamadre despite strong monthly and seasonal variability (Fig. 46). Change points of seasonal flow for the same stations confirmed an abrupt change in the downward and upward directions (Fig. 46). Dry and short rainy season flows in all stations except at the basin outlet showed a significant downward shifts since the early 2000s. Conversely, an abrupt increase of in streamflow has occurred at the basin outlet for the same seasons. The Pettitt test has also shown that hydrological flows during the rainy season remained constant in for most stations (Fig.- 46).

In summary, most stations exhibited a statistically significant change during the short rainy (5 stations) and dry seasons (6 stations). Similarly, trends in the main rainy season and annual flow showed a significant change in 3 to 4 of the stations. Several stations exhibited a change points of monthly streamflow, results not shown here (see Fig. S3). Both upward and downward shifts in streamflow was observed in many stations for the months of January, April, May and August (Fig. S3). However, there was no dominant (increasing or decreasing) trend across the basin. For example, a change points of monthly (January, April, May and August) streamflow is observed in Siluh catchment (Fig.S3). A downward shift of monthly streamflow was has occurred since 1996. The monthly stream flow of Genfel and Geba catchments has significantly declined starting from 2003 and 2004, respectively. In contrast, an upward shift of monthly streamflow was observed at Emabamadre for January, April and May which became significant after 2009.

Interestingly, some of the trends in the upper catchments counterbalance each other when combined in the downstream stations. For example, negative trends during the short rainy season in Siluh, Genfel, and Agulae cancelled out when combined at Geba near Mekelle. A remarkable result was that, although a majoritythe majority -stationsof stations in the upper catchments showed a declining pattern of streamflow, the entire basin flows at the outlet did not show a negative trend., in some cases, the base flow significantly increased after 2009. SThe significant increasing trend during the dry and short rainy season at this station is most likely due to the construction (2009) of Tekeze hydropower dam located 83 km upstream of the station. Change in catchment response from ungagged-ungauged catchments of the basin might also have contributed to increaseing the flow at this station.

The above results are in agreement with previous local studies (e.g. Abraha, 2014; Bizuneh, 2013; Zenebe, 2009) who found strong variability of stream flows in different sub-catchments of the basin. Compared to the neighbouring basin (Upper Blue Nile) studies by Gebremicael et al. (2013) in at four stations and Tessema et al. (2010) in at three stations who found a significant increasing trend of streamflow in short rainy, main rainy, annual flows and a decreasing trend in the dry season flow, the Tekeze basin, particularly in the semi-arid parts of the basin experienced a significant decreasing trend and high variability of streamflow. This variability is expected as land degradation (e.g. deforestation, over cultivation and grazing) in Tekeze basin is more pronounced than other basins in the country (Awulachew et al., 2007; Gebrehiwot et al., 2011; Gebreyohannes et al., 2013; Yazew, 2005; Zenebe, 2009).

4.2.2. Hydrological Alteration Indicators

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Although, the previous analyse is showed the long-term trends of rainfall and streamflow in the Upper Tekeze basin, it could not address the short dynamics-period fluctuations of the hydrology within the catchment, and whether this can explain some of the results given above. The magnitude and duration of annual extreme conditions were also analysed using six IHA parameters (1–day, 3–day and 7–day annual minima and maxima). Results from these extreme conditions indicate a dominant significant decreasing trend in both minimum and maximum daily flows (Table 6). On the other hand, a significant increasing trend of the minimum flow is detected at Embamadre station.

The trends in the rate and frequency of changes in water-hydrological conditions were also explored using rise and fall rate parameters. Accordingly, the rising rate in-of daily flow of all stations remained constant while the daily falling rate has significantly increased in the tributaries and decreased in the basin outlet. It can be seen that the trend of minimum flows

described by 1–day, 3–day and 7–day, is consistent with the trend of monthly and seasonal flows. Moreover, the IHA change point analyses has also shown shifts in minimum and maximum flows during the dry and wet seasons of the catchments. The extreme 1–day and 7–day minimum and maximum flows significantly shifted downward at the Siluh and Geba catchments. In the Tekeze at Embamadre station, the 1–day and 7–day minimum flows significantly increased around 2003, but 1–day and 7–day maximum flows remained unchanged. Illala catchment experienced a decrease of the 1–day minimum and an increase of in the 1–day maximum annual flows with change points at around 2000 and 1995, respectively. Extreme high flows characterized by the 1–day and 7–day annual and maximum flows did not significantly change at the basin outlet which may be also-due to homogenisation of the low flow and peak flow hydrograph after the construction of Tekeze hydropower dambecause of the dam—above the station (see sect. 4.2.1)

4.2.3. Drivers for streamflow variabilities

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Climatic conditions, and the in particular rainfall, as well as human activities in a catchment are the most important factors influencing the hydrological variability of streams. In this study, temporal and spatial analyses of rainfall from both MK and Pettitt tests showed that rainfall over the basin did not significantly change during the period of analyses. Streamflow, in contrast, did exhibit a strong temporal and spatial variability in the basin. This suggests that the change in hydrological flow is not significantly influenced by rainfall. The timing of observed trends in streamflow is not uniform., however, tThis may indicate that the impact of human interference and physiographic characteristics differs from sub-catchment to sub-catchment. Trend analyses is sensitive to the time domain as different results can be obtained for different time periods. In this study, however, change points occurred at different times in most of the sub-catchments even for the same time domain (e.g. Genfel and Agulae). This implies that effect of changes in the underlying surface characteristics could be the physical mechanism behind those variations. Human interventions expressed in terms of water abstraction, implementation of large-scale soil and water conservation, deforestation, and afforestation in the upstream catchments are the more likely driving forces of changes in the flow regimes rather than climatic conditions. For example, Nyssen et al. (2004) and Belay et al. (2014) reported that a strong decrease of forest and bushland has occurred in favour of cultivable and grazing lands from 1960s to early 1990s. However, the rate of deforestation and land degradation in most parts of the basin has slowed down and even started to recover by increasing the coverage of vegetation from late 1990s onwards (Nyssen et al., 2010). The conversion of vegetation cover into agricultural land has increased runoff by 72% and decreased dry season flow by 32% in some parts of the basin. Studies in neighbouring basins (e.g. Upper Blue Nile) also confirmed that conversion of vegetation cover into agriculture and bare land has caused an increasing of surface runoff and decreasing of base flow up to 75% and 50%, respectively.

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Increasing water abstractions, particularly in the semi-arid catchments of the basin, might have caused the decline of streamflow during dry and small rainy seasons. Several studies, e.g., (Alemayehu et al., 2009; Kifle, 2015; Nyssen et al., 2010) have shown that surface and shallow groundwater development and abstraction for irrigation have significantly increased since the mid-2000s, after implementation of intensive catchment management programmes. Moreover, a strong monotonic trend in streamflow without a significant change in rainfall during the rainy season could be attributed to the large-scale soil and water conservation interventions in the upstream watersheds. For example, Nyssen et al. (2010) and Abraha (2014) reported that integrated catchment management and land use change have significantly reduced streamflow in Geba catchment. On the other hand, sSoil and water conservation interventions have significantly increased the availability of groundwater at sub-catchments levelsmaller watershed levels (Alemayehu et al., 2009; Negusse et al., 2013). All these studies are consistent with our findings that observed streamflow alterations in the basin are most likely the result of upper catchments interventions rather than changing patterns of rainfall. Quantifying the impacts of such factors at large scale is beyond the scope of this study and further investigations should be conducted to study the effect of anthropogenic factors on streamflow variability and change at different scales.

It is also essential to point out some limitations in this study. The aAbsence of common records of rainfall and streamflow data, limited coverage of streamflow time series data across the basin, data gaps in the peak flow period, are the major limitations that may have increased uncertainties in the trend analyses. The length of the record period used for the trend analyses of rainfall and streamflow varied from 31–63 and 20–43 years, respectively. Analyses using these different record lengths of data may introduce some discrepancy in the analyses.

5. Conclusion and Recommendations

This study presents a detailed statistical analysies on the existence of trends and point changes of rainfall and streamflow in the Upper Tekeze River basin. The analyses wereas carried out for 21 rainfall and 9 streamflow monitoring stations. Those stations were selected based on the

availability and quality of data from 39 streamflow and more than 70 rainfall stations available in the basin. Linkages between the trends in rainfall and streamflow across the whole basin were carefully examined at different scales. Following these analyses the main driving force for streamflow variability over the basin is deduced.

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Rainfall over the basin has remained constant in the last four decades. The 19 out of the 21 tested stations experienced neither increasing nor decreasing trends during the dry, short rainy, main rainy seasons and annuals at 95 % confidence level. Furthermore, the result of this study clearly showed that monthly rainfall in the majority of the stations experienced high spatial variability compared to the seasonal and annual time scales. In contrast, trend analyses of different hydrological variables showed that streamflow in most stations has changed significantly. A decreasing trend in dry, short, main rainy seasons and annual totals is dominant in six out of the nine stations, located inat the semi-arid areas of the basin. The significant decreasing pattern of streamflow is observed in the eastern and northern part of the basin where land degradation is very high. Only one station, located at the basin outlet, exhibited a significant positive significant trend during both the dry and short rainy seasons. The different trend in this station is likely due to the construction of Tekeze hydropower dam in 2009. The remaining two out of the nine stations stayed constant in all seasons. Findings from both MK and Pettitt tests are consistent in all seasons and stations, but the timing of change points is different for most stations. This could imply that the level of human interference and physiographic characteristics is varying from sub-catchment to sub-catchment and hence the differences in runoff generation response to catchment characteristics.

Surprisingly, our results showsed that there is no linkages/patterns between the trends in rainfall and streamflow in the basin. This suggests that the change in streamflow is influenced by factors other than rainfall. A weak relationship between rainfall and streamflow leads to the conclusion that the significant trends in streamflow could be due to significant changes over time of catchment characteristics, including land use/cover change, catchment management interventions and water abstractions in the upstream. This was also supported by few existing studies as discussed in section 4.2.3.

The findings from this study are useful as a pre-requisite for studying the effects of catchment management dynamics on the hydrological variabilities. Statistical trend analyses investigate s only the trend of historical data without being able to identify the causes of those trends. Therefore, further investigations are needed to verify and quantify the hydrological changes shown in statistical tests by identifying the physical mechanisms behind those changes.

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Table 1. General information of rainfall stations, latitude and longitude, altitude (Alt.) in m.a.s.l, mean annual rainfall (mm yr⁻¹), standard deviation (mm yr⁻¹), and % age-of missing data

Station name	Lat ^o	Long ^o	Alt. (m)	Recording period	Analyses period	Mean	SD	CV (%)	Missing data (%
Mekelle	13.45	39.53	2260	1952-2015	1953-2015	576	141	24	0.0
Mychew	12.69	39.54	2432	1953-2015	1953-2015	697	158	23	6.3
Axum	14.12	38.74	2200	1962-2015	1963-2015	690	159	23	9.4
Gonder	12.60	37.50	2316	1952-2015	1964-2015	1090	195	18	0.0
Adwa	14.16	38.90	1950	1964-2015	1967-2015	705	176	24	7.6
Mykinetal	13.94	38.99	1815	1967-2015	1967-2015	585	129	22	8.1
Shire	14.10	38.28	1920	1963-2015	1968-2015	953	203	21	10.1
Adigrat	14.00	39.27	2470	1970-2015	1970-2015	596	172	29	2.1
Adigudem	13.16	39.13	2100	1975-2015	1971-2015	498	156	31	2.2
E/hamus	14.18	39.56	2700	1971-2015	1971-2015	651	214	33	2.2
Hawzen	13.98	39.43	2255	1971-2015	1971-2015	505	116	23	6.6
Illala	13.52	39.50	2000	1975-2015	1975-2015	563	138	25	4.9
H/Selam	13.65	39.17	2630	1973-2015	1973-2015	685	168	24	0.0
AbiAdi	13.62	39.02	1850	1961-2015	1973-2015	861	246	29	2.3
Samre	13.13	39.13	1920	1967-2015	1978-2015	650	188	29	6.1
D/tabor	11.85	38.00	2969	1974-2015	1974-2015	1502	264	17	2.3
Dengolat	13.19	39.21	1950	1975-2015	1975-2015	617	166	27	2.4
Lalibela	12.03	39.05	2450	1972-2015	1978-2015	789	169	21	5.3
Wukro	13.79	39.60	1995	1962-2015	1985-2015	485	139	29	9.4
Kulmesk	11.93	39.20	2360	1973-2015	1985-2015	668	180	27	3.2
Debarik	13.15	37.90	2850	1955-2015	1984-2015	1104	231	21	6.2

Table 2. Statistical indices for accuracy measurement of satellite rainfall products

Tubic 21 Blandical marces for accurac	y integration of satellite familian	Products
Statistical measure	Equation	Best value
Relative Mean Absolute Error (R _{MAE})	$(1/N\sum (y-x) /(\bar{x})$	0
Nash-Sutcliffe efficiency (N_{SE})	$1 - (\sum (y - x)^2 / (\sum (x - \bar{x})^2)$	1
BIAS	$\sum y/\sum x$	1
Pearson correlation coefficient (r)	$\frac{\sum (x-\bar{x})(y-\bar{y})}{\sqrt{(x-\bar{x})^2}\sqrt{(y-\bar{y})^2}}$	1

Table 3. General Information of the hydrological flow monitoring stations

Station name	Lat.	Long.	Alt.	Catchment area (km²)	Recording period	Analyses period	Annual average	Missin g (%)
Siluh	13.85	39.51	2230	967	1973-2015	1973-2015	flow (m ³ s ⁻¹)	4.6
				,	-, =			
Illala	13.53	39.50	2004	341	1980-2015	1980-2015	0.6	2.8
Genfel	13.80	39.60	1997	733	1992-2015	1992-2015	0.6	2.1
Werie	13.85	39.00	1380	1770	1967-2015	1994-2015	10.1	0.8
Agulae	13.69	39.58	1994	692	1992-2015	1992-2015	1.1	0.3
Geba 1	13.6	39.38	1748	2445	1967-2015	1990-2015	3.9	4.0
Geba 2	13.46	39.02	1370	4590	1994-2015	1994-2015	14.2	0.0
Tekeze 1	12.60	39.19	1490	1002	1994-2015	1994-2015	3.0	3.2
Tekeze 2	13.74	38.20	845	45694	1969-2015	1994-2015	219.5	0.0

Geba 1 and 2 are Geba at Mekelle and Adikumsi and Tekeze 1 and 2 are Tekeze at Kulmesk and Embamadre.

Table 4. Summary results of MK, Z statistics on seasonal and annual rainfall trends. Negative/positive Z value indicates a decreasing/increasing trend and in bold a statistical significant trend at 5 % confidence level ($Z=\pm 1.96$).

station/season —	Rainy season	Dry season	Short rainy	annual
station/season —	Jun to Sep	Oct to Feb	Mar to May	Jan to Dec
Mekelle (AP)	-1.5	1.7	-1.2	-1.8
Mychew	-1.1	1.0	0.9	-0.7
Axum	1.5	0.3	0.3	2.5
Gonder	1.1	-1.7	1.8	0.8
Adwa	4.6	1.6	1.1	0.8
Mykinetal	1.7	-1.4	-1.8	-0.3
Shire	1.6	1.4	1.5	-2.5
Adigrat	-0.4	1.1	1.1	-0.1
Adigudem	0.0	0.1	-0.8	-0.3
Edagahamus	-0.6	-1.2	-1.7	-1.1
Hawzen	0.7	-1.4	-1.1	-0.4
Illala	1.5	0.3	1.6	1.6
Hagereselam	0.2	-0.3	-1.7	-1.0
AbiAdi	0.9	-1.5	1.9	1.9
Debretabour	1.3	1.5	0.1	1.9
Dengolat	-0.6	0.1	1.7	1.9
Lalibela	0.8	-0.6	-1.1	-1.1
Wukro	-1.4	0.8	1.2	1.6
Kulmesk	0.5	-0.3	0.1	0.1
Debark	0.3	0.4	1.8	1.7
Samre	-0.3	-1.9	-0.3	-0.3

Table 5. Summary results of MK, Z statistics on streamflow trends. Negative/positive Z value indicates a decreasing/increasing trend and in bold a statistical significant trend at 5 % confidence level ($Z=\pm 1.96$).

Period	Siluh	Genfel	Agulae	Illala	Werie	Geba 1	Geba 2	Tekeze 1	Tekeze 2
Record	1973- 2015	1992- 2015	1992- 2015	1980- 2015	1994- 2015	1991- 2015	1994- 2015	1994- 2015	1994- 2015
length (yr)	43	23	24	36	22	25	21	21	22
Annual	-4.5	-3.1	-0.9	-0.7	-1.1	-2.1	-2.7	0.6	1.0
Rainy season	-3.1	-3.2	-0.4	0.4	0.2	-2.1	-2.8	1.0	0.1
Short rainy	-2.4	-3.1	-3.5	-3.1	1.1	-1.5	2.7	0.63	3.9
Dry season	-5.0	-3.0	0.3	-3.0	-0.9	-2.2	-3.3	-0.2	3.4
Jan	-5.1	-2.3	-0.7	-2.1	1.4	-2.7	-3.1	0.4	1.6
Feb	-4.5	-2.6	-0.2	-1.4	0.8	-2.8	-0.4	-0.0	0.8
Mar	-5.9	-1.5	-0.5)	-1.6	1.2	-3.3	-0.2	0.4	1.3
Apr	-4.0	-2.4	-4.0	-2.9	0.9	-2.1	-2.6	0.7	2.4
May	-4.8	-2.1	-2.6	-2.2	1.3	-2.1	-0.3	0.7	2.3
Jun	-1.4	-1.9	-0.9	-1.6	0.6	-1.5	-0.2	0.6	1.2
Jul	-4.7	-2.6	-0.7	-1.0	-0.5	-1.6	-0.4	-0.3	-0.1
Aug	-0.9	-2.2	0.8	1.0	-1.2	-2.0	-3.4	0.3	-1.1
Sep	1.1	-1.6	-0.3	1.5	-0.1	-1.4	-0.4	0.6	0.3
Oct	-6.7	-1.8	-0.3	-0.8	3.8	-2.4	-0.6	0.3	0.9
Nov	-4.7	-4.8	-1.2	-1.5	0.7	-1.5	-1.1	0.3	0.7
Dec	-4.7	-2.6	-0.2	-1.5	1.2	-2.4	-3.4	0.2	0.4

Table 6. Summary results of MK, Z statistics on IHA trends. Negative/positive Z value indicates a decreasing/increasing trend and in bold a statistical significant trend at 5 % confidence level ($Z=\pm 1.96$).

IHA parameters	Siluh	Genfel	Agulae	Illala	Werie	Geba 1	Geba 2	Tekeze 1	Tekeze 2
	1973-	1992-	1992-	1980-	1994-	1990-	1994-	1994-	1994-
record	2015	2015	2015	2015	2015	2015	2015	2015	2015
length (yr)	43	23	24	36	22	25	21	21	22
1 4	4.2	2.2	0.7	2.4	0.0	2.6	2.0	0.6	2.0
1-day minimum	-4.3	-3.2	-0.7	-3.4	0.8	-3.6	-3.8	-0.6	2.8
3- day minimum	-3.5	-4.2	-0.8	-3.6	0.8	-3.4	-3.8	-0.5	2.2
7-day minimum	-3.5	-3.5	-1.3	-3.1	0.7	-3.9	-3.5	-0.6	2.1
1-day maximum	-5.7	-1.0	-0.1	0.7	-1.3	-2.7	-1.0	0.8	-1.0
3-day maximum	-5.5	-1.3	-0.7	0.7	-0.9	-2.6	-0.8	-0.1	-1.5
7-day maximum	-6.4	-2.2	-0.4	0.1	-1.3	-2.4	-0.7	-0.6	-1.2
Rise rate	-1.1	-0.4	-1.3	1.0	1.2	-1.5	-1.3	1.0	-0.3
Fall rate	0.8	2.3	3.9	1.6	1.4	2.8	0.8	-2.3)	-3.8

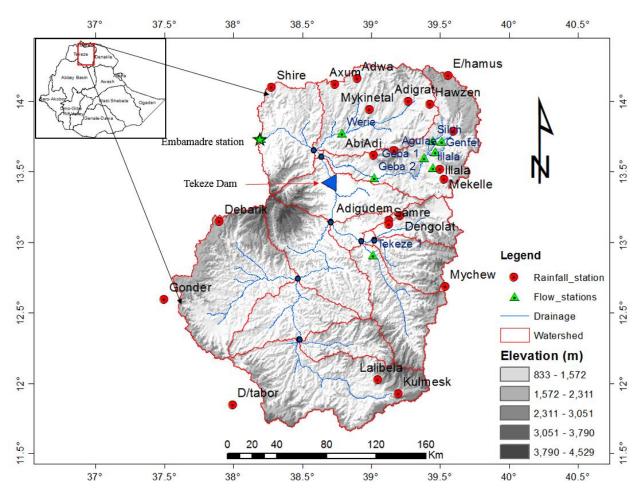


Figure 1. Location and distribution of rainfall and streamflow monitoring stations in the Upper Tekeze-Atbara River Basin.

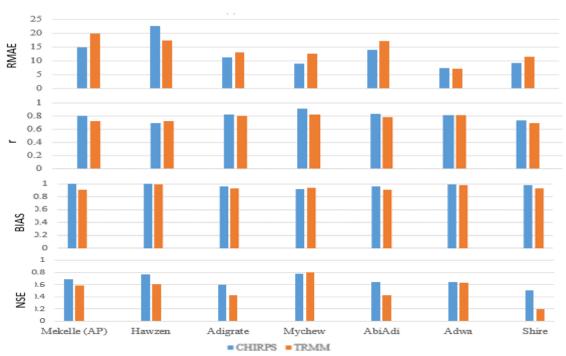


Figure 2. Comparison of satellite rainfall data against measured rainfall data

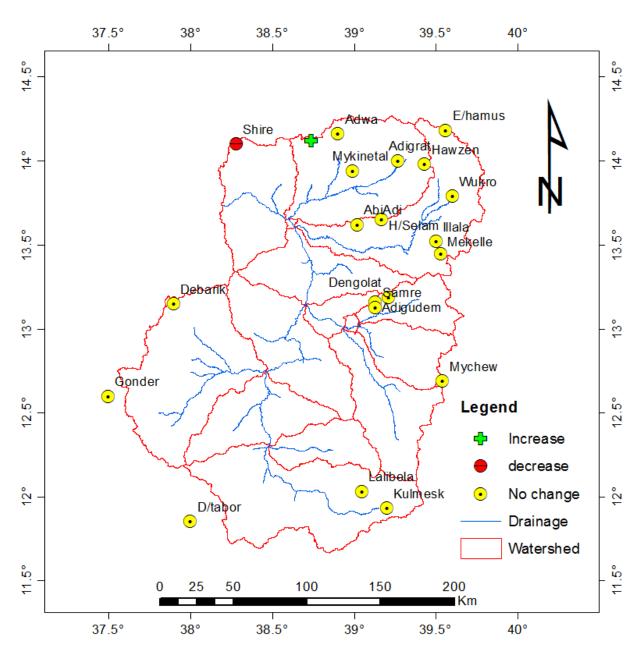


Figure 3. Location of rainfall stations with increasing (+), decreasing (-) and no change (0) trends on annual rainfall

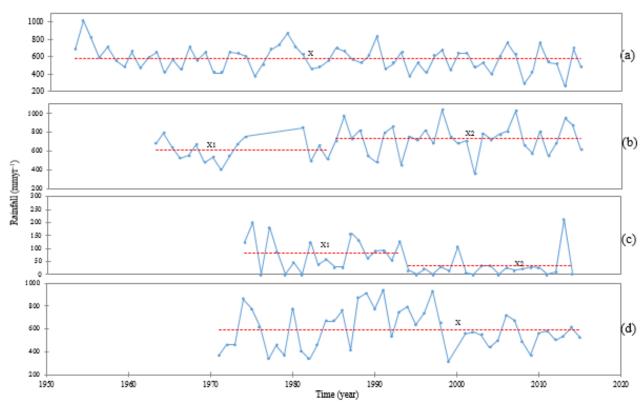


Figure 43. Pettitt homogeneity test of selected rainfall stations (a) annual rainfall in Mekelle (AP), (b) annual rainfall in Axum, (c) short rainy season in AbiAdi, (d) annual rainfall in Adigrat. X1 and X2 are average values of rainfall before and after change point.

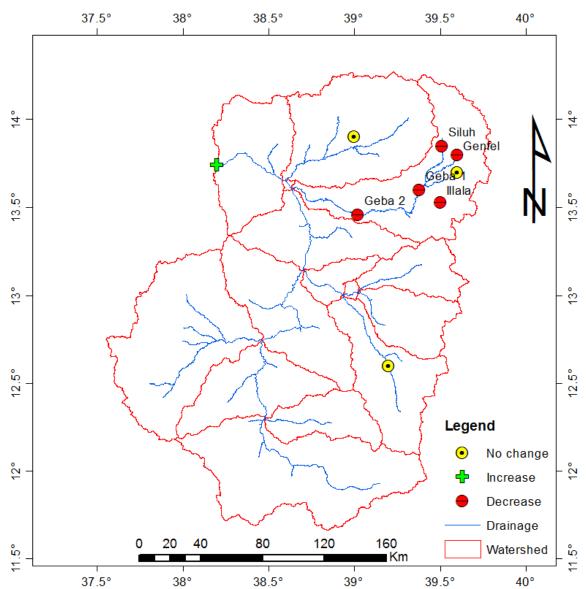


Figure 5. Location of streamflow stations with increasing (+), decreasing (-) and no change (0) trends on dry season flows

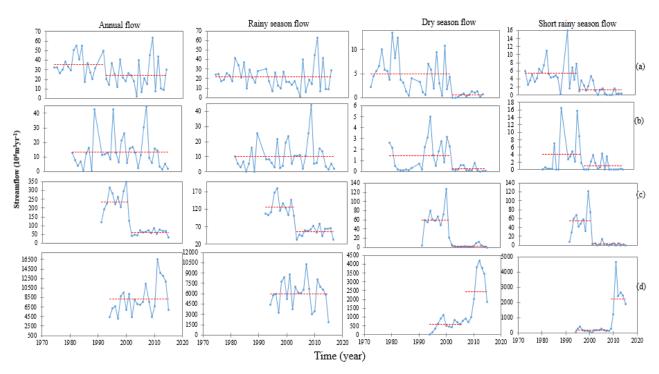


Figure-46. Annual and seasonal streamflow abrupt changes as determined by Pettitt test at (a) Siluh, (b) Illala, (c) Geba 1 and, (d) Embamadre