Developing a representative snow monitoring network in a forested mountain watershed

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Abstract. A challenge in establishing new ground-based stations for monitoring snowpack accumulation and ablation is to locate the sites in areas that represent the key processes affecting snow accumulation and ablation. This is especially challenging in forested montane watersheds where the combined effects of terrain, climate, and land cover affect seasonal snowpack. We present a coupled modelling approach used to objectively identify

10 representative snow monitoring locations in a forested watershed in the western Oregon Cascades mountain range. We used a binary regression tree (BRT) non-parametric statistical model to classify peak snow water equivalent (SWE) based on physiographic landscape characteristics in a average snow year, an above average snow year, and a below average snow year. Training data for the BRT classification were derived using spatially distributed estimates of SWE from a validated physically-based model of snow evolution. The optimal BRT

15 model showed that elevation and land cover type were the most significant drivers of spatial variability in peak SWE across the watershed ($R^2 = 0.93$, p-value < 0.01). Geospatial elevation and land cover data were used to map the BRT-derived snow classes across the watershed. Specific snow monitoring sites were selected randomly within the dominant BRT-derived snow classes to capture the range of spatial variability in snowpack conditions in the McKenzie River Basin. The Forest Elevational Snow Transect (ForEST) is a result of this coupled

20 modelling approach and represents combinations of forested and open land cover types at low, mid, and high elevations. After five years of snowpack monitoring, the ForEST network provides a valuable and detailed dataset of snow accumulation, snow ablation, and snowpack energy balance in forested and open sites from the rain-snow transition zone to the upper seasonal snow zone in the western Oregon Cascades.

1 Introduction

Mountain snowpack is declining as a result of the warming climate (Kunkel et al., 2016; Knowles, 2015; Pederson et al., 2013; Rupp et al., 2013; Pederson et al., 2011; Mote, 2006), subsequently shifting timing (Fritze et al., 2011; Clow, 2010) and volume of streamflow (Woodhouse et al., 2016; Berghuijs et al., 2014; Luce and Holden, 2009) across the western United States. Luce et al., (2013) argue that the declining snowpack is also the result of weakening westerlies leading to a decline in mountain precipitation in the interior West. The volume and seasonality of water produced from these snow-dominated watersheds varies spatially and temporally as a function of precipitation and temperature (Tennant et al., 2015; Barnett et al., 2005; Regonda et al., 2005), as well as local physiographic effects of topography, geology, and vegetation dynamics (Molotch and Meromy, 2014; Clark et al., 2011; Jefferson et al., 2008; Ffolliott et al., 1989).

Montane snow-dominated river basins are topographically complex. Elevation, slope, aspect, and exposure influence snowpack dynamics across a watershed through alterations of precipitation amount and phase (rain vs. snow), wind speed, temperature, and humidity. The degrees to which these physiographic variables control snow persistence vary as functions of snow accumulation and snow ablation, from the plot to regional spatial scales (López-Moreno et al., 2015; Biederman et al., 2014; López-Moreno et al., 2013; Deems et al., 2006; Molotch and Bales, 2005), and from daily to seasonal scales (Fassnacht et al., 2012; Jepsen et al., 2012). In the Pacific Northwest, montane basins are a successional patchwork of variable forest cover driven by forest harvest and replanting, pest infestations, and fire disturbance. In forested regions, snow accumulation and ablation processes are strongly influenced by vegetation structure (Veatch et al., 2009; Musselman et al., 2008; Jost et al., 2007; Trujillo et al., 2007; Sicart et al., 2004; Murray and Buttle, 2003; Pomeroy et al., 2002; Link and Marks, 1999). Both vegetation and topography influence the distribution of solar radiation (Musselman et al., 2015; Musselman et al., 2012; Davis et al., 1997; Dozier, 1980;), snow-surface albedo (Gleason and Nolin, 2016; Gleason et al., 2013; Molotch et al., 2004; Melloh et al., 2002), net longwave radiation (Lundquist et al., 2013; Sicart et al., 2004) , wind speed (Winstral and Marks, 2002) and turbulent fluxes (Burns et al., 2014; Garvelmann et al., 2014; Marks et al., 2008).

Snow water equivalent (SWE) is a critical hydrologic resource in the montane western US that has been actively monitored for decades by the Natural Resources Conservation Service (NRCS). The NRCS currently manages approximately 858 Snowpack Telemetry (SNOTEL) stations across the western US

(http://www.wcc.nrcs.usda.gov/snotel/SNOTEL_brochure.pdf). These stations provide near real-time measurements of SWE, temperature, and precipitation; essential data for operational streamflow forecasts used by water managers who balance a wide range of needs including irrigation, aquatic habitat, hydropower, recreation, and municipal water use. While most SNOTEL sites have been operating since the early 1980s, the data are meant to be used as indices to forecast discharge. These records are valuable but the stations were not designed to be nor are they representative of the total snow volume across a basin (Meromy et al., 2013; Molotch and Bales, 2006a). The SNOTEL monitoring stations in the Oregon Cascades are located within a narrow elevation range (1140–1510 m) that may not capture the inherent variability in the spatial distribution of snow under present day or warmer climate conditions (Nolin, 2012; Brown, 2009).

Modelling has been shown to be an effective means of augmenting remote sensing, and a valuable tool for predicting spatially distributed snow conditions in the rugged, forested, and frequently cloud-covered montane watersheds of the Pacific Northwest (Sproles et al., 2013; Tague and Grant, 2009; Veatch et al., 2009; Luce et al., 1999; Cline et al., 1998). Landscape characteristics have been used to predict snowpack conditions at hillslope scales using non-parametric binary regression tree (BRT) statistical classification models (Molotch et al., 2005; Anderton et al., 2004; Erxleben et al., 2002; Winstral et al., 2002; Balk and Elder, 2000; Elder et al., 1998). Larger scale BRT approaches have also been conducted using remotely sensed snow-covered area and interpolation methods (Molotch and Meromy, 2014; Molotch and Bales, 2006b). However, no study to date has used landscape characteristics in conjunction with modelled and validated physically-based and spatially distributed SWE data to understand physiographic drivers of snow accumulation at broad scales (watersheds > 1000 km^2) or to identify optimal locations for snowpack monitoring. Additionally, most of the research on the physiographic relationships to snow processes has been done in cold-dry continental snowpacks where mid-winter melt events are infrequent and wind redistribution is substantial (Molotch et al., 2005; Erxleben et al., 2002; Winstral et al., 2002; Balk and Elder, 2000). Much less is known about how physiographic conditions influence the temperature sensitive snowpacks in the forested maritime basins of the Pacific Northwest.

To objectively identify optimal site locations to distribute a snow monitoring network which explicitly captures the spatial variability of snow accumulation relative to the physiographic landscape we used a combination of physically-based, statistical, and geospatial models. This paper presents this objective and relatively simple methodology to distribute a snow monitoring network which captures landscape driven spatial variability in snow accumulation and includes four major objectives:

1. Determine the key physiographic drivers of spatial variability in snow accumulation;

2. Classify snow classes in the watershed based on key physiographic drivers using a non-parametric statistical model;

3. Spatially distribute these snow classes across the watershed using a geospatial model;

4. Select site locations for a snow monitoring network which spans the spatial variability in snow water equivalent in the McKenzie River Basin.

2 Methods

2.1 Study Site

The McKenzie River, located in the western Oregon Cascades, is a major tributary of the Willamette River (Figure 1). The McKenzie River Basin (MRB) drains an area of 3,041 km^2, and covers about 12% of the land area in the greater Willamette River Basin. The MRB is a densely forested mountainous watershed, ranging in elevation from 150 m to 3150 m, which is a managed for timber production throughout much of the seasonal snow zone. Brooks et al., (2012), determined that 60-80 % of summer flow in the Willamette River originated from elevations above 1200 m in the Oregon Cascades. The porous basalts in this geologically young landscape allow much of the snowmelt to percolate into groundwater systems (Tague and Grant, 2009; Jefferson et al., 2008; Tague and Grant, 2004). The groundwater-fed McKenzie River provides 25 % of the late season volumetric base flow to the Willamette River at its confluence with the Columbia River (Hulse et al., 2002).

2.2 Data Sources

Gridded data were obtained for physiographic variables shown in the literature to influence snow accumulation and ablation, including elevation, slope, aspect, incoming solar radiation, wind, and three vegetation variables from the following sources for the extent of the MRB. A Digital Elevation Model (DEM) was obtained from the National Elevation Dataset at a 10-m resolution. Slope, aspect, and incoming solar radiation were calculated from the DEM using the Spatial Analyst and Solar Radiation toolboxes in ArcGIS 10.1 (ESRI, Redlands, CA). Upwind contributing area data, which captures the variability in snow deposition as a result of wind redistribution for each cell throughout the watershed (Winstral et al., 2002), was calculated following Molotch et al., (2005). The 2006 National Land Cover Data (NLCD) was used to classify land cover across the watershed (Fry et al., 2011). Land cover data were reclassified into a binary product of forest and open land cover classes. The US Geological Survey (USGS) LANDFIRE Data Distribution Site provided the Existing

Vegetation – Percent Canopy Cover (EVC) data at 30-m spatial resolution. Normalized Difference Vegetation Index (NDVI) data were obtained from the Moderate Resolution Imaging SpectroRadiometer (MODIS) MOD13Q1 – Vegetation Indices, 16-day Land Product for the earliest date possible in April 2009, at a 250-m spatial resolution. Watershed boundaries were defined using the USGS National Hydrography Dataset. Public

5 land ownership data were provided by the Oregon Department of Forestry, and obtained from the website, http://www.oregon.gov/odf/pages/gis/gisdata.aspx. All spatial data were masked to the McKenzie River Basin and converted to the same projection and spatial resolution: NAD83, UTM Zone 10, and a 100-m grid cell size. Spatial data were processed using ArcGIS 10.1 using bilinear interpolation for continuous data and nearest neighbour interpolation for discrete data.

10 Modelled and gridded SWE data across the MRB (Figure 2) were provided by Sproles et al., (2013). These data were developed using a physically-based spatially distributed snow mass and energy balance model, SnowModel (Liston and Elder, 2006). SnowModel uses micrometeorological and topographic data to distribute snow across the landscape accounting for climatic, topographic, and vegetation variability. The model was modified by Sproles et al., (2013) to account for rain/snow precipitation phase partitioning, and snow albedo

15 decay in forested landscapes. This model was calibrated and validated using data from the four SNOTEL sites, meteorological data from the HJ Andrews Long Term Ecological Research site and National Weather Service stations, and Landsat fractional snow covered area data over the sampling period 1989-2009 (Sproles et al., 2013). The model was run at 100-m spatial resolution on a daily time step. We used modelled peak SWE data as the predicted variable in the BRT model. Sproles et al., (2013) showed that 2009 was considered an average snow

20 year (normal snow year) so we used peak SWE from 2009 (five days centred on 04 April 2009) as our reference year. Additionally we used peak SWE from 2008 (five days centred on 24 April 2008) as an above average snow year (high snow year), and peak SWE from 2005 (five days centred on 20 April 2005) as a below average snow year (low snow year).

2.3 Analysis

25 A BRT model was developed to characterize the spatial variability of snow accumulation across the MRB based on independent physiographic variables using the Classified and Regression Trees (CART) software (Salford Systems, San Diego, CA). The BRT model is a hierarchical non-parametric statistical model that characterizes the mean and variance of a dependent variable using a suite of independent explanatory variables. Modelled SWE and physiographic variable data were used as input data for each cell where snow was present

during peak SWE in 2009 (5 day average centred on 04 April 2009). An optimal tree was produced to minimize the standard error of the model, which was pruned to the simplest tree possible within one standard error of the optimal tree and so each terminal node represented at least 1% of the variability in peak SWE. The resultant tree identified 21 terminal nodes that characterized the spatial variability in snow accumulation through combinations of independent drivers into 21 BRT-derived snow classes (Table 1). The BRT model identified elevation, land cover, NDVI, insolation, percent canopy cover, slope, and wind as significant explanatory drivers of the spatial variability of peak SWE (all selected variables had p-values < 0.05 and are listed above in order of significance). Although elevation and land cover were the dominant predictive variables, and all other physiographic variables each explained less than 1% of the variability in peak SWE. In order to reduce the multi-collinearity between related variables and reduce the risk of overfitting the model, we simplified the final optimal model to only include elevation and land cover. Within the CART software, the final optimal BRT model was validated using reserved data from an independent set of 20,000 randomly selected grid cells from within the MRB. The final parameters developed in this optimal tree for peak SWE in a normal snow year (2009), were used to develop equivalent BRT models using peak SWE input for a high snow year (2008), as well as to peak SWE during a low snow year (2005).

Using a Geographic Information Systems (GIS) geospatial model and statistically-derived parameters, the 21 BRT-derived snow classes were spatially distributed across the MRB. The geospatial model used physiographic data to distribute the areal extent of each BRT class across the MRB by assigning cells that met the statistically-derived criteria for each BRT class. Because the BRT-model did not determine a lower elevation limit on snow extent, we excluded areas with an elevation less than 600 m to prevent over-prediction of snow-covered area below elevations where it was observed in the modelled data. Total volumetric SWE (SWE depth × area) was calculated for each BRT class across the watershed, using the mean and variance of SWE, and the spatial extent of each BRT class. To validate the spatial distribution of the BRT derived snow classes, we calculated the overall accuracy of the high snow year (2008) and low snow year (2005) relative to the reference year (2009) snow classes using an error matrix of omission vs. commission statistics (Campbell and Wynne, 2011).

To create a set of feasible locations for the *in situ* snow monitoring network we evaluated the accessibility of locations within the MRB. Using a GIS-based binary selection model, we masked out all private lands and public lands where the presence of endangered Northern Spotted Owl prevented permitted access. To prevent contamination from the road network, but still define accessible site locations, we also identified areas

within 100-500 m of a snowmobile-accessible road. From these accessible areas, the final sites were then randomly selected from each of the dominant BRT-derived snow classes within the seasonal snow zone.

3 Results

Modelled peak SWE for all years had a positively skewed distribution across the range of elevations throughout the MRB (2009; kurtosis=1.8, skewness=1.62; 2008, kurtosis=2.1, skewness=1.7; 2005, kurtosis=1.5, skewness=1.7) with the greatest volume of snow located in the mostly forested area between 1300 and 1500 m in elevation (Figure 3). The final optimal BRT model from the normal snow year (2009) identified elevation and land cover as the dominant drivers of the spatial variability of SWE, and characterized SWE across the MRB into 21 distinct snow classes (2009 BRT model; $R^2 = 0.93$, p-value < 0.01, RMSE = 0.16 m). The final BRT model applied to the high snow year (2008) characterized SWE across the MRB into 21 snow classes with similar spatial extent as during the normal snow year (2008 BRT model; $R^2 = 0.95$, p-value < 0.01, RMSE = 0.18 m) (Figure 4; Table S1). The final BRT model applied to the low snow year (2005) characterized SWE across the MRB into 21 snow classes with similar spatial variability relative to land cover but differing extents relative to elevation than during the normal snow year (2005 BRT model; $R^2 = 0.895$, p-value < 0.01, RMSE = 0.09 m) (Figure 4; Table S2).

Elevation explained the most variance in modelled SWE across the basin, and was the primary driver of all snow classes (2009 BRT model with only elevation; $R^2 = 0.91$, p-value < 0.01). In the middle elevations, land cover was also statistically important in distinguishing snow classes between forested and open land cover types (Figure 4, 5). During the normal snow year (2009), snow classes were distinguished by forest vs. open land cover types across the elevation range from 951-1442 m. This elevation range where the forest vs. open distinction was statistically important was lower during the high snow year (2008) from 949-1299 m, but much higher during the low snow year (2005) from 1193-1747 m. In the high-elevations, above treeline, only elevation was statistically important in classifying the spatial variability in snow accumulation. Snowpack accumulation increased with increasing elevation, resulting in a greater mean SWE per unit area at the highest elevations. Although deep snowpack at the highest elevations only covers a small aerial extent of the MRB, which resulted in decreasing contribution of total basin-wide SWE above approximately 1700 m during the normal and high snow years. In contrast, during the low snow year, the highest elevation classes contributed the most to total basin-wide SWE (Figure 5).

The BRT-derived volumetric SWE estimates had a similar positively skewed distribution across the elevational gradient as the SnowModel-derived SWE data in the MRB (Figure 5). The BRT-derived estimate of 1.49 km^3 total SWE stored in the snowpack on 04 April 2009 within the MRB was less than 1% greater than the SnowModel-derived estimate of 1.48 km^3. The BRT-derived estimate of 1.94 km^3 total SWE stored in the

5   snowpack on 24 April 2008 within the MRB was less than 1% greater than the SnowModel-derived estimate of 1.93 km^3. The BRT-derived estimate of 0.38 km^3 total SWE stored in the snowpack on 20 April 2005 within the MRB was 2.6% less than the SnowModel-derived estimate of 0.39 km^3. The final optimal BRT model from the normal snow year (2009) applied to the high snow year (2008) demonstrated an overall accuracy of 63%, whereas the BRT model applied to the low snow year (2005) demonstrated an overall accuracy of 26% (Table

10   S1, S2). The BRT model performed well across the low and high elevations, where errors of omission and commission were generally lowest (Table S1, S2). Although across the mid-elevations which consist of a patchwork of forest harvest and fire disturbance, were the areas with the greatest error between the BRT models. The high elevations above tree line, were the most consistently classified areas with low error between BRT models. The high error across the mid-elevations was due at least in part to the renumbering of classes when the

15   model is rerun for each year, and therefore these statistics may underrepresent the accuracy of the BRT-model in predicting overall spatial patterns of physiographically derived snow classes between years. The BRT-modelled snow classes captured the spatial variability in peak SWE across the MRB relative to elevation and land cover during an average, above average, and below average snow year and were used to objectively inform the site selection of a snow monitoring network.

20   The geospatial selection model identified 16 of the 21 classes as being accessible during winter (on public land without permit restrictions and within 100-500m of a snowmobile accessible road). The highest elevations in the MRB are far from winter-accessible roads and difficult to monitor due to steep and avalanche prone slopes. Within the area covered by these 16 classes, random site locations were selected within the six most abundant classes across the MRB to capture low, medium, and high elevations, with forested and open land cover

25   classes. The resultant Forest Elevation Snow Transect (ForEST) monitoring network site locations were thus objectively selected to sample across the range of spatial variability in SWE (Figure 4). The ForEST network, composed of six meteorological stations and snow survey transects, was deployed in November 2011, and continues to provide high quality snow and climate data to evaluate snow-forest-climate interactions in the MRB (Figure 6).

4 Discussion

With warming winter temperatures, mountain snowpack in the western US will likely continue to decline with potential impacts to forest health (Albright and Peterson, 2013) and streamflow (Jung and Chang, 2011; Cayan et al., 2010), as well as snow-related recreation and tourism (Gilaberte-Búrdalo et al., 2014; Nolin and Daly, 2006). There remains uncertainty around the magnitude of these impacts (Warren et al., 2011; Maurer, 2007; Xu et al., 2005) thus, it is important that monitoring networks not only capture normal snowpack conditions, but capture the range of variability in peak SWE across the landscape and through time.

Pacific Northwest forests play a key role in affecting snow accumulation and ablation across multiple scales however, most research has been conducted at the stand scale (Storck et al., 2002) or in areas with cold-dry continental snowpacks (Ellis et al., 2013; Pomeroy et al., 2012). By distinguishing snow classes based on forest vs. open land cover across a range of elevations, this study emphasizes the watershed-scale control that vegetation and particularly land cover change relative to timber harvest (and potentially fire disturbance) has on snowpack accumulation in the maritime western Oregon Cascades. During low snow years, the significant influence of forest cover on the spatial variability in snow accumulation moved up in elevation from a normal snow year, suggesting that forest effects may have a more profound influence at higher elevations under future warming climate conditions. Understanding the forest structure effects on snow accumulation and ablation across elevation gradients is increasingly important to help guide decision making by local and regional water and forest managers in response to a changing climate.

We developed a snow monitoring network representative of the spatial variability of peak SWE relative to physiographic landscape characteristics across the MRB for an average, above average, and below average snow year, by coupling a spatially distributed physically-based SnowModel, a BRT statistical classification model, and a geospatial selection model. This objective method is a useful tool in classifying snow characteristics across the landscape to determine representative locations for intelligent snowpack monitoring particularly in physiographically complex landscapes. Although it is an improvement over more commonly used heuristic approaches to site selection, the method incorporates uncertainty as a result of compounding physically-, statistically-, and spatially-based models which justifies caution in implementing these estimates in management decisions. However, the method meets assumptions of non-parametric data analysis, is performed with relative ease, and if data are available for the research basin of interest, it can be well validated. As even physically-based models incorporate inherent empirically-based historically-derived assumptions, there is also uncertainty in using this approach to represent future spatial variability in snow accumulation.

The ForEST network contributes to the existing SNOTEL network to explicitly investigate snow-vegetation-climate interactions across the range of elevations and forest types in the watershed. The ForEST network is unique in that the monitoring site locations were selected based on statistical classification and geospatial analysis, rather than subjective methods that may incorporate bias. The paired forest-open land cover site selection process has already led to important understanding of key sub-canopy snow processes (Storck et al., 2002; Golding and Swanson, 1986). But here, the assumptions driving paired site selection process have been validated using coupled physically-based spatially-distributed snow model input data and non-parametric BRT statistical modelling across a forested montane watershed. After five consecutive years of snow monitoring, we have created a valuable and detailed dataset of snow accumulation, snow ablation, and snowpack energy balance that spans the spatial variability in forest and open land cover types across an elevational gradient (Figure 6).

5 Conclusions

The BRT model characterized peak SWE conditions in an average year, an above average year, and below average year to provide spatially-distributed SWE volume estimates based on physiographic landscape characteristics. This integrated approach informed the distribution of an objective and representative monitoring network that spans the spatial variability in the seasonal snowpack across the MRB (Figure 4). Throughout the maritime Pacific Northwest, it is critical we monitor snow-vegetation interactions across the elevation gradient, particularly at higher elevations where snow-vegetation interactions may be more relevant in low snow years and under a warming climate.

By quantifying the spatial variability in the key drivers of natural resource distribution, researchers can focus on sensitive areas which may not be identified through traditional site selection means. The use of validated model outputs as a predictor of the spatial variability in snow-vegetation interactions is not new (Randin et al., 2014). The novelty of this research stems from the application of the method, where by the coupling of a traditional BRT classification process with a validated physically-based spatially distributed model, we improved snow observational network design in a forested montane watershed.

As the scientific community turns to more complex models to predict ecosystem responses to change, there is still a place for simple modelling approaches to inform scientific research priorities as well as natural resource monitoring and management. Particularly in rugged and densely forested mountain regions, such as the western Oregon Cascades, where there are few alternatives to modelling spatially distributed SWE, this coupled

modelling approach provides a validated hypothesis to guide representative and objective snow monitoring efforts.

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7 References

Albright, W. L., and Peterson, D. L.: Tree growth and climate in the Pacific Northwest, North America: a broad-scale analysis of changing growth environments, Journal of Biogeography, 40, 2119-2133, 2013.

Anderton, S. P., White, S. M., and Alvera, B.: Evaluation of spatial variability in snow water equivalent for a high mountain catchment, Hydrological Processes, 18, 435-453, 10.1002/hyp.1319, 2004.

Balk, B., and Elder, K.: Combining binary decision tree and geostatistical methods to estimate snow distribution in a mountain watershed, Water Resources Research, 36, 13-26, 2000.

Barnett, T. P., Adam, J. C., and Lettenmaier, D. P.: Potential impacts of a warming climate on water availability in snow-dominated regions, Nature, 438, 303-309, 2005.

Berghuijs, W. R., Woods, R. A., and Hrachowitz, M.: A precipitation shift from snow towards rain leads to a decrease in streamflow, Nature Clim. Change, 4, 583-586, 2014.

Biederman, J. A., Brooks, P., Harpold, A., Gochis, D., Gutmann, E., Reed, D., Pendall, E., and Ewers, B.: Multiscale observations of snow accumulation and peak snowpack following widespread, insect-induced lodgepole pine mortality, Ecohydrology, 7, 150-162, 2014.

Brooks, J. R., Wigington, P. J., Phillips, D. L., Comeleo, R., and Coulombe, R.: Willamette River Basin surface water isoscape (δ18O and δ2H): temporal changes of source water within the river, Ecosphere, 3, 1-21, 2012.

Brown, A. L.: Understanding the impact of climate change on snowpack extent and measurement in the Columbia River Basin and nested sub basins, Master's Thesis, Oregon State University, 2009.

Burns, S. P., Molotch, N. P., Williams, M. W., Knowles, J. F., Seok, B., Monson, R. K., Turnipseed, A. A., and Blanken, P. D.: Snow temperature changes within a seasonal snowpack and their relationship to turbulent fluxes of sensible and latent heat, Journal of Hydrometeorology, 15, 117-142, 10.1175/jhm-d-13-026.1, 2014.

Campbell, J. B., and Wynne, R. H.: Introduction to remote sensing. Guilford Press, 2011.

Cayan, D. R., Das, T., Pierce, D. W., Barnett, T. P., Tyree, M., and Gershunov, A.: Future dryness in the southwest US and the hydrology of the early 21st century drought, Proceedings of the National Academy of Sciences, 107, 21271-21276, 2010.

Clark, M. P., Hendrikx, J., Slater, A. G., Kavetski, D., Anderson, B., Cullen, N. J., Kerr, T., Örn Hreinsson, E., and Woods, R. A.: Representing spatial variability of snow water equivalent in hydrologic and land-surface models: A review, Water Resources Research, 47, n/a-n/a, 10.1029/2011WR010745, 2011.

Cline, D. W., Bales, R. C., and Dozier, J.: Estimating the spatial distribution of snow in mountain basins using remote sensing and energy balance modeling, Water Resources Research, 34, 1275-1285, 1998.

Clow, D. W.: Changes in the timing of snowmelt and streamflow in Colorado: a response to recent warming, Journal of Climate, 23, 2293-2306, 2010.

Davis, R. E., Hardy, J. P., Ni, W., Woodcock, J., McKenzie, J. C., Jordan, R., and Li, X.: Variation of snow cover ablation in the boreal forest: A sensitivity study on the effects of conifer canopy, Journal of Geophysical Research. D. Atmospheres, 102, 29, 1997.

Deems, J. S., Fassnacht, S. R., and Elder, K. J.: Fractal distribution of snow depth from lidar data, Journal of Hydrometeorology, 7, 285-297, 2006.

Dozier, J.: A Clear-Sky Spectral Solar Radiation Model, Water Resources Research, 16, 709-718, 1980.

Elder, K., Rosenthal, W., and Davis, R. E.: Estimating the spatial distribution of snow water equivalence in a montane watershed, Hydrological Processes, 12, 1793-1808, 1998.

Ellis, C. R., Pomeroy, J. W., and Link, T. E.: Modeling increases in snowmelt yield and desynchronization resulting from forest gap-thinning treatments in a northern mountain headwater basin, Water Resources Research, 49, 936-949, 10.1002/wrcr.20089, 2013.

Erxleben, J., Elder, K., and Davis, R.: Comparison of spatial interpolation methods for estimating snow distribution in the Colorado Rocky Mountains, Hydrological Processes, 16, 3627-3649, 10.1002/hyp.1239, 2002.

Fassnacht, S., Dressler, K., Hultstrand, D., Bales, R., and Patterson, G.: Temporal inconsistencies in coarse-scale snow water equivalent patterns: Colorado River Basin snow telemetry-topography regressions, Pirineos, 2012.

5   Ffolliott, P. F., Gottfried, G. J., and Jr, M. B. B.: Water yield from forest snowpack management: research findings in Arizona and New Mexico, Water Resources Research, 25, 1999-2007, 1989.

Fritze, H., Stewart, I. T., and Pebesma, E.: Shifts in western North American snowmelt runoff regimes for the recent warm decades, Journal of Hydrometeorology, 12, 989-1006, 2011.

Fry, J. A., Xian, G., Jin, S., Dewitz, J. A., Homer, C. G., Limin, Y., Barnes, C. A., Herold, N. D., and Wickham,

10   J. D.: Completion of the 2006 national land cover database for the conterminous United States, Photogrammetric engineering and remote sensing, 77, 858-864, 2011.

Garvelmann, J., Pohl, S., and Weiler, M.: Variability of observed energy fluxes during rain-on-snow and clear sky snowmelt in a midlatitude mountain environment, Journal of Hydrometeorology, 15, 1220-1237, 10.1175/jhm-d-13-0187.1, 2014.

15   Gilaberte-Búrdalo, M., López-Martín, F., Pino-Otín, M., and López-Moreno, J. I.: Impacts of climate change on ski industry, Environmental Science & Policy, 44, 51-61, 2014.

Gleason, K. E., Nolin, A. W., and Roth, T. R.: Charred forests increase snowmelt: Effects of burned woody debris and incoming solar radiation on snow ablation, Geophysical Research Letters, 40, 4654-4661, 10.1002/grl.50896, 2013.

20   Gleason, K. E., and Nolin, A. W.: Charred forests accelerate snow albedo decay: parameterizing the post-fire radiative forcing on snow for three years following fire, Hydrological Processes, 2016.

Golding, D. L., and Swanson, R. H.: Snow distribution patterns in clearings and adjacent forest, Water Resources Research, 22, 1931-1940, 1986.

Hulse, D. W., Gregory, S., and Baker, J.: Willamette River Basin Atlas: Trajectories of Environmental and

25   Ecological Change, Oregon State University Press, 192 pp., 2002.

Jefferson, A., Nolin, A., Lewis, S., and Tague, C.: Hydrogeologic controls on streamflow sensitivity to climate variation, Hydrological Processes, 22, 4371-4385, 10.1002/hyp.7041, 2008.

Jepsen, S. M., Molotch, N. P., Williams, M. W., Rittger, K. E., and Sickman, J. O.: Interannual variability of snowmelt in the Sierra Nevada and Rocky Mountains, United States: Examples from two alpine watersheds, Water Resources Research, 48, 2012.

Jost, G., Weiler, M., Gluns, D. R., and Alila, Y.: The influence of forest and topography on snow accumulation
5 and melt at the watershed-scale, Journal of Hydrology, 347, 101-115, 10.1016/j.jhydrot.2007.09.006, 2007.

Jung, I. W., and Chang, H.: Assessment of future runoff trends under multiple climate change scenarios in the Willamette River Basin, Oregon, USA, Hydrological Processes, 25, 258-277, 2011.

Knowles, N.: Trends in Snow Cover and Related Quantities at Weather Stations in the Conterminous United States, Journal of Climate, 28, 7518-7528, 2015.

10 Kunkel, K. E., Robinson, D. A., Champion, S., Yin, X., Estilow, T., and Frankson, R. M.: Trends and extremes in northern hemisphere snow characteristics, Current Climate Change Reports, 2, 65-73, 2016.

Link, T., and Marks, D.: Distributed simulation of snowcover mass- and energy-balance in the boreal forest, Hydrological Processes, 13, 2439-+, 10.1002/(sici)1099-1085(199910)13:14/15<2439::aid-hyp866>3.0.co;2-1, 1999.

15 Liston, G. E., and Elder, K.: A distributed snow-evolution modeling system (SnowModel), Journal of Hydrometeorology, 7, 1259-1276, 2006.

López-Moreno, J. I., Fassnacht, S., Heath, J., Musselman, K., Revuelto, J., Latron, J., Morán-Tejeda, E., and Jonas, T.: Small scale spatial variability of snow density and depth over complex alpine terrain: Implications for estimating snow water equivalent, Advances in Water Resources, 55, 40-52, 2013.

20 López-Moreno, J. I., Revuelto, J., Fassnacht, S., Azorín-Molina, C., Vicente-Serrano, S. M., Morán-Tejeda, E., and Sexstone, G.: Snowpack variability across various spatio-temporal resolutions, Hydrological Processes, 29, 1213-1224, 2015.

Luce, C. H., Tarboton, D. G., and Cooley, K. R.: Sub-grid parameterization of snow distribution for an energy and mass balance snow cover model, Hydrological Processes, 13, 1921-1933, 1999.

25 Luce, C. H., and Holden, Z. A.: Declining annual streamflow distributions in the Pacific Northwest United States, 1948–2006, Geophysical Research Letters, 36, n/a-n/a, 10.1029/2009GL039407, 2009.

Luce, C. H., Abatzoglou, J. T., and Holden, Z. A.: The missing mountain water: Slower westerlies decrease orographic enhancement in the Pacific Northwest USA, Science, 10.1126/science.1242335, 2013.

Lundquist, J. D., Dickerson-Lange, S. E., Lutz, J. A., and Cristea, N. C.: Lower forest density enhances snow retention in regions with warmer winters: A global framework developed from plot-scale observations and modeling, Water Resources Research, 49, 6356-6370, 10.1002/wrcr.20504, 2013.

Marks, D., Reba, M., Pomeroy, J., Link, T., Winstral, A., Flerchinger, G., and Elder, K.: Comparing simulated and measured sensible and latent heat fluxes over snow under a pine canopy to improve an energy balance snowmelt model, Journal of Hydrometeorology, 9, 1506-1522, 10.1175/2008jhm874.1, 2008.

Maurer, E. P.: Uncertainty in hydrologic impacts of climate change in the Sierra Nevada, California, under two emissions scenarios, Climatic Change, 82, 309-325, 2007.

Melloh, R. A., Hardy, J. P., Bailey, R. N., and Hall, T. J.: An efficient snow albedo model for the open and sub canopy, Hydrological Processes, 16, 3571-3584, 2002.

Meromy, L., Molotch, N. P., Link, T. E., Fassnacht, S. R., and Rice, R.: Subgrid variability of snow water equivalent at operational snow stations in the western USA, Hydrological Processes, 27, 2383-2400, 2013.

Molotch, N., Colee, M., Bales, R., and Dozier, J.: Estimating the spatial distribution of snow water equivalent in an alpine basin using binary regression tree models: the impact of digital elevation data and independent variable selection, Hydrological Processes, 19, 1459-1479, 2005.

Molotch, N. P., Painter, T. H., Bales, R. C., and Dozier, J.: Incorporating remotely-sensed snow albedo into a spatially-distributed snowmelt model, Geophysical Research Letters, 31, 10.1029/2003gl019063, 2004.

Molotch, N. P., and Bales, R. C.: Scaling snow observations from the point to the grid element: Implications for observation network design, Water Resources Research, 41, W11421, 10.1029/2005wr004229, 2005.

Molotch, N. P., and Bales, R. C.: SNOTEL representativeness in the Rio Grande headwaters on the basis of physiographics and remotely sensed snow cover persistence, Hydrological Processes, 20, 723-739, 10.1002/hyp.6128, 2006a.

Molotch, N. P., and Bales, R. C.: Comparison of ground-based and airborne snow surface albedo parameterizations in an alpine watershed: Impact on snowpack mass balance, Water Resources Research, 42, 10.1029/2005wr004522, 2006b.

Molotch, N. P., and Meromy, L.: Physiographic and climatic controls on snow cover persistence in the Sierra Nevada Mountains, Hydrological Processes, 28, 4573-4586, 2014.

Mote, P. W.: Climate-Driven Variability and Trends in Mountain Snowpack in Western North America*, Journal of Climate, 19, 6209-6220, 2006.

5 Murray, C. D., and Buttle, J. M.: Impacts of clearcut harvesting on snow accumulation and melt in a northern hardwood forest, Journal of Hydrology, 271, 197-212, 2003.

Musselman, K. N., Molotch, N. P., and Brooks, P. D.: Effects of vegetation on snow accumulation and ablation in a mid-latitude sub-alpine forest, Hydrological Processes, 22, 2767-2776, 10.1002/hyp.7050, 2008.

Musselman, K. N., Molotch, N. P., Margulis, S. A., Kirchner, P. B., and Bales, R. C.: Influence of canopy
10 structure and direct beam solar irradiance on snowmelt rates in a mixed conifer forest, Agricultural and Forest Meteorology, 161, 46-56, 10.1016/j.agrformet.2012.03.011, 2012.

Musselman, K. N., Pomeroy, J. W., and Link, T. E.: Variability in shortwave irradiance caused by forest gaps: Measurements, modelling, and implications for snow energetics, Agricultural and Forest Meteorology, 207, 69-82, 2015.

15 Nolin, A. W., and Daly, C.: Mapping at risk" snow in the Pacific Northwest, Journal of Hydrometeorology, 7, 1164-1171, 2006.

Nolin, A. W.: Perspectives on Climate Change, Mountain Hydrology, and Water Resources in the Oregon Cascades, USA, Mountain Research and Development, 32, S35-S46, 10.1659/mrd-journal-d-11-00038.s1, 2012.

Pederson, G. T., Gray, S. T., Woodhouse, C. A., Betancourt, J. L., Fagre, D. B., Littell, J. S., Watson, E.,
20 Luckman, B. H., and Graumlich, L. J.: The Unusual Nature of Recent Snowpack Declines in the North American Cordillera, Science, 333, 332-335, 10.1126/science.1201570, 2011.

Pederson, G. T., Betancourt, J. L., and McCabe, G. J.: Regional patterns and proximal causes of the recent snowpack decline in the Rocky Mountains, US, Geophysical Research Letters, 1-6, 2013.

Pomeroy, J., Fang, X., and Ellis, C.: Sensitivity of snowmelt hydrology in Marmot Creek, Alberta, to forest cover
25 disturbance, Hydrological Processes, 26, 1892-1905, 10.1002/hyp.9248, 2012.

Pomeroy, J. W., Gray, D. M., Hedstrom, N. R., and Janowicz, J. R.: Prediction of seasonal snow accumulation in cold climate forests, Hydrological Processes, 16, 3543-3558, 10.1002/hyp.1228, 2002.

Randin, C. F., Dedieu, J. P., Zappa, M., Long, L., and Dullinger, S.: Validation of and comparison between a semidistributed rainfall–runoff hydrological model (PREVAH) and a spatially distributed snow-evolution model (SnowModel) for snow cover prediction in mountain ecosystems, Ecohydrology, 2014.

Regonda, S. K., Rajagopalan, B., Clark, M., and Pitlick, J.: Seasonal cycle shifts in hydroclimatology over the western United States, Journal of Climate, 18, 372-384, 2005.

Rupp, D. E., Mote, P. W., Bindoff, N. L., Stott, P. A., and Robinson, D. A.: Detection and attribution of observed changes in Northern Hemisphere spring snow cover, Journal of Climate, 26, 6904-6914, 2013.

Serreze, M. C., Clark, M. P., Armstrong, R. L., McGinnis, D. A., and Pulwarty, R. S.: Characteristics of the western United States snowpack from snowpack telemetry (SNOTEL) data, Water Resources Research, 35, 2145-2160, 10.1029/1999wr900090, 1999.

Sicart, J. E., Pomeroy, J. W., Essery, R. L. H., Hardy, J., Link, T., and Marks, D.: A sensitivity study of daytime net radiation during snowmelt to forest canopy and atmospheric conditions, Journal of Hydrometeorology, 5, 774-784, 2004.

Sproles, E. A., Nolin, A. W., Rittger, K., and Painter, T. H.: Climate change impacts on maritime mountain snowpack in the Oregon Cascades, Hydrology and Earth System Sciences, 17, 2581-2597, 10.5194/hess-17-2581-2013, 2013.

Stewart, I. T., Cayan, D. R., and Dettinger, M. D.: Changes in snowmelt runoff timing in western North America under a 'business as usual' climate change scenario, Climatic Change, 62, 217-232, 2004.

Storck, P., Lettenmaier, D. P., and Bolton, S. M.: Measurement of snow interception and canopy effects on snow accumulation and melt in a mountainous maritime climate, Oregon, United States, Water Resources Research, 38, 1223 10.1029/2002wr001281, 2002.

Tague, C., and Grant, G. E.: A geological framework for interpreting the low-flow regimes of Cascade streams, Willamette River Basin, Oregon, Water Resources Research, 40, W0430310.1029/2003wr002629, 2004.

Tague, C., and Grant, G. E.: Groundwater dynamics mediate low-flow response to global warming in snow-dominated alpine regions, Water Resources Research, 45, W0742110.1029/2008wr007179, 2009.

Tennant, C. J., Crosby, B. T., and Godsey, S. E.: Elevation-dependent responses of streamflow to climate warming, Hydrological Processes, 29, 991-1001, 2015.

Trujillo, E., Ramírez, J. A., and Elder, K. J.: Topographic, meteorologic, and canopy controls on the scaling characteristics of the spatial distribution of snow depth fields, Water Resources Research, 43, W07409, 2007.

Veatch, W., Brooks, P. D., Gustafson, J. R., and Molotch, N. P.: 'Quantifying the effects of forest canopy cover on net snow accumulation at a continental, mid-latitude site', Ecohydrology, 2, 115-128, 10.1002/eco.45, 2009.

5   Warren, R., Price, J., Fischlin, A., de la Nava Santos, S., and Midgley, G.: Increasing impacts of climate change upon ecosystems with increasing global mean temperature rise, Climatic Change, 106, 141-177, 2011.

Winstral, A., Elder, K., and Davis, R. E.: Spatial snow modeling of wind-redistributed snow using terrain-based parameters, Journal of Hydrometeorology, 3, 524-538, 2002.

Winstral, A., and Marks, D.: Simulating wind fields and snow redistribution using terrain-based parameters to
10   model snow accumulation and melt over a semi-arid mountain catchment, Hydrological Processes, 16, 3585-3603, 10.1002/hyp.1238, 2002.

Woodhouse, C. A., Pederson, G. T., Morino, K., McAfee, S. A., and McCabe, G. J.: Increasing influence of air temperature on upper Colorado River streamflow, Geophysical Research Letters, 43, 2174-2181, 2016.

Xu, C.-y., Widén, E., and Halldin, S.: Modelling hydrological consequences of climate change—progress and
15   challenges, Advances in Atmospheric Sciences, 22, 789-797, 2005.

8 Tables

Table 1. The binary regression tree (BRT) model characterized SWE within the McKenzie River Basin into 21 snow classes defined using the physiographic parameters elevation (m) and land cover grouped by all land covers (A), forested land cover (F), and open/clear-cut land cover (O) types. The bold lines represent the BRT snow classes used for the ForEST network of snow monitoring stations which have been continuously monitoring snow processes in paired forest and open sites at low, medium and high elevations since November 2011.

BRT Snow class	Normal Snow Year (2009)			High Snow Year (2008)			Low Snow Year (2005)		
	Elevation	Land	Mean SWE / St. Dev.	Elevation	Land	Mean SWE / St. Dev.	Elevation	Land	Mean SWE / St. Dev.
1	≤ 834	A	0.002 / 0.2	≤865	A	0.002 / 0.03	**≤1193**	**A**	**0.001 / 0.01**
2	834-951	A	0.04 / 0.12	865-949	A	0.06 / 0.16	1193-1299	F	0.02 / 0.07
3	951-1067	F	0.11 / 0.21	949-1067	F	0.14 / 0.26	1193-1299	O	0.08 / 0.15
4	951-1067	O	0.24 / 0.28	949-1067	O	0.31 / 0.36	**1299-1422**	**F**	**0.04 / 0.12**
5	**1067-1142**	**F**	**0.26 / 0.28**	**1067-1115**	**F**	**0.31 / 0.34**	1422-1470	F	0.09 / 0.16
6	**1067-1142**	**O**	**0.45 / 0.35**	1115-1142	F	0.45 / 0.36	**1299-1390**	**O**	**0.16 / 0.22**
7	1142-1174	F	0.41 / 0.3	1067-1142	**O**	**0.59 / 0.43**	1390-1470	O	0.29 / 0.28
8	1174-1212	F	0.51 / 0.3	1142-1200	F	0.64 / 0.36	**1470-1496**	**F**	**0.16 / 0.18**
9	1142-1212	O	0.68 / 0.34	1142-1200	O	0.91 / 0.4	1496-1517	F	0.23 / 0.19
10	1212-1265	F	0.64 / 0.29	1200-1235	F	0.83 / 0.33	1517-1536	F	0.28 / 0.19
11	1265-1310	F	0.79 / 0.27	1235-1299	F	1.0 / 0.3	1536-1563	F	0.34 / 0.2
12	1212-1310	O	0.95 / 0.31	1200-1299	O	1.21 / 0.34	**1470-1563**	**O**	**0.54 / 0.29**
13	**1310-1364**	**A**	**0.96 / 0.24**	**1299-1338**	**A**	**1.23 / 0.27**	1563-1620	F	0.44 / 0.19
14	1364-1442	F	1.07 / 0.58	1338-1385	A	1.37 / 0.21	1620-1663	F	0.59 / 0.18
15	1364-1442	O	1.18 / 0.17	1385-1445	A	1.49 / 0.14	1663-1713	F	0.63 / 0.19
16	**1442-1486**	**A**	**1.28 / 0.13**	**1445-1563**	**A**	**1.62 / 0.1**	1713-1747	F	0.75 / 0.19
17	1486-1563	A	1.24 / 0.12	1563-1779	A	1.73 / 0.09	1563-1747	O	0.84 / 0.19
18	1563-1779	A	1.31 / 0.09	1779-1931	A	1.89 / 0.1	1747-1787	A	0.91 / 0.19
19	1779-1910	A	1.43 / 0.09	1931-2016	A	2.17 / 0.15	1787-1866	A	1.04 / 0.12
20	1910-2101	A	1.64 / 0.1	2016-2226	A	2.38 / 0.74	1866-2152	A	1.14 / 0.05
21	>2101	A	1.89 / 0.09	>2226	A	2.66 / 0.12	>2152	A	1.29 / 0.05

9 Figures
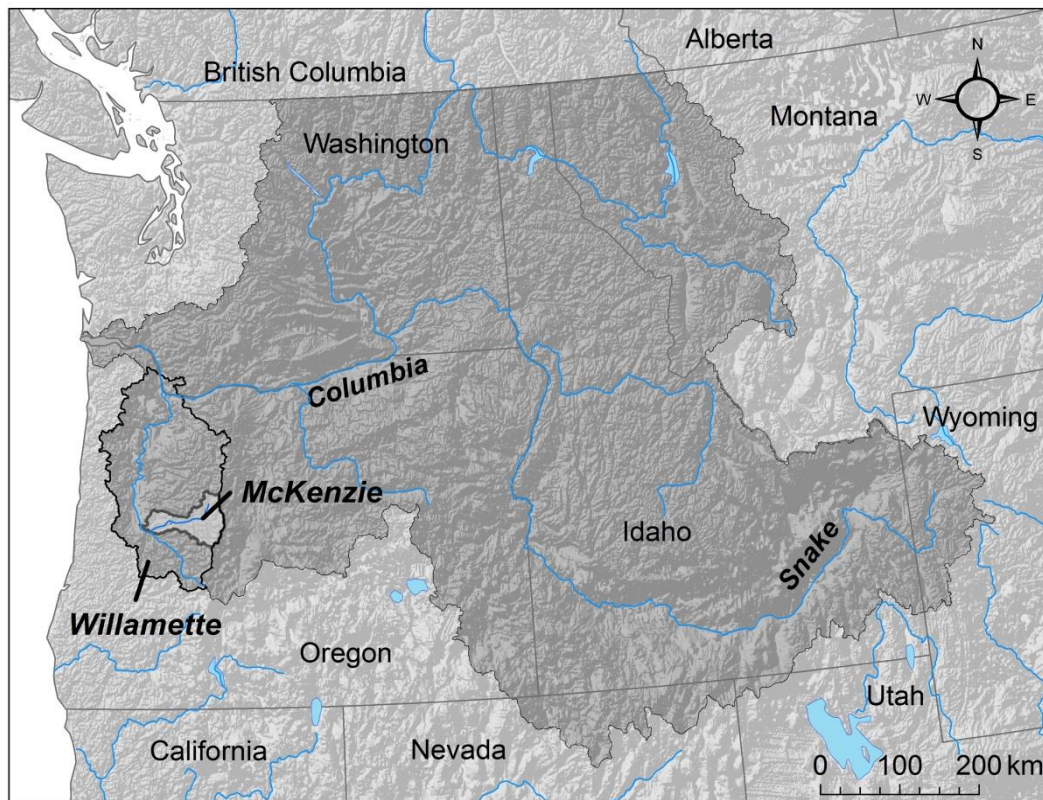
Figure 1. McKenzie River Basin is nested in the Willamette River Basin within the greater Columbia River Basin.
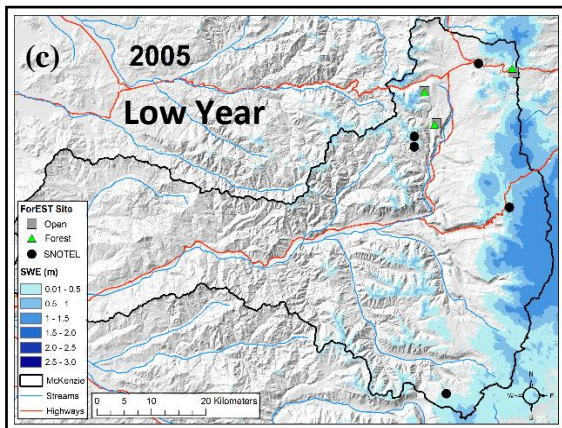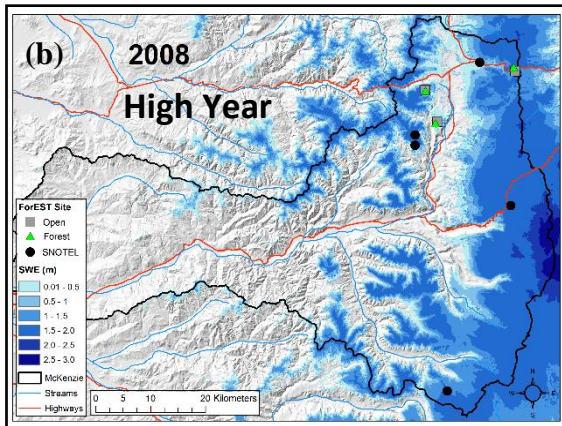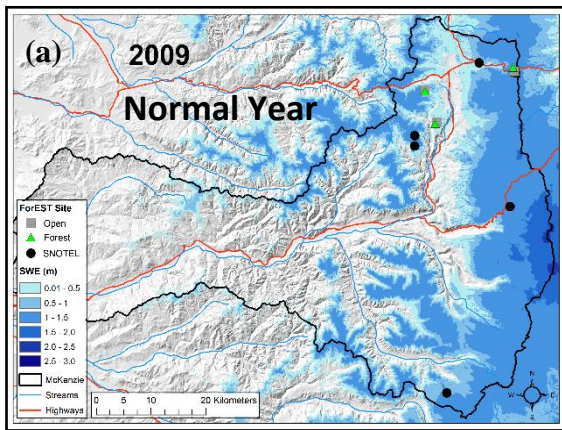
Figure 2. SnowModel-derived snow water equivalent (SWE) (m) is shown in blue for the date of peak SWE from Sproles *et al.,* (2013) for, a) a normal snow year (2009), b) a high snow year (2008), and c) a low snow year (2005), for the modelling domain around the McKenzie River Basin. The locations of the current SNOTEL sites are shown in black circles. The locations of the ForEST sites are shown in grey squares for open sites and green triangles for forested sites.
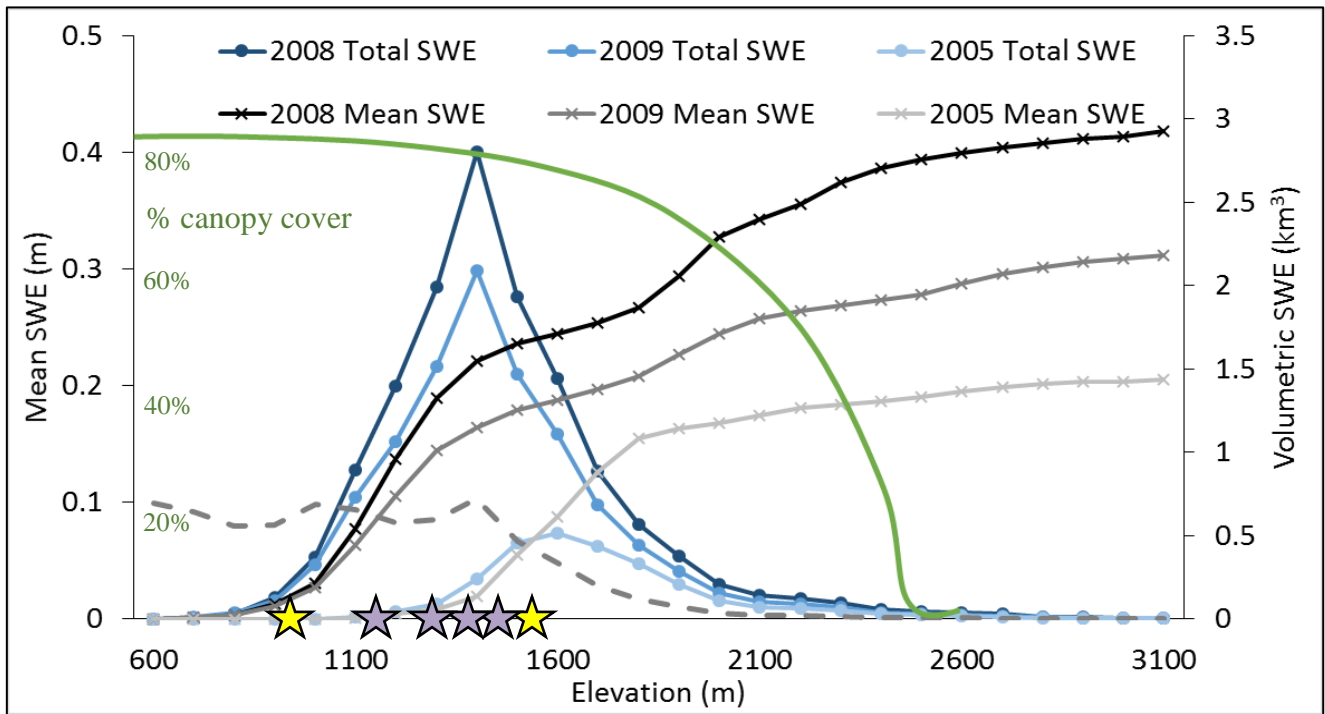
Figure 3. Elevation distribution (using 100 m elevation bands) of SnowModel-derived snow water equivalent (SWE) in the McKenzie River Basin for peak SWE during an average snow year (04 April 2009), an above average snow year (24 April 2008), and a below average snow year (20 April 2005). Mean SWE for each elevation band is shown in greyscale, and total basin wide volumetric SWE is shown in blue scale for 2009, 2008, and 2005. The green line indicates the elevation distribution of the percent canopy cover. The dashed grey line indicates the % / 100 of the area represented by each 100 m elevation band. The area of the greatest volumetric SWE persists in a narrow elevation range which is monitored by four historical and two newly installed (as of 2012) SNOTEL stations (elevations of historical stations shown in purple stars and new stations in yellow stars).
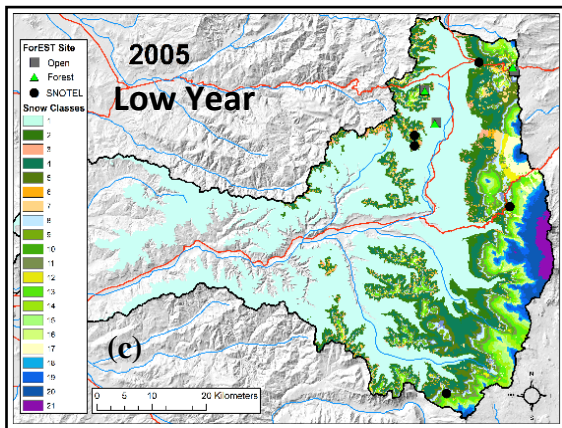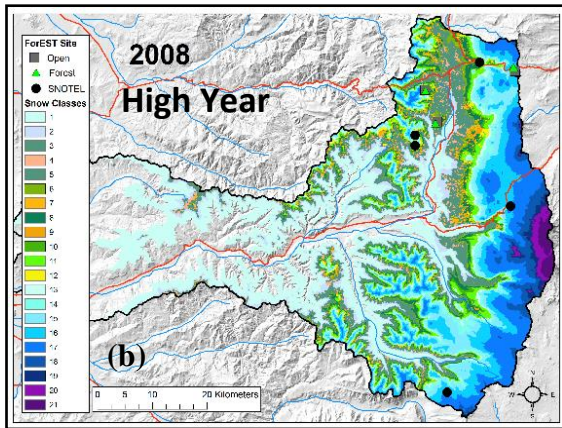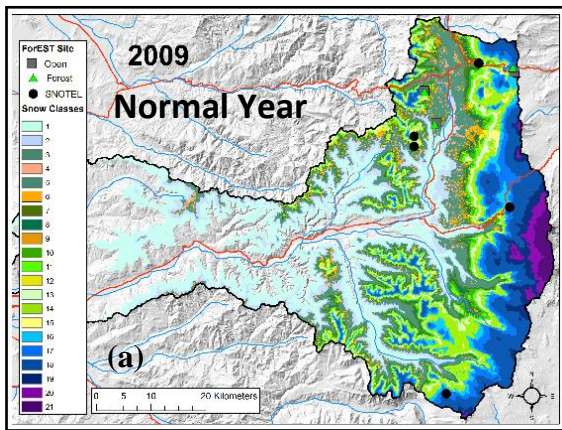
Figure 4. Spatially distributed snow classes derived from the binary regression tree model and geospatial model for, a) a normal snow year (2009), b) a high snow year (2008), and c) a low snow year (2005). Blue/purple colours represent snow classes distributed by elevation for all land covers, green colours represent snow classes distributed by elevation and forest land covers, and orange/yellow colours represent snow classes distributed by elevation and open land covers. The selected locations for the snow monitoring sites were not evenly distributed in space, but were selected to span the range of spatial variability in snow-vegetation-climate interactions.
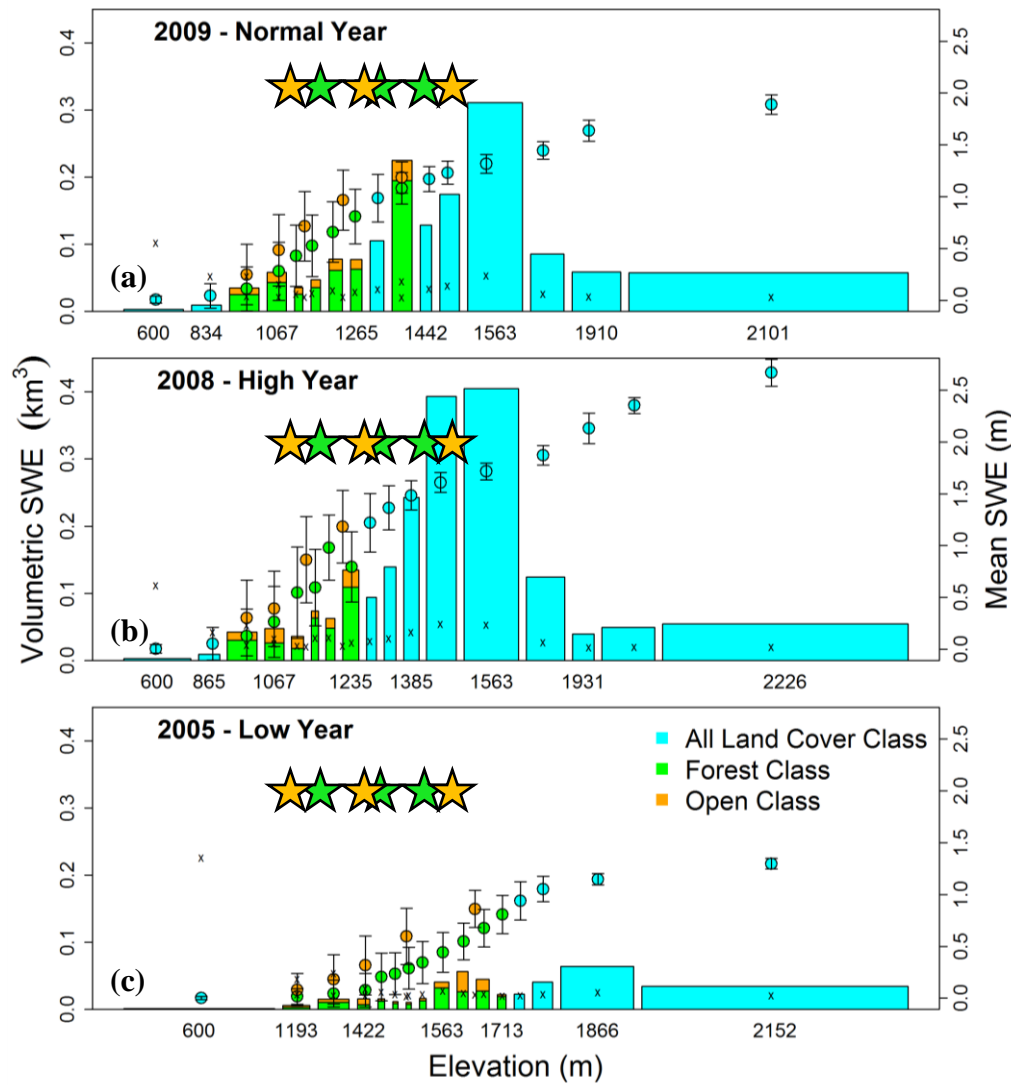
Figure 5. Volumetric SWE (km³) (shown as bar height), across the elevation range where the classes are located (shown as bar width), of the 21 binary regression tree (BRT) derived snow classes for, a) a normal snow year (2009), b) a high snow year (2008), and c) a low snow year (2005) with the minimum elevation for each class labelled on x-axis. Mean SWE (m) (shown in coloured circles), standard deviation (shown as error bars), and area (shown as black x) within each BRT-derived snow class across the elevation range where the classes are located. The elevations of ForEST station locations are shown in orange stars for open sites and green stars for forested sites.
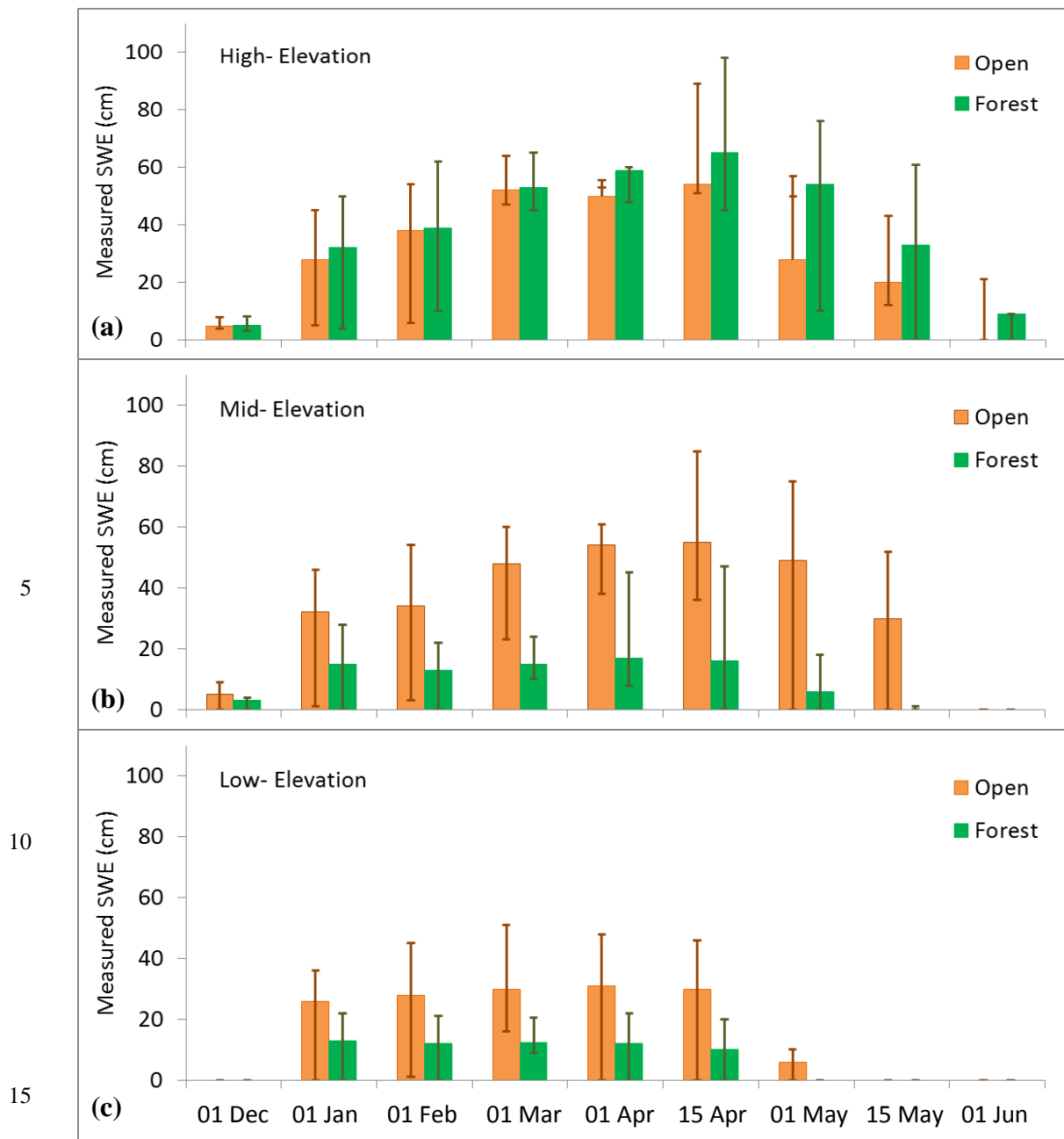
Figure 6. Mean SWE (cm) from snow course measurements collected at the paired open and forested snow monitoring sites in the ForEST network at, (a) high- (1483 and 1467 m), (b) mid- (1335 and 1332 m), and (c) low- (1113 and 1139 m) elevations during the winters of 2012, 2013, and 2014. Orange bars represent mean SWE (cm) in open sites. Green bars represent mean SWE (cm) in forested sites. Error bars indicate the maximum and minimum measured SWE (cm) from 2012, 2013, and 2014.